Algorithm for Calculation of Earth Atmosphere Mass.

According to the I – law of thermodynamics heat change dq depends on the system's internal energy change du and object work change dw, thus total heat energy is dq = du + dw. For ideal gasses, I – law of thermodynamics heat change dq equation is:

$$\frac{dw}{pdv} + vdp = RdT, [J/kg]$$
(1.0)

Here: $R = 287 \ J/(kg \times K)$ – universal gas constant, p – gas pressure, [Pa], $v=1/\rho$ – gas specific volume, $[m^3/kg]$, ρ – gas density, $[kg/m^3]$.

Isothermal -dq=0, dT=0. - slow gas expansion - compression then afterward gas temperature stays unchanged. Earth pressure density change from height is based on **isothermal** gas expansion from height change. The Atmosphere pressure and density are similar units: $\rho(h)\sim\rho(h)$ based on the ideal gas equation condition, for more details see (1.05) equation.

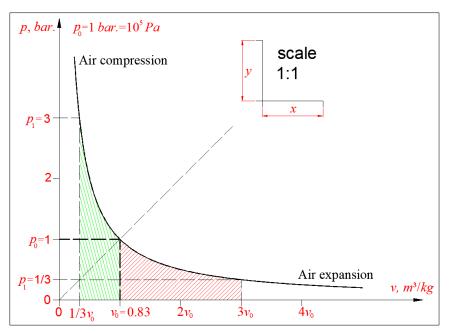


Fig. 1 Isothermal process p=const./v of air compression – expansion.

Equation (1.04) is a square (isosceles) hyperbola math function and it defines a thermodynamic isothermal process of air compression – expansion.

The enthalpy equation of an ideal gas is:

$$pv = RT, [J/kg] \Rightarrow p/\rho = RT. \tag{1.01}$$

First, from ideal gas (1.01) eq. express gas (air) density term and plug this term into a differential equation for compressible gas:

$$\rho = \frac{p}{RT} \Rightarrow dp = \rho g dh \Rightarrow dp = \frac{p}{RT} g dh. \tag{1.02}$$

Multiply the (1.02) equation by 1/p and later integrate the left side from dp in a range from p_0 to p_1 and the right side from dh in a range from h_0 =0 to h:

$$\int_{p_0}^{p} \frac{dp}{p} = \frac{g}{RT} \int_{0}^{h} dh \implies -\ln\left(\frac{p}{p_0}\right) = \frac{gh}{RT} \implies p = p_0 e^{\frac{-gh}{RT}}.$$
(1.03)

Equation (1.03) shows isothermal air pressure decreases from height altitude change p(h). The density of air change $\rho(h)$ from height (altitude) change is:

$$\rho = \rho_0 e^{\frac{-gh}{RT}} \tag{1.04}$$

We can prove that pressure and density are similar units $p \sim \rho$ or $p \sim 1/v$:

$$p = p_0 e^{\frac{-gh}{RT}}; \ v = v_0 e^{\frac{gh}{RT}}; \ p_0 e^{\frac{-gh}{RT}} \cdot v_0 e^{\frac{gh}{RT}} = RT \implies pv = RT = const.$$
 (1.05)

Calculation of average air density based on intensive property of density. Average air temp. is $T=273+20^{\circ}C=293\,K$, at $h_0=0$, thus $\rho(h_0)=1.205$, to $h=10^6$ meters, thus $h\to\infty$, $\rho(h)\to0$, a sum of air density performed by an increment of 1 meter $\delta=h/N=1$ till $h=N\delta$:

$$\overline{\rho} = \frac{\rho_0}{N} \sum_{n=0}^{N=10^6} e^{\frac{-g\delta}{RT}n}$$
 (1.06)

The sum of (1.06) eq. express via infinite geometric series, then $N=10^6 \rightarrow \infty$, and x<1:

$$S = \sum_{n=0}^{N \to \infty} x^n = 1 + x + x^2 + x^3 + \dots = \frac{1}{1 - x}, \ e^{-x} = e^{\frac{-g\delta}{RT}} = 0.999883 < 1$$
 (1.07)

$$\overline{\rho} = \frac{\rho_0}{N} \sum_{n=0}^{N} e^{-xn} = \frac{\rho_0}{N} \sum_{n=0}^{N} e^{\frac{-g\delta}{RT}n} = \frac{\rho_0}{N} \cdot \frac{1}{1 - e^{-g\delta/RT}} = 0.0103298 \ kg/m^3$$
 (1.08)

$$e^{-x} = \sum_{n=0}^{N \to \infty} (-1)^n \frac{x^n}{n!} = 1 - x + \frac{x^2}{2!} - \frac{x^3}{3!} + \frac{x^4}{4!} - \dots$$
 (1.09)

For $x=g\delta/RT=1.183\cdot10^4\approx0$ we approximate (1.09) eq. to such expression:

$$e^{-x} \approx 1 - x \implies e^{\frac{-g\delta}{RT}} \approx 1 - g\delta/RT.$$
 (1.10)

Plug (1.10) eq. into (1.08) eq., we get:

$$\overline{\rho} = \frac{\rho_0}{N} \cdot \frac{1}{1 - e^{-g\delta/RT}} = \frac{\rho_0}{N} \cdot \frac{1}{1 - (1 - g\delta/RT)} = \frac{\rho_0 RT}{gN\delta} = \frac{p_0}{gh}$$
(1.11)

The result of (1.11) eq. is known, see (1.08) eq., thus atmospheric pressure is:

$$p_0 = \rho gh = 1.01335 \cdot 10^5 \cong 10^5 Pa$$
 (1.12)

Atmospheric pressure result in (1.12) eq. is accurate to the actual measured value $-p_0=1.01325\times10^5$ Pa [1]. Mass of air is calculated by summing air density from range $h_0=0$ till $h=10^6$ m, $N=10^6$, over $\delta=h/N=1$ meter increments, and multiplying by sphere volumes in the shape of "onion rings" over $\delta=h/N=1$ meter radius increments starting from a radius of Earth $-R_{eth}=3178000$ m, for more details see (Fig .2):

$$M_{air} = \rho_0 \sum_{n=0}^{N} e^{\frac{-g\delta}{RT}n} \cdot \frac{4}{3} \pi \sum_{n=0}^{N} \left[\left(R_{eth} + \delta(n+1) \right)^3 - \left(R_{eth} + \delta n \right)^3 \right] = 1.318 \times 10^{18} \ kg$$
 (1.13)

The convergence condition to the (1.13) equation is:

$$M_{air} = \lim_{n \to \infty} \sum \rho_n \cdot \sum (V_{n+1} - V_n) \to \text{finite}$$
 (1.14)

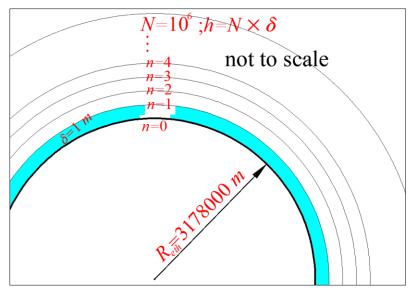


Fig. 1 Sketch for Earth air mass calculation over 1-meter increment.

Error calculation; two facts increase the total mass of the Earth's atmosphere and one factor decreases. Temperature increases starting from h=0, $T(\delta)$ =0.0065 K/m till 100 km height – Karman line also has name Thermosphere, from that point temperature starts again decreasing. And free fall acceleration constant dependable from height:

$$g(\delta) = \frac{GM_{earth}}{(R_{erth} + \delta)^2}, \left[\frac{m}{s^2}\right]$$
 (1.15)

Most of Earth's atmosphere is close to Earth's ground thus $2 \, km$ where atmosphere is dense. A negative factor is a landmass, which is any object that is above sea level [1], mostly high mountains peaks displace Earth's atmosphere. The Earth's mass is $10^{15} \, kg$ bigger than calculated by the (1.13) equation:

$$M_{air} = 1.318 \times 10^{18} kg + 10^{15} kg = 1.319 \times 10^{18} kg$$
 (1.16)

Errors are 1/1000th or 0.1% bigger than calculated.

According to Google search, Earth's mass is M_{air} =5.15×10¹⁸ kg, the ratio is \cong 4 times bigger than calculated here. Chat-gpt inquiry is close to the same result of M_{air} =5.14×10¹⁸ kg.

GitHub Python code: https://github.com/ArthurKarbocius/Earth-s_air_mass