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Friction particles as they move **downstream**. Hence, when

$d\mathbf{V}/dt$ is large, its scalar magnitude can be approximated by the centripetal acceleration V^2/R_T , where R_T is the local radius of curvature of the air trajectories.⁸ Hence, the horizontal equation of motion reduces to the balance of forces in the direction transverse to the flow, i.e.,

$$\frac{V^2}{R_T} = -\nabla\Phi - f\mathbf{k} \times \mathbf{V} \quad (7.16)$$

$$\frac{V^2}{R_T} = -\nabla\Phi - f\mathbf{k} \times \mathbf{V}$$

$$\frac{\partial u}{\partial t} + \frac{\partial u}{\partial x} u + \frac{\partial u}{\partial y} v = -\frac{\partial \Phi}{\partial x} + fv$$

7.2.7 T

The signs of the terms in this three-way balance depend on whether the curvature of the trajectories is cyclonic or anticyclonic, as illustrated in Fig. 7.12. In the cyclonic case, the outward centrifugal force is **positive** there. In the mirror image of the centripetal acceleration (**centrifugal force**) the right reinforcing Coriolis force so that a balance can be

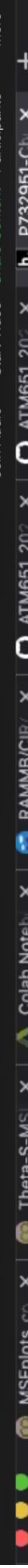
Just as the ship to $\nabla\Phi$ bears a geostrophic pressure surface

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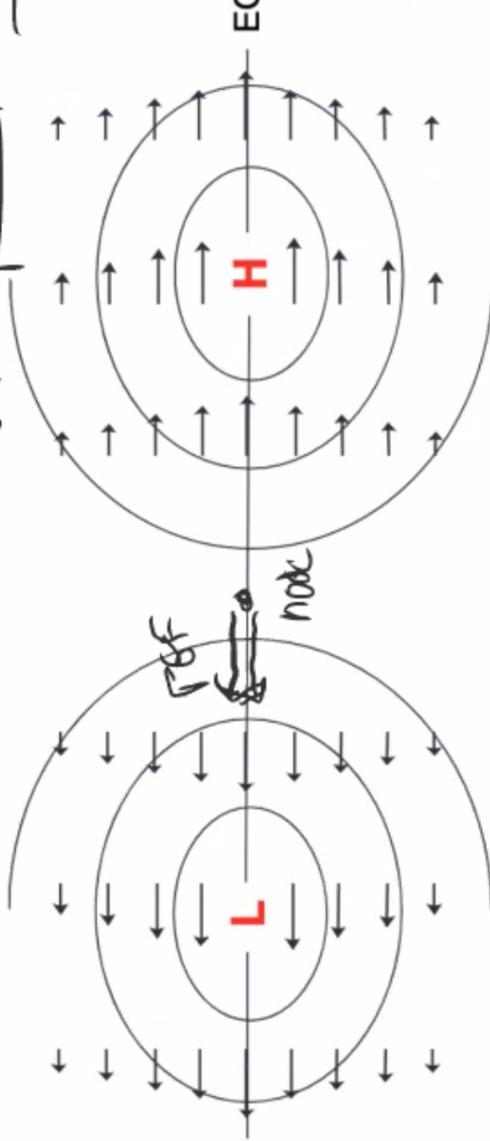
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240%

Sinusoidal) (wave) "Propagation" motion
of a pattern not due to "transport".

$$\frac{\partial u}{\partial t} = -\frac{\partial \Phi}{\partial x} + f_{z0}$$



At node point, ∇F is westward and unopposed.
Therefore, $\frac{\partial u}{\partial t} < 0$.
The wind becomes more like the part of the pattern to its west.
Thus, pattern propagates eastward.

Fig. 7.31 Distribution of wind and geopotential height on a pressure surface in an equatorial Kelvin wave in a coordinate system moving with the wave.

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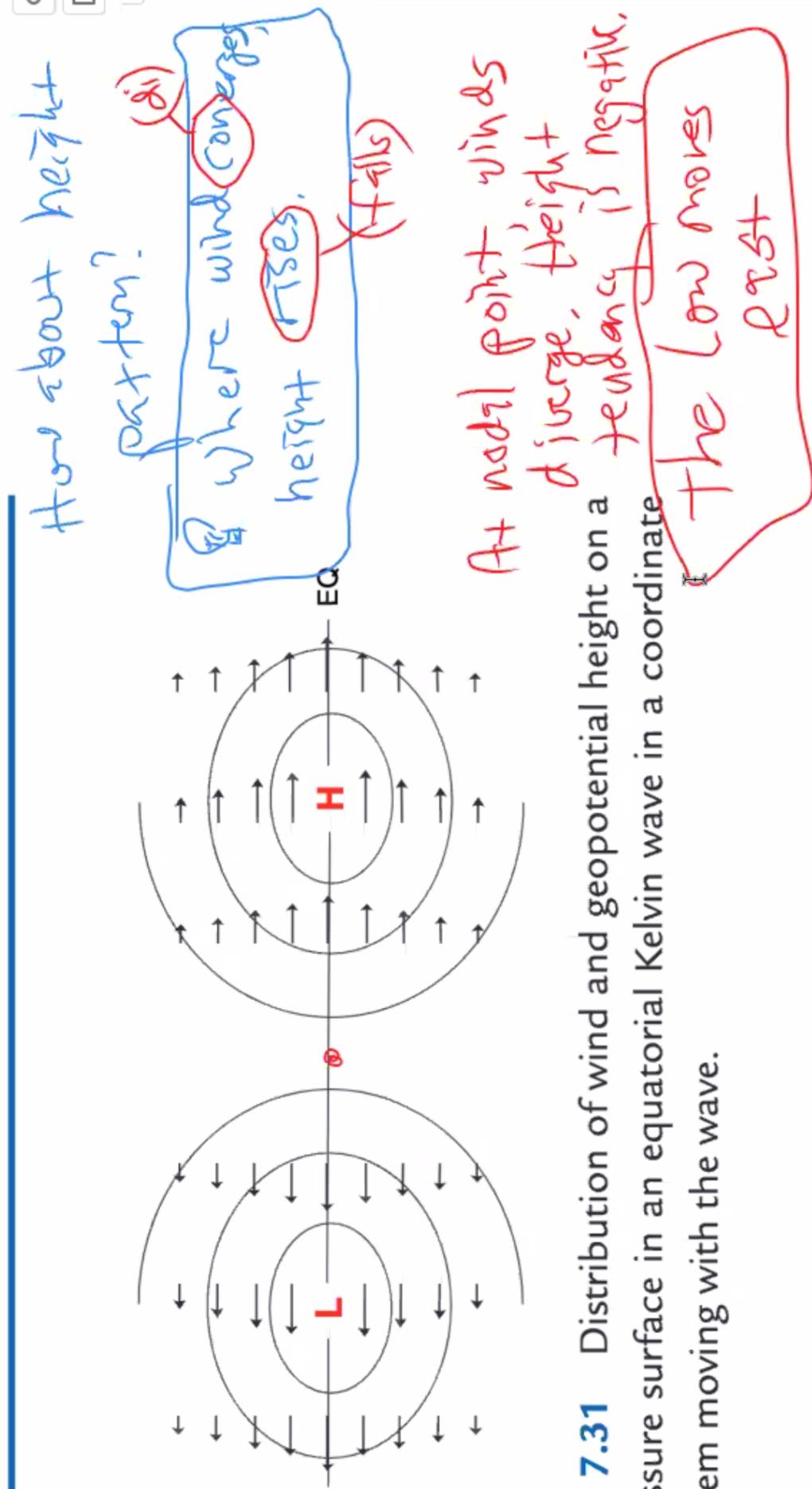


Fig. 7.31 Distribution of wind and geopotential height on a pressure surface in an equatorial Kelvin wave in a coordinate system moving with the wave.



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At a certain station, the 1000-hPa geostrophic wind is blowing from the northeast (050°) at 10 m s^{-1} while the 700-hPa geostrophic wind is blowing from the west (270°) at 30 m s^{-1} . Subsidence is producing adiabatic warming at a rate of $3 \text{ }^\circ\text{C day}^{-1}$ in the 1000 to 700-hPa layer and diabatic heating is negligible.

$$\frac{\partial T}{\partial t} = \text{hor. adv. + (subsidence)} + \text{diabatic heating}$$

3°C/day

$\nabla \cdot \vec{V} T$ need, from $\vec{V} T$

$$\nabla T \cdot \vec{V} T = 0$$

