

Seismic Tomography: solving inverse problems to estimate model parameter(s)

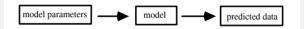
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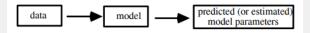
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Forward and Inverse problem



The process of predicting data based on some mathematical (or physical) model with a given set of model parameters. eg: travel time based on reference velocity structure.



The process of predicting (or estimating) the numerical values of a set of model parameters of an assumed model based on observation. eg: invert travel time to determine the velocity structure.





Inverse problem

- We want to find a velocity structure of the Earth from observed seismic data. Given the observed arrival times, find a model that fits them.

$$d = Am$$

 d_i is a data vector (arrival times) and m_j is assumed model vector (eg. slowness or 1/velocity)

- Start with the initial model \mathbf{m}^0 (a guess, eg. 1D Earth velocity model).
- Predictions from starting model

$$d_i^0 = A(m_j^0)$$





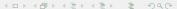
$$d_i^0 = A(m_j^0)$$

- This doesn't predict the actual observation. Hence, we seek to change the initial model by $\Delta \boldsymbol{m}$ amount.

$$m_j^1 = m_j^0 + \Delta m_j^0$$

- Further, our data do not depend linearly on the model parameters.
- Linearize the problem by expanding the data in a Taylor series about the starting model

$$d_i \approx d_i^0 + \sum_j \frac{\partial d}{\partial m}|_{m^0}.\Delta m_j^0$$





$$d_i \approx d_i^0 + \sum_j \frac{\partial d}{\partial m}|_{m^0}.\Delta m_j^0$$

- Let's write this in terms of the difference between the observed data and those predicted,

$$\Delta d_i^0 \equiv d_i - d_i^0 \approx \sum_j \frac{\partial d_i}{\partial m_j}|_{m^0}.\Delta m_j^0$$

 $\Delta \mathbf{d} = \mathbf{G} \Delta \mathbf{m}$

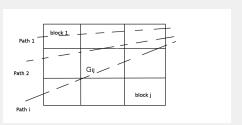
- Where, $\mathbf{G}=G_{ij}=\frac{\partial d_i}{\partial m_j}$ is the partial derivative matrix. This is a linear algebra problem.



- In our seismic tomography example, this matrix would be

$$G_{ij}=rac{\partial T_i}{\partial s_j}$$

 T_i being the travel-time perturbation- i^{th} ray and u_j is the slowness perturbation- j^{th} box



Source: Stein ands Wysession



$$\Delta d = G\Delta m$$

- Because there are many more equations (ray paths) than unknowns (model parameters), the system of equations is overdetermined.
- Also inconsistent because our data contain noise (instrument, measurement related, or unknown velocity structures in earthquake location problem)
- The least-squares problem

$$arg_{min}||\Delta \mathbf{d} - \mathbf{G}\Delta \mathbf{m}||_2$$

has a unique minimum $||\mathbf{r}||_2$ and a unique residual $\mathbf{r} = \Delta \mathbf{d} - \mathbf{G} \Delta \mathbf{m}^*$ iff $\Delta \mathbf{m}^*$ solves the normal equation

$$\mathbf{G}^{\mathsf{T}}\mathbf{G}\Delta\mathbf{m} = \mathbf{G}^{\mathsf{T}}\Delta\mathbf{d}$$





-Then in the second iteration, we use

$$m_{j}^{1} = m_{j}^{0} + \Delta m_{j}^{0}$$
.....
 $m_{j}^{n} = m_{j}^{n-1} + \Delta m_{j}^{n-1}$

- Define a tolerance level and stop iterating.
- In finite precision arithmetics **G**^T**G** may become singular.
- If **G** is dense LAPACK routines (eg. DGELS)
- If **G** is sparse iterative LSQR algorithm



Jupyter notebook

