NATURAL OCCURRENCE OF ELEVATED ARSENIC AND SELENIUM IN GEORGIA REGOLITH: IMPLICATIONS FOR THEIR RELATIVE MOBILITY IN PIEDMONT SOILS

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ABSTRACT

Arsenic concentrations of >100 ppm were detected above the regional value of ~7 ppm in soils near the Georgia Brevard Zone using optical emission spectroscopy (OES). Natural, anthropogenic, and false-positive hypotheses were tested to determine a most likely explanation. Induction coupled plasma mass spectroscopy (ICP-MS), X-ray diffraction, optical microscopy, electron microprobe analysis, and historical aerial photographs were subsequently used to determine the most parsimonious hypothesis. ICP-MS results indicated positive detection and verified accuracy of the OES-measured As and concentrations of many other trace elements such as Se. Arsenopyrite is the primary As- and Se-bearing phase in the underlying mafic schist bedrock. The associated bedrock mineral assemblage suggests a fossil hydrothermal system protolith and subsequent prograde and retrograde moderate metamorphism. As/Se ratios in the 20mthick saprolite are much higher (~2000) than regional baseline values for the SE United States (~15) and the underlying bedrock itself (~45). The high soil-saprolite As/Se ratio empirically supports a theoretical ionic potential basis for greater solubility and transport of Se (relative to As) out of the weathered zone and into rivers. Published As/Se ratios for biomonitor proxies living in rivers that drain through the Brevard Zone (~0.6) further support the idea that As in saprolites of the Piedmont in the SE United States is more conservative in fate and transport than Se.

INTRODUCTION

Arsenic (As) and selenium (Se) in soils of the southeastern United States (SE U.S.) have been generally mapped or delineated by low-density and shallow depth (<20 cm) soil sampling surveys of areas relatively unaffected by human activities (Shacklette and Boerngen, 1984) or by site-specific surveys of areas known to be significantly impacted by human activities (Kukier and others, 2001; Jackson, 1998). Canonical baseline values for As and Se in conterminous eastern U.S. surface soils are reported to be about 7 and 0.5 parts-per-million (ppm), respectively, resulting in an As/Se ratio of about 15 (Shacklette and Boerngen, 1984). Numerous international and national regulations and guidelines for As and Se exposure exist for air, water, and food because they are considered to cause adverse health effects when absorbed in high concentrations (Frankenberger, 2002). Examples of exposure limits include (1) the U.S. Environmental Protection Agency (EPA) recommendation of a maximum As contamination limit of 0.010 ppm for drinking water, (2) the U.S. Occupational Safety and Health Organization (OSHA) recommendation of a permissible As exposure limit of 0.01 mg/m³ over an eight hour period, and (3) the U.S. Food and Drug Administration stipulation of As limits at 0.5 to 2.0 ppm for certain animal by-products treated with As-bearing drugs (Anonymous ATSDR report, 2006). The Georgia Environmental Protection Division (EPD) notes that releases resulting in As and Se concentrations in excess of 41 and 36 ppm, respectively requires notification (Georgia State rule 391-3-19-04). Concentrations of As and Se in soils are not specifically regulated by these guidelines, but are factored into the assessment of EPA's Na-

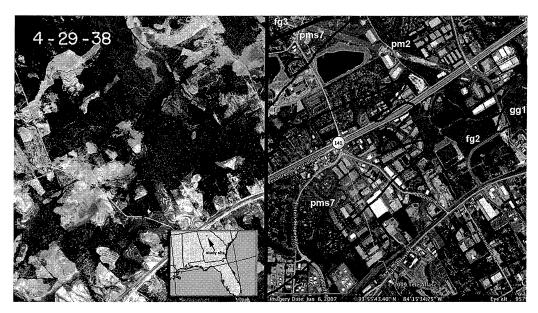


Figure 1. Location map (inset) and aerial photographs of study site. Left image is 1938 photo showing primary land use as forested with remnants of agricultural tilling. Right photo shows 2007 use as mixed residential and industrial parks. Bold dark line in NW corner of right photo delineates SE extent of the Brevard Zone. Lines on right photo delineate formation contacts as published by (Dicken, 2005). All rocks are Precambrian-Cambrian with lithology noted by following symbols: fg2 -Biotitic Gneiss Undifferentiated; gg1 - Granitic Gneiss; pms7 - Button Mica Schist undifferentiated; fg3 - Biotitic Gneiss/Mica Schist/ Amphibolite; pm2 - Metagraywacke/ Mica Schist. Sample location is at 33° 55′ 43.40″ N, 84° 15′ 34.75″ W. Right photo modified from Google Earth © Tele Atlas.

U.S. Comprehensive Environmental Response, Compensation, and Liability Act (CERCLA). As part of due diligence associated with performing CERCLA-prompted Phase I and II environmental site assessments (standardized by EPA), soil As and Se concentrations in excess of regional background levels appear to be the threshold to report a potential for threat to human health and/or the environment. Georgia EPD release thresholds for reporting As and Se are about 6 and 72 times regional baseline values, respectively.

As and Se concentrations observed in excess of the regional background can be attributed to multiple factors. Anthropogenic factors and natural factors are the two main categories for explaining or finding elevated levels in soils. Human activities such as pesticide application, wood preservative production, coal fly ash production, and acid-sulfide mine tailing accumulations account for the major sources of

anthropogenic As and Se contaminant sites in the SE United States (Frankenberger, 2002). The purpose of this study is to document a natural occurrence of elevated As and Se concentrations in soils and to provide empirical insights into their relative fate and transport in the near surface weathering environment of the Piedmont SE United States.

BACKGROUND

The study site is located in the Piedmont of Georgia, Gwinnett County (Figure 1 inset) and was initially identified in response to reconnaissance work associated with a phase 1 assessment, in which As-bearing soils were noted in surface auger samples tested by optical emission spectroscopy (OES) (see more detailed comments below in methods selection). Causes for elevated As concentrations in the soil were considered with multiple working hypotheses, which included (1) a human source or (2) a nat-

ural source, or (3) false positive results. Common anthropogenic sources of As in the Georgia Piedmont include pesticide applications associated with control of the cotton boll weevil around the 1920's (Williams and others, 2005; Haney and others, 2009) and herbicide control for cotton growing, which occurred into the 1990's (Bednar and others, 2002). Historical aerial photography of the study site indicates it was utilized for agricultural purposes as far back as the 1930's, as evidenced by the plowing contours seen in Figure 1 (approximate center of left photo). The area has since rapidly developed into industrial parks and is now proximal to residential areas and major rail and road transportation lines (Figure 1, approximate center of right photo). Arsenopyrite has been reported in association with gold mining and tailings (EPA Report, 2003), the latter of which has been known to occur in north Georgia. If such tailings were transported to the site and dumped, then this represents a possible As and Se source.

Quadrangle-scale geologic mapping for this area indicates the study site is situated on lower Paleozoic metamorphic rocks with SW-NE trending structures (Higgins, 1968). These rocks have experienced numerous cycles of intrusion, folding, faulting, and deformation with the most recent influences of metamorphism associated with final stage collisions of Eurasian and North American Plates (200 to 230 Ma) and subsequent mafic dike emplacement related to cratonic rifting. Figure 1 shows approximate locations of formation contacts in the study area, as well as the trace of the Brevard Zone. The Brevard Zone is a narrow SW-NE trending feature with numerous interpretations and likely owes its origins to multiple mechanisms of thrust faulting, strike-slip faulting, overturned folding, and the development of complex features associated with prograde and retrograde metamorphism (Crawford and Kath, 2001). The study site (Figure 1) is mapped as "Button schist", a term noted by Higgins (1966) for the pattern of two distinct subparallel cleavages that appear upon weathering. These rocks occur throughout the Brevard Zone and are described as having crystalline texture resulting from met-

amorphic recrystallization under conditions of high viscosity, directed pressure, and some recrystallized after granulation (Reed, 1970).

METHODS

Impetus for study was based on prior surface sampling of soils (<5 m) and using OES analysis via EPA method 6010C (Jones and others, 1987; Hassan and Loux, 1990), with some As concentrations reported above the regional baseline value of ~7 ppm. Reported OES-determined As concentrations ranged from 4 to 375 ppm. A two-point drilling program was conducted to continuously core from the surface through the saprolite and into the rock basement. Two hand-auger samples from approximately the same locations were also taken to a depth of approximately 1.5 m below the surface. The auger samples were immediately tested using Induction Coupled Plasma Mass Spectrometry (ICP-MS) because of the potential for false positive As testing by ICP-OES in the presence of Al-rich materials (Jones and others, 1987) and because the regional bedrock and soils are known to be aluminous (Higgins, 1966). ICP-MS technology is capable of simultaneous determination of up to 80 elements in a liquid sample in a single run of a few minutes. The mass-selective detector is extremely sensitive, particularly for heavier elements, giving very low, down to parts per trillion, detection limits. The detector system is also relatively immune to many of the chemical and spectroscopic interferences that plague ICP-OES systems (Hassan and Loux, 1990). Seventy-three elements were determined on individually extracted samples using a proprietary Perkin-Elmer Elan 6000 ICP-MS semi-quantitative element scan called "Total scanQuant." The instrument was also calibrated for As, Se and several other metals of environmental interest (including EPA regulated metals) and run using a proprietary calibrated technique called "fullQuant".

Cores were sampled approximately every meter (Table 1) and hand ground for ICP-MS and X-ray diffraction powder analysis (XRD) using a zirconium mortar and pestle with alcohol as a grinding agent. For soils, sediments,

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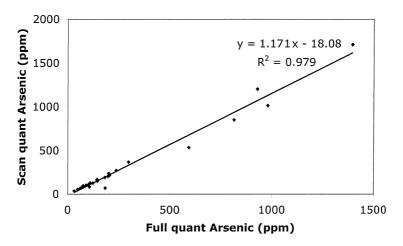


Figure 2. Quantitative analysis of As concentration in cores #1 and #2 sampled in Gwinnett County, GA. Axes depict duplicate analyses using two different proprietary quantification schemes. Line represents least-squares minimization fit to the data assuming a linear model. Cross plot using Se values produces similar coefficients.

rocks and organics, a boiling nitric acid digest was used to extract metals (EPA method 3050A). This involves digestion of the sample in concentrated nitric acid on a hot plate, followed by hydrogen peroxide to further oxidize organics and solubilize metals (Edgell, 1988). This works less well for soils and rock material and very well for highly organic samples (e.g., plant tissue). The concentrations reported herein therefore, may reflect slightly varied values relative to the total mass of the soil or rock because less soluble silica and aluminum oxide phases may not be totally dissolved. The hand auger near-surface soil samples measured by ICP-MS showed similar arsenic values to those measured by ICP-OES. It was therefore concluded that the measured arsenic concentrations using ICP-OES are accurate and grounds for nullifying the false positive hypothesis is justified.

The powdered bedrock material for XRD analysis was transferred to 30 x 30 mm mount and pressed to minimize transparency and preferred orientation. Data were collected using a Scintag diffractometer, with Co K α radiation, a 250 mm goniometer circle, $2^{\circ}/4^{\circ}$ primary and scattering slits, $0.5^{\circ}/0.3^{\circ}$ scattering and receiving slits, 40 kV and 35 mA, a step size of 0.01°, and a scan rate of 2° per minute. Selected pieces

of the bedrock schist from the unweathered portion of the core were cut, polished, and prepared for electron microprobe analysis (EMPA) using wavelength dispersive spectroscopy (WDS) and backscatter secondary electron imaging (BSE). A representative thick section was cut perpendicular to foliation and fractures in the schist. The polished mount was carbon coated for EMPA-WDS using a JEOL JXA-8600 Superprobe. Beam current was 15 nA and accelerating voltage was 15 keV. Grains were analyzed for Fe, As, and, Se using natural and synthetic mineral standards.

RESULTS

Coring resulted in the retrieval of 23 and 25m of continuous sampling for cores #1 and #2, respectively (Table 1). Saprolite thicknesses for each core are approximately 20m and 22m, respectively and approximately 3m of solid bedrock core were recovered from each drill hole. Foliations of the core are oriented with a dip of about 25°, which is slightly less than the regional dips of 36°-40° SE, published by Higgins and Crawford (2007). Table 1 includes depths sampled for ICP-MS analysis, with the solid line indicating sampling above and below the bedrock-saprolite interface.

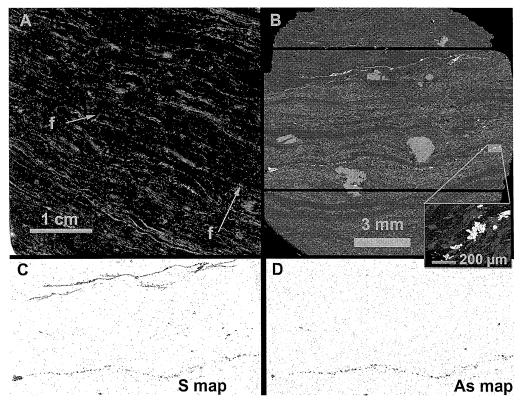


Figure 3. Micrographic images of schist from core #2 (depth 21.8 m). A) Reflective light image of polished core section cut perpendicular to dipping foliations. With exception to occasional cross-cutting brittle fractures (f), all primary mineral fabrics appear to follow the foliation, B) BSE image reveals electron dense (brighter) minerals following foliation trends. Horizontal lines depict mapped areas shown below. Inset shows higher magnification view of arsenopyrite grains. C) Sulfur WDS map of same area shown above in BSE. Two dark bands (one in the top and other in bottom) highlight concentrations of pyrite (top) and arsenopyrite (bottom), D) Arsenic WDS map of same area shown above in BSE, but only lower dark band highlights concentration of arsenopyrite. Note As blebs follow foliations, while S follows both blebs in foliations and pyrite in fractures.

ICP-MS duplicate analyses for As using the different quantitative approaches reveals values are reproducible over the range from 37 to 1710 ppm and having near unity correlation (Figure 2). A plot using Se scanQuant and fullQuant values (not shown) reveals linear trend coefficients similar to As, with the main difference being the concentration range, which varied from 0.03 to 6.12 ppm for Se. Semi-quantitative scanQuant ICP-MS results in Table 1 are reported to three significant figures. The focus of this study is the source and relative fate of As and Se, therefore discussion of other elemental trends is beyond the study scope, however data are provided for future workers.

XRD and optical analyses of the bedrock schist from the core hole bottom reveals a mafic minerals assemblage, with quartz, biotite, muscovite, garnet, clinochlore, albite, laumontite, and pyrite occurring as major phases (Figures 3 and 4). Foliation and fracture fabrics dominate the rock's texture. Included in the foliations are most all the above mineral grains, which appear to follow slight undulating mm-scale waves (Figures 3A and 3B). Cross-cutting brittle fractures comprises a fabric that is often filled with pyrite and laumontite. These fracture-fills clearly post-date the foliation and are indicative of retrograde metamorphism of a fossil hydrothermal system (Deer and others, 2004). Sulfur and

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orted as Table 1. ICP-MS FullQuant results for elemental concentrations using EPA extraction method 3010A. Concentrations

Table 1. ICP-MS FullQuant results for elemental concentrations using EPA extraction method 3010A. ppm (mgkg).Solid line marks samples from above and below the saprolite/bedrock contact.	Quant res ne marks	resul rks sa	ts tor or makes	elen fo	nen m a	a c 3000	once an	sults for elemental concentrations using EPA extraction metho samples from above and below the saprolite/bedrock contact.	s using the sap	y EPA prolite	extr /bed	actior rock o	meth	ر بر بر	010A		cen	tratic	Concentrations are reported as	e reg	orte	g d as
Core 1																						
Depth (m) Sc Ti V	င်	Ē	£	ပိ	Z	ទ	Z	Ga Ge	As Se	e B	S _d	ş	>	Z	Q N	No	Pd	Ag C	Cd	S	Sb	
0.6 2.54 733 139	44.6	257	43200	13	27	22	115	15 0.14	70 0.92	2 0.52	119	4.5	43.9	14.6	0.07	0.19 0.	0.03 0.	0.02 0.	0.04 0.08	8 1.05	0.01	
1.8 1.35 555 102	42.4	732	48244	12	24	102	20	12 0.13	237 0.82	12 0.38	125	2.3	32.2	22.3	0.08	0.00	0.89 0.	0.04 0.	0.07 0.05	5 0.37	0.01	
3.1 1.13 470 79	38.4	699	42964	23	26	49	89	10 0.12	163 0.61	11 0.27	84	3.0	21.3	15.9	0.07	0.14 0.	0.01 0.	0.03 0.	0.07 0.04	4 0.45	0.00	
4.3 2.50 648 155	50.5	1446	45588	29	39	20	123	15 0.15	130 0.68	8 0.94	106	4.7	32.1	16.7	0.05	0.00	0.04 0.	0.05 0.	0.14 0.07	7 0.19	0.01	
5.8 2.68 732 122	62.8	334	39349	12	33	2	145	15 0.12	56 0.58	8 0.82	123	4.7	20.1	14.2	0.05	0.14 0	0.03 0.	0.02 0.	0.05 0.07	7 0.22	0.00	
7.6 0.30 234 37	20.5	67	26473	10	თ	4	40	4 0.07	97 0.51	1 0.46	196	3.3	14.0	12.6	0.11	0.38 0	0.01 0.	0.02 0.	0.08 0.01	1 0.11	0.00	_
8.5 0.81 534 66	30.4	298	36390	7	24	45	91	90.06	103 0.30	0 0.37	81	7.7	10.6	19.6	0.11	0.22 0	0.71 0.	0.04 0.	0.15 0.03	3 0.23	3 0.00	
10.1 0.82 509 86	40.1	424	46291	8	35	38	33	9 0.07	162 0.42	12 0.44	103	7.3	11.6	14.2	0.12	0.19 0.	0.44 0.	0.06 0.	0.23 0.03	3 0.23	3 0.01	
11.6 1.97 746 158	32.0	999	45844	24	20	=======================================	149	13 0.11	110 0.40	0 0.48	109	9.4	11.6	11.6	0.06	0.35 0.	0.49 0.	0.08 0.	0.17 0.07	7 0.40	0.00	_
13.7 2.20 482 164	40.4	1057	72604	. 22	34	4	180	13 0.25	271 0.94 0.49	4 0.48	84	10.3	31.3	11.3	0.05	0.61 0.	0.40 0.	0.13 0.	0.99 0.06	6 0.15	0.02	
15.3 1.20 546 118	47.8	282	48175	25	29	53	61	10 0.10	218 0.54	4 0.31	97	10.1	17.4	6.6	0.09	0.65 0	0.03 0.	0.07 0.	0.27 0.04	4 0.38	3 0.01	
19.2 0.77 412 88	39.7	98	33720	21	37	48	26	7 0.07	534 0.79	9 0.46	82	4.6	13.4	5.8	0.10	1.48 0.	0.25 0.	0.14 0.	0.06 0.03	3 0.46	0.02	
20.7 1.07 509 95	53.7	. 62	36521	31	41	68	38	10 0.04	1203 0.73	3 0.71	26	2.8	6.2	8.6	0.16	0.91 0.	0.01	0.28 0.	0.10 0.04	4 0.65	0.05	
21.4 0.30 143 29	40.0	∞	245774	20	177	132	2	7 0.14	86 3.1	3.13 0.84	o	44.0	13.6	11.6 0.21	0.21	1.16 0.	0.10 0.	0.24 1.	1.73 0.04	4 0.83	3 0.87	
21.8 0.80 69 9	8.5	48	37004	20	66	9	16	23 1.39	37 6.16	6 2.30	3	242.0	383.7	4.	0.01	0.48 0.	0.08 0.	0.12 0.	0.08 0.01	1 0.05	0.12	
Core 2																						
Depth (m) Sc Ti V	ပ်	Ē	æ	ပိ	Z	ಪ	ď	Ga Ge	As Se	Ä	S.	ķ	>	Z	Q N	Mo	Pd A	Ag C	Cd	S	Sp	
0.6 1.64 415 118	45.1	206	48436	15	35	85	20	12 0.08	86 0.79	9 0.62	31	2.6	9.3	9.2	0.21	0.58 0.	0.02 0.	0.13 0.	0.04 0.05	5 0.43	3 0.03	
	109.2	673	43796	21	22	48	55	14 0.08	152 0.91	11.28	40	9.0	27.6	13.9	0.12 (0.15 0.	0.47 0.	0.03 0.	0.06 0.07	7 0.60	0.01	
3.1 1.96 332 88	35.4	621	34396	α	თ	36	64	12 0.06	1014 0.60	0 0.36	44	0.3	35.5	21.2	0.11	0.00	0.81 0.	0.02 0.	0.10 0.06	6 0.29	0.01	
4.3 2.47 838 215	77.9	1094	59365	20	8	20	160	15 0.09	852 0.64	4 0.52	88	0.8	47.2	16.3	0.06	0.00 0.	0.01 0.	0.03 0.	0.24 0.10	0 0.37	0.01	
5.5 0.41 62 8	6.7	116	34848	7	25	20	58	3 0.04	72 0.64	4 0.51	12	6.4	10.3	8.2	0.05	0.00	0.03 0.	0.03 0.	0.30 0.02	2 0.26	3 0.23	
6.7 1.02 531 66	34.2	292	34264	7	7	51	1 4	9 0.08	368 0.39	9 0.30	123	/	15.7	14.0	0.14	0.31 0.	0.00 0.	0.02 0.0	0.09 0.03	3 0.65	0.02	
8.2 0.88 138 45	18.1	444	25501	4		7	181	9 0.04	170 0.52	2 0.23	27	2.4	21.8	16.1	0.03	0.06 0.	0.02 0.	0.02 0.	0.41 0.05	5 0.11	0.00	
9.8 0.43 17 62		/	40494	29	£	4	105	7 0.05	193 0.67	7 0.19	16	5.1	28.8	7.7	0.01	0.26 0.	0.27 0.	0.03 0.3	0.24 0.02	2 0.03	0.01	
11.9 1.32 131 116		620	48508	8	19	51	129	9 0.12	1712 0.71 0.54	1 0.54	56	13.6	46.3	12.1	0.02	0.00	0.06 0.	0.25 1.	1.76 0.03	3 0.08	3 0.02	
14.3 0.64 343 59			28994			%	153	7 0.04	206 0.34	4 0.25	20	6.4	9.6	13.4	0.13	0.16 0.	0.03 0.	0.07 0.	0.19 0.03	3 0.12	0.01	
16.5 1.87 530 115	46.9	671	36045	27	23	41	8	9 0.06	128 0.31	11 0.25	58	5.2	14.5	13.5	0.11	0.31 0.	0.03 0.	0.03 0.	0.14 0.04	4 0.29	0.01	
23.2 1.11 431 102	50.9	208	50735	23	20	118	9	11 0.08	89 0.92	2 0.39	38	3.4	8.7	5.6	0.17	49 0	03 0.	16 0.	5.6 0.17 0.49 0.03 0.16 0.05 0.04	4 0.51	0.03	

Table 1 continued. ICP-MS results for elemental concentrations using EPA extraction method 3010A. Concentrations are reported as ppm (mg/kg). Solid line marks samples from above and below the saprolite/bedrock contact.

Core 1												
Depth (m) Te	Cs Ba	Ë	ပ္ပံ	ፚ	N	Sm	Eu	Gd	Tb D	Dy Ho	Er Tm	Yb Lu Hf Ta W Hg Ti Pb Bi Th U
0.6 0.03 0.53	9.2 508	9.99	16.3	16.0	62.0	12.8	3.0 1	13.5	2.0 1	.3 2.1	1 5.7 0.8	4,9 0,7 0,7 0,02 0,08 0,00 0,85 12.8 0,19 6,65 0,70
1.8 0.03 0.28	8.4 212	43,4	41.6	11.3	45.0	9,5	2.1	10.0	3.	3.2 1.5	5 4.2 0.6	3.8 0.6 0.9 0.02 0.09 0.01 1.02 16.8 0.37 6.90 1.57
3.1 0.05 0.16	8.1 255	31.5	31.9	8.3	34.1	6.9	6.	7.2	0.1	5.7 1.0	2.9 0.4	2.6 0.4 0.6 0.01 1.32 0.03 0.80 21.9 0.37 8.27 1.39
4.3 0.01 0.23	4.7 526	54,4	105.6	14.7	61.6	12.5	2.7	12.2	3.	8.6 1.6	3 4.3 0.6	3.7 0.6 0.7 0.02 0.10 0.03 0.90 29.5 0.24 7.26 0.00
5.8 0.02 0.27	5.2 543	24.9	32.0	6.9	29.3	6.1	9.	6.1	0.9	4.8 0.9	9 2.5 0.3	2.1 0.3 0.6 0.01 0.08 0.01 0.82 8.2 0.15 6.94 0.61
7.6 0.03 0.06	4.0 148	40.1	82.2	10.7	43.3	8,4	1.7	8.1	1.1	5.0 0.8	3 1.9 0.3	1,7 0,3 0,6 0,01 0,06 0,00 1,11 8,9 0,37 6,56 1,60
8.5 0.06 0.03	4.4 311	14.0	33.1	4.3	17.5	3.6	0.7	3.6	0.5	2.8 0.5	5 1.5 0.2	1,4 0,2 0,8 0,01 0,05 0,01 0,64 13,7 0,35 8,41 0,00
10.1 0.04 0.08	4.3 297	13.9	37.4	4.8	19.5	4.2	1.0	3.9	9.0	3.4 0.7	7 1.8 0.3	1.8 0.3 0.6 0.01 0.06 0.03 0.68 17.6 0.19 11.28 2.19
11.6 0.04 0.04	5.9 500	32.1	75.3	9.0	36.3	7.7	£.	6.9	0.8	3.8 0.6	3 1.4 0.2	1.1 0.2 0.5 0.01 0.14 0.02 0.78 14.4 0.36 5.90 0.00
13.7 0.05 0.07 10.6 444	10.6 444	174.8	320.4	34.1	138.7	27.6	4.7 2	27.2	2.9 1	11.3 1.	7 4.0 0.5	3.3 0.5 0.5 0.02 1.31 0.03 0.52 75.0 0.54 5.48
15.3 0.04 0.04	5.7 264	76.4	137.7	15.5	63.2	12.1	2.3	11.8	1,5	6.2 0.9	9 2.0 0.2	1.4 0.2 0.4 0.01 0.40 0.03 0.73 20.1 0.75 9.04 0.00
19.2 0.09 0.01	4.0 165	51.3	114.3	13.7	55.2	10.6	2.1	10.1	1.2	5.3 0.7	7 1.5 0.2	0.9 0.1 0.3 0.01 2.21 0.02 0.63 9.9 1.09 10.93 2.84
20.7 0.16 0.00	4.2 240	15.2	34.3	4.2	16.8	3.4	0.7	3.2	0.4	2.2 0.3	3 0.8 0.1	0.5 0.1 0.4 0.01 0.12 0.04 0.76 9.1 1.81 11.23 2.61
21.4 0.00 0.01 4.9 26	4.9 26	62.4	133.9	16.1	67.3	13.5	2.8	13.6	1.7 (6.0 9.9	3 1.8 0.2	1.3 0.2 0.5 0.00 5.09 0.27 0.65 20.7 0.60 43.78 1.40
21.8 0.01 0.00 18.2	18.2 50	2218.8 4	4963.0 8	589.4 2	2307.9	348.7 6	68.5 32	322.1 3	36.0 13	37.8 18.3	3 37.4 4.1	25.4 3.0 1.2 0.13 0.30 0.13 0.58 18.6 0.17 37.79 15.70
Core 2												
Depth (m) Te	Cs Ba	La	లి	ď	Š	Sm	Eu	. pg	To of	Dy Ho	Er Tm	Yb Lu Hf Ta W Hg Ti Pb Bi Th
0.6 0.03 0.42	1.9 87	10.4	36.0	3.0	12.7	2.8	9.0		0.5	2.6 0.5	5 1.3 0.2	1.0 0.1 0.4 0.01 0.34 0.02 0.50 16.0 0.34 8.90 1.16
1.8 0.04 1.24	3.0 255	40.3	69.4	12.6	53.7	7.	2.4	10.0	4.	7.8 1.	5 4.0 0.6	3.8 0.5 0.7 0.02 0.69 0.16 0.46 34.0 0.25 12.74
3.1 0.34 0.94	2.1 243	40.2	41.1	10.8	43.9	6.6	2.2		1.6	9.6 1.8	3 5.3 0.8	4.9 0.7 1.0 0.01 0.09 0.01 0.34 28.3 0.78 9.58
4.3 0.05 0.80	4.6 388	45.0	44.3	12.3	51.6	11.8	2.9	12.9	2.1 1.	12.7 2.4	4 6.8 0.9	6.0 0.9 0.7 0.02 0.18 0.02 0.61 33.3 0.57 8.57 1.84
5.5 0.01 0.06	1.2 50	18.6	35.3	5.1	20.9	4.5	0.9	4.5		3.0 0.5	5 1.3 0.2	1.1 0.2 0.3 0.01 0.92 0.03 0.15 10.0 0.15 8.10 0.00
6.7 0.10 0.16	6.0 203	21.5	21.3	6.1	26.0	5.7	<u>4</u> 5	5.3	0.8	4.5 0.8	3 2.3 0.3	2.0 0.3 0.6 0.01 0.11 0.02 0.66 22.7 0.47 8.49
8.2 0.02 0.10	0.9 120	28.9	28.8	7.9	33.0	7.1	1.5	7.2		6.3 1.2	2 3.1 0.4	2.6 0.4 0.7 0.01 0.05 0.02 0.16 65.6 0.45 6.21 3.35
9.8 0.05 0.23	0.7 156	39.4	163.6	10.1	40.1	8.0	ا 6	8.9	1.3	7.7 1	4 3.8 0.5	3.1 0.4 0.3 0.01 0.34 0.03 0.49 56.3 0.26 6.59
11.9 0.09 0.08	1.2 267	108.8	162.8	27.5	106.5	21.5	4.0	19.2	2.7 14	14.6 2.6	5 6.8 0.9	5.5 0.7 0.6 0.03 0.08 0.02 0.14 60.3 0.82 7.12
14.3 0.03 0.02	2.1 302	10.7	29.2	3.6	15.6	3.6	0.7	3.6	0.5	3.0 0.6	3 1.5 0.2	1.3 0.2 0.6 0.01 0.04 0.00 0.32 28.0 0.27 8.46 1
16.5 0.00 0.01	4.5 452	29.4	70.0	8.6	36.1	7.8	1.4	7.4	1.0	5.1 0.8	3 2.0 0.3	1.5 0.2 0.7 0.01 0.51 0.02 0.45 8.0 0.16 7.71
23.2 0.04 0.00	2.2 105	10.9	24.4	3.0	12.3	2.7	9.0	3.1	0.4	2.6 0.4	1.1 0.1	0.8 0.1 0.2 0.00 0.16 0.02 0.64 6.7 0.29 9.48 1.17

pears to be mobile and subject to transport away from the weathering zone.

The behavior of naturally occurring As in arsenopyrite and its oxidized forms can be understood in terms of location within the bedrock/ saprolite/soil profile and available electron acceptors in the weathering profile. The oxidation from As(III) in arsenopyrite to As(V) is a wellstudied process with associations to microbially mediated lithotrophy during weathering (Ehrlich, 1964; Strawn and others, 2002; Yu Yunmei and others, 2004). An important consequence of this oxidation reaction is the generation of acidity and the production of arsenate complexes (Walker and others, 2005). The latter of which behave much like soil phosphate complexes that are know to associate with Feoxide surfaces. (Qafoku and others, 1999; Filippi and others, 2007; Neel and others, 2003). Given the abundance of hematite and goethite in these piedmont soils, the retention of As relative to bedrock concentrations in this study can be parsimoniously explained as a natural occurrence related to the original bedrock. The heterogeneous vertical distribution As and its concordance of mass balance between bedrock and weathered horizons are in contrast to known studies of anthropogenic As occurrences. For example, Kukier and others (2001) amended similar Georgia Piedmont Cecil soils with As-bearing fly ash. In this anthropogenic case, a signature distribution of As is characterized by a high near the surface followed by a gradational decrease at depth to an order of magnitude lower level.

The behavior of naturally occurring Se, as it is released during arsenopyrite oxidation, is not as well documented as As. Se in arsenopyrite most likely occupies sulfur sites in arsenopyrite as a reduced selenide form. Upon exposure to electron acceptors such as oxygenated ground waters, the Se reacts to form the oxidized states of Se^(IV) and Se^(VI), with the latter selenate form being more common in weathering environments (Strawn and others, 2002). Comparison of As and Se ionic potentials [IP = defined as the ratio of ionic charge (Z) to ionic radius (r)] predicts their solubility behavior to be different (Pauling, 1948). Assuming the respective

values of $r_{As(V)}$ = 0.46Å and $r_{Se(VI)}$ = 0.42, then respectively $IP_{As(V)}$ = 10.9 and $IP_{Se(VI)}$ = 14.2. These IP values indicate that both As and Se should complex as soluble radicals. The greater IP value of Se^(VI) indicates it is more soluble than As, which is consistent with the observations of this study. Studies of both As and Se for the Georgia Piedmont are not extensive, however Peltier and others (2008) have indirectly evaluated heavy metal abundance in watersheds by using the Asiatic clam (Corbicula fluminea) as a biomonitor. The studies specific interest is the contributions of trace elements associated with different point sources and land uses in a large river. In particular they studied the tributaries of the Chattahoochee River, whose main channel in the Georgia Piedmont is largely controlled by the trace of the Brevard Zone. Realizing that the Chattahoochee integrates a watershed larger than the Brevard Zone, analysis of As and Se in nine samples from each of fifteen river sites reveals an As/Se ratio of 0.6, which is much smaller than the ratios of about ~40 for the bedrock and ~2000 for the saprolite at the study site. If the primary As and Se signature of the Georgia Piedmont is controlled by arsenopyrite similar to that in this study area, then both the saprolite and the river waters (as proxied by a biomonitor of Peltier and others, (2008)) provide a good model for the partitioning of the two elements in the Piedmont of SE United States.

CONCLUSIONS

Pyrite and arsenopyrite occur in a mafic schist of the Piedmont of Gwinnett County, Georgia associated with a fossil hydrothermal system and the Brevard Zone. The pyrite is texturally associated with late stage retrograde metamorphic laumontite, which occurs along brittle fractures and within foliations. Arsenopyrite appears as discreet grains within and associated with foliations formed or retained during prograde metamorphism. Analysis of the As and Se in the bedrock shows they reside in the arsenopyrite resulting in bulk rock concentrations above regional baseline values. Analysis of the As and Se in overlying saprolite and

soil indicates they are weathered from the arsenopyrite but have different fates. Arsenic is conserved most likely as arsenate complexes adsorbed to abundant Fe-oxyhydroxide surfaces in the saprolite/soil. Selenium is transported out of the weathering profile and presumed to be carried off by the rivers. This scenario is directly supported by the high As/Se ratios measured in the saprolite of this study and indirectly by the very low As/Se ratios measured in biomonitor proxies of river waters draining the Brevard Zone (Peltier and others, 2008). Ascribing high concentrations of As and Se to anthropogenic, natural, or false-positive factors requires analysis of bedrock, saprolite, soil, and methodology (in aluminous terrains). In cases where As concentrations are high at the surface and decrease to regional baseline values, then an anthropogenic cause might be ascribed. In the case of this study, where there are no discrete vertical trends and mass balancing of underlying bedrock can account for saprolite concentrations, naturally elevated As levels are possible on small spatial scales in the Piedmont soils of the SE United States.

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