

Improvements in modelling the diurnal temperature amplitude in ACCESS

www.cawcr.gov.au



Eva Kowalczyk, Lauren Stevens,
Ian Harman and Vanessa Haverd



Australian Government
Bureau of Meteorology

The Centre for Australian Weather and Climate Research
A partnership between CSIRO and the Bureau of Meteorology



Diurnal Temperature Range

$$\text{DTR} = \sum_{d=1}^{N_d} (\max T_d - \min T_d) / N_d$$

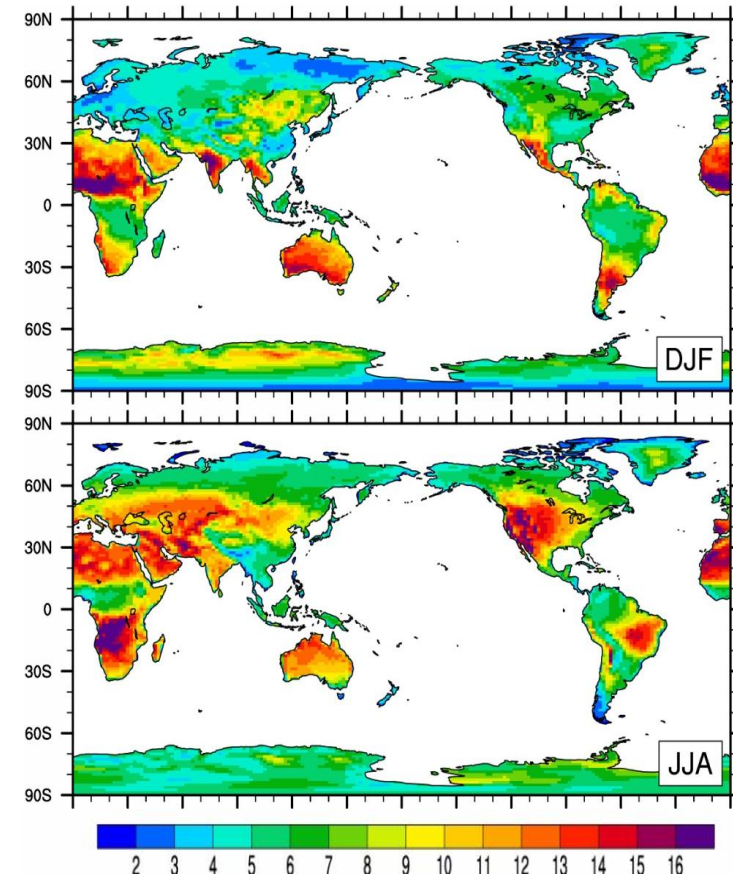
$\max T_d$ and $\min T_d$ are the daily max, min
and N_d is of days

$\max T_d$ is largely determined by:

- shortwave forcing, cloud cover & precipitation, stability of the atmosphere
- evapotranspiration, soil moisture availability, surface parameters affecting surface exchange processes (albedo & roughness)

$\min T_d$ is largely determined by:

- incoming longwave affected by cloud cover, stability of the atmosphere, roughness length

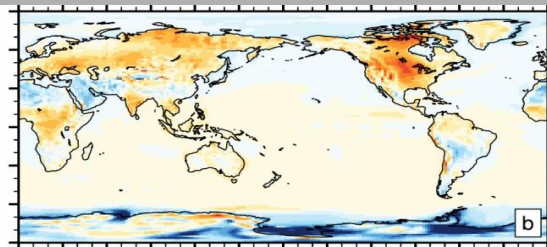
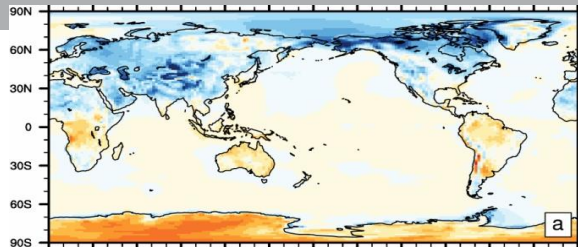


Seasonal Mean Screen Temperature Bias (K)



DJF

JJA

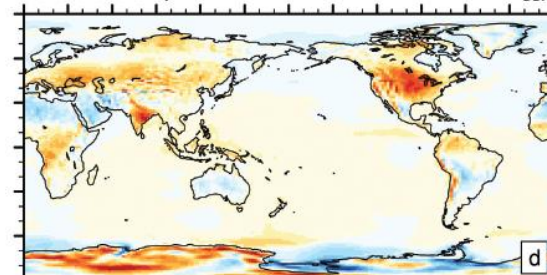
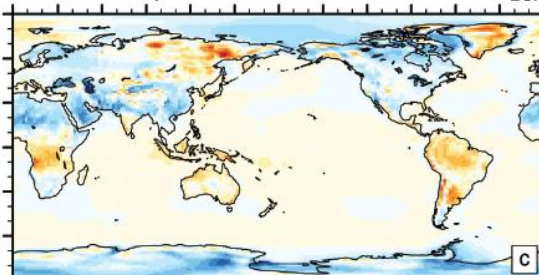


ACCESS1.3 amip

DJF

ACCESS1.3 amip

JJA

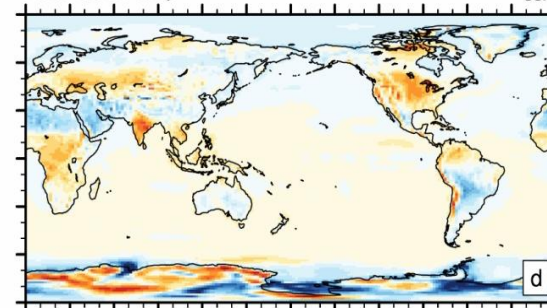
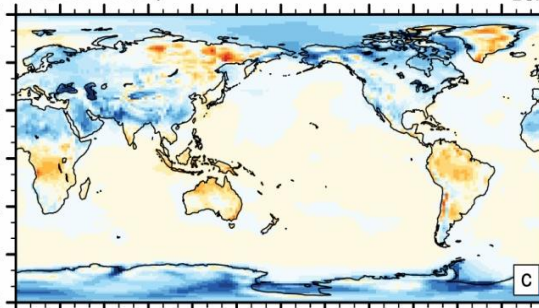


ACCESS1.1 amip

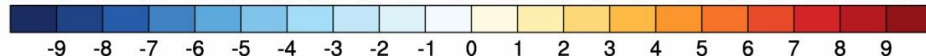
DJF

ACCESS1.1 amip

JJA



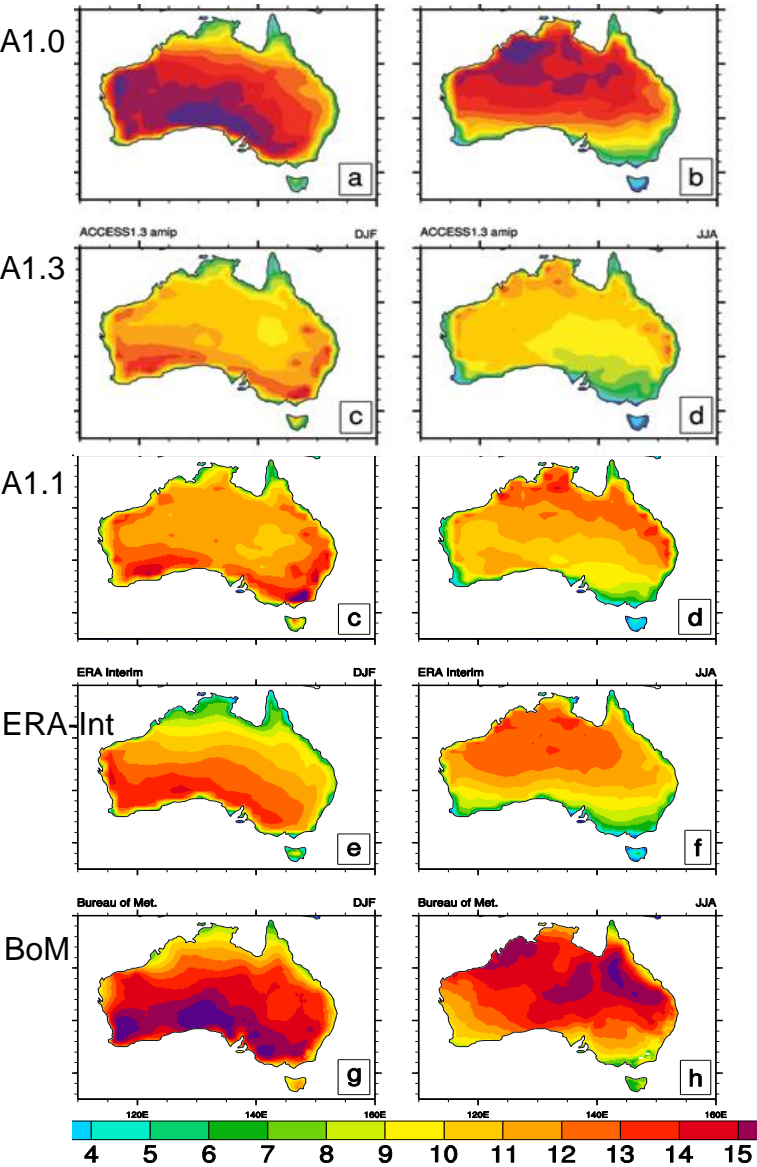
Model	Mean land temp	Mean precip
1.0	8.73	2.13
1.1	8.58	2.20
1.3	9.15	2.57



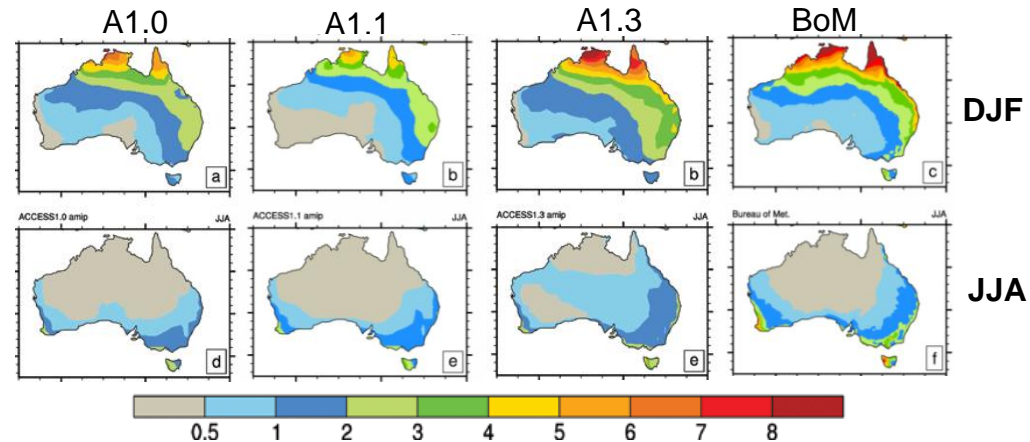
Impact of precipitation on diurnal amplitude in ACCESS



Diurnal amplitude for DJF left column and JJA right column



DFJ and JJA Precipitation



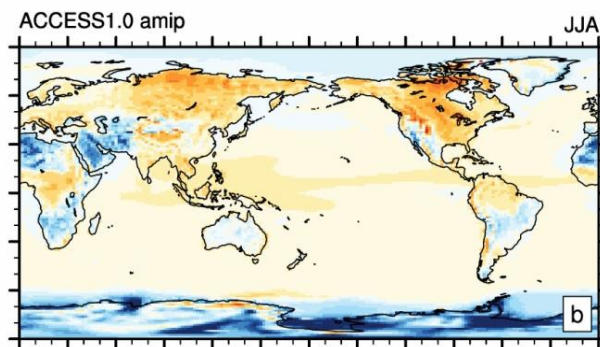
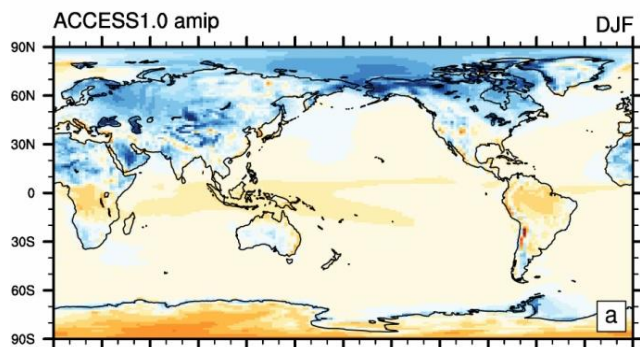
In A1.3 HadGEM3' settings include PC2 cloud scheme



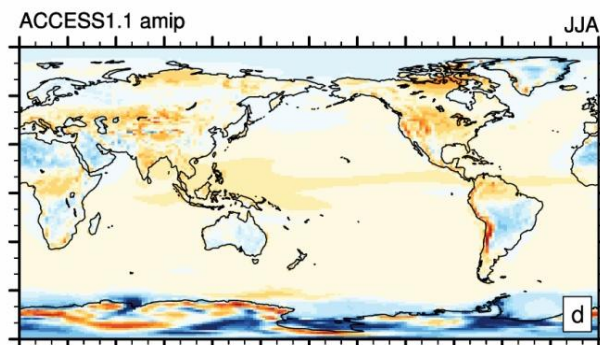
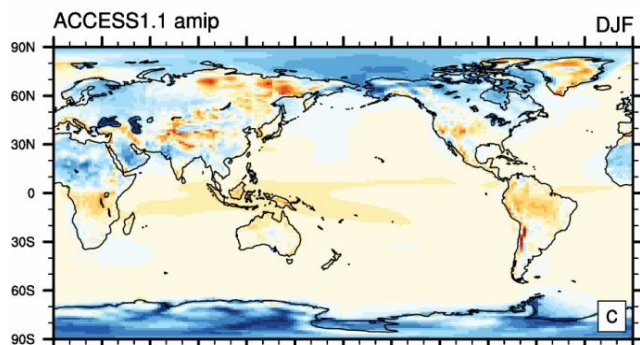
Australian Government
Bureau of Meteorology



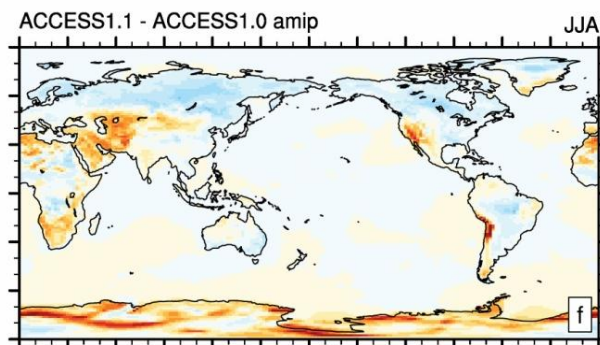
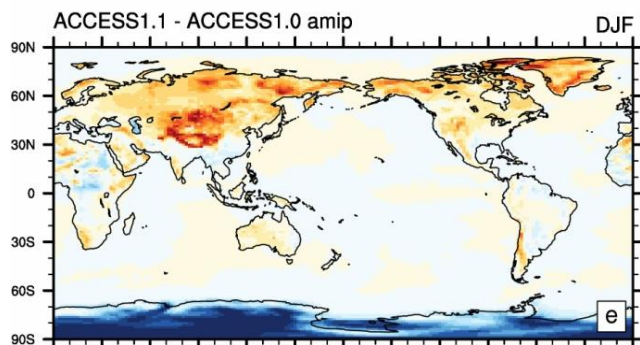
Seasonal minimum screen temperature biases for A1.0 and A1.1 AMIP simulations



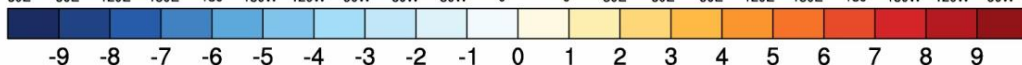
ACCESS1.0 evaluated against ERA-Interim for DJF (left column) and JJA (right column).



ACCESS1.1 simulation



Tmin difference, ACCESS1.1 minus ACCESS1.0



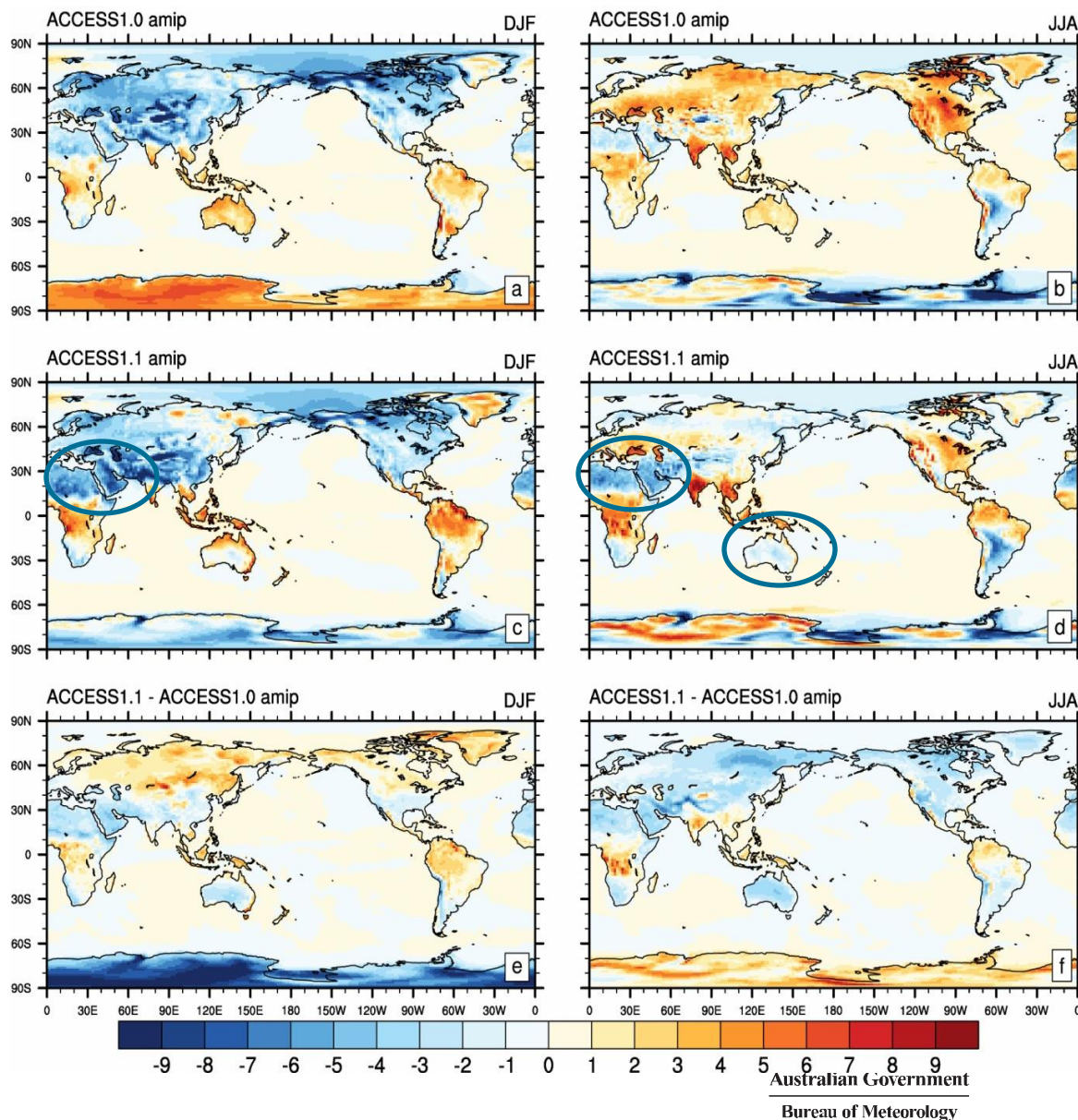
Seasonal maximum screen temperature biases for A1.0 and A1.1 AMIP simulations



ACCESS1.0 evaluated against ERA-Interim for DJF (left column) and JJA (right column).

ACCESS1.1 simulation

Tmax difference, ACCESS1.1 minus ACCESS1.0



Model improvements



- Evapotranspiration and soil moisture availability
 - Changes to restrict transpiration/photosynthesis in dry conditions (Vanessa Haverd)
 - Setting more stringent criteria to maintain water balance
 - Soil evaporation cut off below the soil wilting value
 - Increased surface runoff in areas of high topography
 - Increased speed of drainage for coarse texture soil
- Surface roughness length for bare ground
 - New formula for bare ground roughness length as a function of friction velocity and leaf area index (Ian Harman)

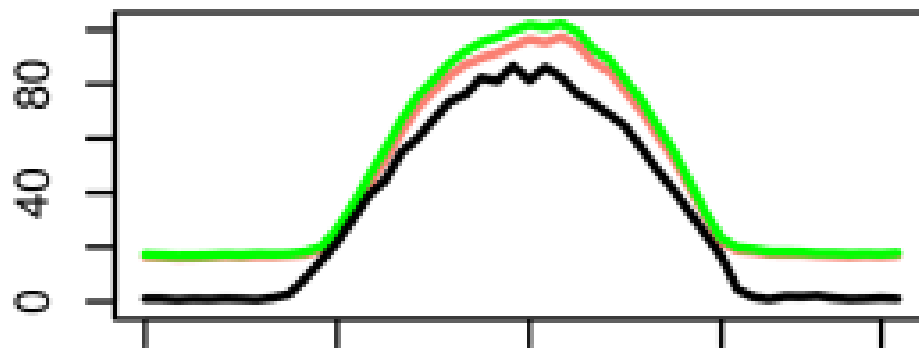


Rescaling the minimum stomatal resistance with soil availability modifier (Venessa Harverd)

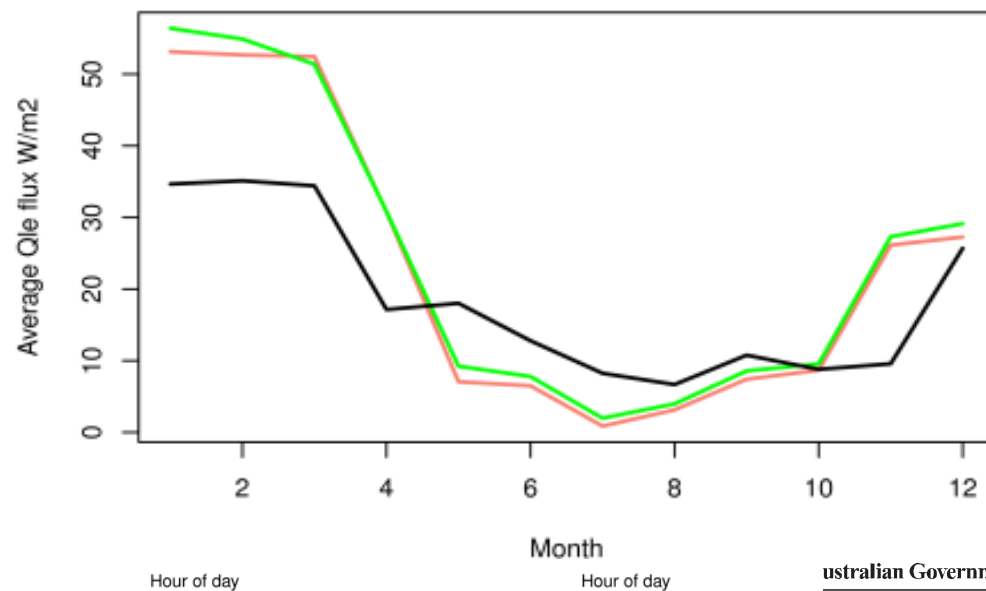


Kruger DJF Qle

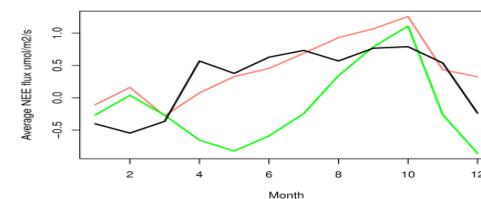
Qle flux W/m²



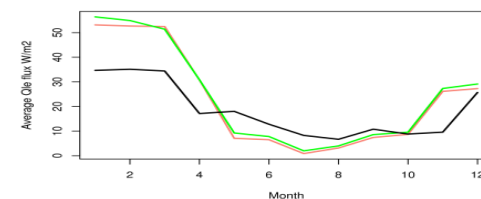
Kruger Qle



Kruger NEE



Kruger Qle



Roughness length (z_0) of sand / bare ground

Constant value of $z_{0\text{soil}}$ in UM/CABLE produced cool bias in T_{mean} over arid and semi arid areas:

- low value of $z_0 = 1.e-6$

gives reasonable T_{min} but T_{max} and T_{mean} are underestimated

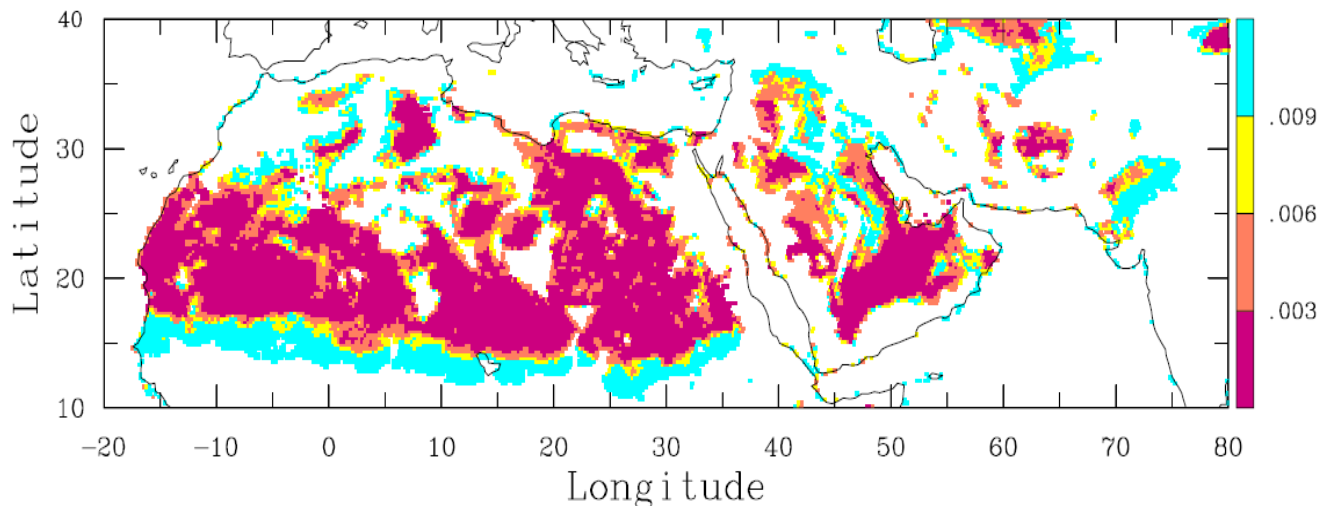
- high value of $z_0 = 1.e-2$

gives reasonable T_{max} but T_{min} and T_{mean} is severely underestimated.

D09205

PRIGENT ET AL.: GLOBAL ROUGHNESS LENGTHS FROM SATELLITE

Std z_0 (cm) for a year from ERS scatterometer



Estimation of the aerodynamic roughness length in arid and semi-arid regions over the globe with the ERS scatterometer. Prigent et al., 2005, JGR

Figure 9. Standard deviation of the ERS z_0 calculated for 1993 over Africa and Middle East.

Roughness length (z_0) of sand / bare ground

Saltation is downwind movement of particles in a series of jumps or skips.

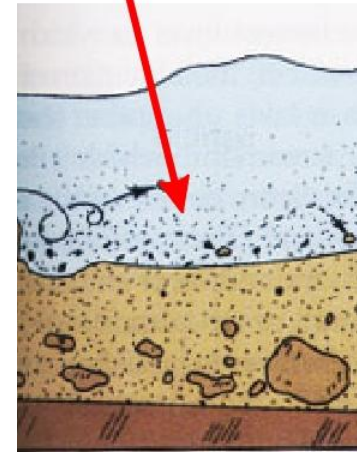
- advected mass flow of sand increases with friction velocity U_*^3
- vertical velocity of the grain is proportional to U_*
- the height to which particle rises is proportional to U_*^2 / g

So the soil roughness length is also proportional to

$$U_*^2 / g$$

Chamberlain BLM 1983

Saltation

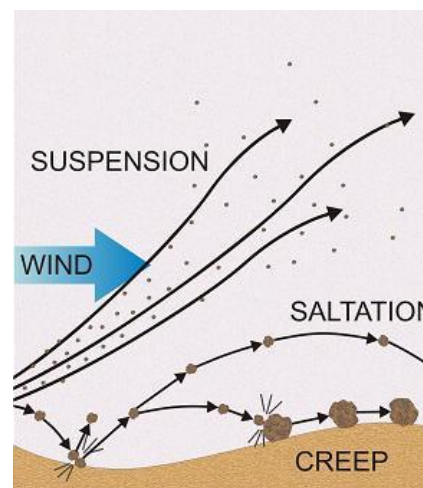
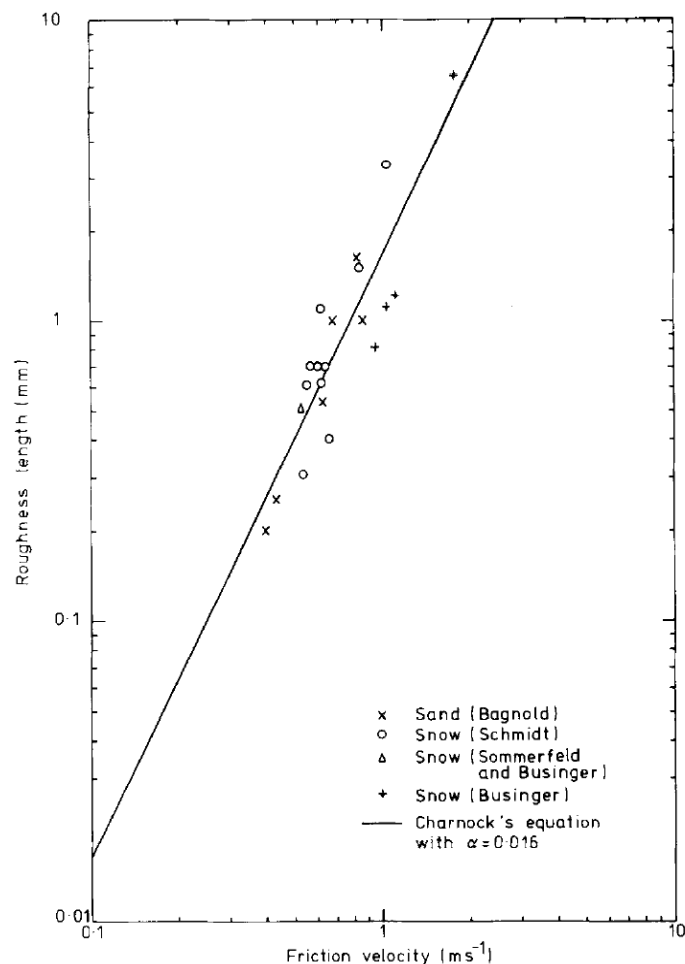


ROUGHNESS LENGTH OF SEA, SAND, AND SNOW

BLM 1983, A. CHAMBERLAIN



ROUGHNESS LENGTH OF SEA, SAND, AND SNOW



$$z_0 = a * (U_* * U_*) / \text{grav} + b * \min(1, \text{lai})$$

where $a = 0.02$, $b = 0.01$

Fig. 2. Roughness length during blowing of sand or snow.



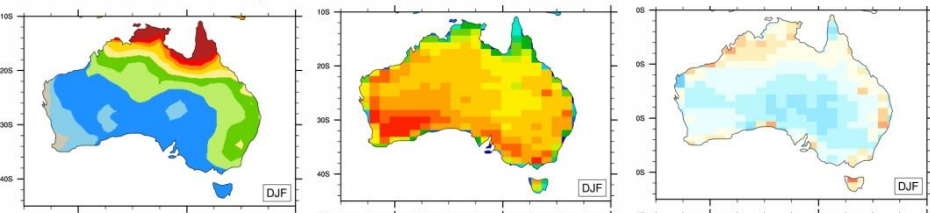
Australian Government
Bureau of Meteorology

The Centre for Australian Weather and Climate Research
A partnership between CSIRO and the Bureau of Meteorology

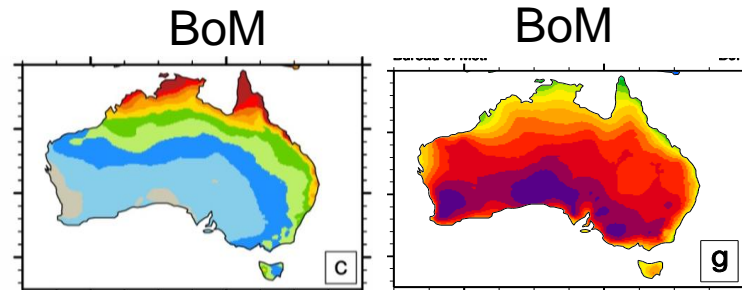
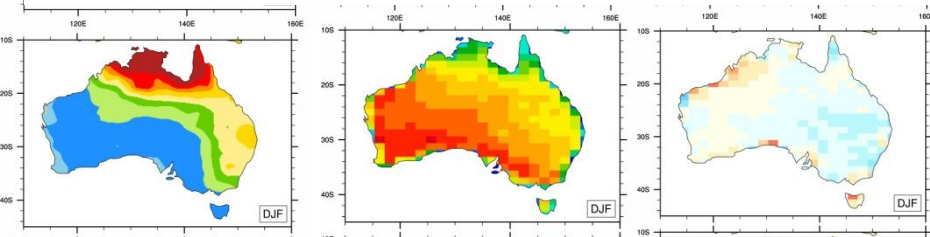


CSIRO

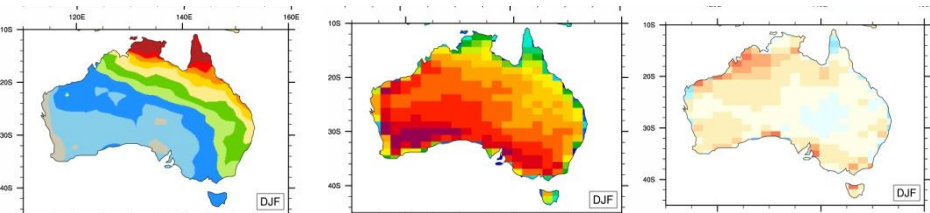
DTR over Australia in UM/CableV2.0 & V2.2 DJF



original V2.0

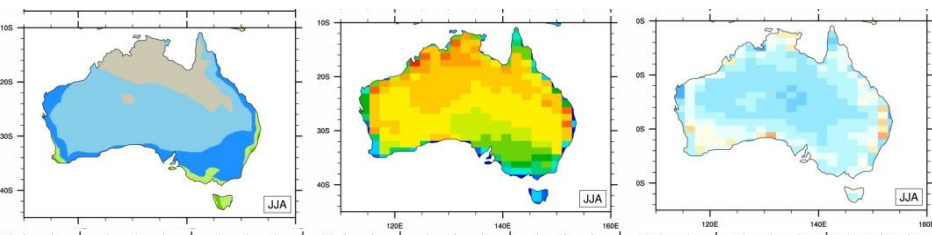


hydrology related updates V2.0

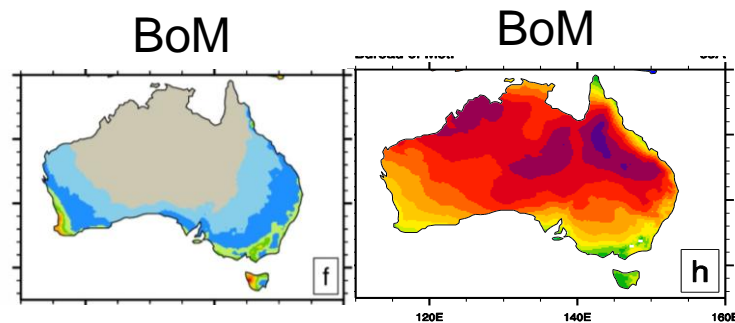
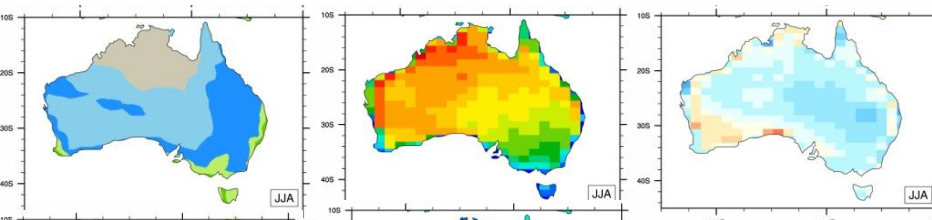


new roughness length V2.2

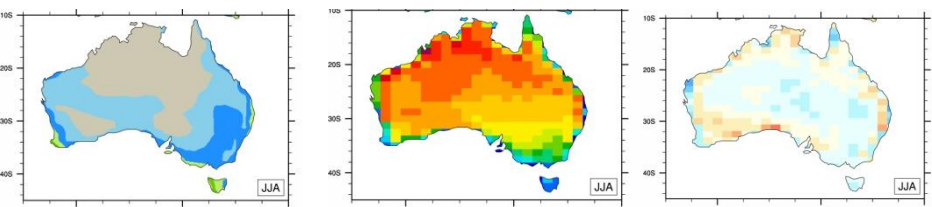
DTR over Australia in UM/CableV2.0 & V2.2 JJA



original V2.0



hydrology related updates V2.0



new roughness length V2.2

The Centre for Australian Weather and Climate Research
A partnership between CSIRO and the Bureau of Meteorology



Future work related to Tmax and Tmin



Amplitude bias V2.2

