

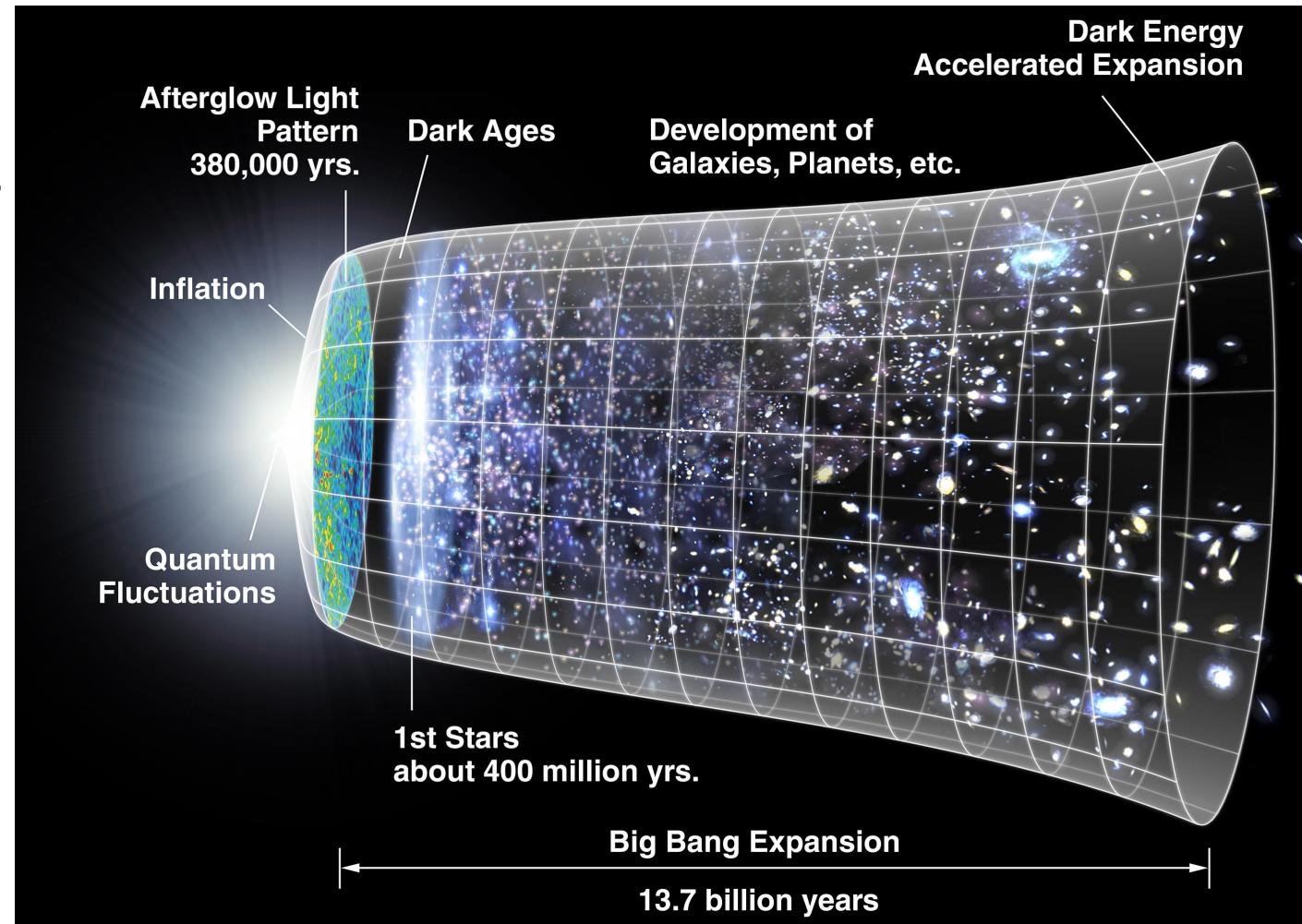
*Fundamentals of Earth Sciences*  
*(ESO 213A)*

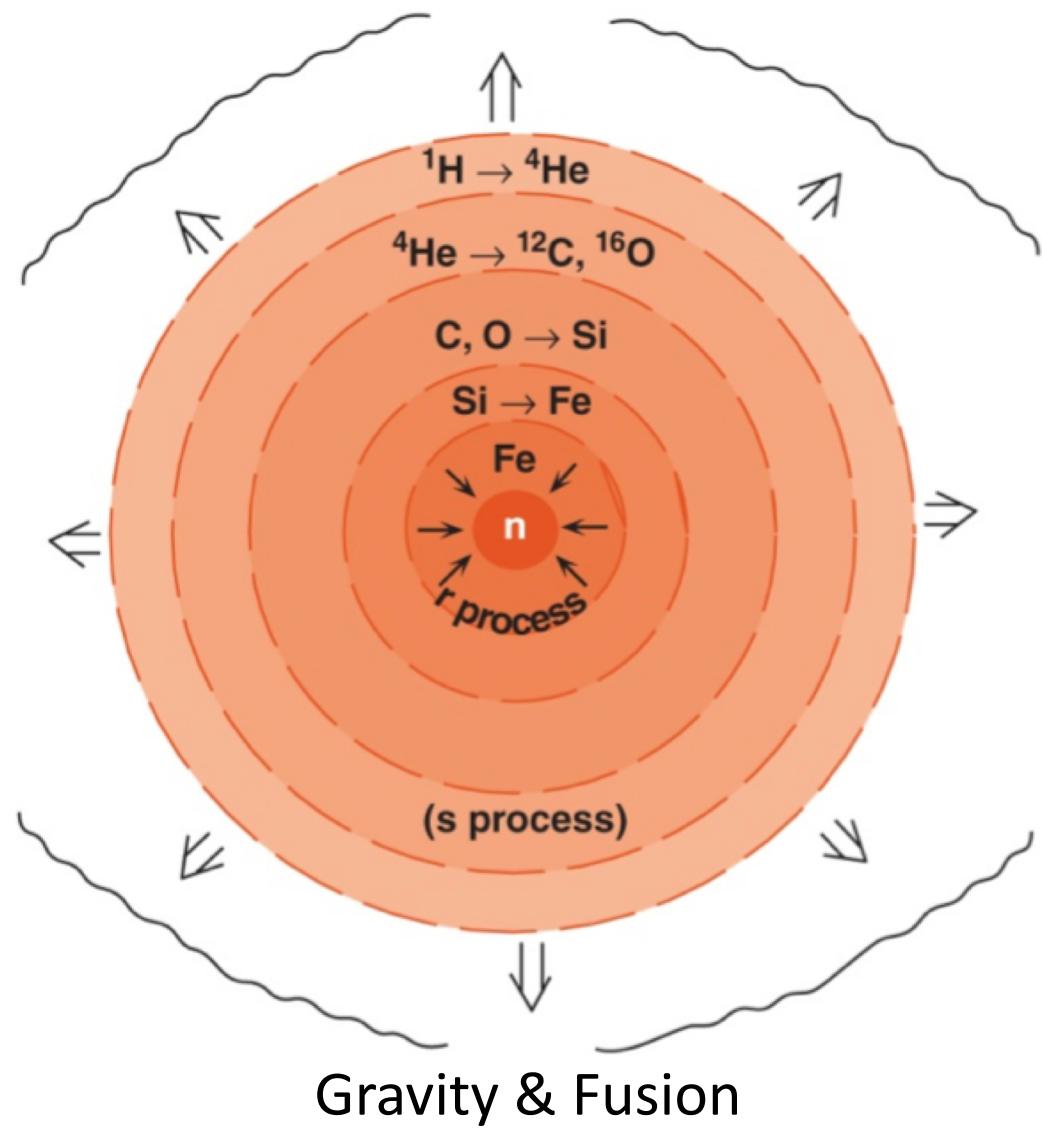
Dibakar Ghosal  
Department of Earth Sciences

Lecture 3: Origin of Earth, Moon and Atmosphere

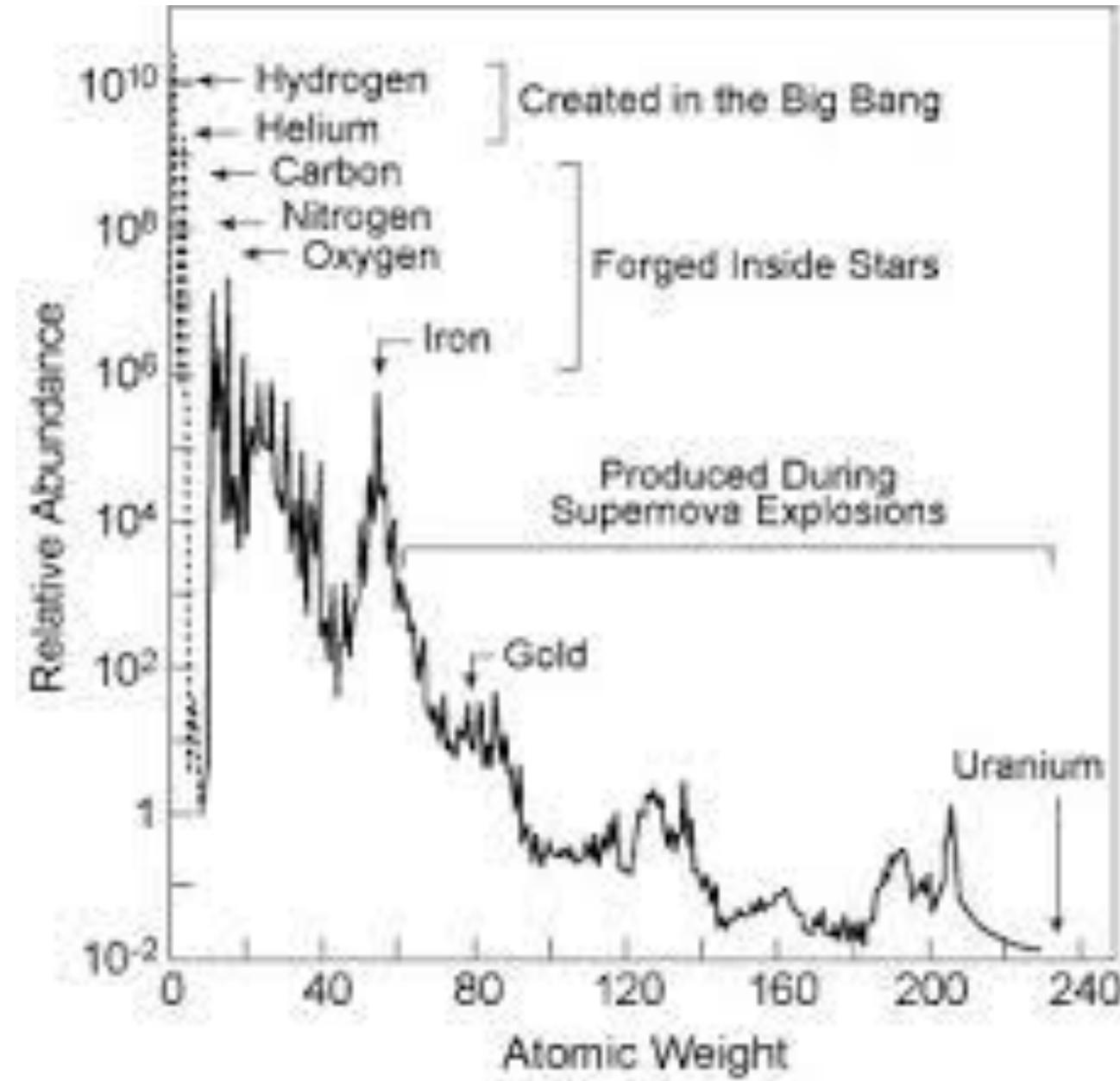
# Last Class: Review

- Big Bang Theory (13.7 billions of years ago)
- Nucleosynthesis (BigBang or Primordial and Stellar)
- Supernova Explosion
- Origin of our Solar System (4.57 billions of years ago)





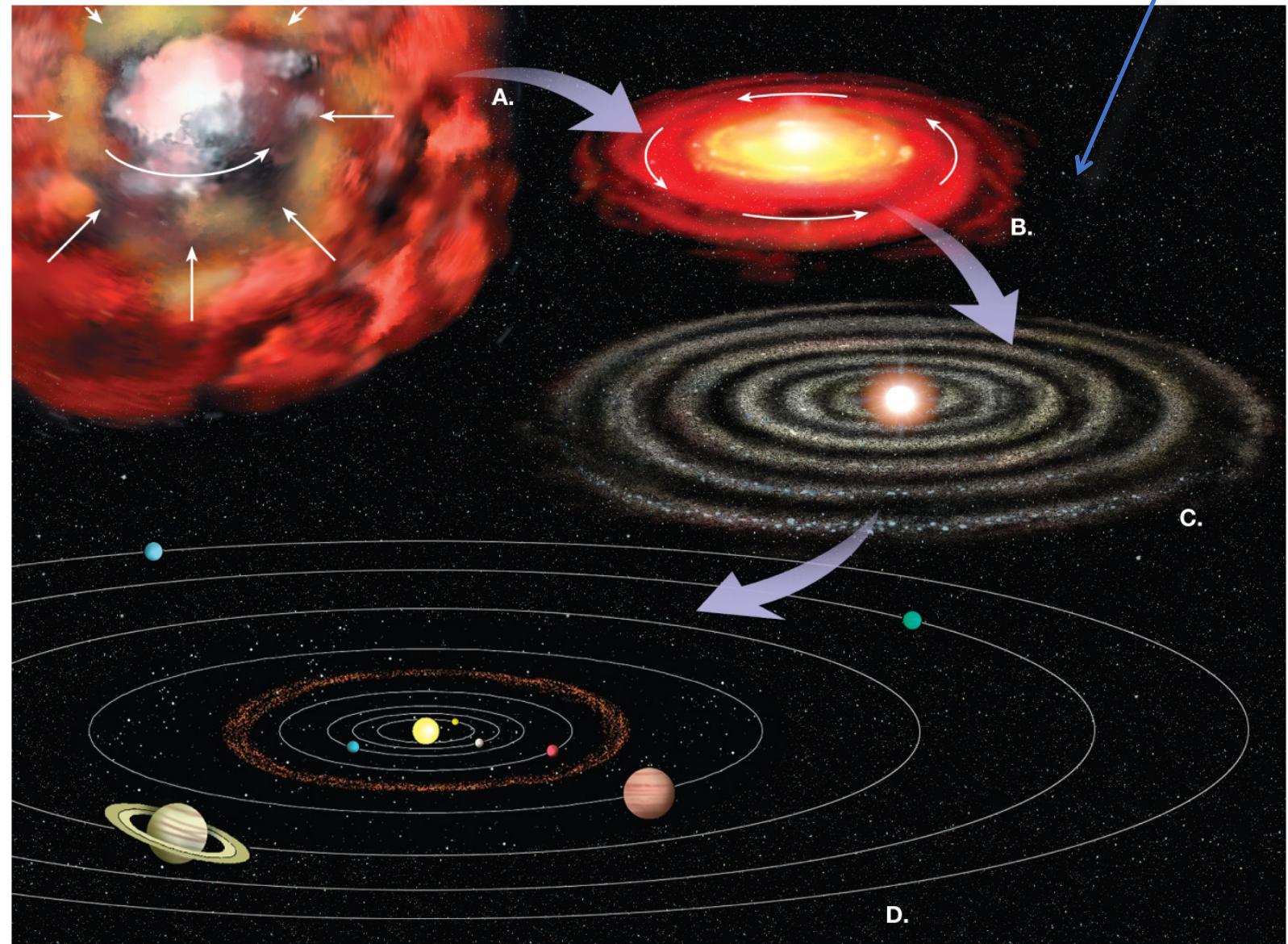
Stellar Nucleosynthesis && Neutron capture (s-r) processes



# Formation of Solar System

4.57 billions years ago

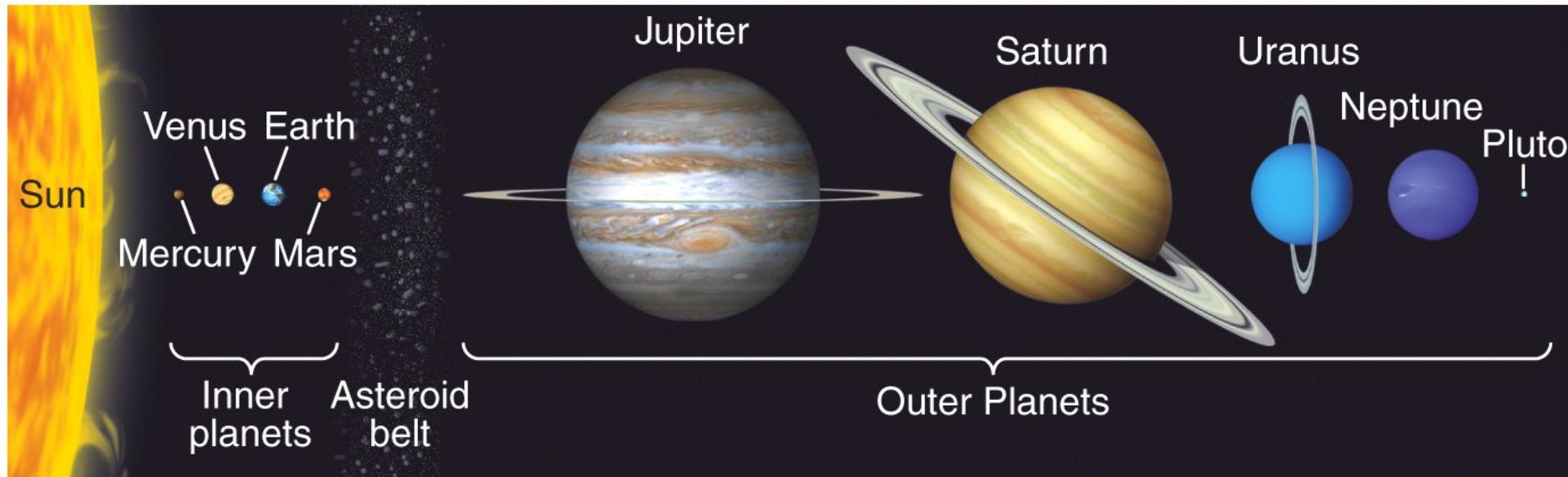
1. Cosmic cloud formed from dust of previous Supernovae
2. Gravity pulls particle closer and increase rotation
3. Nuclear reaction begins ~ 5 billion years ago at the centre of the cloud due to high concentration of material
4. Remaining particle continues to come closer and **accretion** took place
5. Eventually larger particles accreted and formed 'planets'



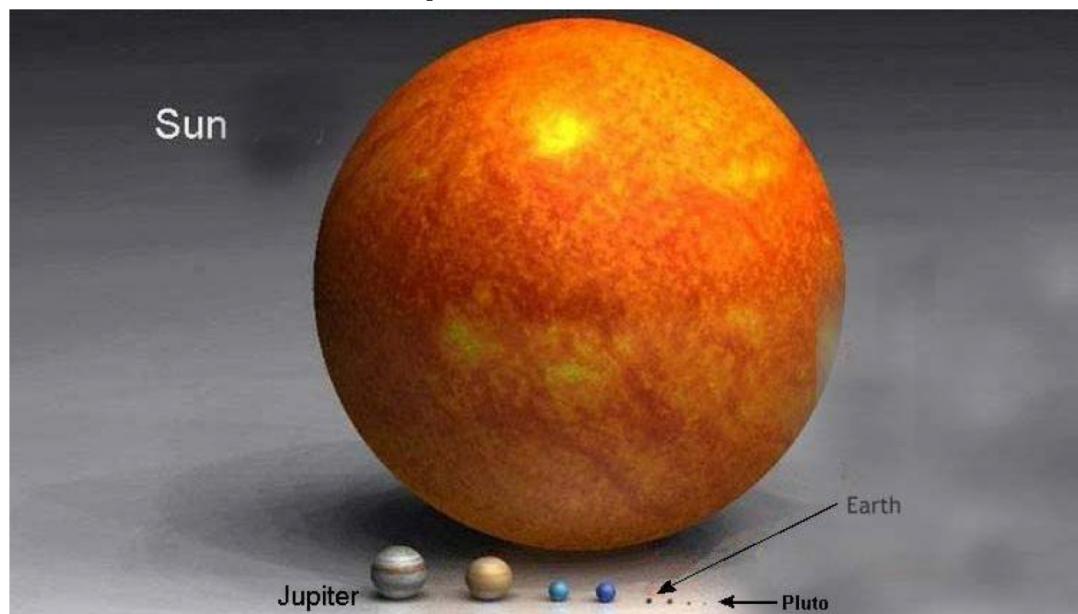
Dust >rocks >planetesimals >embryos >planets

*The Nebular Theory*

# Today's Solar System



## Sizes of the planets relative to Sun



Sun-planet distance  
(relative to Earth: AU)

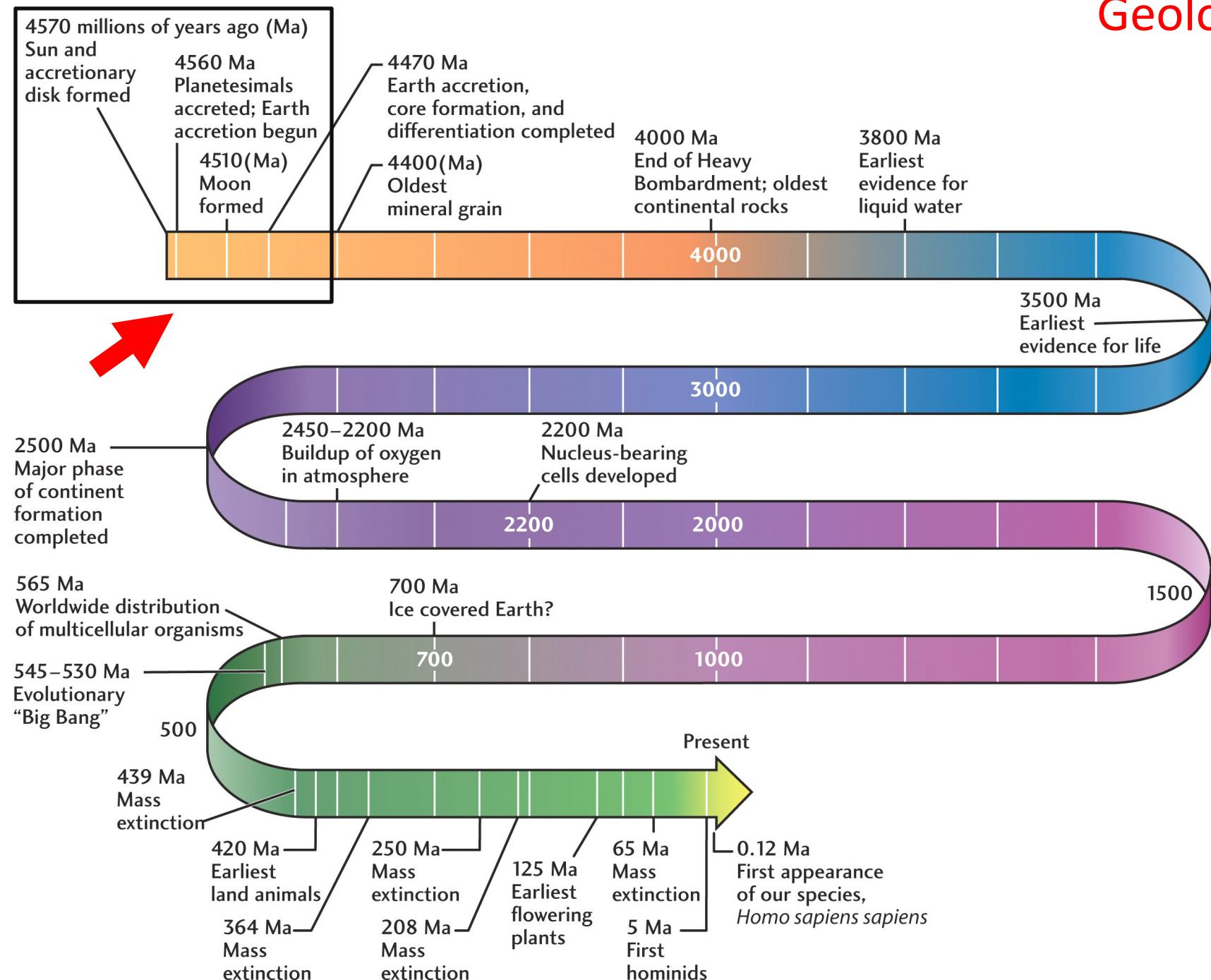
Mercury	0.4	AU
Venus	0.7	
Earth	1.0	
Mars	1.5	
Jupiter	5.2	
Saturn	9.5	
Uranus	19	
Neptune	30	

1 AU = 150 million km

Two different types of planets in Solar system:

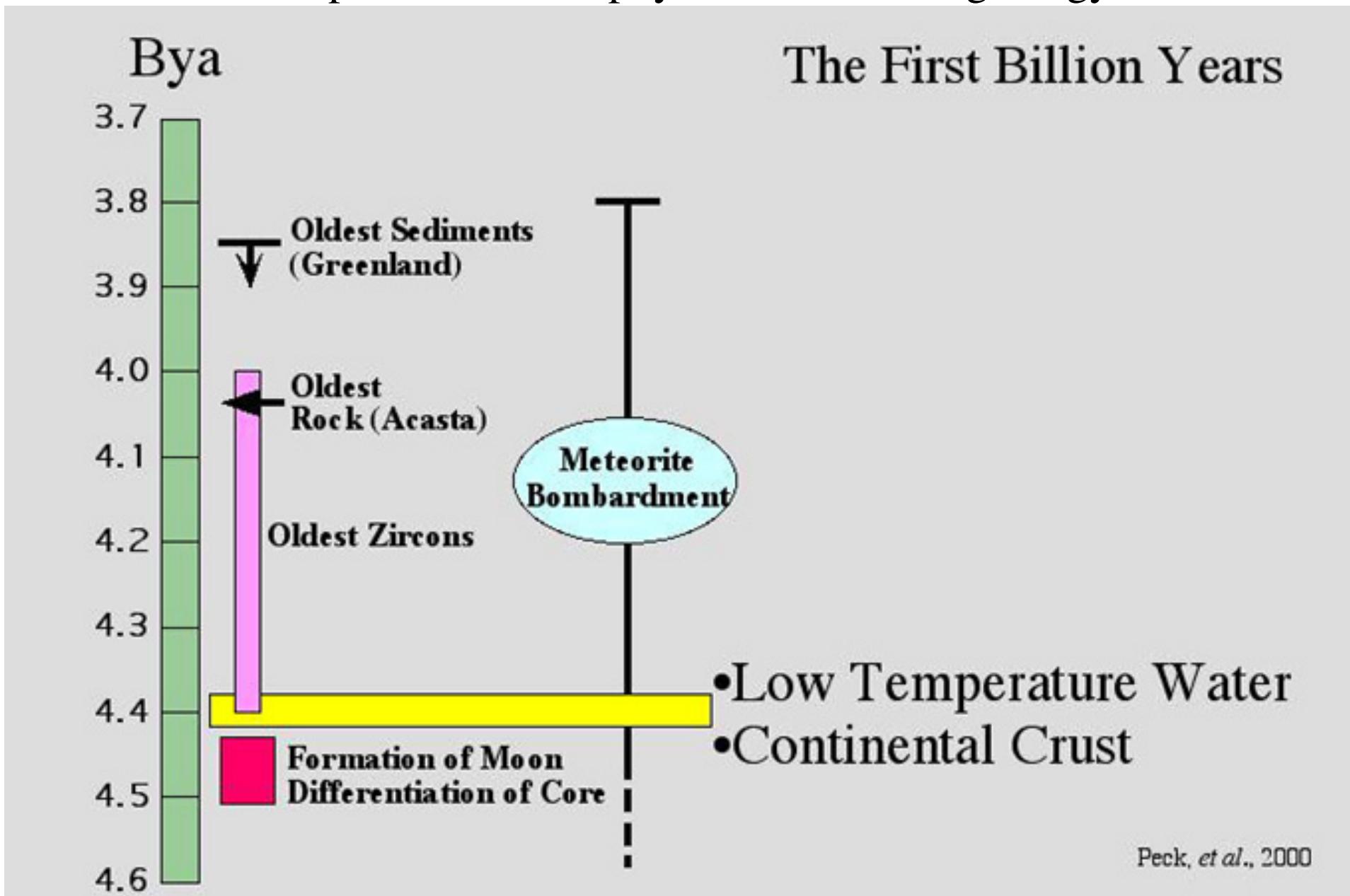
1. Terrestrial planet:  
Mercury, Venus, Earth, Mars  
(Small rocky, dense, closer to Sun)
2. Jovian planets  
Jupiter, Saturn, Uranus, Neptune  
(larger, more gaseous, less dense, away from Sun)

# Geologic Time Scale



# Connecting the dots: From planet formation to early Earth

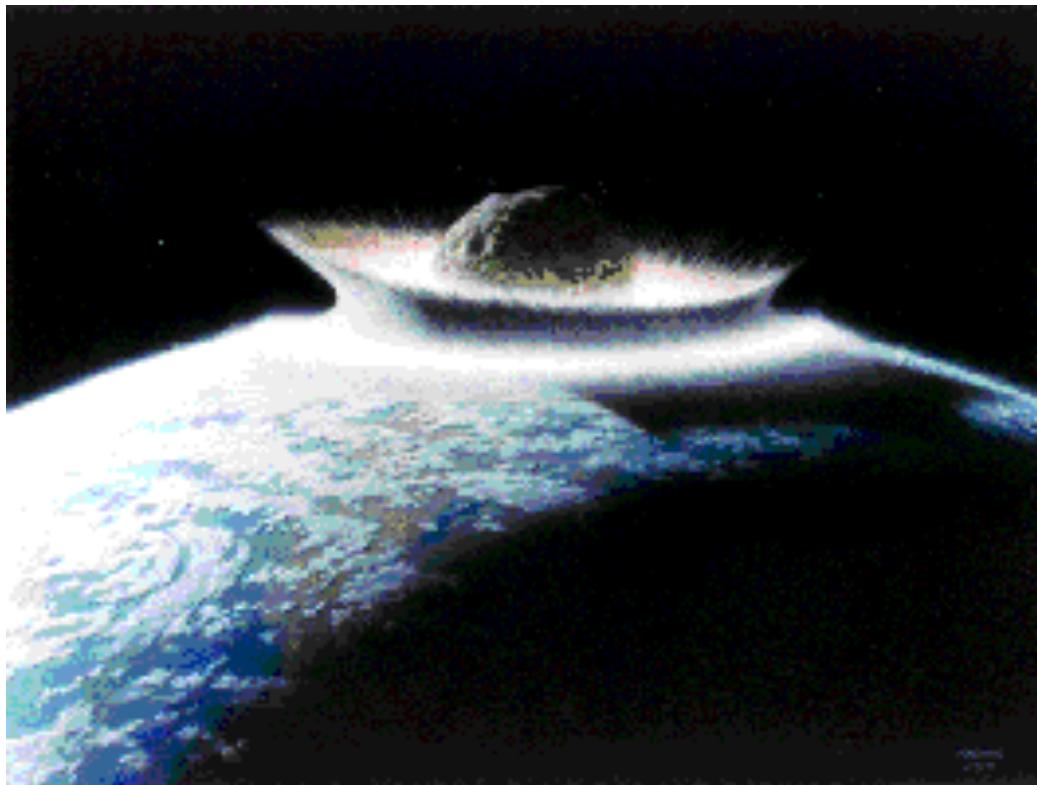
## Computational astrophysics meets field geology!



# Bombardment From Space

For the first 50 million years of its existence, the surface of the Earth was repeatedly pulverized by asteroids and comets of all sizes

One of these collisions formed the Moon



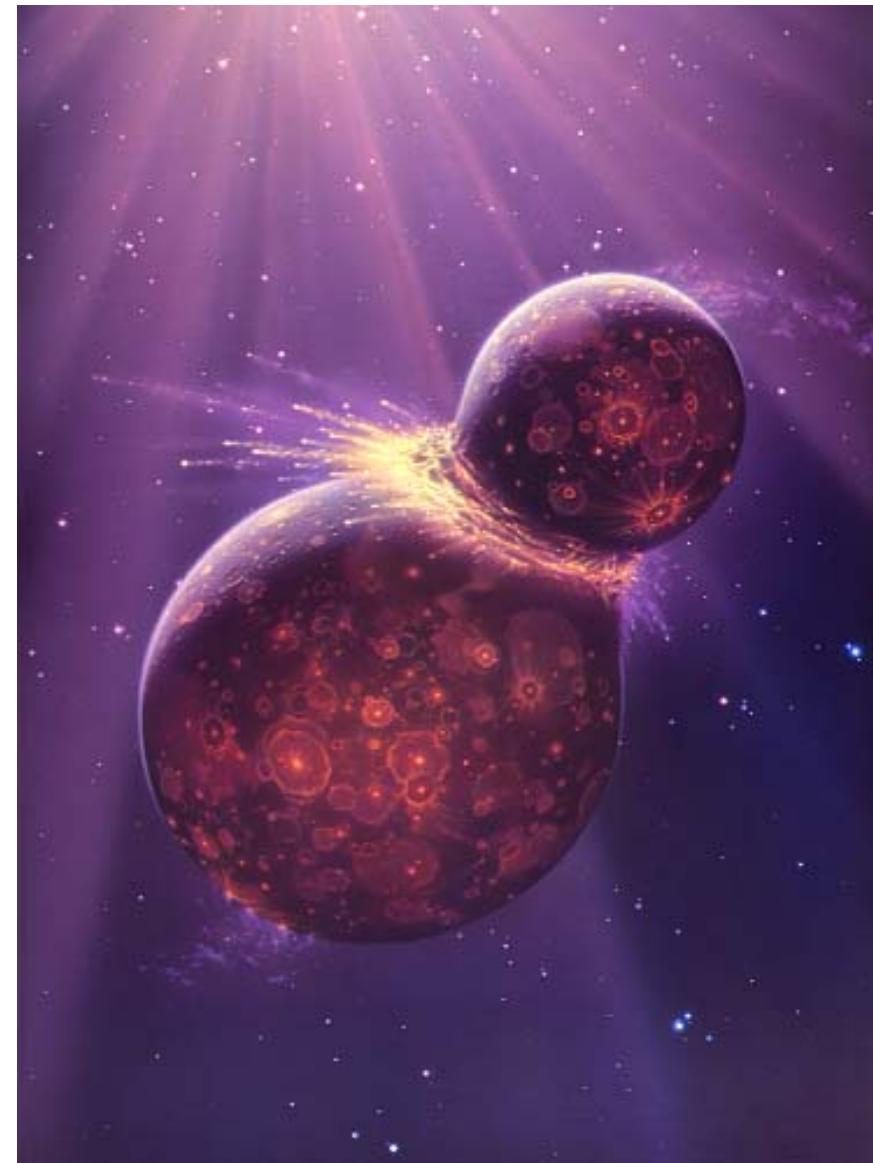


**Chelyabinsk meteor** entered Earth's atmosphere over the southern Ural region in Russia on 15 February 2013

# Formation of the Moon

**Giant Impact Hypothesis** predicts that around 50 million years after the initial creation of Earth, a planet about the size of Mars collided with Earth

This idea was first proposed about 30 years ago, but it took calculations by modern high-speed computers to prove the feasibility

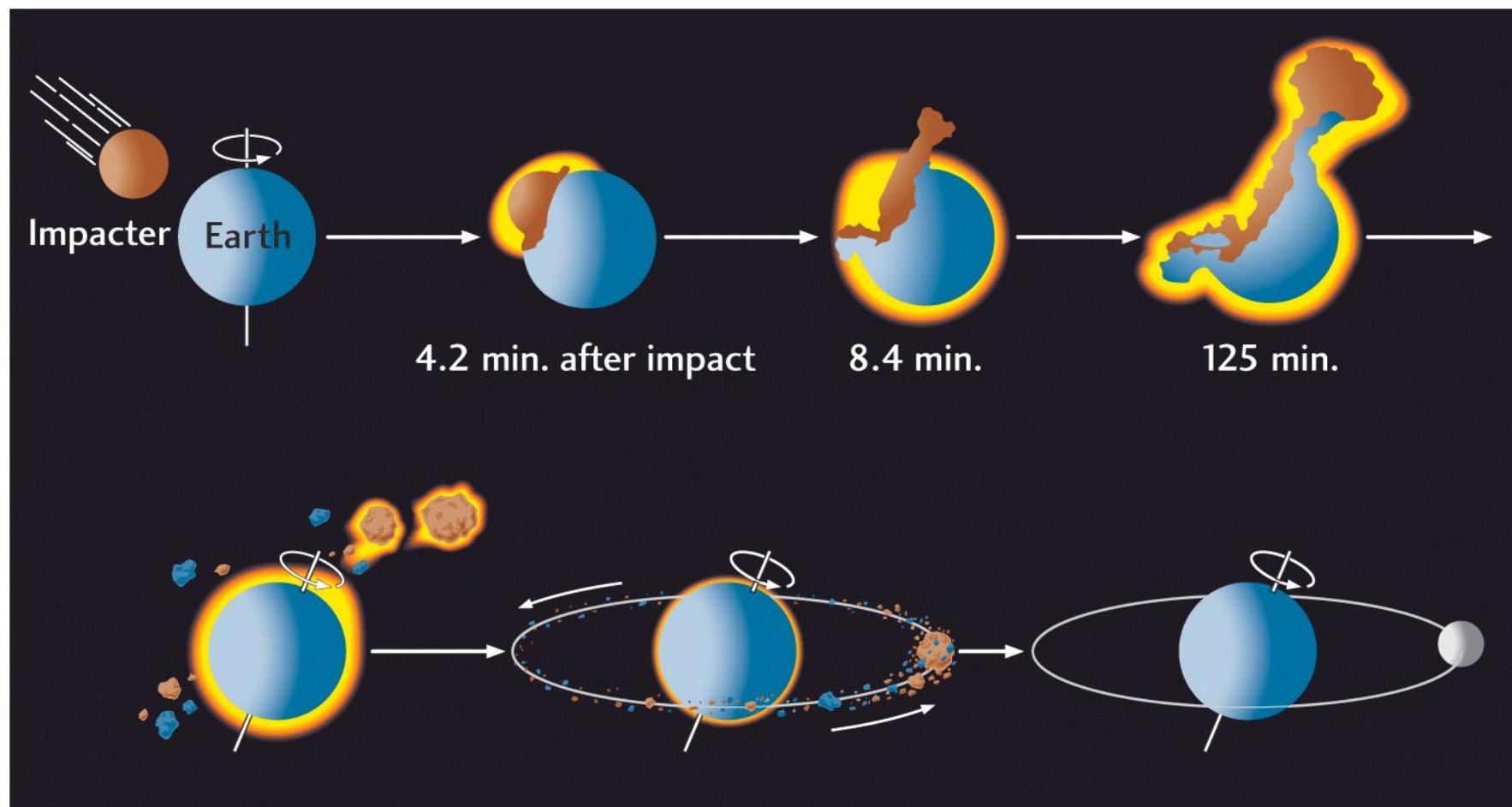


Why Moon Formation is Important?

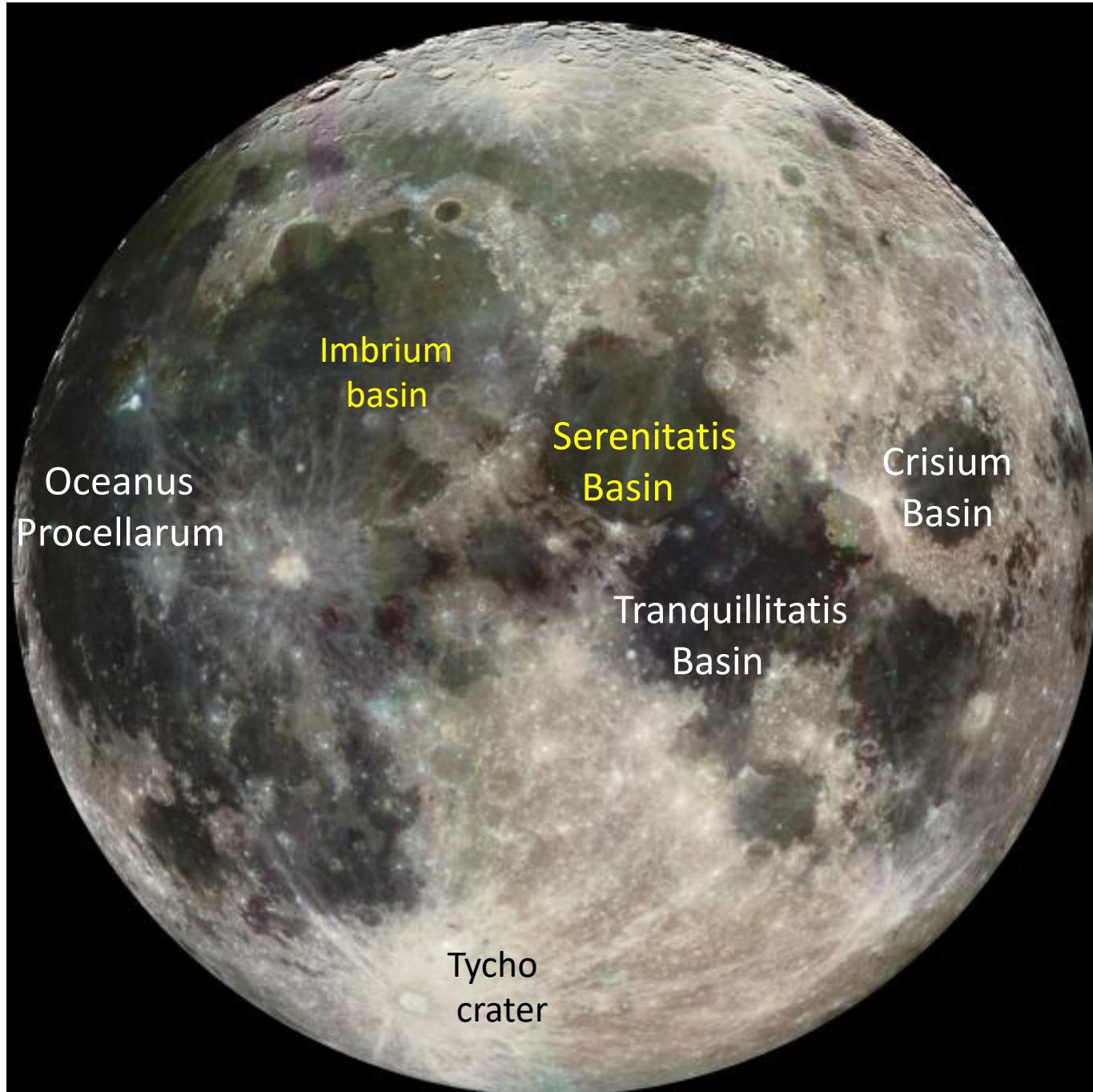
# Formation of the Moon

This collision had to be very spectacular!

A considerable amount of material was blown off into space, but most fell back onto the Earth

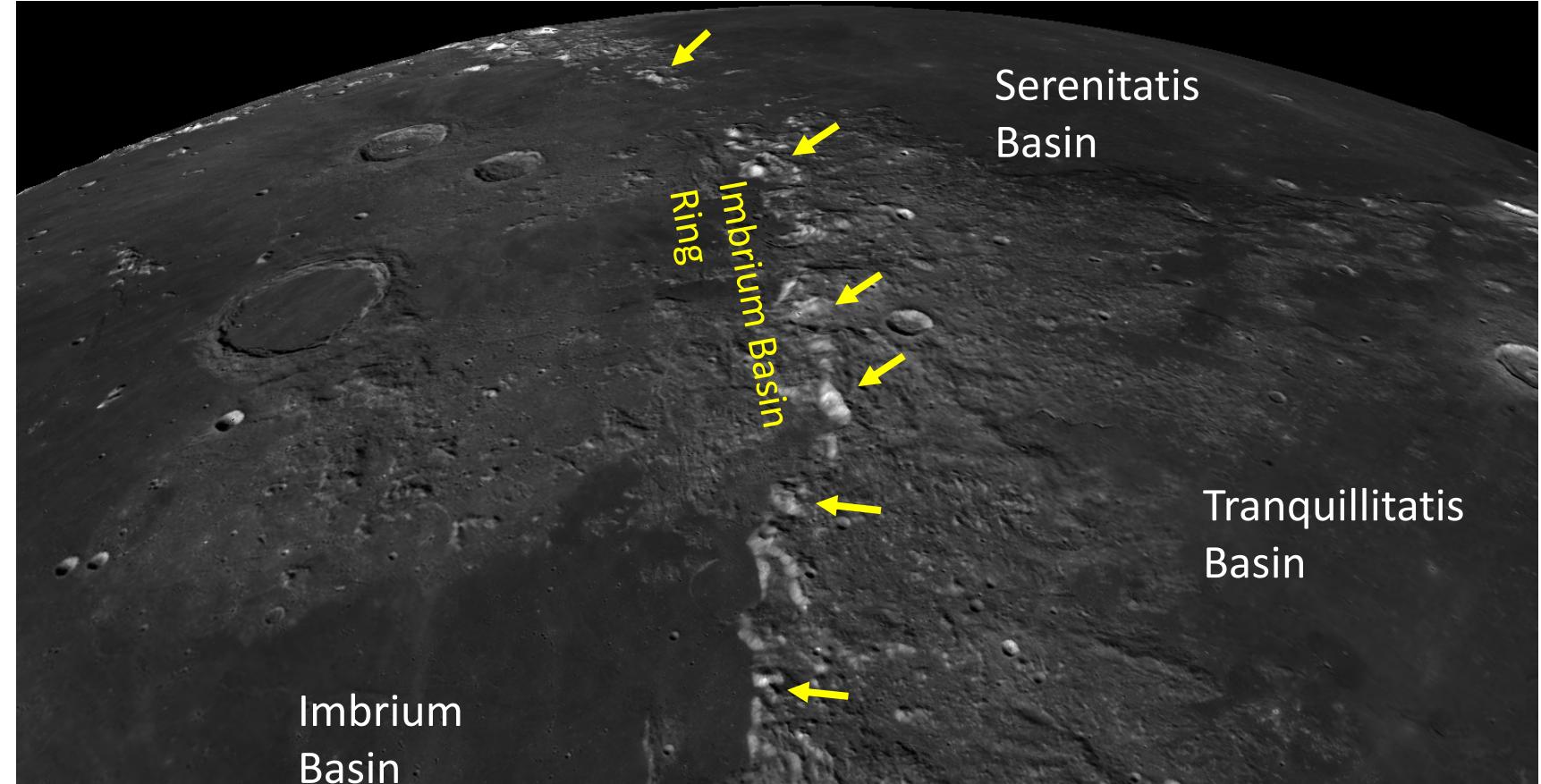


Moon from Galileo spacecraft. The distinct bright ray crater at the bottom of the image is the Tycho impact crater. The dark areas are lava rock filled impact basins: Oceanus Procellarum (more than 2,500 km), Mare Imbrium (diameter of 1145 km), Mare Serenitatis (diameter 674 km) and Mare Tranquillitatis (diameter 873 km), and Mare Crisium (diameter 556 km). This picture contains images through the Violet, 756 nm, 968 nm filters.



Both Earth and Moon  
were struck by numerous  
large asteroids and comets  
in their early history. These  
impacts produced deep  
basins up to 2500 km  
across surrounded by high  
rings of mountains on the  
Moon and are visible to  
the human eye as  
prominent circular  
structures.  
A view of the mountains  
that surround the Imbrium  
impact basin. Lava flow  
observed both side of the  
ridge.

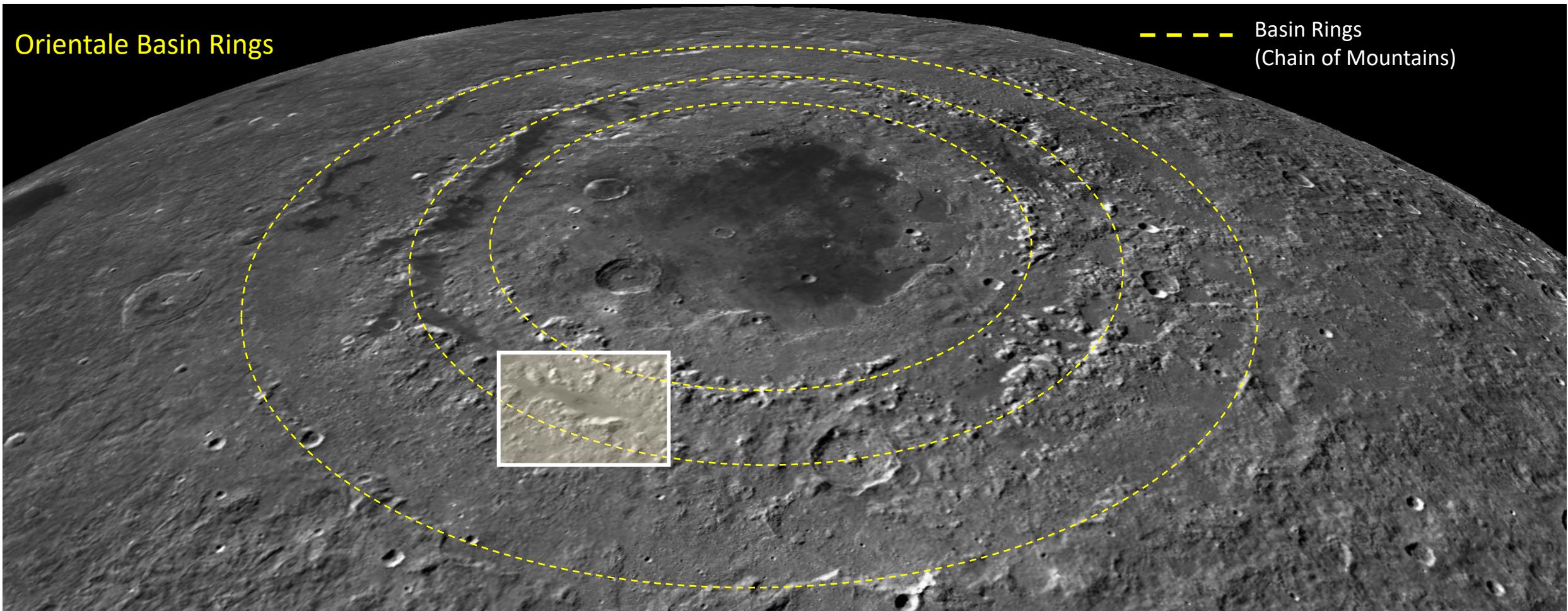
# Lunar Impact Basins



Courtesy of Prof. D. Dhingra

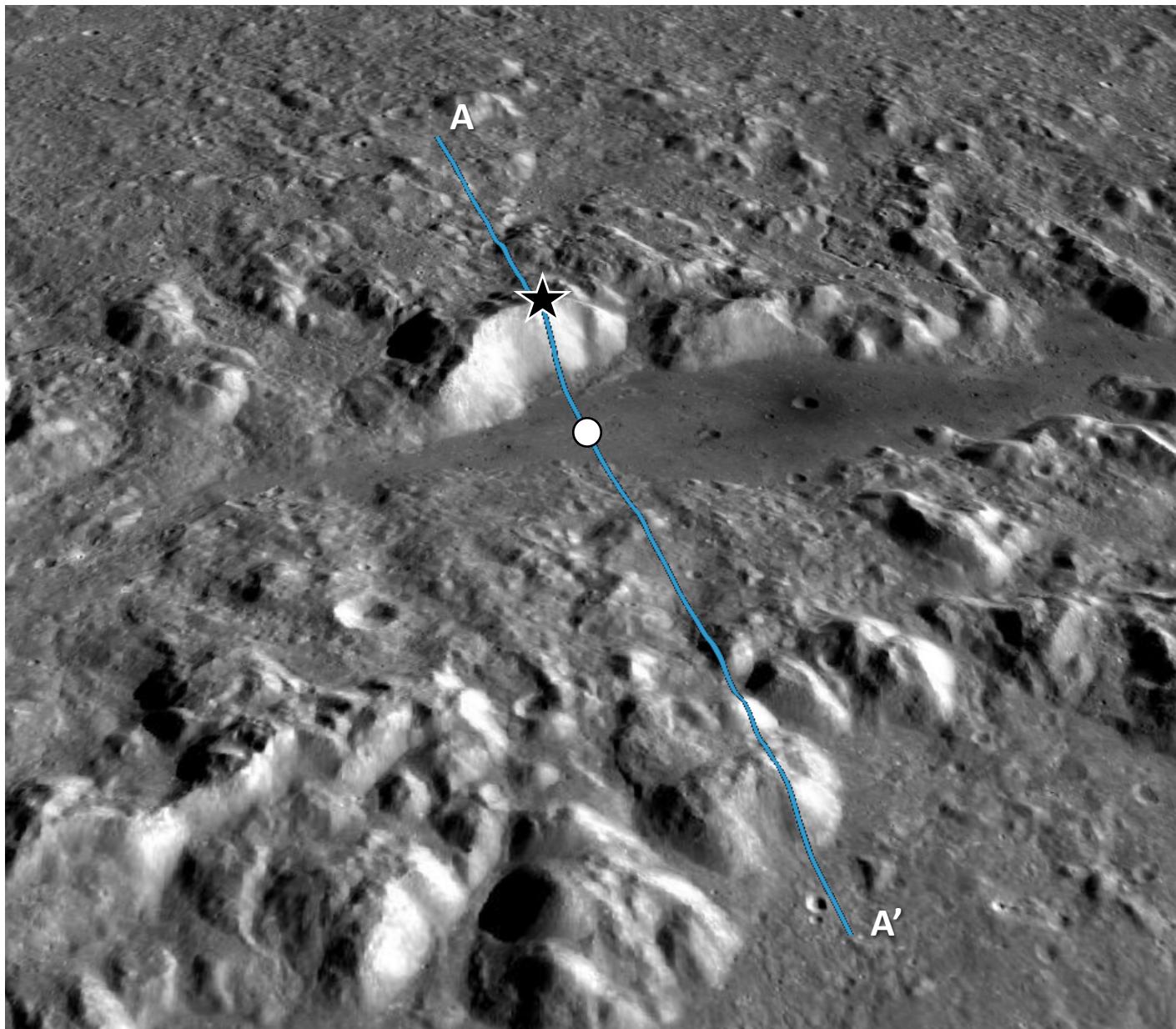
*Big, frequent impacts until 3.8 billion years ago  
Impact events continue on all moons and planets today*

# Lunar Impact Basins



Courtesy of Prof. D. Dhingra

Three mountain rings surround the Orientale impact basin. Diameter of the outer ring is ~900 km. Both the Imbrium and the Orientale impacts occurred around 3.8 billion years ago.



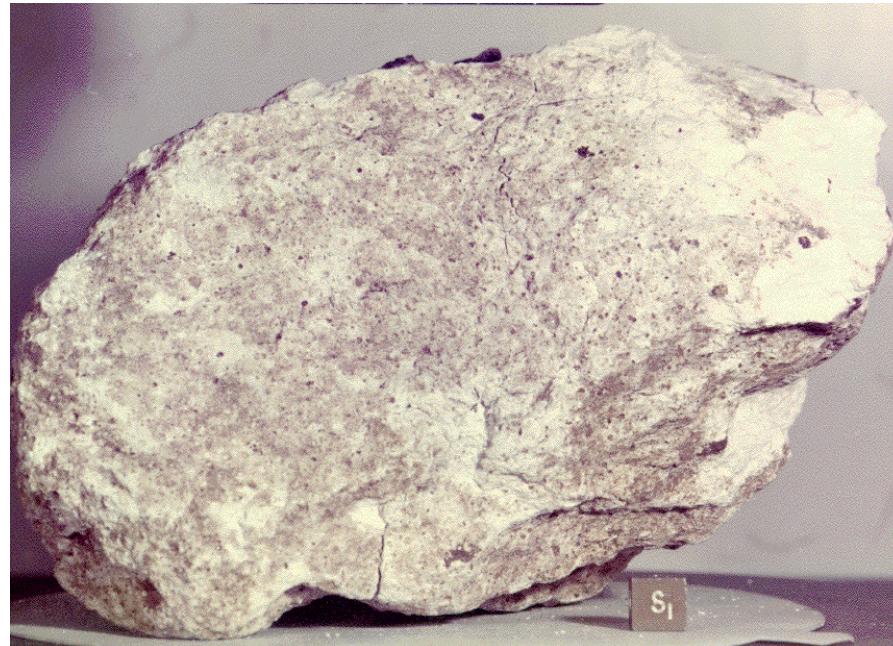
How high are the basin rings?



Height of mountain: 5.4 km !

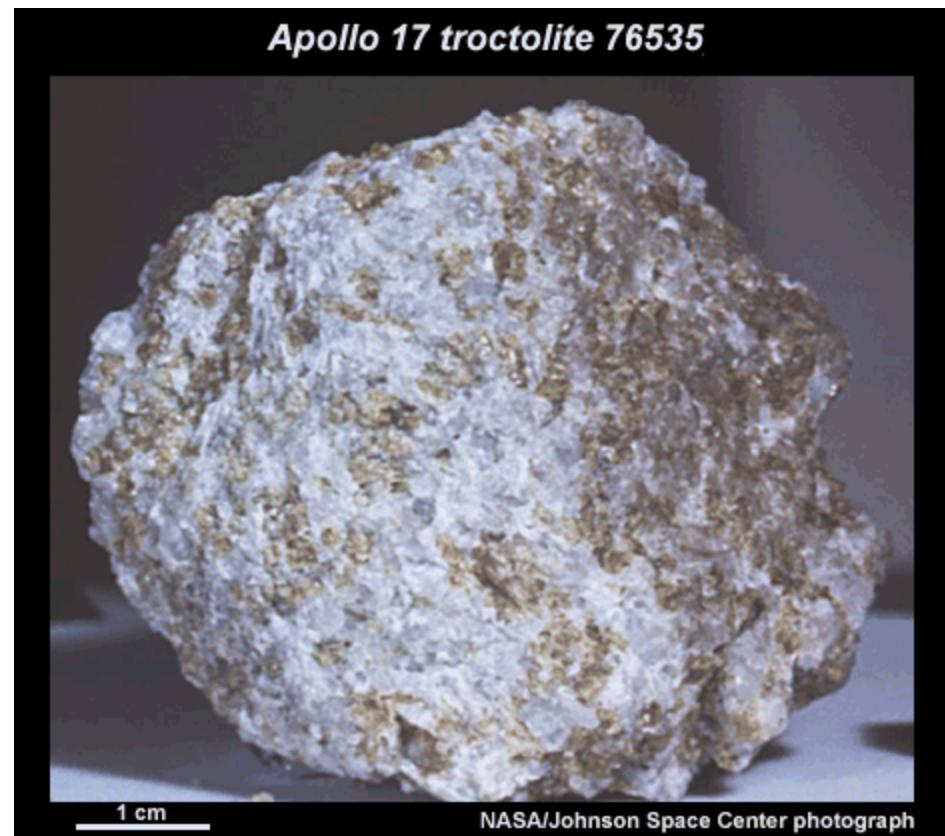
# Magma Ocean Rocks

These are two examples of rocks that crystallized from the lunar magma ocean. Both are made primarily of the mineral plagioclase, which gives the lunar highlands its light gray color. The troctolite also contains some greenish olivine grains (troctolite may be later intrusion).



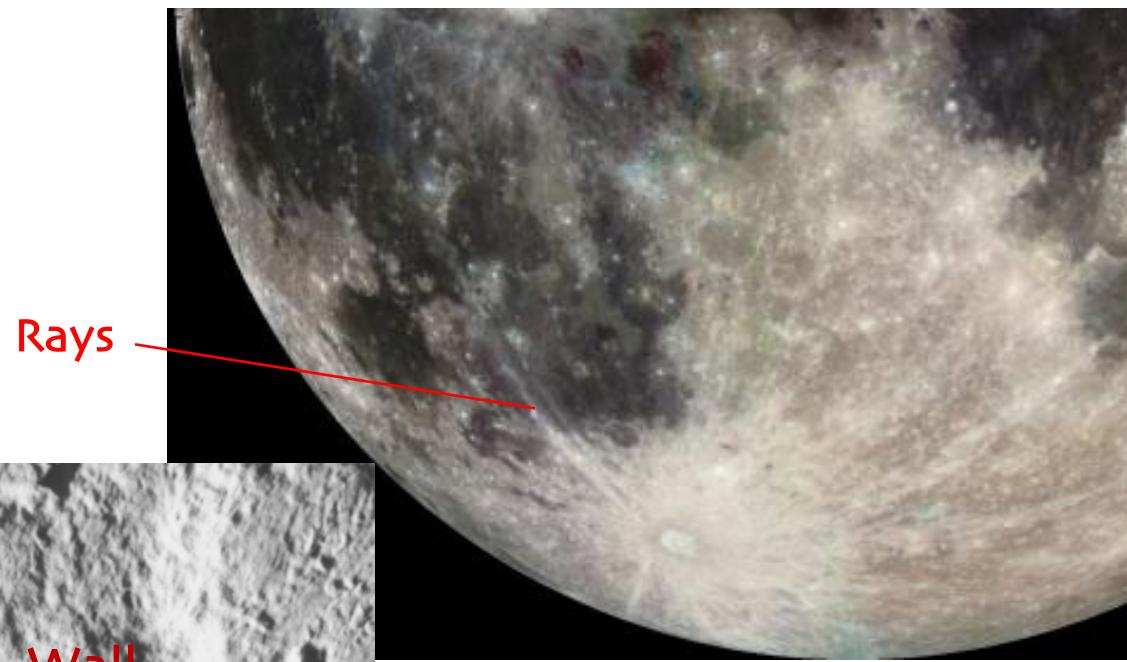
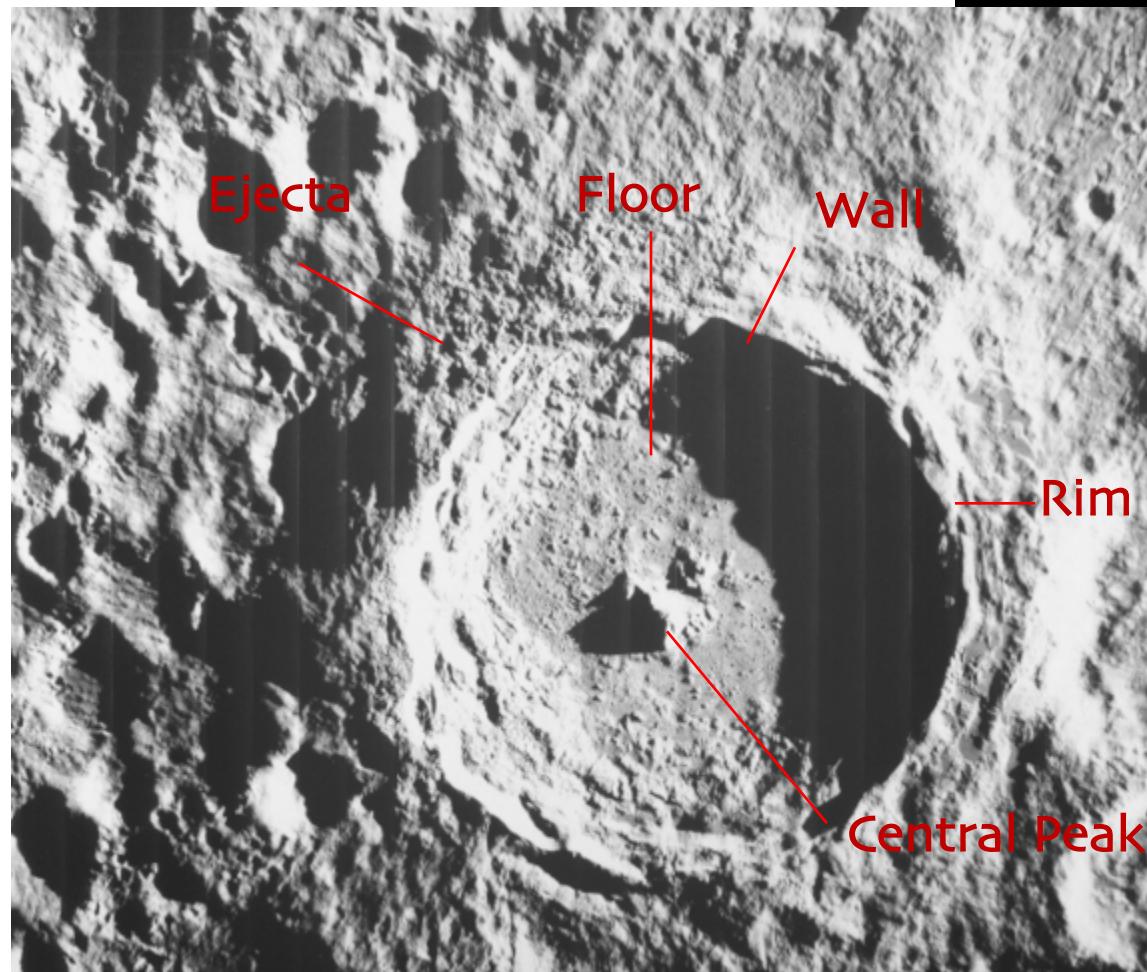
60025  
Anorthosite  
4.44-4.51 Ga  
95% Plag Feld (anorthite)  
Magma Ocean!  
No Water!

76535  
Troctolite  
4.2-4.3 Ga



NASA/Johnson Space Center photograph

This image clearly shows both the central peak and terracing in the walls of Tycho. Tycho is in the lunar highlands, and the terrain surrounding the crater is quite rugged. [Multispectral images obtained by the Clementine spacecraft](#) show that the central peak has a different composition than the surrounding material, presumably because the central peak is composed of material that originated at greater depths in the Moon's crust. (*Lunar Orbiter image V-125M.*)



Tycho Crater

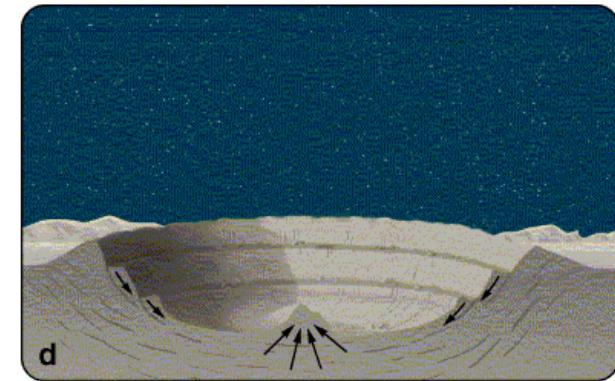
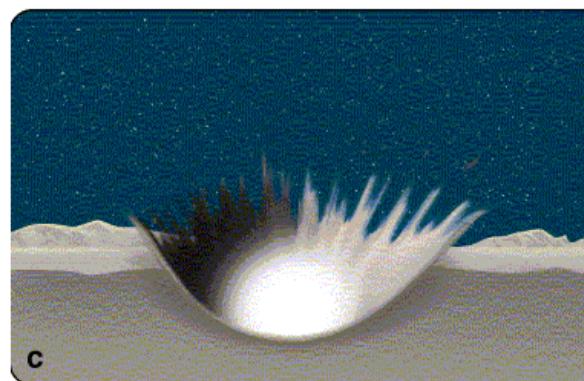
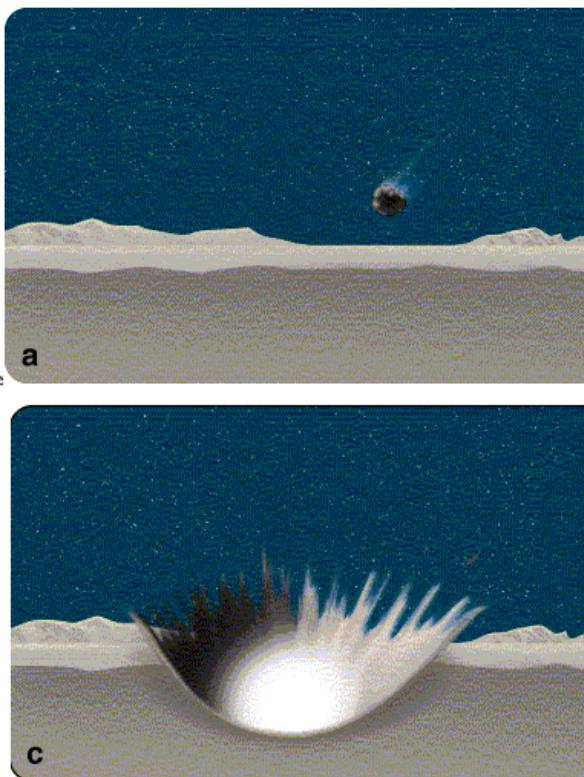
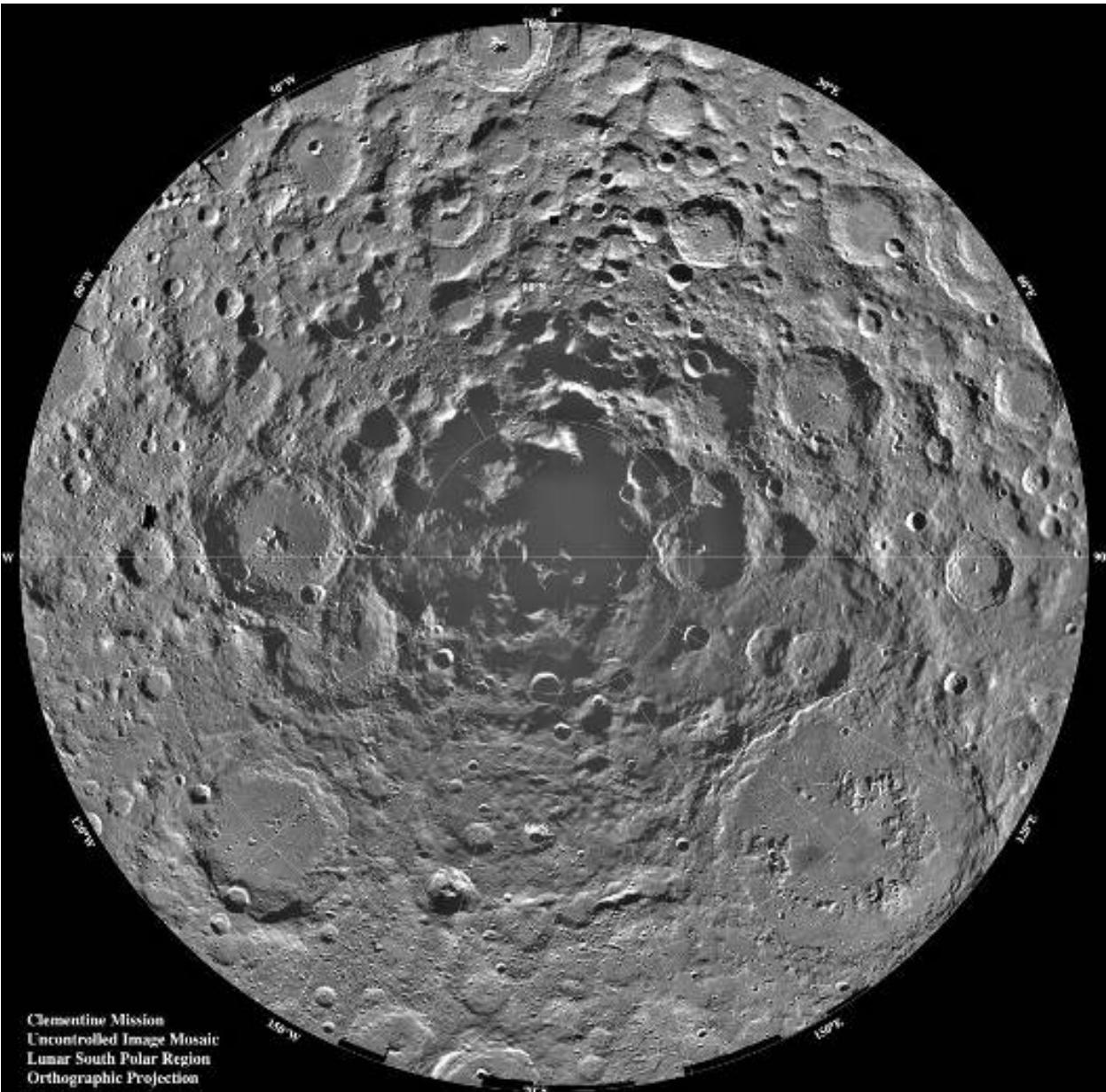
Young – 100 MY

85 kilometers  
diameter

Fresh (rays) = young

# Impacts ...“the most fundamental process on the terrestrial planets...”

*Eugene and Carolyn Shoemaker*



## Lunar Volcanism

Portions of the Moon's interior remained hot enough to produce magma for more than a billion years after it formed. Molten rock flowed onto the lunar surface through cracks in the crust, spreading out and filling the low regions in the impact basins. The lava cooled quickly, forming the fine-grained, dark rocks — basalt — sampled during the Apollo missions. The dark areas seen on the Moon are basaltic lava plains 4.2 to 1 billion. Volcanism in circular Mare Imbrium (left image). At right, shadows reveal the edges of a long lava flow from the lower left to the upper right of the image. The volcanism in Mare Imbrium occurred about 3.3 billion years ago (7 am on our clock). Because of its small size, the Moon cooled quickly and was mostly dead volcanically by 3 billion years ago, although limited volcanism in isolated regions is thought to have occurred as recently as 1 to 2 billion years ago.

## Lunar Geologic History Mare Volcanism



Mare Imbrium



SW Mare Imbrium

## Moon Becomes Geologically Inactive

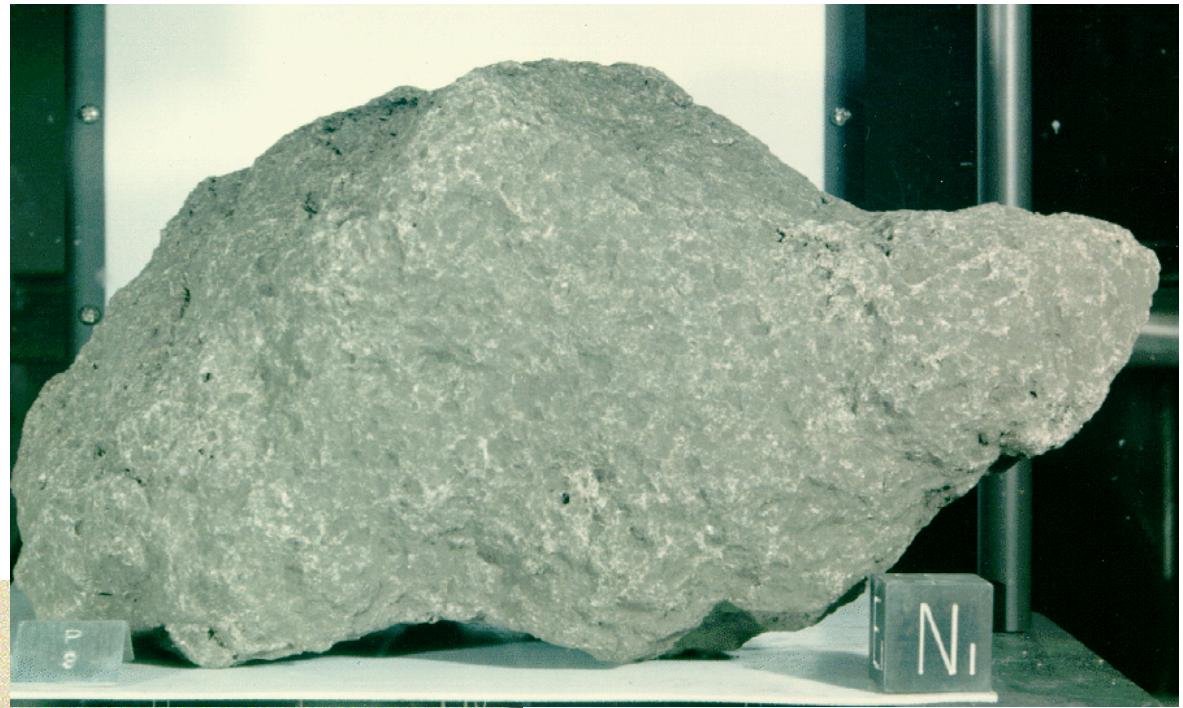
Lunar volcanism decreased significantly by 3 billion years ago and ceased completely by about 1 billion years ago as the interior of this small body cooled.

Volcanism *after* impacts – most before 3 Ga (to 1 Ga)

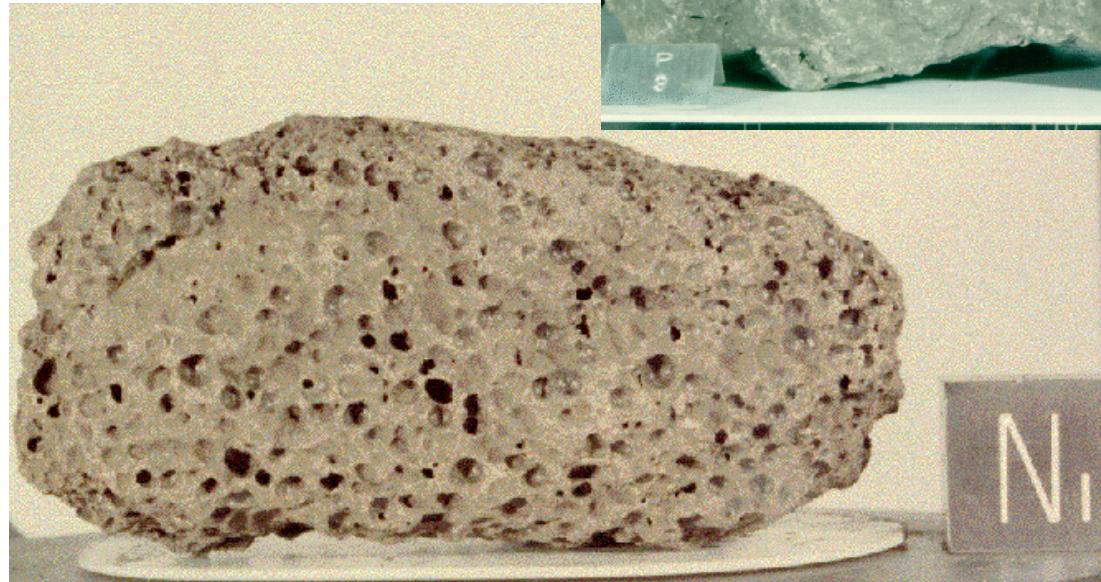
These rocks are typical of lunar volcanic rocks. Collected on Apollo 15, both are 3.3 billion year old basalts, similar to those produced by volcanos such as Hawaii on Earth. The lower image (sample 15016) contained some type of gas, possibly carbon monoxide, which formed the round holes known as vesicles.

## Lunar Basalts

15555

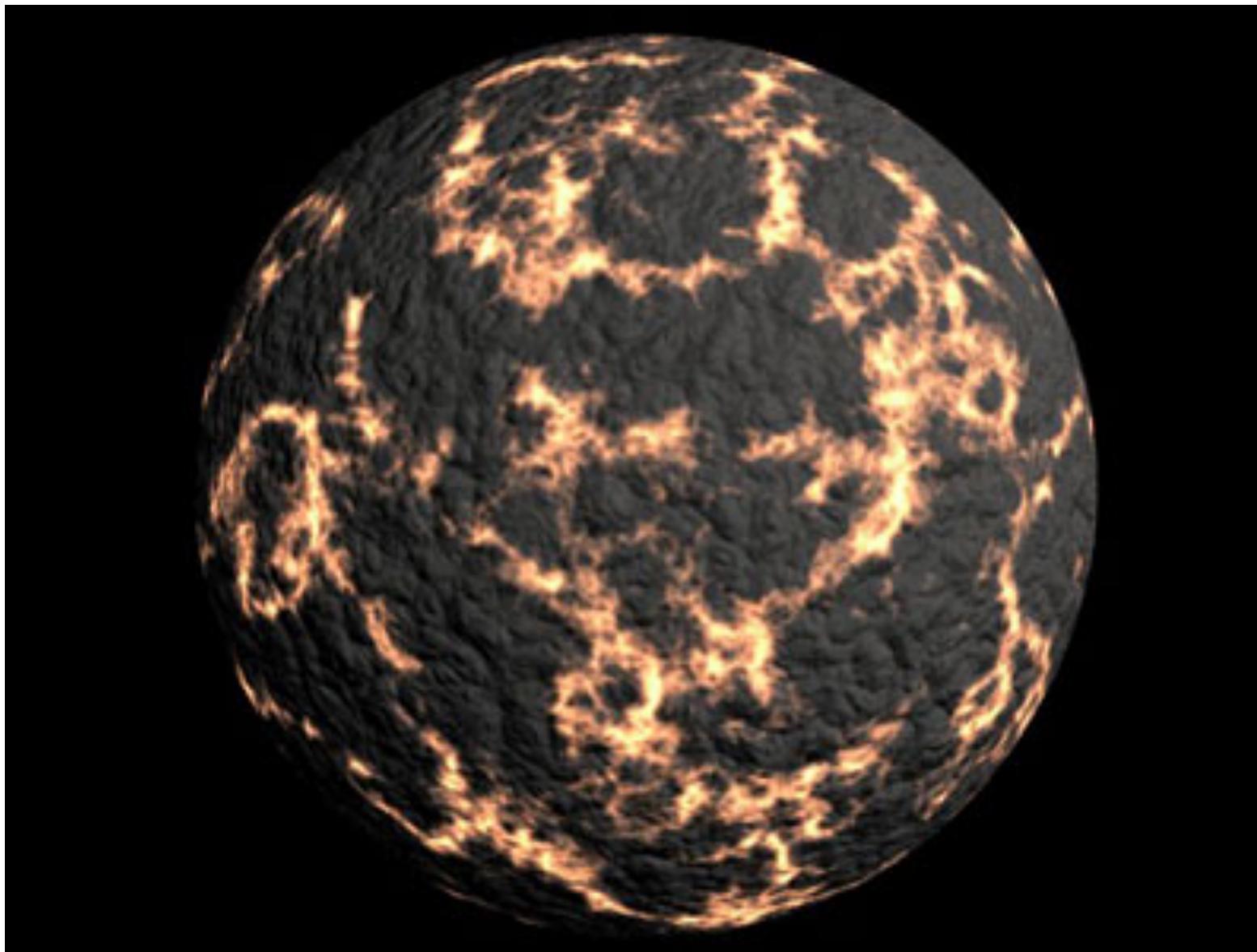


3.3 Ga



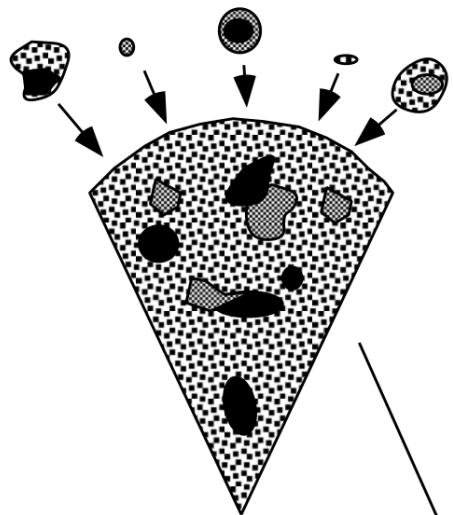
15016

# The Early Earth Heats Up (Magma Ocean)

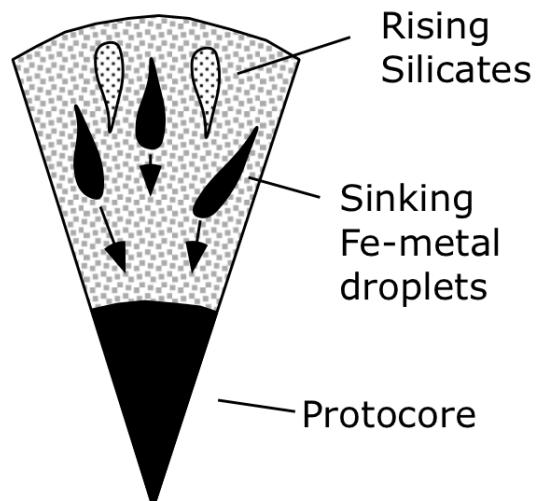
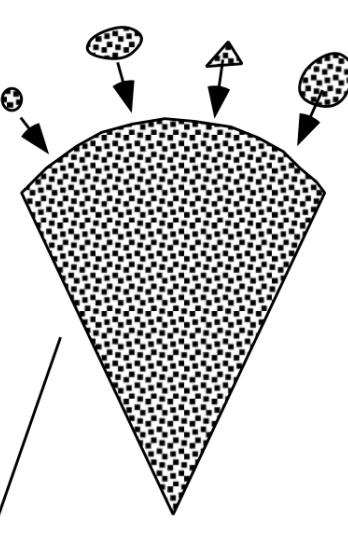


# Origin of the Earth's Internal Layering

Heterogeneous Accretion



Homogeneous Accretion



## Origin of the earth by accretion of planet-forming materials.

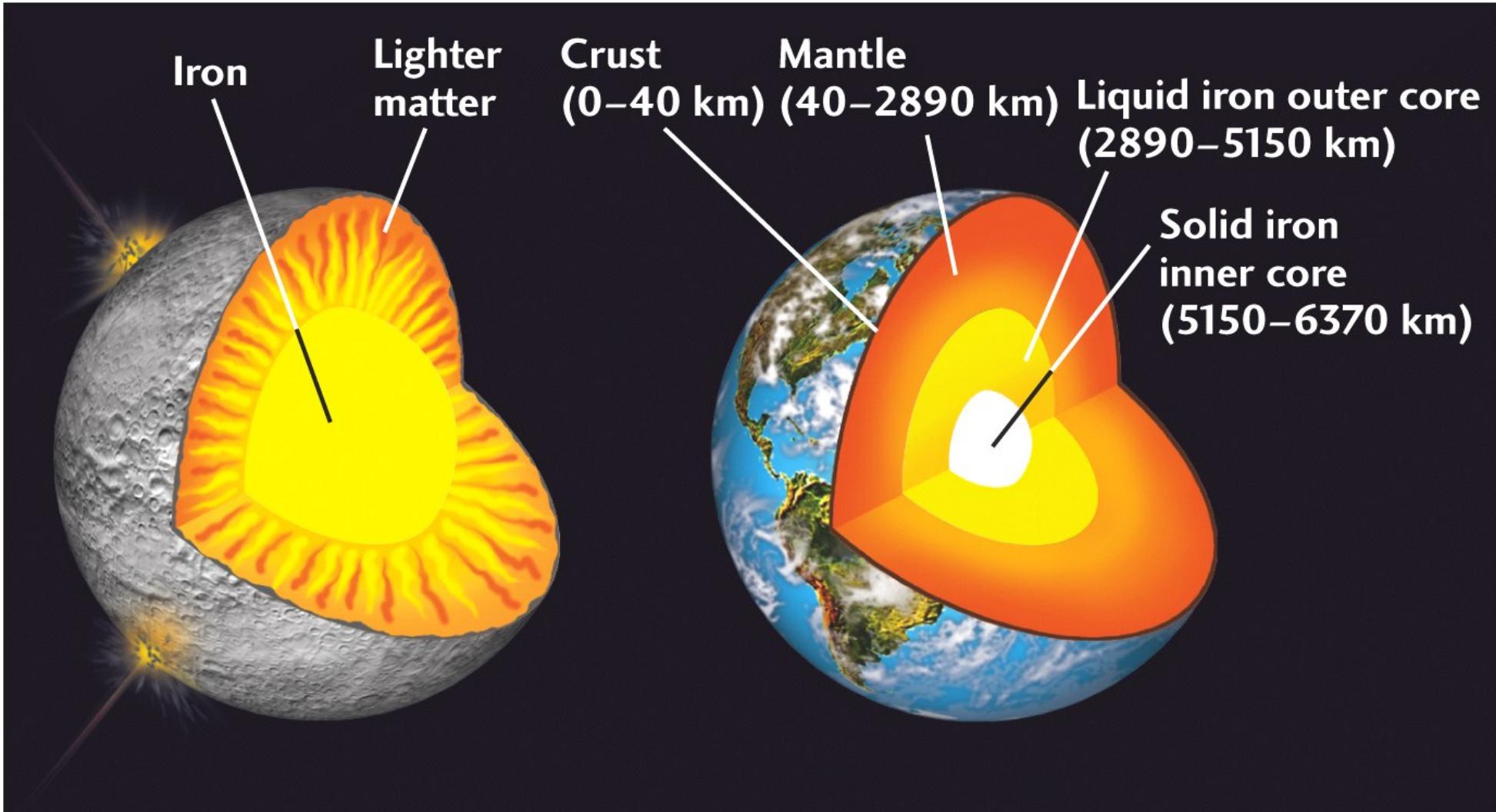
Terrestrial (inner) planets undergo differentiation process that includes melting and separation into layers

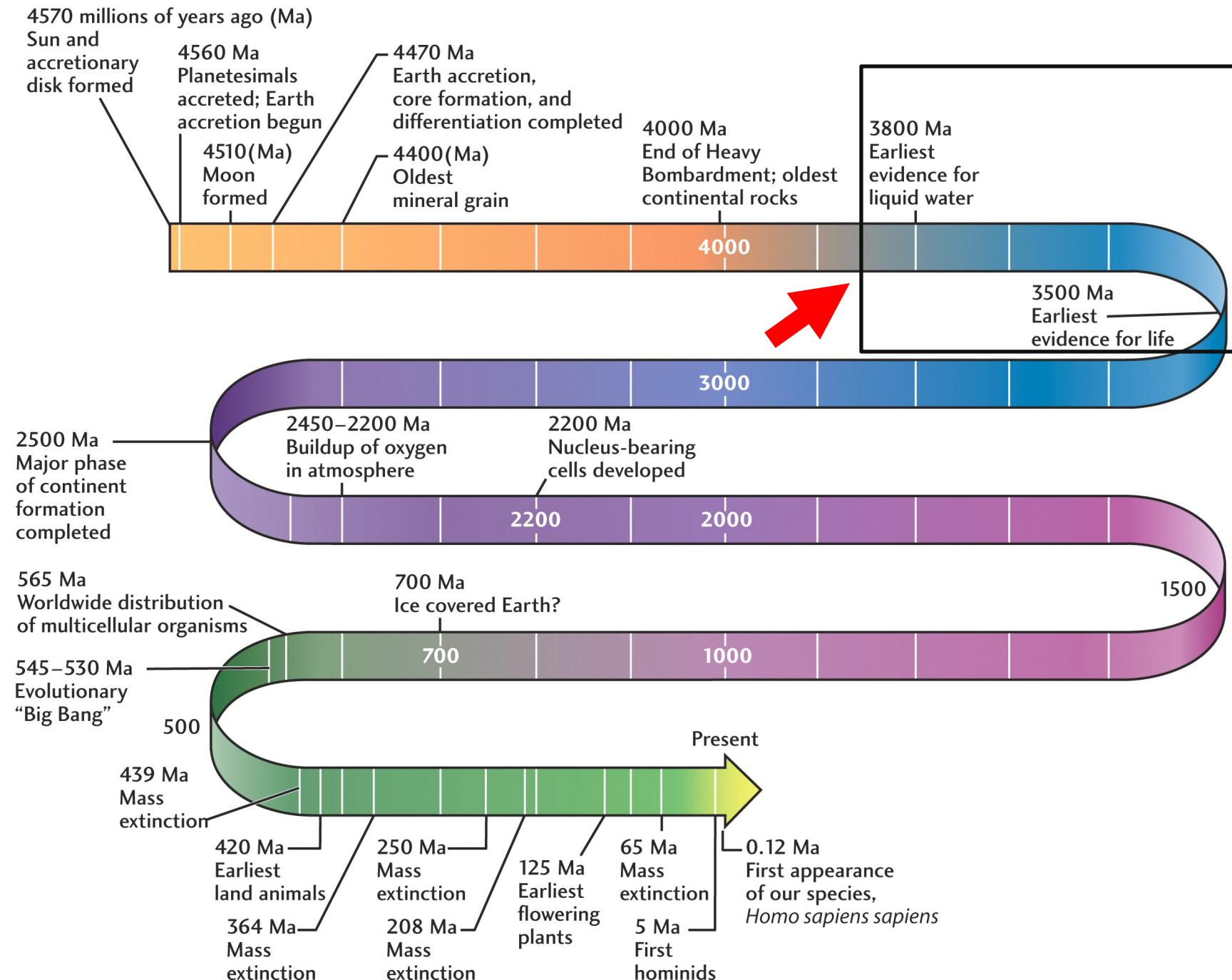
As rocks undergo collision during accretion phase and produce heat causing melting and eventually forms different layers

Denser material sinks to centre of the planet and lighter ones rises to surface

# Global Chemical Differentiation

This global chemical differential was completed by ~4.3 billion years ago, and the Earth had developed a inner and outer core, a mantle and crust





# The Evolving Atmosphere

Right after its creation, the Earth is thought to have had a thin atmosphere composed primarily of helium (He) and hydrogen (H) gases



The Earth's gravity could not hold these light gases and they easily escaped into outer space

Today, H and He are very rare in our atmosphere

# The Evolving Atmosphere



For the next several hundred million years, volcanic out-gassing began to create a thicker atmosphere composed of a wide variety of gases mostly composed of  $\text{CO}_2$ ,  $\text{H}_2\text{O}$  and  $\text{NH}_3$ . The released gases were probably similar to those created by modern volcanic eruptions.

# Creating the Oceans



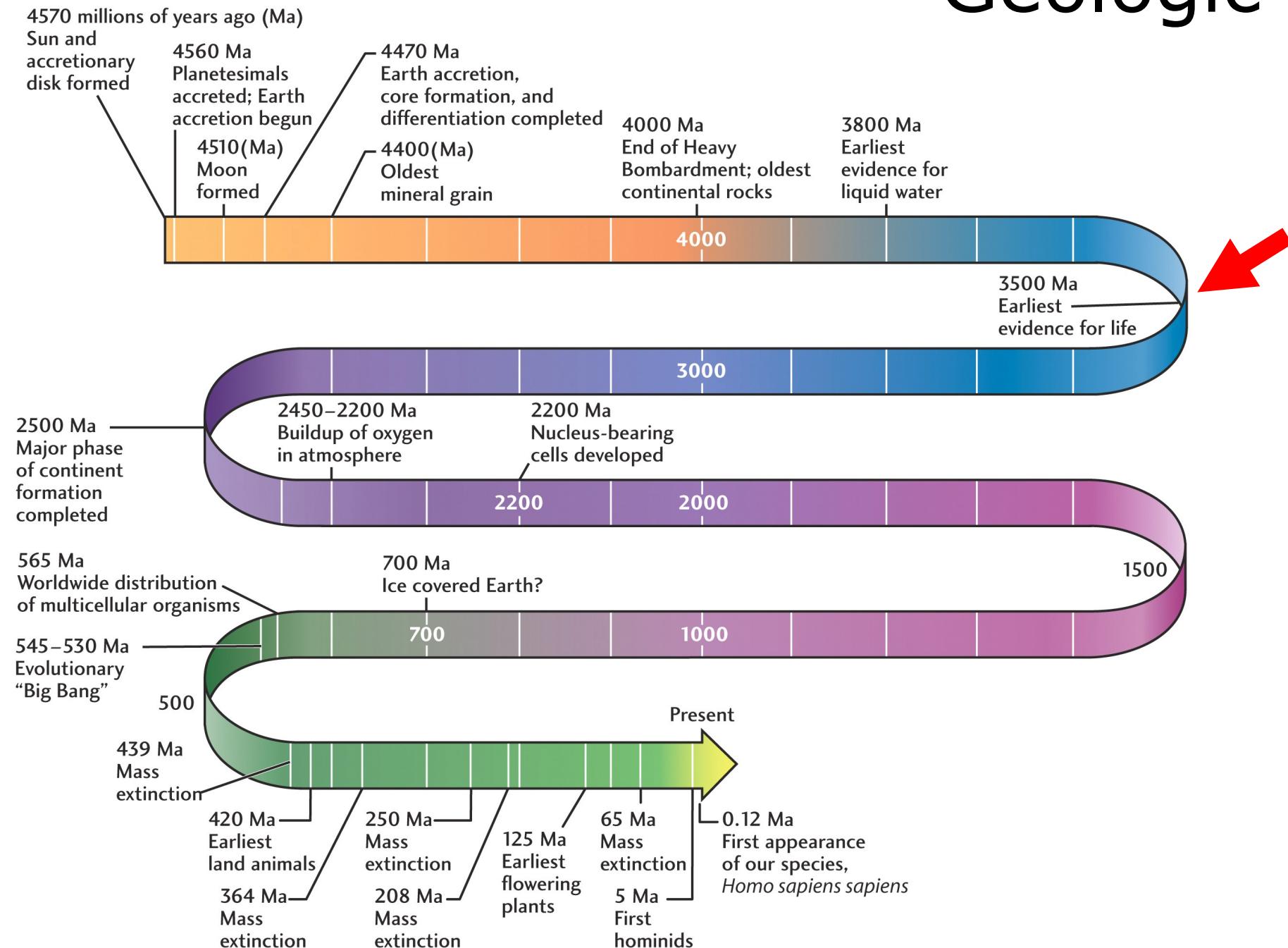
By 3.5 billion years ago, when the Earth was a billion years old, it had a thick atmosphere composed of CO<sub>2</sub>, methane, water vapor and other volcanic gases

Earth cooled with time.

It is hypothesized that water vapor escaping from the interior of the Earth via countless volcanic eruptions produced rain for millions of years and eventually created oceans (this took hundreds of millions of years)

But O<sub>2</sub> is still Missing!

# Geologic Time



# Cyanobacteria

Cyanobacteria, commonly called blue-green algae, is a phylum of bacteria that obtain their energy through photosynthesis



This was the first life on Earth



CO<sub>2</sub> and N<sub>2</sub> dissolved into Oceans and Cyanobacteria developed in oceans which started photosynthesis using sunlight, H<sub>2</sub>O and CO<sub>2</sub>.

# Evidence of Early Life

The 3.5 billion year old fossilized algae mats, which are called stromatolites, are considered to be the earliest known life on earth. Stromatolites are formed in shallow seas or lagoons when millions of cyanobacteria (a primitive type of bacteria) live together in a colony.



They are found in 'Shark bay' lagoon in Western Australia



Stromatolites are mound like bio-accretionary structures. Microbial mats by Cyanobacteria developed on top surfaces of stromatolitic deposits and then was trapped by cycle of sedimentation. Eventually another layer of microbial mat developed on top of the earlier layer and again was trapped by sedimentation. Thus stromatolites grew preserving history.

# Banded Iron Formations (BIF)

How do we know that there was no oxygen in the early Earth atmosphere?

It is hypothesized that the banded iron layers were formed in sea water as the result of free oxygen released by photosynthetic cyanobacteria combining with dissolved iron in the oceans to form insoluble iron oxides, which precipitated out, forming a thin layer on the seafloor

The structures consist of repeated thin layers of iron oxides, either magnetite or hematite, alternating with bands of iron-poor shale and chert



# Banded Iron Formations (BIF)



BIFs are primarily found in very old sedimentary rocks, ranging from over 3 to 1.8 billion years in age

Iron rich basalt leached iron into acidic ocean formed due to dissolved CO<sub>2</sub>. Photosynthesis of Cyanobacteria produced free oxygen which oxidized iron (Fe<sup>+2</sup> and Fe<sup>+3</sup>) to produce red iron (III) oxide (Fe<sub>2</sub>O<sub>3</sub>) on the ocean floor

# Banded Iron Formations



Water must have been replenished with fresh, Iron enriched ocean water cyclically, leading to fresh deposition of BIF

BIF indicates seasonality with iron-rich layers and also indicates possibility of a period of more sunlight availability with increasing photosynthesis, alternating with iron poor layers with low availability of sunlight with less photosynthesis

# Oxygen Evolution in the Atmosphere

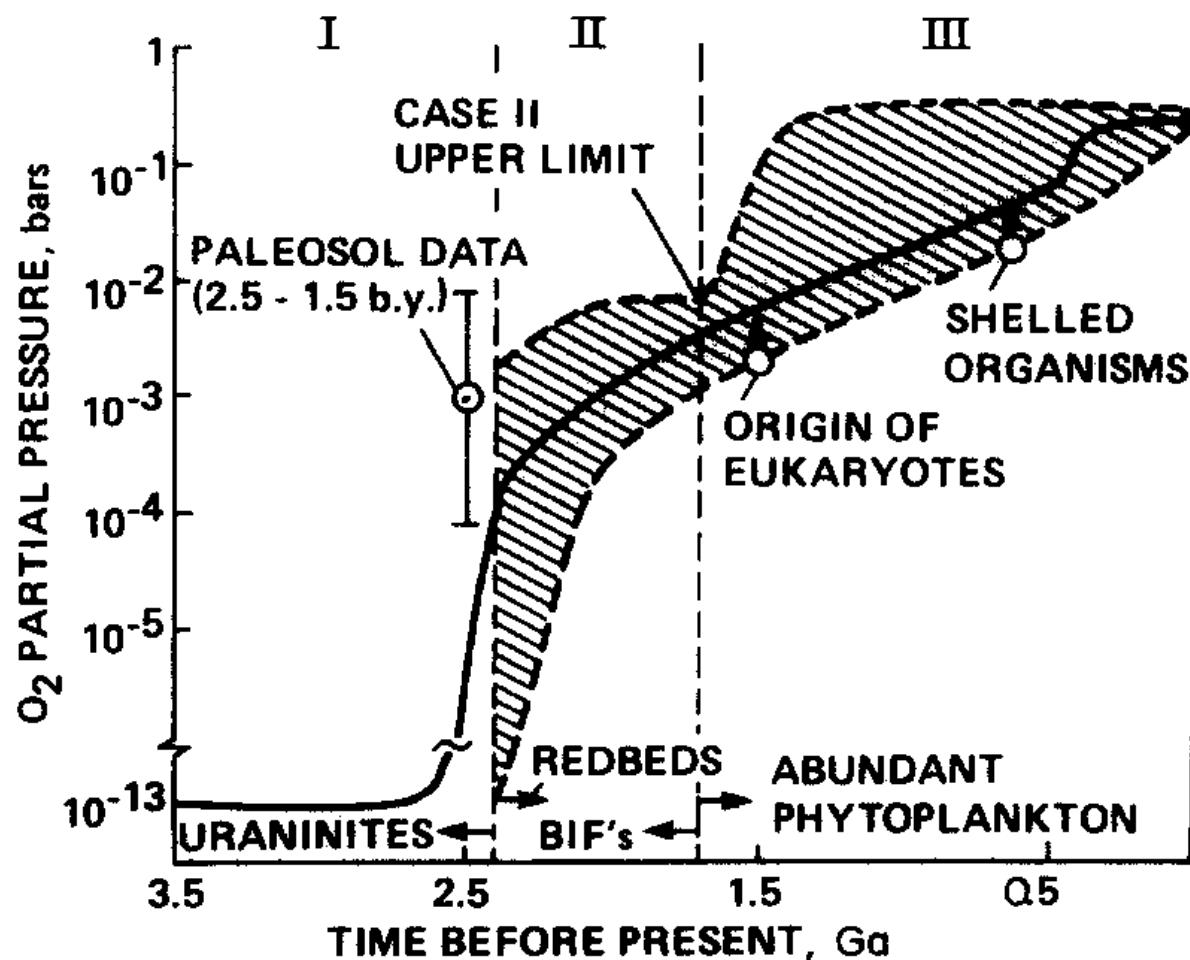


Fig. 7. Estimated change in  $pO_2$  over geologic time. Again, the solid curve is a best guess, and the shaded area represents the range of uncertainty. The point labelled 'paleosol data' is derived from the results of Holland and Zbinden (1986), assuming that  $pCO_2 = 0.05$  bar.

Cyanobacteria were considered to be the source of oxygen and was responsible for BIF precipitation between 3.5-1.8 Ba.

Free O<sub>2</sub> released into ocean participated in Fe oxidation until all Fe was depleted

Increasing O<sub>2</sub> production eventually released into atmosphere and the environment shifted to irreversible oxidizing state from a reducing state

