# Multivariate Spatial Extreme Value Analysis of Reconstructed Coastal Sea-Level Time-Series

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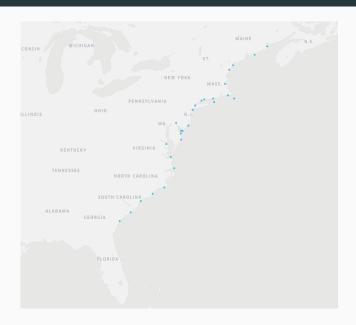
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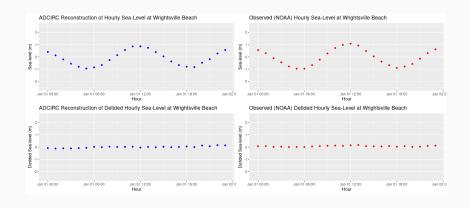
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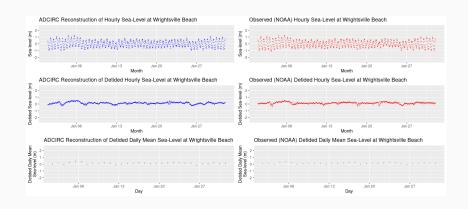
#### Introduction

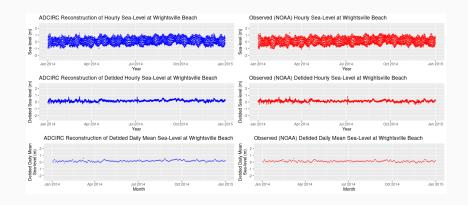
- Research aim: To understand how sea-level extremes along the U.S. East Coast have varied over space and time in the past and to make predictions for how these patterns will evolve in the future under different climate change projections.
- 2. Tools: Oceanography, climate-science, statistical methodology (i.e. extreme value theory. spatial statistics)
- 3. Data:
  - (a) Hourly sea-level time-series taken from NOAA observation stations along the U.S. East Coast over a 40-year period.
  - (b) Corresponding model-generated (ADCIRC) reconstruction of historic sea-level time-series.

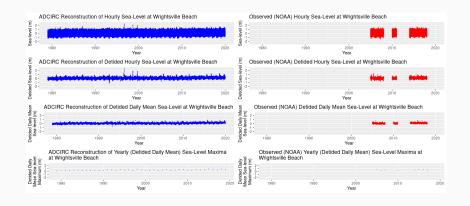
#### **NOAA Sea-Level Observation Stations**

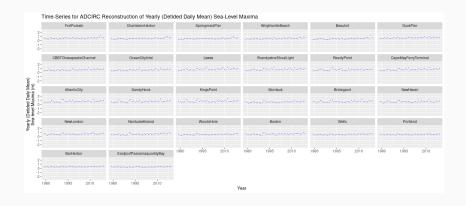


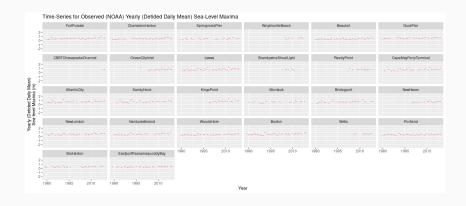




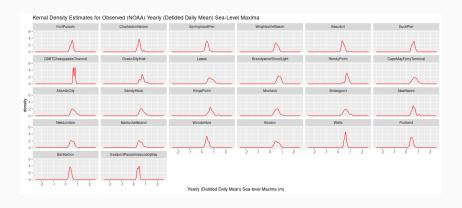




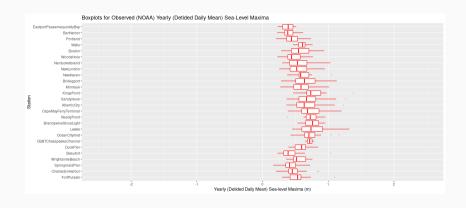












## Background on Univariate Extreme Value Theory

### Generalized Extreme Value (GEV) Distribution

Let  $X_1, X_2, ..., X_n$  be i.i.d. random variables with common CDF F and let  $M_n := \max\{X_1, ..., X_n\}$ .

Then, when appropriately centered and scaled,  $M_n$  converges in distribution to a member of the GEV family:

$$G(z) := \exp\left\{-\left[1 + \xi\left(\frac{z - \mu}{\sigma}\right)\right]_{+}^{-\frac{1}{\xi}}\right\} \tag{1}$$

- $\mu \in \mathbb{R}$  is the location parameter
- $\cdot$   $\sigma \in \mathbb{R}^+$  is the scale parameter
- $\xi \in \mathbb{R}$  is the shape parameter

## Background on Univariate Extreme Value Theory

#### r-Year Return Level

The GEV parameters can be used to characterize extremes via the quantile function of the GEV.

For  $Z \sim GEV(\mu, \sigma, \xi)$ , the 1-p quantile of Z (i.e. the value that is exceeded with probability p) is given by:

$$Z_{p}(\mu, \sigma, \xi) := \begin{cases} \mu - \frac{\sigma}{\xi} [1 - \{-\log(1-p)\}^{-\xi}] & \xi \neq 0 \\ \mu - \sigma \log\{-\log(1-p)\} & \xi = 0 \end{cases}$$
 (2)

When each year contains exactly one block the r-year return level is given by  $Z_{r-1}(\mu, \sigma, \xi)$ 

# Background on Univariate Extreme Value Theory

#### Inference

The GEV parameters are usually estimated by a likelihood based method.

In particular, for  $Z_1,...,Z_m \stackrel{iid}{\sim} GEV(\mu,\sigma,\xi)$  the maximum likelihood estimate (MLE) of  $\theta:=(\mu,\sigma,\xi)^t$  is:

$$\hat{\theta} := (\hat{\mu}, \hat{\sigma}, \hat{\xi})^{t} := \underset{\mu, \sigma, \xi}{\operatorname{argmax}} \left\{ \ell(\mu, \sigma, \xi | Z_{1}, ..., Z_{m}) \right\}$$
(3)

where  $\ell(\mu, \sigma, \xi | Z_1, ..., Z_m)$  is the log-likelihood of  $Z_1, ..., Z_m$ .

The r-year return level estimate is simply  $Z_{r-1}(\hat{\mu}, \hat{\sigma}, \hat{\xi})$ .

## **Modeling Questions**

- 1. How do the GEV parameters and r-year return levels depend upon their spatial location?
- 2. How should one take into account dependence between annual maxima?
- 3. How should the NOAA data be used to validate the ADCIRC reconstruction?
- 4. How should one incorporate global climatic information?

#### **Methods: Overview**

Idea (Russell et al. 2019): Multivariate spatial extreme value model fit by a 2-stage inference procedure.

- a. Independently model the yearly (detided daily mean) sea-level maxima at each station using the GEV distribution.
  - b. Perform inference via MLE.
- a. Model the MLE output from stage 1 as a multi-dimensional Gaussian process with measurement error.
  - b. Perform inference via MLE.

The output of stage 2 can then be used to spatially interpolate the GEV parameters and return-levels along the coastline via Kriging.

#### **Methods: Latent Process**

Let Y(s) be the yearly (detided daily mean) maximum sea-level at location  $s \in \mathcal{D} \subset \mathbb{R}^2$  and assume

$$Y(s) \sim GEV(\mu(s), \sigma(s), \xi(s))$$

To characterize how the sea-level extremes vary spatially define, at both observed and unobserved locations, the latent Gaussian process

$$\theta(\mathsf{s}) = \beta + \eta(\mathsf{s}) \tag{4}$$

for  $\theta(s) := (\mu(s), \log(\sigma(s)), \xi(s))^t$ .

Here  $\beta$  is a vector of mean parameter values over  $\mathcal{D}$  and  $\eta(s)$  a vector of spatially correlated random effects.

#### **Methods: Latent Process**

The spatially correlated random effects are defined by the relation

$$\eta(\mathsf{s}) := \mathsf{A}\delta(\mathsf{s})$$
(5)

where A is a lower-triangular matrix and  $\delta(s)$  is a vector of independent second-order stationary Gaussian processes with mean 0 and covariance function

$$Cov(\delta_i(\mathbf{s}), \delta_i(\mathbf{s}')) = \exp\left(\frac{-||\mathbf{s} - \mathbf{s}'||}{\rho_i}\right)$$
 (6)

for  $s, s' \in \mathcal{D}$  where  $\rho_i > 0$  is the range parameter.

#### Methods: Latent Process with Measurement Error

For NOAA station  $l \in \{1, ..., 26\}$ , let  $\hat{\theta}(\mathbf{s}_l)$  be the point-wise MLE for the GEV distribution associated with  $Y(\mathbf{s}_l)$ . We assume that

$$\hat{\boldsymbol{\theta}}(\mathbf{s}_l) = \boldsymbol{\theta}(\mathbf{s}_l) + \epsilon(\mathbf{s}_l) \tag{7}$$

where  $\epsilon(\mathbf{s}_l)$  is estimation error that is independent of  $\eta$ .

Thus, the latent process with measurement error at station l is

$$\hat{\theta}(s_l) = \theta(s_l) + \epsilon(s_l)$$

$$= \beta + \eta(s_l) + \epsilon(s_l)$$

$$= \beta + A\delta(s_l) + \epsilon(s_l)$$
(8)

# Methods: Latent Process with Measurement Error

Now, let

 $\boldsymbol{\Theta} := (\boldsymbol{\theta}(s_1), ..., \boldsymbol{\theta}(s_{26}))^t$ 

and

$$\hat{\mathbf{\Theta}} := \left(\hat{\boldsymbol{\theta}}(\mathsf{s}_1), ..., \hat{\boldsymbol{\theta}}(\mathsf{s}_{26})\right)^{\mathsf{T}}$$

$$\epsilon := \left(\epsilon(\mathsf{s}_1),...,\epsilon(\mathsf{s}_{26})\right)^t \sim \mathcal{N}_{78}(\mathsf{0},\mathsf{W})$$

where W is unknown and estimated via a regularized non-parametric bootstrap procedure (i.e.  $\tilde{W}$ ) and let

$$\mathsf{Cov}(\hat{oldsymbol{\Theta}}) := oldsymbol{\Sigma}_{oldsymbol{
ho}, \mathsf{A}}.$$

(11)

(9)

#### Methods: Latent Process with Measurment Error

Thus,

$$\hat{\mathbf{\Theta}} = \mathbf{\Theta} + \boldsymbol{\epsilon} \tag{12}$$

and hence

$$\hat{\boldsymbol{\Theta}} \sim \mathcal{N}_{78} \left( \mathbf{1}_{26} \otimes \boldsymbol{\beta}, \boldsymbol{\Sigma}_{\mathsf{A}, \boldsymbol{\rho}} + \tilde{W} \right) \tag{13}$$

Therefore, given  $\hat{\Theta}$  (i.e. the output from the 1st stage of inference) and  $\tilde{W}$ , we can obtain  $\hat{\beta}$ ,  $\hat{\rho}$  and  $\hat{A}$  via MLE.

# Methods: Kriging (Gaussian Process Regression)

Given  $\hat{\beta}$ ,  $\hat{\rho}$  and  $\hat{A}$  (i.e. the output from the 2nd stage of inference) we then interpolate  $\hat{\theta}(s)$  over  $s \in \mathcal{D}$  via Kriging:

$$\begin{split} \boldsymbol{\hat{\theta}}(\mathbf{s}) &= \boldsymbol{\hat{\beta}} + Cov(\boldsymbol{\hat{\theta}}(\mathbf{s}), \boldsymbol{\hat{\Theta}}) (\boldsymbol{\Sigma}_{\boldsymbol{\hat{\rho}}, \hat{A}} + \tilde{W})^{-1} (\boldsymbol{\hat{\Theta}} - \boldsymbol{1}_{26} \otimes \boldsymbol{\hat{\beta}}) \\ Var(\boldsymbol{\hat{\theta}}(\mathbf{s})) &= Var(\boldsymbol{\hat{\theta}}(\mathbf{s})) - Cov(\boldsymbol{\hat{\theta}}(\mathbf{s}), \boldsymbol{\hat{\Theta}}) (\boldsymbol{\Sigma}_{\boldsymbol{\hat{\rho}}, \hat{A}} + \tilde{W})^{-1} Cov(\boldsymbol{\hat{\theta}}(\mathbf{s}), \boldsymbol{\hat{\Theta}})^{t} \end{split}$$

and compute the 100-Year Return Level estimates for the Yearly (Detided Daily Mean) Sea Level Maxima:

$$Z_{100^{-1}}(s) = Z_{100^{-1}}(\hat{\boldsymbol{\theta}}(s))$$
$$Var(Z_{100^{-1}}(s)) = \nabla Z_{p}(\hat{\boldsymbol{\theta}}(s))^{t} Var(\hat{\boldsymbol{\theta}}(s)) \nabla Z_{p}(\hat{\boldsymbol{\theta}}(s))$$

# Preliminary Results: Stage 1 Output (ADCIRC)

GEV Parameter Estimate at Each Station (ADCIRC)				
station	location	log.scale	shape	
FortPulaski	0.53	-2.14	0.00	
CharlestonHarbor	0.47	-2.20	-0.05	
SpringmaidPier	0.44	-2.18	-0.08	
WrightsvilleBeach	0.45	-2.00	-0.14	
Beaufort	0.42	-2.34	0.13	
DuckPier	0.61	-2.15	-0.15	
CBBTChesapeakeChannel	0.60	-2.10	-0.05	
OceanCityInlet	0.58	-2.02	-0.03	
Lewes	0.72	-1.79	-0.03	
BrandywineShoalLight	0.68	-1.81	-0.11	
ReedyPoint	0.63	-1.93	-0.08	
CapeMayFerryTerminal	0.67	-1.80	-0.13	
AtlanticCity	0.67	-1.89	0.02	

# Preliminary Results: Stage 1 Output (ADCIRC)

GEV Parameter Estimates at Each Station (ADCIRC)				
station	location	log.scale	shape	
SandyHook	0.66	-1.88	0.00	
KingsPoint	0.69	-1.84	0.06	
Montauk	0.54	-1.94	-0.13	
Bridegport	0.63	-1.94	-0.06	
NewHaven	0.61	-1.99	-0.09	
NewLondon	0.54	-2.07	-0.10	
NantucketIsland	0.49	-2.02	0.00	
WoodsHole	0.44	-2.33	-0.14	
Boston	0.53	-1.89	-0.25	
Wells	0.46	-2.29	-0.21	
Portland	0.45	-2.43	-0.26	
BarHarbor	0.39	-2.78	-0.07	
EastportPassamaquoddyBay	0.35	-2.92	-0.04	

# Preliminary Results: Stage 1 Output (NOAA)

GEV Parameter Estimates at Each Station (NOAA)				
station	location	log.scale	shape	
FortPulaski	0.467	-2.159	-0.031	
CharlestonHarbor	0.405	-2.140	-0.087	
Beaufort	0.356	-2.406	0.156	
DuckPier	0.549	-2.197	-0.248	
Lewes	0.660	-1.820	0.055	
CapeMayFerryTerminal	0.642	-1.744	-0.167	
AtlanticCity	0.582	-1.892	0.032	
SandyHook	0.611	-1.857	-0.010	
Bridegport	0.576	-1.738	-0.126	
NewLondon	0.481	-1.910	-0.182	
NantucketIsland	0.473	-1.987	0.075	
WoodsHole	0.400	-2.181	-0.166	
Boston	0.508	-1.783	-0.216	
Portland	0.403	-2.238	-0.190	

# Preliminary Results: Stage 2 Output (ADCIRC)

# Gaussian Process Parameter Estimates (ADCIRC)

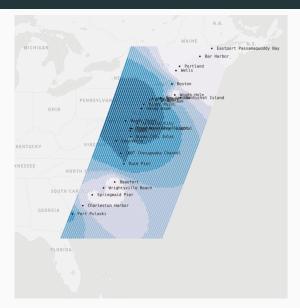
$\hat{\beta}$	$\hat{ ho}$		$\hat{A}$	
0.50	37.67	0.10	0.00	0
-2.28	145.60	0.19	-0.15	0
-0.09	144.62	0.05	-0.04	0

# Preliminary Results: Stage 2 Output (NOAA)

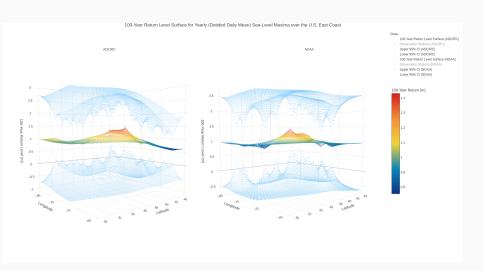
Gaussian Process Parameter Estimates (NOAA)

$\hat{\beta}$	$\hat{ ho}$	$\hat{A}$		
0.48	21.84	0.09	0	0
-2.08	163.79	0.18	0	0
-0.07	129.20	0.03	0	0

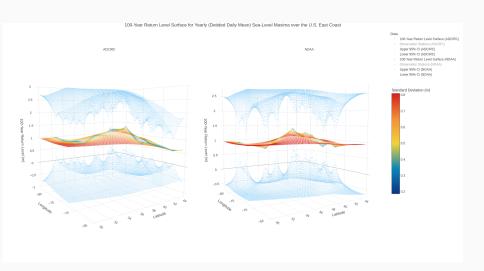
# Preliminary Results: 100-Year Return Level Heat Map for Yearly (Detided Daily Mean) Sea-Level Maxima



# Preliminary Results: 100-Year Return Level Surface for Yearly (Detided Daily Mean) Sea-Level Maxima



# Preliminary Results: 100-Year Return Level Surface for Yearly (Detided Daily Mean) Sea-Level Maxima



# Next Steps & Future Work

#### Next Steps

 Use the 100-year return level surface based on the NOAA data to improve the performance of the corresponding surface based on the ADCIRC reconstruction.

#### **Future Work**

- 1. Introduce global climatic covariate(s) in the 1st stage of inference.
- 2. Examine how the r-year return-level surface along the coastline changes as a function of these covariates.

#### References

- An Introduction to Statistical Modeling of Extreme Values. Coles. Springer Series in Statistics Springer-Verlag, London, 2001.
- Investigating the association between late spring Gulf of Mexico sea surface temperatures and US Gulf Coast precipitation extremes with focus on Hurricane Harvey. Russel, Riser, Smith and Kunkel. Environmetrics 31(1), 2019.