



## Atmospheric Effects Modelling

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### Fermat's principle :

Fermat's principle, also known as the principle of least time, is the link between ray optics and wave optics. In its original "strong" form, Fermat's principle states that the path taken by a ray between two given points is the path that can be traveled in the least time. In order to be true in all cases, this statement must be weakened by replacing the "least" time with a time that is "stationary" with respect to variations of the path — so that a deviation in the path causes, at most, a second-order change in the traversal time. To put it loosely, a ray path is surrounded by close paths that can be traversed in very close times. It can be shown that this technical definition corresponds to more intuitive notions of a ray, such as a line of sight or the path of a narrow beam. If points A and B are given, a wavefront expanding from A sweeps all possible ray paths radiating from A, whether they pass through B or not. If the wavefront reaches point B, it sweeps not only the ray path(s) from A to B, but also an infinitude of nearby paths with the same endpoints. Fermat's principle describes any ray that happens to reach point B; there is no implication that the ray "knew" the quickest path or "intended" to take that path.

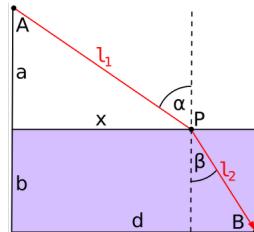


Figure 2: Beam refraction from A to B

since GPS signals are electromagnetic signals so they refract in layers of atmosphere.

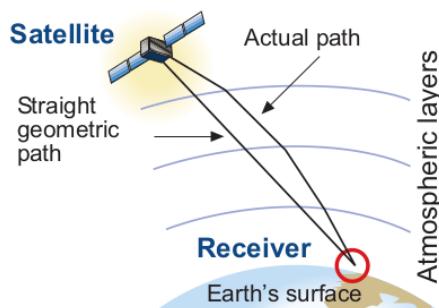


Figure 3: Path of signal in atmosphere layers

Equation of distance is Euclidean, but ray path is not so it is one obligation to correct ray path to Euclidean line.

$$\Delta = \int_{raypath} n dl - \int_{straightline} dl = \int (n - 1) dl$$

There are two parts in atmosphere , Ionosphere and Troposphere which their delays and effects are given as follows :

## 1 Ionospheric Delay

The ionosphere is that part of the terrestrial atmosphere that extends from about 60 km up to more than 2000 km. As its name implies, it contains a partially ionised medium, as a result of solar X- and Extreme UltraViolet (EUV) rays in the solar radiation and the incidence of charged particles. The propagation speed of GNSS electromagnetic signals in the iono- sphere depends on its electron density (see below), which is typically driven by two main processes. During the day, the Sun's radiation ionises neutral atoms to produce free electrons and ions. During the night, the recombination process prevails, where free electrons are recombined with ions to produce neutral particles, which leads to a reduction in the electron density.

refractive index of the iono- sphere can be approximated :

$$\begin{aligned} n_p h &= 1 - \frac{40.3}{f^2} N_e \\ n_g r &= 1 + \frac{40.3}{f^2} N_e \\ \Delta_{ph,f}^{iono} &= -\frac{40.3}{f^2} \int N_e dl = -\alpha_f STEC \\ \Delta_{gr,f}^{iono} &= +\frac{40.3}{f^2} \int N_e dl = \alpha_f STEC \end{aligned}$$

### 1.1 Ionospheric Models for Single-Frequency Receivers

Single-frequency receivers need to apply a model to remove the ionospheric refraction, which can reach up to few tens of metres, depending on the elevation of rays and the ionospheric conditions.

#### 1.1.1 Klobuchar Model

1- Calculate the Earth-centred angle

$$\psi = \pi/2 - E - \arcsin\left(\frac{R_E}{R_E+h} \cos(E)\right)$$

2- Compute the latitude of the IPP

$$\phi_I = \arcsin(\sin(\phi_u)\cos(\psi) + \cos(\phi_u)\sin(\psi)\cos(A))$$

3- Compute the longitude of the IPP

$$\lambda_I = \lambda_u + \frac{\psi\sin(A)}{\cos(\phi_I)}$$

4- Find the geomagnetic latitude of the IPP

$$\phi_m = \arcsin(\sin(\phi_I)\sin(\phi_P) + \cos(\phi_I)\cos(\phi_P)\cos(\lambda_I - \lambda_P))$$

pole coordinate lat and long = 78.3 and 291.0 deg.

5- Find the local time at the IPP

$$t = 43200\lambda_I/\pi + t_{GPS} \text{ and } \lambda_i \text{ radians and } t \text{ in second}$$

6- Compute the amplitude of ionospheric delay

$$A_I = \sum_{n=0}^3 \alpha_n (\phi_m/\pi)^n \text{ (seconds)}$$

if  $A_I < 0 \rightarrow A_I = 0$ .

7- Compute the period of ionospheric delay

$$P_I = \sum_{n=0}^3 \beta_n (\phi_m/\pi) \text{ (seconds)}$$

if  $P_I < 72000 \rightarrow P_I = 72000$ .

8- Compute the phase of ionospheric delay

$$X_I = \frac{2\pi(t-50400)}{P_I}$$

9- Compute the slant factor (ionospheric mapping function)

$$F = [1 - (\frac{R_E}{R_E+h}\cos(E))]$$

10- Compute the ionospheric time delay

$$I_1 = \begin{cases} [5.10^{-9} + A_I \cos(X_I)]F & \text{if } |X_I| < \pi/2 \\ 5.10^{-9}F & \text{if } |X_I| \geq \pi/2 \end{cases}$$

Result :

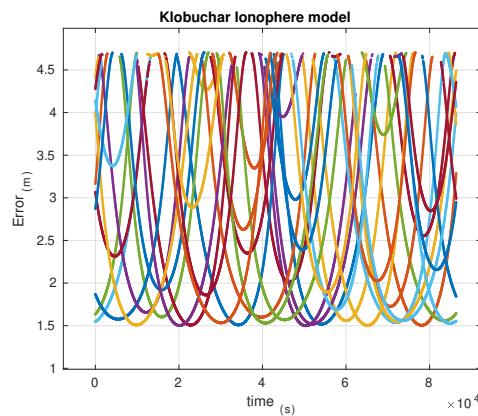


Figure 4: Klobuchar model IZU satation in 2015

## 2 Tropospheric Delay

The troposphere is the atmospheric layer between Earth's surface and an altitude of about 60 km. The effect of the troposphere on the GNSS signals appears as an extra delay in the measurement of the signal travelling from the satellite to the receiver. This delay depends on the temperature, pressure and humidity as well as the transmitter and receiver antenna locations.

$$Tr = \int (n - 1)dl = \int N_{hydr} + N_{wet}$$

Hydrostatic component delay: This is caused by the dry gases present in the troposphere (78 varies with local temperature and atmospheric pressure in a quite predictable manner, although its variation is less than 1 hours. The error caused by this component is about 2.3 m in the zenith direction and 10 m for lower elevations

Wet component delay: This is caused by the water vapour and condensed water in the form of clouds and, therefore, it depends on the weather conditions. The excess delay is small in this case, only some tens of centimetres, but this component varies faster than the hydrostatic component and in a quite random way, thus being very difficult to model.

### 2.1 Collins Model

there is a common mapping function for the wet and dry troposphere :

$$\begin{aligned} Tr(E) &= (Tr_{z,d} + Tr_{z,w})M(E) \\ M(E) &= \frac{1.001}{\sqrt{0.002001 + \sin^2(E)}} \\ \zeta(\phi, D) &= \zeta_0(\phi) - \Delta\zeta(\phi)\cos\left[\frac{2\pi(D - D_{min})}{365.25}\right] \\ Tr_{z_0,d} &= \frac{10^{-6}k_1R_dP}{g_m}, \quad Tr_{z_0,w} = \frac{10^{-k_2Rde}}{(\lambda+1)g_m - \beta R_dT} \\ Tr_{z,d} &= [1 - \frac{\beta H}{T}]^{g/(R_d\beta)} Tr_{z_0,d} \\ Tr_{z,w} &= [1 - \frac{\beta H}{T}]^{(\lambda+1)g/(R_d\beta)-1} Tr_{z_0,w} \end{aligned}$$

$$k_1 = 77.604K/mbar, \quad k_2 = 382000K^2/mbar, \quad R_d = 287.054J/(kgK), \quad g_m = 9.784m/s^2, \\ g = 9.8m/s^2$$

### 2.2 Hopfield

$$\begin{aligned} ZHD &= (0.62291/T + 0.0023081)P \\ ZWD &= (555.7 + 1.79210^{-4}\exp((T - 273)/22.9))e/T^2 \\ Tr_h &= (ZHD + ZWD)M \end{aligned}$$

## 2.3 Sustamiinen

$$g = 1 - 0.0026 \cos d(2\phi) - 0.00000028 h_0$$

$$pr = 1013.25; Tr = 18 + 273$$

$$hr = 0; Hr = 50/100$$

$$p = pr(1 - 0.0000226(h_0 - hr))^5 \cdot 225$$

$$T = Tr - 0.0065(h_0 - hr)$$

$$H = H_r e^{(-0.0006396(h_0 - hr))}$$

$$Tr_{dw} = \frac{0.002277}{g}(P + (1255./T + 0.05)e)$$

$$Tr_s = (Tr_{dw})M$$

Result :

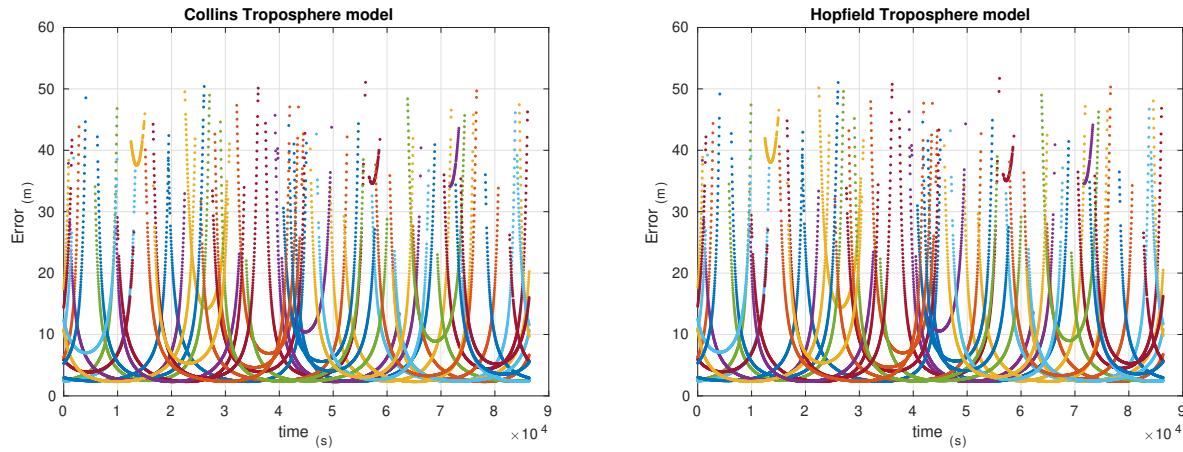


Figure 5: Collins and Hopfield model MIZU satation in 2015

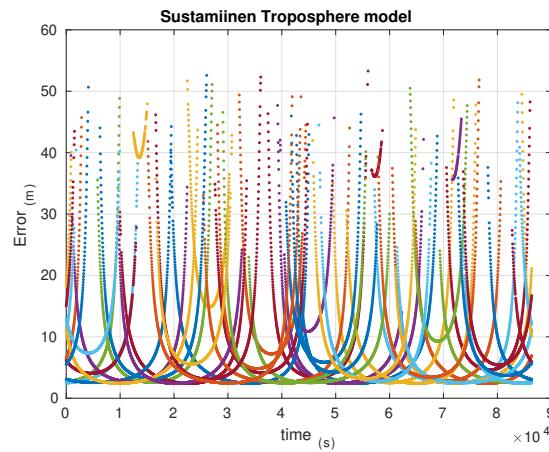


Figure 6: Sustamiinen model MIZU satation in 2015