

## Ensemble Kalman filter parameter estimation of ocean optical properties for reduced biases in a coupled general circulation model

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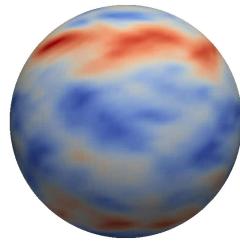
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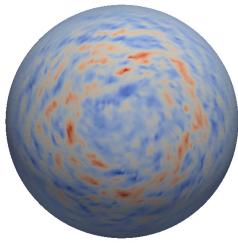


# Problem: Model Biases Limiting Predictability

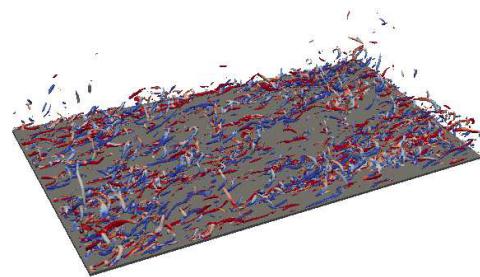
- In GCMs unresolved subgrid processes are parameterised. Many parameters are known with little precision, contributing to **model bias**.
- E.g. ENSO dynamics more difficult to represent in coupled GCMs with parameterised **air-sea fluxes** as opposed to prescribed BCs in O/A GCMs.
- **Typical approach:** Run various simulations, with manually tuned parameter values, and determine which produces the most realistic results.
- **Potential solution:** Calculate subgrid parameters from higher resolution *truth* simulation, such that lower resolution model matches the *true* stats.



Kitsios et. al. (2012, 2019, JAS)



Kitsios et. al. (2013, OM)

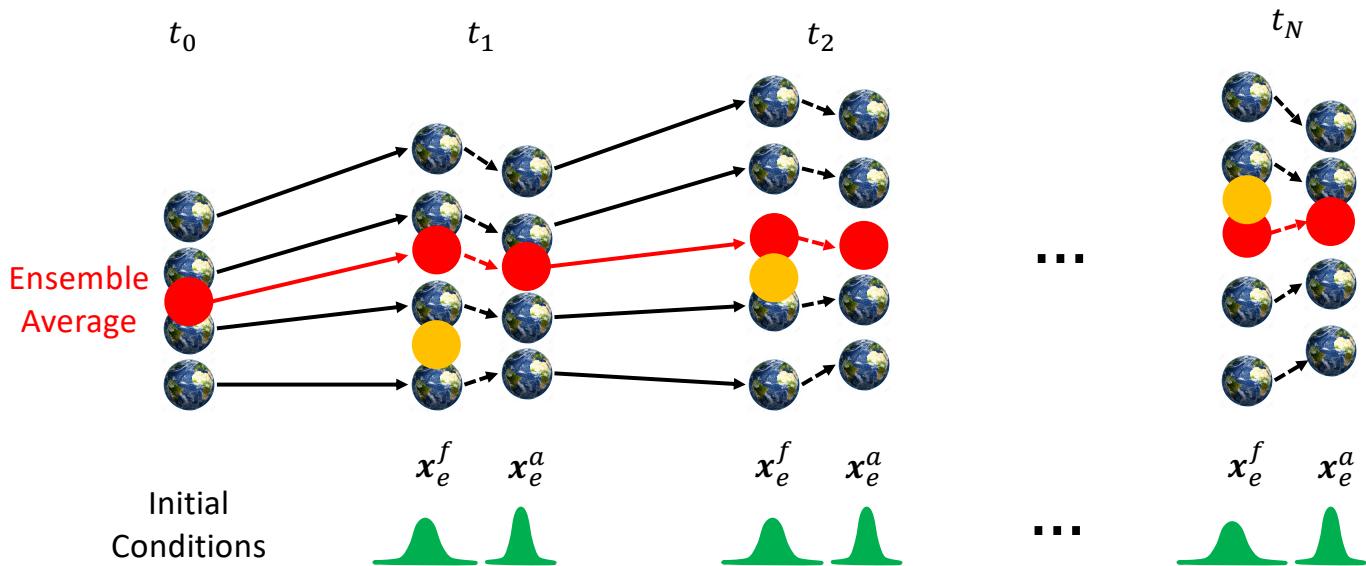


Kitsios et. al. (2017, C&F)

- What if a *truth* case is unavailable, or too computationally expensive?
- **Present approach:** Use the Ensemble Transform Kalman Filter, to estimate parameters governing the air-sea fluxes - specifically the ocean albedo ( $\alpha$ ), and the shortwave e-folding length scale in the ocean ( $L_{SW}$ ).

# Illustration of the Ensemble Kalman Filter

- Data assimilation (DA) modifies imperfect simulations of reality with a series of incomplete and possibly noisy measurements ( $y$ ), ideally resulting in a better representation of the true system state ( $x$ ).



$$\mathbf{x}_e^a(t) = \mathbf{x}_e^f(t) + \mathbf{K} [\mathbf{y}_e(t) - \mathbf{H}\mathbf{x}_e^f(t)] \text{ for ensemble member } e \text{ where}$$

$$\mathbf{K} = \mathbf{P}^f \mathbf{H}^T [\mathbf{H} \mathbf{P}^f \mathbf{H}^T + \mathbf{R}]^{-1} \text{ with observational operator } \mathbf{H},$$

observational error covariance  $\mathbf{R}$ , and state error covariance  $\mathbf{P}^f$ .

# Climate re-Analysis & Forecast Ensemble (CAFE) system

## Data Assimilation System:

- ETKF designed for high dimensional geophysical models (enkf-c).
- **96 ensemble members** (i.e. climate models).
- $\approx 10^8$  model states, with many parameters.
- $\approx 10^4$  CPUs required to simulate the GCM ensemble and solve EnKF.

## Observations:

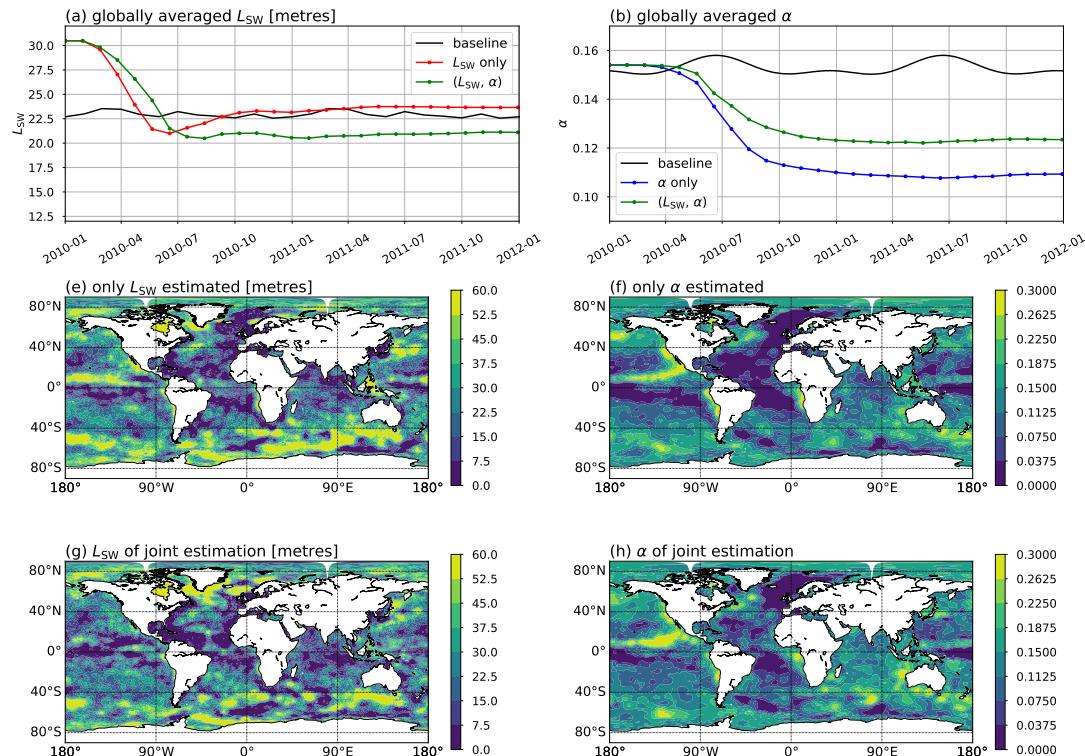
- Satellite and sub-surface observations of the ocean
- JRA55 reanalysis treated as atmospheric observations
- Satellite sea-ice concentration observations from Norwegian and Danish Meteorological Institutes.

## General Circulation Model:

- Modified version of the Geophysical Fluid Dynamics Laboratory Coupled Model version 2.1
- Atmosphere:  $2^\circ$  latitude ;  $2.5^\circ$  longitude ; 24 vertical levels,
- Ocean:  $1^\circ$  longitude; finer latitudinal in specific regions; 50 vertical levels.
- Sea-ice: ocean horizontal resolution; 5 ice thickness categories.

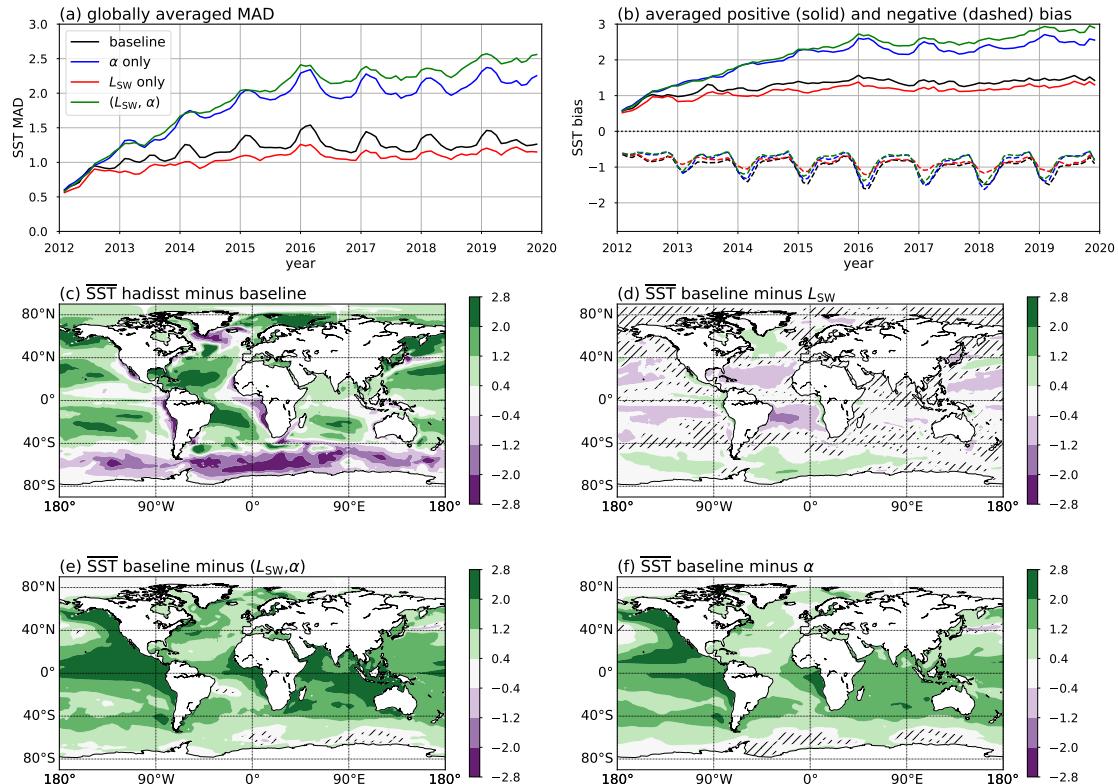
# Spatio-temporal Properties of the Parameters

- DA undertaken from 1/1/2010 to 31/12/2011, on a 28-day cycle, with parameter estimation cases improving fit to observations.
- Parameters have a weak dependence upon time, with distribution more representative of biases than ocean optical properties.



# Out-of-sample Multi-year Monthly Average SST Bias

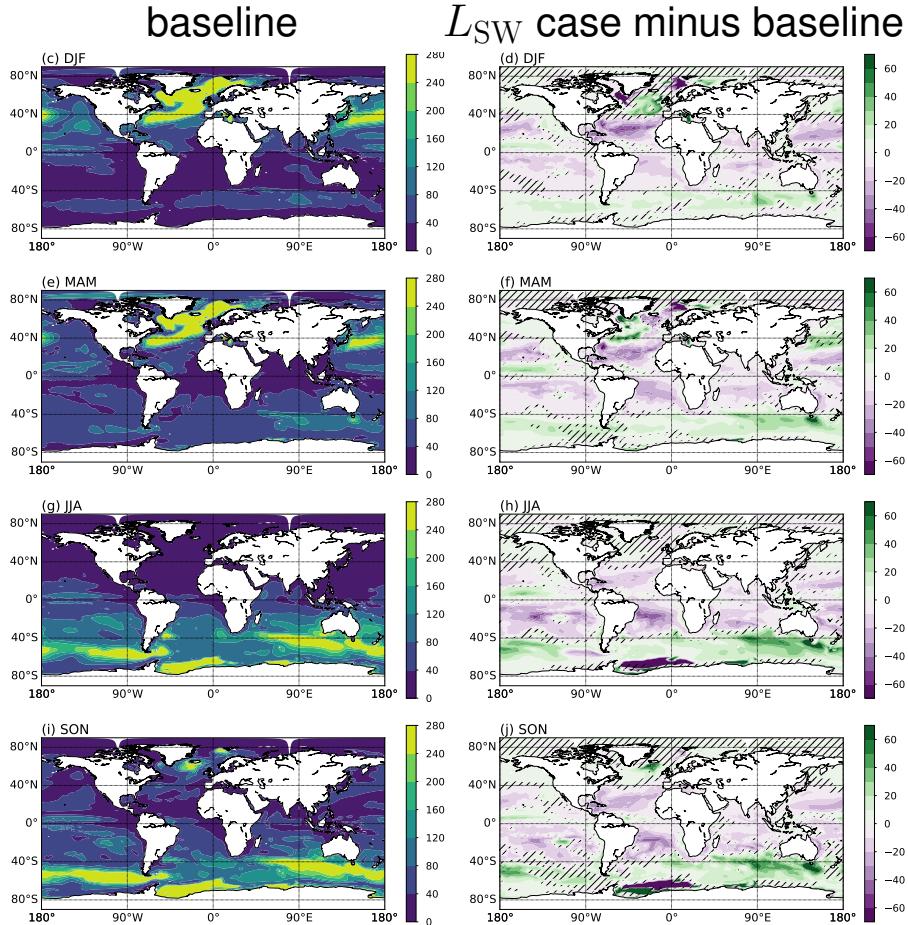
- Only  $L_{SW}$  case has reduced globally averaged mean absolute deviation (MAD,  $|y - H\langle x^f \rangle|$ ) and positive and negative SST bias ( $y - H\langle x^f \rangle$ ).
- Also true for sea-ice concentration and atmospheric jet position.



## Concluding Remarks

- Ensemble Transform Kalman filter used to estimate spatio-temporally varying parameter maps of **ocean albedo** and the **ocean shortwave e-folding length scale** ( $L_{SW}$ ) on the basis of short term 28-days forecasts of a coupled climate atmosphere / ocean / sea-ice GCM.
- Estimated parameter maps in these data assimilation experiments resemble known model biases, and **improve the in-sample fit of the coupled GCM** to a network of real world observations during 2010-2011 (see additional slides).
- However, only the multi-year forecast using the individually estimated map of  $L_{SW}$  systematically **improves out-of-sample skill** during 2012-2020.
- To our knowledge, this is the first attempt at undertaking parameter estimation on GCMs of this size and complexity.
- Whilst the parameter maps appear aphysical, there is a precedent for aphysical parameters in turbulence modelling (e.g. anisotropic and at times negative eddy viscosities),  $L_{SW}$  and  $\alpha$  presumably not optimal, nor perhaps even an appropriate, selection of parameters.
- In current work we are estimating turbulence mixing parameters.

# Changes to ocean mixed layer depth (MLD) per season



- MLD is seasonal because winter sea-ice formation increases salinity and hence density (left column).
- For  $L_{SW}$  case, parameters are time independent, but influence on ocean structure with respect to the baseline is seasonal (right column).
- Low (high) values of  $L_{SW}$  produce shallower (deeper) MLD - qualitatively similar to  $L_{SW}$  map.

# Parameterisation of the Shortwave Flux

- Initial tests indicated that of the air-sea flux parameters, those governing the shortwave flux ( $F_{SW}$ ) exhibited the greatest influence.
- Estimating parameters of the other fluxes were found to be either ineffective, or produced numerical instabilities - indicative of model form error.

## Representation of shortwave radiation:

- In the atmosphere solved using an energy balance model.
- The air-sea flux at the interface governed by

$$F^{SW} = F_0^{SW}(1 - \alpha), \text{ where ocean albedo } \alpha(\lambda, \phi, t) = \frac{0.037}{1.1\theta(\lambda, \phi, t)^{1.4} + 0.15}$$

and  $\theta$  is the zenith angle of the Sun.

- Shortwave flux at a given depth  $z$  of the ocean given by

$$\frac{F^{SW}(\lambda, \phi, z, t)}{F^{SW}(\lambda, \phi, z = 0, t)} = \exp\left(\frac{-z}{L_{SW}(\lambda, \phi, t)}\right) \text{ where } L_{SW} \text{ is an e-folding length}$$

- Here we estimated  $\alpha$  and  $L_{SW}$ , simultaneously and individually.

# Ensemble Kalman Filter (EnKF) of Evensen (1996)

- Data assimilation (DA) modifies imperfect simulations of reality with a series of incomplete and possibly noisy measurements ( $\mathbf{y}$ ), ideally resulting in a better representation of the true system state ( $\mathbf{x}$ ).

**System structure:** Non-linear with Gaussian process and observational noise.

$$\mathbf{x}(t) = \mathcal{F}(\mathbf{x}(t-1)) + \mathbf{w}(t), \text{ where } \mathbf{w}(t) \sim \mathcal{N}(\mathbf{0}, \mathbf{Q}(t))$$

$$\mathbf{y}(t) = \mathcal{H}(\mathbf{x}(t)) + \mathbf{v}(t), \text{ where } \mathbf{v}(t) \sim \mathcal{N}(\mathbf{0}, \mathbf{R}(t))$$

**EnKF Forecast step:** The ensemble of forecast states ( $\mathbf{x}_e^f$ ) evolved in time and associated error covariance ( $\mathbf{P}^f$ ) is sampled

$$\mathbf{x}_e^f(t) = \mathcal{F}(\mathbf{x}_e^a(t-1))$$

$$\mathbf{P}^f(t) = \mathbf{A}\mathbf{A}^T/(m-1), \text{ columns of } \mathbf{A} \text{ contains } \mathbf{x}_e^f - \langle \mathbf{x}^f \rangle$$

with  $m$  the number of ensemble members, and  $\langle \cdot \rangle$  the ensemble average.

**EnKF Analysis step:** Sub-optimally minimises  $\text{trace}(\mathbf{P}^a)$ , where  $\mathbf{P}^a$  is the error covariance of the analysed / corrected state ( $\mathbf{x}_e^a$ ).

$$\mathbf{x}_e^a(t) = \mathbf{x}_e^f(t) + \mathbf{K} [\mathbf{y}_e(t) - \mathcal{H}(\mathbf{x}_e^f(t))] \text{ for ensemble member } e$$

$$\mathbf{K} = \mathbf{P}^f \mathbf{H}^T [\mathbf{H} \mathbf{P}^f \mathbf{H}^T + \mathbf{R}]^{-1}, \text{ where } \mathbf{H} \text{ tangent linear operators of } \mathcal{H}$$

# Ensemble Kalman Filter for Parameter Estimation

- Redefine the state vector to include both the model states ( $\mathbf{x}_S^f$ ) and parameters ( $\mathbf{x}_\psi^f$ ), and associated observational operator (i.e. a satellite cannot observe a model parameter)

$$\mathbf{x}_e^f \equiv \begin{bmatrix} \mathbf{x}_S^f \\ \mathbf{x}_\psi^f \end{bmatrix}_e, \text{ and } \mathbf{H} = [\mathbf{H}_S \ \mathbf{0}]$$

- Substituting above redefinitions

$$\langle \mathbf{x}^a \rangle = \langle \mathbf{x}^f \rangle + \mathbf{K} \left[ \mathbf{y} - \mathbf{H}_S \langle \mathbf{x}_S^f \rangle \right], \text{ with } \mathbf{K} = \begin{bmatrix} \mathbf{P}_{SS}^f \mathbf{H}_S^T \\ \mathbf{P}_{\psi S}^f \mathbf{H}_S^T \end{bmatrix} \left[ \mathbf{H}_S \mathbf{P}_{SS}^f \mathbf{H}_S^T + \mathbf{R} \right]^{-1}$$

- Kalman gain for the states is unchanged, but gain for the parameters is proportional to the covariance between the states and parameters  $\mathbf{P}_{\psi S}^f$ .
- For parameter estimation, the choice of forecast window ( $t_1 - t_0$ ) is not obvious. In general, as the forecast window increases,  $\mathbf{P}_{\psi S}^f$  decreases, associated  $\mathbf{K}$  decreases, but the bias  $\mathbf{y} - \mathbf{H}_S \langle \mathbf{x}_S^f \rangle$ , increases.

## In-sample DA error metrics

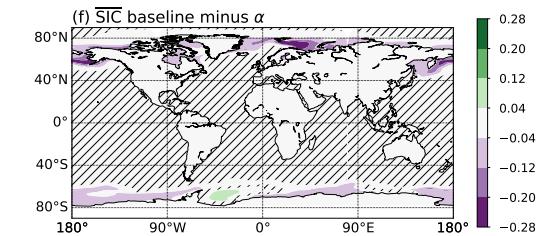
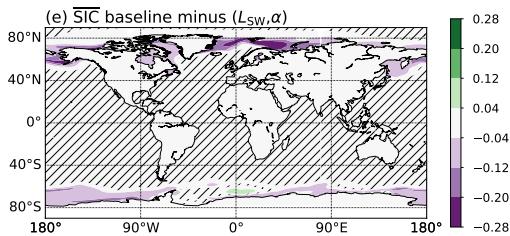
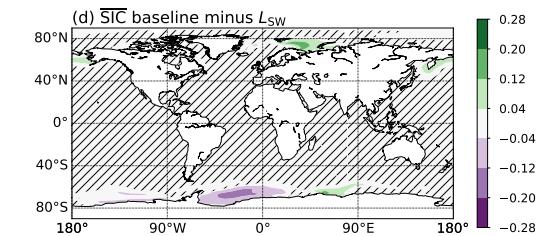
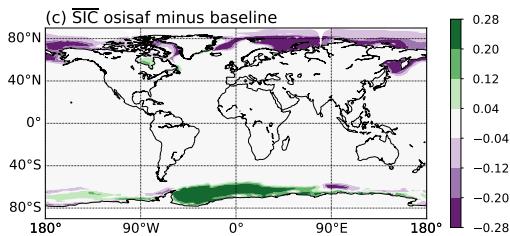
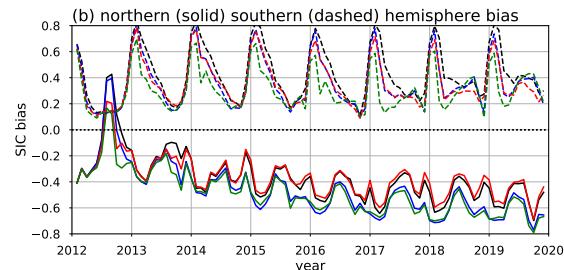
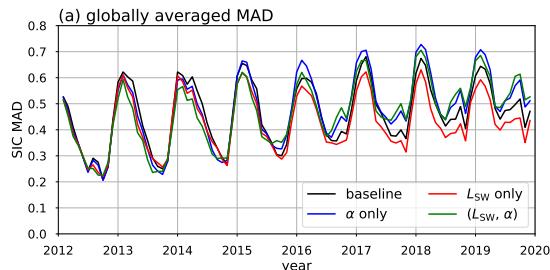
- Globally averaged mean absolute deviation (MAD,  $|y - \mathbf{H}\langle x^f \rangle|$ ) and bias ( $y - \mathbf{H}\langle x^f \rangle$ ) averaged over the second year of the DA experiments to remove “spin up” period.
- Note, the error is evaluated prior to the analysis step, therefore, it is a true measure of the forecast error.
- For selected key observations the estimated cases exhibit reduced in-sample error for the bold error measures.

		baseline	$L_{SW}$ only	$\alpha$ only	$(L_{SW}, \alpha)$
sea-ice concentration ( $m^2/m^2$ )	MAD	0.077	<b>0.071</b>	<b>0.075</b>	<b>0.072</b>
	bias	0.024	<b>0.021</b>	<b>0.022</b>	<b>0.017</b>
sea surface temperature ( $^{\circ}\text{K}$ )	MAD	0.643	<b>0.634</b>	<b>0.634</b>	<b>0.634</b>
	bias	-0.022	<b>-0.014</b>	0.083	0.110
sub-surface ocean temperature ( $^{\circ}\text{K}$ )	MAD	0.708	0.711	<b>0.702</b>	<b>0.694</b>
	bias	-0.106	<b>-0.098</b>	<b>-0.070</b>	<b>-0.041</b>

- However, since state and parameter estimation are intrinsically linked, one must assess the parameters in the freely running model during an out-of-sample period (i.e. 2012 to 2020).

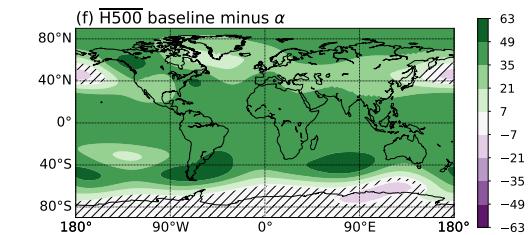
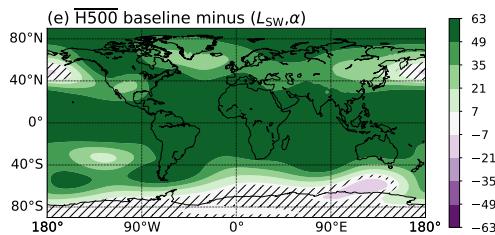
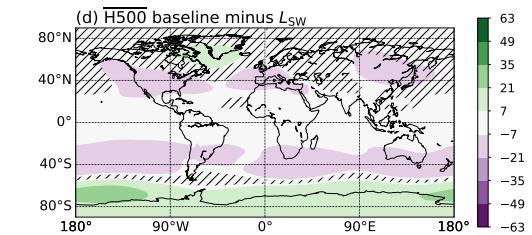
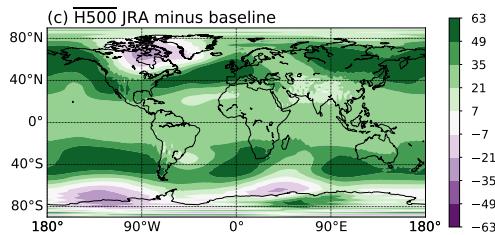
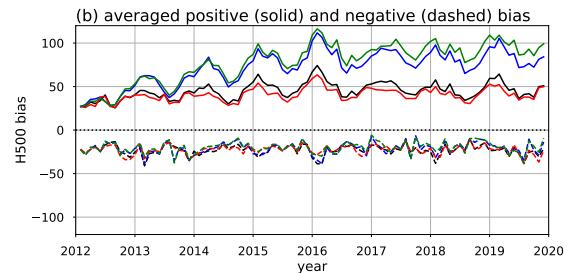
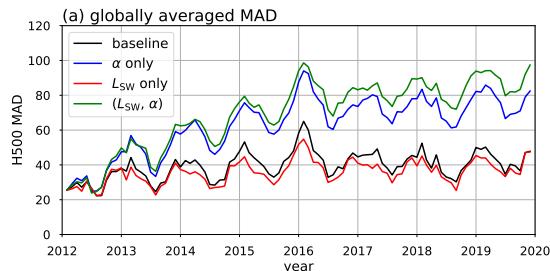
# Out-of-sample Multi-year Monthly Average SIC Bias

- Only  $L_{SW}$  case has reduced MAD, and bias in both hemispheres.



# Out-of-sample Multi-year Monthly Average H500 Bias

- Only  $L_{SW}$  case has systematically reduced MAD, and H500 bias.



# Thank You

**CSIRO Oceans and Atmosphere**

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