

# Assessment of GHR SST and OISST Datasets in Identification of Marine Heat-Waves and Heat-Spikes

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**Abstract**—This study assesses the performance of two satellite-derived products, viz. Group for High-Resolution Sea-Surface Temperature (GHR SST) and Optimum Interpolation Sea-Surface Temperature (OISST), in detecting the Marine Heat-Waves (MHWs) and Marine Heat-Spikes (MHSs) after a detailed validation with in situ instruments in the Bay of Bengal (BoB). A significantly good correlation and low Root Mean Square Error (RMSE) between GHR SST (OISST) and the Research Moored Array for African-Asian-Australian Monsoon Analysis and Prediction buoy at 15°N and 90°E (RM15) were found to be 0.97 (0.96) and 0.30 (0.35)°C, respectively. In comparison with three near-coast Wave-Rider-Buoys (WRBs), GHR SST (OISST) shows limited ability to reproduce daily temperature variability with RMSEs ranging from 0.61 °C to 1.07 °C (0.51 °C–1.08 °C), respectively. GHR SST captures signals below 1–2 month periodicity better than OISST, however, both products efficiently capture low-frequency variability (> 4 months). Finally, regarding the number of MHW counts and duration, the GHR SST (36 and 382 days) and OISST (33 and 294 days) overestimate these quantities with respect to RM15 (25 and 287 days). This results in RM15 having a higher duration per MHW event (11.48 days) than GHR SST (10.61 days) and OISST (8.91 days). GHR SST exhibited its highest counts (duration) of 7 (82 days) during 2020 (2010), whereas RM15 showed only 3 (55 days) counts (duration) during 2020 (2010). Moreover, the mean daily intensities from satellite-derived SSTs (0.86 °C) are comparable to RM15 (0.87 °C). Therefore, our study suggests that the satellite-derived SSTs are suitable to capture MHW characteristics in the open-ocean region of the BoB, but their use in the coastal regions should be judiciously done.

**Index Terms**—Group for High-Resolution Sea-Surface Temperature (GHR SST), Marine Heat-Spikes (MHS), Marine Heat-Waves (MHW), Optimum Interpolation Sea-Surface Temperature (OISST), Sea-Surface Temperature (SST), wavelet analysis.

## I. INTRODUCTION

**M**OST of the heat transfer occurring at the air-sea interface are the manifestation of variability in Sea-Surface Temperature (SST) [1]. SST also defines the stability of the ocean surface by modulating its stratification, affecting the biological productivity of the ocean [2]. Moreover, SST helps in defining the coastal upwelling zones [3], [4]. Hence,

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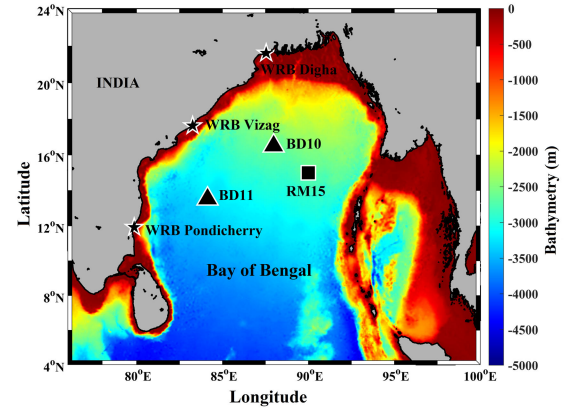


Fig. 1. Map of the BoB with marked locations of RM15 (square), WRBs (star), and OMNI Buoys (triangle). Colormap represents bathymetry in meters.

accurate global SST data, whether from satellite observations or in situ measurements, is vital for understanding physical processes, monitoring climate change, and their assessment.

The Indian Ocean (IO), specifically the Bay of Bengal (BoB), harbors the Indian summer monsoon and cyclones, which are majorly modulated by local SST [5]. It was found that the episodes of anomalously high SST (higher than 90th percentile threshold), regarded as Marine Heat-Waves (MHWs) [6], modulates monsoon characteristics over the Indian mainland [7] and cyclonic activity over the BoB [8]. These MHWs are found to be linked to major climatic modes over the IO [7], [9]. MHWs are the global warming response of oceans and have major implications on climate change [10]. Therefore, to detect MHWs or Marine Heat-Spikes (MHSs) [6], and to understand its impact as an ocean stressor, accurate and precise measurements of SST is required. Along with accurate data, high spatio-temporal resolution is essential to understand the response of MHWs to various spatial-scale processes and its correspondence to the ocean-current system giving a decent idea of the spatial structures of MHWs.

In the era of satellite-based observations, SST measurements are taken in the form of brightness temperature which is then converted to actual temperature using various algorithms [11]. To reduce errors or biases, this data is generally objectively analyzed and optimally interpolated. Examples of such products are the Optimally Interpolated SST (OISST) and Group for High-Resolution SST (GHR SST) which are widely used over the global oceans [12], [13]. These satellite products have been extensively utilized to detect and understand the

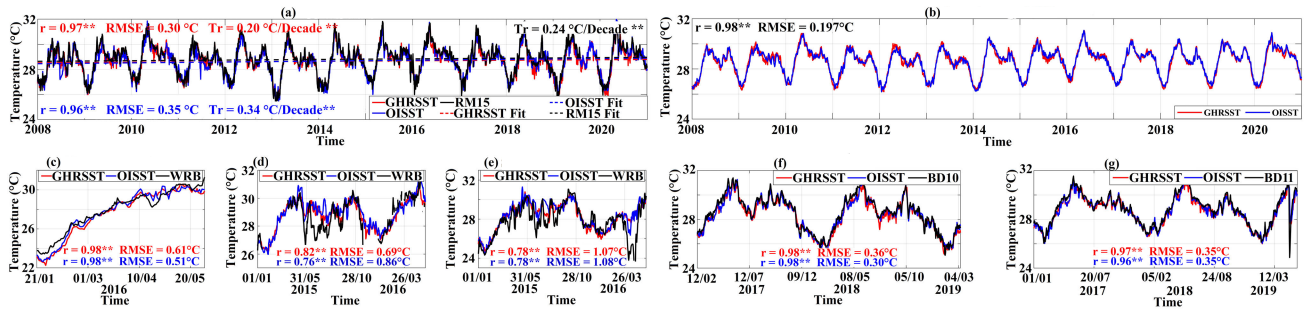


Fig. 2. Timeseries of GHR SST (red) and OISST (blue) against in situ (black). (a) RM15, (c) WRBD, (d) WRBP, (e) WRBV, (f) BD10, and (g) BD11. (a) Dashed red, blue, and black lines represent trends (Tr) with values written in the same color. In (a) and (c)–(g) correlation coefficient values represented by “r” and RMSE between GHR SST and in situ are given in red colored text, whereas blue color is used for OISST and in situ. (b) Represents the area-averaged timeseries of GHR SST and OISST over the BoB, with correlation and RMSE written in black-colored text. “\*\*\*” against the values of r and Tr indicates more than 99% significance.

signatures of the MHWs over different regions of the world [14], [15], including the BoB [7], [9]. Gupta et al. [9] found a good correlation between GHR SST and OISST for the daily intensity of MHWs. However, the studies have not made objective comparisons of the suitability of these satellite products in MHW analysis with respect to in situ observations. Hence, herein we aim to assess the suitability of the GHR SST and OISST in identifying MHWs and MHSs, highlighting the strengths and weaknesses, in comparison to an in situ measurement. A precursor of this comparative analysis is a rigorous validation of the satellite-based SSTs against various in situ observations.

## II. DATA AND METHODOLOGY

### A. Data

The daily GHR SST and OISST datasets for the region of BoB (period: 2008–2020) are distributed by the Physical Oceanography Distributed Active Archive Center (<https://podaac.jpl.nasa.gov/>) and Physical Science Laboratory (<https://psl.noaa.gov/data/gridded/>). Daily SST data at Research Moored Array for African-Asian-Australian Monsoon Analysis and Prediction buoy present at 15°N and 90°E (RM15) [16], for the period of 2008–2020, are obtained from <https://www.pmel.noaa.gov/tao/drupal/disdel/>. 30-min SST data of Wave-Rider-Buoy (WRB) at Digha (WRBD), Pondicherry (WRBP), and Vizag (WRBV) are acquired from the Indian National Center for Ocean Information Services (INCOIS) for the period of 01-Jan-2016 to 31-Mar-2016, 01-Jan-2015 to 30-Mar-2016, and 01-Jan-2015 to 31-Mar-2016, respectively. Hourly SST of BoB Deep (BD) ocean Network for Northern Indian (OMNI) moored buoys, BD10 and BD11, are also obtained from the INCOIS for the period of 12-Feb-2017 to 10-Mar-2019 and 01-Jan-2017 to 27-May-2019, respectively [17]. This study uses the RM15 dataset for MHW identification due to its comparatively long-term availability than any other aforementioned in situ dataset (Fig. 1).

### B. Methodology

For comparing GHR SST and OISST against in situ data, four nearest neighbors about the site of in situ equipment were averaged. The one or two-day missing data for GHR SST and

RM15 has been filled using the linear interpolation method. Outliers have been removed from the WRB data using the median-based outlier detection algorithm present in MATLAB.

To understand the SST variability, comparative wavelet analysis has been performed to highlight the differences in the dominant modes of variability in GHR SST, OISST, and RM15. It is a signal processing technique that decomposes a signal into different frequency components over time [18], [19]. This helps us analyze how the contribution of these components to the total signal varies over time. The wavelet transforms of a 1D-signal (in this case the SST) result in a 2-D pattern of localized power (or variance [20]) in time-frequency space. Relatively higher power of a particular frequency at a specific time indicates a relatively strong presence of that particular frequency component at that time. The patterns outside the Cone of Influence are not considered as they suffer from edge effects [20]. In our study, SST datasets were also filtered for <1, 1–2, and 2–4 months of periodicity range, using the Butterworth-bandpass filter, available in MATLAB.

MHWs are defined as an extreme temperature event where the SST is higher than seasonally varying 90th percentile “threshold” (1) for at least five consecutive days. The discrete heatwave events with a gap of less than 2 days are considered as a single event. Temperatures that are higher than the threshold but do not satisfy the criteria of MHWs are called MHSs [6]. The “threshold” ( $T_{90}$ ; following [9]) has been defined using the daily climatology ( $\mu$ ) and standard deviation ( $\sigma$ ) of SST data computed for the period of 2008–2020 as

$$T_{90} = z \cdot \sigma + \mu \quad (\text{where, } z = 1.282). \quad (1)$$

MHWs have three important attributes, namely, counts—total number of distinct events in a given period of time, duration—total number of days experiencing MHWs conditions in a given period of time, and Daily Intensity (DI)—difference of observed temperature from climatology when MHW or MHS condition is satisfied.

## III. RESULTS AND DISCUSSION

### A. Validation

On comparing GHR SST and OISST with RM15 for 2008–2020, a good correspondence between them is found. A significant ( $p < 0.01$ ) high correlation coefficient ( $r$ ) of 0.97 (0.96) was found between GHR SST (OISST) and RM15,

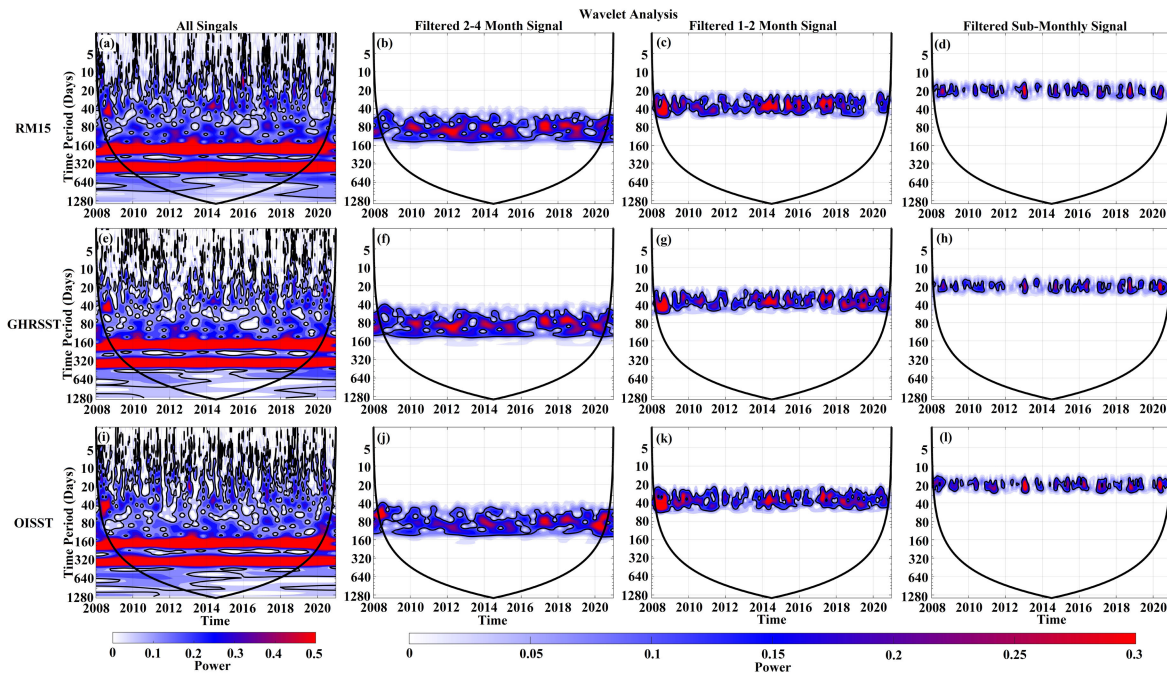


Fig. 3. Wavelet analysis of SST timeseries for (a)–(d) RM15, (e)–(h) GHR SST, and (i)–(l) OISST for all signals, 2–4 months, 1–2 month and sub-monthly scale filtered signals with background representing the power of the signals and solid black curve representing the Cone of Influence.

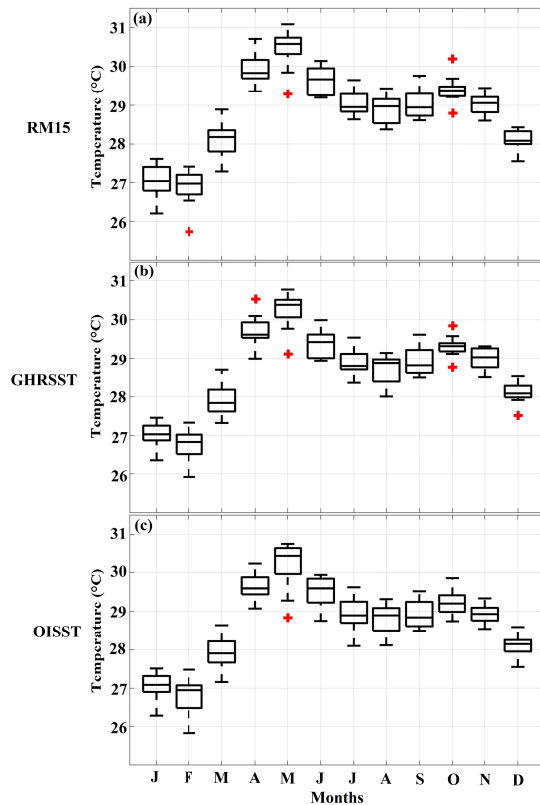


Fig. 4. Box plot of monthly climatology. (a) RM15. (b) GHR SST. (c) OISST.

having a low Root Mean Square Error (RMSE) of  $0.30^{\circ}\text{C}$  ( $0.35^{\circ}\text{C}$ ) [Fig. 2(a)]. The “per decade” trend shown by RM15, GHR SST, and OISST are  $0.24^{\circ}\text{C}$ ,  $0.20^{\circ}\text{C}$ , and  $0.34^{\circ}\text{C}$  ( $p < 0.01$ ), respectively. The area-averaged (for the BoB) timeseries of OISST and GHR SST show a high correlation of  $0.98$  ( $p < 0.01$ ), with an RMSE of  $0.20^{\circ}\text{C}$  [Fig. 2(b)]. Relatively low correlations and high RMSEs are observed near

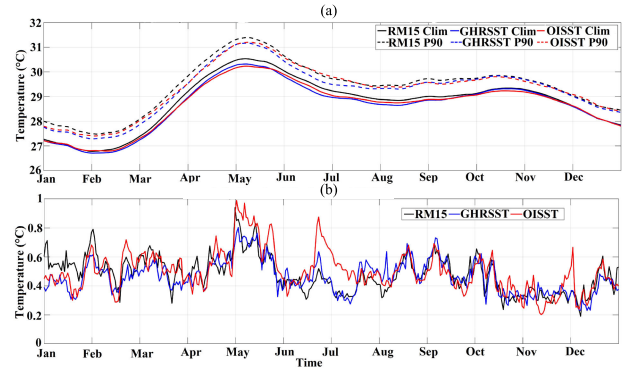


Fig. 5. (a) Timeseries of daily climatology (Clim; solid lines) and 90th percentile threshold (P90; dashed lines) are shown. (b) Time series of daily standard deviation. Black color is used for RM15, red for OISST, and blue for GHR SST.

coastal regions as indicated by WRBD, WRBP, and WRBV against GHR SST and OISST [Fig. 2(c)–(e)]. Again, good correlations and low RMSEs of satellite SST with OMNI buoys (BD10 and BD11) are found [Fig. 2(f) and (g)]. This analysis suggests that GHR SST and OISST perform better in the open ocean compared to coastal regions; therefore, they should be used after proper validation for studying coastal processes.

Wavelet analysis of SST at RM15 [Fig. 3(a)–(d)] are shown in comparison to GHR SST [Fig. 3(e)–(h)] and OISST [Fig. 3(i)–(l)]. For 2–4 month filtered data, GHR SST shows almost a similar pattern as that of RM15 with some underestimation of the power in 2014, whereas OISST deviates both in the pattern and power of the signal, suggesting GHR SST captures processes occurring at a period of 2–4 months more accurately than OISST. For 1–2 months filtered signal, different patterns of dominant signal were observed both in GHR SST and OISST during various years when compared to



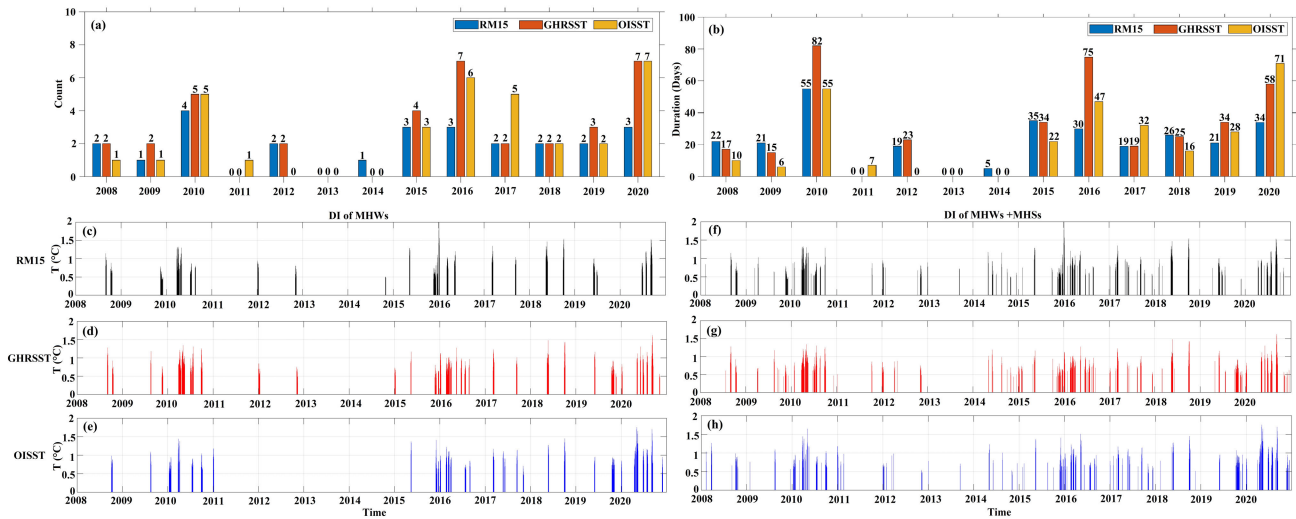


Fig. 6. Bar plot of the number of MHW (a) counts and (b) duration (in days) in a year. Blue bar, red bar, and green bar represent RM15, GHRSSST, and OISST respectively. DI of MHWs detected using (c) RM15, (d) GHRSSST, and (e) OISST. (f)–(h) are similar to (c)–(e) but for DI of MHWs + MHSs respectively.

RM15, with OISST deviating more than GHRSSST. Similarly, in sub-monthly filtered data, GHRSSST shows similar patterns as RM15, but OISST is unable to capture it in some of the years. Signals above 2–4 months periodicity are captured by both GHRSSST and OISST accurately. Overall, this analysis suggests that GHRSSST is more efficient than OISST in capturing signals of different periodicity over the years. Moreover, GHRSSST and OISST might show deviations, with respect to RM15, in identifying MHW characteristics owing to their limitation in capturing sub-monthly signals, as MHWs exist for at least 5-days and are primarily driven by sub-monthly processes [9].

### B. MHW and MHS Identification

Monthly climatology of OISST has the largest spread about the median followed by RM15 and GHRSSST [Fig. 4(a)–(c)], which can also be corroborated by the daily standard deviation [Fig. 5(b)]. To detect the MHW characteristics, it is required to generate daily climatology, standard deviation, and the corresponding threshold [Fig. 5(a) and (b)]. Daily climatology of RM15 is highest followed by OISST and GHRSSST; whereas daily threshold of RM15 is highest for almost all the time except for the period of May–July in which the threshold of OISST dominates. This is attributed to the significantly higher values of the standard deviation of OISST compared to RM15, which are used in the calculation of threshold (1).

The MHWs from RM15, GHRSSST, and OISST are found to be higher in both counts and duration during the years 2010, 2016, and 2020, albeit with some differences in number. RM15 showed MHW counts and duration of 4 and 55 days, 3 and 30 days, and 3 and 34 days in the aforementioned years, whereas GHRSSST (OISST) showed 5 and 82 days (5 and 55 days), 7 and 75 days (6 and 47 days), and 7 and 58 days (7 and 71 days), respectively. Also, these were the years with the most intense MHWs as well [9]. However, OISST in 2020 showed a significantly higher duration of MHWs (71 days) when compared to RM15 (34 days), whereas GHRSSST has a duration (58 days) lying between the OISST

and RM15. On the other hand, GHRSSST showed higher duration in the years 2010 and 2016 when compared to both RM15 and OISST. Additionally, the years 2011, 2013, and 2014 showed inconspicuous or zero MHW events which are more or less consistent with all the datasets. In the year 2012, both GHRSSST and RM15 showed an equal number of counts (two), and similar duration (19 days and 23 days respectively), whereas OISST showed no heatwave event. After 2014, all years have experienced frequent and intense MHWs which is also consistent with Gupta et al. [9].

The yearly statistics of MHWs characteristics from satellite products showed overestimation in counts and duration when compared to RM15 [Fig. 6(a) and (e)]. In contrast, the DI has similar values whenever the MHWs were captured by all three datasets [Fig. 6(b)–(d)]. The total count and duration of MHWs using RM15 were found to be 25 and 287 days respectively for the study period. In contrast, GHRSSST (OISST) overestimates both count and duration with a total of 36 (33) and 382 (294) days respectively. Owing to the values above, the average duration per count comes out to be 11.48 days, 10.61 days, and 8.91 days for the RM15, GHRSSST, and OISST respectively. The statistically significant ( $p < 0.01$ ) correlation of the counts and duration between the RM15 and GHRSSST (OISST) were 0.85 (0.74) and 0.88 (0.74) respectively. The mean DI from RM15, GHRSSST, and OISST were similar with a value of 0.87 °C, 0.82 °C, and 0.90 °C, respectively. The correlations ( $p < 0.01$ ) of the DI of RM15 against GHRSSST and OISST were observed to be 0.72 and 0.57, respectively.

To understand the source of the overestimation of MHW count and duration by GHRSSST and OISST compared to RM15, MHSs have been plotted with MHWs [Fig. 6(f)–(h)]. It was found that the duration (DI) of MHWs and MHSs when taken together are 418 days (0.84 °C), 490 days (0.79 °C), and 413 days (0.88 °C) for RM15, GHRSSST, and OISST respectively, suggesting a reduced difference in MHW + MHS duration. Therefore, the discrepancy in MHW count and duration comes from the fact that these MHSs do not persist longer than 5-days to be categorized as MHWs at RM15,

whereas in GHR SST and OISST, there is sustenance of MHSs for a longer duration resulting in more MHWs.

#### IV. CONCLUSION

This study assesses the applicability of GHR SST and OISST datasets in identifying MHWs in comparison to moored buoy observations in the BoB. A good agreement between the satellite-derived SST and in situ SST has been found to hold in the open ocean region (maximum RMSE of 0.36 °C across all in situ measurements). However, GHR SST and OISST showed higher RMSEs when compared to WRBs present near the coastal regions (minimum RMSE of 0.51 °C across all the WRB locations), and therefore, GHR SST and OISST datasets are not good estimates of in situ temperatures near the coastal region. Hence, they should be used after extensive validation for studying coastal processes. Consequently, these satellite datasets are suggested to be improved with respect to capturing coastal SST variabilities. Moreover, GHR SST and OISST are unable to capture the signals having periodicity <2 months with the same power and frequency as that of RM15; however, GHR SST captures 2–4 months signals more accurately. Wavelet analysis suggested that GHR SST is better than OISST in capturing signals with different frequencies. Moreover, monthly and daily climatology and daily standard deviations for GHR SST, OISST, and RM15 were slightly different from one another resulting in differing 90th percentile thresholds calculated for these datasets. MHWs, which are becoming more frequent and hazardous to aquatic life in the BoB [9], are also being overestimated by GHR SST (count = 36 and duration = 382 days) and OISST (count = 33 and duration = 294 days) due to sustenance of high temperature for longer period than RM15 (count = 25 and duration = 287 days). However, the duration of MHWs and MHSs taken together shows a better match in all the cases showing the efficacy of GHR SST and OISST in detecting heat extremes.

This study showed the applicability of the OISST and GHR SST in accurately identifying the MHS and MHW characteristics in the BoB. Therefore, these datasets can be utilized for understanding the spatial distribution of MHWs along with their formation and dissipation mechanism. This analysis also motivates us to investigate the reasons for the recent increment in the long-sustained MHW events in the region. In addition, the results can further be utilized for validation of any model-simulated SSTs in this region. Furthermore, there is a scope to look into the vertical extension of MHWs in this region using observations or models.

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