

180256: metagranodiorite, Hubert Soak

Location and sampling

SCOTT (SG52-6), BATES (4646)
MGA Zone 52, 488577E 7143422N

Sampled on 24 September 2003

The sample was taken from an outcrop about 1 km southeast of Hubert Soak.

Tectonic unit/relations

The unit sampled is a moderately to strongly foliated seriate to porphyritic metamorphosed granodiorite that contains rare rounded K-feldspar phenocrysts up to 4 cm in length, and pyroxene-rich clots up to 2 cm long that are strongly flattened parallel to the foliation. The granodiorite is part of the Miturtu Monzogranite, within the Mesoproterozoic Pitjantjatjara Supersuite of the Musgrave Complex (Howard et al., in prep.), and was sampled in order to constrain the age of Miturtu plutonism.

Petrographic description

The visually estimated primary mineralogy of this sample includes 15% quartz, 40% K-feldspar, 20% plagioclase, 10% clinopyroxene, 7–8% orthopyroxene, 5% garnet, 2–3% opaque oxide minerals, and accessory minerals (apatite, biotite, and zircon). The quartz and feldspars are anhedral and, apart from rare plagioclase to 8 mm long, are mostly less than 4 mm in size. K-feldspar is micropertite, with very narrow exsolution lamellae of plagioclase. The mafic minerals lie in subparallel lenses up to 7 mm long, and comprise granular clinopyroxene, less abundant granular to prismatic orthopyroxene, granular opaque oxide minerals, and apatite. Partial rims of fine-grained garnet enclose most of these mafic aggregates, especially adjacent to opaque oxide minerals and orthopyroxene. There are also rare flakes of biotite containing symplectites of quartz. Zircon, and possible monazite, forms grains 0.1–0.5 mm in diameter within the mafic aggregates. Some orthopyroxene grains are veined and rimmed by small amounts of brown clay minerals, but most of the minerals reflect granulite facies metamorphism, with garnet rims possibly formed during isobaric cooling from peak metamorphic conditions.

Zircon morphology

Zircons isolated from this sample range from subhedral to well-rounded, and are clear and colourless. They are up to 250 μm long (although some are fragments of larger crystals), with aspect ratios up to 4:1. Many are pervasively cracked. Most crystals consist of central cores (in some cases showing evidence of alteration or resorption) in

which concentric growth zoning is prominent, surrounded by relatively featureless rims. A cathodoluminescence image of representative zircons is shown in Figure 1.

Analytical details

This sample was analysed on 21 November 2004 (using SHRIMP-B) and 9 December 2004 (using SHRIMP-A). Analyses 1.1 to 24.2 were obtained during the first session, together with 13 analyses of the CZ3 standard that indicated a $^{238}\text{U}/^{206}\text{Pb}^*$ calibration uncertainty of 1.80% (1σ). Analyses 26.1 to 55.2 were obtained during the second session, together with 15 analyses of the CZ3 standard. Following the deletion of one standard analysis as an outlier, the remaining 14 analyses indicated a $^{238}\text{U}/^{206}\text{Pb}^*$ calibration uncertainty of 1.71% (1σ).

Common-Pb corrections were initially applied using Broken Hill common-Pb isotopic compositions for all analyses, and these data are presented in Table 1. However, owing to the existence of a strong negative correlation between f_{204} and $^{207}\text{Pb}^*/^{206}\text{Pb}^*$ dates (corrected for common Pb) in each of the analytical sessions (Fig. 2), explicit correction for common Pb was avoided for the purpose of weighted mean date calculations (see below).

Results

Sixty-seven analyses were obtained from 55 zircons, with 12 grains (3, 5, 6, 18, 22, 24, 25, 35, 43, 47, 50, and 55) each analysed twice. Results are listed in Table 1, and shown in Figure 3.

Interpretation

With the exception of analysis 6.1, all analyses yield $^{207}\text{Pb}/^{206}\text{Pb}$ dates (uncorrected for common Pb) of approximately 1180 Ma. However, owing to the existence of a strong negative correlation between common-Pb corrected $^{207}\text{Pb}^*/^{206}\text{Pb}^*$ dates and f_{204} (Fig. 2), the weighted mean date of the zircons is not based on $^{207}\text{Pb}^*/^{206}\text{Pb}^*$ ratios corrected for common Pb by reference to measured $^{204}\text{Pb}/^{206}\text{Pb}$. Instead the weighted mean date is determined from the intersection of a regression line (Fig. 3), anchored at contemporaneous crustal common Pb ($^{207}\text{Pb}/^{206}\text{Pb} = 0.924$ at 1180 Ma; Stacey and Kramers, 1975), with the concordia curve. The analyses can be divided into three groups based on their uncorrected $^{238}\text{U}/^{206}\text{Pb}$ and $^{207}\text{Pb}/^{206}\text{Pb}$ ratios and their discordance values.

Group 1 comprises 63 broadly concordant analyses of 52 zircons that have $^{238}\text{U}/^{206}\text{Pb}$ and $^{207}\text{Pb}/^{206}\text{Pb}$ ratios within error of a regression line (anchored at $^{238}\text{U}/^{206}\text{Pb} = 0$, $^{207}\text{Pb}/^{206}\text{Pb} = 0.924$) that intersects the concordia curve at 1176 ± 6 Ma (MSWD = 1.02).

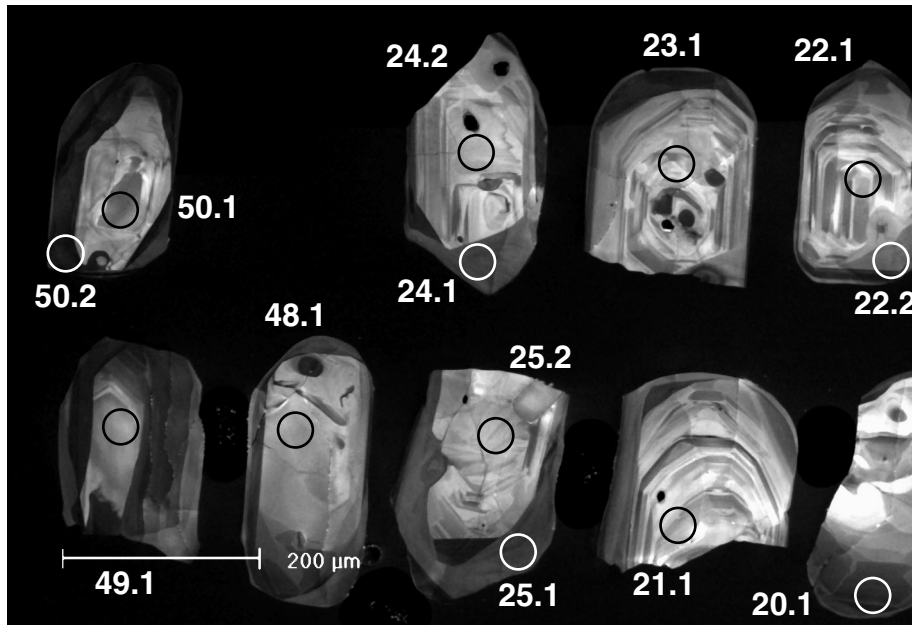


Figure 1. Cathodoluminescence image of representative zircons from sample 180256: metagranodiorite, Hubert Soak. Numbered circles represent approximate positions of analysis sites

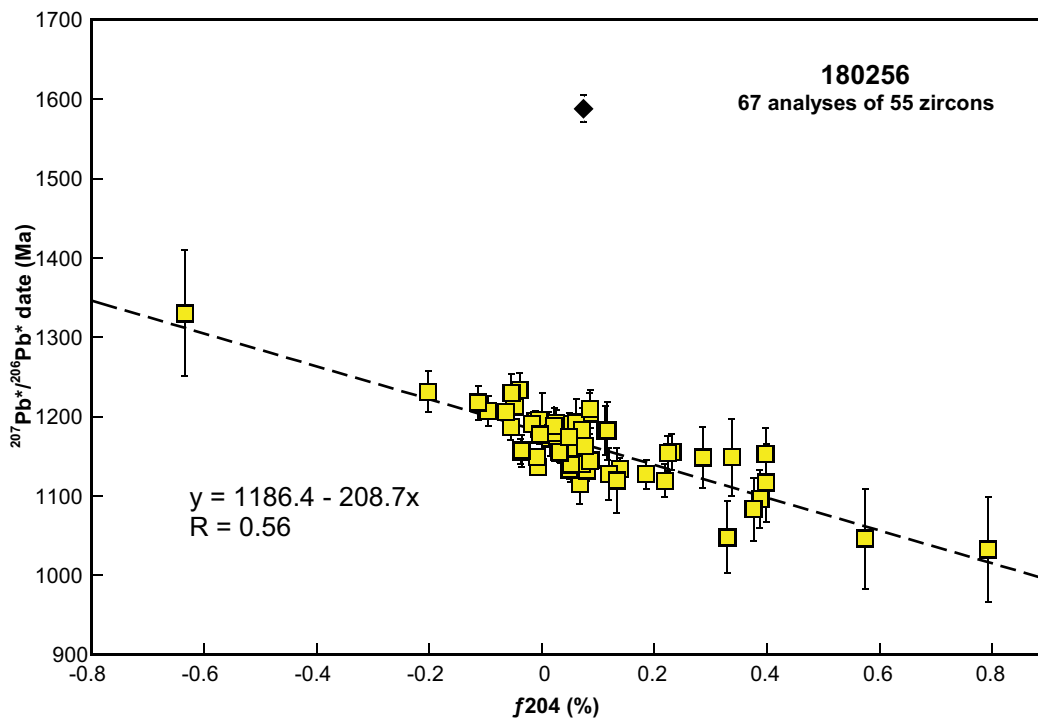


Figure 2. Correlation between f_{204} and $^{207}\text{Pb}^*/^{206}\text{Pb}^*$ date (corrected for common Pb) for sample 180256: metagranodiorite, Hubert Soak. Dashed line depicts a regression through all data (excluding analysis 6.1, shown by filled diamond). The equation of the regression line is included, and R is Pearson's correlation coefficient

Table 1 (continued)

Grain -spot	U (ppm)	Th (ppm)	Th/U	f_{204} (%)	$^{238}\text{U}/^{206}\text{Pb}$ $\pm 1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$ $\pm 1\sigma$	$^{238}\text{U}/^{206}\text{Pb}$ $\pm 1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$ $\pm 1\sigma$	$^{238}\text{U}/^{206}\text{Pb}$ $\pm 1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$ $\pm 1\sigma$	$^{238}\text{U}/^{206}\text{Pb}$ age (Ma) $\pm 1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$ age (Ma) $\pm 1\sigma$	D_{isc} (%)
40.1	249	121	0.50	0.086	5.014 ± 0.090	0.07866 ± 0.00049	5.018 ± 0.090	0.07790 ± 0.00055	5.018 ± 0.090	0.07790 ± 0.00055	1171 ± 19	1144 ± 14	-2.4
41.1	453	208	0.47	-0.006	4.931 ± 0.087	0.07989 ± 0.00041	4.931 ± 0.087	0.07994 ± 0.00041	4.931 ± 0.087	0.07994 ± 0.00041	1190 ± 19	1195 ± 10	0.4
42.1	307	120	0.40	0.033	5.088 ± 0.089	0.07859 ± 0.00047	5.090 ± 0.089	0.07831 ± 0.00053	5.090 ± 0.089	0.07831 ± 0.00053	1156 ± 19	1155 ± 13	-0.1
43.1	432	79	0.51	-0.017	4.943 ± 0.087	0.07960 ± 0.00040	4.943 ± 0.087	0.07975 ± 0.00040	4.943 ± 0.087	0.07975 ± 0.00040	1188 ± 19	1191 ± 10	0.3
43.2	133	72	0.62	-0.113	5.172 ± 0.094	0.07983 ± 0.00074	5.167 ± 0.094	0.08082 ± 0.00090	5.167 ± 0.094	0.08082 ± 0.00090	1141 ± 19	1217 ± 22	6.3
44.1	92	91	1.02	0.119	5.046 ± 0.094	0.07831 ± 0.00086	5.052 ± 0.094	0.07726 ± 0.00128	5.052 ± 0.094	0.07726 ± 0.00128	1164 ± 20	1128 ± 33	-3.2
45.1	121	148	1.26	0.072	5.105 ± 0.093	0.08005 ± 0.00077	5.109 ± 0.093	0.07942 ± 0.00113	5.109 ± 0.093	0.07942 ± 0.00113	1152 ± 19	1183 ± 28	2.6
18.2	448	236	0.54	0.026	4.850 ± 0.085	0.08002 ± 0.00066	4.851 ± 0.085	0.07979 ± 0.00069	4.851 ± 0.085	0.07979 ± 0.00069	1208 ± 19	1192 ± 17	-1.4
46.1	349	175	0.52	0.031	4.979 ± 0.087	0.07860 ± 0.00043	4.981 ± 0.087	0.07832 ± 0.00050	4.981 ± 0.087	0.07832 ± 0.00050	1180 ± 19	1155 ± 13	-2.1
47.1	73	54	0.77	0.133	5.142 ± 0.098	0.07811 ± 0.00099	5.148 ± 0.098	0.07694 ± 0.00157	5.148 ± 0.098	0.07694 ± 0.00157	1144 ± 20	1120 ± 41	-2.2
47.2	397	195	0.51	0.021	4.997 ± 0.089	0.07948 ± 0.00042	4.998 ± 0.089	0.07930 ± 0.00044	4.998 ± 0.089	0.07930 ± 0.00044	1176 ± 19	1180 ± 11	0.3
22.2	253	150	0.61	0.077	5.164 ± 0.091	0.07928 ± 0.00054	5.168 ± 0.091	0.07860 ± 0.00069	5.168 ± 0.091	0.07860 ± 0.00069	1140 ± 18	1162 ± 17	1.9
25.2	90	67	0.77	0.022	5.072 ± 0.094	0.07981 ± 0.00090	5.073 ± 0.095	0.07962 ± 0.00096	5.073 ± 0.095	0.07962 ± 0.00096	1160 ± 20	1188 ± 24	2.3
48.1	114	75	0.68	-0.063	4.941 ± 0.091	0.07982 ± 0.00080	4.938 ± 0.091	0.08038 ± 0.00084	4.938 ± 0.091	0.08038 ± 0.00084	1189 ± 20	1206 ± 21	1.5
49.1	114	107	0.97	-0.202	4.963 ± 0.091	0.07964 ± 0.00081	4.953 ± 0.091	0.08141 ± 0.00107	4.953 ± 0.091	0.08141 ± 0.00107	1186 ± 20	1231 ± 26	3.7
50.1	456	207	0.47	-0.009	5.048 ± 0.088	0.07799 ± 0.00040	5.047 ± 0.088	0.07807 ± 0.00041	5.047 ± 0.088	0.07807 ± 0.00041	1165 ± 19	1149 ± 10	-1.4
50.2	128	155	1.26	-0.036	4.983 ± 0.091	0.07808 ± 0.00078	4.981 ± 0.091	0.07840 ± 0.00080	4.981 ± 0.091	0.07840 ± 0.00080	1179 ± 20	1157 ± 20	-1.9
51.1	134	91	0.70	0.086	5.043 ± 0.092	0.08127 ± 0.00094	5.047 ± 0.092	0.08051 ± 0.00098	5.047 ± 0.092	0.08051 ± 0.00098	1165 ± 19	1210 ± 24	3.7
52.1	320	140	0.45	0.048	4.984 ± 0.088	0.07951 ± 0.00049	4.987 ± 0.088	0.07909 ± 0.00055	4.987 ± 0.088	0.07909 ± 0.00055	1178 ± 19	1174 ± 14	-0.3
53.1	69	52	0.77	0.337	5.034 ± 0.097	0.08105 ± 0.00113	5.051 ± 0.098	0.07807 ± 0.00190	5.051 ± 0.098	0.07807 ± 0.00190	1164 ± 21	1149 ± 48	-1.4
54.1	82	64	0.81	0.329	4.993 ± 0.095	0.07716 ± 0.00098	5.009 ± 0.095	0.07424 ± 0.00167	5.009 ± 0.095	0.07424 ± 0.00167	1173 ± 20	1048 ± 45	-12.0
55.1	13	28	2.22	-0.634	4.946 ± 0.134	0.08010 ± 0.00249	4.915 ± 0.134	0.08565 ± 0.00352	4.915 ± 0.134	0.08565 ± 0.00352	1194 ± 30	1330 ± 80	10.3
55.2	418	251	0.62	-0.004	4.947 ± 0.088	0.07916 ± 0.00047	4.946 ± 0.088	0.07920 ± 0.00047	4.946 ± 0.088	0.07920 ± 0.00047	1187 ± 19	1177 ± 12	-0.8

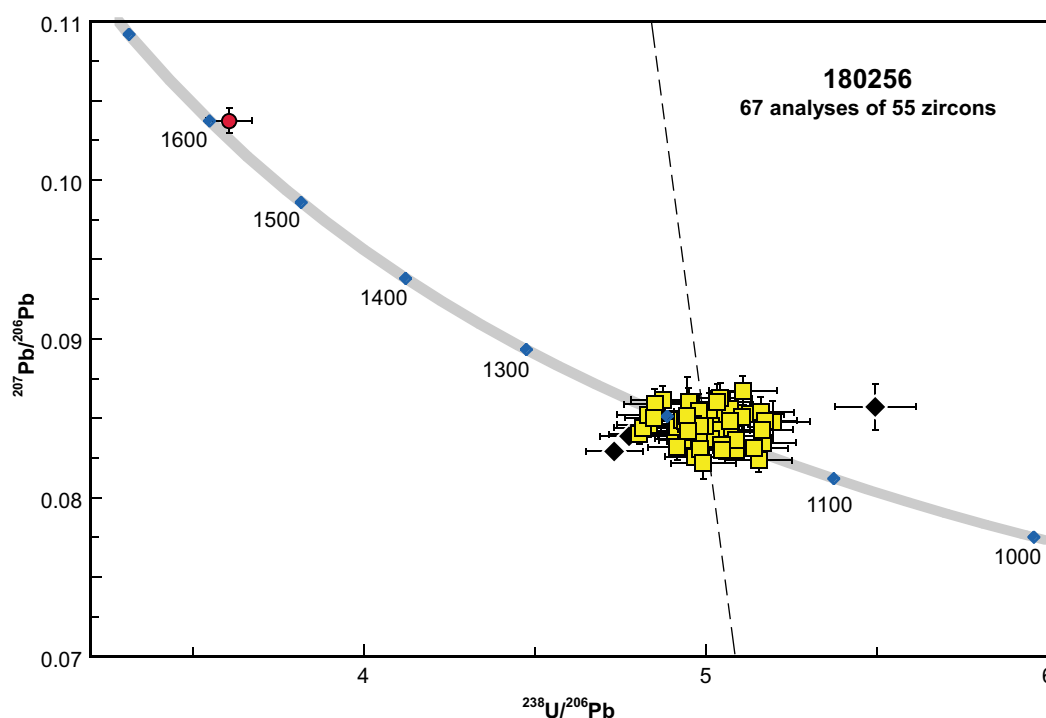


Figure 3. U–Pb analytical data, not corrected for common Pb, for sample 180256: metagranodiorite, Hubert Soak. Squares denote Group 1 (igneous crystallization and metamorphism); filled diamonds denote Group 2 (xenocrystic material or radiogenic Pb loss); filled circle denotes analysis 6.1 (xenocrystic core). Dashed line depicts a regression line defined by the data in Group 1, and anchored at contemporaneous crustal common Pb ($^{207}\text{Pb}/^{206}\text{Pb} = 0.924$ at 1180 Ma; Stacey and Kramers, 1975)

Group 2 comprises three analyses of three zircons (16.1, 27.1, and 33.1) that lie beyond uncertainty of the 1176 ± 6 Ma regression line. Analyses 27.1 and 33.1 are reversely discordant and lie on the old side of the regression line, whereas analysis 16.1 is normally discordant and lies on the younger side of the regression line.

The remaining analysis (6.1) is concordant and has a $^{207}\text{Pb}^*/^{206}\text{Pb}^*$ date of 1588 ± 17 Ma (1σ).

The vast majority of analyses sited within the core and rim zircon identified in the cathodoluminescence images (Fig. 1) yielded uncorrected $^{238}\text{U}/^{206}\text{Pb}$ and $^{207}\text{Pb}/^{206}\text{Pb}$ ratios within uncertainty of the 1176 ± 6 Ma regression line. This implies that igneous crystallization of the precursor granodiorite and subsequent granulite facies metamorphism were not significantly separated in time. Consequently the date of 1176 ± 6 Ma indicated by the uncorrected $^{238}\text{U}/^{206}\text{Pb}$ and $^{207}\text{Pb}/^{206}\text{Pb}$ ratios of the 63 analyses in Group 1 is interpreted as the best estimate of the age of igneous crystallization of the precursor granodiorite, and a maximum age for granulite facies metamorphism. Analysis 6.1 is from a zircon core with anomalously low cathodoluminescence response and well-developed oscillatory zoning, and is interpreted as a xenocrystic core.

Analyses 27.1 and 33.1 yielded uncorrected $^{238}\text{U}/^{206}\text{Pb}$ and $^{207}\text{Pb}/^{206}\text{Pb}$ ratios slightly older than the Group 1 regression line, and possibly incorporated minor amounts

of older (xenocrystic) zircon. Analysis 16.1 yielded uncorrected $^{238}\text{U}/^{206}\text{Pb}$ and $^{207}\text{Pb}/^{206}\text{Pb}$ ratios younger than the Group 1 regression line, and is interpreted to be of a zircon that has undergone minor post-metamorphic loss of radiogenic Pb.

References

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- STACEY, J. S., and KRAMERS, J. D., 1975, Approximation of terrestrial lead isotope evolution by a two-stage model: *Earth and Planetary Science Letters*, v. 26, p. 207–221.

Recommended reference for this publication

- BODORKOS, S., LOVE, G. J., and WINGATE, M. T. D., 2006, 180256: metagranodiorite, Hubert Soak; Geochronology dataset 649, in *Compilation of geochronology data, June 2006 update*: Western Australia Geological Survey.

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