



# **climlab-0.3 Documentation**

***Release 0.3.0.1***

**Moritz Kreuzer**

April 05, 2016



<b>1</b>	<b>Introduction</b>	<b>3</b>
1.1	What is climlab?	3
1.2	What's new in version 0.3?	3
1.3	Implementation of models	3
1.4	Documentation	4
<b>2</b>	<b>Download</b>	<b>5</b>
2.1	Code	5
2.2	Dependencies	5
2.3	Installation	6
<b>3</b>	<b>Architecture</b>	<b>7</b>
3.1	Process	7
3.2	Domain	10
3.3	Axis	11
3.4	Accessibility	11
<b>4</b>	<b>Models</b>	<b>13</b>
4.1	Energy Balance Model	13
4.2	Other Models	15
<b>5</b>	<b>Tutorials</b>	<b>17</b>
5.1	Preconfigured Energy Balance Models	17
5.2	Boltzmann Outgoing Longwave Radiation	22
5.3	Budyko Transport for Energy Balance Models	26
5.4	Distribution of insolation	31
5.5	The seasonal cycle of surface temperature	38
5.6	Ice - Albedo Feedback and runaway glaciation	45
<b>6</b>	<b>Application Programming Interface</b>	<b>55</b>
6.1	Subpackages	55
6.2	Inheritance Diagram	119
<b>7</b>	<b>References</b>	<b>121</b>
<b>8</b>	<b>License</b>	<b>123</b>
8.1	climlab	123
8.2	Documentation	123
<b>9</b>	<b>Contact</b>	<b>125</b>
9.1	climlab package	125
9.2	climlab documentation	125
<b>10</b>	<b>Indices and tables</b>	<b>127</b>
	<b>Bibliography</b>	<b>129</b>



Contents:



## INTRODUCTION

### 1.1 What is climlab?

*climlab* is a flexible engine for process-oriented climate modeling. It is based on a very general concept of a model as a collection of individual, interacting processes. *climlab* defines a base class called *Process*, which can contain an arbitrarily complex tree of sub-processes (each also some sub-class of *Process*). Every climate process (radiative, dynamical, physical, turbulent, convective, chemical, etc.) can be simulated as a stand-alone process model given appropriate input, or as a sub-process of a more complex model. New classes of model can easily be defined and run interactively by putting together an appropriate collection of sub-processes.

Most of the actual computation for simpler model components use vectorized `numpy` array functions. It should run out-of-the-box on a standard scientific Python distribution, such as Anaconda or Enthought Canopy.

### 1.2 What's new in version 0.3?

New in version 0.3, *climlab* now includes Python wrappers for more numerically intensive processes implemented in Fortran code (specifically the CAM3 radiation module). These require a Fortran compiler on your system, but otherwise have no other library dependencies. *climlab* uses a compile-on-demand strategy. The compiler is invoked automatically as necessary when a new process is created by the user.

### 1.3 Implementation of models

Currently, *climlab* has out-of-the-box support and documented examples for:

- 1D radiative and radiative-convective single column models, with various radiation schemes:
  - Grey Gas
  - Simplified band-averaged models (4 bands each in longwave and shortwave)
  - One GCM-level radiation module (CAM3)
- 1D diffusive energy balance models
- Seasonal and steady-state models
- Arbitrary combinations of the above, for example:
  - 2D latitude-pressure models with radiation, horizontal diffusion, and fixed relative humidity
- orbital / insolation calculations
- boundary layer sensible and latent heat fluxes

---

**Note:** For more details about the implemented Energy Balance Models, see the [Models](#) (page 13) chapter.

---

## 1.4 Documentation

This documentation currently only covers all Energy Balance Model relevant parts of the code, which is just a part of the package. The whole package may be covered in a later release of the documentation.



## DOWNLOAD

### 2.1 Code

Stables releases as well as the current development version can be found on github:

- [Stable Releases](#)
- [Development Version](#)

### 2.2 Dependencies

*climlab* is written in [Python 2.7](#) and requires following *Python* packages to run:

- [Numpy](#)
- [Scipy](#)
- [NetCDF4](#)

**Optional packages:**

- [Jupyter](#) (to run Jupyter notebooks containing tutorials and introduction for *climlab*)
- [Matplotlib](#) (plotting library)

The packages have to be installed on your machine. They can either be downloaded, compiled and installed individually.

Otherwise Python distributions like [Anaconda](#) or [Enthought Canopy](#) can be used which already include many popular Python packages.

#### 2.2.1 Setup environment with Anaconda

An example is given here how to set up a python environment with [Anaconda](#).

A new environment named `climlab_env` with all above packages is created like this:

```
$ conda create --name climlab_env numpy scipy netcdf4 jupyter matplotlib
```

All new packages which will be installed are displayed including their version number. Press `y` to proceed.

After the environment is build, it can be activated with

```
$ source activate climlab_env
```

and can be deactivated with

```
$ source deactivate
```

To install *climlab* in the new environment follow the steps below.

## 2.3 Installation

### 2.3.1 Stable Release

With the Python package `pip` which collects the current version from the [Python Package Index](#), climlab can be easily installed on a machine through the terminal command

```
$ pip install climlab
```

### 2.3.2 Development Version

Otherwise the package can be downloaded from the above referred link and installed manually through running from the package directory

```
$ python setup.py install
```

for a regular system wide installation.

In case you want to develop new code, run following command (which also has an uninstall option):

```
$ python setup.py develop
```

## ARCHITECTURE

The backbone of the *climlab* architecture are *Processes* and their relatives *TimeDependentProcesses*. As all relevant procedures and events that can be modelled with *climlab* are expressed in *Processes*, they build the basic structure of the package.

For example if you want to model the incoming solar radiation on Earth, *climlab* implements it as a *Process* namely in the *Diagnostic Process \_Insolation* (page 103) (or one of it's daughter classes to be specific).

Or the emitted energy of a surface can be computed through the *Boltzmann* (page 96) class which is also a *climlab Process* and implements the Stefan Boltzmann Law for a grey body. Like that all events and procedures that *climlab* can model are organized in *Processes*

---

**Note:** Also the implementation of a whole model, for example an Energy Balance Model (*EBM* (page 73)) is also an instance of the *Process* (page 83) class in *climlab*.

For more about models, see the *climlab Models* (page 13) chapter.

---

A *Process* that represents a whole model will have a couple of *subprocesses* which will be *Processes* themselves. They represent a certain part of the model, for example the albedo or the insolation component. More details about subprocesses can be found below.

A **Process** is always defined on a **Domain** which itself is based on **Axes** or a single **Axis**. The following section will give a basic introduction about their role in the package, their dependencies and their implementation.

### 3.1 Process

A process is an instance of the class *Process* (page 83). Most processes are timedependent and therefore instance of the daughter class *TimeDependentProcess* (page 89).

#### 3.1.1 Basic Dictionaries

A *climlab.Process* object has several iterable dictionaries (*dict*) of named, gridded variables:

- **process.state** contains the *process*'s state variables, which are usually time-dependent and which are major quantities that identify the condition and status of the *process*. This can be the (surface) temperature of a model for instance.
- **process.input** contains boundary conditions and other gridded quantities independent of the *process*. This dictionary is often set by a parent *process*.
- **process.param** contains parameter of the *Process* or model. Basically this is the same as *process.input* but with scalar entries.
- **process.tendencies** is an iterable dictionary of time-tendencies ( $d/dt$ ) for each state variable defined in *process.state*.

---

**Note:** A non `TimeDependentProcess` (but instance of `Process` (page 83)) does not have this dictionary.

---

- **`process.diagnostics`** contains any quantity derived from current state. In an Energy Balance Model this dictionary can have entries like `'ASR'`, `'OLR'`, `'icelat'`, `'net_radiation'`, `'albedo'` or `'insolation'`.
- **`process.subprocess`** holds subprocesses of the *process*. More about subprocesses is described below.

The *process* is fully described by contents of *state*, *input* and *param* dictionaries. *tendencies* and *diagnostics* are always computable from current state.

### 3.1.2 Subprocesses

Subprocesses are representing and modelling certain components of the parent process. A model consists of many subprocesses which are usually defined on the same state variables, domains and axes as the parent process, at least partially.

**Example** The subprocess tree of an EBM may look like this:

```

model_EBM          #<head process>
  diffusion         #<subprocess>
  LW                #<subprocess>
  albedo            #<subprocess>
    iceline         #<sub-subprocess>
    cold_albedo     #<sub-subprocess>
    warm_albedo     #<sub-subprocess>
    insolation      #<subprocess>

```

It can be seen that subprocesses can have subprocesses themselves, like `albedo` in this case.

A subprocess is same as the parent process an instance of the `Process` (page 83) class. So it has dictionaries and attributes with same names as it's parent process. Not necessary all will be the same or having the same entries, but a subprocess has at least the basic dictionaries and attributes created during initialization of the `Process` (page 83) instance.

Every *subprocess* should work independently of its *parent process* given appropriate *input*.

**Example** Investigating an individual *process* (possibly with its own *subprocesses*) isolated from it's parent can be done through:

```

newproc = climlab.process_like(procname.subprocess['subprocname'])
newproc.compute()

```

Thereby anything in the *input* dictionary of `'subprocname'` will remain fixed.

### 3.1.3 Process Integration over time

A `TimeDependentProcess` (page 89) can be integrated over time to see how the state variables and other diagnostic variables vary in time.

#### Time Dependency of a State Variable

For a state variable  $S$  which is dependendet on processes  $P_A$ ,  $P_B$ , ... the time dependency can be written as

$$\frac{dS}{dt} = \underbrace{P_A(S)}_{S \text{ tendency by } P_A} + \underbrace{P_B(S)}_{S \text{ tendency by } P_B} + \dots$$

When state variable  $S$  is discretized over time like

$$\frac{dS}{dt} = \frac{\Delta S}{\Delta t} = \frac{S(t_1) - S(t_0)}{t_1 - t_0} = \frac{S_1 - S_0}{\Delta t}$$

the state tendency can be calculated through

$$\Delta S = [P_A(S) + P_B(S) + \dots] \Delta t$$

and the new state of  $S$  after one timestep  $\Delta t$  is then:

$$S_1 = S_0 + \left[ \underbrace{P_A(S)}_{S \text{ tendency by } P_A} + \underbrace{P_B(S)}_{S \text{ tendency by } P_B} + \dots \right] \Delta t$$

So the new state of  $S$  is calculated through multiplying the process tendencies of  $S$  with the timestep and adding them up to the previous state of  $S$ .

### Time Dependency of an Energy Budget

The time dependency of an EBM energy budget is very similar to the above noted equations, just differing in a heat capacity factor  $C$ . The state variable is temperature  $T$  in this case, which is altered by subprocesses  $SP_A$ ,  $SP_B$ , ...

$$\begin{aligned} \frac{dE}{dt} &= C \frac{dT}{dt} = \underbrace{SP_A(T)}_{\text{heating-rate of } SP_A} + \underbrace{SP_B(T)}_{\text{heating-rate of } SP_B} + \dots \\ \Leftrightarrow \frac{dT}{dt} &= \underbrace{\frac{SP_A(T)}{C}}_{T \text{ tendency by } SP_A} + \underbrace{\frac{SP_B(T)}{C}}_{T \text{ tendency by } SP_B} + \dots \end{aligned}$$

Therefore the new state of  $T$  after one timestep  $\Delta t$  can be written as:

$$T_1 = T_0 + \underbrace{\left[ \frac{SP_A(T)}{C} + \frac{SP_B(T)}{C} + \dots \right] \Delta t}_{\substack{\text{compute()} \\ \text{step\_forward()}}}$$

The integration procedure is implemented in multiple nested function calls. The top functions for model integration are explained here, for details about computation of subprocess tendencies, see [Classification of Subprocess Types](#) (page 10) below.

- **`compute()` (page 90)** is a method that computes tendencies  $d/dt$  for all state variables

- it returns a dictionary of tendencies for all state variables

Temperature tendencies are  $\frac{SP_A(T)}{C}$ ,  $\frac{SP_B(T)}{C}$ , ... in this case, which are summed up like:

$$\text{tendencies}(T) = \frac{SP_A(T)}{C} + \frac{SP_B(T)}{C} + \dots$$

- the keys for this dictionary are same as keys of state dictionary

As temperature  $T$  is the only state variable in this energy budget, the tendencies dictionary also just has the one key, representing state variable  $T$ .

- the tendency dictionary holds the total tendencies for each state including all subprocesses

In case subprocess  $SP_A$  itself has subprocesses, their  $T$  tendencies get included in tendency computation by `compute()` (page 90).

- the method only computes  $d/dt$  but **does not apply changes** (which is done by `step_forward()` (page 92))

- therefore the method is relatively independent of numerical scheme
- method **will update** variables in `proc.diagnostic` dictionary. Therefore it will also **gather all diagnostics** from the *subprocesses*
- `step_forward()` (page 92) updates the state variables
  - it calls `compute()` (page 90) to get current tendencies
  - the method multiplies state tendencies with the timestep and adds them up to the state variables
- `integrate_years()` (page 91) etc will automate time-stepping through calling the `step_forward` (page 92) method multiple times. It also does the computation of time-average diagnostics.
- `integrate_converge()` (page 91) calls `integrate_years()` (page 91) as long as the state variables keep changing over time.

**Example** to integrate a climlab EBM model over time can look like this:

```
import climlab
model = climlab.EBM()

# integrate the model for one year
model.integrate_years(1)
```

## Classification of Subprocess Types

Processes can be classified in types: *explicit*, *implicit*, *diagnostic* and *adjustment*. This makes sense as subprocesses may have different impact on state variable tendencies (*diagnostic* processes don't have a direct influence for instance) or the way their tendencies are computed differ (*explicit* and *implicit*).

Therefore the `compute()` (page 90) method handles them separately as well as in specific order. It calls private `_compute()` methods that are specified in daughter classes of *Process* (page 83) namely *DiagnosticProcess* (page 81), *EnergyBudget* (page 81) (which are explicit processes) or *ImplicitProcess* (page 83).

The description of `compute()` (page 90) reveals the details how the different process types are handled:

The function first computes all diagnostic processes as they may effect all the other processes (such as change in solar distribution). After all the diagnostic processes don't produce any tendencies directly. Subsequently all tendencies and diagnostics for all explicit processes are computed.

Tendencies due to implicit and adjustment processes need to be calculated from a state that is already adjusted after explicit alteration. So the explicit tendencies are applied to the states temporarily. Now all tendencies from implicit processes are calculated through matrix inversions and same like the explicit tendencies applied to the states temporarily. Subsequently all instantaneous adjustments are computed.

Then the changes made to the states from explicit and implicit processes are removed again as this `compute()` (page 90) function is supposed to calculate only tendencies and not applying them to the states.

Finally all calculated tendencies from all processes are collected for each state, summed up and stored in the dictionary `self.tendencies`, which is an attribute of the time-dependent-process object for which the `compute()` (page 90) method has been called.

## 3.2 Domain

A *Domain* defines an area or spatial base for a climlab *Process* (page 83) object. It consists of axes which are *Axis* (page 55) objects that define the dimensions of the *Domain*.

In a *Domain* the heat capacity of grid points, bounds or cells/boxes is specified.

There are daughter classes *Atmosphere* (page 57) and *Ocean* (page 58) of the private *\_Domain* (page 59) class implemented which themselves have daughter classes *SlabAtmosphere* (page 58) and *SlabOcean* (page 59).

Every *Process* (page 83) needs to be defined on a *Domain*. If none is given during initialization but latitude *lat* is specified, a default *Domain* is created.

Several methods are implemented that create *Domains* with special specifications. These are

- *single\_column()* (page 61)
- *zonal\_mean\_column()* (page 62)
- *box\_model\_domain()* (page 60)

## 3.3 Axis

An *Axis* (page 55) is an object where information of a *\_Domain* (page 59)'s spacial dimension are specified.

These include the *type* of the axis, the *number of points*, location of *points* and *bounds* on the spatial dimension, magnitude of bounds differences *delta* as well as their *unit*.

The *axes* of a *\_Domain* (page 59) are stored in the dictionary *axes*, so they can be accessed through *dom.axes* if *dom* is an instance of *\_Domain* (page 59).

## 3.4 Accessibility

For convenience with interactive work, each subprocess 'name' should be accessible as *proc.subprocess.name* as well as the regular way through the subprocess dictionary *proc.subprocess['name']*. Note that *proc* is an instance of the *Process* (page 83) class here.

### Example

```
import climlab
model = climlab.EBM()

# quick access
longwave_subp = model.subprocess['LW']

# regular path
longwave_subp = model.subprocess.LW
```

*climlab* will remain (as much as possible) agnostic about the data formats. Variables within the dictionaries will behave as *numpy.ndarray* objects.

Grid information and other domain details are accessible as attributes of each process. These attributes are *lat*, *lat\_bounds*, *lon*, *lon\_bounds*, *lev*, *lev\_bounds*, *depth* and *depth\_bounds*.

### Example the latitude points of a *process* object that is describing an EBM model

```
import climlab
model = climlab.EBM()

# quick access
lat_points = model.lat

# regular path
lat_points = model.domains['Ts'].axes['lat'].points
```

Shortcuts like *proc.lat* will work where these are unambiguous, which means there is only a single axis of that type in the process.

Many variables will be accessible as process attributes *proc.name*. This restricts to unique field names in the above dictionaries.

**Warning:** There may be other dictionaries that do have name conflicts: e.g. dictionary of tendencies `proc.tendencies`, with same keys as `proc.state`. These will **not be accessible** as `proc.name`, but **will be accessible** as `proc.dict_name.name` (as well as regular dictionary interface `proc.dict_name['name']`).



## MODELS

As indicated in the [Introduction](#) (page 3) *climlab* can implement different types of models out of the box. Here we focus on Energy Balance Models which are referred to as EBMs.

### 4.1 Energy Balance Model

Currently there are three “standard” Energy Balance Models implemented in the *climlab* code. These are [EBM](#) (page 73), [EBM\\_seasonal](#) (page 79) and [EBM\\_annual](#) (page 78), which are explained below.

Let’s first give an overview about different (sub)processes that are implemented:

#### 4.1.1 EBM Subprocesses

##### Insolation

- [FixedInsolation](#) (page 102) defines a constant solar factor for all spatial points of the domain.

$$S(lat) = S_{\text{input}}$$

- [P2Insolation](#) (page 103) characterizes a parabolic solar distribution over the domain’s latitude on basis of the second order Legendre Polynomial  $P_2$ :

$$S(lat) = \frac{S_0}{4} [1 + s_2 P_2(\sin lat)]$$

- [DailyInsolation](#) (page 101) computes the daily solar insolation for each latitude off the domain on the basis of orbital parameters and astronomical formulas.
- [AnnualMeanInsolation](#) (page 99) computes a latitudewise yearly mean for solar insolation on the basis of orbital parameters and astronomical formulas.

##### Albedo

- [ConstantAlbedo](#) (page 110) defines constant albedo values at all spatial points of the domain:

$$\alpha(lat) = a_0$$

- [P2Albedo](#) (page 111) initializes parabolic distributed albedo values across the domain on basis of the second order Legendre Polynomial  $P_2$ :

$$\alpha(lat) = a_0 + a_2 P_2(\sin lat)$$

- [Iceline](#) (page 111) determines which part of the domain is covered with ice according to a given freezing temperature.
- [StepFunctionAlbedo](#) (page 113) implements an albedo step function in dependence of the surface temperature through using instances of the above described albedo classes as subprocesses.

## Outgoing Longwave Radiation

- **[AplusBT](#) (page 93)** calculates the Outgoing Longwave Radiation (*OLR*) in form of a linear dependence of surface temperature  $T$  like

$$OLR = A + B \cdot T$$

- **[AplusBT\\_CO2](#) (page 95)** calculates OLR same as [AplusBT](#) (page 93) but uses parameters  $A$  and  $B$  dependent of the atmospheric CO2 concentration  $c$ .

$$OLR = A(c) + B(c) \cdot T$$

- **[Boltzmann](#) (page 96)** calculates OLR after the Stefan-Boltzmann law for a grey body like

$$OLR = \sigma \epsilon T^4$$

## Energy Transport

These classes calculate the transport of energy  $H(\varphi)$  across the latitude  $\varphi$  in an Energy Budget noted as:

$$C(\varphi) \frac{dT(\varphi)}{dt} = R \downarrow(\varphi) - R \uparrow(\varphi) + H(\varphi)$$

- **[MeridionalDiffusion](#) (page 70)** calculates the energy transport in a diffusion like process along the temperature gradient:

$$H(\varphi) = \frac{D}{\cos \varphi} \frac{\partial}{\partial \varphi} \left( \cos \varphi \frac{\partial T(\varphi)}{\partial \varphi} \right)$$

- **[BudykoTransport](#) (page 66)** calculates the energy transport for each latitude  $\varphi$  in relation to the global mean temperature  $\bar{T}$ :

$$H(\varphi) = -b[T(\varphi) - \bar{T}]$$

### 4.1.2 EBM templates

The preconfigured Energy Balance Models [EBM](#) (page 14), [EBM\\_seasonal](#) (page 14) and [EBM\\_annual](#) (page 15) use the described subprocesses above:

#### EBM

The [EBM](#) (page 73) class sets up a typical Energy Balance Model with following subprocesses:

- Outgoing Longwave Radiation (OLR) parameterization through [AplusBT](#) (page 93)
- solar insolation parameterization through [P2Insolation](#) (page 103)
- albedo parameterization in dependence of temperature through [StepFunctionAlbedo](#) (page 113)
- energy diffusion through [MeridionalDiffusion](#) (page 70)

#### EBM\_seasonal

The [EBM\\_seasonal](#) (page 79) class implements Energy Balance Models with realistic daily insolation. It uses following subprocesses:

- Outgoing Longwave Radiation (OLR) parameterization through [AplusBT](#) (page 93)
- solar insolation parameterization through [DailyInsolation](#) (page 101)
- albedo parameterization in dependence of temperature through [StepFunctionAlbedo](#) (page 113)
- energy diffusion through [MeridionalDiffusion](#) (page 70)

## EBM\_annual

The [EBM\\_annual](#) (page 78) class that implements Energy Balance Models with annual mean insolation. It uses following subprocesses:

- Outgoing Longwave Radiation (OLR) parameterization through [AplusBT](#) (page 93)
- solar insolation parameterization through [AnnualMeanInsolation](#) (page 99)
- albedo parameterization in dependence of temperature through [StepFunctionAlbedo](#) (page 113)
- energy diffusion through [MeridionalDiffusion](#) (page 70)

---

**Note:** For information how to set up individual models or modify instances of the classes above, see the [Tutorials](#) (page 17) chapter.

---

## 4.2 Other Models

As noted in the [Introduction](#) (page 3) more model types are implemented in climlab but not covered in the documentation yet.



## TUTORIALS

For a learning-by-doing approach, a couple of Tutorials come with the *climlab* package. They can be found in the package's *courseware* folder. They are in the Jupyter Notebook format. These Notebooks have been used for teaching some basics of climate science and had also the initial purpose to document the *climlab* package.

### Example usage

The notebooks are self-describing, and should all run out-of-the-box once the package is installed, e.g:

```
jupyter notebook Insolation.ipynb
```

### Notebooks

## 5.1 Preconfigured Energy Balance Models

In this document the basic use of *climlab*'s preconfigured EBM class is shown.

Contents are how to

- setup an EBM model
- show and access subprocesses
- integrate the model
- access and plot various model variables
- calculate the global mean of the temperature

```
In [1]: # import header

%matplotlib inline
import numpy as np
import matplotlib.pyplot as plt
import climlab
from climlab import constants as const
from climlab.domain.field import global_mean
```

### 5.1.1 Model Creation

The regular path for the EBM class is *climlab.model.ebm.EBM* but it can also be accessed through *climlab.EBM*

An EBM model instance is created through

```
In [2]: # model creation
        ebm_model = climlab.EBM()
```

By default many parameters are set during initialization:

```
num_lat=90, S0=const.S0, A=210., B=2., D=0.55, water_depth=10., Tf=-10,
a0=0.3, a2=0.078, ai=0.62, timestep=const.seconds_per_year/90., T0=12.,
T2=-40
```

For further details see the climlab documentation.

Many of the input parameters are stored in the following dictionary:

```
In [3]: # print model parameters
        ebm_model.param

Out[3]: {'A': 210.0,
        'B': 2.0,
        'D': 0.555,
        'S0': 1365.2,
        'Tf': -10.0,
        'a0': 0.3,
        'a2': 0.078,
        'ai': 0.62,
        'timestep': 350632.51200000005,
        'water_depth': 10.0}
```

The model consists of one state variable (surface temperature) and a couple of defined subprocesses.

```
In [4]: # print model states and subprocesses
        print ebm_model

climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.P2Insolation'>
```

### 5.1.2 Model subprocesses

The subprocesses are stored in a dictionary and can be accessed through

```
In [5]: # access model subprocesses
        ebm_model.subprocess.keys()

Out[5]: ['diffusion', 'LW', 'albedo', 'insolation']
```

So to access the time type of the Longwave Radiation subprocess for example, type:

```
In [6]: # access specific subprocess through dictionary
        ebm_model.subprocess['LW'].time_type

Out[6]: 'explicit'
```

### 5.1.3 Model integration

The model time dictionary shows information about all the time related content and quantities.

```
In [7]: # accessing the model time dictionary
        ebm_model.time

Out[7]: {'day_of_year_index': 0,
        'days_elapsed': 0,
        'days_of_year': array([ 0.          ,  4.05824667,  8.11649333, 12.17474
        16.23298667, 20.29123333, 24.34948   , 28.40772667,
        32.46597333, 36.52422   , 40.58246667, 44.64071333,
        48.69896   , 52.75720667, 56.81545333, 60.8737   ,
        64.93194667, 68.99019333, 73.04844   , 77.10668667,
        81.16493333, 85.22318   , 89.28142667, 93.33967333,
        97.39792   , 101.45616667, 105.51441333, 109.57266   ,
        113.63090667, 117.68915333, 121.7474   , 125.80564667,
        129.86389333, 133.92214   , 137.98038667, 142.03863333,
        146.09688   , 150.15512667, 154.21337333, 158.27162   ,
        162.32986667, 166.38811333, 170.44636   , 174.50460667,
        178.56285333, 182.6211   , 186.67934667, 190.73759333,
        194.79584   , 198.85408667, 202.91233333, 206.97058   ,
        211.02882667, 215.08707333, 219.14532   , 223.20356667,
        227.26181333, 231.32006   , 235.37830667, 239.43655333,
        243.4948   , 247.55304667, 251.61129333, 255.66954   ,
        259.72778667, 263.78603333, 267.84428   , 271.90252667,
        275.96077333, 280.01902   , 284.07726667, 288.13551333,
        292.19376   , 296.25200667, 300.31025333, 304.3685   ,
        308.42674667, 312.48499333, 316.54324   , 320.60148667,
        324.65973333, 328.71798   , 332.77622667, 336.83447333,
        340.89272   , 344.95096667, 349.00921333, 353.06746   ,
        357.12570667, 361.18395333]),
        'num_steps_per_year': 90.0,
        'steps': 0,
        'timestep': 350632.512000000005,
        'years_elapsed': 0}
```

To integrate the model forward in time different methods are available:

```
In [8]: # integrate model for a single timestep
        ebm_model.step_forward()
```

The model time step has increased from 0 to 1:

```
In [9]: ebm_model.time['steps']
```

```
Out[9]: 1
```

```
In [10]: # integrate model for a 50 days
         ebm_model.integrate_days(50.)
```

Integrating for 12 steps, 50.0 days, or 0.136895462792 years.  
Total elapsed time is 0.144444444444 years.

```
In [11]: # integrate model for two years
         ebm_model.integrate_years(1.)
```

Integrating for 90 steps, 365.2422 days, or 1.0 years.  
Total elapsed time is 1.144444444444 years.

```
In [12]: # integrate model until solution converges
         ebm_model.integrate_converge()
```

Total elapsed time is 9.144444444444 years.

## Plotting model variables

A couple of interesting model variables are stored in a dictionary named `diagnostics`. It has following entries:

```
In [13]: ebm_model.diagnostics.keys()
```

```
Out[13]: ['ASR', 'OLR', 'icelat', 'net_radiation', 'albedo', 'insolation']
```

They can be accessed respectively to the keys through `ebm_model.diagnostics['ASR']`. Most of them can be also accessed directly as model attributes like:

```
In [14]: ebm_model.icelat
```

```
Out[14]: array([-70.,  70.])
```

The following code does the plotting for some model variables.

```
In [15]: # creating plot figure
fig = plt.figure(figsize=(15,10))

# Temperature plot
ax1 = fig.add_subplot(221)
ax1.plot(ebm_model.lat, ebm_model.Ts)

ax1.set_xticks([-90, -60, -30, 0, 30, 60, 90])
ax1.set_xlim([-90, 90])
ax1.set_xlabel('latitude')
ax1.set_ylabel('surface temperature (degC)', fontsize=12)
ax1.grid()

# Albedo plot
ax2 = fig.add_subplot(223, sharex = ax1)
ax2.plot(ebm_model.lat, ebm_model.albedo)

ax2.set_ylabel('albedo', fontsize=12)
ax2.set_ylim([0, 1])
ax2.grid()

# Net Radiation plot
ax3 = fig.add_subplot(222, sharex = ax1)
ax3.plot(ebm_model.lat, ebm_model.OLR, label='OLR',
        color='cyan')
ax3.plot(ebm_model.lat, ebm_model.ASR, label='ASR',
        color='magenta')
ax3.plot(ebm_model.lat, ebm_model.ASR-ebm_model.OLR,
        label='net radiation',
        color='red')

ax3.set_ylabel('radiation (W/m$^2$)', fontsize=12)
ax3.legend(loc='best')
ax3.grid()

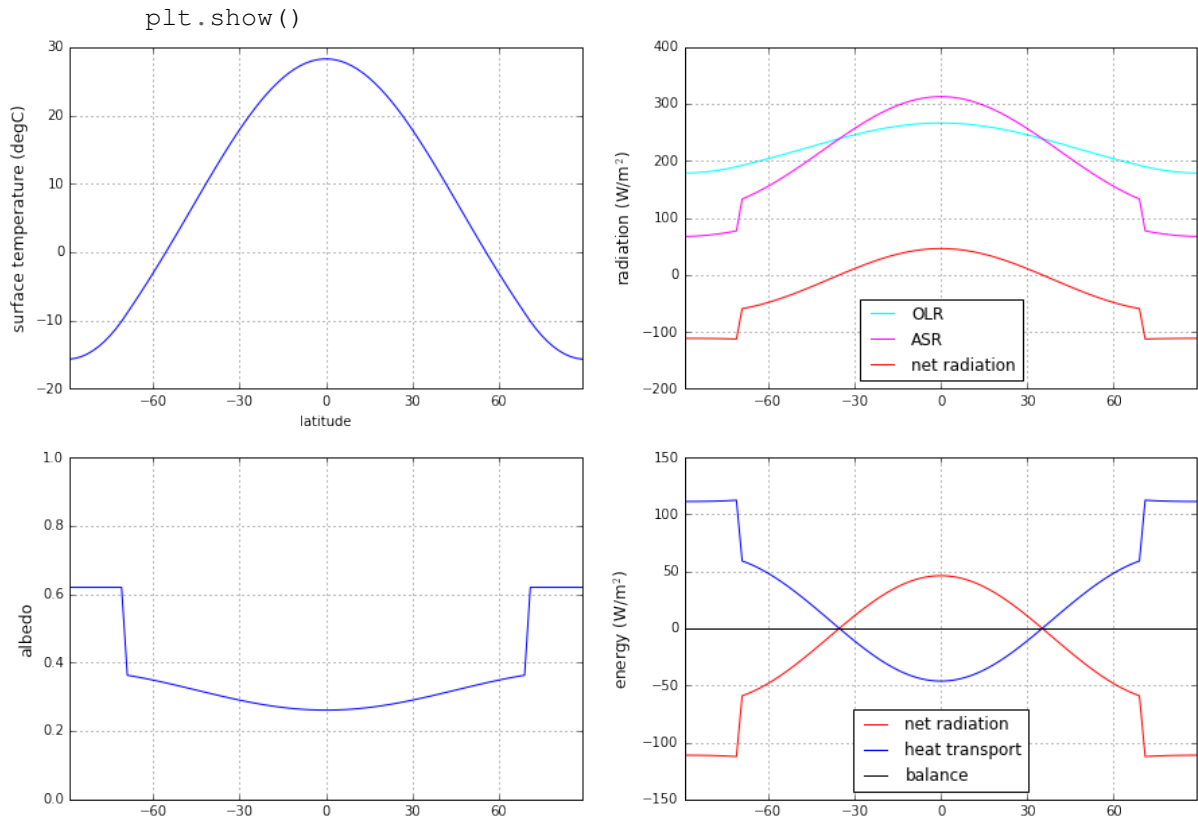
# Energy Balance plot
net_rad = np.squeeze(ebm_model.net_radiation)
transport = ebm_model.heat_transport_convergence()

ax4 = fig.add_subplot(224, sharex = ax1)
ax4.plot(ebm_model.lat, net_rad, label='net radiation',
        color='red')
ax4.plot(ebm_model.lat, transport, label='heat transport',
        color='blue')
```



```
ax4.plot(ebm_model.lat, net_rad+transport, label='balance',
        color='black')

ax4.set_ylabel('energy (W/m$^2$)', fontsize=12)
ax4.legend(loc='best')
ax4.grid()
```



### 5.1.4 Global mean temperature

The model's state dictionary has following entries:

```
In [16]: ebm_model.state.keys()
```

```
Out [16]: ['Ts']
```

So the surface temperature can usually be accessed through `ebm_model.state['Ts']` but is also available as a model attribute: `ebm_model.Ts`

The global mean of the model's surface temperature can be calculated through

```
In [17]: # calculate global mean temperature
         global_mean(ebm_model.Ts)
```

```
Out [17]: Field(14.288135944994657)
```

Note that in the **header** the `global_mean` method has been **imported**!

```
In [18]: print 'The global mean temperature is %s degC with a model ice edge at %s deg.'
         % (np.round(global_mean(ebm_model.Ts), 2), np.max(ebm_model.icelat))
```

The global mean temperature is 14.29 degC with a model ice edge at 70.0 deg.

## 5.2 Boltzmann Outgoing Longwave Radiation

In this document an Energy Balance Model (EBM) is set up with the Outgoing Longwave Radiation (OLR) parameterized through the Stefan Boltzmann radiation of a grey body.

$$OLR(\varphi) = \sigma \cdot \varepsilon \cdot T_s(\varphi)^4$$

```
In [1]: # import header

%matplotlib inline
import numpy as np
import matplotlib.pyplot as plt
import climlab
from climlab import constants as const
from climlab.domain.field import global_mean
```

### 5.2.1 Model Creation

An EBM model instance is created through

```
In [2]: # model creation
ebm_boltz = climlab.EBM()
```

The model is set up by default with a linearized OLR parameterization (A+BT).

```
In [3]: # print model states and subprocesses
print ebm_boltz
```

```
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.P2Insolation'>
```

### 5.2.2 Create new subprocess

The creation of a subprocess needs some information from the model, especially on which model state the subprocess should be defined on.

```
In [4]: # create Boltzmann subprocess
LW_boltz = climlab.radiation.Boltzmann(eps=0.69, tau=0.95,
                                       state=ebm_boltz.state,
                                       **ebm_boltz.param)
```

Note that the model's **whole state dictionary** is given as **input** to the subprocess. In case only the temperature field `ebm_boltz.state['Ts']` would be given, a new state dictionary would be created which holds the surface temperature with the key `'default'`. That raises an error as the Boltzmann process refers the temperature with key `'Ts'`.

Now the new OLR subprocess has to be merged into the model. Therefore the `AplusBT` subprocess has to be removed first.

```
In [5]: # remove the old longwave subprocess
        ebm_boltz.remove_subprocess('LW')

        # add the new longwave subprocess
        ebm_boltz.add_subprocess('LW', LW_boltz)
```

Note that the new OLR subprocess has to have the **same key** “LW” as the old one, as the model refers to this key for radiation balance computation.

That is why the old process has to be removed before the new one is added.

```
In [6]: print ebm_boltz

climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.Boltzmann.Boltzmann'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.P2Insolation'>
```

## 5.2.3 Model integration & Plotting

```
In [7]: # integrate model for a single timestep
        ebm_boltz.step_forward()

In [8]: # creating plot figure
        fig = plt.figure(figsize=(15,10))

        # Temperature plot
        ax1 = fig.add_subplot(221)
        ax1.plot(ebm_boltz.lat, ebm_boltz.Ts)

        ax1.set_xticks([-90, -60, -30, 0, 30, 60, 90])
        ax1.set_xlim([-90, 90])
        ax1.set_xlabel('latitude')
        ax1.set_ylabel('surface temperature (degC)', fontsize=12)
        ax1.grid()

        # Albedo plot
        ax2 = fig.add_subplot(223, sharex = ax1)
        ax2.plot(ebm_boltz.lat, ebm_boltz.albedo)

        ax2.set_ylabel('albedo', fontsize=12)
        ax2.set_ylim([0, 1])
        ax2.grid()

        # Net Radiation plot
        ax3 = fig.add_subplot(222, sharex = ax1)
        ax3.plot(ebm_boltz.lat, ebm_boltz.OLR, label='OLR',
                color='cyan')
        ax3.plot(ebm_boltz.lat, ebm_boltz.ASR, label='ASR',
                color='magenta')
        ax3.plot(ebm_boltz.lat, ebm_boltz.ASR-ebm_boltz.OLR,
                label='net radiation',
```

```

        color='red')

ax3.set_ylabel('net radiation (W/m$^2$)', fontsize=12)
ax3.legend(loc='best')
ax3.grid()

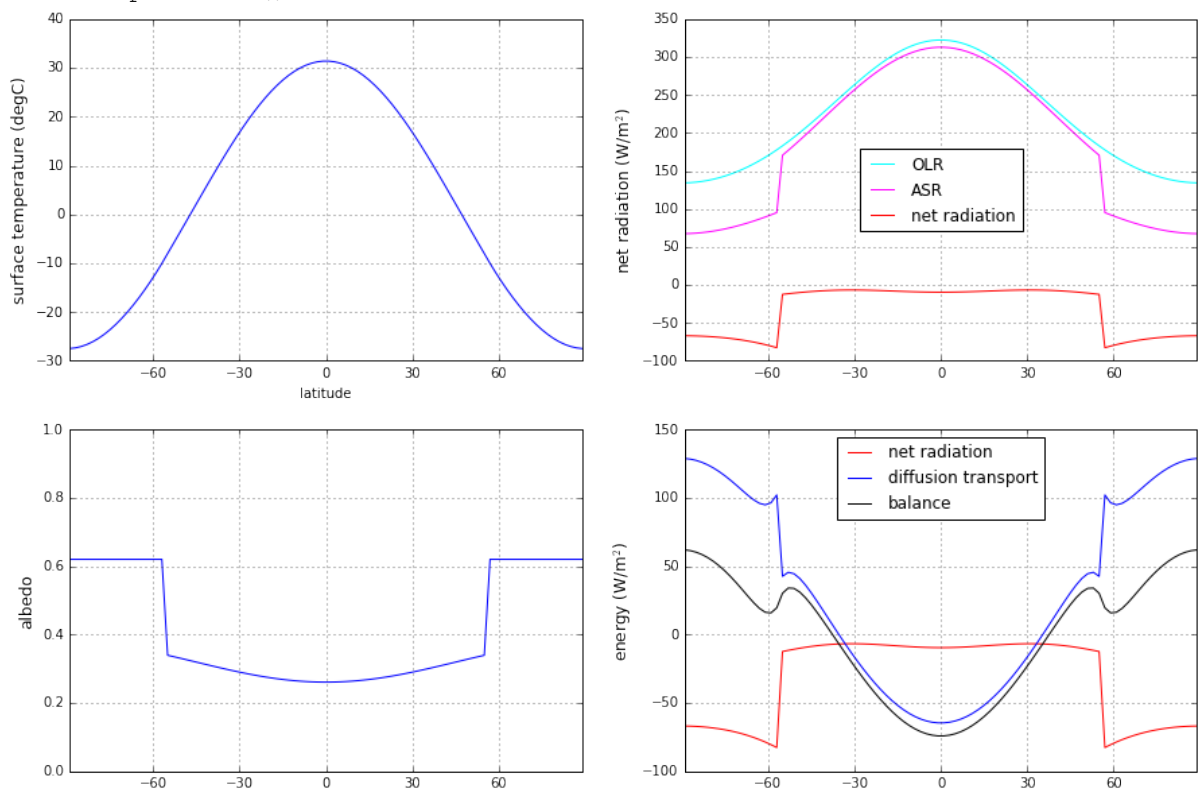
# Energy Balance plot
net_rad = np.squeeze(ebm_boltz.net_radiation)
transport = ebm_boltz.heat_transport_convergence()

ax4 = fig.add_subplot(224, sharex = ax1)
ax4.plot(ebm_boltz.lat, net_rad, label='net radiation',
        color='red')
ax4.plot(ebm_boltz.lat, transport, label='diffusion transport',
        color='blue')
ax4.plot(ebm_boltz.lat, net_rad+transport, label='balance',
        color='black')

ax4.set_ylabel('energy (W/m$^2$)', fontsize=12)
ax4.legend(loc='best')
ax4.grid()

```

```
plt.show()
```



```

In [9]: # integrate model until solution converges
        ebm_boltz.integrate_converge()

```

Total elapsed time is 5.011111111111 years.

```

In [10]: # creating plot figure
         fig = plt.figure(figsize=(15,10))

```

```
# Temperature plot
ax1 = fig.add_subplot(221)
ax1.plot(ebm_boltz.lat, ebm_boltz.Ts)

ax1.set_xticks([-90, -60, -30, 0, 30, 60, 90])
ax1.set_xlim([-90, 90])
ax1.set_xlabel('latitude')
ax1.set_ylabel('surface temperature (degC)', fontsize=12)
ax1.grid()

# Albedo plot
ax2 = fig.add_subplot(223, sharex = ax1)
ax2.plot(ebm_boltz.lat, ebm_boltz.albedo)

ax2.set_ylabel('albedo', fontsize=12)
ax2.set_ylim([0, 1])
ax2.grid()

# Net Radiation plot
ax3 = fig.add_subplot(222, sharex = ax1)
ax3.plot(ebm_boltz.lat, ebm_boltz.OLR, label='OLR',
        color='cyan')
ax3.plot(ebm_boltz.lat, ebm_boltz.ASR, label='ASR',
        color='magenta')
ax3.plot(ebm_boltz.lat, ebm_boltz.ASR-ebm_boltz.OLR,
        label='net radiation',
        color='red')

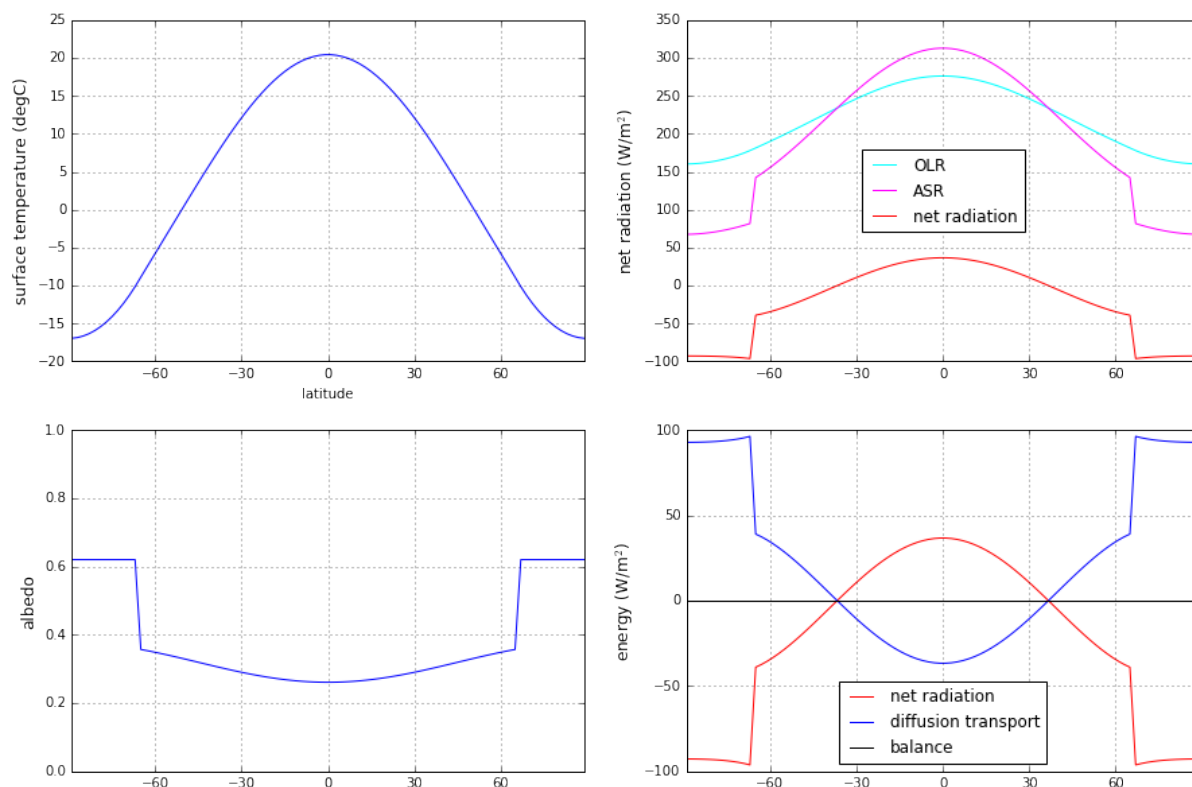
ax3.set_ylabel('net radiation (W/m$^2$)', fontsize=12)
ax3.legend(loc='best')
ax3.grid()

# Energy Balance plot
net_rad = np.squeeze(ebm_boltz.net_radiation)
transport = ebm_boltz.heat_transport_convergence()

ax4 = fig.add_subplot(224, sharex = ax1)
ax4.plot(ebm_boltz.lat, net_rad, label='net radiation',
        color='red')
ax4.plot(ebm_boltz.lat, transport, label='diffusion transport',
        color='blue')
ax4.plot(ebm_boltz.lat, net_rad+transport, label='balance',
        color='black')

ax4.set_ylabel('energy (W/m$^2$)', fontsize=12)
ax4.legend(loc='best')
ax4.grid()

plt.show()
```



## 5.2.4 Global mean temperature

```
In [11]: print 'The global mean temperature is %s degC with a model ice edge at %s deg.' % (np.round(global_mean(ebm_boltz.Ts), 2), np.max(ebm_boltz.icelat))
```

The global mean temperature is 8.87 degC with a model ice edge at 66.0 deg.

## 5.3 Budyko Transport for Energy Balance Models

In this document an Energy Balance Model (EBM) is set up with the energy transport parameterized through the **budyko type parameterization** term (instead of the default diffusion term), which characterizes the local energy flux through the difference between local temperature and global mean temperature.

$$H(\varphi) = -b[T(\varphi) - \bar{T}]$$

where  $T(\varphi)$  is the surface temperature across the latitude  $\varphi$ ,  $\bar{T}$  the global mean temperature and  $H(\varphi)$  is the transport of energy in an Energy Budget noted as:

$$C(\varphi) \frac{dT(\varphi)}{dt} = R \downarrow(\varphi) - R \uparrow(\varphi) + H(\varphi)$$

```
In [1]: # import header

%matplotlib inline
import numpy as np
import matplotlib.pyplot as plt
import climlab
from climlab import constants as const
from climlab.domain.field import global_mean
```

### 5.3.1 Model Creation

An EBM model instance is created through

```
In [2]: # model creation
        ebm_budyko= climlab.EBM()
```

The model is set up by default with a meridional diffusion term.

```
In [3]: # print model states and subprocesses
        print ebm_budyko
```

```
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.P2Insolation'>
```

### 5.3.2 Create new subprocess

The creation of a subprocess needs some information from the model, especially on which model state the subprocess should be defined on.

```
In [4]: # create Budyko subprocess
        budyko_transp = climlab.dynamics.BudykoTransport(b=3.81,
                                                         state=ebm_budyko.state,
                                                         **ebm_budyko.param)
```

Note that the model's **whole state dictionary** is given as **input** to the subprocess. In case only the temperature field `ebm_budyko.state['Ts']` is given, a new state dictionary would be created which holds the surface temperature with the key 'default'. That raises an error as the budyko transport process refers the temperature with key 'Ts'.

Now the new transport subprocess has to be merged into the model. The diffusion subprocess has to be removed.

```
In [5]: # add the new transport subprocess
        ebm_budyko.add_subprocess('budyko_transport',budyko_transp)

        # remove the old diffusion subprocess
        ebm_budyko.remove_subprocess('diffusion')
```

```
In [6]: print ebm_budyko
```

```
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  budyko_transport: <class 'climlab.dynamics.budyko_transport.BudykoTransport'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
```

```
warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
insolation: <class 'climlab.radiation.insolation.P2Insolation'>
```

### 5.3.3 Model integration & Plotting

```
In [7]: # integrate model for a single timestep
        ebm_budyko.step_forward()

In [8]: # creating plot figure
        fig = plt.figure(figsize=(15,10))

        # Temperature plot
        ax1 = fig.add_subplot(221)
        ax1.plot(ebm_budyko.lat, ebm_budyko.Ts)

        ax1.set_xticks([-90, -60, -30, 0, 30, 60, 90])
        ax1.set_xlim([-90, 90])
        ax1.set_xlabel('latitude')
        ax1.set_ylabel('surface temperature (degC)', fontsize=12)
        ax1.grid()

        # Albedo plot
        ax2 = fig.add_subplot(223, sharex = ax1)
        ax2.plot(ebm_budyko.lat, ebm_budyko.albedo)

        ax2.set_ylabel('albedo', fontsize=12)
        ax2.set_ylim([0,1])
        ax2.grid()

        # Net Radiation plot
        ax3 = fig.add_subplot(222, sharex = ax1)
        ax3.plot(ebm_budyko.lat, ebm_budyko.OLR, label='OLR',
                color='cyan')
        ax3.plot(ebm_budyko.lat, ebm_budyko.ASR, label='ASR',
                color='magenta')
        ax3.plot(ebm_budyko.lat, ebm_budyko.ASR-ebm_budyko.OLR,
                label='net radiation',
                color='red')

        ax3.set_ylabel('radiation (W/m$^2$)', fontsize=12)
        ax3.legend(loc='best')
        ax3.grid()

        # Energy Balance plot
        net_rad = ebm_budyko.net_radiation
        transport = ebm_budyko.subprocess['budyko_transport'].heating_rate['Ts']

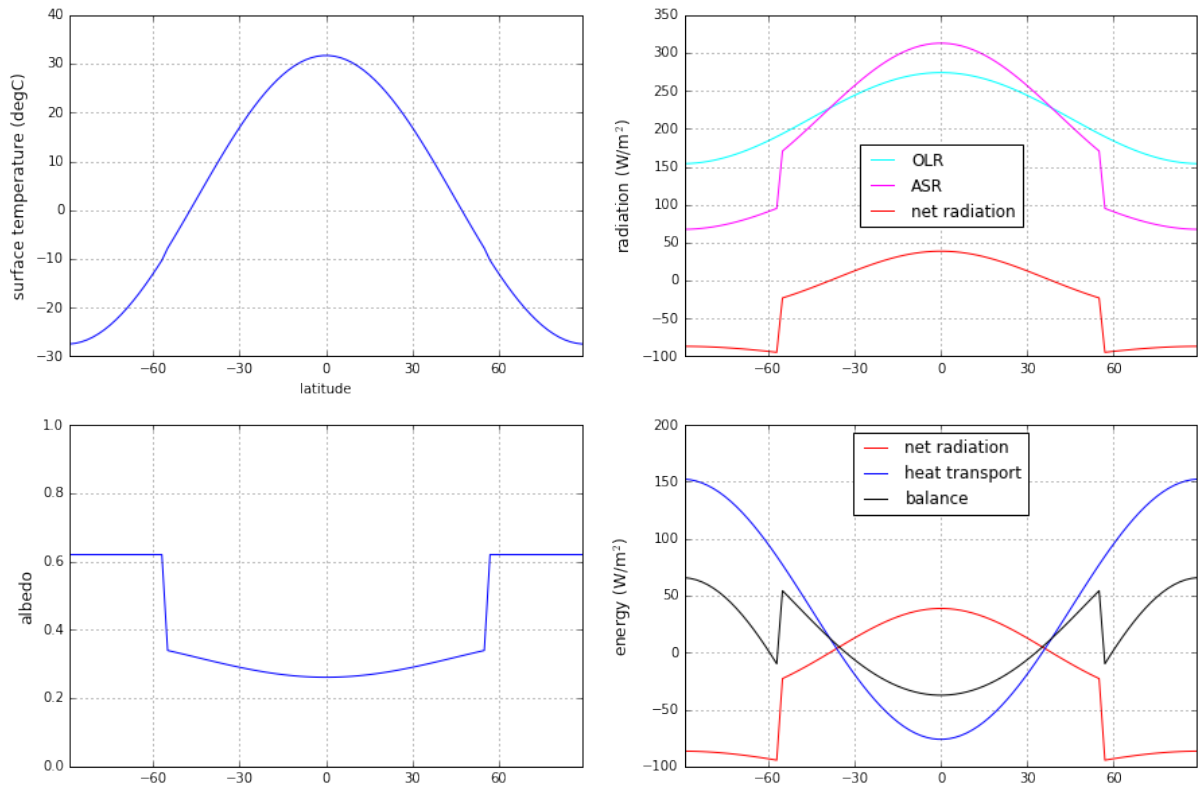
        ax4 = fig.add_subplot(224, sharex = ax1)
        ax4.plot(ebm_budyko.lat, net_rad, label='net radiation',
                color='red')
        ax4.plot(ebm_budyko.lat, transport, label='heat transport',
                color='blue')
        ax4.plot(ebm_budyko.lat, net_rad+transport, label='balance',
                color='black')

        ax4.set_ylabel('energy (W/m$^2$)', fontsize=12)
        ax4.legend(loc='best')
```



```
ax4.grid()
```

```
plt.show()
```



```
In [9]: # integrate model until solution converges
ebm_budyko.integrate_converge()
```

Total elapsed time is 7.011111111111 years.

```
In [10]: # creating plot figure
fig = plt.figure(figsize=(15,10))

# Temperature plot
ax1 = fig.add_subplot(221)
ax1.plot(ebm_budyko.lat, ebm_budyko.Ts)

ax1.set_xticks([-90, -60, -30, 0, 30, 60, 90])
ax1.set_xlim([-90, 90])
ax1.set_xlabel('latitude')
ax1.set_ylabel('surface temperature (degC)', fontsize=12)
ax1.grid()

# Albedo plot
ax2 = fig.add_subplot(223, sharex = ax1)
ax2.plot(ebm_budyko.lat, ebm_budyko.albedo)

ax2.set_ylabel('albedo', fontsize=12)
ax2.set_ylim([0,1])
ax2.grid()

# Net Radiation plot
ax3 = fig.add_subplot(222, sharex = ax1)
ax3.plot(ebm_budyko.lat, ebm_budyko.OLR, label='OLR',
```

```

                                color='cyan')
ax3.plot(ebm_budyko.lat, ebm_budyko.ASR, label='ASR',
                                color='magenta')
ax3.plot(ebm_budyko.lat, ebm_budyko.ASR-ebm_budyko.OLR,
                                label='net radiation',
                                color='red')

ax3.set_ylabel('radiation (W/m$^2$)', fontsize=12)
ax3.legend(loc='best')
ax3.grid()

# Energy Balance plot
net_rad = ebm_budyko.net_radiation
transport = ebm_budyko.subprocess['budyko_transport'].heating_rate['Ts']

ax4 = fig.add_subplot(224, sharex = ax1)
ax4.plot(ebm_budyko.lat, net_rad, label='net radiation',
                                color='red')

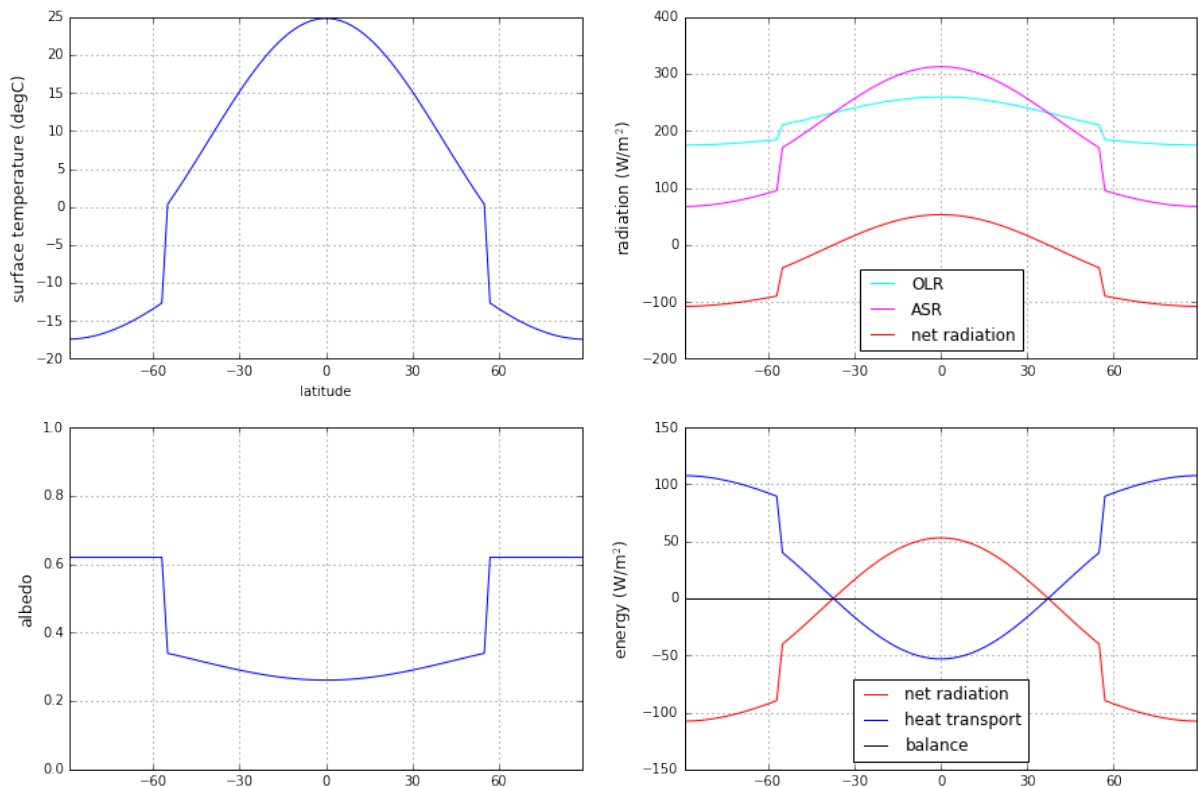
ax4.plot(ebm_budyko.lat, transport, label='heat transport',
                                color='blue')

ax4.plot(ebm_budyko.lat, net_rad+transport, label='balance',
                                color='black')

ax4.set_ylabel('energy (W/m$^2$)', fontsize=12)
ax4.legend(loc='best')
ax4.grid()

```

```
plt.show()
```



### 5.3.4 Global mean temperature

```
In [11]: print 'The global mean temperature is %s degC with a model ice edge at %s deg.'
          %(np.round(global_mean(ebm_budyko.Ts),2), np.max(ebm_budyko.icelat))
```

The global mean temperature is 10.87 degC with a model ice edge at 56.0 deg.

## 5.4 Distribution of insolation

Here are some examples calculating daily average insolation at different locations and times.

These all use a function called `daily_insolation` in the module `insolation.py` to do the calculation. The code calculates daily average insolation anywhere on Earth at any time of year for a given set of orbital parameters.

To look at past orbital variations and their effects on insolation, we use the module `orbital.py` which accesses tables of values for the past 5 million years. We can easily lookup parameters for any point in the past and pass these to `daily_insolation`.

```
In [1]: %matplotlib inline
import numpy as np
import matplotlib.pyplot as plt
import netCDF4 as nc
from climlab import constants as const
from climlab.solar.insolation import daily_insolation
from climlab.solar.orbital import OrbitalTable
```

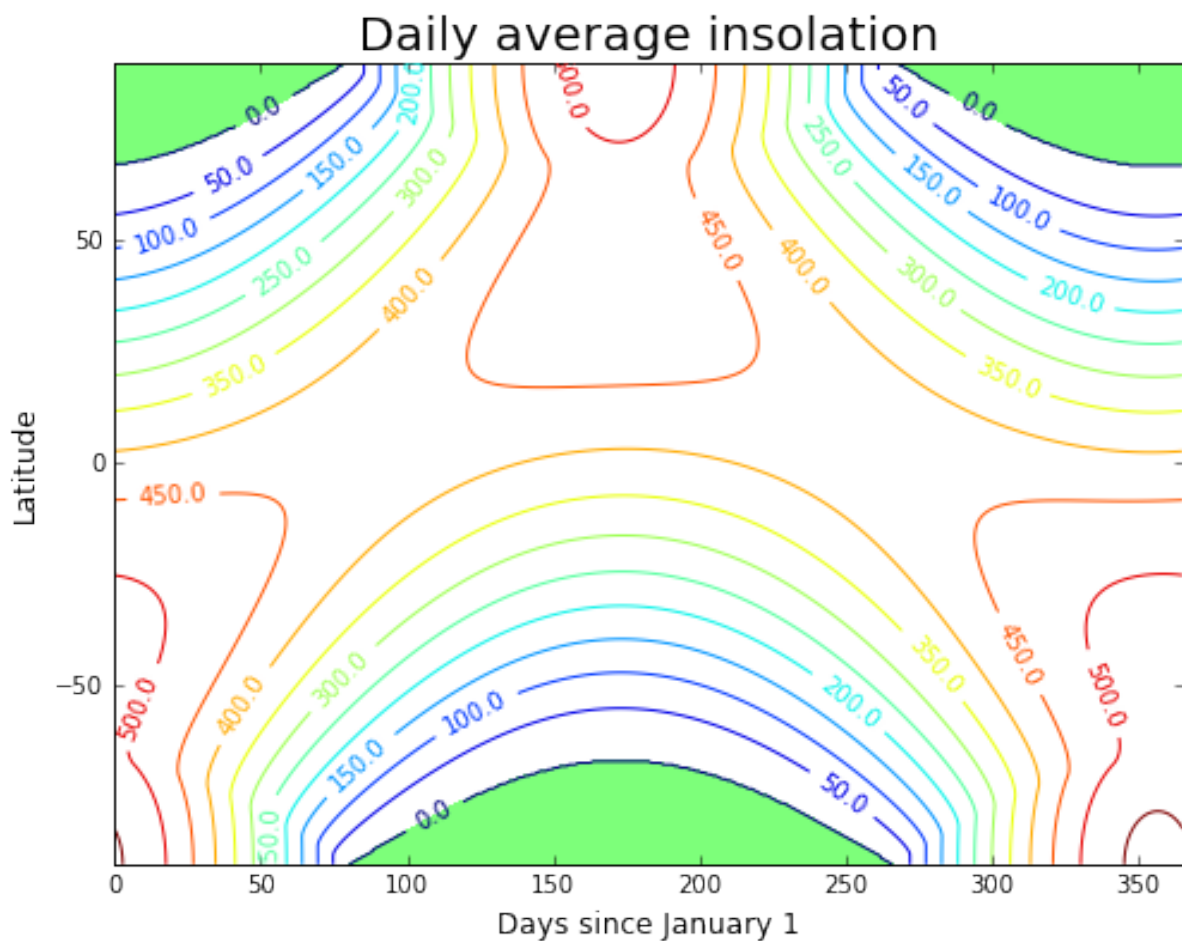
### 5.4.1 Present-day orbital parameters

Calculate an array of insolation over the year and all latitudes (for present-day orbital parameters).

```
In [2]: lat = np.linspace( -90., 90., 500. )
        days = np.linspace(0, const.days_per_year, 365. )
        Q = daily_insolation( lat, days )
```

And make a contour plot of `Q` as function of latitude and time of year.

```
In [3]: ax = plt.figure( figsize=(8,6) ).add_subplot(111)
        CS = ax.contour( days, lat, Q , levels = np.arange(0., 600., 50.) )
        ax.clabel(CS, CS.levels, inline=True, fmt='%r', fontsize=10)
        ax.set_xlabel('Days since January 1', fontsize=12 )
        ax.set_ylabel('Latitude', fontsize=12 )
        ax.set_title('Daily average insolation', fontsize=20 )
        ax.contourf ( days, lat, Q, levels=[-500., 0.] )
        plt.show()
```

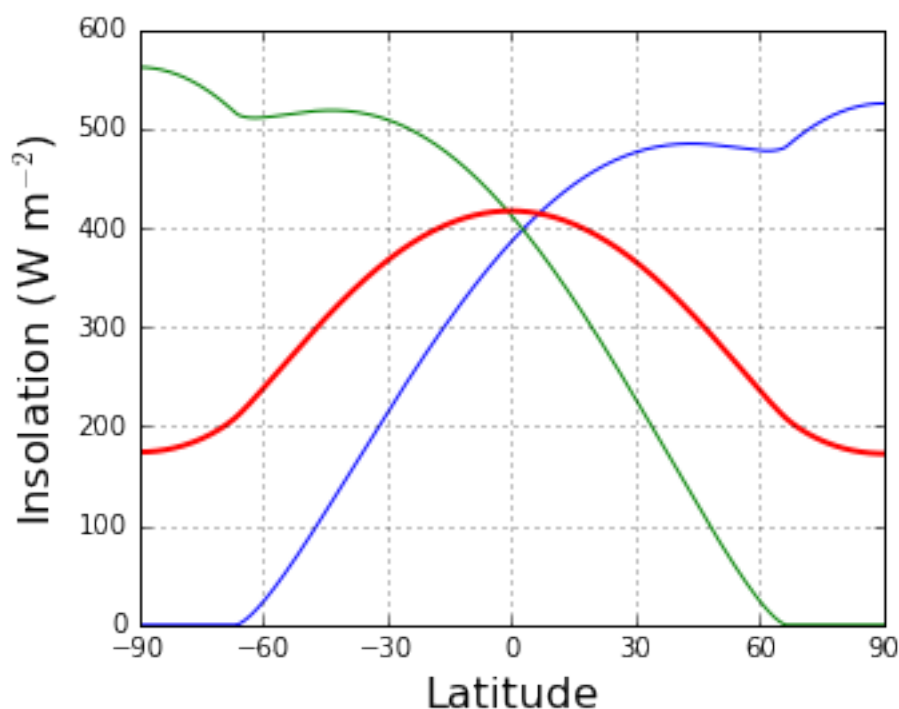


Take the area-weighted global, annual average of  $Q_{\text{net}}$ :

```
In [4]: print np.sum( np.mean( Q, axis=1 ) * np.cos( np.deg2rad(lat) ) ) / np.sum( np.cos( np.deg2rad(lat) ) )
341.384184481
```

Also plot the zonally averaged insolation at a few different times of the year:

```
In [5]: summer_solstice = 170
        winter_solstice = 353
        ax = plt.figure( figsize=(5,4) ).add_subplot(111)
        ax.plot( lat, Q[:,(summer_solstice, winter_solstice)] );
        ax.plot( lat, np.mean(Q, axis=1), linewidth=2 )
        ax.set_xbound(-90, 90)
        ax.set_xticks( range(-90,100,30) )
        ax.set_xlabel('Latitude', fontsize=16 );
        ax.set_ylabel('Insolation (W m$^{-2}$)', fontsize=16 );
        ax.grid()
        plt.show()
```



### 5.4.2 Past orbital parameters

The `orbital.py` code allows us to look up the orbital parameters for Earth over the last 5 million years.

Make reference plots of the variation in the three orbital parameter over the last 1 million years

```
In [6]: kyears = np.arange( -1000., 1.)
        table = OrbitalTable()
        orb = table.lookup_parameters( kyears )
```

Loading Berger and Loutre (1991) orbital parameter data from file `/home/moritz/anaconda2`

The Python object `orb` now holds 1 million years worth of orbital data, total of 1001 data points for each element: eccentricity `ecc`, obliquity angle `obliquity`, and solar longitude of perihelion `long_peri`.

```
In [23]: orb
```

```
Out[23]: {'ecc': array([ 0.035765,  0.036953,  0.038114, ...,  0.018024,  0.017644,
                        0.017236]),
          'long_peri': array([ 122.46,  138.29,  154.17, ...,  247.23,  264.26,  281.37]),
          'obliquity': array([ 23.778,  23.835,  23.877, ...,  23.697,  23.573,  23.446])}
```

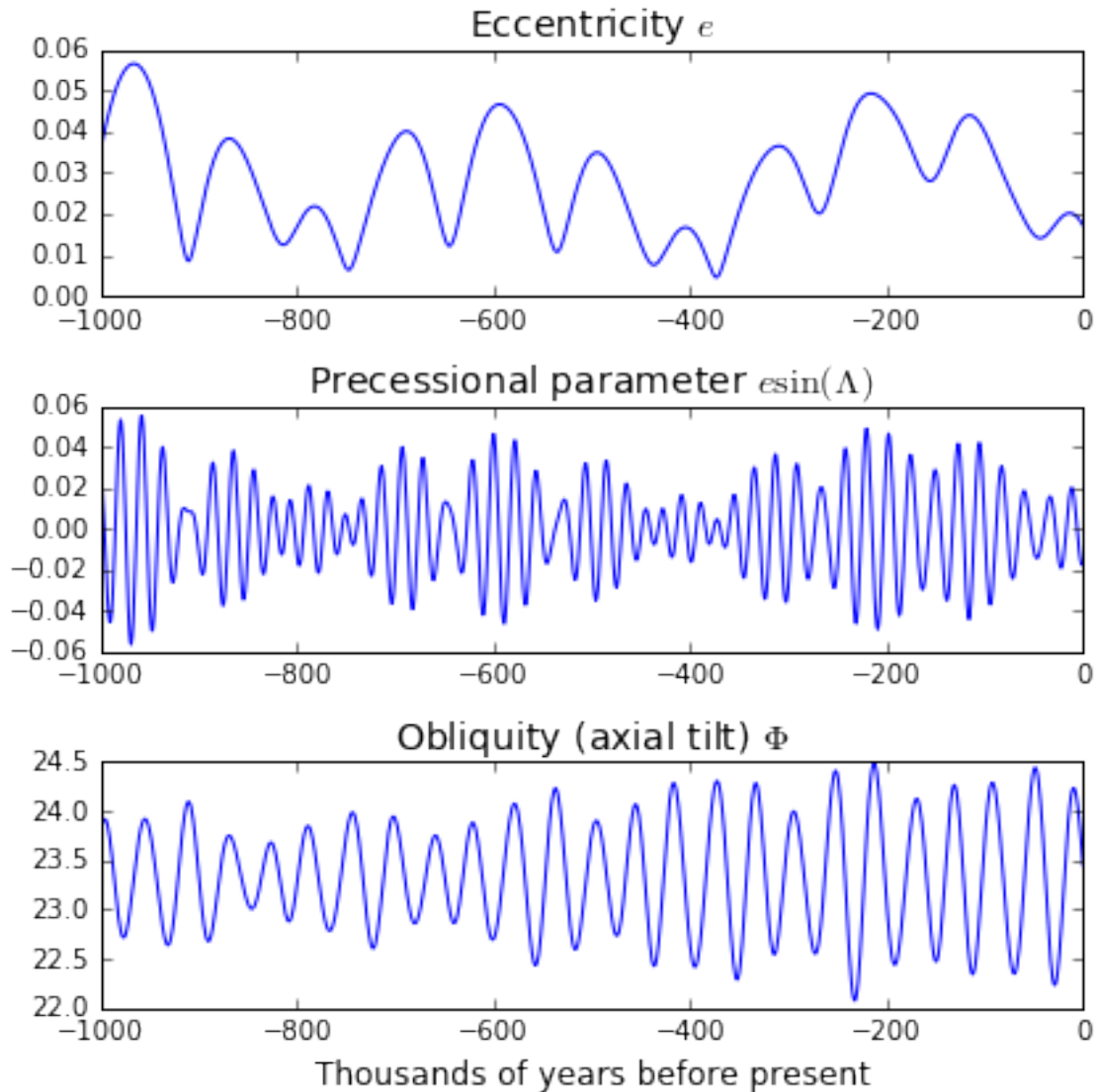
```
In [25]: print [np.shape( orb['ecc'] ),
                np.shape( orb['long_peri'] ),
                np.shape( orb['obliquity'] ) ]
```

```
[(1001,), (1001,), (1001,)]
```

```
In [8]: fig = plt.figure( figsize = (6,6) )
        ax1 = fig.add_subplot(3,1,1)
        ax1.plot( kyears, orb['ecc'] )
        ax1.set_title('Eccentricity $e$', fontsize=14 )
        ax2 = fig.add_subplot(3,1,2)
        ax2.plot( kyears, orb['ecc'] * np.sin( np.deg2rad( orb['long_peri'] ) ) )
        ax2.set_title('Precessional parameter $e \sin(\Lambda)$', fontsize=14 )
        ax3 = fig.add_subplot(3,1,3)
        ax3.plot( kyears, orb['obliquity'] )
```

```
ax3.set_title('Obliquity (axial tilt)  $\Phi$ ', fontsize=14 )
ax3.set_xlabel( 'Thousands of years before present', fontsize=12 )

plt.tight_layout()
plt.show()
```



### Annual mean insolation

Create a large array of insolation over the whole globe, whole year, and for every set of orbital parameters.

```
In [9]: lat = np.linspace(-90, 90, 181)
        days = np.linspace(1., 50.)/50 * const.days_per_year
        Q = daily_insolation(lat, days, orb)
        print Q.shape

(181, 50, 1001)

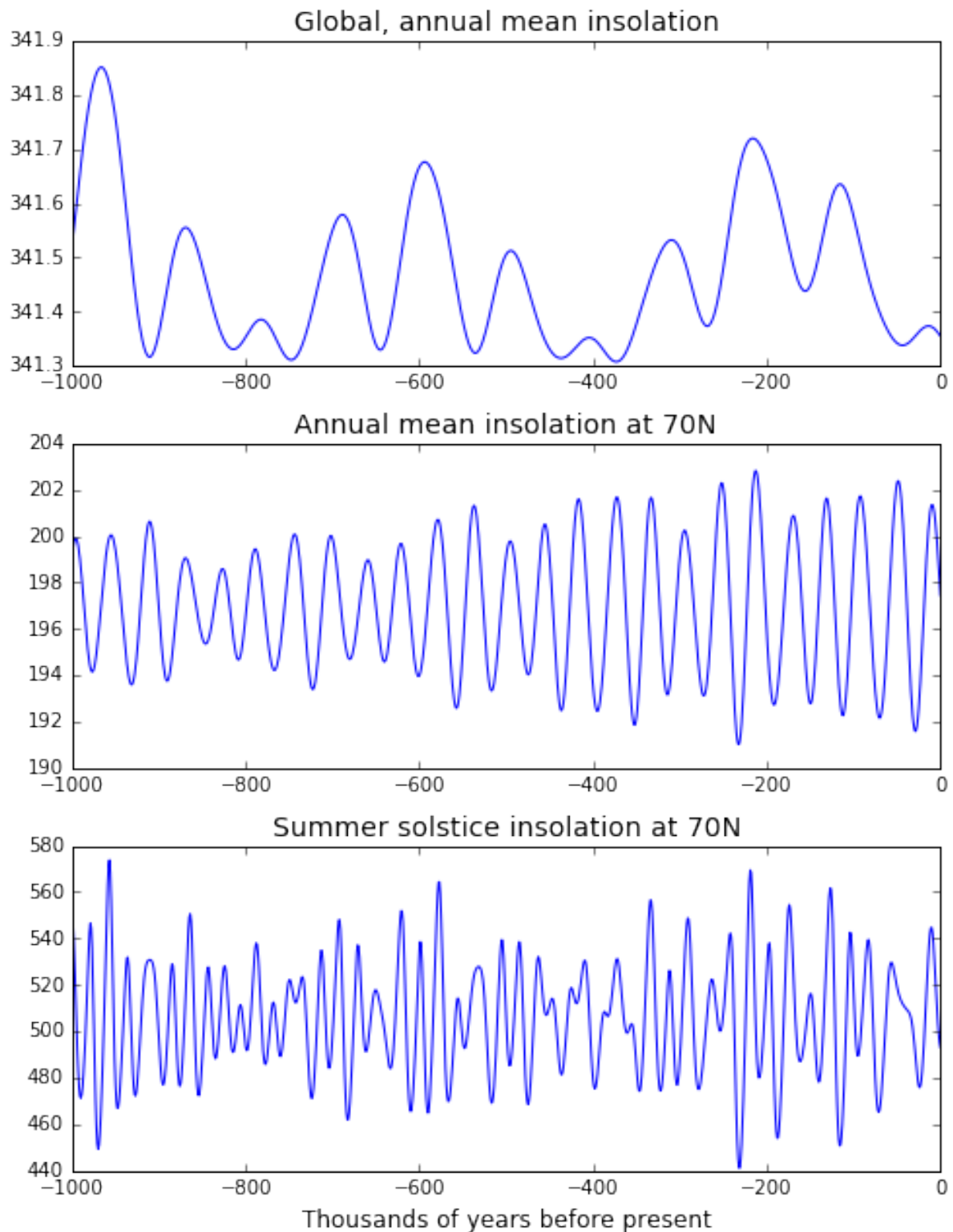
In [10]: Qann = np.mean(Q, axis=1) # time average over the year
         print Qann.shape
         Qglobal = np.empty_like( kyears )
         for n in range( kyears.size ): # global area-weighted average
             Qglobal[n] = np.sum( Qann[:,n] * np.cos( np.deg2rad(lat) ) ) \
```

```
                / np.sum( np.cos( np.deg2rad(lat) ) )  
    print Qglobal.shape  
(181, 1001)  
(1001,)
```

We are going to create a figure showing past time variations in three quantities:

1. Global, annual mean insolation
2. Annual mean insolation at high northern latitudes
3. Summer solstice insolation at high northern latitudes

```
In [32]: fig = plt.figure( figsize = (7,9) , dpi=80 )  
  
    ax1 = fig.add_subplot(3,1,1)  
    ax1.plot( kyears, Qglobal )  
    ax1.set_title('Global, annual mean insolation', fontsize=14 )  
    ax1.ticklabel_format( useOffset=False )  
  
    ax2 = fig.add_subplot(3,1,2)  
    ax2.plot( kyears, Qann[160,:] )  
    ax2.set_title('Annual mean insolation at 70N', fontsize=14 )  
  
    ax3 = fig.add_subplot(3,1,3)  
    ax3.plot( kyears, Q[160,23,:] )  
    ax3.set_xlabel( 'Thousands of years before present', fontsize=12 )  
    ax3.set_title('Summer solstice insolation at 70N', fontsize=14 )  
  
    plt.tight_layout()  
    plt.show()
```



And comparing with the plots of orbital variations above, we see that

1. Global annual mean insolation variations on with eccentricity (slow), and the variations are very small!
2. Annual mean insolation varies with obliquity (medium). Annual mean insolation does NOT depend on precession!
3. Summer solstice insolation at high northern latitudes is affected by both precession and obliquity. The variations are large.



## Insolation changes between the Last Glacial Maximum and the end of the last ice age

Last Glacial Maximum or “LGM” occurred around 23,000 years before present, when the ice sheets were at their greatest extent. By 10,000 years ago, the ice sheets were mostly gone and the last ice age was over. Let’s plot the changes in the seasonal distribution of insolation from 23 kyrs to 10 kyrs.

```
In [12]: # present-day orbital parameters
orb_0 = table.lookup_parameters( 0 )

# orbital parameters for 10 kyrs before present
orb_10 = table.lookup_parameters( -10 )

# 23 kyrs before present
orb_23 = table.lookup_parameters( -23 )

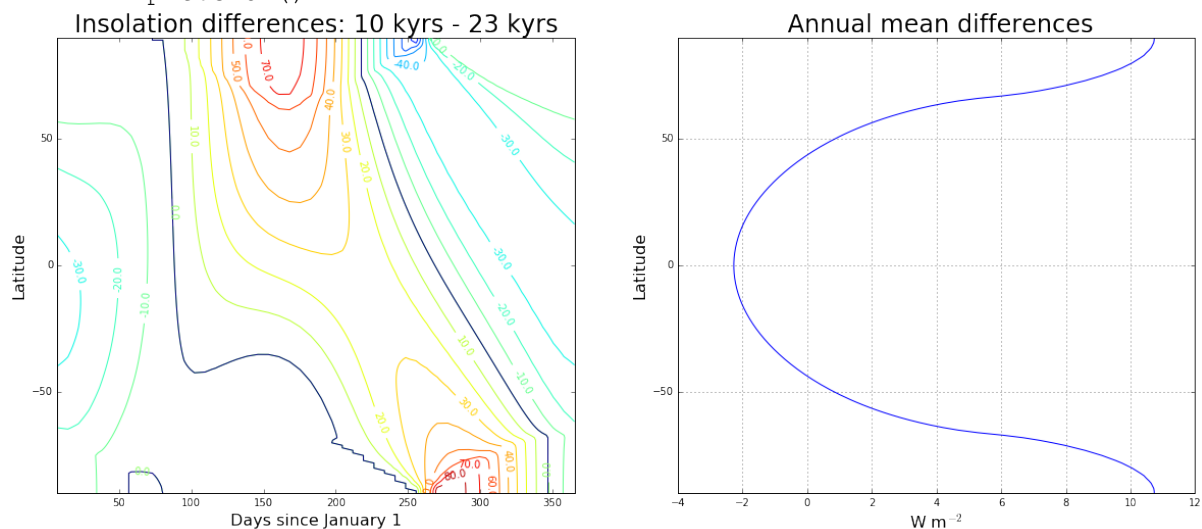
# insolation arrays for each of the three sets of orbital parameters
Q_0 = daily_insolation( lat, days, orb_0 )
Q_10 = daily_insolation( lat, days, orb_10 )
Q_23 = daily_insolation( lat, days, orb_23 )

In [13]: fig = plt.figure( figsize=(20,8) )

ax1 = fig.add_subplot(1,2,1)
Qdiff = Q_10 - Q_23
CS1 = ax1.contour( days, lat, Qdiff, levels = np.arange(-100., 100., 10.) )
ax1.clabel(CS1, CS1.levels, inline=True, fmt='%r', fontsize=10)
ax1.contour( days, lat, Qdiff, levels = [0.], color = 'k' )
ax1.set_xlabel('Days since January 1', fontsize=16 )
ax1.set_ylabel('Latitude', fontsize=16 )
ax1.set_title('Insolation differences: 10 kyrs - 23 kyrs', fontsize=24 )

ax2 = fig.add_subplot(1,2,2)
ax2.plot( np.mean( Qdiff, axis=1 ), lat )
ax2.set_xlabel('W m-2', fontsize=16 )
ax2.set_ylabel('Latitude', fontsize=16 )
ax2.set_title('Annual mean differences', fontsize=24 )
ax2.set_ylim((-90,90))
ax2.grid()

plt.show()
```



The annual mean plot shows a classic obliquity signal: at 10 kyrs, the axis close to its maximum tilt, around 24.2°. At 23 kyrs, the tilt was much weaker, only about 22.7°. In the annual mean, a stronger tilt means more sunlight to

the poles and less to the equator. This is very helpful if you are trying to melt an ice sheet.

Finally, take the global average of the difference:

```
In [14]: print np.sum( np.mean(Qdiff,axis=1) * np.cos( np.deg2rad(lat) ) ) \
          / np.sum( np.cos(np.deg2rad(lat)) )

0.00651043078327
```

This confirms that the difference is tiny (and due to very small changes in the eccentricity). **Ice ages are driven by seasonal and latitudinal redistributions of solar energy**, NOT by changes in the total global amount of solar energy!

## 5.5 The seasonal cycle of surface temperature

Look at the observed seasonal cycle in the NCEP reanalysis data.

Read in the necessary data from the online server.

The catalog is here: <http://www.esrl.noaa.gov/psd/thredds/dodsC/Datasets/ncep.reanalysis.derived/catalog.html>

```
In [1]: %matplotlib inline
import numpy as np
import matplotlib.pyplot as plt
import netCDF4 as nc
from climlab import constants as const
from climlab.model import ebm
from climlab.solar.insolation import daily_insolation

datapath = "http://ramadda.atmos.albany.edu:8080/repository/.opendap/latest/Top/U
endstr = "/entry.das"

In [2]: ncep_url = "http://www.esrl.noaa.gov/psd/thredds/dodsC/Datasets/ncep.reanalysis.
ncep_air = nc.Dataset( ncep_url + "pressure/air.mon.1981-2010.ltm.nc" )
ncep_Ts = nc.Dataset( ncep_url + "surface_gauss/skt.sfc.mon.1981-2010.ltm.nc" )
lat_ncep = ncep_Ts.variables['lat'][:]; lon_ncep = ncep_Ts.variables['lon'][:];
Ts_ncep = ncep_Ts.variables['skt'][:];
print Ts_ncep.shape

(12, 94, 192)
```

Load the topography file from CESM, just so we can plot the continents.

```
In [3]: topo = nc.Dataset( datapath + 'som_input/USGS-gtopo30_1.9x2.5_remap_c050602.nc'
lat_cesm = topo.variables['lat'][:];
lon_cesm = topo.variables['lon'][:];
```

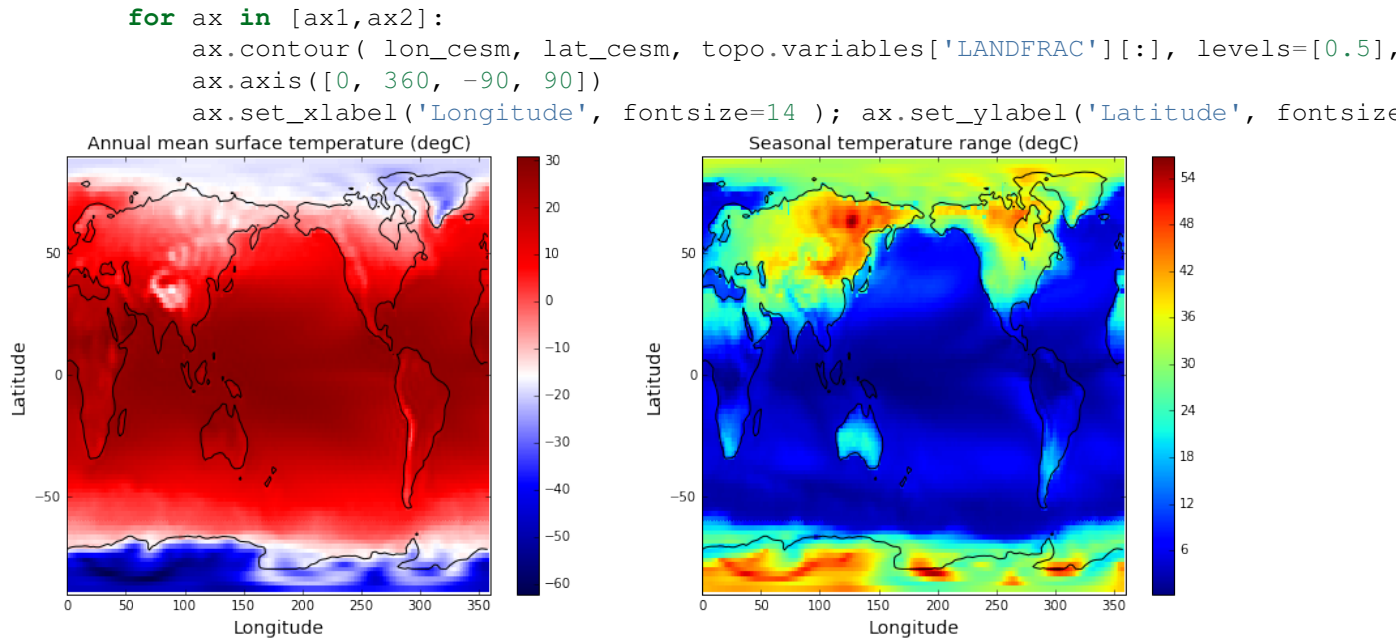
Make two maps: one of annual mean surface temperature, another of the seasonal range (max minus min).

```
In [4]: maxTs = np.max(Ts_ncep,axis=0)
minTs = np.min(Ts_ncep,axis=0)

In [5]: fig = plt.figure( figsize=(16,6) )

ax1 = fig.add_subplot(1,2,1)
cax1 = ax1.pcolormesh( lon_ncep, lat_ncep, np.mean(Ts_ncep, axis=0), cmap=plt.cm
cbar1 = plt.colorbar(cax1)
ax1.set_title('Annual mean surface temperature (degC)', fontsize=14 )

ax2 = fig.add_subplot(1,2,2)
cax2 = ax2.pcolormesh( lon_ncep, lat_ncep, maxTs - minTs )
cbar2 = plt.colorbar(cax2)
ax2.set_title('Seasonal temperature range (degC)', fontsize=14)
```



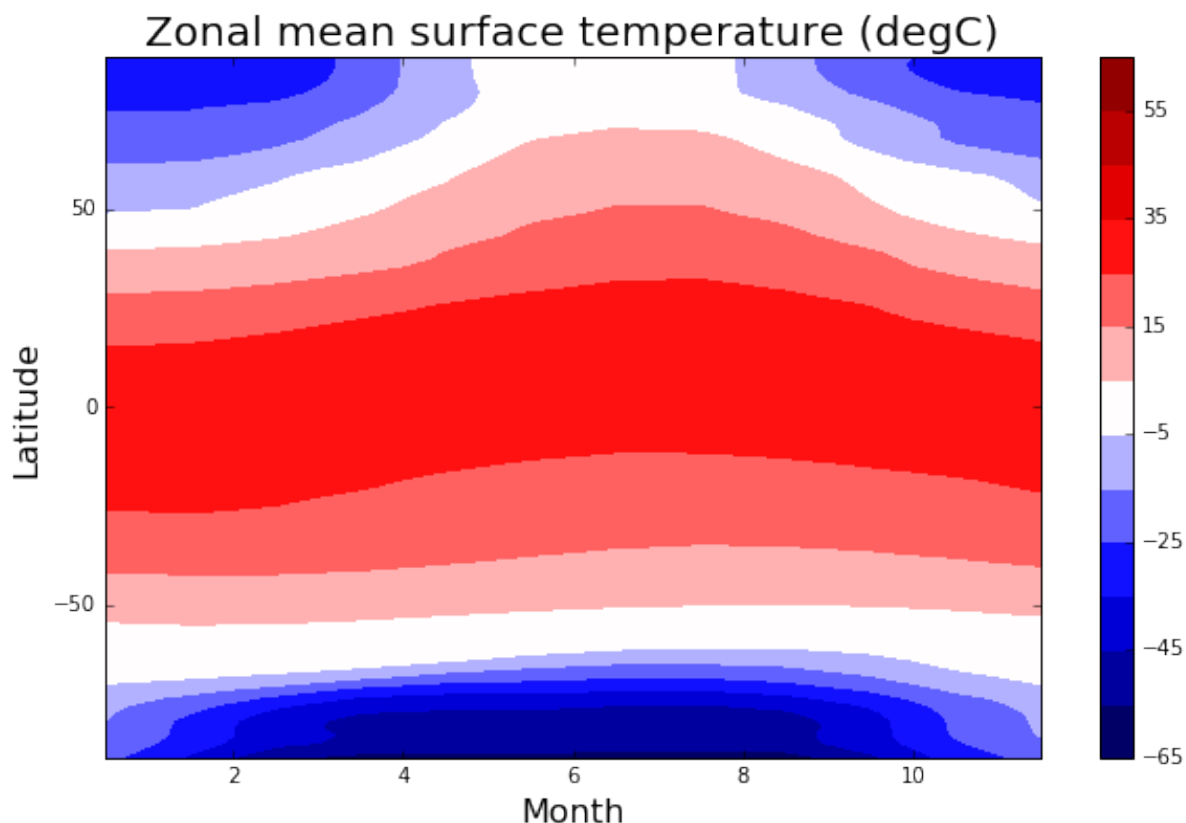
Make a contour plot of the zonal mean temperature as a function of time of year

```

In [6]: [-65, -55, -45, -35, -25, -15, -5, 5, 15, 25, 35, 45, 55, 65]
Out[6]: [-65, -55, -45, -35, -25, -15, -5, 5, 15, 25, 35, 45, 55, 65]
In [7]: np.arange(-65, 75, 10)
Out[7]: array([-65, -55, -45, -35, -25, -15, -5, 5, 15, 25, 35, 45, 55, 65])
In [8]: Tmax = 65; Tmin = -Tmax; delT = 10
        clevels = np.arange(Tmin, Tmax+delT, delT)

fig_zonobs = plt.figure( figsize=(10,6) )
ax = fig_zonobs.add_subplot(111)
cax = ax.contourf( np.arange(12)+0.5, lat_ncep, np.transpose(np.mean( Ts_ncep, a
        cmap=plt.cm.seismic, vmin=Tmin, vmax=Tmax )
ax.set_xlabel('Month', fontsize=16)
ax.set_ylabel('Latitude', fontsize=16 )
cbar = plt.colorbar(cax)
ax.set_title('Zonal mean surface temperature (degC)', fontsize=20)
plt.show()

```



### 5.5.1 Exploring the amplitude of the seasonal cycle with an EBM

We are looking at the 1D (zonally averaged) energy balance model with diffusive heat transport. The equation is

$$C \frac{\partial T(\phi, t)}{\partial t} = (1 - \alpha) Q(\phi, t) - (A + BT(\phi, t)) + \frac{K}{\cos \phi} \frac{\partial}{\partial \phi} \left( \cos \phi \frac{\partial T}{\partial \phi} \right)$$

and the code in `climlab.model.ebm.py` solves this equation numerically.

One handy feature of `climlab` process code: the function `integrate_years()` automatically calculates the time averaged temperature. So if we run it for exactly one year, we get the annual mean temperature saved in the field `T_timeave`.

We will look at the seasonal cycle of temperature in three different models with different heat capacities (which we express through an equivalent depth of water in meters):

```
In [9]: model1 = ebm.EBM_seasonal()
        model1.integrate_years(1, verbose=True)

        water_depths = np.array([2., 10., 50.])

        num_depths = water_depths.size
        Tann = np.empty( [model1.lat.size, num_depths] )
        models = []

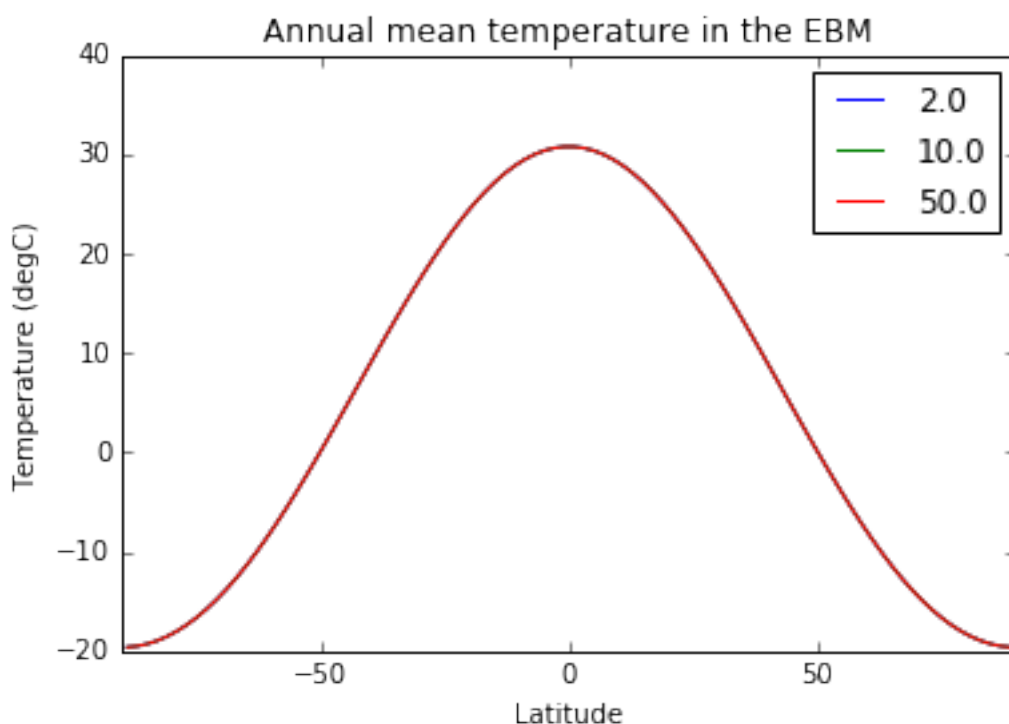
        for n in range(num_depths):
            models.append(ebm.EBM_seasonal(water_depth=water_depths[n]))
            models[n].integrate_years(20., verbose=False)
            models[n].integrate_years(1., verbose=False)
            Tann[:,n] = np.squeeze(models[n].timeave['Ts'])
```

Integrating for 90 steps, 365.2422 days, or 1 years.  
Total elapsed time is 1.0 years.

All models should have the same annual mean temperature:

```
In [10]: lat = modell.lat

plt.plot(lat, Tann)
plt.xlim(-90,90)
plt.xlabel('Latitude')
plt.ylabel('Temperature (degC)')
plt.title('Annual mean temperature in the EBM')
plt.legend( water_depths.astype(str) )
plt.show()
```



There is no automatic function in the `ebm.py` code to keep track of minimum and maximum temperatures (though we might add that in the future!)

Instead we'll step through one year "by hand" and save all the temperatures.

```
In [11]: num_steps_per_year = int(modell.time['num_steps_per_year'])
Tyear = np.empty((lat.size, num_steps_per_year, num_depths))
for n in range(num_depths):
    for m in range(num_steps_per_year):
        models[n].step_forward()
        Tyear[:,m,n] = np.squeeze(models[n].state['Ts'])
```

Make a figure to compare the observed zonal mean seasonal temperature cycle to what we get from the EBM with different heat capacities:

```
In [12]: fig = plt.figure( figsize=(20,10) )

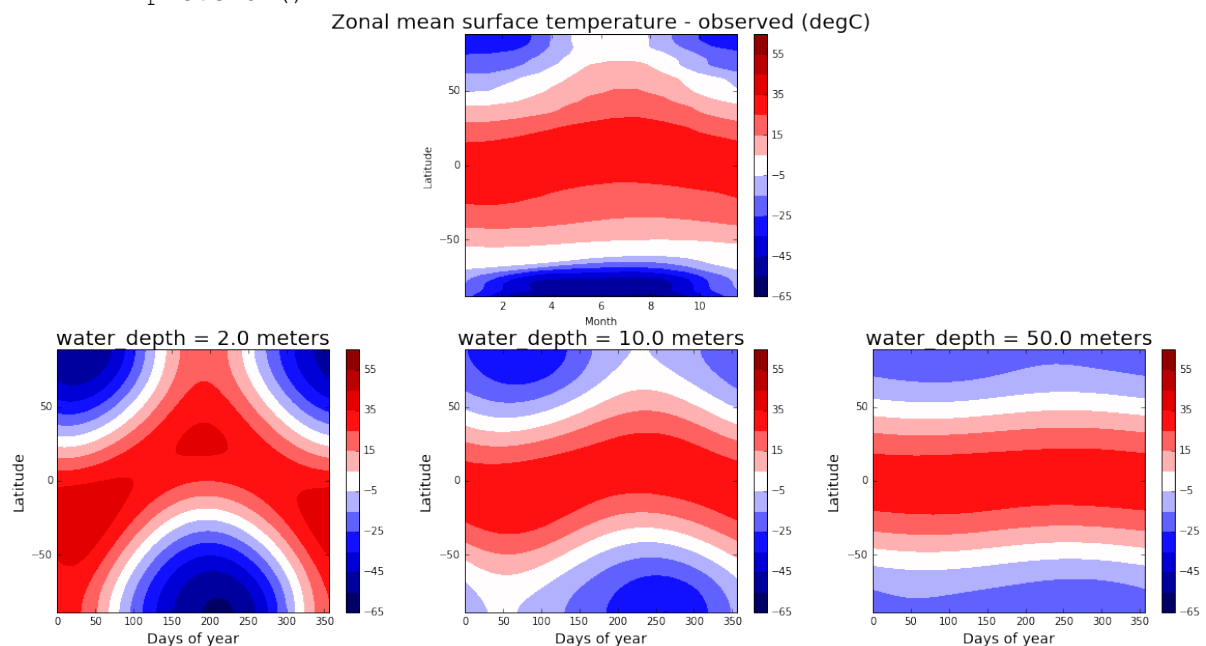
ax = fig.add_subplot(2,num_depths,2)
cax = ax.contourf( np.arange(12)+0.5, lat_ncep, np.transpose(np.mean( Ts_ncep,
                                                                    cmap=plt.cm.seismic, vmin=Tmin, vmax=Tmax )
ax.set_xlabel('Month')
ax.set_ylabel('Latitude')
cbar = plt.colorbar(cax)
ax.set_title('Zonal mean surface temperature - observed (degC)', fontsize=20)
```

```

for n in range(num_depths):
    ax = fig.add_subplot(2,num_depths,num_depths+n+1)
    cax = ax.contourf( 4*np.arange(num_steps_per_year), lat, Tyear[:, :, n], levels=levels,
                      cmap=plt.cm.seismic, vmin=Tmin, vmax=Tmax )
    cbar1 = plt.colorbar(cax)
    ax.set_title('water_depth = ' + str(models[n].param['water_depth']) + ' meters')
    ax.set_xlabel('Days of year', fontsize=14 )
    ax.set_ylabel('Latitude', fontsize=14 )

#fig.set_title('Temperature in seasonal EBM with various water depths', fontsize=14)
plt.show()

```



Which one looks more realistic? Depends a bit on where you look. But overall, the observed seasonal cycle matches the 10 meter case best. The effective heat capacity governing the seasonal cycle of the zonal mean temperature is closer to 10 meters of water than to either 2 or 50 meters.

## Making an animation of the EBM solutions

```
In [13]: fpath = '/Users/Brian/Dropbox/PythonStuff/ebm_seasonal_frames/'
```

```

fig = plt.figure( figsize=(20,5) )
#for m in range(model2.time['num_steps_per_year']):
for m in range(1):
    thisday = m * models[0].param['timestep'] / const.seconds_per_day
    Q = daily_insolation( lat, thisday )
    for n in range(num_depths):
        c1 = 'b'
        ax = fig.add_subplot(1,num_depths,n+1)
        ax.plot( lat, Tyear[:, m, n], c1 )
        ax.set_title('water_depth = ' + str(models[n].param['water_depth']) + ' meters')
        ax.set_xlabel('Latitude', fontsize=14 )
        if n is 0:
            ax.set_ylabel('Temperature', fontsize=14, color=c1 )
        ax.set_xlim([-90, 90])
        ax.set_ylim([-60, 60])
        for tl in ax.get_yticklabels():
            tl.set_color(c1)

```

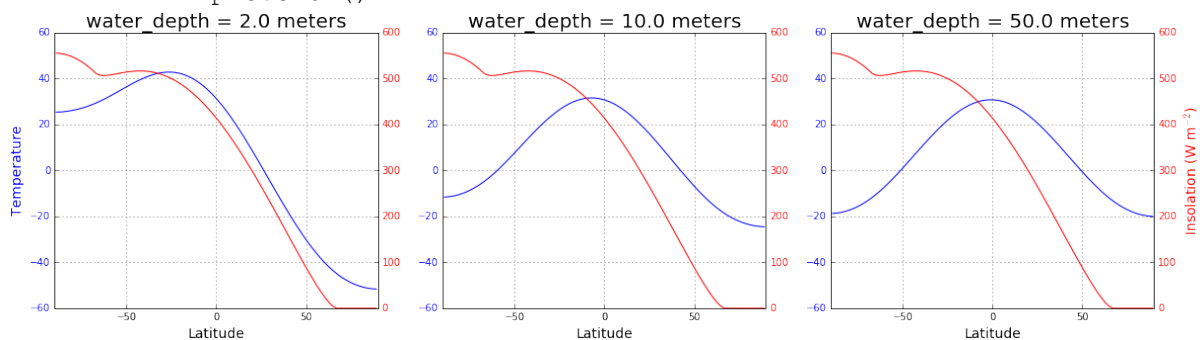
```

ax.grid()

c2 = 'r'
ax2 = ax.twinx()
ax2.plot( lat, Q, c2)
if n is 2:
    ax2.set_ylabel('Insolation (W m$^{-2}$)', color=c2, fontsize=14)
for t1 in ax2.get_yticklabels():
    t1.set_color(c2)
ax2.set_xlim([-90,90])
ax2.set_ylim([0,600])

filename = fpath + 'ebm_seasonal' + str(m).zfill(4) + '.png'
#fig.savefig(filename)
#fig.clear()
plt.show()

```



## 5.5.2 The seasonal cycle for a planet with 90° obliquity

The EBM code uses our familiar `insolation.py` code to calculate insolation, and therefore it's easy to set up a model with different orbital parameters. Here is an example with **very** different orbital parameters: 90° obliquity. We looked at the distribution of insolation by latitude and season for this type of planet in the last homework.

```

In [14]: orb_highobl = {'ecc':0., 'obliquity':90., 'long_peri':0.}
        print orb_highobl
        model_highobl = ebm.EBM_seasonal(orb=orb_highobl)
        print model_highobl.param['orb']

```

```

'long_peri': 0.0, 'ecc': 0.0, 'obliquity': 90.0
'long_peri': 0.0, 'ecc': 0.0, 'obliquity': 90.0

```

Repeat the same procedure to calculate and store temperature throughout one year, after letting the models run out to equilibrium.

```

In [15]: Tann_highobl = np.empty( [lat.size, num_depths] )
        models_highobl = []

        for n in range(num_depths):
            models_highobl.append(ebm.EBM_seasonal(water_depth=water_depths[n], orb=orb_highobl))
            models_highobl[n].integrate_years(40., verbose=False)
            models_highobl[n].integrate_years(1.)
            Tann_highobl[:,n] = np.squeeze(models_highobl[n].timeave['Ts'])

        Tyear_highobl = np.empty([lat.size, num_steps_per_year, num_depths])
        for n in range(num_depths):
            for m in range(num_steps_per_year):
                models_highobl[n].step_forward()
                Tyear_highobl[:,m,n] = np.squeeze(models_highobl[n].state['Ts'])

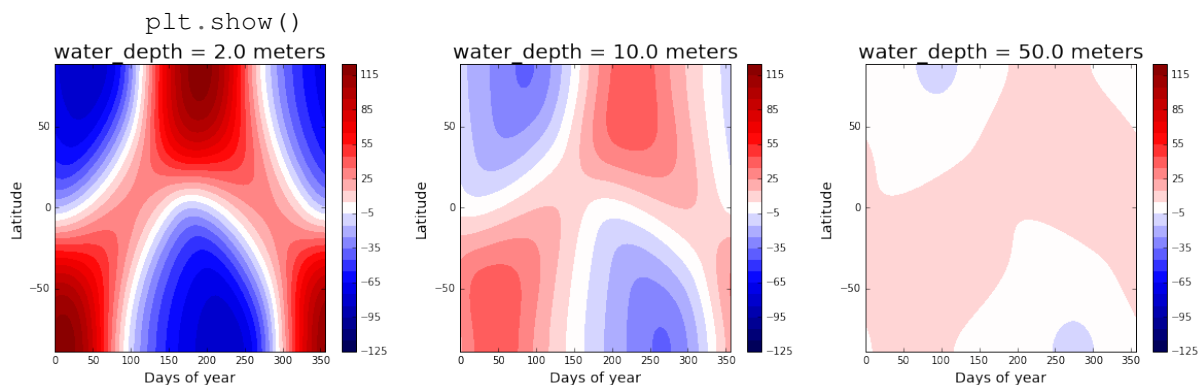
```

```
Integrating for 90 steps, 365.2422 days, or 1.0 years.
Total elapsed time is 41.0 years.
Integrating for 90 steps, 365.2422 days, or 1.0 years.
Total elapsed time is 41.0 years.
Integrating for 90 steps, 365.2422 days, or 1.0 years.
Total elapsed time is 41.0 years.
```

And plot the seasonal temperature cycle same as we did above:

```
In [16]: fig = plt.figure( figsize=(20,5) )
        Tmax_highobl = 125; Tmin_highobl = -Tmax_highobl; delT_highobl = 10
        clevels_highobl = np.arange(Tmin_highobl, Tmax_highobl+delT_highobl, delT_highobl)

        for n in range(num_depths):
            ax = fig.add_subplot(1,num_depths,n+1)
            cax = ax.contourf( 4*np.arange(num_steps_per_year), lat, Tyear_highobl[:,n],
                             levels=clevels_highobl, cmap=plt.cm.seismic, vmin=Tmin_highobl, vmax=Tmax_highobl)
            cbar1 = plt.colorbar(cax)
            ax.set_title('water_depth = ' + str(models_highobl[n].param['water_depth']))
            ax.set_xlabel('Days of year', fontsize=14 )
            ax.set_ylabel('Latitude', fontsize=14 )
```



Note that the temperature range is much larger than for the Earth-like case above (but same contour interval, 10 degC).

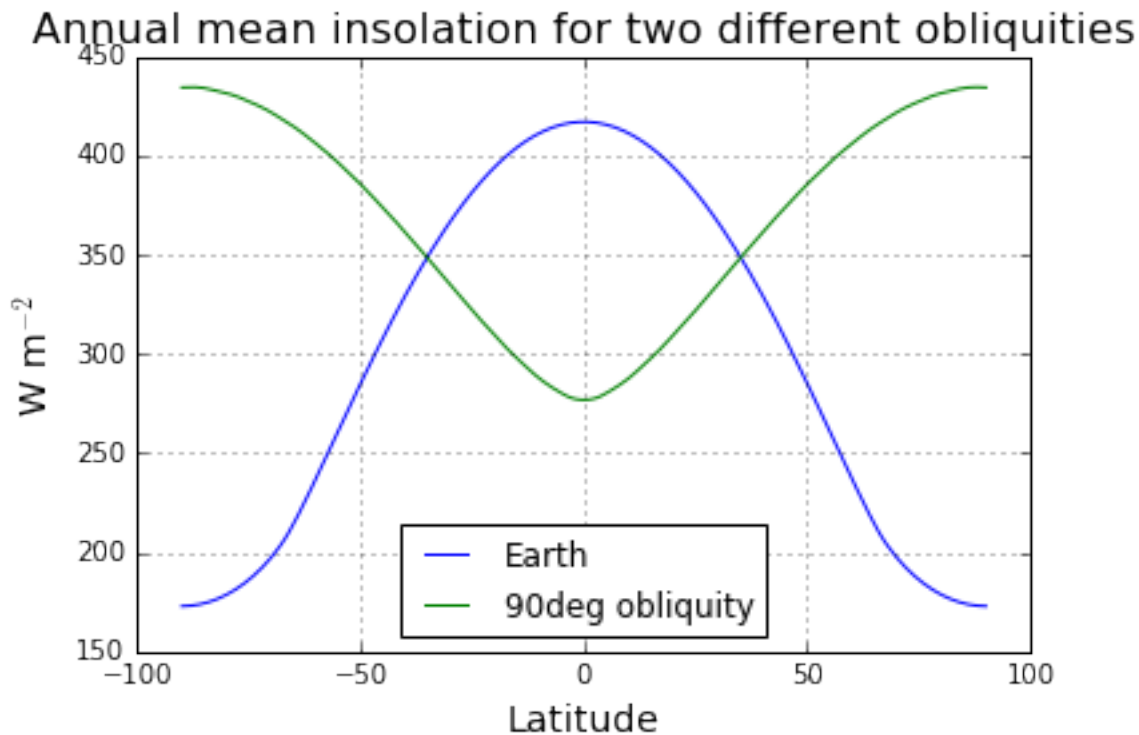
Why is the temperature so uniform in the north-south direction with 50 meters of water?

To see the reason, let's plot the annual mean insolation at 90° obliquity, alongside the present-day annual mean insolation:

```
In [17]: lat2 = np.linspace(-90, 90, 181)
        days = np.linspace(1., 50.) / 50 * const.days_per_year
        Q_present = daily_insolation( lat2, days )
        Q_highobl = daily_insolation( lat2, days, orb_highobl )
        Q_present_ann = np.mean( Q_present, axis=1 )
        Q_highobl_ann = np.mean( Q_highobl, axis=1 )

In [18]: plt.plot( lat2, Q_present_ann, label='Earth' )
        plt.plot( lat2, Q_highobl_ann, label='90deg obliquity' )
        plt.grid()
        plt.legend(loc='lower center')
        plt.xlabel('Latitude', fontsize=14 )
        plt.ylabel('W m$^{-2}$', fontsize=14 )
        plt.title('Annual mean insolation for two different obliquities', fontsize=16)
        plt.show()
```





Though this is a bit misleading, because our model prescribes an increase in albedo from the equator to the pole. So the absorbed shortwave gradients look even more different.

In [ ]:

## 5.6 Ice - Albedo Feedback and runaway glaciation

Here we will use the 1-dimensional diffusive Energy Balance Model (EBM) to explore the effects of albedo feedback and heat transport on climate sensitivity.

```
In [1]: %matplotlib inline
import numpy as np
import matplotlib.pyplot as plt
import climlab
from climlab import constants as const
from climlab import legendre
from climlab.domain.field import global_mean
```

### 5.6.1 Annual-mean model with albedo feedback: adjustment to equilibrium

A version of the EBM in which albedo adjusts to the current position of the ice line, wherever  $T < T_f$

```
In [2]: modell = climlab.EBM_annual( num_points = 180, a0=0.3, a2=0.078, ai=0.62)
print modell
```

climlab Process of type <class 'climlab.model.ebm.EBM\_annual'>.

State variables and domain shapes:

Ts: (90, 1)

The subprocess tree:

```
top: <class 'climlab.model.ebm.EBM_annual'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
```

```
albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
iceline: <class 'climlab.surface.albedo.Iceline'>
cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
insolation: <class 'climlab.radiation.insolation.AnnualMeanInsolation'>
```

```
In [3]: modell.integrate_years(5)
        Tequil = np.array(modell.Ts)
        ALBequil = np.array(modell.albedo)
        OLRequil = np.array(modell.OLR)
        ASRequil = np.array(modell.ASR)
```

Integrating for 450 steps, 1826.211 days, or 5 years.  
Total elapsed time is 5.0 years.

Let's look at what happens if we perturb the temperature – make it 20°C colder everywhere!

```
In [4]: modell.Ts -= 20.
        modell.compute_diagnostics()
```

Let's take a look at how we have just perturbed the absorbed shortwave:

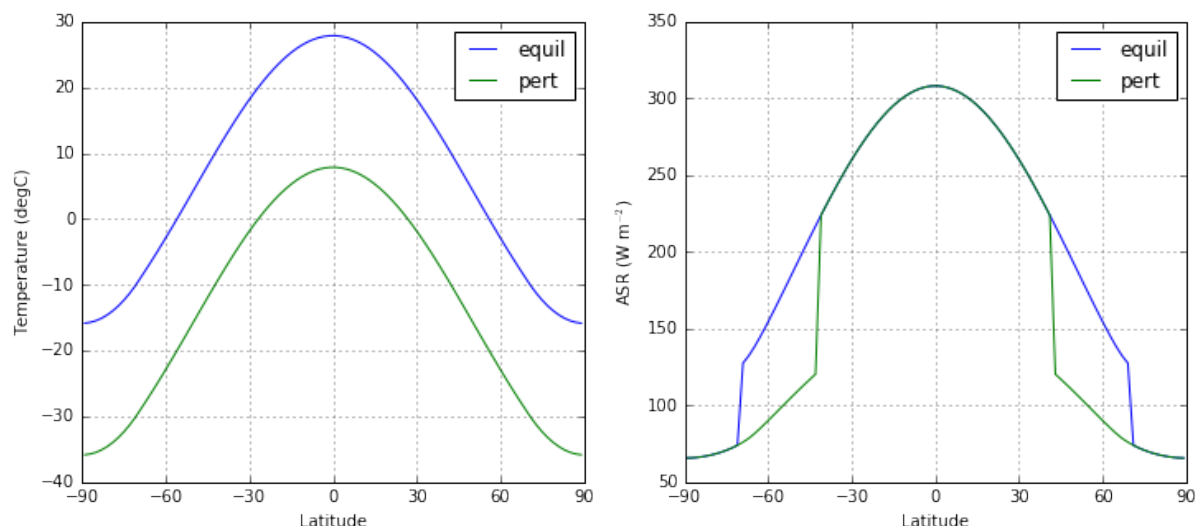
```
In [5]: my_ticks = [-90, -60, -30, 0, 30, 60, 90]
        lat = modell.domains['Ts'].axes['lat'].points

        fig = plt.figure( figsize=(12,5) )

        ax1 = fig.add_subplot(1,2,1)
        ax1.plot(lat, Tequil, label='equil')
        ax1.plot(lat, modell.Ts, label='pert' )
        ax1.grid()
        ax1.legend()
        ax1.set_xlim(-90,90)
        ax1.set_xticks(my_ticks)
        ax1.set_xlabel('Latitude')
        ax1.set_ylabel('Temperature (degC)')

        ax2 = fig.add_subplot(1,2,2)
        ax2.plot( lat, ASRequil, label='equil')
        ax2.plot( lat, modell.ASR, label='pert' )
        ax2.grid()
        ax2.legend()
        ax2.set_xlim(-90,90)
        ax2.set_xticks(my_ticks)
        ax2.set_xlabel('Latitude')
        ax2.set_ylabel('ASR (W m$^{-2}$)')

        plt.show()
```



So there is less absorbed shortwave now, because of the increased albedo. The global mean difference is:

```
In [6]: global_mean( modell.ASR - ASRequil )
```

```
Out[6]: Field(-20.37046205447195)
```

Less shortwave means that there is a tendency for the climate to cool down even more! In other words, the shortwave feedback is **positive**.

Recall that the net feedback for the EBM can be written

$$\lambda = -B + \frac{\Delta\langle(1-\alpha)Q\rangle}{\Delta\langle T\rangle}$$

where the second term is the change in the absorbed shortwave per degree global mean temperature change.

Plugging these numbers in gives

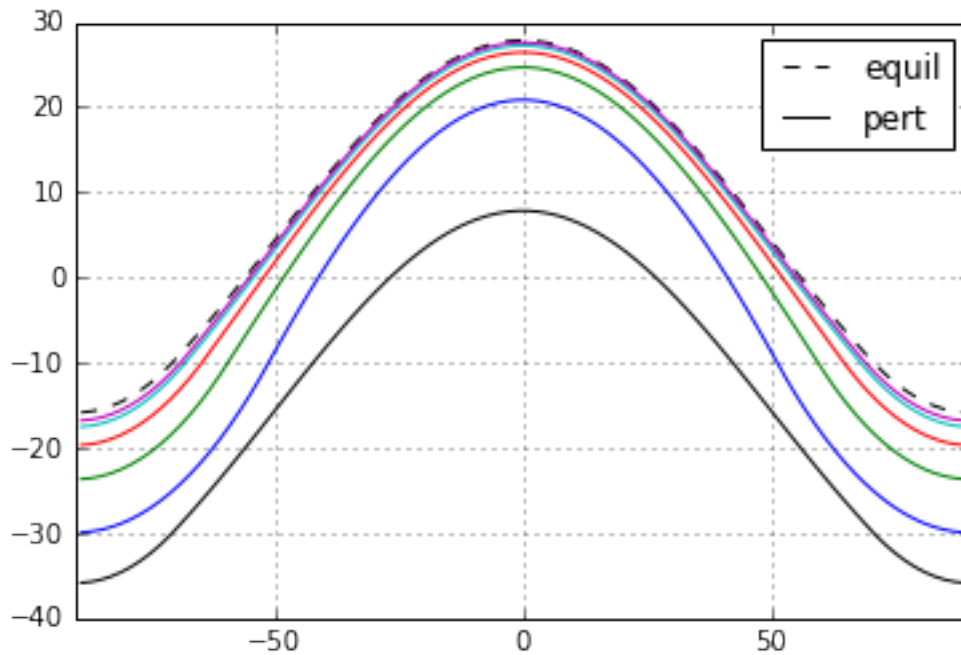
$$\lambda = -2 + \frac{-20.4}{-20} = -2 + 1 = -1 \text{ W m}^{-2} \text{ } ^\circ\text{C}^{-1}$$

The feedback is negative, as we expect! The tendency to warm up from reduced OLR outweighs the tendency to cool down from reduced ASR. A negative net feedback means that the system will relax back towards the equilibrium.

Let's let the temperature evolve one year at a time and add extra lines to the graph:

```
In [7]: plt.plot( lat, Tequil, 'k--', label='equil' )
        plt.plot( lat, modell.Ts, 'k-', label='pert' )
        plt.grid()
        plt.xlim(-90,90)
        plt.legend()

        for n in range(5):
            modell.integrate_years(years=1.0, verbose=False)
            plt.plot(lat, modell.Ts)
```



Temperature drifts back towards equilibrium, as we expected!

What if we cool the climate **so much** that the entire planet is ice covered?

```
In [8]: modell.Ts -= 40.  
        modell.compute_diagnostics()
```

Look again at the change in absorbed shortwave:

```
In [9]: global_mean( modell.ASR - ASRequil )  
Out[9]: Field(-108.99200830729608)
```

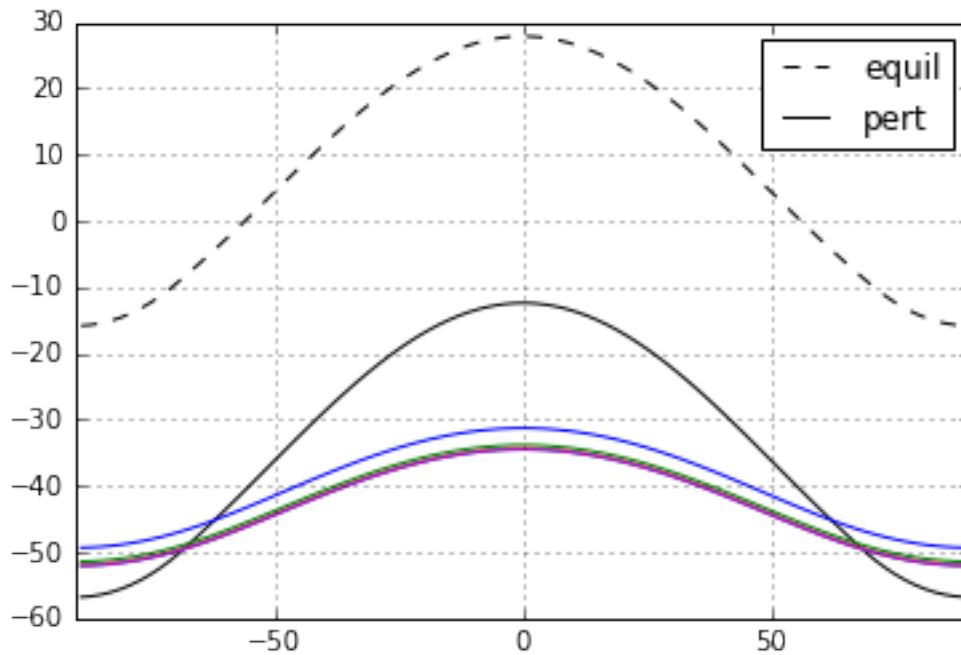
It's much larger because we've covered so much more surface area with ice!

The feedback calculation now looks like

$$\lambda = -2 + \frac{-109}{-40} = -2 + 2.7 = +0.7 \text{ W m}^{-2} \text{ }^{\circ}\text{C}^{-1}$$

What? Looks like the **positive** albedo feedback is so strong here that it has outweighed the **negative** longwave feedback. What will happen to the system now? Let's find out...

```
In [10]: plt.plot( lat, Tequil, 'k--', label='equil' )  
         plt.plot( lat, modell.Ts, 'k-', label='pert' )  
         plt.grid()  
         plt.xlim(-90,90)  
         plt.legend()  
  
         for n in range(5):  
             modell.integrate_years(years=1.0, verbose=False)  
             plt.plot(lat, modell.Ts)
```



Something **very different** happened! The climate drifted towards an entirely different equilibrium state, in which the entire planet is cold and ice-covered.

We will refer to this as the **SNOWBALL EARTH**.

Note that the warmest spot on the planet is still the equator, but it is now about  $-33^{\circ}\text{C}$  rather than  $+28^{\circ}\text{C}$ !

## 5.6.2 Here Comes the Sun! Where is the ice edge?

The ice edge in our model is always where the temperature crosses  $T_f = -10^{\circ}\text{C}$ . The system is at **equilibrium** when the temperature is such that there is a balance between ASR, OLR, and heat transport convergence everywhere.

Suppose that sun was hotter or cooler at different times (in fact it was significantly cooler during early Earth history). That would mean that the solar constant  $S_0 = 4Q$  was larger or smaller. We should expect that the temperature (and thus the ice edge) should increase and decrease as we change  $S_0$ .

$S_0$  during the Neoproterozoic Snowball Earth events is believed to be about 93% of its present-day value, or about  $1270 \text{ W m}^{-2}$ .

We are going to look at how the **equilibrium** ice edge depends on  $S_0$ , by integrating the model out to equilibrium for lots of different values of  $S_0$ . We will start by slowly decreasing  $S_0$ , and then slowly increasing  $S_0$ .

```
In [11]: model2 = climlab.EBM_annual(num_points = 360, a0=0.3, a2=0.078, ai=0.62)
```

```
In [12]: S0array = np.linspace(1400., 1200., 200.)
         #S0array = np.linspace(1400., 1200., 10.)
         #print S0array
```

```
In [13]: model2.integrate_years(5)
```

Integrating for 450 steps, 1826.211 days, or 5 years.  
Total elapsed time is 5.0 years.

```
In [14]: print model2.icelat
```

```
[-70.  70.]
```

```
In [15]: icelat_cooling = np.empty_like(S0array)
         icelat_warming = np.empty_like(S0array)
```

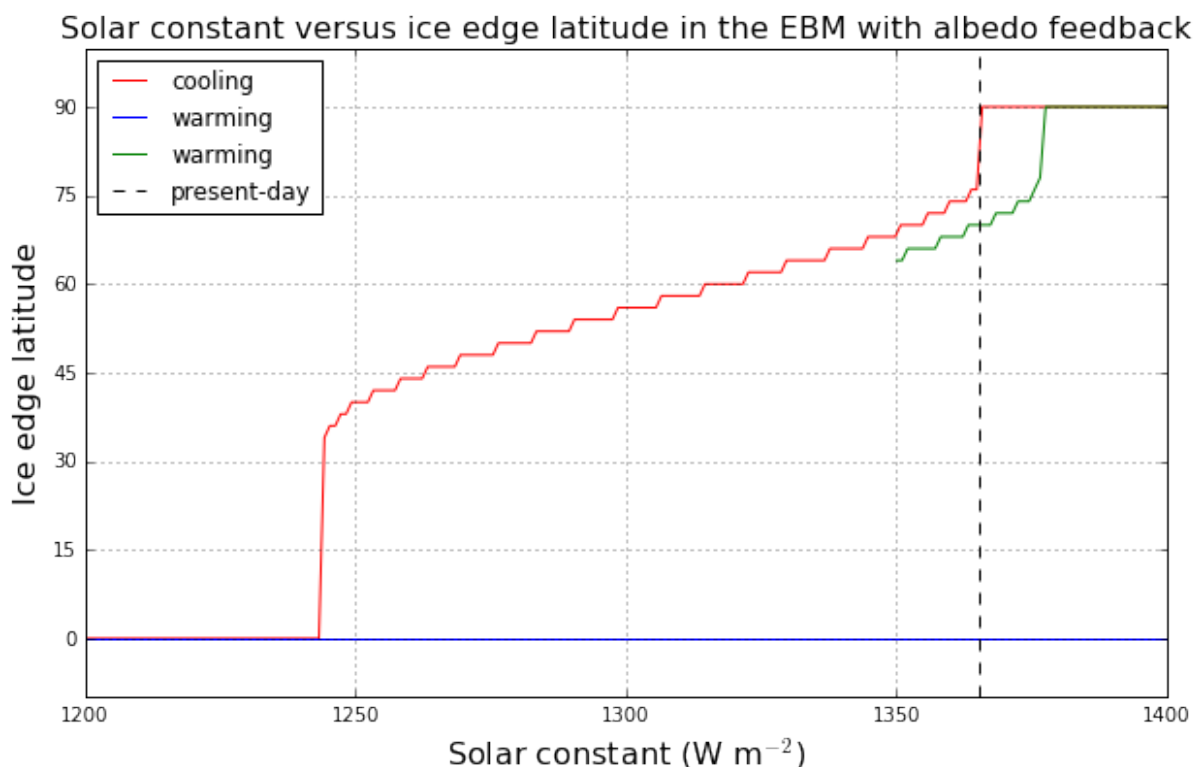
```
In [16]: # First cool....
        for n in range(S0array.size):
            model2.subprocess['insolation'].S0 = S0array[n]
            model2.integrate_years(10, verbose=False)
            icelat_cooling[n] = np.max(model2.icelat)
        # Then warm...
        for n in range(S0array.size):
            model2.subprocess['insolation'].S0 = np.flipud(S0array)[n]
            model2.integrate_years(10, verbose=False)
            icelat_warming[n] = np.max(model2.icelat)
```

For completeness: also start from present-day conditions and warm up.

```
In [17]: model3 = climlab.EBM_annual(num_points = 360, a0=0.3, a2=0.078, ai=0.62)
        S0array3 = np.linspace(1350., 1400., 50.)
        #S0array3 = np.linspace(1350., 1400., 5.)
        icelat3 = np.empty_like(S0array3)
```

```
In [18]: for n in range(S0array3.size):
        model3.subprocess['insolation'].S0 = S0array3[n]
        model3.integrate_years(10, verbose=False)
        icelat3[n] = np.max(model3.icelat)
```

```
In [19]: fig = plt.figure( figsize=(10,6) )
        ax = fig.add_subplot(111)
        ax.plot(S0array, icelat_cooling, 'r-', label='cooling' )
        ax.plot(S0array, icelat_warming, 'b-', label='warming' )
        ax.plot(S0array3, icelat3, 'g-', label='warming' )
        ax.set_ylim(-10,100)
        ax.set_yticks((0,15,30,45,60,75,90))
        ax.grid()
        ax.set_ylabel('Ice edge latitude', fontsize=16)
        ax.set_xlabel('Solar constant (W m$^{-2}$)', fontsize=16)
        ax.plot( [const.S0, const.S0], [-10, 100], 'k--', label='present-day' )
        ax.legend(loc='upper left')
        ax.set_title('Solar constant versus ice edge latitude in the EBM with albedo feedback')
        plt.show()
```



There are actually up to 3 different climates possible for a given value of  $S_0$ !

### How to un-freeze the Snowball

The graph indicates that if the Earth were completely frozen over, it would be perfectly happy to stay that way even if the sun were brighter and hotter than it is today.

Our EBM predicts that (with present-day parameters) the equilibrium temperature at the equator in the Snowball state is about  $-33^\circ\text{C}$ , which is much colder than the threshold temperature  $T_f = -10^\circ\text{C}$ . How can we melt the Snowball?

We need to increase the available energy sufficiently to get the equatorial temperatures above this threshold! That is going to require a much larger increase in  $S_0$  (could also increase the greenhouse gases, which would have a similar effect)!

Let's crank up the sun to  $1830 \text{ W m}^{-2}$  (about a 34% increase from present-day).

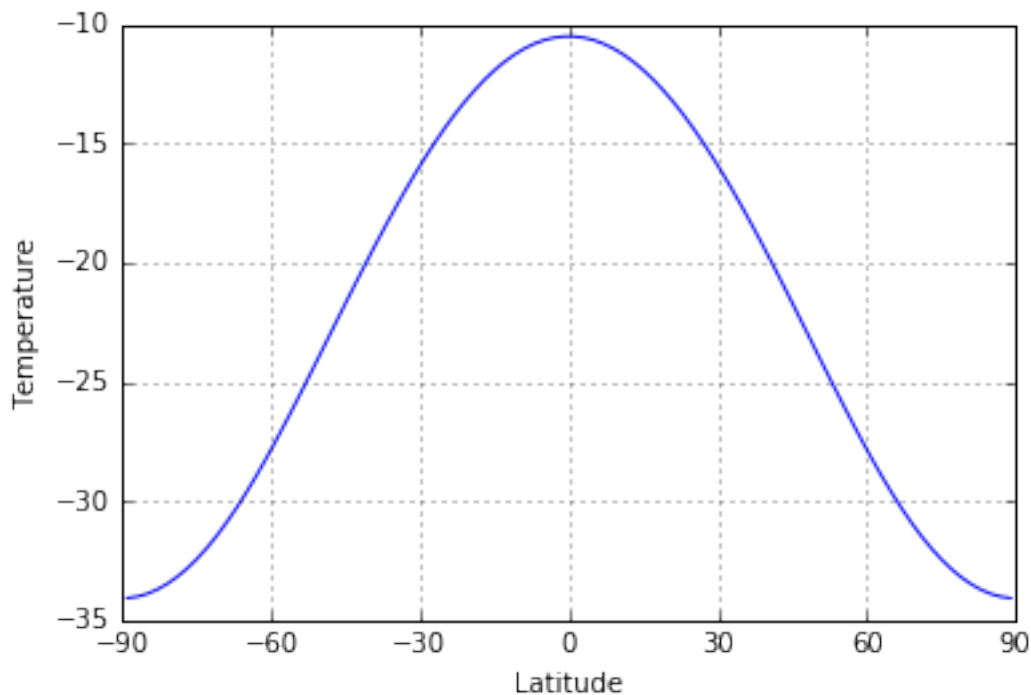
```
In [20]: from climlab.process.process import process_like

model4 = process_like(model2) # initialize with cold Snowball temperature
model4.subprocess['insolation'].S0 = 1830.
model4.integrate_years(40)

#lat = model4.domains['Ts'].axes['lat'].points
plt.plot(model4.lat, model4.Ts)
plt.xlim(-90, 90)
plt.ylabel('Temperature')
plt.xlabel('Latitude')
plt.grid()
plt.xticks(my_ticks)
plt.show()

print('The ice edge is at ' + str(model4.icelat) + 'degrees latitude.')
```

Integrating for 3600 steps, 14609.688 days, or 40 years.  
Total elapsed time is 4044.99999998 years.



The ice edge is at `[-0. 0.]`degrees latitude.

Still a Snowball... but just barely! The temperature at the equator is just below the threshold.

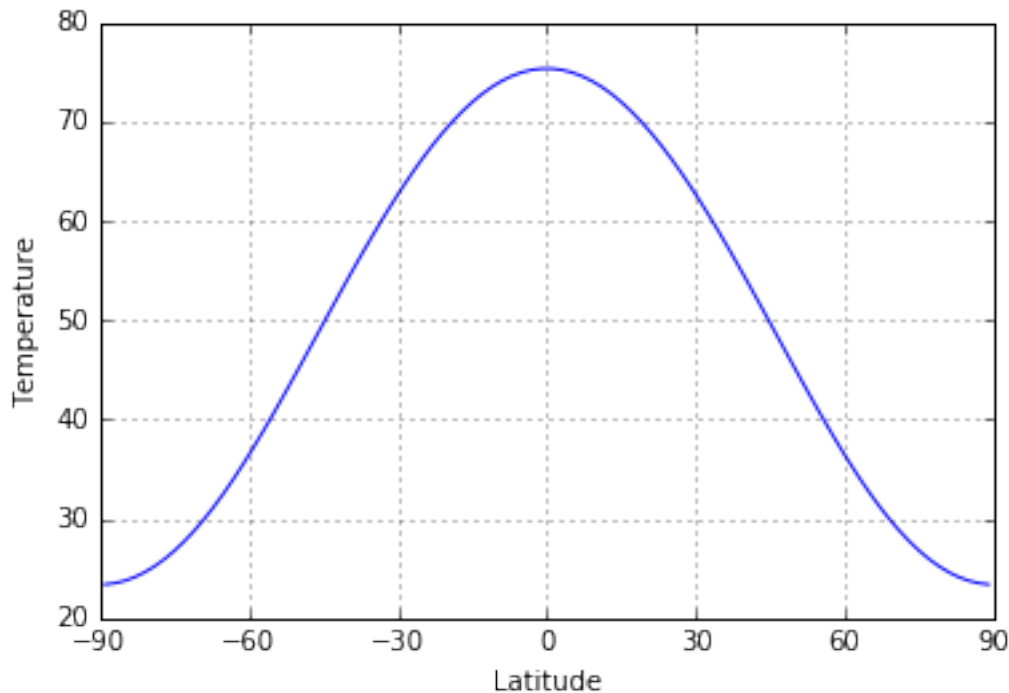
Try to imagine what might happen once it starts to melt. The solar constant is huge, and if it weren't for the highly reflective ice and snow, the climate would be really really hot!

We're going to increase  $S_0$  one more time...

```
In [21]: model4.subprocess['insolation'].S0 = 1845.  
         model4.integrate_years(10)  
  
         plt.plot(lat, model4.state['Ts'])  
         plt.xlim(-90,90)  
         plt.ylabel('Temperature')  
         plt.xlabel('Latitude')  
         plt.grid()  
         plt.xticks(my_ticks)  
         plt.show()
```

Integrating for 900 steps, 3652.422 days, or 10 years.  
Total elapsed time is 4054.99999998 years.





Suddenly the climate looks very very different again! The global mean temperature is

```
In [22]: print( model4.global_mean_temperature() )
58.171701295
```

A roasty 60°C, and the poles are above 20°C. A tiny increase in  $S_0$  has led to a very drastic change in the climate.

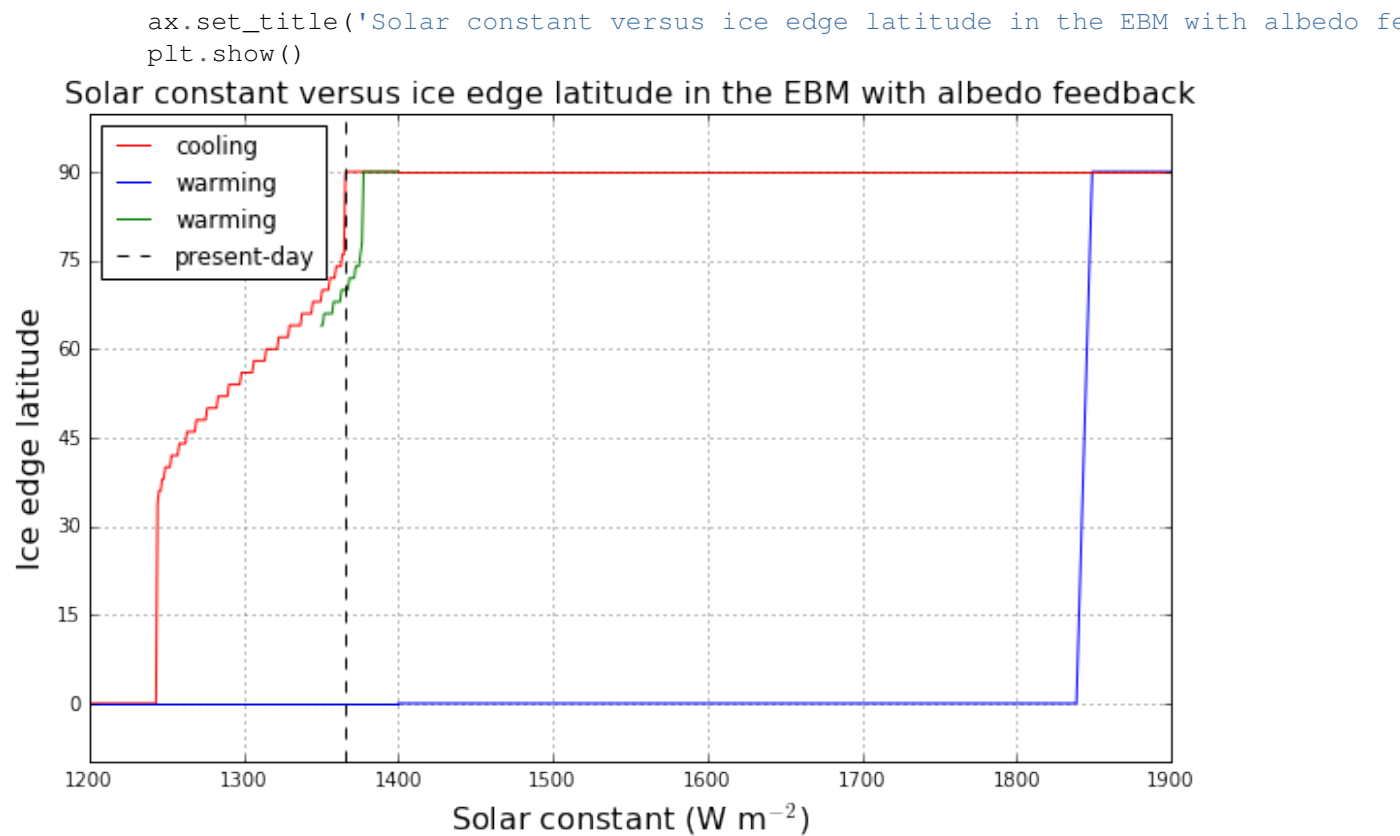
```
In [23]: S0array_snowballmelt = np.linspace(1400., 1900., 50)
        icelat_snowballmelt = np.empty_like(S0array_snowballmelt)
        icelat_snowballmelt_cooling = np.empty_like(S0array_snowballmelt)

        for n in range(S0array_snowballmelt.size):
            model2.subprocess['insolation'].S0 = S0array_snowballmelt[n]
            model2.integrate_years(10, verbose=False)
            icelat_snowballmelt[n] = np.max(model2.diagnostics['icelat'])

        for n in range(S0array_snowballmelt.size):
            model2.subprocess['insolation'].S0 = np.flipud(S0array_snowballmelt)[n]
            model2.integrate_years(10, verbose=False)
            icelat_snowballmelt_cooling[n] = np.max(model2.diagnostics['icelat'])
```

Now we will complete the plot of ice edge versus solar constant.

```
In [24]: fig = plt.figure( figsize=(10,6) )
        ax = fig.add_subplot(111)
        ax.plot(S0array, icelat_cooling, 'r-', label='cooling' )
        ax.plot(S0array, icelat_warming, 'b-', label='warming' )
        ax.plot(S0array3, icelat3, 'g-', label='warming' )
        ax.plot(S0array_snowballmelt, icelat_snowballmelt, 'b-' )
        ax.plot(S0array_snowballmelt, icelat_snowballmelt_cooling, 'r-' )
        ax.set_ylim(-10,100)
        ax.set_yticks((0,15,30,45,60,75,90))
        ax.grid()
        ax.set_ylabel('Ice edge latitude', fontsize=16)
        ax.set_xlabel('Solar constant (W m$^{-2}$)', fontsize=16)
        ax.plot( [const.S0, const.S0], [-10, 100], 'k--', label='present-day' )
        ax.legend(loc='upper left')
```



The upshot:

- For extremely large  $S_0$ , the only possible climate is a hot Earth with no ice.
- For extremely small  $S_0$ , the only possible climate is a cold Earth completely covered in ice.
- For a large range of  $S_0$  including the present-day value, more than one climate is possible!
- Once we get into a Snowball Earth state, getting out again is rather difficult!

In [ ]:

## APPLICATION PROGRAMMING INTERFACE

### 6.1 Subpackages

#### 6.1.1 climlab.domain package

climlab.domain.axis module

Axis

```
class climlab.domain.axis.Axis (axis_type='abstract',      num_points=10,      points=None,
                                bounds=None)
```

Bases: `object`

Creates a new climlab Axis object.

An *Axis* (page 55) is an object where information of a spacial dimension of a *Domain* (page 59) are specified.

These include the *type* of the axis, the *number of points*, location of *points* and *bounds* on the spatial dimension, magnitude of bounds differences *delta* as well as their *unit*.

The *axes* of a *Domain* (page 59) are stored in the dictionary `axes`, so they can be accessed through `dom.axes` if `dom` is an instance of *Domain* (page 59).

#### Initialization parameters

An instance of *Axis* is initialized with the following arguments (*for detailed information see Object attributes below*):

##### Parameters

- **axis\_type** (*str*) – information about the type of axis
- **num\_points** (*int*) – number of points on axis
- **points** (*array*) – array with specific points (optional)
- **bounds** (*array*) – array with specific bounds between points (optional)

**Raises** `ValueError` if `axis_type` is not one of the valid types or their euivalents (see below).

**Raises** `ValueError` if `points` are given and not array-like.

**Raises** `ValueError` if `bounds` are given and not array-like.

## Object attributes

Following object attributes are generated during initialization:

### Variables

- **axis\_type** (*str*) – Information about the type of axis. Valid axis types are:
  - 'lev'
  - 'lat'
  - 'lon'
  - 'depth'
  - 'abstract' (default)
- **num\_points** (*int*) – number of points on axis
- **units** (*str*) – Unit of the axis. During initialization the unit is chosen from the `defaultUnits` dictionary (see below).
- **points** (*array*) – array with all points of the axis (grid)
- **bounds** (*array*) – array with all bounds between points (staggered grid)
- **delta** (*array*) – array with spatial differences between bounds

### Axis Types

A couple of differing axis type strings are rendered to valid axis types. Alternate forms are listed here:

- 'lev'
  - 'p'
  - 'press'
  - 'pressure'
  - 'P'
  - 'Pressure'
  - 'Press'
- 'lat'
  - 'Latitude'
  - 'latitude'
- 'lon'
  - 'Longitude'
  - 'longitude'
- 'depth'
  - 'Depth'
  - 'waterDepth'
  - 'water\_depth'
  - 'slab'

The default units are:

```
defaultUnits = {'lev': 'mb',  
                'lat': 'degrees',  
                'lon': 'degrees',  
                'depth': 'meters',  
                'abstract': 'none'}
```

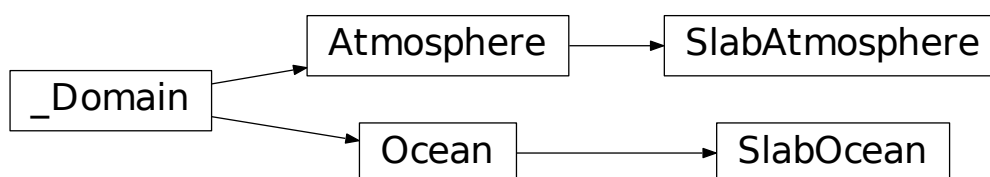
If bounds are not given during initialization, **default end points** are used:

```
defaultEndPoints = {'lev': (0., climlab.constants.ps),
                    'lat': (-90., 90.),
                    'lon': (0., 360.),
                    'depth': (0., 10.),
                    'abstract': (0, num_points)}
```

**Example** Creation of a standalone Axis:

[illegible]

## climlab.domain.domain module



```
class climlab.domain.domain.Atmosphere (**kwargs)
    Bases: climlab.domain.domain._Domain (page 59)
```

Class for the implementation of an Atmosphere Domain.

## Object attributes

Additional to the parent class `_Domain` (page 59) the following object attribute is modified during initialization:

Variables **domain\_type** (*str*) – is set to 'atm'

**Example** Setting up an Atmosphere Domain:

```
>>> import climlab
>>> atm_ax = climlab.domain.Axis(axis_type='pressure', num_points=10)
>>> atm_domain = climlab.domain.Atmosphere(axes=atm_ax)

>>> print atm_domain
climlab Domain object with domain_type=atm and shape=(10,)

>>> atm_domain.axes
{'lev': <climlab.domain.axis.Axis object at 0x7fe5b8ef8e10>}

>>> atm_domain.heat_capacity
array([ 1024489.79591837,  1024489.79591837,  1024489.79591837,
        1024489.79591837,  1024489.79591837,  1024489.79591837,
        1024489.79591837,  1024489.79591837,  1024489.79591837,
        1024489.79591837])
```

### **set\_heat\_capacity()**

Sets the heat capacity of the Atmosphere Domain.

Calls the utils heat capacity function *atmosphere()* (page 115) and gives the delta array of grid points of it's level axis `self.axes['lev'].delta` as input.

**Variables** *heat\_capacity* (page 115) (*array*) – the ocean domain's heat capacity over the 'lev' Axis.

**class** `climlab.domain.domain.Ocean` (*\*\*kwargs*)

Bases: `climlab.domain.domain._Domain` (page 59)

Class for the implementation of an Ocean Domain.

### **Object attributes**

Additional to the parent class `_Domain` (page 59) the following object attribute is modified during initialization:

**Variables** `domain_type` (*str*) – is set to 'ocean'

**Example** Setting up an Ocean Domain:

```
>>> import climlab
>>> ocean_ax = climlab.domain.Axis(axis_type='depth', num_points=5)
>>> ocean_domain = climlab.domain.Ocean(axes=ocean_ax)

>>> print ocean_domain
climlab Domain object with domain_type=ocean and shape=(5,)

>>> ocean_domain.axes
{'depth': <climlab.domain.axis.Axis object at 0x7fe5b8f102d0>}

>>> ocean_domain.heat_capacity
array([ 8362600.,  8362600.,  8362600.,  8362600.,  8362600.])
```

### **set\_heat\_capacity()**

Sets the heat capacity of the Ocean Domain.

Calls the utils heat capacity function *ocean()* (page 116) and gives the delta array of grid points of it's depth axis `self.axes['depth'].delta` as input.

### **Object attributes**

During method execution following object attribute is modified:

**Variables** *heat\_capacity* (page 115) (*array*) – the ocean domain's heat capacity over the 'depth' Axis.

**class** climlab.domain.domain.**SlabAtmosphere** (*axes=<climlab.domain.axis.Axis object>*,  
*\*\*kwargs*)  
 Bases: *climlab.domain.domain.Atmosphere* (page 57)

A class to create a SlabAtmosphere Domain by default.

Initializes the parent *Atmosphere* (page 57) class for with a simple axis for a Slab Atmosphere created by *make\_slabatm\_axis()* (page 60) which has just 1 cell in height by default.

**Example** Creating a SlabAtmosphere Domain:

```
>>> import climlab
>>> slab_atm_domain = climlab.domain.SlabAtmosphere()

>>> print slab_atm_domain
climlab Domain object with domain_type=atm and shape=(1,)

>>> slab_atm_domain.axes
{'lev': <climlab.domain.axis.Axis object at 0x7fe5c4281610>}

>>> slab_atm_domain.heat_capacity
array([ 10244897.95918367])
```

**class** climlab.domain.domain.**SlabOcean** (*axes=<climlab.domain.axis.Axis object>*,  
*\*\*kwargs*)  
 Bases: *climlab.domain.domain.Ocean* (page 58)

A class to create a SlabOcean Domain by default.

Initializes the parent *Ocean* (page 58) class for with a simple axis for a Slab Ocean created by *make\_slabocean\_axis()* (page 61) which has just 1 cell in depth by default.

**Example** Creating a SlabOcean Domain:

```
>>> import climlab
>>> slab_ocean_domain = climlab.domain.SlabOcean()

>>> print slab_ocean_domain
climlab Domain object with domain_type=ocean and shape=(1,)

>>> slab_ocean_domain.axes
{'depth': <climlab.domain.axis.Axis object at 0x7fe5c42814d0>}

>>> slab_ocean_domain.heat_capacity
array([ 41813000.])
```

**class** climlab.domain.domain.**\_Domain** (*axes=None, \*\*kwargs*)  
 Bases: *object*

Private parent class for *Domains*.

A *Domain* defines an area or spatial base for a climlab *Process* (page 83) object. It consists of axes which are *Axis* (page 55) objects that define the dimensions of the *Domain*.

In a *Domain* the heat capacity of grid points, bounds or cells/boxes is specified.

There are daughter classes *Atmosphere* (page 57) and *Ocean* (page 58) of the private *\_Domain* (page 59) class implemented which themselves have daughter classes *SlabAtmosphere* (page 58) and *SlabOcean* (page 59).

Several methods are implemented that create *Domains* with special specifications. These are

- *single\_column()* (page 61)
- *zonal\_mean\_column()* (page 62)
- *box\_model\_domain()* (page 60)

## Initialization parameters

An instance of `_Domain` is initialized with the following arguments:

**Parameters** `axes` (dict or [Axis](#) (page 55)) – Axis object or dictionary of Axis object where domain will be defined on.

## Object attributes

Following object attributes are generated during initialization:

### Variables

- `domain_type` (*str*) – Set to 'undefined'.
- `axes` (*dict*) – A dictionary of the domains axes. Created by `_make_axes_dict()` (page 60) called with input argument `axes`
- `numdims` (*int*) – Number of [Axis](#) (page 55) objects in `self.axes` dictionary.
- `ax_index` (*dict*) – A dictionary of domain axes and their corresponding index in an ordered list of the axes with:
  - 'lev' or 'depth' is last
  - 'lat' is second last
- `shape` (*tuple*) – Number of points of all domain axes. Order in tuple given by `self.ax_index`.
- `heat_capacity` (page 115) (*array*) – the domain's heat capacity over axis specified in function call of `set_heat_capacity()` (page 60)

`_make_axes_dict` (*axes*)

Makes an axes dictionary.

---

**Note:** In case the input is `None`, the dictionary `{ 'empty' : None }` is returned.

---

## Function-call argument

**Parameters** `axes` (dict or single instance of [Axis](#) (page 55) object or `None`) – axes input

**Raises** `ValueError` if input is not an instance of `Axis` class or a dictionary of `Axis` objects

**Returns** dictionary of input axes

**Return type** `dict`

`set_heat_capacity` ()

A dummy function to set the heat capacity of a domain.

*Should be overridden by daughter classes.*

`climlab.domain.domain.box_model_domain` (*num\_points=2, \*\*kwargs*)

Creates a box model domain (a single abstract axis).

**Parameters** `num_points` (*int*) – number of boxes [default: 2]

**Returns** Domain with single axis of type 'abstract' and `self.domain_type = 'box'`

**Return type** `_Domain` (page 59)

## Example

```
>>> from climlab import domain
>>> box = domain.box_model_domain(num_points=2)

>>> print box
climlab Domain object with domain_type=box and shape=(2,)
```



`climlab.domain.domain.make_slabatm_axis(num_points=1)`

Convenience method to create a simple axis for a slab atmosphere.

#### Function-call argument

**Parameters** `num_points` (*int*) – number of points for the slabatmosphere Axis [default: 1]

**Returns** an Axis with `axis_type='lev'` and `num_points=num_points`

**Return type** [Axis](#) (page 55)

#### Example

```
>>> import climlab
>>> slab_atm_axis = climlab.domain.make_slabatm_axis()

>>> print slab_atm_axis
Axis of type lev with 1 points.

>>> slab_atm_axis.axis_type
'lev'

>>> slab_atm_axis.bounds
array([ 0., 1000.])

>>> slab_atm_axis.units
'mb'
```

`climlab.domain.domain.make_slabocean_axis(num_points=1)`

Convenience method to create a simple axis for a slab ocean.

#### Function-call argument

**Parameters** `num_points` (*int*) – number of points for the slabocean Axis [default: 1]

**Returns** an Axis with `axis_type='depth'` and `num_points=num_points`

**Return type** [Axis](#) (page 55)

#### Example

```
>>> import climlab
>>> slab_ocean_axis = climlab.domain.make_slabocean_axis()

>>> print slab_ocean_axis
Axis of type depth with 1 points.

>>> slab_ocean_axis.axis_type
'depth'

>>> slab_ocean_axis.bounds
array([ 0., 10.])

>>> slab_ocean_axis.units
'meters'
```

`climlab.domain.domain.single_column(num_lev=30, water_depth=1.0, lev=None, **kwargs)`

Creates domains for a single column of atmosphere overlying a slab of water.

Can also pass a pressure array or pressure level axis object specified in `lev`.

If argument `lev` is not `None` then function tries to build a level axis and `num_lev` is ignored.

#### Function-call argument

##### Parameters

- **num\_lev** (*int*) – number of pressure levels (evenly spaced from surface to TOA) [default: 30]
- **water\_depth** (*float*) – depth of the ocean slab [default: 1.]
- **lev** (*Axis* (page 55) or pressure array) – specification for height axis (optional)

**Raises** `ValueError` if *lev* is given but neither *Axis* nor pressure array.

**Returns** a list of 2 Domain objects (slab ocean, atmosphere)

**Return type** list of *SlabOcean* (page 59), *SlabAtmosphere* (page 58)

**Example**

```
>>> from climlab import domain

>>> sfc, atm = domain.single_column(num_lev=2, water_depth=10.)

>>> print sfc
climlab Domain object with domain_type=ocean and shape=(1,)

>>> print atm
climlab Domain object with domain_type=atm and shape=(2,)
```

`climlab.domain.domain.zonal_mean_column(num_lat=90, num_lev=30, water_depth=10.0, lat=None, lev=None, **kwargs)`

Creates two Domains with one water cell, a latitude axis and a level/height axis.

- *SlabOcean*: one water cell and a latitude axis above (similar to *zonal\_mean\_surface* (page 62))
- *Atmosphere*: a latitude axis and a level/height axis (two dimensional)

**Function-call argument**

**Parameters**

- **num\_lat** (*int*) – number of latitude points on the axis [default: 90]
- **num\_lev** (*int*) – number of pressure levels (evenly spaced from surface to TOA) [default: 30]
- **water\_depth** (*float*) – depth of the water cell (slab ocean) [default: 10.]
- **lat** (*Axis* (page 55) or latitude array) – specification for latitude axis (optional)
- **lev** (*Axis* (page 55) or pressure array) – specification for height axis (optional)

**Raises** `ValueError` if *lat* is given but neither *Axis* nor latitude array.

**Raises** `ValueError` if *lev* is given but neither *Axis* nor pressure array.

**Returns** a list of 2 Domain objects (slab ocean, atmosphere)

**Return type** list of *SlabOcean* (page 59), *Atmosphere* (page 57)

**Example**

```
>>> from climlab import domain

>>> sfc, atm = domain.zonal_mean_column(num_lat=36, num_lev=10)

>>> print sfc
climlab Domain object with domain_type=ocean and shape=(36, 1)

>>> print atm
climlab Domain object with domain_type=atm and shape=(36, 10)
```

```
climlab.domain.domain.zonal_mean_surface(num_lat=90, water_depth=10.0, lat=None,
                                          **kwargs)
```

Creates a Domain with one water cell and a latitude axis above.

Domain has a single heat capacity according to the specified water depth.

### Function-call argument

#### Parameters

- **num\_lat** (*int*) – number of latitude points on the axis [default: 90]
- **water\_depth** (*float*) – depth of the water cell (slab ocean) [default: 10.]
- **lat** (*Axis* (page 55) or latitude array) – specification for latitude axis (optional)

**Raises** `ValueError` if *lat* is given but neither *Axis* nor latitude array.

**Returns** surface domain

**Return type** *SlabOcean* (page 59)

#### Example

```
>>> from climlab import domain
>>> sfc = domain.zonal_mean_surface(num_lat=36)

>>> print sfc
climlab Domain object with domain_type=ocean and shape=(36, 1)
```

## climlab.domain.field module



**class** `climlab.domain.field.Field`

Bases: `numpy.ndarray`

Custom class for climlab gridded quantities, called Field

This class behaves exactly like `numpy.ndarray` but every object has an attribute called `self.domain` which is the domain associated with that field (e.g. state variables).

### Initialization parameters

An instance of `Field` is initialized with the following arguments:

#### Parameters

- **input\_array** (*array*) – the array which the Field object should be initialized with
- **domain** (*\_Domain* (page 59)) – the domain associated with that field (e.g. state variables)

### Object attributes

Following object attribute is generated during initialization:

**Variables** `domain` (page 57) (*\_Domain* (page 59)) – the domain associated with that field (e.g. state variables)

### Example

```
>>> import climlab
>>> import numpy as np
>>> from climlab import domain
>>> from climlab.domain import field

>>> # distribution of state
>>> distr = np.linspace(0., 10., 30)
>>> # domain creation
>>> sfc, atm = domain.single_column()
>>> # build state of type Field
>>> s = field.Field(distr, domain=atm)

>>> print s
[ 0.          0.34482759  0.68965517  1.03448276  1.37931034
 1.72413793  2.06896552  2.4137931  2.75862069  3.10344828
 3.44827586  3.79310345  4.13793103  4.48275862  4.82758621
 5.17241379  5.51724138  5.86206897  6.20689655  6.55172414
 6.89655172  7.24137931  7.5862069  7.93103448  8.27586207
 8.62068966  8.96551724  9.31034483  9.65517241 10.          ]

>>> print s.domain
climlab Domain object with domain_type=atm and shape=(30,)

>>> # can slice this and it preserves the domain
>>> # a more full-featured implementation would have intelligent
>>> # slicing like in iris
>>> s.shape == s.domain.shape
True
>>> s[:1].shape == s[:1].domain.shape
False

>>> # But some things work very well. E.g. new field creation:
>>> s2 = np.zeros_like(s)

>>> print s2
[ 0.  0.  0.  0.  0.  0.  0.  0.  0.  0.  0.  0.  0.  0.  0.  0.  0.  0.
 0.  0.  0.  0.  0.  0.  0.  0.  0.  0.  0.  0.]

>>> print s2.domain
climlab Domain object with domain_type=atm and shape=(30,)
```

`climlab.domain.field.global_mean(field)`

Calculates the latitude weighted global mean of a field with latitude dependence.

**Parameters** `field` (`Field` (page 63)) – input field

**Raises** `ValueError` if input field has no latitude axis

**Returns** latitude weighted global mean of the field

**Return type** `float`

**Example** initial global mean temperature of EBM model:

```
>>> import climlab
>>> from climlab.domain.field import global_mean

>>> model = climlab.EBM()

>>> global_mean(model.Ts)
Field(11.997968598413685)
```

## climlab.domain.initial module

Convenience routines for setting up initial conditions.

`climlab.domain.initial.column_state` (*num\_lev=30, num\_lat=1, lev=None, lat=None, water\_depth=1.0*)

Sets up a state variable dictionary consisting of temperatures for atmospheric column (*T<sub>atm</sub>*) and surface mixed layer (*T<sub>s</sub>*).

Surface temperature is always 288 K. Atmospheric temperature is initialized between 278 K at lowest altitude and 200 at top of atmosphere according to the number of levels given.

### Function-call arguments

#### Parameters

- **num\_lev** (*int*) – number of pressure levels (evenly spaced from surface to top of atmosphere) [default: 30]
- **num\_lat** (*int*) – number of latitude points on the axis [default: 1]
- **lev** (*Axis* (page 55) or pressure array) – specification for height axis (optional)
- **lat** (*array*) – size of array determines dimension of latitude
- **water\_depth** (*float*) – irrelevant

**Returns** dictionary with two temperature *Field* (page 63) for atmospheric column *T<sub>atm</sub>* and surface mixed layer *T<sub>s</sub>*

**Return type** `dict`

#### Example

```
>>> from climlab.domain import initial
>>> T_dict = initial.column_state()

>>> print T_dict
{'Tatm': Field([ 200.          , 202.68965517, 205.37931034, 208.06896552,
                210.75862069, 213.44827586, 216.13793103, 218.82758621,
                221.51724138, 224.20689655, 226.89655172, 229.5862069 ,
                232.27586207, 234.96551724, 237.65517241, 240.34482759,
                243.03448276, 245.72413793, 248.4137931 , 251.10344828,
                253.79310345, 256.48275862, 259.17241379, 261.86206897,
                264.55172414, 267.24137931, 269.93103448, 272.62068966,
                275.31034483, 278.          ]), 'Ts': Field([ 288.] )}
```

`climlab.domain.initial.surface_state` (*num\_lat=90, water\_depth=10.0, T0=12.0, T2=-40.0*)

Sets up a state variable dictionary for a zonal-mean surface model (e.g. basic EBM).

Returns a single state variable *T<sub>s</sub>*, the temperature of the surface mixed layer, initialized by a basic temperature and the second Legendre polynomial.

### Function-call arguments

#### Parameters

- **num\_lat** (*int*) – number of latitude points on the axis [default: 90]
- **water\_depth** (*float*) – irrelevant
- **T0** (*float*) – base value for initial temperature
  - unit °C
  - default value: 12
- **T2** (*float*) – factor for 2nd Legendre polynomial *P<sub>2</sub>* (page 117) to calculate initial temperature

- unit: dimensionless
- default value: 40

**Returns** dictionary with temperature *Field* (page 63) for surface mixed layer Ts

**Return type** dict

**Example**

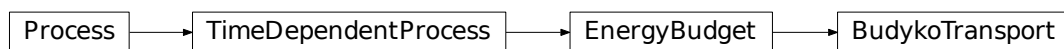
```
>>> from climlab.domain import initial
>>> import numpy as np

>>> T_dict = initial.surface_state(num_lat=36)

>>> print np.squeeze(T_dict['Ts'])
[-27.88584094 -26.97777479 -25.18923361 -22.57456133 -19.21320344
 -15.20729309 -10.67854785 -5.76457135 -0.61467228  4.61467228
  9.76457135  14.67854785  19.20729309  23.21320344  26.57456133
 29.18923361  30.97777479  31.88584094  31.88584094  30.97777479
 29.18923361  26.57456133  23.21320344  19.20729309  14.67854785
  9.76457135  4.61467228 -0.61467228 -5.76457135 -10.67854785
 -15.20729309 -19.21320344 -22.57456133 -25.18923361 -26.97777479
 -27.88584094]
```

## 6.1.2 climlab.dynamics package

**climlab.dynamics.budyko\_transport module**



**class** climlab.dynamics.budyko\_transport.**BudykoTransport** (*b=3.81, \*\*kwargs*)

Bases: *climlab.process.energy\_budget.EnergyBudget* (page 81)

calculates the 1 dimensional heat transport as the difference between the local temperature and the global mean temperature.

**Parameters** *b* (*float*) – budyko transport parameter

- unit: W / (m<sup>2</sup> °C)
- default value: 3.81

As BudykoTransport is a *Process* (page 83) it needs a state do be defined on. See example for details.

**Computation Details:**

In a global Energy Balance Model

$$C \frac{dT}{dt} = R \downarrow - R \uparrow - H$$

with model state *T*, the energy transport term *H* can be described as

$$H = b[T - \bar{T}]$$

where *T* is a vector of the model temperature and  $\bar{T}$  describes the mean value of *T*.

For further information see [*Budyko1969*] (page 129).

**Example** Budyko Transport as a standalone process:

```

import climlab
from climlab.dynamics.budyko_transport import BudykoTransport
from climlab import domain
from climlab.domain import field
from climlab.utils.legendre import P2
import numpy as np
import matplotlib.pyplot as plt

# create domain
sfc = domain.zonal_mean_surface(num_lat = 36)

lat = sfc.lat.points
lat_rad = np.deg2rad(lat)

# define initial temperature distribution
T0 = 15.
T2 = -20.
Ts = field.Field(T0 + T2 * P2(np.sin(lat_rad)), domain=sfc)

# create BudykoTransport process
budyko_transp = BudykoTransport(state=Ts)

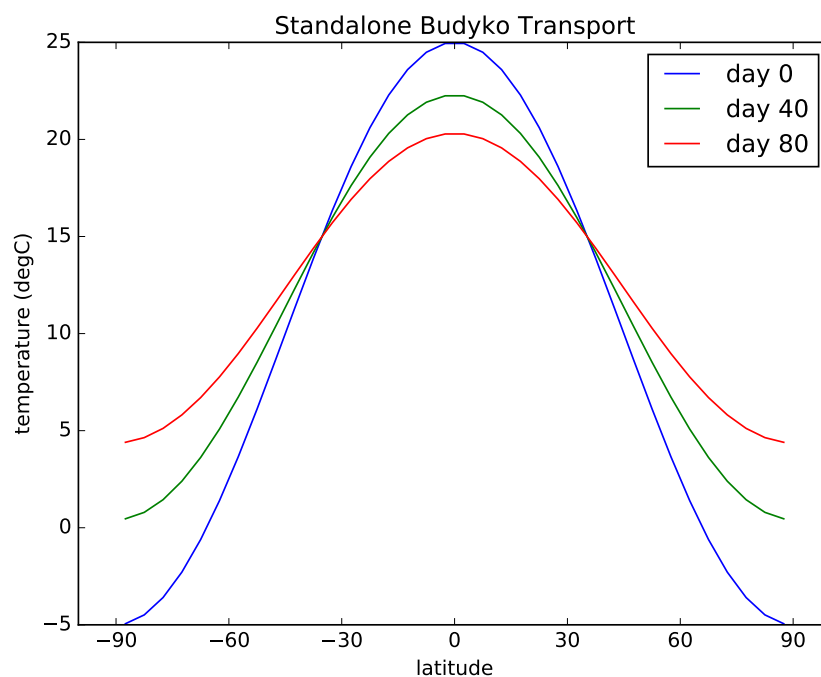
### Integrate & Plot ###

fig = plt.figure()
ax = fig.add_subplot(111)

for i in np.arange(0,3,1):
    ax.plot(lat, budyko_transp.default, label='day %s' % (i*40))
    budyko_transp.integrate_days(40.)

ax.set_title('Standalone Budyko Transport')
ax.set_xlabel('latitude')
ax.set_xticks([-90,-60,-30,0,30,60,90])
ax.set_ylabel('temperature (degC)')
ax.legend(loc='best')
plt.show()

```



**b**

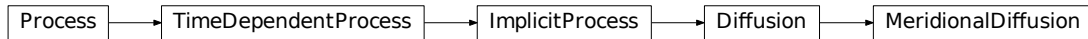
the budyko transport parameter in unit  $\frac{W}{m^2K}$

**Getter** returns the budyko transport parameter

**Setter** sets the budyko transport parameter

**Type** float

## climlab.dynamics.diffusion module



**class** `climlab.dynamics.diffusion.Diffusion` ( $K=None$ ,  $diffusion\_axis=None$ ,  
 $use\_banded\_solver=False$ ,  $**kwargs$ )  
 Bases: `climlab.process.implicit.ImplicitProcess` (page 83)

A parent class for one dimensional implicit diffusion modules.

Solves the one dimensional heat equation

$$\frac{dT}{dt} = \frac{d}{dy} \left[ K \cdot \frac{dT}{dy} \right]$$

### Initialization parameters

#### Parameters

- **K** (*float*) – the diffusivity parameter in units of  $\frac{[length]^2}{time}$  where length is the unit of the spatial axis on which the diffusion is occurring.
- **diffusion\_axis** (*str*) – dictionary key for axis on which the diffusion is occurring in process's domain axes dictionary
- **use\_banded\_solver** (*bool*) – input flag, whether to use `scipy.linalg.solve_banded()` instead of `numpy.linalg.solve()`

---

**Note:** The banded solver `scipy.linalg.solve_banded()` is faster than `numpy.linalg.solve()` but only works for one dimensional diffusion.

---

### Object attributes

Additional to the parent class `ImplicitProcess` (page 83) following object attributes are generated or modified during initialization:

#### Variables

- **param** (*dict*) – parameter dictionary is extended by diffusivity parameter K (unit:  $\frac{[length]^2}{time}$ )
- **use\_banded\_solver** (*bool*) – input flag specifying numerical solving method (given during initialization)
- **diffusion\_axis** (*str*) – dictionary key for axis where diffusion is occurring: specified during initialization or output of method `_guess_diffusion_axis()` (page 71)
- **K\_dimensionless** (*array*) – diffusion parameter K multiplied by the timestep and divided by mean of diffusion axis delta in the power of two. Array has the size of diffusion axis bounds.  $K_{dimensionless}[i] = K \frac{\Delta t}{(\Delta bounds)^2}$



- **diffTriDiag** (*array*) – tridiagonal diffusion matrix made by `_make_diffusion_matrix()` (page 72) with input `self.K_dimensionless`

**Example** Here is an example showing implementation of a vertical diffusion. It shows that a subprocess can work on just a subset of the parent process state variables.

```
import climlab
from climlab.dynamics.diffusion import Diffusion
import matplotlib.pyplot as plt

c = climlab.GreyRadiationModel()
K = 0.5
d = Diffusion(K=K, state = {'Tatm':c.state['Tatm']}, **c.param)

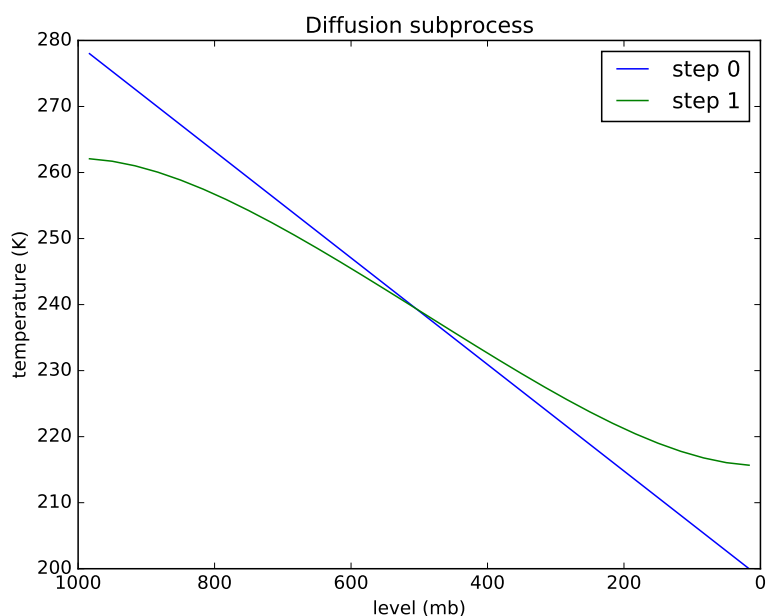
c.add_subprocess('diffusion',d)

### Integrate & Plot ###

fig = plt.figure()
ax = fig.add_subplot(111)

ax.plot(c.lev, c.state['Tatm'], label='step 0')
c.step_forward()
ax.plot(c.lev, c.state['Tatm'], label='step 1')

ax.invert_xaxis()
ax.set_title('Diffusion subprocess')
ax.set_xlabel('level (mb)')
#ax.set_xticks([])
ax.set_ylabel('temperature (K)')
ax.legend(loc='best')
plt.show()
```



#### `_implicit_solver()`

Invertes and solves the matrix problem for diffusion matrix and temperature T.

The method is called by the `_compute()` (page 83) function of the *ImplicitProcess* (page 83) class and solves the matrix problem

$$A \cdot T_{\text{new}} = T_{\text{old}}$$

for diffusion matrix  $A$  and corresponding temperatures.  $T_{\text{old}}$  is in this case the current state variable which already has been adjusted by the explicit processes.  $T_{\text{new}}$  is the new state of the variable. To derive the temperature tendency of the diffusion process the adjustment has to be calculated and multiplied with the timestep which is done by the `_compute()` (page 83) function of the `ImplicitProcess` (page 83) class.

This method calculates the matrix inversion for every state variable and calling either `solve_implicit_banded()` or `numpy.linalg.solve()` dependent on the flag `self.use_banded_solver`.

#### Variables

- **state** (*dict*) – method uses current state variables but does not modify them
- **use\_banded\_solver** (*bool*) – input flag whether to use `_solve_implicit_banded()` (page 73) or `numpy.linalg.solve()` to do the matrix inversion
- **diffTriDiag** (*array*) – the diffusion matrix which is given with the current state variable to the method solving the matrix problem

**class** `climlab.dynamics.diffusion.MeridionalDiffusion` ( $K=None$ ,  $**kwargs$ )

Bases: `climlab.dynamics.diffusion.Diffusion` (page 68)

A parent class for Meridional diffusion processes.

Calculates the energy transport in a diffusion like process along the temperature gradient:

$$H(\varphi) = \frac{D}{\cos \varphi} \frac{\partial}{\partial \varphi} \left( \cos \varphi \frac{\partial T(\varphi)}{\partial \varphi} \right)$$

for an Energy Balance Model whose Energy Budget can be noted as:

$$C(\varphi) \frac{dT(\varphi)}{dt} = R \downarrow(\varphi) - R \uparrow(\varphi) + H(\varphi)$$

#### Initialization parameters

An instance of `MeridionalDiffusion` is initialized with the following arguments:

**Parameters**  $K$  (*float*) – diffusion parameter in units of  $1/s$

#### Object attributes

Additional to the parent class `Diffusion` (page 68) which is initialized with `diffusion_axis='lat'`, following object attributes are modified during initialization:

#### Variables

- **K\_dimensionless** (*array*) – As  $K_{\text{dimensionless}}$  has been computed like  $K_{\text{dimensionless}} = K \frac{\Delta t}{(\Delta \text{bounds})^2}$  with  $K$  in units  $1/s$ , the  $\Delta(\text{bounds})$  have to be converted from deg to rad to make the array actually dimensionless. This is done during initialization.
- **diffTriDiag** (*array*) – the diffusion matrix is recomputed with appropriate weights for the meridional case by `_make_meridional_diffusion_matrix()` (page 72)

**Example** Meridional Diffusion of temperature as a stand-alone process:

```
import numpy as np
import climlab
from climlab.dynamics.diffusion import MeridionalDiffusion
from climlab.utils import legendre
```

```

sfc = climlab.domain.zonal_mean_surface(num_lat=90, water_depth=10.)
lat = sfc.lat.points
initial = 12. - 40. * legendre.P2(np.sin(np.deg2rad(lat)))

# make a copy of initial so that it remains unmodified
Ts = climlab.Field(np.array(initial), domain=sfc)

# thermal diffusivity in W/m**2/degC
D = 0.55

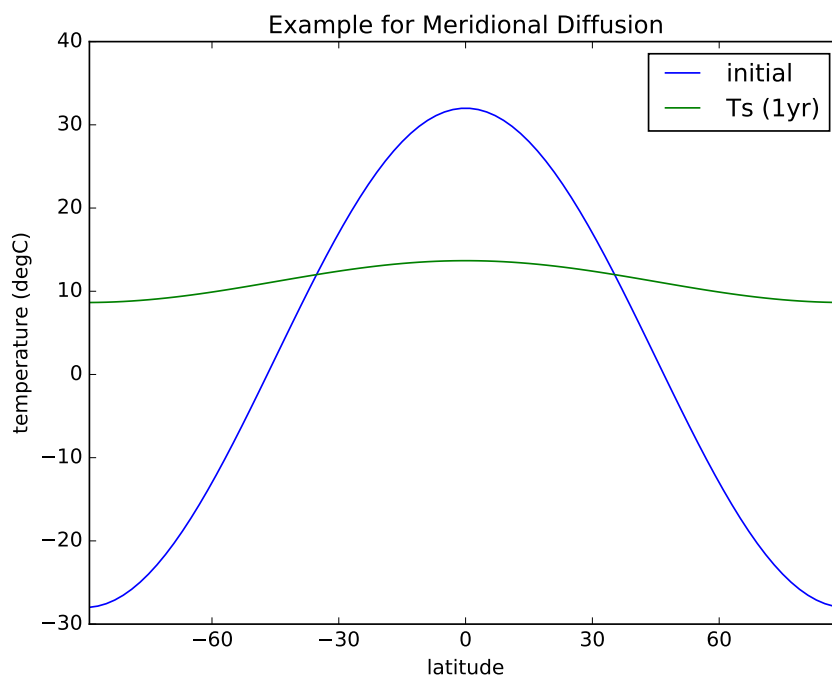
# meridional diffusivity in 1/s
K = D / sfc.heat_capacity
d = MeridionalDiffusion(state=Ts, K=K)

d.integrate_years(1.)

import matplotlib.pyplot as plt

fig = plt.figure()
ax = fig.add_subplot(111)
ax.set_title('Example for Meridional Diffusion')
ax.set_xlabel('latitude')
ax.set_xticks([-90, -60, -30, 0, 30, 60, 90])
ax.set_ylabel('temperature (degC)')
ax.plot(lat, initial, label='initial')
ax.plot(lat, Ts, label='Ts (1yr)')
ax.legend(loc='best')
plt.show()

```



`climlab.dynamics.diffusion._guess_diffusion_axis` (*process\_or\_domain*)

Scans given process, domain or dictionary of domains for a diffusion axis and returns appropriate name.

In case only one axis with length > 1 in the process or set of domains exists, the name of that axis is returned. Otherwise an error is raised.

**Parameters** `process_or_domain` (*Process* (page 83), *\_Domain* (page 59) or *dict* of domains) – input from where diffusion axis should be guessed

**Raises** `ValueError` if more than one diffusion axis is possible.

**Returns** name of the diffusion axis

**Return type** `str`

`climlab.dynamics.diffusion._make_diffusion_matrix` (*K*, *weight1=None*, *weight2=None*)

Builds the general diffusion matrix with dimension  $n \times n$ .

---

**Note:**  $n$  = number of points of diffusion axis  $n + 1$  = number of bounds of diffusion axis

---

### Function-all argument

#### Parameters

- **K** (*array*) – dimensionless diffusivities at cell boundaries (*size: 1xn+1*)
- **weight1** (*array*) – *weight\_1* (*size: 1xn+1*)
- **weight2** (*array*) – *weight\_2* (*size: 1xn*)

**Returns** completely listed tridiagonal diffusion matrix (*size: nxn*)

**Return type** `array`

---

**Note:** The elements of array *K* are acutally dimensionless:

$$K[i] = K_{\text{physical}} \frac{\Delta t}{(\Delta y)^2}$$

where  $K_{\text{physical}}$  is in unit  $\frac{\text{length}^2}{\text{time}}$

---

The diffusion matrix is build like the following

$$\text{diffTriDiag} = \begin{bmatrix} 1 + \frac{w_{1,1}K_1}{w_{2,0}} & -\frac{w_{1,1}K_1}{w_{2,0}} & 0 & & & \dots \\ -\frac{w_{1,1}K_1}{w_{2,1}} & 1 + \frac{w_{1,1}K_1 + w_{1,2}K_2}{w_{2,1}} & -\frac{w_{1,2}K_2}{w_{2,1}} & 0 & & \dots \\ 0 & -\frac{w_{1,2}K_2}{w_{2,2}} & 1 + \frac{w_{1,2}K_2 + w_{1,3}K_3}{w_{2,2}} & -\frac{w_{1,3}K_3}{w_{2,2}} & & \dots \\ & \ddots & \ddots & \ddots & \ddots & \ddots \\ 0 & 0 & \dots & -\frac{w_{1,n-2}K_{n-2}}{w_{2,n-2}} & 1 + \frac{w_{1,n-2}K_{n-2} + w_{1,n-1}K_{n-1}}{w_{2,n-2}} & \dots \\ 0 & 0 & \dots & 0 & -\frac{w_{1,n-1}K_{n-1}}{w_{2,n-1}} & \dots \end{bmatrix}$$

where

$$\begin{aligned} K &= [K_0, K_1, K_2, \dots, K_{n-1}, K_n] \\ w_1 &= [w_{1,0}, w_{1,1}, w_{1,2}, \dots, w_{1,n-1}, w_{1,n}] \\ w_2 &= [w_{2,0}, w_{2,1}, w_{2,2}, \dots, w_{2,n-1}] \end{aligned}$$

`climlab.dynamics.diffusion._make_meridional_diffusion_matrix` (*K*, *lataxis*)

Calls `_make_diffusion_matrix()` (page 72) with appropriate weights for the meridional diffusion case.

#### Parameters

- **K** (*array*) – dimensionless diffusivities at cell boundaries of diffusion axis *lataxis*
- **lataxis** (*axis* (page 55)) – latitude axis where diffusion is occuring

Weights are computed as the following:

$$\begin{aligned} w_1 &= \cos(\text{bounds}) \\ &= [\cos(b_0), \cos(b_1), \cos(b_2), \dots, \cos(b_{n-1}), \cos(b_n)] \\ w_2 &= \cos(\text{points}) \\ &= [\cos(p_0), \cos(p_1), \cos(p_2), \dots, \cos(p_{n-1})] \end{aligned}$$

when bounds and points from `lataxis` are written as

$$\begin{aligned}\text{bounds} &= [b_0, b_1, b_2, \dots, b_{n-1}, b_n] \\ \text{points} &= [p_0, p_1, p_2, \dots, p_{n-1}]\end{aligned}$$

Giving this input to `_make_diffusion_matrix()` (page 72) results in a matrix like:

$$\text{diffTriDiag} = \begin{bmatrix} 1 + \frac{\cos(b_1)K_1}{\cos(p_0)} & -\frac{\cos(b_1)K_1}{\cos(p_0)} & 0 & & & \dots \\ -\frac{\cos(b_1)K_1}{\cos(p_1)} & 1 + \frac{\cos(b_1)K_1 + \cos(b_2)K_2}{\cos(p_1)} & -\frac{\cos(b_2)K_2}{\cos(p_1)} & & & \dots \\ 0 & -\frac{\cos(b_2)K_2}{\cos(p_2)} & 1 + \frac{\cos(b_2)K_2 + \cos(b_3)K_3}{\cos(p_2)} & -\frac{\cos(b_3)K_3}{\cos(p_2)} & & \dots \\ & \ddots & \ddots & \ddots & \ddots & \ddots \\ 0 & 0 & \dots & -\frac{\cos(b_{n-2})K_{n-2}}{\cos(p_{n-2})} & 1 + \frac{\cos(b_{n-2})K_{n-2} + \cos(b_{n-1})K_{n-1}}{\cos(p_{n-1})} & -\frac{\cos(b_{n-1})K_{n-1}}{\cos(p_n)} \\ 0 & 0 & \dots & 0 & -\frac{\cos(b_{n-1})K_{n-1}}{\cos(p_n)} & 1 + \frac{\cos(b_{n-1})K_{n-1}}{\cos(p_n)} \end{bmatrix}$$

`climlab.dynamics.diffusion._solve_implicit_banded(current, banded_matrix)`

Uses a banded solver for matrix inversion of a tridiagonal matrix.

Converts the complete listed tridiagonal matrix ( $nxn$ ) into a three row matrix ( $3xn$ ) and calls `scipy.linalg.solve_banded()`.

#### Parameters

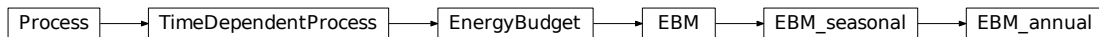
- **current** (*array*) – the current state of the variable for which matrix inversion should be computed
- **banded\_matrix** (*array*) – complete diffusion matrix (*dimension: nxn*)

**Returns** output of `scipy.linalg.solve_banded()`

**Return type** *array*

## 6.1.3 climlab.model package

### climlab.model.ebm module



**class** `climlab.model.ebm.EBM`(*num\_lat=90, S0=1365.2, A=210.0, B=2.0, D=0.555, water\_depth=10.0, Tf=-10.0, a0=0.3, a2=0.078, ai=0.62, timestep=350632.51200000005, T0=12.0, T2=-40.0, \*\*kwargs*)

Bases: `climlab.process.energy_budget.EnergyBudget` (page 81)

A parent class for all Energy-Balance-Model classes.

This class sets up a typical EnergyBalance Model with following subprocesses:

- Outgoing Longwave Radiation (OLR) parameterization through [AplusBT](#) (page 93)
- solar insolation parameterization through [P2Insolation](#) (page 103)
- albedo parameterization in dependence of temperature through [StepFunctionAlbedo](#) (page 113)
- energy diffusion through [MeridionalDiffusion](#) (page 70)

#### Initialization parameters

An instance of EBM is initialized with the following arguments (*for detailed information see Object attributes below*):

#### Parameters

- **num\_lat** (*int*) – number of equally spaced points for the latitude grid. Used for domain initialization of *zonal\_mean\_surface* (page 62)
  - default value: 90
- **S0** (*float*) – solar constant
  - unit:  $\frac{\text{W}}{\text{m}^2}$
  - default value: 1365.2
- **A** (*float*) – parameter for linear OLR parameterization *AplusBT* (page 93)
  - unit:  $\frac{\text{W}}{\text{m}^2}$
  - default value: 210.0
- **B** (*float*) – parameter for linear OLR parameterization *AplusBT* (page 93)
  - unit:  $\frac{\text{W}}{\text{m}^2 \text{ } ^\circ\text{C}}$
  - default value: 2.0
- **D** (*float*) – diffusion parameter for Meridional Energy Diffusion *MeridionalDiffusion* (page 70)
  - unit:  $\frac{\text{W}}{\text{m}^2 \text{ } ^\circ\text{C}}$
  - default value: 0.555
- **water\_depth** (*float*) – depth of *zonal\_mean\_surface* (page 62) domain, which the heat capacity is dependent on
  - unit: meters
  - default value: 10.0
- **Tf** (*float*) – freezing temperature
  - unit:  $^\circ\text{C}$
  - default value: -10.0
- **a0** (*float*) – base value for planetary albedo parameterization *StepFunctionAlbedo* (page 113)
  - unit: dimensionless
  - default value: 0.3
- **a2** (*float*) – parabolic value for planetary albedo parameterization *StepFunctionAlbedo* (page 113)
  - unit: dimensionless
  - default value: 0.078
- **ai** (*float*) – value for ice albedo parameterization in *StepFunctionAlbedo* (page 113)
  - unit: dimensionless
  - default value: 0.62
- **timestep** (*float*) – specifies the EBM's timestep
  - unit: seconds
  - default value:  $(365.2422 * 24 * 60 * 60) / 90$   
-> (90 timesteps per year)
- **T0** (*float*) – base value for initial temperature
  - unit  $^\circ\text{C}$

- default value: 12
- **T2** (*float*) – factor for 2nd Legendre polynomial *P2* (page 117) to calculate initial temperature
  - unit: dimensionless
  - default value: 40

### Object attributes

Additional to the parent class *EnergyBudget* (page 81) following object attributes are generated and updated during initialization:

#### Variables

- **param** (*dict*) – The parameter dictionary is updated with a couple of the initialization input arguments, namely 'S0', 'A', 'B', 'D', 'Tf', 'water\_depth', 'a0', 'a2' and 'ai'.
- **domains** (*dict*) – If the object's domains and the state dictionaries are empty during initialization a domain sfc is created through *zonal\_mean\_surface()* (page 62). In the meantime the object's domains and state dictionaries are updated.
- **subprocess** (*dict*) – Several subprocesses are created (see above) through calling *add\_subprocess()* (page 84) and therefore the subprocess dictionary is updated.
- **topdown** (*bool*) – is set to False to call subprocess compute methods first. See also *TimeDependentProcess* (page 89).
- **diagnostics** (*dict*) – is initialized with keys: 'OLR', 'ASR', 'net\_radiation', 'albedo' and 'icelat' through *init\_diagnostic()* (page 85).

**Example** Creation and integration of the preconfigured Energy Balance Model:

```
>>> import climlab
>>> model = climlab.EBM()

>>> model.integrate_years(2.)
Integrating for 180 steps, 730.4844 days, or 2.0 years.
Total elapsed time is 2.0 years.
```

For more information how to use the EBM class, see the *Tutorials* (page 17) chapter.

### **diffusive\_heat\_transport()**

Compute instantaneous diffusive heat transport in unit PW on the staggered grid (bounds) through calculating:

$$H(\varphi) = -2\pi R^2 \cos(\varphi) D \frac{dT}{d\varphi} \approx -2\pi R^2 \cos(\varphi) D \frac{\Delta T}{\Delta \varphi}$$

**Return type** array of size np.size(self.lat\_bounds)

### **global\_mean\_temperature()**

Convenience method to compute global mean surface temperature.

Calls *global\_mean()* (page 64) method which for the object attribute Ts which calculates the latitude weighted global mean of a field.

**Example** Calculating the global mean temperature of initial EBM temperature:

```
>>> import climlab
>>> model = climlab.EBM(T0=14., T2=-25)

>>> model.global_mean_temperature()
Field(13.99873037400856)
```

**heat\_transport()**

Returns instantaneous heat transport in unit PW on the staggered grid (bounds) through `callinaag.diffusive_heat_transport()` (page 75).

**Example**

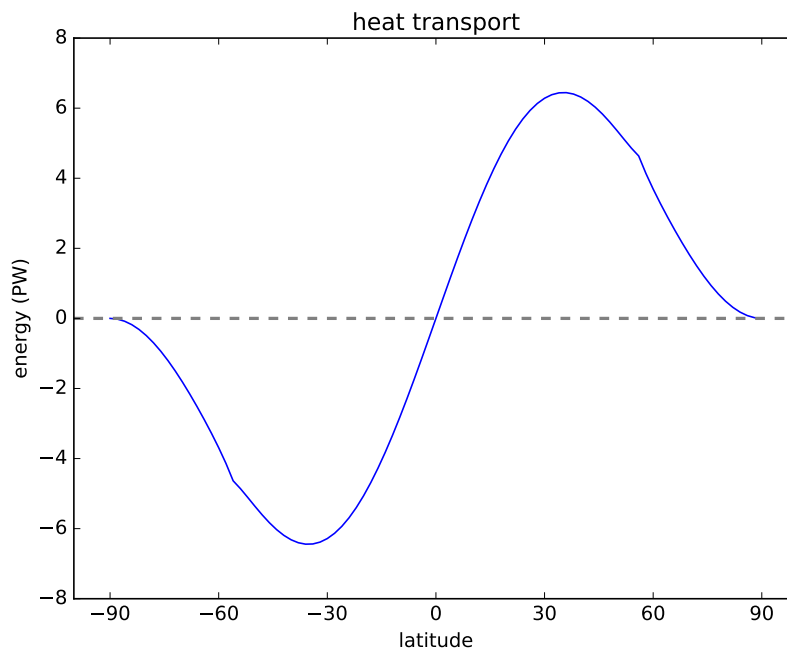
```
import climlab
import matplotlib.pyplot as plt

# creating & integrating model
model = climlab.EBM()
model.step_forward()

# plot
fig = plt.figure()
ax = fig.add_subplot(111)

bounds = model.domains['Ts'].axes['lat'].bounds
ax.plot(bounds, model.heat_transport())

ax.set_title('heat transport')
ax.set_xlabel('latitude')
ax.set_xticks([-90, -60, -30, 0, 30, 60, 90])
ax.set_ylabel('energy (PW)')
plt.axhline(linewidth=2, color='grey', linestyle='dashed')
plt.show()
```

**heat\_transport\_convergence()**

Returns instantaneous convergence of heat transport.

$$h(\varphi) = -\frac{1}{2\pi R^2 \cos(\varphi)} \frac{dH}{d\varphi} \approx -\frac{1}{2\pi R^2 \cos(\varphi)} \frac{\Delta H}{\Delta \varphi}$$

$h$  is the *dynamical heating rate* in unit  $\text{W}/\text{m}^2$  which is the convergence of energy transport into each latitude band, namely the difference between what's coming in and what's going out.

**Example**



```

import climlab
import matplotlib.pyplot as plt

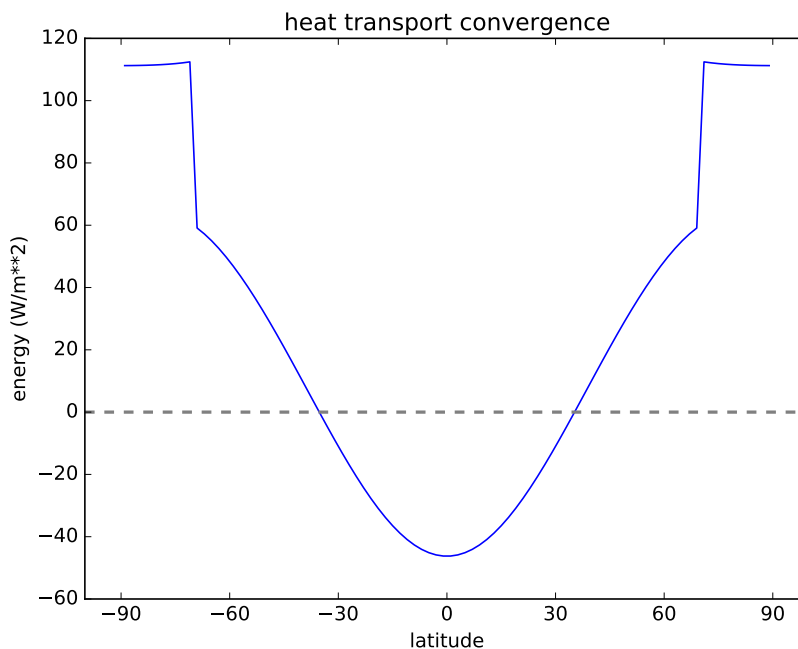
# creating & integrating model
model = climlab.EBM()
model.integrate_converge()

# plot
fig = plt.figure()
ax = fig.add_subplot(111)

ax.plot(model.lat, model.heat_transport_convergence())

ax.set_title('heat transport convergence')
ax.set_xlabel('latitude')
ax.set_xticks([-90, -60, -30, 0, 30, 60, 90])
ax.set_ylabel('energy (W/m**2)')
plt.axhline(linewidth=2, color='grey', linestyle='dashed')
plt.show()

```



#### **inferred\_heat\_transport()**

Calculates the inferred heat transport by integrating the TOA energy imbalance from pole to pole.

The method is calculating

$$H(\varphi) = 2\pi R^2 \int_{-\pi/2}^{\varphi} \cos\phi R_{TOA} d\phi$$

where  $R_{TOA}$  is the net radiation at top of atmosphere.

**Returns** total heat transport on the latitude grid in unit PW

**Return type** array of size `np.size(self.lat_lat)`

**Example**

```

import climlab
import matplotlib.pyplot as plt

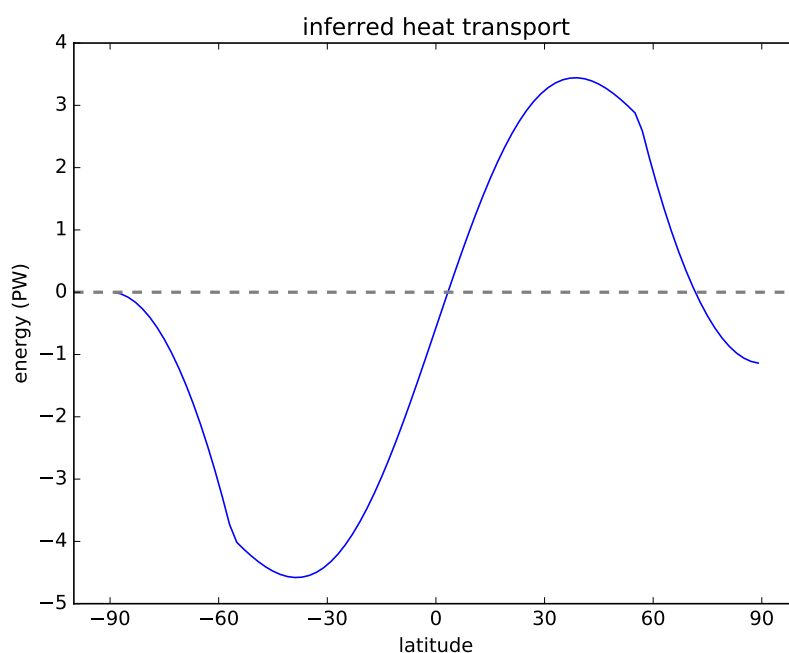
```

```
# creating & integrating model
model = climlab.EBM()
model.step_forward()

# plot
fig = plt.figure()
ax = fig.add_subplot(111)

ax.plot(model.lat, model.inferred_heat_transport())

ax.set_title('inferred heat transport')
ax.set_xlabel('latitude')
ax.set_xticks([-90, -60, -30, 0, 30, 60, 90])
ax.set_ylabel('energy (PW)')
plt.axhline(linewidth=2, color='grey', linestyle='dashed')
plt.show()
```



**class** climlab.model.ebm.**EBM\_annual** (\*\*kwargs)

Bases: [climlab.model.ebm.EBM\\_seasonal](#) (page 79)

A class that implements Energy Balance Models with annual mean insolation.

The annual solar distribution is calculated through averaging the [DailyInsolation](#) (page 101) over time which has been used in the parent class [EBM\\_seasonal](#) (page 79). That is done by the subprocess [AnnualMeanInsolation](#) (page 99) which is more realistic than the [P2Insolation](#) (page 103) module used in the classical [EBM](#) (page 73) class.

According to the parent class [EBM\\_seasonal](#) (page 79) the model will not have an ice-albedo feedback, if albedo ice parameter 'ai' is not given. For details see there.

#### Object attributes

Following object attributes are updated during initialization:

**Variables** `subprocess` (*dict*) – subprocess 'insolation' is overwritten by [AnnualMeanInsolation](#) (page 99)

**Example** The [EBM\\_annual](#) (page 78) class uses a different insolation subprocess than the [EBM](#) (page 73) class:

```
>>> import climlab
>>> model_annual = climlab.EBM_annual()

>>> print model_annual
```

```
climlab Process of type <class 'climlab.model.ebm.EBM_annual'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM_annual'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.AnnualMeanInsolation'>
```

**class** climlab.model.ebm.**EBM\_seasonal** (*a0=0.33, a2=0.25, ai=None, \*\*kwargs*)

Bases: [climlab.model.ebm.EBM](#) (page 73)

A class that implements Energy Balance Models with realistic daily insolation.

This class is inherited from the general [EBM](#) (page 73) class and uses the insolation subprocess [DailyInsolation](#) (page 101) instead of [P2Insolation](#) (page 103) to compute a realistic distribution of solar radiation on a daily basis.

If argument for ice albedo 'ai' is not given, the model will not have an albedo feedback.

An instance of `EBM_seasonal` is initialized with the following arguments:

#### Parameters

- **a0** (*float*) – base value for planetary albedo parameterization [StepFunctionAlbedo](#) (page 113)  
– default value: 0.33
- **a2** (*float*) – parabolic value for planetary albedo parameterization [StepFunctionAlbedo](#) (page 113)  
– default value: 0.25
- **ai** (*float*) – value for ice albedo parameterization in [StepFunctionAlbedo](#) (page 113)  
– default value: None

#### Object attributes

Following object attributes are updated during initialization:

#### Variables

- **param** (*dict*) – The parameter dictionary is updated with 'a0' and 'a2'.
- **subprocess** (*dict*) – subprocess 'insolation' is overwritten by [DailyInsolation](#) (page 101).

if 'ai' is not given:

#### Variables

- **param** (*dict*) – 'ai' and 'Tf' are removed from the parameter dictionary (initialized by parent class [EBM](#) (page 73))
- **subprocess** (*dict*) – subprocess 'albedo' is overwritten by [P2Albedo](#) (page 111).

if 'ai' is given:

#### Variables

- **param** (*dict*) – The parameter dictionary is updated with 'ai'.

- **subprocess** (*dict*) – subprocess 'albedo' is overwritten by `StepFunctionAlbedo` (page 113) (which basically has been there before but now is updated with the new albedo parameter values).

**Example** The annual distribution of solar insolation:

```
import climlab
from climlab.utils import constants as const
import numpy as np
import matplotlib.pyplot as plt

# creating model
model = climlab.EBM_seasonal()
model.step_forward()

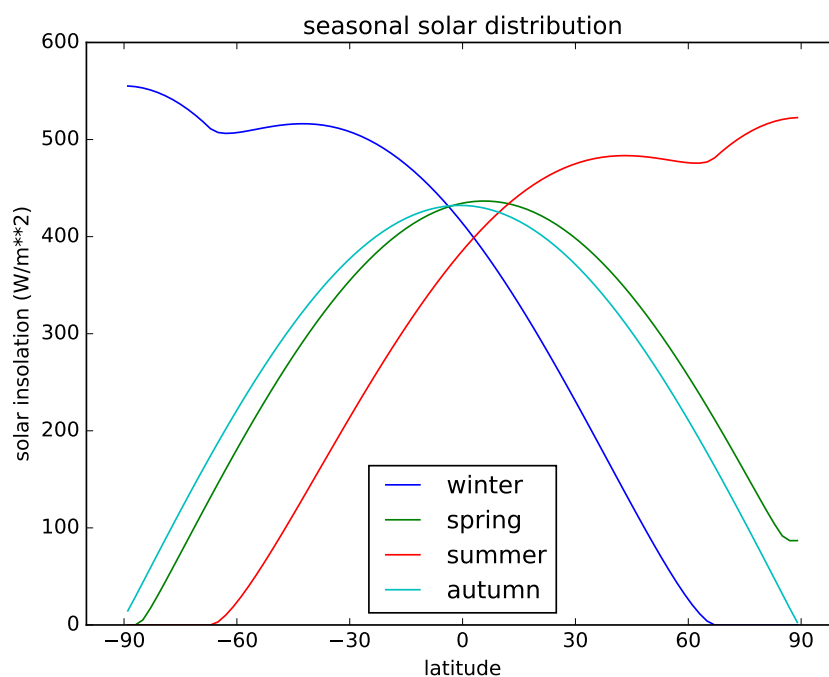
solar = model.subprocess['insolation'].insolation

# plot
fig = plt.figure()
ax = fig.add_subplot(111)

season_days = const.days_per_year/4

for season in ['winter', 'spring', 'summer', 'autumn']:
    ax.plot(model.lat, solar, label=season)
    model.integrate_days(season_days)

ax.set_title('seasonal solar distribution')
ax.set_xlabel('latitude')
ax.set_xticks([-90, -60, -30, 0, 30, 60, 90])
ax.set_ylabel('solar insolation (W/m**2)')
ax.legend(loc='best')
plt.show()
```



## 6.1.4 climlab.process package

### climlab.process.diagnostic module



**class** climlab.process.diagnostic.**DiagnosticProcess** (\*\*kwargs)

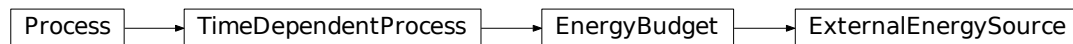
Bases: [climlab.process.time\\_dependent\\_process.TimeDependentProcess](#) (page 89)

A parent class for all processes that are strictly diagnostic, namely no time dependence.

During initialization following attribute is set:

**Variables** `time_type` (*str*) – is set to 'diagnostic'

### climlab.process.energy\_budget module



**class** climlab.process.energy\_budget.**EnergyBudget** (\*\*kwargs)

Bases: [climlab.process.time\\_dependent\\_process.TimeDependentProcess](#) (page 89)

A parent class for explicit energy budget processes.

This class solves equations that include a heat capacity term like  $C \frac{dT}{dt} = \text{flux convergence}$

In an Energy Balance Model with model state  $T$  this equation will look like this:

$$C \frac{dT}{dt} = R_{\downarrow} - R_{\uparrow} - H$$

$$\frac{dT}{dt} = \frac{R_{\downarrow}}{C} - \frac{R_{\uparrow}}{C} - \frac{H}{C}$$

Every EnergyBudget object has a `heating_rate` dictionary with items corresponding to each state variable. The heating rate accounts the actual heating of a subprocess, namely the contribution to the energy budget of  $R_{\downarrow}$ ,  $R_{\uparrow}$  and  $H$  in this case. The temperature tendencies for each subprocess are then calculated through dividing the heating rate by the heat capacity  $C$ .

#### Initialization parameters

An instance of EnergyBudget is initialized with the forwarded keyword arguments `**kwargs` of the corresponding children classes.

#### Object attributes

Additional to the parent class TimeDependentProcess following object attributes are generated or modified during initialization:

##### Variables

- **time\_type** (*str*) – is set to 'explicit'
- **heating\_rate** (*dict*) – energy share for given subprocess in unit W/m<sup>2</sup> stored in a dictionary sorted by model states

**class** climlab.process.energy\_budget.**ExternalEnergySource** (*\*\*kwargs*)

Bases: [climlab.process.energy\\_budget.EnergyBudget](#) (page 81)

A fixed energy source or sink to be specified by the user.

### Object attributes

Additional to the parent class [EnergyBudget](#) (page 81) the following object attribute is modified during initialization:

**Variables** **heating\_rate** (*dict*) – energy share dictionary for this subprocess is set to zero for every model state.

After initialization the user should modify the fields in the heating\_rate dictionary, which contain heating rates in unit  $\text{W/m}^2$  for all state variables.

**Example** Creating an Energy Balance Model with a uniform external energy source of  $10 \text{ W/m}^2$  for all latitudes:

```
>>> import climlab
>>> from climlab.process.energy_budget import ExternalEnergySource
>>> import numpy as np

>>> # create model & external energy subprocess
>>> model = climlab.EBM(num_lat=36)
>>> ext_en = ExternalEnergySource(state= model.state,**model.param)

>>> # modify external energy rate
>>> ext_en.heating_rate.keys()
['Ts']

>>> np.squeeze(ext_en.heating_rate['Ts'])
Field([ -0., -0., -0., -0., -0., -0., -0., -0., -0.,  0.,  0.,  0.,  0.,
         0.,  0.,  0.,  0.,  0.,  0.,  0.,  0.,  0.,  0.,  0.,  0.,
         0., -0., -0., -0., -0., -0., -0., -0., -0., -0., -0.])

>>> ext_en.heating_rate['Ts'][:]=10

>>> np.squeeze(ext_en.heating_rate['Ts'])
Field([ 10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,
        10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,
        10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,  10.,
        10.,  10.,  10.])

>>> # add subprocess to model
>>> model.add_subprocess('ext_energy',ext_en)

>>> print model
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (36, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  ext_energy: <class 'climlab.process.energy_budget.ExternalEnergySource'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.P2Insolation'>
```

## climlab.process.implicit module



**class** `climlab.process.implicit.ImplicitProcess` (*\*\*kwargs*)

Bases: `climlab.process.time_dependent_process.TimeDependentProcess` (page 89)

A parent class for modules that use implicit time discretization.

During initialization following attributes are initialized:

### Variables

- **time\_type** (*str*) – is set to 'implicit'
- **adjustment** (*dict*) – the model state adjustments due to this implicit subprocess

### `_compute()`

Computes the state variable tendencies in time for implicit processes.

To calculate the new state the `_implicit_solver()` method is called for daughter classes. This however returns the new state of the variables, not just the tendencies. Therefore the adjustment is calculated which is the difference between the new and the old state and stored in the object's attribute `adjustment`.

Calculating the new model states through solving the matrix problem already includes the multiplication with the timestep. The derived adjustment is divided by the timestep to calculate the implicit subprocess tendencies, which can be handled by the `compute()` (page 90) method of the parent `TimeDependentProcess` (page 89) class.

**Variables** **adjustment** (*dict*) – holding all state variables' adjustments of the implicit process which are the differences between the new states (which have been solved through matrix inversion) and the old states.

## climlab.process.process module

Process

**class** `climlab.process.process.Process` (*state=None, domains=None, subprocess=None, lat=None, lev=None, num\_lat=None, num\_levels=None, input=None, \*\*kwargs*)

Bases: `object`

A generic parent class for all climlab process objects. Every process object has a set of state variables on a spatial grid.

For more general information about *Processes* and their role in climlab, see [Process](#) (page 7) section climlab-architecture.

### Initialization parameters

An instance of `Process` is initialized with the following arguments (*for detailed information see Object attributes below*):

#### Parameters

- **state** (`Field` (page 63)) – spatial state variable for the process. Set to `None` if not specified.
- **domains** (`_Domain` (page 59) or dict of `_Domain` (page 59)) – domain(s) for the process
- **subprocess** (`Process` (page 83) or dict of `Process` (page 83)) – subprocess(es) of the process
- **lat** (`array`) – latitudinal points [optional]
- **lev** – altitudinal points [optional]
- **num\_lat** (`int`) – number of latitudinal points [optional]
- **num\_levels** (`int`) – number of altitudinal points [optional]
- **input** (`dict`) – collection of input quantities

#### Object attributes

Additional to the parent class `Process` (page 83) following object attributes are generated during initialization:

#### Variables

- **domains** (`dict`) – dictionary of process `_Domain` (page 59)
- **state** (`dict`) – dictionary of process states (of type `Field` (page 63))
- **param** (`dict`) – dictionary of model parameters which are given through `**kwargs`
- **diagnostics** (`dict`) – a dictionary with all diagnostic variables
- **\_input\_vars** (`dict`) – collection of input quantities like boundary conditions and other gridded quantities
- **creation\_date** (`str`) – date and time when process was created
- **subprocess** (dict of `Process` (page 83)) – dictionary of subprocesses of the process

#### `add_input` (`inputlist`)

Updates the process's list of inputs.

**Parameters** `inputlist` (`list`) – list of names of input variables

#### `add_subprocess` (`name`, `proc`)

Adds a single subprocess to this process.

#### Parameters

- **name** (`string`) – name of the subprocess
- **proc** (`Process` (page 83)) – a `Process` object

**Raises** `ValueError` if `proc` is not a process

**Example** Replacing an albedo subprocess through adding a subprocess with same name:

```
>>> from climlab.model.ebm import EBM_seasonal
>>> from climlab.surface.albedo import StepFunctionAlbedo

>>> # creating EBM model
>>> ebm_s = EBM_seasonal()

>>> print ebm_s
```



```

climlab Process of type <class 'climlab.model.ebm.EBM_seasonal'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM_seasonal'>
    diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
    LW: <class 'climlab.radiation.AplusBT.AplusBT'>
    albedo: <class 'climlab.surface.albedo.P2Albedo'>
    insolation: <class 'climlab.radiation.insolation.DailyInsolation'>

```

```

>>> # creating and adding albedo feedback subprocess
>>> step_albedo = StepFunctionAlbedo(state=ebm_s.state, **ebm_s.param)
>>> ebm_s.add_subprocess('albedo', step_albedo)
>>>
>>> print ebm_s

```

```

climlab Process of type <class 'climlab.model.ebm.EBM_seasonal'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM_seasonal'>
    diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
    LW: <class 'climlab.radiation.AplusBT.AplusBT'>
    albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
        iceline: <class 'climlab.surface.albedo.Iceline'>
        cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
        warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
    insolation: <class 'climlab.radiation.insolation.DailyInsolation'>

```

### **add\_subprocesses** (*procdict*)

Adds a dictionary of subprocesses to this process.

Calls *add\_subprocess()* (page 84) for every process given in the input-dictionary. It can also pass a single process, which will be given the name *default*.

**Parameters** *procdict* (*dict*) – a dictionary with process names as keys

### **depth**

Property of depth points of the process.

**Getter** Returns the points of axis 'depth' if available in the process's domains.

**Type** array

**Raises** ValueError if no 'depth' axis can be found.

### **depth\_bounds**

Property of depth bounds of the process.

**Getter** Returns the bounds of axis 'depth' if available in the process's domains.

**Type** array

**Raises** ValueError if no 'depth' axis can be found.

### **init\_diagnostic** (*name*, *value=0.0*)

Defines a new diagnostic quantity called *name* and initialize it with the given *value*.

Quantity is accessible and settable in two ways:

- as a process attribute, i.e. *proc.name*
- as a member of the diagnostics dictionary, i.e. *proc.diagnostics['name']*

**Parameters**

- **name** (*str*) – name of diagnostic quantity to be initialized
- **value** (*array*) – initial value for quantity - accepts also type float, int, etc. (*default:0.*)

**Example** Add a diagnostic CO2 variable to an energy balance model:

```
>>> import climlab
>>> model = climlab.EBM()

>>> # initialize CO2 variable with value 280 ppm
>>> model.init_diagnostic('CO2',280)

>>> # access variable directly or through diagnostic dictionary
>>> model.CO2
280
>>> model.diagnostics.keys()
['ASR', 'CO2', 'net_radiation', 'icelat', 'OLR', 'albedo']
```

### **input**

dictionary with all input variables

That can be boundary conditions and other gridded quantities independent of the *process*

**Getter** Returns the content of `self._input_vars`.

**Type** dict

### **lat**

Property of latitudinal points of the process.

**Getter** Returns the points of axis 'lat' if available in the process's domains.

**Type** array

**Raises** `ValueError` if no 'lat' axis can be found.

### **lat\_bounds**

Property of latitudinal bounds of the process.

**Getter** Returns the bounds of axis 'lat' if available in the process's domains.

**Type** array

**Raises** `ValueError` if no 'lat' axis can be found.

### **lev**

Property of altitudinal points of the process.

**Getter** Returns the points of axis 'lev' if available in the process's domains.

**Type** array

**Raises** `ValueError` if no 'lev' axis can be found.

### **lev\_bounds**

Property of altitudinal bounds of the process.

**Getter** Returns the bounds of axis 'lev' if available in the process's domains.

**Type** array

**Raises** `ValueError` if no 'lev' axis can be found.

### **lon**

Property of longitudinal points of the process.

**Getter** Returns the points of axis 'lon' if available in the process's domains.

**Type** array

**Raises** `ValueError` if no `'lon'` axis can be found.

#### `lon_bounds`

Property of longitudinal bounds of the process.

**Getter** Returns the bounds of axis `'lon'` if available in the process's domains.

**Type** `array`

**Raises** `ValueError` if no `'lon'` axis can be found.

#### `remove_diagnostic(name)`

Removes a diagnostic from the `process.diagnostic` dictionary and also delete the associated process attribute.

**Parameters** `name (str)` – name of diagnostic quantity to be removed

**Example** Remove diagnostic variable `'icelat'` from energy balance model:

```
>>> import climlab
>>> model = climlab.EBM()

>>> # display all diagnostic variables
>>> model.diagnostics.keys()
['ASR', 'OLR', 'net_radiation', 'albedo', 'icelat']

>>> model.remove_diagnostic('icelat')
>>> model.diagnostics.keys()
['ASR', 'OLR', 'net_radiation', 'albedo']

>>> # Watch out for subprocesses that may still want
>>> # to access the diagnostic 'icelat' variable !!!
```

#### `remove_subprocess(name)`

Removes a single subprocess from this process.

**Parameters** `name (string)` – name of the subprocess

**Example** Remove albedo subprocess from energy balance model:

```
>>> import climlab
>>> model = climlab.EBM()

>>> print model
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.P2Insolation'>

>>> model.remove_subprocess('albedo')

>>> print model
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
```

```
diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
LW: <class 'climlab.radiation.AplusBT.AplusBT'>
insolation: <class 'climlab.radiation.insolation.P2Insolation'>
```

**set\_state** (*name, value*)

Sets the variable name to a new state value.

#### Parameters

- **name** (*string*) – name of the state
- **value** (*Field* (page 63) or *array*) – state variable

**Raises** `ValueError` if state variable value is not having a domain.

**Raises** `ValueError` if shape mismatch between existing domain and new state variable.

**Example** Resetting the surface temperature of an EBM to  $-5^{\circ}\text{C}$  on all latitudes:

```
>>> import climlab
>>> from climlab import Field
>>> import numpy as np

>>> # setup model
>>> model = climlab.EBM(num_lat=36)

>>> # create new temperature distribution
>>> initial = -5 * ones(size(model.lat))
>>> model.set_state('Ts', Field(initial, domain=model.domains['Ts']))

>>> np.squeeze(model.Ts)
Field([-5., -5., -5., -5., -5., -5., -5., -5., -5., -5., -5., -5., -5.,
      -5., -5., -5., -5., -5., -5., -5., -5., -5., -5., -5., -5., -5.,
      -5., -5., -5., -5., -5., -5., -5., -5., -5., -5.])
```

`climlab.process.process.get_axes` (*process\_or\_domain*)

Returns a dictionary of all Axis in a domain or dictionary of domains.

**Parameters** **process\_or\_domain** (*Process* (page 83) or *Domain* (page 59)) – a process or a domain object

#### Raises

**exc** `TypeError` if input is not or not having a domain

**Returns** dictionary of input's Axis

**Return type** `dict`

#### Example

```
>>> import climlab
>>> from climlab.process.process import get_axes

>>> model = climlab.EBM()

>>> get_axes(model)
{'lat': <climlab.domain.axis.Axis object at 0x7ff13b9dd2d0>,
 'depth': <climlab.domain.axis.Axis object at 0x7ff13b9dd310>}
```

`climlab.process.process.process_like` (*proc*)

Copys the given process.

The creation date is updated.

**Parameters** **proc** (*Process* (page 83)) – process

**Returns** new process identical to the given process

Return type [Process](#) (page 83)

### Example

```
>>> import climlab
>>> from climlab.process.process import process_like

>>> model = climlab.EBM()
>>> model.subprocess.keys()
['diffusion', 'LW', 'albedo', 'insolation']

>>> albedo = model.subprocess['albedo']
>>> albedo_copy = process_like(albedo)

>>> albedo.creation_date
'Thu, 24 Mar 2016 01:32:25 +0000'

>>> albedo_copy.creation_date
'Thu, 24 Mar 2016 01:33:29 +0000'
```

## climlab.process.time\_dependent\_process module



**class** climlab.process.time\_dependent\_process.**TimeDependentProcess** (*time\_type='explicit', timestep=None, top-down=True, \*\*kwargs*)

Bases: [climlab.process.process.Process](#) (page 83)

A generic parent class for all time-dependent processes.

`TimeDependentProcess` is a child of the [Process](#) (page 83) class and therefore inherits all those attributes.

### Initialization parameters

An instance of `TimeDependentProcess` is initialized with the following arguments (*for detailed information see Object attributes below*):

#### Parameters

- **timestep** (*float*) – specifies the timestep of the object
- **time\_type** (*str*) – how time-dependent-process should be computed. Set to 'explicit' by default.
- **topdown** (*bool*) – whether generate *process\_types* in regular or in reverse order. Set to True by default.

### Object attributes

Additional to the parent class [Process](#) (page 83) following object attributes are generated during initialization:

#### Variables

- **has\_process\_type\_list** (*bool*) – information whether attribute *process\_types* (which is needed for *compute()* (page 90) and build in *\_build\_process\_type\_list()*) exists or not. Attribute is set to 'False' during initialization.
- **topdown** (*bool*) – information whether the list *process\_types* (which contains all processes and sub-processes) should be generated in regular or in reverse order. See *\_build\_process\_type\_list()*.
- **timeave** (*dict*) – a time averaged collection of all states and diagnostic processes over the timeperiod that *integrate\_years()* (page 91) has been called for last.
- **tendencies** (*dict*) – computed difference in a timestep for each state. See *compute()* (page 90) for details.
- **time\_type** (*str*) – how time-dependent-process should be computed. Possible values are: 'explicit', 'implicit', 'diagnostic', 'adjustment'.
- **time** (*dict*) –  
**a collection of all time-related attributes of the process.** The dictionary contains following items:
  - 'timestep': see initialization parameter
  - 'num\_steps\_per\_year': see *set\_timestep()* (page 92) and *timestep()* (page 92) for details
  - 'day\_of\_year\_index': counter how many steps have been integrated in current year
  - 'steps': counter how many steps have been integrated in total,
  - 'days\_elapsed': time counter for days,
  - 'years\_elapsed': time counter for years,
  - 'days\_of\_year': array which holds the number of numerical steps per year, expressed in days

#### **compute()**

Computes the tendencies for all state variables given current state and specified input.

The function first computes all diagnostic processes as they may effect all the other processes (such as change in solar distribution). After all the diagnostic processes don't produce any tendencies directly. Subsequently all tendencies and diagnostics for all explicit processes are computed.

Tendencies due to implicit and adjustment processes need to be calculated from a state that is already adjusted after explicit alteration. So the explicit tendencies are applied to the states temporarily. Now all tendencies from implicit processes are calculated through matrix inversions and same like the explicit tendencies applied to the states temporarily. Subsequently all instantaneous adjustments are computed.

Then the changes made to the states from explicit and implicit processes are removed again as this *compute()* (page 90) function is supposed to calculate only tendencies and not applying them to the states.

Finally all calculated tendencies from all processes are collected for each state, summed up and stored in the dictionary *self.tendencies*, which is an attribute of the time-dependent-process object for which the *compute()* (page 90) method has been called.

#### **Object attributes**

During method execution following object attributes are modified:

##### **Variables**

- **tendencies** (*dict*) – dictionary that holds tendencies for all states is calculated for current timestep through adding up tendencies from explicit, implicit and adjustment processes.
- **diagnostics** (*dict*) – process diagnostic dictionary is updated by diagnostic dictionaries of subprocesses after computation of tendencies.

**compute\_diagnostics** (*num\_iter=3*)

Compute all tendencies and diagnostics, but don't update model state. By default it will call `compute()` 3 times to make sure all subprocess coupling is accounted for. The number of iterations can be changed with the input argument.

**integrate\_converge** (*crit=0.0001, verbose=True*)

Integrates the model until model states are converging.

#### Parameters

- **crit** (*float*) – exit criteria for difference of iterated solutions
- **verbose** (*bool*) – information whether total elapsed time should be printed.

#### Example

```
>>> import climlab
>>> model = climlab.EBM()

>>> model.global_mean_temperature()
Field(11.997968598413685)

>>> model.integrate_converge()
Total elapsed time is 10.0 years.

>>> model.global_mean_temperature()
Field(14.288155406577301)
```

**integrate\_days** (*days=1.0, verbose=True*)

Integrates the model forward for a specified number of days.

It converts the given number of days into years and calls `integrate_years()` (page 91).

#### Parameters

- **days** (*float*) – integration time for the model in days
- **verbose** (*bool*) – information whether model time details should be printed.

#### Example

```
>>> import climlab
>>> model = climlab.EBM()

>>> model.global_mean_temperature()
Field(11.997968598413685)

>>> model.integrate_days(80.)
Integrating for 19 steps, 80.0 days, or 0.219032740466 years.
Total elapsed time is 0.211111111111 years.

>>> model.global_mean_temperature()
Field(11.873680783355553)
```

**integrate\_years** (*years=1.0, verbose=True*)

Integrates the model by a given number of years.

#### Parameters

- **years** (*float*) – integration time for the model in years

- **verbose** (*bool*) – information whether model time details should be printed.

It calls `step_forward()` (page 92) repetitively and calculates a time averaged value over the integrated period for every model state and all diagnostics processes.

#### Example

```
>>> import climlab
>>> model = climlab.EBM()

>>> model.global_mean_temperature()
Field(11.997968598413685)

>>> model.integrate_years(2.)
Integrating for 180 steps, 730.4844 days, or 2.0 years.
Total elapsed time is 2.0 years.

>>> model.global_mean_temperature()
Field(13.531055349437258)
```

**set\_timestep** (*timestep=86400.0, num\_steps\_per\_year=None*)

Calculates the timestep in unit seconds and calls the setter function of `timestep()` (page 92)

#### Parameters

- **timestep** (*float*) – the amount of time over which `step_forward()` (page 92) is integrating in unit seconds
- **num\_steps\_per\_year** (*float*) – a number of steps per calendar year

If the parameter `num_steps_per_year` is specified and not `None`, the timestep is calculated accordingly and therefore the given input parameter `timestep` is ignored.

**step\_forward()**

Updates state variables with computed tendencies.

Calls the `compute()` (page 90) method to get current tendencies for all process states. Multiplied with the timestep and added up to the state variables is updating all model states.

#### Example

```
>>> import climlab
>>> model = climlab.EBM()

>>> # checking time step counter
>>> model.time['steps']
0

>>> # stepping the model forward
>>> model.step_forward()

>>> # step counter increased
>>> model.time['steps']
1
```

**timestep**

The amount of time over which `step_forward()` (page 92) is integrating in unit seconds.

**Getter** Returns the object timestep which is stored in `self.param['timestep']`.

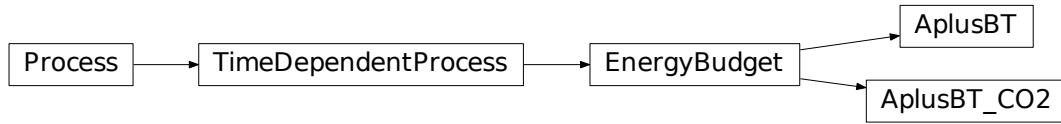
**Setter** Sets the timestep to the given input. See also `set_timestep()` (page 92).

**Type** float



## 6.1.5 climlab.radiation package

### climlab.radiation.AplusBT module



**class** climlab.radiation.AplusBT.**AplusBT** ( $A=200.0$ ,  $B=2.0$ , *\*\*kwargs*)  
 Bases: [climlab.process.energy\\_budget.EnergyBudget](#) (page 81)

The simplest linear longwave radiation module.

Calculates the Outgoing Longwave Radation (OLR)  $R \uparrow$  as

$$R \uparrow = A + B \cdot T$$

where  $T$  is the state variable.

Should be invoked with a single temperature state variable only.

#### Initialization parameters

An instance of `AplusBT` is initialized with the following arguments:

##### Parameters

- **A** (*float*) – parameter for linear OLR parameterization
  - unit:  $\frac{\text{W}}{\text{m}^2}$
  - default value: 200.0
- **B** (*float*) – parameter for linear OLR parameterization
  - unit:  $\frac{\text{W}}{\text{m}^2 \cdot ^\circ\text{C}}$
  - default value: 2.0

#### Object attributes

Additional to the parent class [EnergyBudget](#) (page 81) following object attributes are generated or modified during initialization:

##### Variables

- **A** (page 94) (*float*) – calls the setter function of `A()` (page 94)
- **B** (page 94) (*float*) – calls the setter function of `B()` (page 94)
- **diagnostics** (*dict*) – key 'OLR' initialized with value: [Field](#) (page 63) of zeros in size of `self.Ts`
- **OLR** ([Field](#) (page 63)) – the subprocess attribute `self.OLR` is created with correct dimensions

**Warning:** This module currently works only for a single state variable!

**Example** Simple linear radiation module (stand alone):

```
>>> import climlab

>>> # create a column atmosphere and scalar surface
```

```
>>> sfc, atm = climlab.domain.single_column()

>>> # Create a state variable
>>> Ts = climlab.Field(15., domain=sfc)

>>> # Make a dictionary of state variables
>>> s = {'Ts': Ts}

>>> # create process
>>> olr = climlab.radiation.AplusBT(state=s)

>>> print olr
climlab Process of type <class 'climlab.radiation.AplusBT.AplusBT'>.
State variables and domain shapes:
  Ts: (1,)
The subprocess tree:
top: <class 'climlab.radiation.AplusBT.AplusBT'>

>>> # to compute tendencies and diagnostics
>>> olr.compute()

>>> # or to actually update the temperature
>>> olr.step_forward()

>>> print olr.state
{'Ts': Field([ 5.69123176])}
```

#### A

Property of AplusBT parameter A.

**Getter** Returns the parameter A which is stored in attribute `self._A`

**Setter**

- sets parameter A which is addressed as `self._A` to the new value
- updates the parameter dictionary `self.param['A']`

**Type** float

**Example**

```
>>> import climlab
>>> model = climlab.EBM()

>>> # getter
>>> model.subprocess['LW'].A
210.0
>>> # setter
>>> model.subprocess['LW'].A = 220
>>> # getter again
>>> model.subprocess['LW'].A
220

>>> # subprocess parameter dictionary
>>> model.subprocess['LW'].param['A']
220
```

#### B

Property of AplusBT parameter B.

**Getter** Returns the parameter B which is stored in attribute `self._B`

**Setter**

- sets parameter B which is addressed as `self._B` to the new value

- updates the parameter dictionary `self.param['B']`

**Type** float

**class** `climlab.radiation.AplusBT.AplusBT_CO2` ( $CO_2=300.0$ , *\*\*kwargs*)

Bases: `climlab.process.energy_budget.EnergyBudget` (page 81)

Linear longwave radiation module considering  $CO_2$  concentration.

This radiation subprocess is based in the idea to linearize the Outgoing Longwave Radiation (OLR) emitted to space according to the surface temperature (see [AplusBT](#) (page 93)).

To consider a the change of the greenhouse effect through range of  $CO_2$  in the atmosphere, the parameters A and B are computed like the following:

$$A(c) = -326.4 + 9.161c - 3.164c^2 + 0.5468c^3$$

$$B(c) = 1.953 - 0.04866c + 0.01309c^2 - 0.002577c^3$$

where  $c = \log \frac{p}{300}$  and  $p$  represents the concentration of  $CO_2$  in the atmosphere.

For further reading see [\[CaldeiraKasting1992\]](#) (page 129).

### Initialization parameters

An instance of `AplusBT_CO2` is initialized with the following argument:

**Parameters** `CO2` (*float*) – The concentration of  $CO_2$  in the atmosphere. Referred to as  $p$  in the above given formulas.

- unit: ppm (parts per million)
- default value: `300.0`

### Object attributes

Additional to the parent class [EnergyBudget](#) (page 81) following object attributes are generated or updated during initialization:

#### Variables

- `CO2` (page 96) (*float*) – calls the setter function of `CO2()` (page 96)
- `diagnostics` (*dict*) – the subprocess's diagnostic dictionary `self.diagnostic` is initialized through calling `self.init_diagnostic('OLR', 0. * self.Ts)`
- `OLR` ([Field](#) (page 63)) – the subprocess attribute `self.OLR` is created with correct dimensions

**Example** Replacing an the regular `AplusBT` subprocess in an energy balance model:

```
>>> import climlab
>>> from climlab.radiation.AplusBT import AplusBT_CO2

>>> # creating EBM model
>>> model = climlab.EBM()

>>> print model
```

```
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
  iceline: <class 'climlab.surface.albedo.Iceline'>
```

```

cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
insolation: <class 'climlab.radiation.insolation.P2Insolation'>

```

```

>>> # creating and adding albedo feedback subprocess
>>> LW_CO2 = AplusBT_CO2(CO2=400, state=model.state, **model.param)

>>> # overwriting old 'LW' subprocess with same name
>>> model.add_subprocess('LW', LW_CO2)

>>> print model

```

```

climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT_CO2'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.P2Insolation'>

```

## CO2

Property of AplusBT\_CO2 parameter CO2.

**Getter** Returns the CO2 concentration which is stored in attribute `self._CO2`

**Setter**

- sets the CO2 concentration which is addressed as `self._CO2` to the new value
- updates the parameter dictionary `self.param['CO2']`

**Type** float

## climlab.radiation.Boltzmann module



**class** `climlab.radiation.Boltzmann.Boltzmann` (*eps*=0.65, *tau*=0.95, *\*\*kwargs*)

Bases: `climlab.process.energy_budget.EnergyBudget` (page 81)

A class for black body radiation.

Implements a radiation subprocess which computes longwave radiation with the Stefan-Boltzmann law for black/grey body radiation.

According to the Stefan Boltzmann law the total power radiated from an object with surface area  $A$  and temperature  $T$  (in unit Kelvin) can be written as

$$P = A\epsilon\sigma T^4$$

where  $\epsilon$  is the emissivity of the body.

As the *EnergyBudget* (page 81) of the Energy Balance Model is accounted in unit energy/area (W/m<sup>2</sup>) the energy budget equation looks like this:

$$C \frac{dT}{dt} = R \downarrow - R \uparrow - H$$

The *Boltzmann* (page 96) radiation subprocess represents the outgoing radiation  $R \uparrow$  which then can be written as

$$R \uparrow = \varepsilon \sigma T^4$$

with state variable  $T$ .

### Initialization parameters

An instance of `Boltzmann` is initialized with the following arguments:

#### Parameters

- **eps** (*float*) – emissivity of the planet’s surface which is the effectiveness in emitting energy as thermal radiation
  - unit: dimensionless
  - default value: 0.65
- **tau** (*float*) – transmissivity of the planet’s atmosphere which is the effectiveness in transmitting the longwave radiation emitted from the surface
  - unit: dimensionless
  - default value: 0.95

### Object attributes

During initialization both arguments described above are created as object attributes which calls their setter function (see below).

#### Variables

- **eps** (page 98) (*float*) – calls the setter function of `eps()` (page 98)
- **tau** (page 98) (*float*) – calls the setter function of `tau()` (page 98)
- **diagnostics** (*dict*) – the subprocess’s diagnostic dictionary `self.diagnostic` is initialized through calling `self.init_diagnostic('OLR', 0. * self.Ts)`
- **OLR** (*Field* (page 63)) – the subprocess attribute `self.OLR` is created with correct dimensions

**Example** Replacing an the regular `AplusBT` subprocess in an energy balance model:

```
>>> import climlab
>>> from climlab.radiation.Boltzmann import Boltzmann

>>> # creating EBM model
>>> model = climlab.EBM()

>>> print model
```

```
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
```

```

albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
iceline: <class 'climlab.surface.albedo.Iceline'>
cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
insolation: <class 'climlab.radiation.insolation.P2Insolation'>

```

```

>>> # creating and adding albedo feedback subprocess
>>> LW_boltz = Boltzmann(eps=0.69, tau=0.98, state=model.state, **model.param)

>>> # overwriting old 'LW' subprocess with same name
>>> model.add_subprocess('LW', LW_boltz)

>>> print model

```

```

climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
LW: <class 'climlab.radiation.Boltzmann.Boltzmann'>
albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
iceline: <class 'climlab.surface.albedo.Iceline'>
cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
insolation: <class 'climlab.radiation.insolation.P2Insolation'>

```

### eps

Property of emissivity parameter.

**Getter** Returns the albedo value which is stored in attribute `self._eps`

#### Setter

- sets the emissivity which is addressed as `self._eps` to the new value
- updates the parameter dictionary `self.param['eps']`

**Type** float

### tau

Property of the transmissivity parameter.

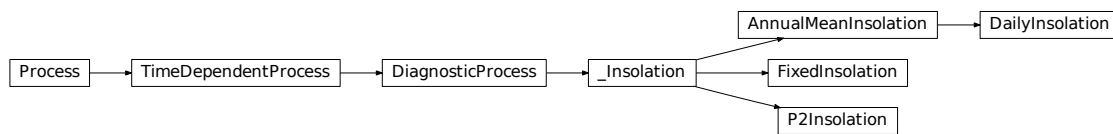
**Getter** Returns the albedo value which is stored in attribute `self._tau`

#### Setter

- sets the emissivity which is addressed as `self._tau` to the new value
- updates the parameter dictionary `self.param['tau']`

**Type** float

## climlab.radiation.insolation module



```

class climlab.radiation.insolation.AnnualMeanInsolation (S0=1365.2,
                                                         orb={'long_peri':
                                                         281.37, 'ecc': 0.017236,
                                                         'obliquity': 23.446},
                                                         **kwargs)

```

Bases: `climlab.radiation.insolation._Insolation` (page 103)

A class for latitudewise solar insolation averaged over a year.

This class computes the solar insolation for each day of the year and latitude specified in the domain on the basis of orbital parameters and astronomical formulas.

Therefor it uses the method `daily_insolation()` (page 105). For details how the solar distribution is dependend on orbital parameters see there.

The mean over the year is calculated from data given by `daily_insolation()` (page 105) and stored in the object's attribute `self.insolation`

### Initialization parameters

#### Parameters

- **S0** (*float*) – solar constant
  - unit:  $\frac{W}{m^2}$
  - default value: 1365.2
- **orb** (*dict*) – a dictionary with three orbital parameters (as provided by `OrbitalTable` (page 107)):
  - 'ecc' - eccentricity
    - \* unit: dimensionless
    - \* default value: 0.017236
  - 'long\_peri' - longitude of perihelion (precession angle)
    - \* unit: degrees
    - \* default value: 281.37
  - 'obliquity' - obliquity angle
    - \* unit: degrees
    - \* default value: 23.446

### Object attributes

Additional to the parent class `_Insolation` (page 103) following object attributes are generated and updated during initialization:

#### Variables

- **insolation** (page 99) (*Field* (page 63)) – the solar distribution is calculated as a `Field` on the basis of the `self.domains['default']` domain and stored in the attribute `self.insolation`.
- **orb** (page 100) (*dict*) – initialized with given argument `orb`

**Example** Create regular EBM and replace standard insolation subprocess by AnnualMeanInsolation:

```
>>> import climlab
>>> from climlab.radiation import AnnualMeanInsolation

>>> # model creation
>>> model = climlab.EBM()

>>> print model
```

```
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.P2Insolation'>
```

```
>>> # catch model domain for subprocess creation
>>> sfc = model.domains['Ts']

>>> # create AnnualMeanInsolation subprocess
>>> new_insol = AnnualMeanInsolation(domains=sfc, **model.param)

>>> # add it to the model
>>> model.add_subprocess('insolation', new_insol)

>>> print model
```

```
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.AnnualMeanInsolation'>
```

### orb

Property of dictionary for orbital parameters.

orb contains: (for more information see *OrbitalTable* (page 107))

- 'ecc' - eccentricity [unit: dimensionless]
- 'long\_peri' - longitude of perihelion (precession angle) [unit: degrees]
- 'obliquity' - obliquity angle [unit: degrees]

**Getter** Returns the orbital dictionary which is stored in attribute `self._orb`.

**Setter**



- sets orb which is addressed as `self._orb` to the new value
- updates the parameter dictionary `self.param['orb']` and
- calls method `_compute_fixed()`

**Type** dict

**class** `climlab.radiation.insolation.DailyInsolation` ( $S_0=1365.2$ , `orb`={`'long_peri'`: 281.37, `'ecc'`: 0.017236, `'obliquity'`: 23.446}, `**kwargs`)

Bases: `climlab.radiation.insolation.AnnualMeanInsolation` (page 99)

A class to compute latitudewise daily solar insolation for specific days of the year.

This class computes the solar insolation on basis of orbital parameters and astronomical formulas.

Therefor it uses the method `daily_insolation()` (page 105). For details how the solar distribution is dependent on orbital parameters see there.

### Initialization parameters

#### Parameters

- **`S0`** (*float*) – solar constant
  - unit:  $\frac{\text{W}}{\text{m}^2}$
  - default value: 1365.2
- **`orb`** (*dict*) – a dictionary with orbital parameters:
  - `'ecc'` - eccentricity
    - \* unit: dimensionless
    - \* default value: 0.017236
  - `'long_peri'` - longitude of perihelion (precession angle)
    - \* unit: degrees
    - \* default value: 281.37
  - `'obliquity'` - obliquity angle
    - \* unit: degrees
    - \* default value: 23.446

### Object attributes

Additional to the parent class `_Insolation` (page 103) following object attributes are generated and updated during initialization:

#### Variables

- **`insolation`** (page 99) (*Field* (page 63)) – the solar distribution is calculated as a Field on the basis of the `self.domains['default']` domain and stored in the attribute `self.insolation`.
- **`orb`** (page 100) (*dict*) – initialized with given argument `orb`

**Example** Create regular EBM and replace standard insolation subprocess by `DailyInsolation`:

```
>>> import climlab
>>> from climlab.radiation import DailyInsolation

>>> # model creation
>>> model = climlab.EBM()
```

```
>>> print model
```

```
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.P2Insolation'>
```

```
>>> # catch model domain for subprocess creation
>>> sfc = model.domains['Ts']

>>> # create DailyInsolation subprocess and add it to the model
>>> model.add_subprocess('insolation', DailyInsolation(domains=sfc, **model.param))

>>> print model
```

```
climlab Process of type <class 'climlab.model.ebm.EBM'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.DailyInsolation'>
```

**class** `climlab.radiation.insolation.FixedInsolation` ( $S_0=341.3$ , *\*\*kwargs*)

Bases: `climlab.radiation.insolation._Insolation` (page 103)

A class for fixed insolation at each point of latitude off the domain.

The solar distribution for the whole domain is constant and specified by a parameter.

#### Initialization parameters

**Parameters**  $S_0$  (*float*) – solar constant

- unit:  $\frac{\text{W}}{\text{m}^2}$
- default value:  $\text{const.S}_0/4 = 341.2$

#### Example

```
>>> import climlab
>>> from climlab.radiation.insolation import FixedInsolation

>>> model = climlab.EBM()
>>> sfc = model.Ts.domain

>>> fixed_ins = FixedInsolation(S0=340.0, domains=sfc)

>>> print fixed_ins
climlab Process of type <class 'climlab.radiation.insolation.FixedInsolation'>.
```

```
State variables and domain shapes:
The subprocess tree:
top: <class 'climlab.radiation.insolation.FixedInsolation'>
```

**class** climlab.radiation.insolation.**P2Insolation** ( $S_0=1365.2$ ,  $s_2=-0.48$ , **\*\*kwargs**)

Bases: [climlab.radiation.insolation.\\_Insolation](#) (page 103)

A class for parabolic solar distribution over the domain's latitude on the basis of the second order Legendre Polynomial.

Calculates the latitude dependent solar distribution as

$$S(\varphi) = \frac{S_0}{4} (1 + s_2 P_2(x))$$

where  $P_2(x) = \frac{1}{2}(3x^2 - 1)$  is the second order Legendre Polynomial and  $x = \sin(\varphi)$ .

### Initialization parameters

#### Parameters

- **S0** (*float*) – solar constant
  - unit:  $\frac{\text{W}}{\text{m}^2}$
  - default value: 1365.2
- **s2** (*float*) – factor for second legendre polynomial term
  - default value: -0.48

#### Example

```
>>> import climlab
>>> from climlab.radiation.insolation import P2Insolation

>>> model = climlab.EBM()
>>> sfc = model.Ts.domain

>>> p2_ins = P2Insolation(S0=340.0, s2=-0.5, domains=sfc)

>>> print p2_ins
climlab Process of type <class 'climlab.radiation.insolation.P2Insolation'>.
State variables and domain shapes:
The subprocess tree:
top: <class 'climlab.radiation.insolation.P2Insolation'>
```

### s2

Property of second legendre polynomial factor s2.

s2 in following equation:

$$S(\varphi) = \frac{S_0}{4} (1 + s_2 P_2(x))$$

**Getter** Returns the s2 parameter which is stored in attribute `self._s2`.

#### Setter

- sets s2 which is addressed as `self._S0` to the new value
- updates the parameter dictionary `self.param['s2']` and
- calls method `_compute_fixed()`

**Type** float

**class** `climlab.radiation.insolation._Insolation` ( $S_0=1365.2$ , *\*\*kwargs*)  
Bases: `climlab.process.diagnostic.DiagnosticProcess` (page 81)

A private parent class for insolation processes.

Calling `compute()` will update `self.insolation` with current values.

### Initialization parameters

An instance of `_Insolation` is initialized with the following arguments (*for detailed information see Object attributes below*):

**Parameters**  $S_0$  (*float*) – solar constant

- unit:  $\frac{\text{W}}{\text{m}^2}$
- default value: 1365.2

### Object attributes

Additional to the parent class `DiagnosticProcess` (page 81) following object attributes are generated and updated during initialization:

#### Variables

- ***insolation*** (page 99) (**Field** (page 63)) – the array is initialized with zeros of the size of `self.domains['sfc']` or `self.domains['default']`.
- ***S0*** (page 104) (*float*) – initialized with given argument  $S_0$
- ***diagnostics*** (*dict*) – key 'insolation' initialized with value: **Field** (page 63) of zeros in size of `self.domains['sfc']` or `self.domains['default']`
- ***insolation*** (page 99) – the subprocess attribute `self.insolation` is created with correct dimensions

---

**Note:** `self.insolation` should always be modified with `self.insolation[:] = ...` so that links to the insolation in other processes will work.

---

### $S_0$

Property of solar constant  $S_0$ .

The parameter  $S_0$  is stored using a python property and can be changed through `self.S0 = newvalue` which will also update the parameter dictionary.

**Warning:** changing `self.param['S0']` will not work!

**Getter** Returns the  $S_0$  parameter which is stored in attribute `self._S0`.

#### Setter

- sets  $S_0$  which is addressed as `self._S0` to the new value
- updates the parameter dictionary `self.param['S0']` and
- calls method `_compute_fixed()`

**Type** float

## 6.1.6 climlab.solar package

### climlab.solar.insolation module

This module contains general-purpose routines for computing incoming solar radiation at the top of the atmosphere.

Currently, only daily average insolation is computed.

---

**Note:** Ported and modified from MATLAB code `daily_insolation.m`

*Original authors:*

Ian Eisenman and Peter Huybers, Harvard University, August 2006

Available online at [http://eisenman.ucsd.edu/code/daily\\_insolation.m](http://eisenman.ucsd.edu/code/daily_insolation.m)

---

If using calendar days, solar longitude is found using an approximate solution to the differential equation representing conservation of angular momentum (Kepler's Second Law). Given the orbital parameters and solar longitude, daily average insolation is calculated exactly following [Berger1978] (page 129). Further references: [Berger1991] (page 129).

```
climlab.solar.insolation.daily_insolation(lat, day, orb={'long_peri': 281.37, 'ecc':
                                                    0.017236, 'obliquity': 23.446}, S0=None,
                                           day_type=1)
```

Compute daily average insolation given latitude, time of year and orbital parameters.

Orbital parameters can be computed for any time in the last 5 Myears with `lookup_parameters()` (page 107) (see example below).

### Function-call argument

#### Parameters

- **lat** (*array*) – Latitude in degrees (-90 to 90).
- **day** (*array*) – Indicator of time of year. See argument `day_type` for details about format.
- **orb** (*dict*) – a dictionary with three members (as provided by `OrbitalTable` (page 107))
  - `'ecc'` - eccentricity
    - \* unit: dimensionless
    - \* default value: 0.017236
  - `'long_peri'` - longitude of perihelion (precession angle)
    - \* unit: degrees
    - \* default value: 281.37
  - `'obliquity'` - obliquity angle
    - \* unit: degrees
    - \* default value: 23.446
- **S0** (*float*) – solar constant
  - unit: W/m<sup>2</sup>
  - default value: 1365.2
- **day\_type** (*int*) – Convention for specifying time of year (+/- 1,2) [optional].
  - day\_type=1 (default):** day input is calendar day (1-365.24), where day 1 is January first. The calendar is referenced to the vernal equinox which always occurs at day 80.
  - day\_type=2:** day input is solar longitude (0-360 degrees). Solar longitude is the angle of the Earth's orbit measured from spring equinox (21 March). Note that calendar days and solar longitude are not linearly related because, by Kepler's Second Law, Earth's angular velocity varies according to its distance from the sun.

**Raises** `ValueError` if `day_type` is neither 1 nor 2

### Returns

Daily average solar radiation in unit  $\text{W/m}^2$ .

Dimensions of output are `(lat.size, day.size, ecc.size)`

### Return type `array`

Code is fully vectorized to handle array input for all arguments.

Orbital arguments should all have the same sizes. This is automatic if computed from `lookup_parameters()` (page 107)

**Example** to compute the timeseries of insolation at 65N at summer solstice over the past 5 Myears:

```
from climlab.solar.orbital import OrbitalTable
from climlab.solar.insolation import daily_insolation

# import orbital table
table = OrbitalTable()

# array with specified kyears
years = np.linspace(-5000, 0, 5001)

# orbital parameters for specified time
orb = table.lookup_parameters( years )

# insolation values for past 5 Myears at 65N at summer solstice
S65 = daily_insolation( 65, 172, orb )
```

For more information about computation of solar insolation see the [Tutorials](#) (page 17) chapter.

```
climlab.solar.insolation.solar_longitude( day, orb={'long_peri': 281.37,
                                                    'ecc': 0.017236, 'obliquity': 23.446},
                                           days_per_year=None)
```

Estimates solar longitude from calendar day.

Method is using an approximation from [\[Berger1978\]](#) (page 129) section 3 ( $\lambda = 0$  at spring equinox).

### Function-call arguments

#### Parameters

- **day** (*array*) – Indicator of time of year.
- **orb** (*dict*) – a dictionary with three members (as provided by `OrbitalTable` (page 107))
  - `'ecc'` - eccentricity
    - \* unit: dimensionless
    - \* default value: 0.017236
  - `'long_peri'` - longitude of perihelion (precession angle)
    - \* unit: degrees
    - \* default value: 281.37
  - `'obliquity'` - obliquity angle
    - \* unit: degrees
    - \* default value: 23.446
- **days\_per\_year** (*float*) – number of days in a year (optional) (default: 365.2422)  
Reads the length of the year from [constants](#) (page 114) if available.

**Returns** solar longitude `lambda_long` in dimension“( day.size, ecc.size )“

**Return type** `array`

Works for both scalar and vector orbital parameters.

## climlab.solar.orbital module



This module defines the class `OrbitalTable` (page 107) which holds orbital data, and includes a method `lookup_parameters()` (page 107) which interpolates the orbital data for a specific year (- works equally well for arrays of years).

The base class `OrbitalTable()` (page 107) is designed to work with 5 Myears of orbital data (**eccentricity, obliquity, and longitude of perihelion**) from [\[Berger1991\]](#) (page 129).

Data will be read from the file `orbit91`, which was originally obtained from <ftp://ftp.ncdc.noaa.gov/pub/data/paleo/insolation/> If the file isn't found locally, the module will attempt to read it remotely from the above URL.

A subclass `LongOrbitalTable()` (page 107) works with La2004 orbital data for -51 to +21 Myears as calculated by [\[Laskar2004\]](#) (page 129). See <http://vo.imcce.fr/insola/earth/online/earth/La2004/README.TXT>

**class** `climlab.solar.orbital.LongOrbitalTable`

Bases: `climlab.solar.orbital.OrbitalTable` (page 107)

Loads orbital parameter tables for -51 to +21 Myears.

**Based on calculations by** [\[Laskar2004\]](#) (page 129) <http://vo.imcce.fr/insola/earth/online/earth/La2004/README.TXT>

Usage is identical to parent class `OrbitalTable()` (page 107).

**class** `climlab.solar.orbital.OrbitalTable`

Invoking `OrbitalTable()` will load 5 million years of orbital data from [\[Berger1991\]](#) (page 129) and compute linear interpolants.

The data can be accessed through the method `lookup_parameters()` (page 107).

### Object attributes

Following object attributes are generated during initialization:

#### Variables

- **kyear** (*array*) – time table with negative values are before present (*unit*: kyears)
- **ecc** (*array*) – eccentricity over time (*unit*: dimensionless)
- **long\_peri** (*array*) – longitude of perihelion (precession angle) (*unit*: degrees)
- **obliquity** (*array*) – obliquity angle (*unit*: degrees)
- **kyear\_min** (*float*) – minimum value of time table (*unit*: kyears)
- **kyear\_max** (*float*) – maximum value of time table (*unit*: kyears)

**lookup\_parameters** (*kyear=0*)

Look up orbital parameters for given kyear measured from present.

---

**Note:** Input `kyear` is thousands of years after present. For years before present, use `kyear < 0`.

---

### Function-call argument

**Parameters** `kyear` (*array*) – Time for which orbital parameters should be given. Will handle scalar or vector input (for multiple years).

### Returns

a three-member dictionary of orbital parameters:

- `'ecc'`: eccentricity (dimensionless)
- `'long_peri'`: longitude of perihelion relative to vernal equinox (degrees)
- `'obliquity'`: obliquity angle or axial tilt (degrees).

Each member is an array of same size as `kyear`.

**Return type** `dict`

## climlab.solar.orbital\_cycles module

OrbitalCycles

```
class climlab.solar.orbital_cycles.OrbitalCycles (model, kyear_start=-
20.0, kyear_stop=0.0, segment_length_years=100.0, or-
bital_year_factor=1.0, ver-
bose=True)
```

Automatically integrates a process through changes in orbital parameters.

`OrbitalCycles` is a module for setting up long integrations of `climlab` processes over orbital cycles.

The duration between integration start and end time is partitioned in time segments over which the orbital parameters are held constant. The process is integrated over every time segment and the process state `Ts` is stored for each segment.

The storage arrays are saving:

- **current model state** at end of each segment
- **model state averaged** over last integrated year of each segment
- **global mean** of averaged model state over last integrated year of each segment

---

**Note:** Input `kyear` is thousands of years after present. For years before present, use `kyear < 0`.

---

### Initialization parameters

#### Parameters

- **model** (*TimeDependentProcess* (page 89)) – a time dependent process
- **kyear\_start** (*float*) – integration start time.

As time reference is present, argument should be `< 0` for time before present.



- *unit*: kiloyears
- *default value*: -20.
- **kyear\_stop** (*float*) – integration stop time.  
As time reference is present, argument should be  $\leq 0$  for time before present.
- *unit*: kiloyears
- *default value*: 0.
- **segment\_length\_years** (*float*) – is the length of each integration with fixed orbital parameters. (default: 100.)
- **orbital\_year\_factor** (*float*) – is an optional speed-up to the orbital cycles. (default: 1.)
- **verbose** (*bool*) – prints product of calculation and information about computation progress if set to True (default).

### Object attributes

Following object attributes are generated during initialization:

#### Variables

- **model** (*TimeDependentProcess* (page 89)) – timedependent process to be integrated
- **kyear\_start** (*float*) – integration start time
- **kyear\_stop** (*float*) – integration stop time
- **segment\_length\_years** (*float*) – length of each integration with fixed orbital parameters
- **orbital\_year\_factor** (*float*) – speed-up factor to the orbital cycles
- **verbose** (*bool*) – print flag
- **num\_segments** (*int*) – number of segments with fixed orbital parameters, calculated through:

$$num_{seg} = \frac{-(kyear_{start} - kyear_{stop}) * 1000}{seglength * orb_{factor}}$$

- **T\_segments\_global** (*array*) – storage for global mean temperature for final year of each segment
- **T\_segments** (*array*) – storage for actual temperature at end of each segment
- **T\_segments\_annual** (*array*) – storage for timeaveraged temperature over last year of segment  
dimension: (size(Ts), num\_segments)
- **orb\_kyear** (*array*) – integration start time of all segments
- **orb** (page 100) (*dict*) – orbital parameters for last integrated segment

**Example** Integration of an energy balance model for 10,000 years with corresponding orbital parameters:

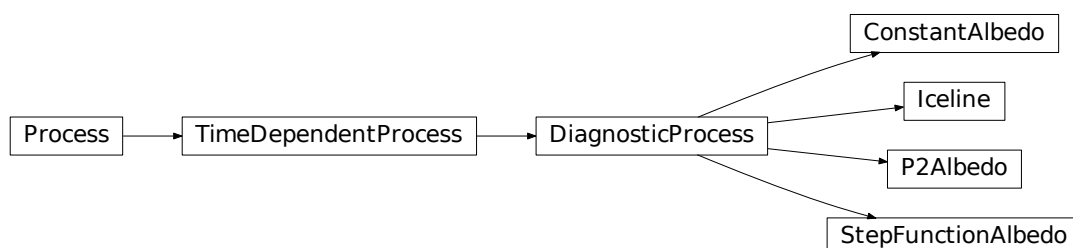
```
from climlab.model.ebm import EBM_seasonal
from climlab.solar.orbital_cycles import OrbitalCycles
from climlab.surface.albedo import StepFunctionAlbedo
ebm = EBM_seasonal()
print ebm
```

```
# add an albedo feedback
albedo = StepFunctionAlbedo(state=ebm.state, **ebm.param)
ebm.add_subprocess('albedo', albedo)

# start the integration
# run for 10,000 orbital years, but only 1,000 model years
experiment = OrbitalCycles(ebm, kyear_start=-20, kyear_stop=-10, orbital_year_factor=
```

## 6.1.7 climlab.surface package

### climlab.surface.albedo module



**class** `climlab.surface.albedo.ConstantAlbedo` (*albedo=0.33, \*\*kwargs*)  
 Bases: `climlab.process.diagnostic.DiagnosticProcess` (page 81)

A class for constant albedo values at all spatial points of the domain.

#### Initialization parameters

**Parameters** `albedo` (*float*) – albedo values

- unit: dimensionless
- default value: 0.33

#### Object attributes

Additional to the parent class `DiagnosticProcess` (page 81) following object attributes are generated and updated during initialization:

**Variables** `albedo` (page 110) (`Field` (page 63)) – attribute to store the albedo value. During initialization the `albedo()` (page 110) setter is called.

**Example** Creation of a constant albedo subprocess on basis of an EBM domain:

```
>>> import climlab
>>> from climlab.surface.albedo import ConstantAlbedo

>>> # model creation
>>> model = climlab.EBM()

>>> sfc = model.domains['Ts']

>>> # subprocess creation
>>> const_alb = ConstantAlbedo(albedo=0.3, domains=sfc, **model.param)
```

Uniform prescribed albedo.

#### `albedo`

Property of albedo value.

**Getter** Returns the albedo value which is stored in diagnostic dict  
`self.diagnostic['albedo']`

#### Setter

- sets albedo which is addressed as `diagnostics['albedo']` to the new value through creating a Field on the basis of domain `self.domain['default']`
- updates the parameter dictionary `self.param['albedo']`

**Type** Field

**class** `climlab.surface.albedo.Iceline` ( $T_f=-10.0$ , *\*\*kwargs*)

Bases: `climlab.process.diagnostic.DiagnosticProcess` (page 81)

A class for an Iceline subprocess.

Depending on a freezing temperature it calculates where on the domain the surface is covered with ice, where there is no ice and on which latitude the ice-edge is placed.

#### Initialization parameters

**Parameters**  $T_f$  (*float*) – freezing temperature where sea water freezes and surface is covered with ice

- unit: °C
- default value: -10

#### Object attributes

Additional to the parent class `DiagnosticProcess` (page 81) following object attributes are generated and updated during initialization:

##### Variables

- **param** (*dict*) – The parameter dictionary is updated with the input argument ' $T_f$ '.
- **diagnostics** (*dict*) – key '`icelat`' initialized
- **icelat** (*array*) – the subprocess attribute `self.icelat` is created

#### `find_icelines()`

Finds iceline according to the surface temperature.

This method is called by the private function `_compute()` and updates following attributes according to the freezing temperature `self.param['Tf']` and the surface temperature `self.param['Ts']`:

#### Object attributes

##### Variables

- **noice** (*Field* (page 63)) – a Field of booleans which are `True` where  $T_s \geq T_f$
- **ice** (*Field* (page 63)) – a Field of booleans which are `True` where  $T_s < T_f$
- **icelat** (*array*) – an array with two elements indicating the ice-edge latitudes
- **diagnostics** (*dict*) – key '`icelat`' is updated according to object attribute `self.icelat` during modification

**class** `climlab.surface.albedo.P2Albedo` ( $a_0=0.33$ ,  $a_2=0.25$ , *\*\*kwargs*)

Bases: `climlab.process.diagnostic.DiagnosticProcess` (page 81)

A class for parabolic distributed albedo values across the domain on basis of the second order Legendre Polynomial.

Calculates the latitude dependent albedo values as

$$\alpha(\varphi) = a_0 + a_2 P_2(x)$$

where  $P_2(x) = \frac{1}{2}(3x^2 - 1)$  is the second order Legendre Polynomial and  $x = \sin(\varphi)$ .

## Initialization parameters

### Parameters

- **a0** (*float*) – basic parameter for albedo function
  - unit: dimensionless
  - default value: 0.33
- **a2** (*float*) – factor for second legendre polynomial term in albedo function
  - unit: dimensionless
  - default value: 0.25

### Object attributes

Additional to the parent class *DiagnosticProcess* (page 81) following object attributes are generated and updated during initialization:

### Variables

- **a0** (page 112) (*float*) – attribute to store the albedo parameter a0. During initialization the *a0()* (page 112) setter is called.
- **a2** (page 112) (*float*) – attribute to store the albedo parameter a2. During initialization the *a2()* (page 112) setter is called.
- **diagnostics** (*dict*) – key 'albedo' initialized
- **albedo** (page 110) (*Field* (page 63)) – the subprocess attribute *self.albedo* is created with correct dimensions (according to *self.lat*)

**Example** Creation of a parabolic albedo subprocess on basis of an EBM domain:

```
>>> import climlab
>>> from climlab.surface.albedo import P2Albedo

>>> # model creation
>>> model = climlab.EBM()

>>> # modify a0 and a2 values in model parameter dictionary
>>> model.param['a0']=0.35
>>> model.param['a2']= 0.10

>>> # subprocess creation
>>> p2_alb = P2Albedo(domains=model.domains['Ts'], **model.param)

>>> p2_alb.a0
0.33
>>> p2_alb.a2
0.1
```

### a0

Property of albedo parameter a0.

**Getter** Returns the albedo parameter value which is stored in attribute *self.\_a0*

### Setter

- sets albedo parameter which is addressed as *self.\_a0* to the new value
- updates the parameter dictionary *self.param['a0']*
- calls method *\_compute\_fixed()*

**Type** float

### a2

Property of albedo parameter a2.

**Getter** Returns the albedo parameter value which is stored in attribute `self._a2`

**Setter**

- sets albedo parameter which is addressed as `self._a2` to the new value
- updates the parameter dictionary `self.param['a2']`
- calls method `_compute_fixed()`

**Type** float

**class** `climlab.surface.albedo.StepFunctionAlbedo` (*Tf=-10.0, a0=0.3, a2=0.078, ai=0.62, \*\*kwargs*)

Bases: `climlab.process.diagnostic.DiagnosticProcess` (page 81)

A step function albedo subprocess.

This class itself defines three subprocesses that are created during initialization:

- 'iceline' - [Iceline](#) (page 111)
- 'warm\_albedo' - [P2Albedo](#) (page 111)
- 'cold\_albedo' - [ConstantAlbedo](#) (page 110)

**Initialization parameters**

**Parameters**

- **Tf** (*float*) – freezing temperature for Iceline subprocess
  - unit: °C
  - default value: -10
- **a0** (*float*) – basic parameter for P2Albedo subprocess
  - unit: dimensionless
  - default value: 0.3
- **a2** (*float*) – factor for second legendre polynomial term in P2Albedo subprocess
  - unit: dimensionless
  - default value: 0.078
- **ai** (*float*) – ice albedo value for ConstantAlbedo subprocess
  - unit: dimensionless
  - default value: 0.62

Additional to the parent class [DiagnosticProcess](#) (page 81) following object attributes are generated/updated during initialization:

**Variables**

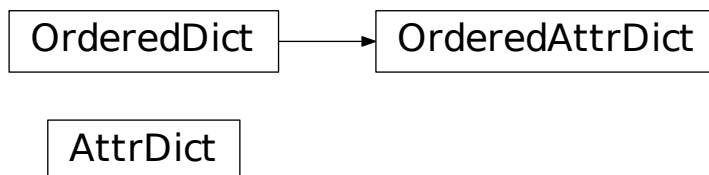
- **param** (*dict*) – The parameter dictionary is updated with a couple of the initialization input arguments, namely 'Tf', 'a0', 'a2' and 'ai'.
- **topdown** (*bool*) – is set to `False` to call subprocess compute method first
- **diagnostics** (*dict*) – key 'albedo' initialized
- **albedo** (page 110) ([Field](#) (page 63)) – the subprocess attribute `self.albedo` is created

**Example** Creation of a step albedo subprocess on basis of an EBM domain:

```
>>> import climlab
>>> from climlab.surface.albedo import StepFunctionAlbedo
>>>
>>> model = climlab.EBM(a0=0.29, a2=0.1, ai=0.65, Tf=-2)
>>>
>>> step_alb = StepFunctionAlbedo(state= model.state, **model.param)
>>>
>>> print step_alb
climlab Process of type <class 'climlab.surface.albedo.StepFunctionAlbedo'>.
State variables and domain shapes:
  Ts: (90, 1)
The subprocess tree:
top: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
  iceline: <class 'climlab.surface.albedo.Iceline'>
  cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
  warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
```

## 6.1.8 climlab.utils package

### climlab.utils.attr\_dict module



```
class climlab.utils.attr_dict.AttrDict (*args, **kwargs)
```

Bases: dict

Constructs a dict object with attribute access to data.

```
class climlab.utils.attr_dict.OrderedAttrDict (*args, **kwargs)
```

Bases: collections.OrderedDict

Constructs an OrderedDict object with attribute access to data.

### climlab.utils.constants module

Contains a collection of physical constants for the atmosphere and ocean.

```
import numpy as np

a = 6.373E6          # Radius of Earth (m)
Lhvap = 2.5E6        # Latent heat of vaporization (J / kg)
Lhsub = 2.834E6      # Latent heat of sublimation (J / kg)
Lhfus = Lhsub - Lhvap # Latent heat of fusion (J / kg)
cp = 1004.           # specific heat at constant pressure for dry air (J / kg / K)
Rd = 287.            # gas constant for dry air (J / kg / K)
kappa = Rd / cp
Rv = 461.5           # gas constant for water vapor (J / kg / K)
cpv = 1875.          # specific heat at constant pressure for water vapor (J / kg / K)
Omega = 2 * np.math.pi / 24. / 3600. # Earth's rotation rate, (s*(-1))
```

```

g = 9.8          # gravitational acceleration (m / s**2)
kBoltzmann = 1.3806488E-23 # the Boltzmann constant (J / K)
c_light = 2.99792458E8    # speed of light (m/s)
hPlanck = 6.62606957E-34 # Planck's constant (J s)
# sigma = 5.67E-8 # Stefan-Boltzmann constant (W / m**2 / K**4)
# sigma derived from fundamental constants
sigma = (2*np.pi**5 * kBoltzmann**4) / (15 * c_light**2 * hPlanck**3)

S0 = 1365.2      # solar constant (W / m**2)
# value is consistent with Trenberth and Fasullo, Surveys of Geophysics 2012

ps = 1000.       # approximate surface pressure (mb or hPa)

rho_w = 1000.    # density of water (kg / m**3)
cw = 4181.3     # specific heat of liquid water (J / kg / K)

tempCtoK = 273.15 # 0degC in Kelvin
tempKtoC = -tempCtoK # 0 K in degC
mb_to_Pa = 100.   # conversion factor from mb to Pa

# Some useful time conversion factors
seconds_per_minute = 60.
minutes_per_hour = 60.
hours_per_day = 24.

# the length of the "tropical year" -- time between vernal equinoxes
# This value is consistent with Berger (1978)
# "Long-Term Variations of Daily Insolation and Quaternary Climatic Changes"
days_per_year = 365.2422
seconds_per_hour = minutes_per_hour * seconds_per_minute
minutes_per_day = hours_per_day * minutes_per_hour
seconds_per_day = hours_per_day * seconds_per_hour
seconds_per_year = seconds_per_day * days_per_year
minutes_per_year = seconds_per_year / seconds_per_minute
hours_per_year = seconds_per_year / seconds_per_hour
# average lengths of months based on dividing the year into 12 equal parts
months_per_year = 12.
seconds_per_month = seconds_per_year / months_per_year
minutes_per_month = minutes_per_year / months_per_year
hours_per_month = hours_per_year / months_per_year
days_per_month = days_per_year / months_per_year

area_earth = 4 * np.math.pi * a**2

# present-day orbital parameters, in the same format generated by orbital.py
orb_present = {'ecc': 0.017236, 'long_peri': 281.37, 'obliquity': 23.446}

```

## climlab.utils.heat\_capacity module

Routines for calculating heat capacities for grid boxes.

`climlab.utils.heat_capacity.atmosphere(dp)`

Returns heat capacity of a unit area of atmosphere, in units J/m\*\*2 / K.

$$C_a = \frac{c_p \cdot dp \cdot f_{\text{mb-to-Pa}}}{g}$$

where

variable	value	unit	description
$C_a$	<i>output</i>	J/m <sup>2</sup> /K	heat capacity for atmospheric cell
$c_p$	1004.	J/kg/K	specific heat at constant pressure for dry air
$dp$	<i>input</i>	mb	pressure for atmospheric cell
$f_{\text{mb-to-Pa}}$	100	Pa/mb	conversion factor from mb to Pa
$g$	9.8	m/s <sup>2</sup>	gravitational acceleration

**Function-call argument**

**Parameters** `dp` (*array*) – pressure intervals (*unit*: mb)

**Returns** the heat capacity for atmosphere cells corresponding to pressure input (*unit*: J / m\*\*2 / K)

**Return type** *array*

**Example** Calculate atmospheric heat capacity for pressure intervals of 1, 10, 100 mb:

```
>>> from climlab.utils import heat_capacity

>>> pressure_interval = array([1,10,100]) # in mb
>>> heat_capacity.atmosphere(pressure_interval) # in J / m**2 / K
array([ 10244.89795918, 102448.97959184, 1024489.79591837])
```

`climlab.utils.heat_capacity.ocean(dz)`

Returns heat capacity of a unit area of water, in units J / m\*\*2 / K.

$$C_o = \rho_w \cdot c_w \cdot dz$$

where

variable	value	unit	description
$C_o$	<i>output</i>	J/m <sup>2</sup> /K	heat capacity for oceanic cell
$c_w$	4181.3	J/kg/K	specific heat of liquid water
$dz$	<i>input</i>	m	water depth of oceanic cell
$\rho_w$	1000.	kg/m <sup>3</sup>	density of water

**Function-call argument**

**Parameters** `dz` (*array*) – water depth of ocean cells (*unit*: m)

**Returns** the heat capacity for ocean cells corresponding to depth input (*unit*: J / m\*\*2 / K)

**Return type** *array*

**Example** Calculate atmospheric heat capacity for pressure intervals of 1, 10, 100 m:

```
>>> from climlab.utils import heat_capacity

>>> pressure_interval = array([1,10,100]) # in m
>>> heat_capacity.ocean(pressure_interval) # in J / m**2 / K
array([ 4.18130000e+06, 4.18130000e+07, 4.18130000e+08])
```

`climlab.utils.heat_capacity.slab_ocean(water_depth)`

Returns heat capacity of a unit area slab of water, in units of J / m\*\*2 / K.

Takes input argument `water_depth` and calls `ocean()` (page 116)

**Function-call argument**

**Parameters** `float` – water depth of slab ocean (*unit*: m)

**Returns** the heat capacity for slab ocean cell (*unit*: J / m\*\*2 / K)

**Return type** *float*



## climlab.utils.legendre module

Can calculate the first several Legendre polynomials, along with (some of) their first derivatives.

`climlab.utils.legendre.P0(x)`

$$P_0(x) = 1$$

`climlab.utils.legendre.P1(x)`

$$P_1(x) = x$$

`climlab.utils.legendre.P2(x)`

The second Legendre polynomial.

$$P_2(x) = \frac{1}{2}(3x^2 - 1)$$

`climlab.utils.legendre.Pn(x)`

Calculate Legendre polynomials P0 to P28 and returns them in a dictionary Pn.

**Parameters** *x* (*float*) – argument to calculate Legendre polynomials

**Return** Pn dictionary which contains order of Legendre polynomials (from 0 to 28) as keys and the corresponding evaluation of Legendre polynomials as values.

**Return type** dict

`climlab.utils.legendre.Pnprime(x)`

Calculates first derivatives of Legendre polynomials and returns them in a dictionary Pnprime.

**Parameters** *x* (*float*) – argument to calculate first derivate of Legendre polynomials

**Return** Pn dictionary which contains order of Legendre polynomials (from 0 to 4 and even numbers until 14) as keys and the corresponding evaluation of first derivative of Legendre polynomials as values.

**Return type** dict

## climlab.utils.walk module

`climlab.utils.walk.process_tree(top, name='top')`

Creates a string representation of the process tree for process top.

This method uses the `walk_processes()` (page 118) method to create the process tree.

**Parameters**

- **top** (*Process* (page 83)) – top process for which process tree string should be created
- **name** (*str*) – name of top process

**Returns** string representation of the process tree

**Return type** str

**Example**

```
>>> import climlab
>>> from climlab.utils import walk

>>> model = climlab.EBM()
>>> proc_tree_str = walk.process_tree(model, name='model')

>>> print proc_tree_str
```

```
model: <class 'climlab.model.ebm.EBM'>
  diffusion: <class 'climlab.dynamics.diffusion.MeridionalDiffusion'>
  LW: <class 'climlab.radiation.AplusBT.AplusBT'>
  albedo: <class 'climlab.surface.albedo.StepFunctionAlbedo'>
    iceline: <class 'climlab.surface.albedo.Iceline'>
    cold_albedo: <class 'climlab.surface.albedo.ConstantAlbedo'>
    warm_albedo: <class 'climlab.surface.albedo.P2Albedo'>
  insolation: <class 'climlab.radiation.insolation.P2Insolation'>
```

`climlab.utils.walk.walk_processes` (*top*, *topname='top'*, *topdown=True*, *ignore-Flag=False*)

Generator for recursive tree of climlab processes

Starts walking from climlab process *top* and generates a complete list of all processes and sub-processes that are managed from *top* process. *level* indicates the rank of specific process in the process hierarchy:

---

**Note:**

- **level 0: top process**

- **level 1: sub-processes of top process**

- \* **level 2: sub-sub-processes of top process (=subprocesses of level 1 processes)**

---

The method is based on `os.walk()`.

**Parameters**

- **top** (*Process* (page 83)) – top process from where walking should start
- **topname** (*str*) – name of top process
- **topdown** (*bool*) – whether generate *process\_types* in regular or in reverse order. Set to `True` by default.
- **ignoreFlag** (*bool*) – whether *topdown* flag should be ignored or not

**Returns** name (*str*), proc (process), level (*int*)

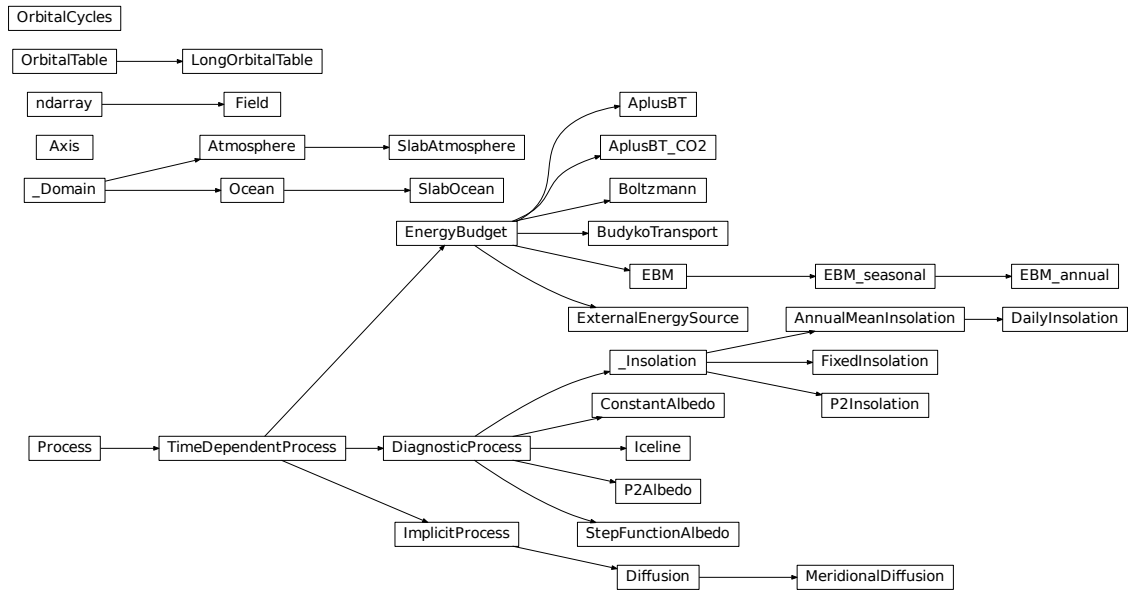
**Example**

```
>>> import climlab
>>> from climlab.utils import walk

>>> model = climlab.EBM()

>>> for name, proc, top_proc in walk.walk_processes(model):
...     print name
...
top
diffusion
LW
iceline
cold_albedo
warm_albedo
albedo
insolation
```

## 6.2 Inheritance Diagram





REFERENCES



## 8.1 climlab

The climlab Python package is licensed under MIT License:

```
The MIT License (MIT)

Copyright (c) 2015 Brian E. J. Rose

Permission is hereby granted, free of charge, to any person obtaining a copy
of this software and associated documentation files (the "Software"), to deal
in the Software without restriction, including without limitation the rights
to use, copy, modify, merge, publish, distribute, sublicense, and/or sell
copies of the Software, and to permit persons to whom the Software is
furnished to do so, subject to the following conditions:

The above copyright notice and this permission notice shall be included in all
copies or substantial portions of the Software.

THE SOFTWARE IS PROVIDED "AS IS", WITHOUT WARRANTY OF ANY KIND, EXPRESS OR
IMPLIED, INCLUDING BUT NOT LIMITED TO THE WARRANTIES OF MERCHANTABILITY,
FITNESS FOR A PARTICULAR PURPOSE AND NONINFRINGEMENT. IN NO EVENT SHALL THE
AUTHORS OR COPYRIGHT HOLDERS BE LIABLE FOR ANY CLAIM, DAMAGES OR OTHER
LIABILITY, WHETHER IN AN ACTION OF CONTRACT, TORT OR OTHERWISE, ARISING FROM,
OUT OF OR IN CONNECTION WITH THE SOFTWARE OR THE USE OR OTHER DEALINGS IN THE
SOFTWARE.
```

## 8.2 Documentation





## CONTACT

### 9.1 climlab package

The climlab package has been developed by Brian Rose:

**Brian E. J. Rose**  
Department of Atmospheric and Environmental Sciences  
University at Albany  
[brose@albany.edu](mailto:brose@albany.edu)

Bug reports can be reported through the [issue tracker](#) on github.

### 9.2 climlab documentation

The documentation has been built by Moritz Kreuzer using [Sphinx](#). Based on some commentary strings in the source code and a couple of Jupyter Notebooks, this documentation has been developed.

Currently it covers only the Energy Balance Model relevant parts of the package.

**Moritz Kreuzer**  
Potsdam Institut for Climate Impact Research (PIK)  
Potsdam, Germany  
[kreuzer@pik-potsdam.de](mailto:kreuzer@pik-potsdam.de)



## INDICES AND TABLES

- `genindex`
- `modindex`
- `search`



## BIBLIOGRAPHY

- [Berger1978] Berger A. 1978. “Long-term variations of daily insolation and Quaternary climatic changes.” *Journal of Atmospheric Science* 35(12):2362-2367.
- [Berger1991] Berger A./Loutre M.F. 1991. “Insolation values for the climate of the last 10 million years.” *Quaternary Science Reviews* 10(4):297-317.
- [Budyko1969] Budyko, M. I. 1969. “The effect of solar radiation variations on the climate of the Earth.” *Tellus* 21(5):611–619.
- [CaldeiraKasting1992] Caldeira, Ken/Kasting, James. 1992. “Susceptibility of the early Earth to irreversible glaciation caused by carbon dioxide clouds.” *Nature* 359:226-228.
- [Laskar2004] Laskar, J./P. Robutel/F. Joutel/M. Gastineau/A. C. M. Correia/B. Levrard. 2004. “A long-term numerical solution for the insolation quantities of the Earth.” *Astronomy & Astrophysics* 428:261–285.