

Lecture II

Climate Archives, Data, and Models

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How do we know what we know about past climate?

Recorders of 'Natural climate experiments': Archives



Historical
records

Coral reefs

Pollen

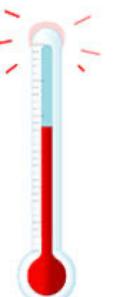
Ice-cores

Tree rings

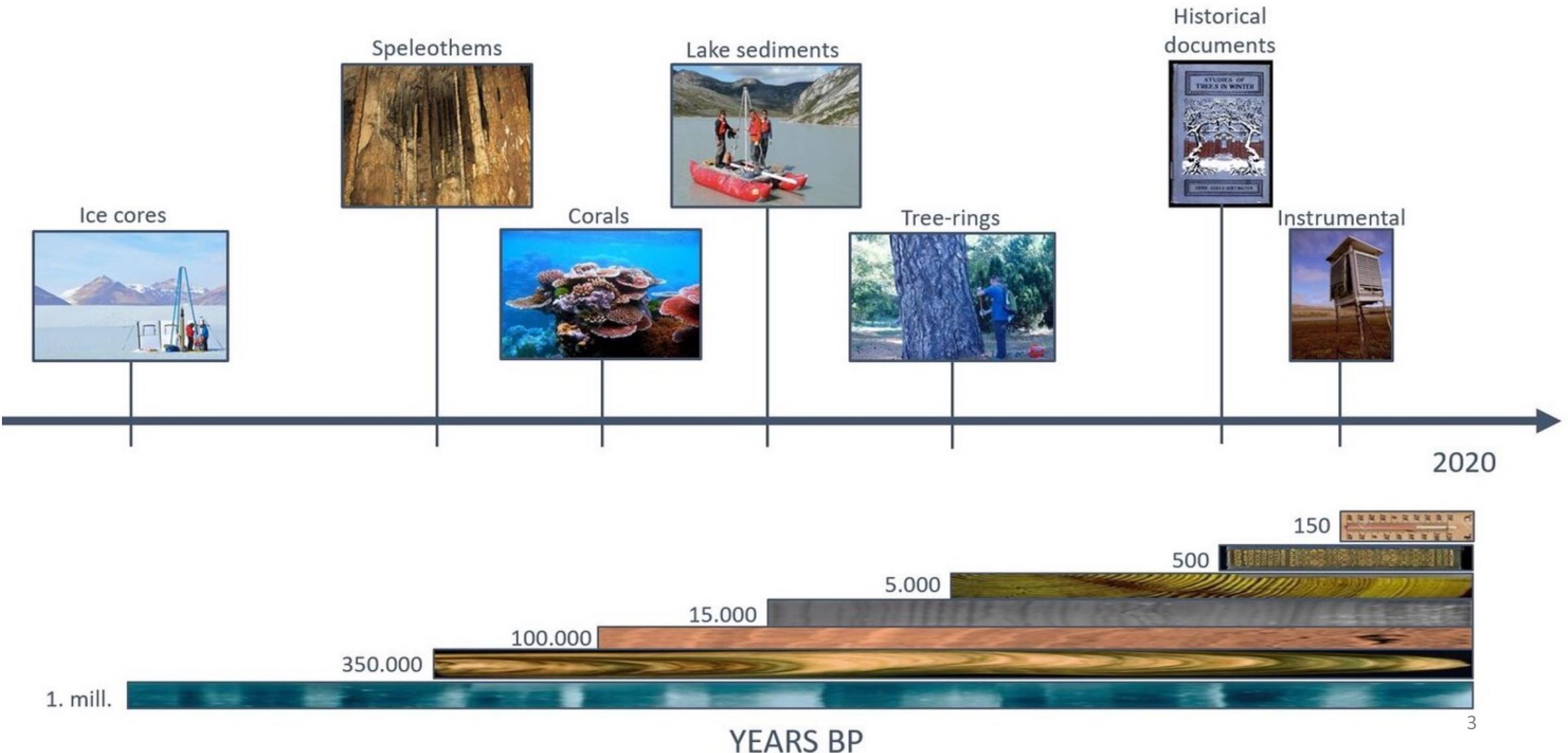
Carbonate
caves

Proxy data are preserved physical characteristics of the environment

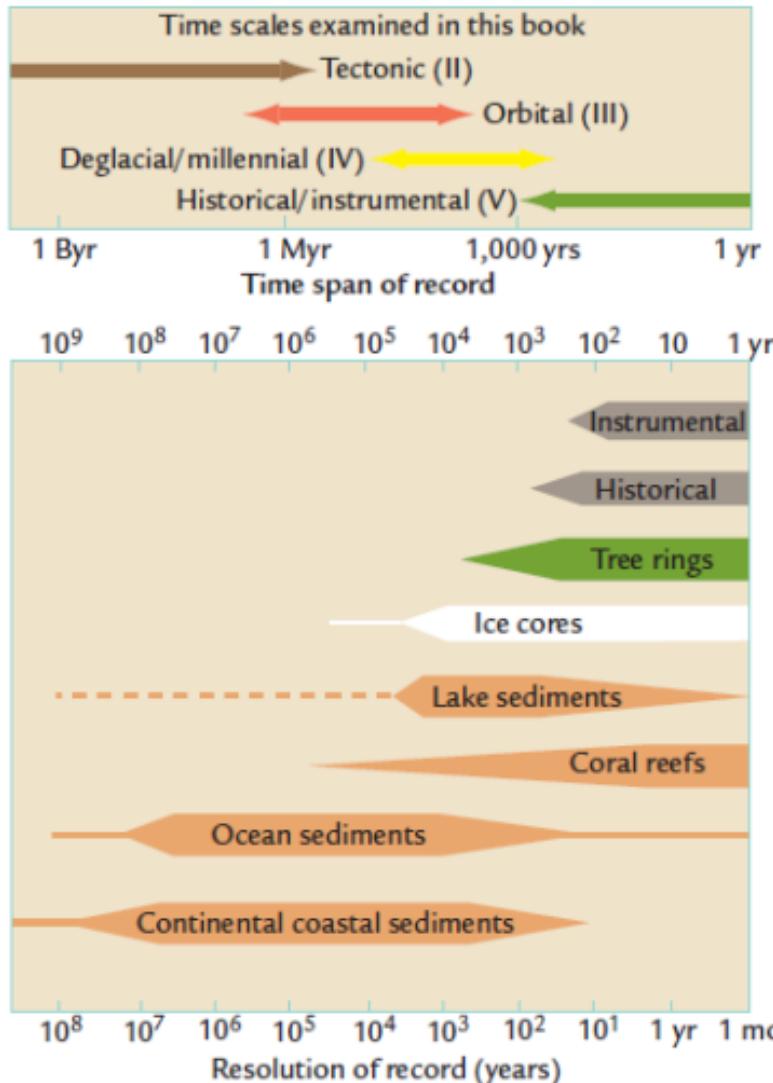
Paleoclimatology : study of past climates



Recorders of ‘Natural climate experiments’: Archives

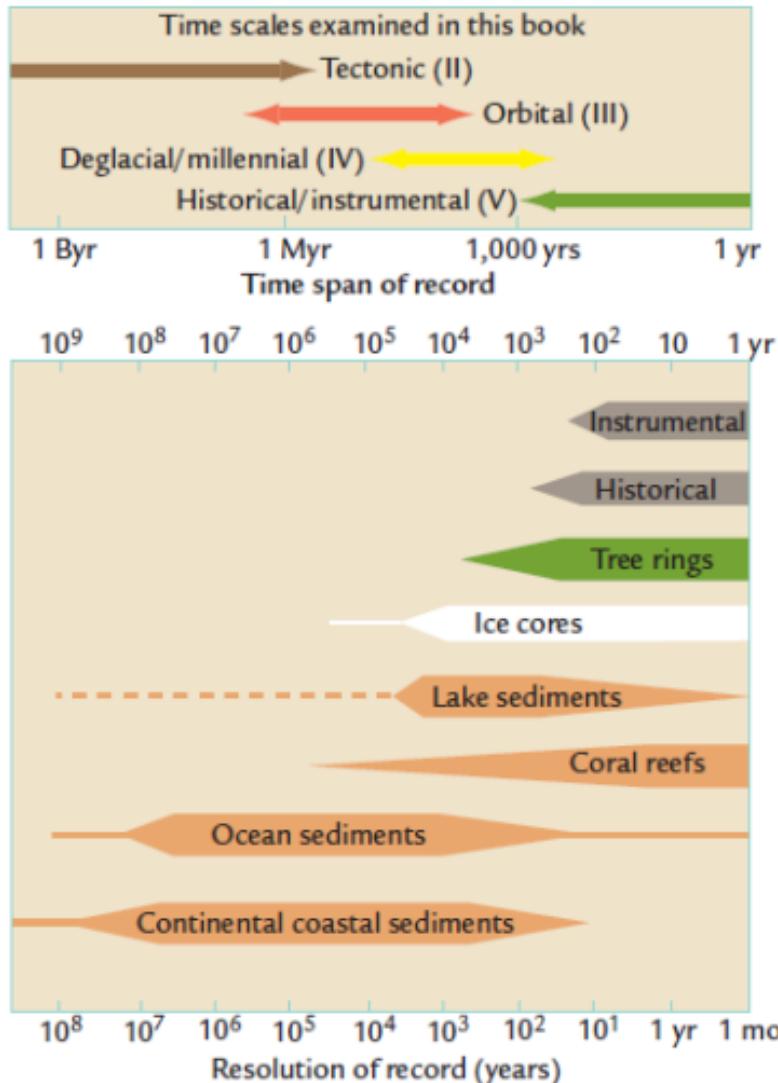


Direct climate information vs proxy records



- Direct measurements have limited timeframe to record the full range pf natural variability in the climate system
- Historical archives are qualitative and incomplete
- Proxy data: natural archives of longer term climate variability

Direct climate information vs proxy records

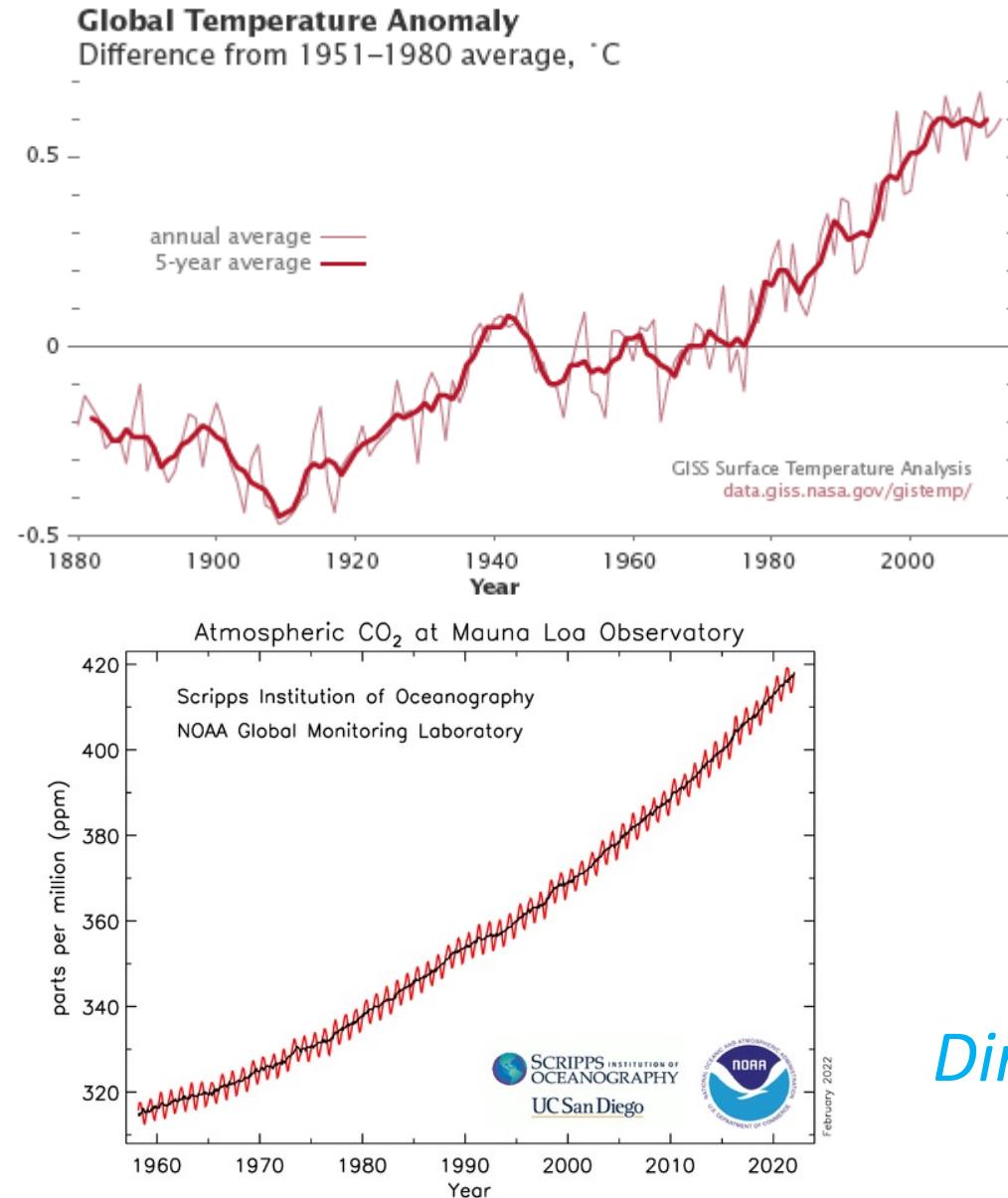


- Sediment archives (major climate archive for over 99% Earth's geological time)
- Ice cores
- Tree rings
- Corals
- Speleothems

Proxies

- Geological-geochemical proxies, which are measurements of mass movements of materials through the climate system, either as discrete (physical) particles or in dissolved (chemical) form.
- Biotic proxies: which are based on changes in composition of plant and animal groups,

Instrumental records



Direct climate information

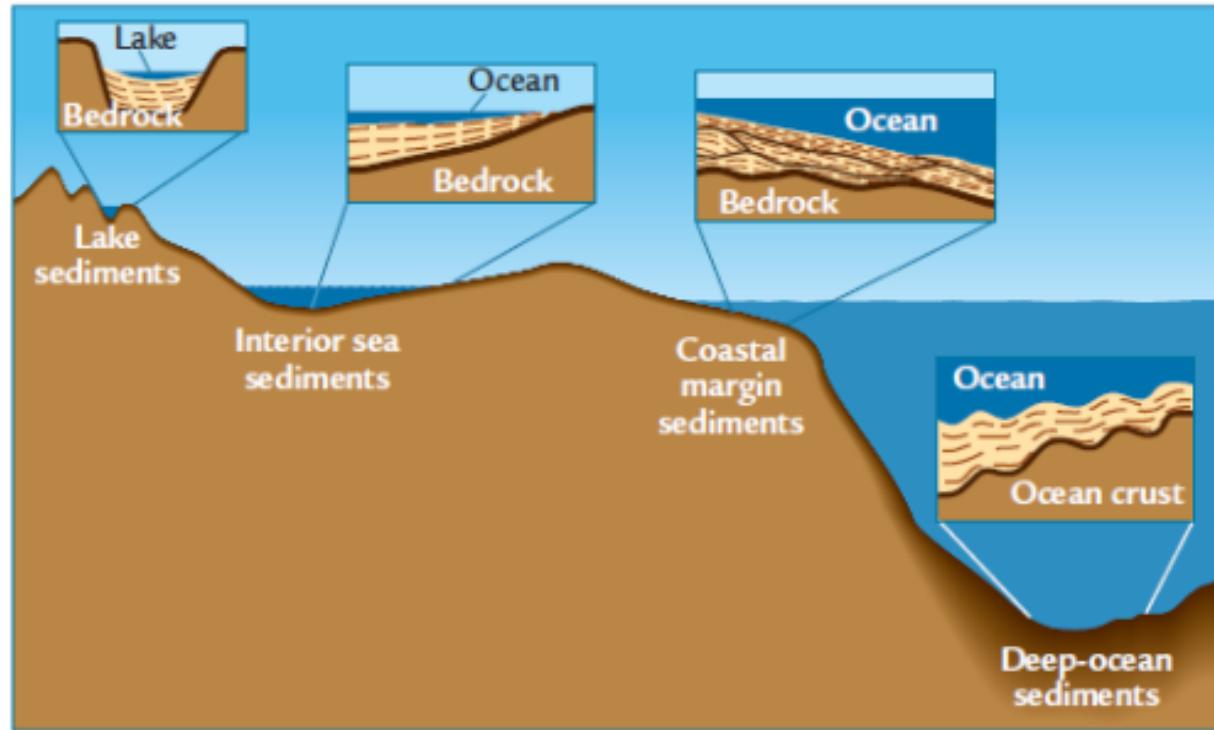
Historical records

- Written accounts
 - El Niño events recorded from late 1500's
 - Crop harvest, migrations, spring blooms
 - Hurricane landfall
- Artwork
 - Snow/ice in temperate locations during Little Ice Age

Direct climate information



I. Sedimentary archives

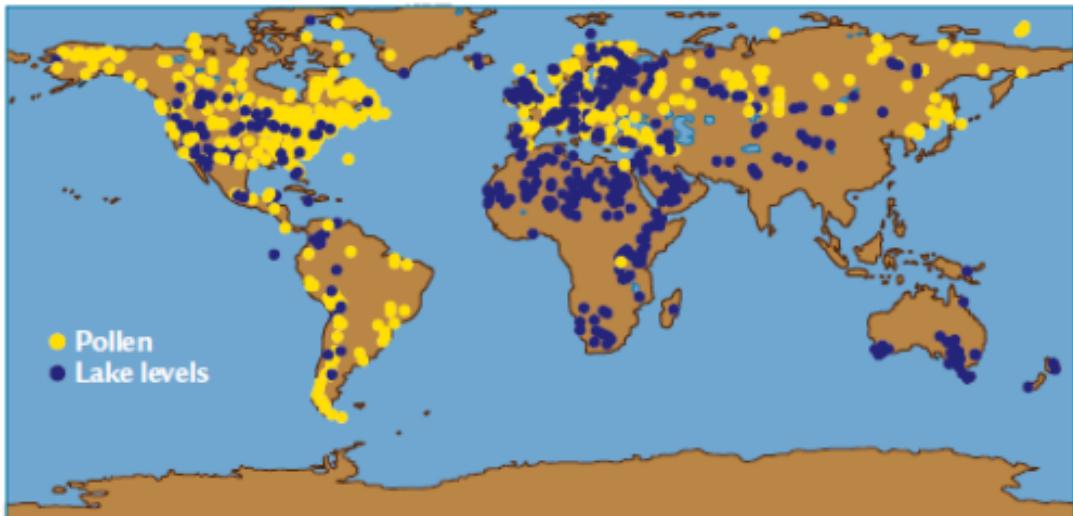


How far back in time can we reconstruct climate from ocean archives?

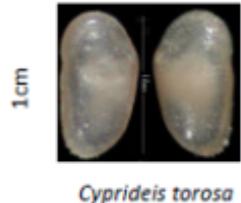
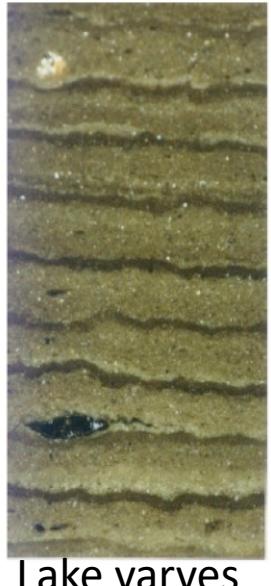
~170 Myr – ocean crust gets destroyed at convergent plate boundaries

- Rainfall and runoff it produces erode/weather the exposed rocks on the continents and transport the debris either physical (granular) and chemical (dissolved) forms.
- Most of the sediments are carried to the oceans, some deposited on quieter waters (eg. lakes), which eventually is carried to the ocean by long intervals of erosion and tectonic processes.

Lake sediments



NPDC, Boulder



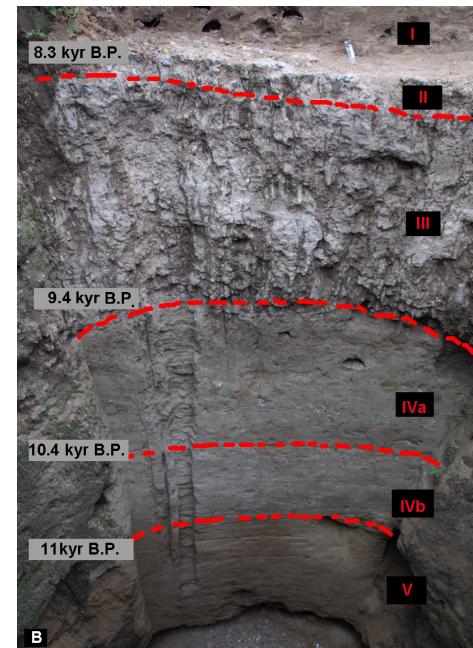
Lake varves

- Record terrestrial climate variations (temperature, precipitation).
- Varved sediments – annual banding caused by seasonal changes in productivity, sediment input
 - Summer – biological, organic rich
 - Winter – clastic runoff
- Ostracod shells chemistry – rainfall and temperature changes (Oxygen isotopes trace element composition)
- Pollen/C isotopes – vegetation changes

Lake and Ocean sedimentary archives



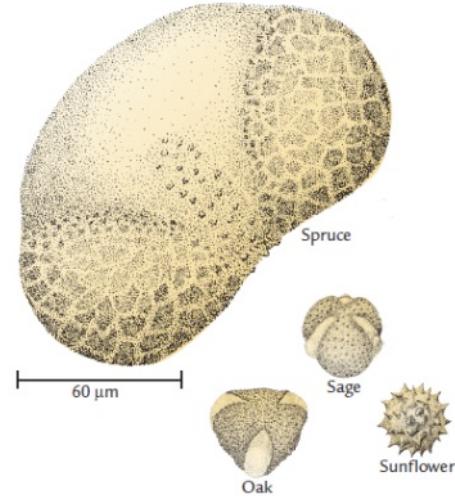
Sediment cores onboard the research ship



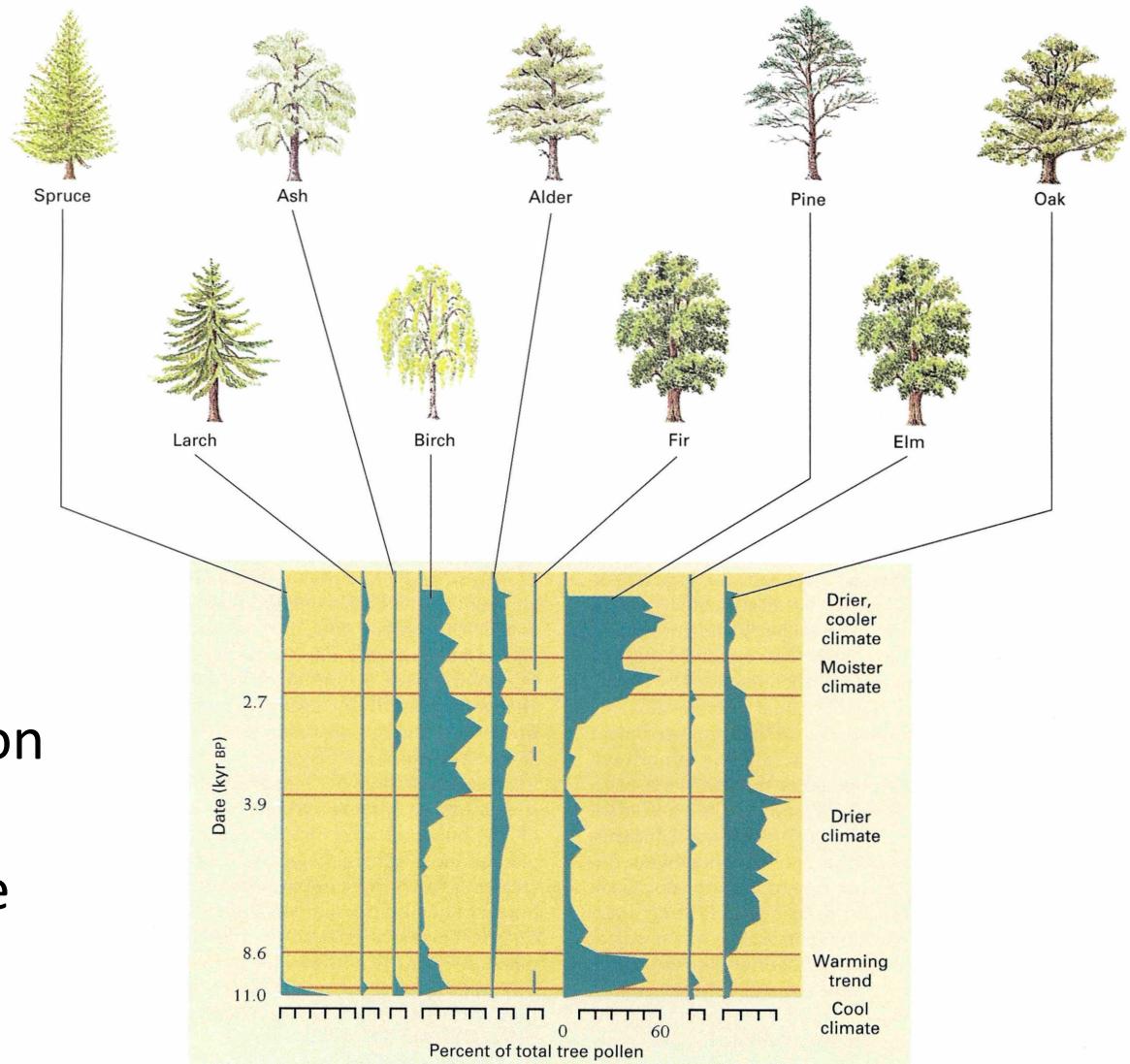
Lake sediment sampling from a cross sectional wall from a dug trench



Pollen analysis in terrestrial sediments



- A proxy indicator of climate on land
- Can be used to reconstruct past climates on **tectonic timescales**
- Presence of a single temperature sensitive form can indicate the climate e.g. Pagan climate from Palm vegetation in northern latitudes

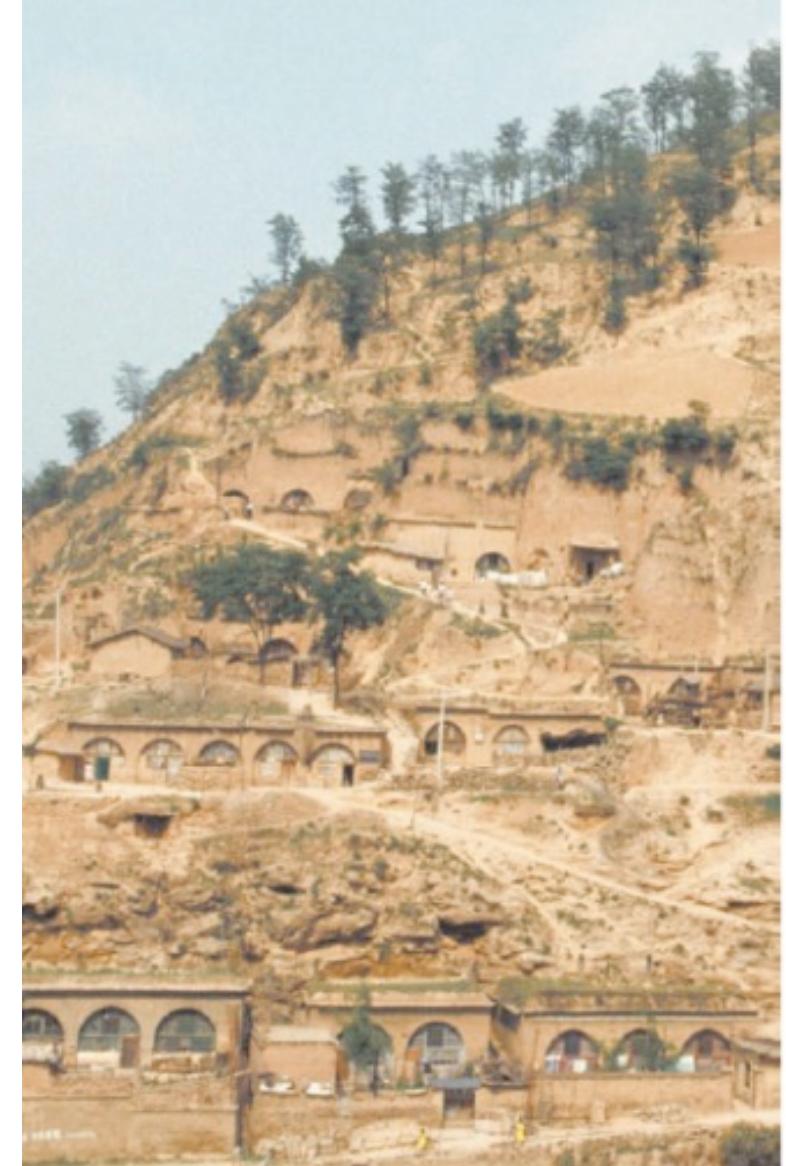


Loess sediments

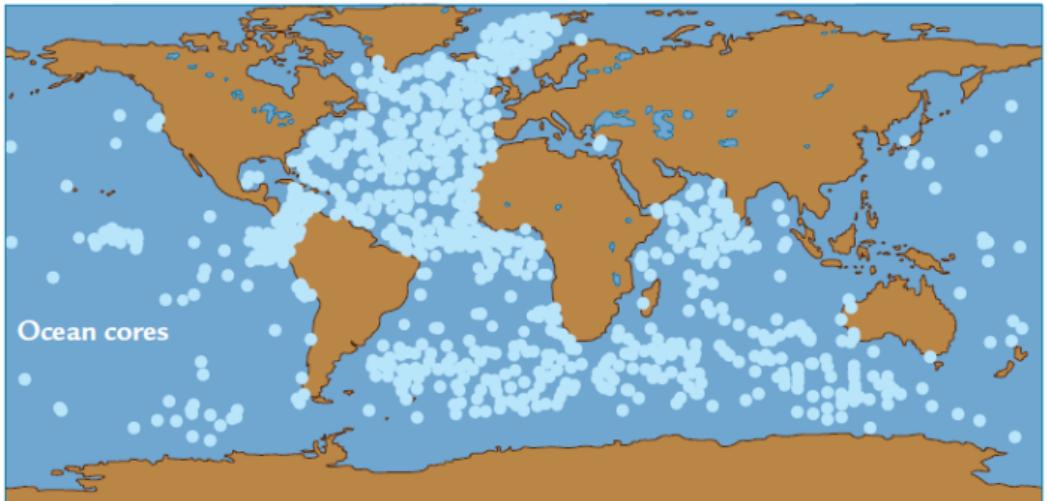
- Coarse (silt-sized) wind blown sediment sequences called loess.
- China has some excellent loess repositories of the last 3 million years.
- Grain size analysis of the entire exposed section of the loess deposition indicator of glacial-interglacial climate changes.

Soil – fine grain size (wetter period)

Loess sediment – coarse , wind blown (dry period)



Ocean archives



NPDC, Boulder

IODP - International Ocean Discovery Program



Geological and geochemical data

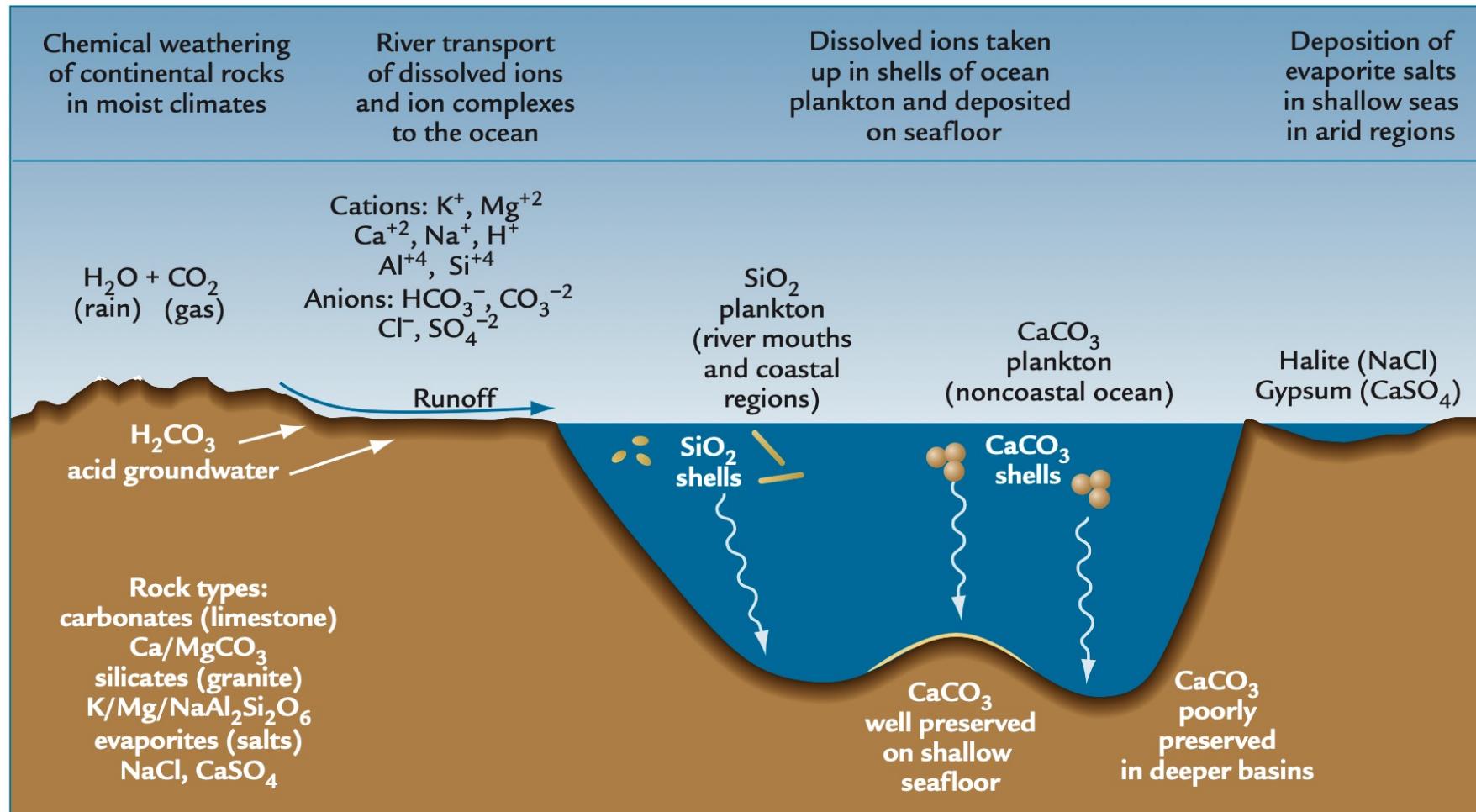
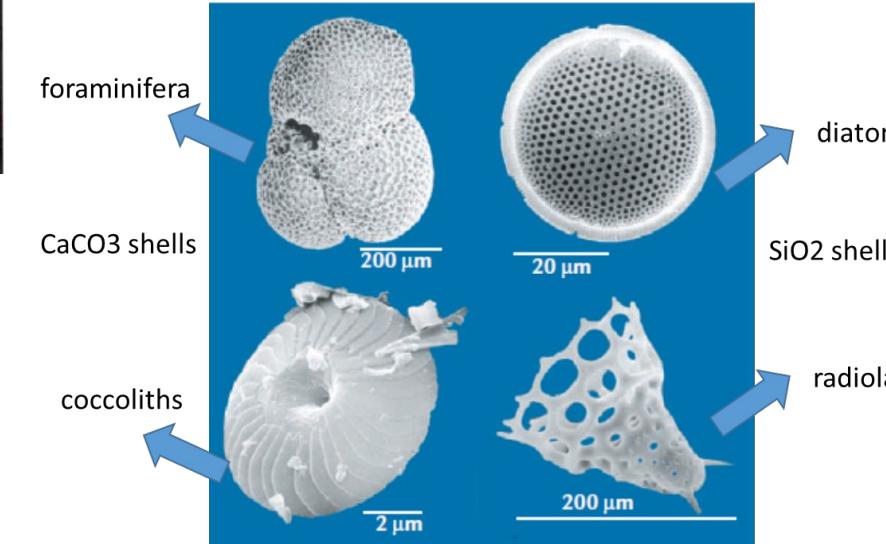
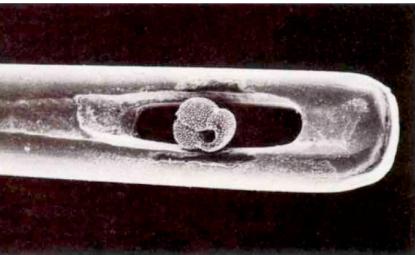
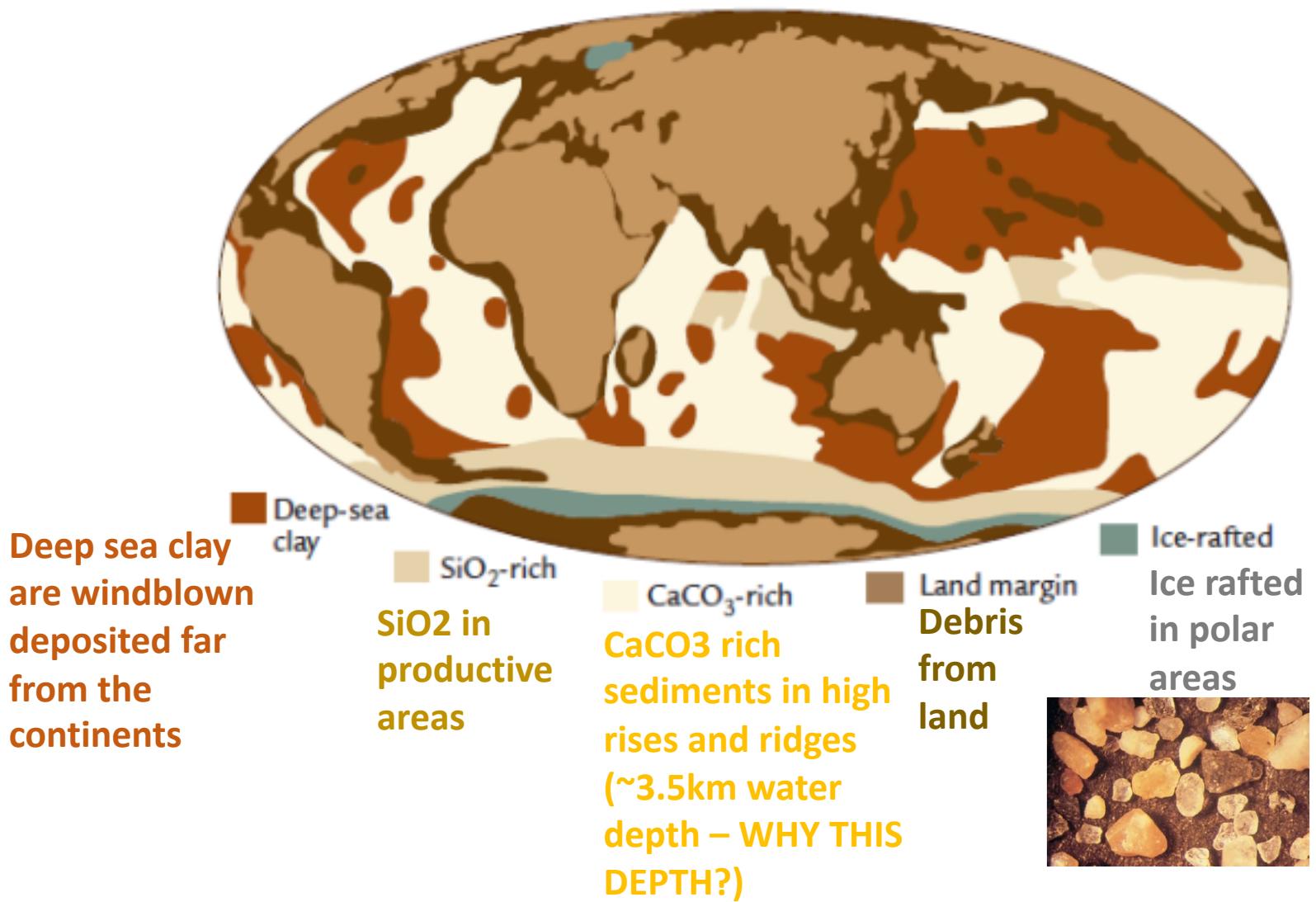


FIGURE 3-17

Chemical weathering, transport, and deposition

Chemical weathering slowly attacks rocks on land and sends dissolved ions into rivers for transport to the ocean. Ocean plankton incorporate some of the dissolved ions in their shells, which fall to the seafloor and form part of the geologic record. Some dissolved ions are also deposited in shallow salty evaporating pools on continental margins where the climate is dry.

Ocean archives



Plankton: a proxy indicator of climate in the ocean

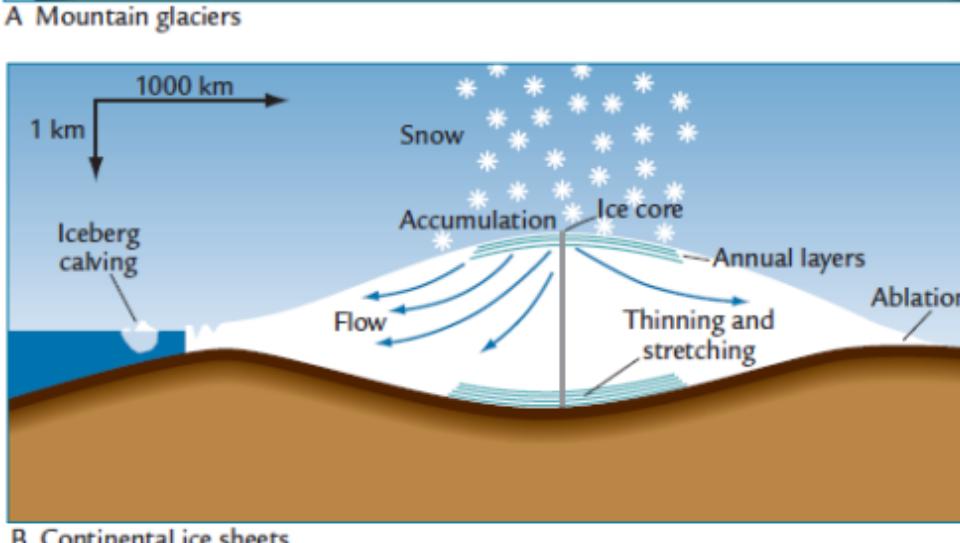
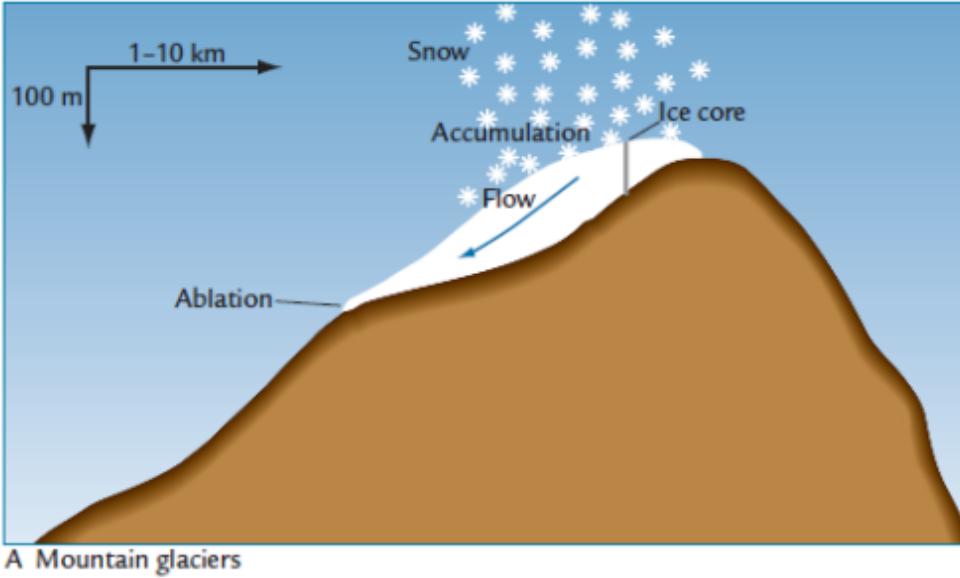
Four types of shelled remains of plankton are common in ocean sediments:

1. CaCO₃ shells are represented by sand-sized planktic foraminifera (upper left) and small clay-sized coccoliths (lower left);
2. SiO₂ shells include silt-sized diatoms (upper right) and sand-sized radiolaria (lower right).

For comparison, small grains of sand are 60 µm or larger in diameter.

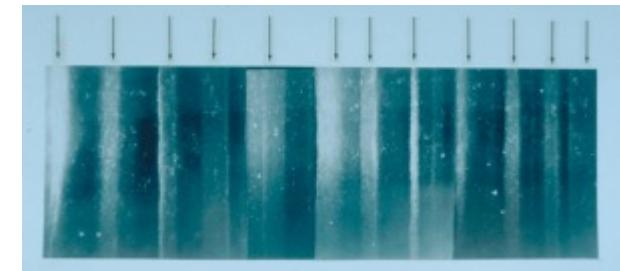


Ice archives



Located on high latitudes and altitudes

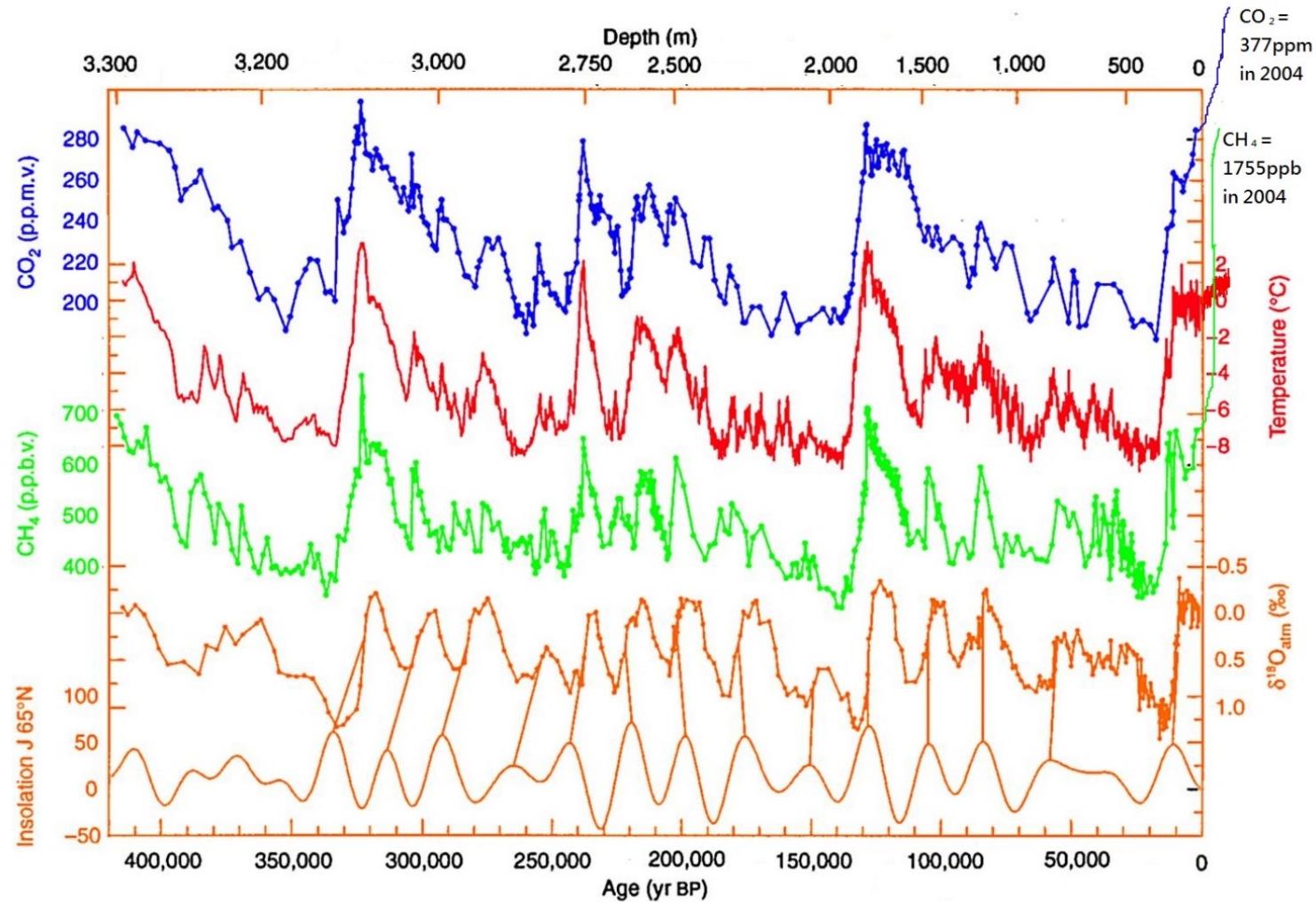
- H₂O in ice – Oxygen isotopes –temp, precipitation
- Dust amount – global dryness, wind
- Air bubbles – Actual samples of trapped air to determine past concentrations of gases such as CO₂, CH₄
- Layers in the ice



Ice archives

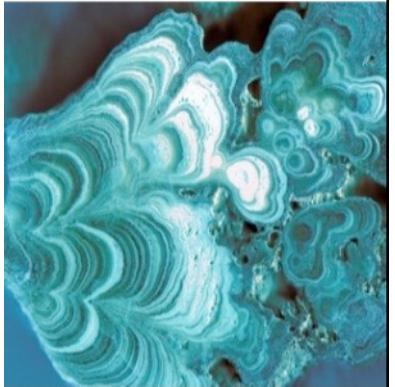
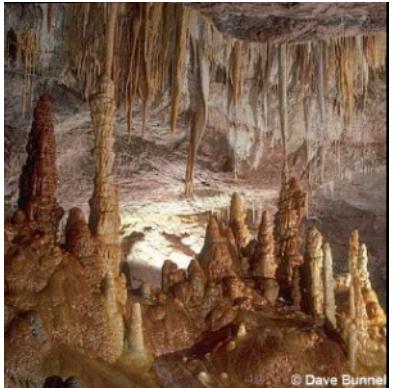
- Ice recovered from the Antarctic ice sheet now dates back to 800,000 years, and ice from the Greenland ice sheet dates to just beyond 125,000 years.

Vostok ice core record in East Antarctica

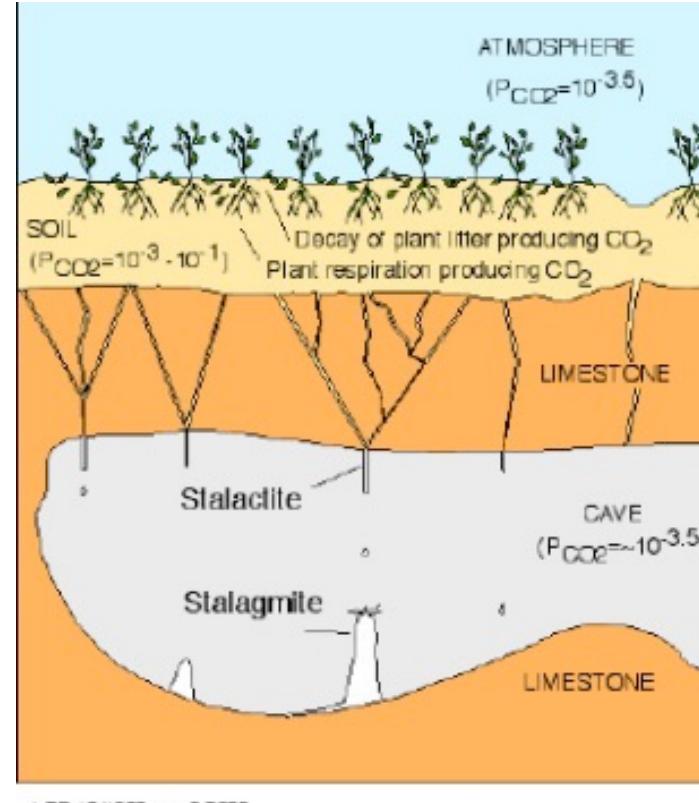


Speleothems

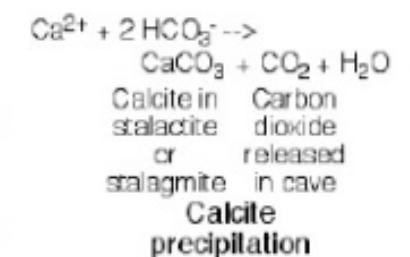
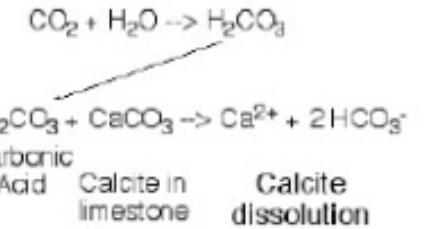
Calcium carbonate cave deposits – Stalagmites and Stalactites



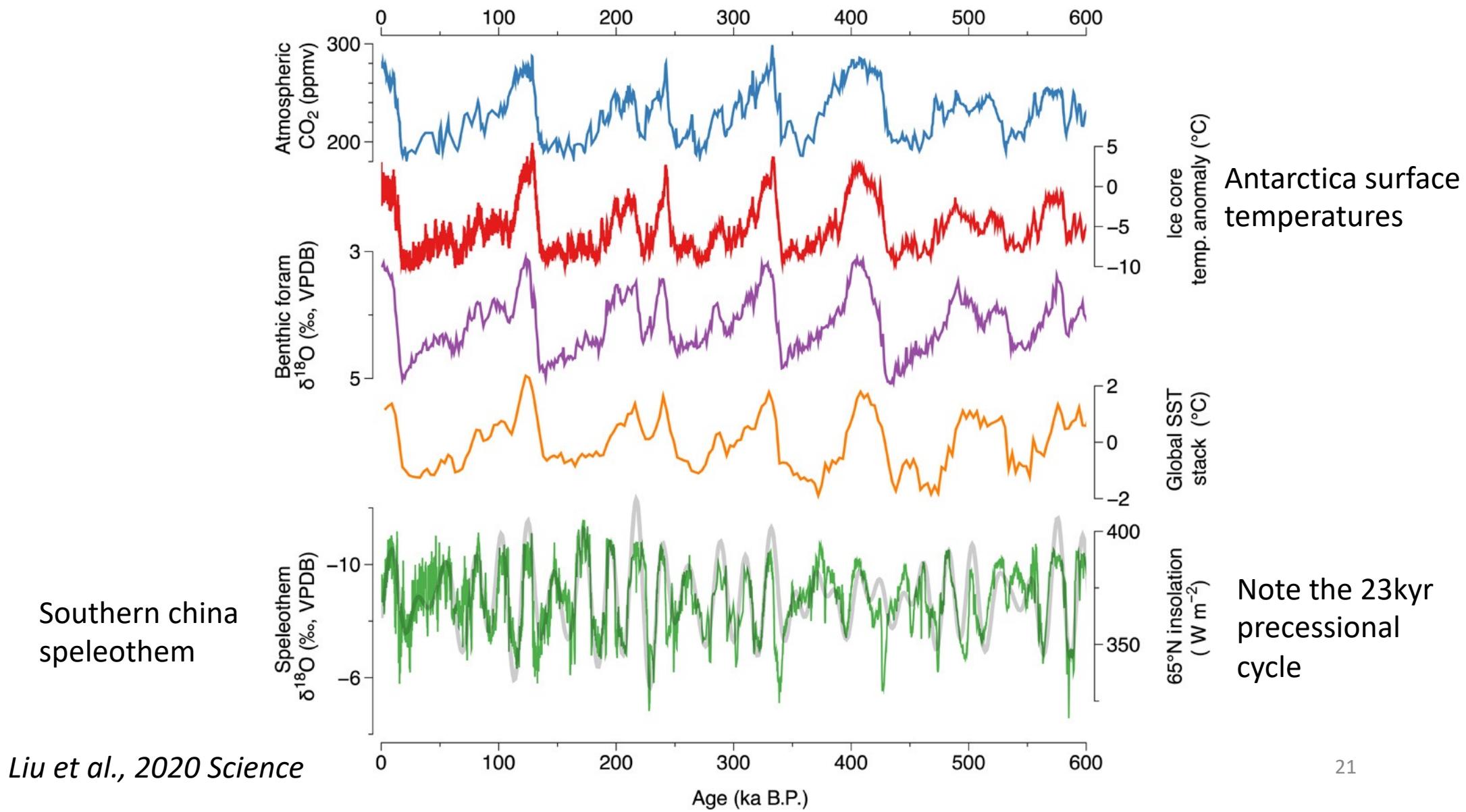
d₁₈O of speleothems
is a reflection of the
groundwater d₁₈O
and ultimately the
d₁₈O of the rainfall in
the region



Release of CO₂ from dripwater in
cave drives precipitation of calcite
to make stalactites and stalagmites.



Chinese speleothem vs global records



Other climate archives

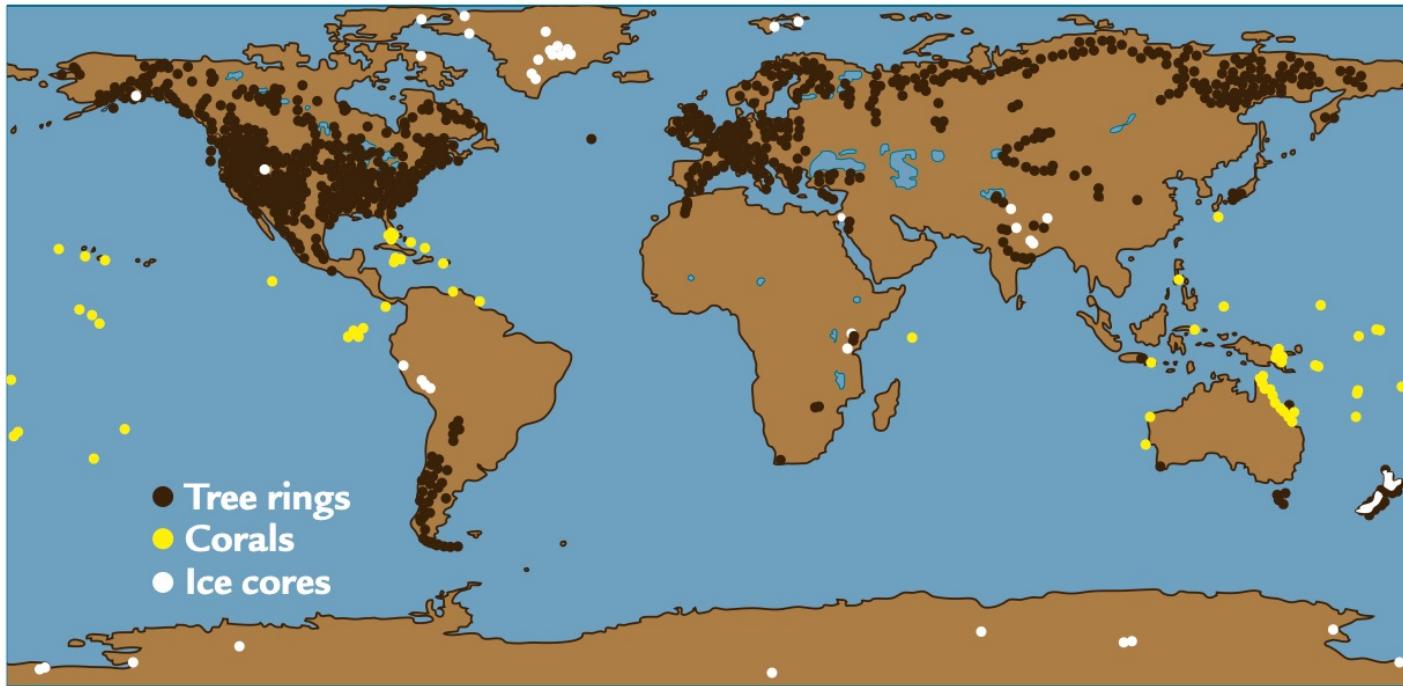
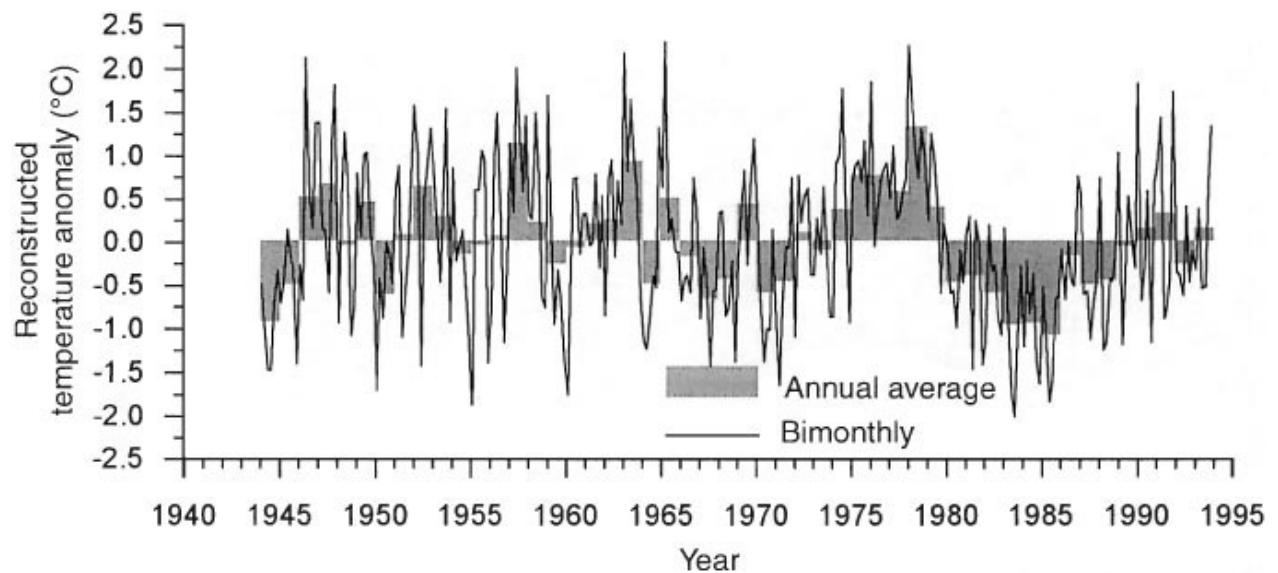
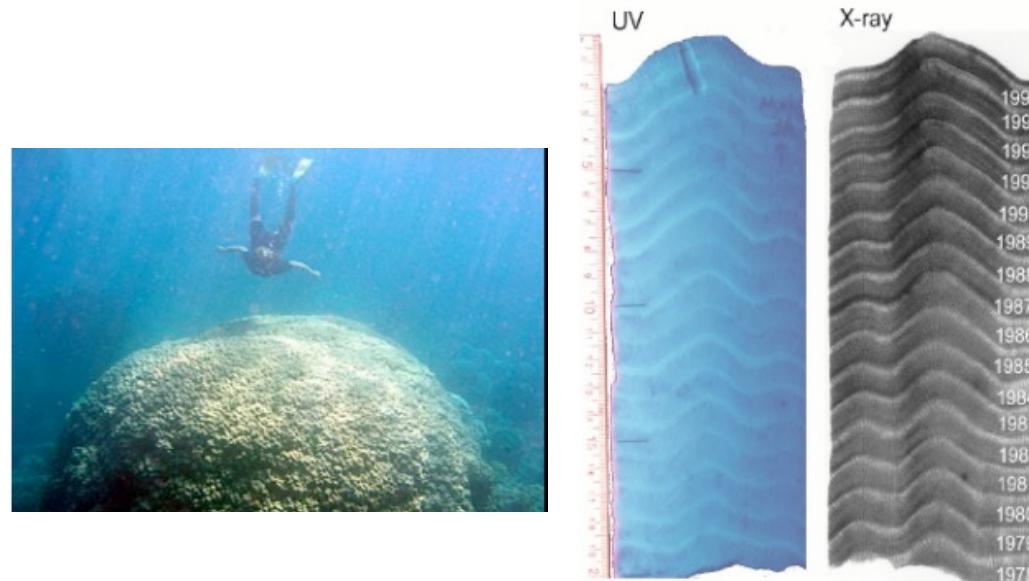


FIGURE 3-6
Ice cores, corals, and tree rings
Ice cores, corals, and tree rings are
archives of climate change in more recent
Earth history. (NATIONAL PALEOCLIMATE DATA
CENTER, NGDC, BOULDER, CO.)

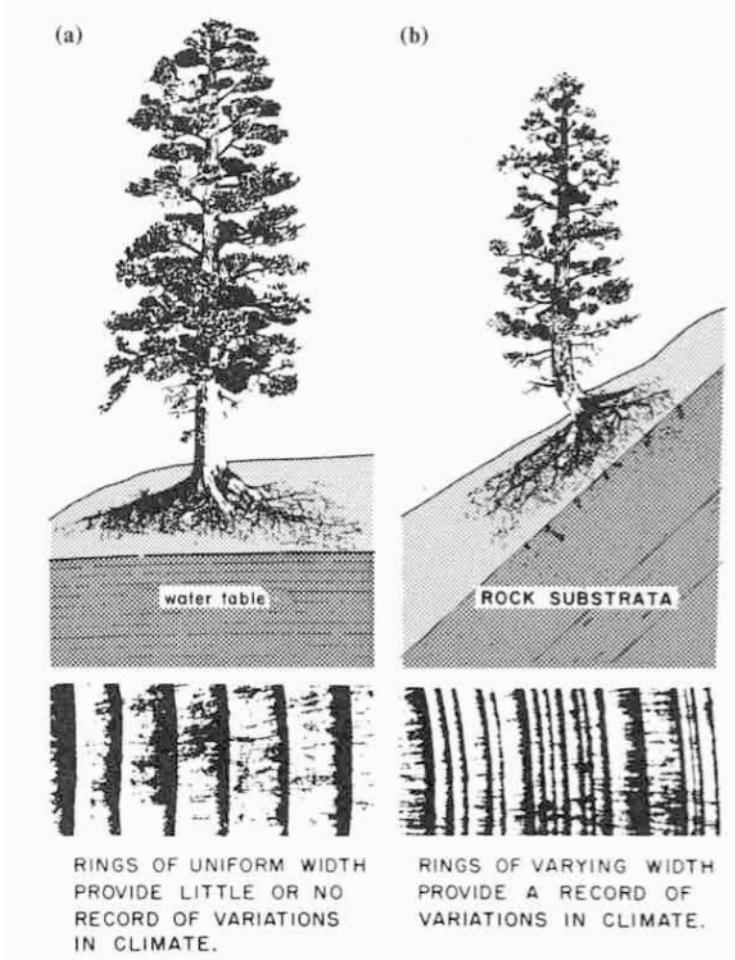
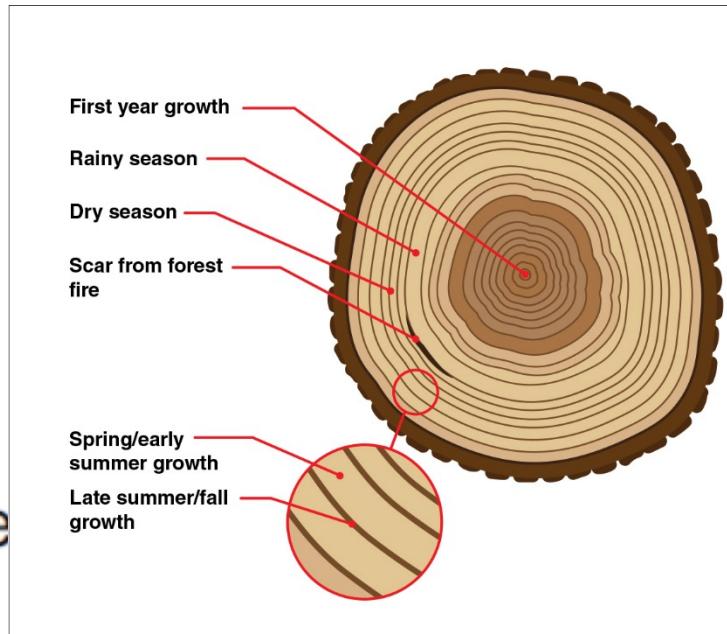
Corals

- Skeletons made of aragonite (CaCO_3)
- Dated using annual density band U-Th, ^{14}C
- Recorders of tropical sea surface conditions:
 - Temperature, Salinity
- Oxygen Isotopes record a combination of temp and salinity
- Strontium/Calcium (Sr/Ca) records mainly temperature

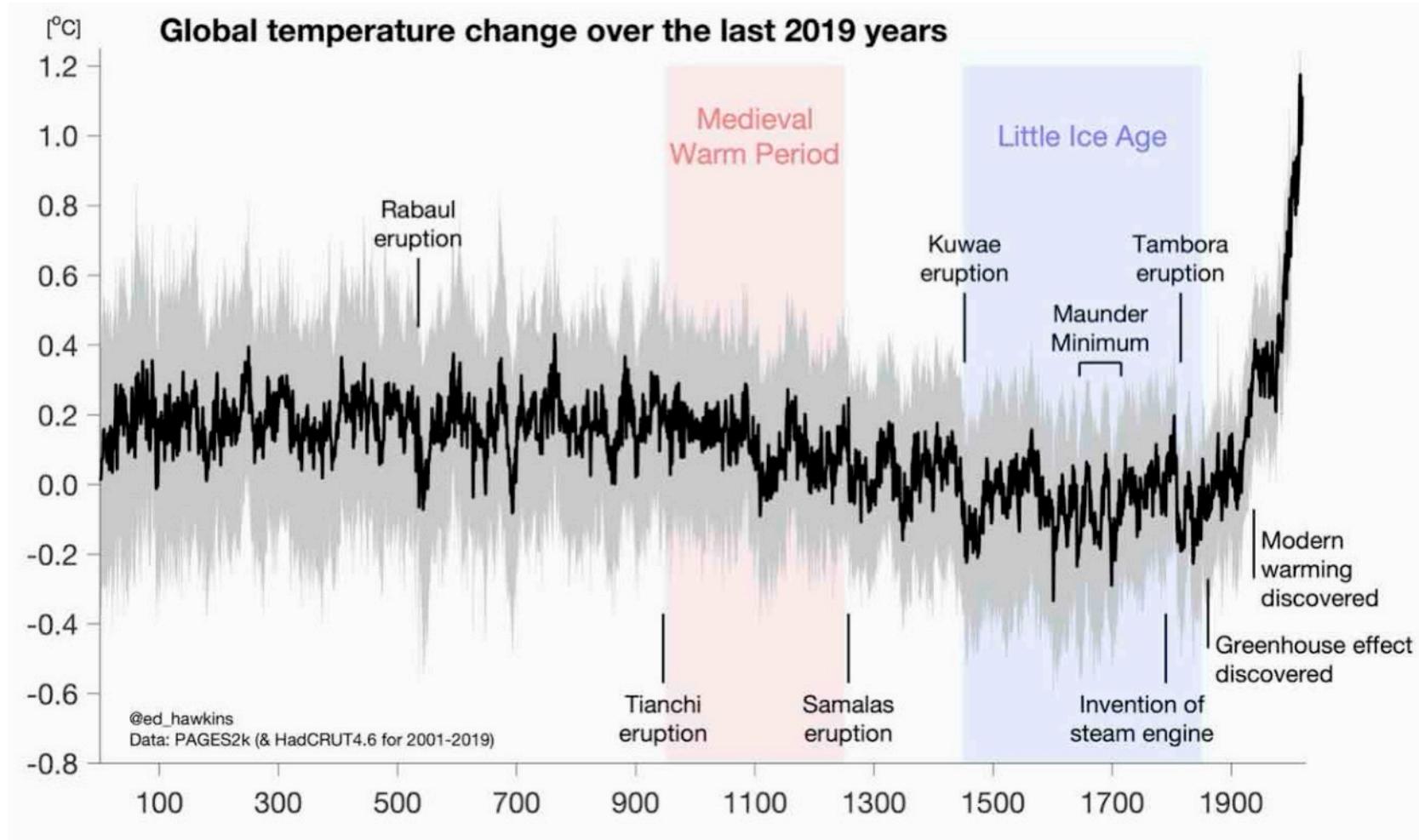


Tree rings

- Predominantly temperate (mid-latitudes)
- Dated using radiocarbon
- Thick bands during growing season, thin bands during cold/dry months
- Varying widths of growth bands reflect temperature or precipitation



Global temperature change over two millennia



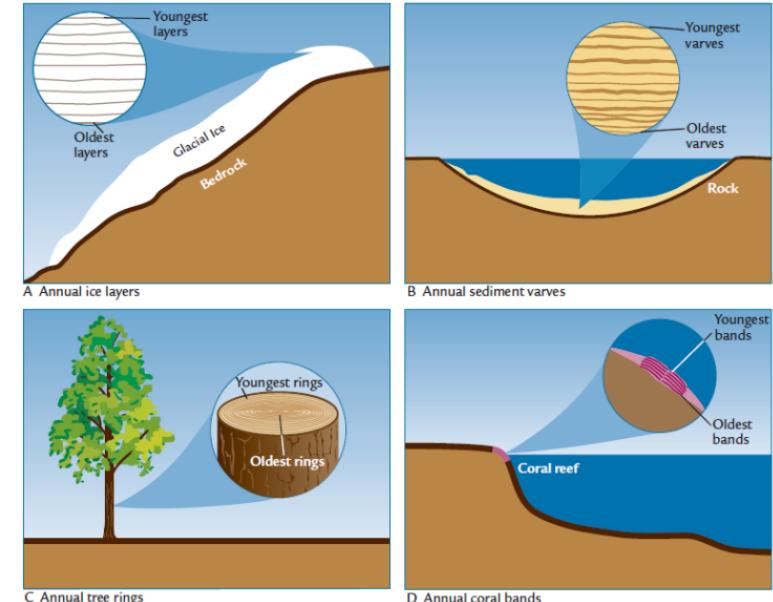
How do we date these proxy records?

Counting annual layers

Table 3-1 Radioactive Decay Used to Date Climate Records

Parent isotope	Daughter isotope	Half-life	Useful for ages:	Useful for dating:
Rubidium-87 (^{87}Rb)	Strontium-87 (^{87}Sr)	47 Byr	100 Myr	Granites
Uranium-238 (^{238}U)	Lead-206 (^{206}Pb)	4.5 Byr	>100 Myr	Many rocks
Uranium-235 (^{235}U)	Lead-207 (^{207}Pb)	0.7 Byr	>100 Myr	Many rocks
Potassium-40 (^{40}K)	Argon-40 (^{40}Ar)	1.3 Byr	>100,000 years	Basalts
Thorium 230 (^{230}Th)	Radon-226* (^{226}Ra)	75,000 years	<400,000 years	Corals
Carbon-14 (^{14}C)	Nitrogen-14* (^{14}N)	5,780 years	<50,000 years	Anything that contains carbon

*Daughter in this case is a gas that has escaped and cannot be measured.



Correlating records with orbital cycles – ‘tuning’

- Timing is accurately known from astronomical calculations and physical processes like cyclical growth and decay of high-latitude ice-sheets and monsoon variations in low-latitudes
- Absolute dating covering many Myr

Confluence of disciplines – Paleoclimatology

