

Norwegian University  
of Life Sciences

**Masters Thesis 2024 30 ECTS**

Faculty of Science and Technology

# **A comparative study of soil temperature models, including machine learning models**

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Master of Science in Data Science

## Forword

Jeg vil takke alle som var rundt meg under skrivingen, for uten dere så ville jeg ikke ha klart å fullføre.

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## Acronyms

K | L | M | N | S

**K**

**Kilden**

Norwegian Institute of Bioeconomy Research Kilden. 6

**L**

**LMT**

Norwegian Institute of Bioeconomy Research LandbruksMeteorologisk service. 6, 9

**M**

**MET**

The Norwegian Meteorological Institute. 6, 9

**MSTL**

Multiple Seasonal-Trend decomposition using LOESS. 9

**N**

**NMBU**

The Norwegian University of Life Science. 12

**S**

**STL**

Seasonal-Trend decomposition using LOESS. 9

### Abstract

This study focuses on 3 models that have been used in the literature to predict soil temperatures. The depths chosen as targets are 10cm, and 20cm in 4 regions; Innlandet, Østfold, Vestfold, and Trøndelag. In each region there are 4 stations with self-draining or saturated soils.

### Oppsummering

Denne studien ser på modeller som predikerer jordtemperaturer

*Keywords:* Soil temperature, Machine learning, regression, hour, ...

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## 1 Introduction

In agriculture soil temperature is one of the important parameters to put into consideration when thinking about pest prevention, conservation, and yield prediction. The reasoning for this is that knowing the soil temperature is knowing climate change [1], water management [2], yield [3], nitrogen processes [4] in the soil, calculation of plant-growth [5], when seeds start to sprout [5], potential flooding and erosions[6], and predicting when insect eggs hatch that were laid last winter. Being able to predict the soil temperature into the future will be a huge advantage for farmers, civilians, and scientists.

If it's important, why don't institutions measure it everywhere? There are several reasons for this, but a common reason is that it's expensive to install new equipment on old weather stations. Sometimes the weather station do have the sensors in the fields reading soil temperature at given levels, but due to technical misadventures and unforeseen phenomenons there might be gaps or misreadings that need to be replaced with approximations or NULL values<sup>1</sup>. There are algorithms, models, and statistical tools to interpolate these missing values but they have their drawbacks. For instance approximation by global mean, which is a common method used in timeseries[7]. This method is preserved global statistics, however does not represent local changes. Further more for a good estimation of soil temperature it is useful to include exogenous<sup>2</sup> features.

There has been done research into heat conductivity in soil that has lead to differential equations[8], however these equations[8, 9] are computationally expensive and difficult to simulate, or calculate[4]. To add to the complexity the heat dynamics change depending on soil temperature as they change the physical stucture of the soil

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There are also massive developments in types of models, one of which are ASPER[10] that combines logical statments with deep learning models to aceve better or similar results to "non-logical" deep learning models, but on fewer samples. A study has been preformed with the model which shows that the model can reduce the number of samples/observations needed by a factor of 1/1000[11]. After an intervju with the study resercher (Machot), although the model needs a strict ruleset it is possible to incorporate baysioan statitics to make the model more geneneral for more applications by weaking the ruleset and imply that the rulse given might not be 100% accurate and can be relaxed.

Furter more there are investigation into introducing randomness to the model to improve prediction. As an example is a study into Brownian fractional motion to predict temperature fluxiations at the Campi Flegrei Caldera area.

<sup>1</sup>These values are different from 0 as they represent "no data" and can't be used to do calculations.

<sup>2</sup>Variable that can affect the model, but is not not directly described by the model.

In this study 4 methods will be compared and evaluated for the sake of further research into interpolation of missing data in northic countries based on as few features as possible. This study has chosen 2 types of models; Analytical, and Data-Driven models. There will also be base models to compare against, one for each model type.

## 2 Theory

This section discusses the theory behind the models used in the study, there are a few base models (section 2.1, and 2.4) based on the interpretation of the papers hand picked for this study

### 2.1 Linear regression

The regression model will be for the sake of convenience be expressed as the following expression

$$\left(\vec{F}(\mathbf{A})\right) \vec{\beta} = \vec{y} + \vec{\varepsilon}$$

Where  $\vec{F}$  is a vector function with following domain  $\vec{F} : \mathbb{R}^{m \times n} \rightarrow \mathbb{R}^{m \times p}$  where  $m, n, p \in \mathbb{N}$ ,  $\mathbf{A}$  is the data in matrix form with dimensions  $\mathbb{R}^{m \times n}$ ,  $\vec{\beta}$  is the regression terms,  $\vec{y}$  is the target (TJM10 or TJM20), and  $\vec{\varepsilon}$  is the residual error.

This basic model to express the linearity of the components to soil temperature. This will function as the base model for regression models.

The  $\vec{F}$  is not important, just that your data is shaped by a function.

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### 2.2 Plauborg linear regression model with Fourier terms

An improvement over an time independent linear regression model would be a time dependent linear regression model that takes not only current time into account of the calculations but also previous measurements. It is current knowledge that soil temperatures depends on previous temperatures and metrological phenomena. In the paper “Simple model for 10 cm soil temperature in different soils with short grass” the author chose to extend the features from only air temperature at current time to include also previous days of year and the air temperature from those days. This means the following F function that Plauborg used would be

$$\vec{F} := [air_t, air_{t-1}, air_{t-2}, air_{t-3}, \sin(\omega t), \cos(\omega t), \sin(2 * \omega t), \cos(2 * \omega t)]^T$$

Where  $air_t$  is the air temperature at time  $t$  expressed in day of the year (0-365),  $\omega$  is the angular frequency in radians per hour or radians per day, depending on the time unit. The sine/cosine elements in the F function represent the variations through the day by fitting  $\vec{\beta}$  to the yearly variation. To adapt the authors model to an hourly time unit would be to either

1. Extend the F function to include a larger  $\omega$  coefficient to reflect hourly oscillations in conjunction with daily fluctuations
2. Refit the Fourier terms with a larger  $\omega$  coefficient to make the oscillations more representative of daily temperature changes.

The larger coefficient could be expressed as  $\pi/12$  while the smaller  $\omega$  for daily values would be rescaled to  $2\pi/365$ .

The problem with this approach would be Fourier's Sine-Cosine series approximation which would suggest that Plauborg's method could be subject to overfitting with addition of more terms. On the other hand it gives us a way to compute the coefficients  $\alpha_i$  and  $\gamma_i$  for sine and cosine terms respectively, though it would be more numerically stable with a pseudo-inverse computation or a max log likelihood approach.



### 2.3 Rankin's finite difference method of simplified heat flow in snow covered soil

A more direct method based on laws of physics develop by Karvonen involves forming a Finite Difference Method (FDM) around point of interest with simplifications to the equations described in *A model for predicting the effect of drainage on soil moisture, soil temperature and crop yield*. A team of researchers collaborating with the original author found an algorithm by making simplifications to the general differential equations forming a iterative 2-step procedure seen at the procedure 1.

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#### Algorithm 1: Rankin algorithm

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**Data:**  $D, f_d$   
**Result:**  $T_Z$   
1  $\alpha_t \leftarrow \frac{\partial T / \partial t}{\partial^2 T / \partial z^2}$ ;  
2 **for**  $t \in T$  **do**  
3      $T_*^{t+1} \leftarrow T_Z^t + \Delta t \times \frac{\alpha_t}{(2Z)^2} \times (T_{air}^t - T_Z^t)$ ;  
4      $T_Z^{t+1} \leftarrow T_*^{t+1} * e^{-f_d \times D}$ ;  
5 **end for**

---

Where  $\alpha_t = K_T / C_A$  is the Thermal diffusivity from Fourier's law in thermodynamics,  $K_T$  is average soil thermal conductivity,  $C_A$  is the apparent heat capacity, and  $f_d$  is the damping parameter that has to be empirically derived however for this study it will be estimated from the data through the following estimation

$$f_d \approx \frac{-\ln \left( \frac{T_Z^{t+1}}{T_Z^t + \Delta t \frac{\alpha_t}{(2Z)^2} (T_{air}^t - T_Z^t)} \right)}{2D}$$

The approximation used in the algorithmn 1 assumes that  $K_T$  is not dependend on depth . To make the approximation of  $\alpha_t$  more accurate the inclusion of rain ( $\theta$ ) to introduce variation can be approximated with

$$\alpha_t \approx \frac{b_1 + b_2\theta + b_3\sqrt{\theta}}{a_1 + a_2\theta}$$

proposed by Kodešová *et al.*[13]<sup>3</sup>. To make the computation easier of this Padé-Puiseux<sup>4</sup> approximation hybrid we will realize that  $\alpha_t$  is expressed by

$$\frac{b_1 + b_2\theta + b_3\sqrt{\theta}}{a_1 + a_2\theta} \approx \alpha_t \approx \frac{(T_z^{t+1} - T_{air}) * (2z)^2}{(T_{air} - T_z^t) * \Delta t}$$

Thereby only needing a linear regression of two F-functions;  $F_1 = [1, \theta, \sqrt{\theta}]^T$  and  $F_2 = [1, \theta]^T$  rather than a three step approximation. This algorithm (algorithm 1) will approximate the following integral

---

<sup>3</sup>This representation was not proposed by the author however the linear approximations was proposed to approximate  $K_T$  and  $C_A$  respectfully. Since  $\theta \propto m_w$  we can substitute water content with rain in mm since the area is constant and during all messurement the soil type will be the same, however this would need to be reestimated if a station contains a different soil type as the constant has a wide range of values[13].

<sup>4</sup>Padé Approximation is a of the form  $\frac{\sum_{i=0}^{\infty} c_i x^i}{\sum_{j=0}^{\infty} c_j x^j}$  and a Puiseux series is a  $\sum_{j=N}^{\infty} c_j x^{j/N}$

$$T = \int_{t_0}^{t_{max}} \frac{K_T}{C_A} \frac{\partial^2 T}{\partial z^2} dt$$

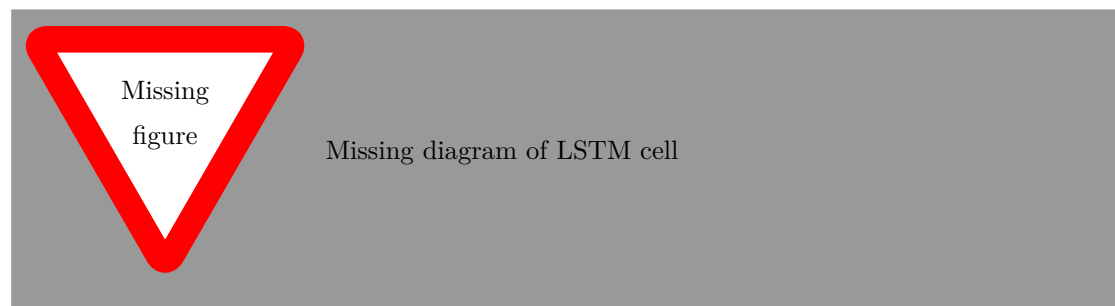
via a Finite Difference Method, although other methods are possible with higher accuracy<sup>5</sup>. **Must verify for this case!** This study will use the FDM used by the author for the purpose of making the results in this study comparable with the study presented in the paper “A simple model for predicting soil temperature in snow-covered and seasonally frozen soil.”

For initial values this study are utilizing 2 methods under different assumptions:

$$T_z^0 \approx \frac{k \exp(D)}{1 + \exp(D) \times (k - 1)} \times T_{air}$$

Where  $k$  is  $K_T * \Delta t / (C_A * (2Z)^2)$ , and  $D$  is  $-f_d * Snow_{Depth}$ . This assumes constant air temperature above a constant layer of snow, though unrealistic since air temperature has a tendency to change during the day due to solar radiation and other climate factors that can cool down or heat up the air. Another problem is the fact that the snow level remains the same which is also untrue.

## 2.4 Long Short Term Memory model



When modeling soil temperature it is important to know the previous hours or days to predict the next timestep, for this a natural selection for a data driven model is a recurrent network. This type of network makes prediction based on previous timesteps in the data, however the longer timespan the model takes into account the less important are the earlier timesteps in the data. To combat this there was developed an improved model called Long-Short Term Memory model[15] that deploys a memory cell that feeds information from earlier timesteps to the later ones. To make sure that redundant information or unimportant information doesn't get feed forward there are also forgetting gates that removes some of the newly learned patterns and integrates it into the memory cell.

The most common problem in Neural networks is the vanishing gradient problem where updating the first few layers of a large network becomes exponentially more difficult since the adjustments get smaller and smaller for each layer towards the start rather than the reverse. Long Short Term-Memory changes this by carrying information from the previous cells forward thereby allowing updating earlier cells with bigger impact than the standard approach[15]. LSTM is part of a family of Recurrent Neural Networks that passes information to other cells in the same layer.

<sup>5</sup>For example fourth degree Runge-Kutta method[14] which converges quicker than forward-Euler method or FDM.

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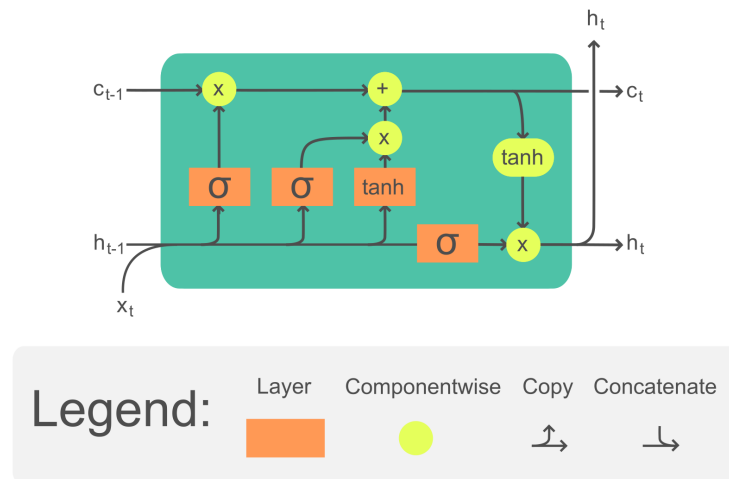


Figure 1: LSTM cell Artist: Chevalier [16]

## 2.5 Attention aware LSTM model

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### 3 Method

#### 3.1 Source of data

For this comparative study the following data sources will be used

1. LMT
2. Xgeo
3. Kilden
4. MET

#### 3.2 Dataset

The dataset is chosen from four regions in Norway; Innlandet, Vestfold, Trøndelag, and Østfold. From each region are four stations picked:

Region	Name	ID	Drain type	Soile category	Texture	MET name	Latitude	Longdetude
Innlandet	Apelsvoll	11	Selvdrenert	CM	17	SN11500	60,70024	10,86952
Innlandet	Fåvang	17	Selvdrenert	CM	15	SN13150	61,45822	10,1872
Innlandet	Ilseeng	26	Selvdrenert	PH	17	SN12180	60,80264	11,20298
Innlandet	Kise	27	Vannmettet	GL	99	SN12550	60,77324	10,80569
Trøndelag	Kvithamar	57	Vannmettet	ST	16	SN69150	63,48795	10,87994
Trøndelag	Frosta	15	Selvdrenert	LP	13	SN69655	63,56502	10,69298
Trøndelag	Mære	34	Selvdrenert	RG	14	SN71320	63,94244	11,42527
Trøndelag	Rissa	39	Vannmettet	PL	13	SN71320	63,58569	9,97007
Vestfold	Lier	30	Vannmettet	ST	16	SN19940	59,79084	10,25962
Vestfold	Sande	42	Vannmettet	ST	16	SN26990	59,6162	10,22339
Vestfold	Tjølling	50	Selvdrenert	AR	13	SN27780	59,04641	10,12513
Vestfold	Ramnes	38	Vannmettet	ST	16	SN27315	59,38081	10,2397
Østfold	Rakkestad	37	Vannmettet	ST	18	SN3290	59,38824	11,39042
Østfold	Rygge	41	Selvdrenert	AR	13	SN17380	59,39805	10,75427
Østfold	Tomb	52	Vannmettet	ST	16	SN17050	59,31893	10,81449
Østfold	Øsaker	118	Vannmettet	ST	18	SN3370	59,31936	11,04221

Table 1: Station information from stations used in this study. The texture class is defined in this article: <https://nibio.no/tema/jord/jordkartlegging/jordsmonnkart/dominerende-tekstur-i-overflatesjikt>

All stations are sampled from the date<sup>6</sup> 03-01 to 10-31 from 2016 to 2020. The features rain (RR), mean soil temperature at 10cm (TJM10), mean soil temperature at 20cm (TJM20), and air temperature at 2m (TM) are sampled from the LMT database. The snow parameter is sampled from MET via Xgeo for imputed values in areas where there are no measured values. The soil type, and soil texture is sampled from Kilden from Norwegian Institute of Bioeconomy Research.

##### 3.2.1 Selection process

The selection process for finding these station can be compiled into these steps

1. Recommendation from Norwegian Institute of Bioeconomy Research
2. Compute the missing values in the data
3. Missing values analyse

<sup>6</sup>Format month-day

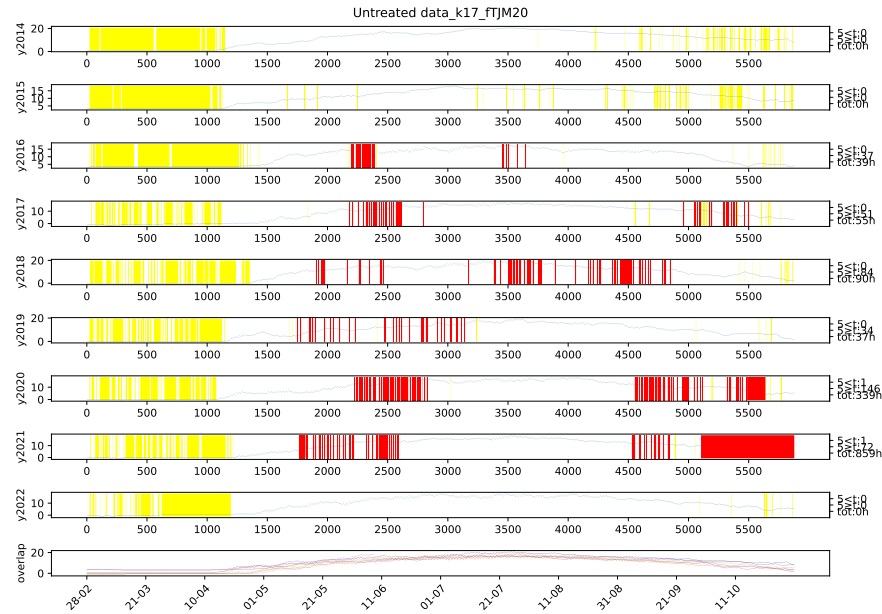


Figure 2: Visual representation of missing values at station 17 from 2014 to 2022 at the parameter "TJM20". The left numbers indicated how many hours that are missing and how many of them are shorter than or longer than 5 hours. The yellow markings indicate possible outliers based on the given year, all markings was checked if they were actual outliers. The red colouring indicate missing values in the data (represented in the data with code "NULL").

4. Searching LMT database for alternative station candidates if current data is insufficient
5. If some station was replaced the repeat step 2

The plots of stations follow a simple representation where the y-axis represent the year and the x-axis represent the index of the data as all tables are taken from the same period. A circle represent a singular na values, while a band represent a series of 2 or more missing values. The colours represents the features used in this comparative study. This representation of the missing values will indicate sesonal, and systematic removal of data and give an overall indication of how much data is missing. To get further insight into the data a report is generated in parallel to the plots describing precise date and time of all values and which other parameter values is also missing values in the same period. See appendix ?? for the full detail of the report generation and appendix A for na-plots of the station chosen for this study.

### 3.2.2 Collection of data

The method used was a powershell<sup>7</sup> script that called the respective institutions servers using the "curl" program<sup>8</sup> to send an http request for the timeseries starting from 2014 to 2020 in the interval 1 of May to 31 of October. Code for data collection can be viewed in appendix ?. The data is stores as an either a csv file or a json file for easy retrieval and manual control of values.

<sup>7</sup>Version 7.3.11

<sup>8</sup>curl 8.4.0 (Windows) libcurl/8.4.0 Schannel WinIDN

FROST	Description
Stations with rain	Requested rain data in millimeters.
Station ID	Sendt a request to LMT for station information using their remote API.
LMT	Description
Meteorological data	Requested soil temperature from 10cm depth, and 20cm depth and air temperature (2m), from 2014-03-01 to 2022-10-31.

Table 2: Description of what was requested from each server (FROST part of MET, LMT).

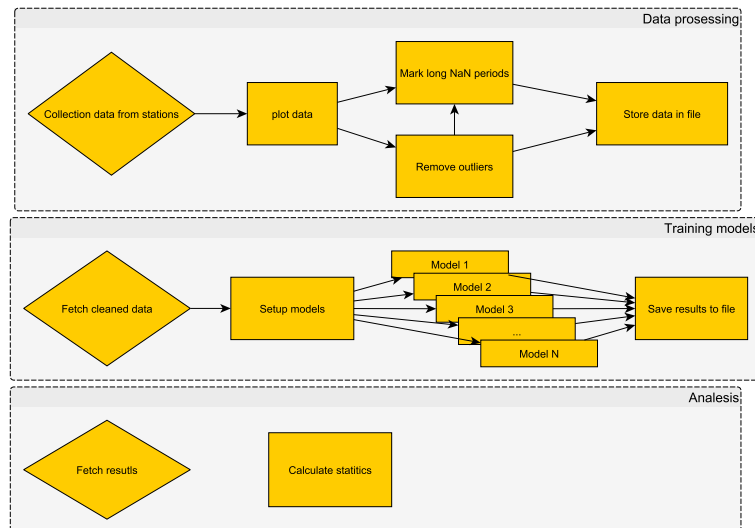


Figure 3: A surface level diagram of the

### 3.2.3 Labeling of stations between Nibio and MET

Since Nibio and MET have different names for the same stations one must compile a list that converts Nibio ID to MET ID. This was performed with these requests shown in table 2 where ID is the Nibio Id for the given station, Frost.ID is the MET id, ID. latitude is the latitude gathered from Nibio, ID.longitude is the longitude gathered from Nibio. These variables can be swapped out for the relevant station.

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caption

### 3.2.4 Storage of data

The storage of the data is done through two data structures; Hashmap and DataFrame from the package pandas. The transformation of data is done with a costume datatype called "DataFile-Handler" which is converted to a module for convenience. The keys for the hashmap is chosen by the naming of the data files and the pattern given to the class. To escalete modeling the data will also be exported to a binary file for faster retrieval.

MORE  
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diagram

**Technical overview of custom data structure** The data structure used to store the data from the different stations is called "DataFileHandler" and stores the data in a tree-structure where indexes are dictated by the filename. It has several built-in functions to assist with data partitioning, and merging of data. This makes it easier to move and store all 846 720 observations from 16 station from all 4 regions<sup>9</sup>. Further more, the data structure has iteration functions so it is compatible with python's built in loops, and print functions.

### 3.3 Data cleaning and treatment

To use the data in this study it must be cleaned and treated for training. The following methods were picked common practice in litterateur with new methods based on the decomposition of the data in the from of Seasonal-Trend decomposition using LOESS (STL)[17]<sup>10</sup>.

#### 3.3.1 Outlier detection and removal

Though the data fetched from LMT is treated and controlled the external data from MET might not be, and this research project incorporated raw, untreated data from LMT to fill inn missing values.

The method to quickly find obvious outliers was to look at the following condition

$$\left| \frac{|\Delta T| - (|\Delta T|)}{\text{var}(|\Delta T|)} \right| > 4$$

This condition looks at the absolute difference between consecutive measurements and calculates the z-score for each observation. It is expected that the change in temperature can't be too rapid. Further methods used to highlight potential outliers is

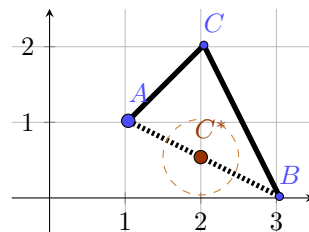


Figure 4: An simple outlier detection method utilizing a simple line to estimate where the expected point ( $C^*$ ) is supposed to be. If observed point C falls outside the tolerance level (red dotted circle) then it is marked as an outlier.

#### 3.3.2 Missing value imputation

The data has missing values, in particular during early Fall when there were sub-zero temperatures meaning any rain measurements done during this period would have unpredictable fluctuations since at negative temperatures water can freeze, get clogged up with residual bio-material from the surrounding area

<sup>9</sup>there are 4 stations per region.

<sup>10</sup>In this study we expand this for multiple seasons using Multiple Seasonal-Trend decomposition using LOESS (MSTL)[18], but the theory of this imputation method remains the same.

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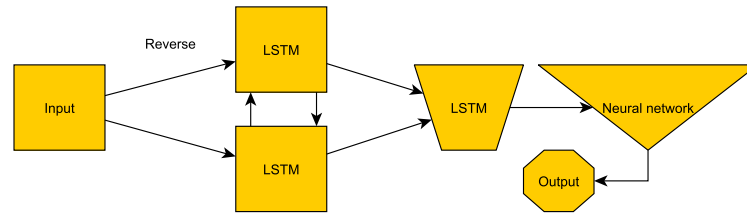


Figure 5

### 3.4 Setup of models

The models are set up in according to the relevant paper the model is fetched from, alternatively reuse the code made by the author. When importing the data to the model there will be modifying to the original code to facilitate for the model as far as it goes. Any modifications will be in the appendix under section ?? . For the convenience of the reader all code is using the sklearn estimator class to make all the models discusses in this study more user friendly and compatible with sklearn's other functions. The details of the models will be discussed in section 2, this section discusses the setup and implementation of the models.<sup>11</sup>

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#### 3.4.1 Basic Linear model

The linear model (sec 2.1) utilises in the study is created from the python model sklearn (or scikit-learn according to python's package manager)

#### 3.4.2 Plauborg

The Plauborg regression will be formulated as a linear regression problem so that the Linear-Regression function in the Sci-kit module can be used. For the parameters used in the paper[12] the F function defined in section 2.2 will be formulated with loops to give rise 3 more parameters for fine-tuning the model.

#### 3.4.3 BiLSTM

Soil temperatures are dependent on earlier timestamps meaning that to make a good prediction one needs to include temperature from  $t, t - 1, \dots, t - k$  to make a decent prediction. As a base model it is usefull to evaluate the data both forwards and backwards to find features that is only notisable in a sertain direction. The BiLSTM defined is crafted with Tensorflow's Keras module for ease of use.

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#### 3.4.4 Attention-aware ILSTM

It is a know fact, since 1849<sup>12</sup>, that to know previous weather patterns will greatly improve prediction accuracy. To improve the accurasy even more a model can focus one specific patterns that has a big impact on the prediction. The ILSTM attempts to do this by including a new thecnick in data science called attention[19] that takes a collection of data and gives each element a weight associated with importance. When that paper was published it was focused

<sup>11</sup>Caution to the reader; The code used was run on the Linux subsystem (Debian) on windows due to the fact that the current version of tensorflow can't run on Windows.

<sup>12</sup>First weather prediction made by Joseph Henry in 1849



on translation between English and German, however the paper published by Li *et al.* uses this novel teknick to do both time and feature importance and from that make a prediction.

### 3.5 Metrics

The metrics used in this study are

• Mean Squere Error • Mean Absolute Error • Explained Variance • bias • Log Condition number • digit sensitivity

Soil temperatur as a diffirenet behavior than air temperature since energy (temperature) though the soil gets dampen and delayed. Since the data used in this study has outliers that was not cought during datatreatment, which has been addressed, the author of this study desided to include two more metrics that are not usually included in the evaluation; The log condition number, and digit sensitivity. Both metrics are based on the calculation of the condition number defined as

$$\kappa = \lim_{\varepsilon \rightarrow 0^+} \sup_{|\partial x| \leq \varepsilon} \frac{|f(x + \partial x) - f(x)|}{|f(x)|} * \frac{|x|}{|\partial x|} \quad (1)$$

This is not feasible to calculate since infinite calculations with infinitesimal numbers is not possible as per April 29, 2024for simulation approach<sup>13</sup>. Therefore this paper uses algorithm 2 to approximate  $\kappa$  for all the models.

---

**Algorithm 2:** Method for calculating  $\kappa$ .  $\mathcal{U}$  is a uniform random distrebuton in a range.

---

**Data:** Data  
**Result:**  $\log(\kappa)$

```

1 Let  $\kappa_f$  be the function 1;
2  $\kappa \leftarrow 0$ ;
3 for  $i \in 1 \dots |Data|$  do
4    $\partial x \leftarrow \mathcal{U}_{[-\sqrt{\varepsilon/|Data|}, \sqrt{\varepsilon/|Data|}]}$ ;
5    $k \leftarrow$  calculate with  $\kappa_f$  from  $x$  and  $x + \partial x$ ;
6   if  $k > \kappa$  then
7      $\kappa \leftarrow k$ ;
8   end if
9 end for
10 return  $\kappa$ 
```

---

The digit sensitivity is included to give an intitve understanding of  $\kappa$  and is computed simply as  $\log_e(\kappa) + 1$ . This number tells us the significant digit generated from the model. If the number is less than 0 then its the ith digit after the decimal point.

For the rest of the metrics, they are defined as follows

- $RMSE = \sqrt{\frac{\sum (y_{pred} - y_{truth})^2}{n}}$
- $MAE = \frac{\sum |y_{pred} - y_{truth}|}{n}$
- $bias = \frac{\sum (y_{pred} - y_{truth})}{n}$
- $Explained\ variance = 1 - \frac{\sum (y_{pred} - y_{truth})^2}{\sum (y_{pred} - \bar{y})^2}$

---

<sup>13</sup>This calculation is possible for some models, for instance linear regression models when converted to the form  $A\vec{\beta} = \vec{y}$ .

Where  $\bar{y}$  is the mean of the target,  $y_{\text{pred}}$  is the predicted data, and  $y_{\text{truth}}$  is the observed soil temperature.

### 3.6 Use of Artificial Intelligence in this paper

In this paper there has been used Artificial Intelligence (AI), specifically Bing Chat / Copilot hosted by Microsoft Cooperation with special agreement with The Norwegian University of Life Science (NMBU), for the following purposes:

1. Formalising sentences and rephrasing sentences.
2. Spellchecking
3. Code generation of basic concepts and structures (tree traversal, template for generic classes)

It is important to emphasize that our engagement with AI have been actively curated and verified with known information. All code underwent rigorous manual inspection within a dedicated testing environment. Furthermore, no confidential or sensitive information was shared with the AI; our interactions focused solely on broad topics and general inquiries. To validate the accuracy of AI-generated responses, we cross-referenced them with established research papers and textbooks.



## 4 Results

### 4.1 Linear regression vs Plauborg

The result of the modelling (table 5 to 8) show that modelig soil temperature without the inclusion of time is an innefficient, and innaccurat method of predicting soil temperatures. The global measure for the linear regression has an average error of  $2.3C^{\circ} \pm 4.23C^{\circ}$  while the global measure of the Plauborg daily model has an average error of  $0.6C^{\circ} \pm 1.96C^{\circ}$ . Further more Plauborg has an hight  $R^2$  value indicating that it follows the changes in the soil better than just scaling the air temperature.

Summerize  
results

Make sec-  
tions

Include  
tables

### 4.2 Modification of Plauborg

IN this study the original model, that was trained for daily values was converted to predict hourly data to see if the same formulation could be used to make predictions. When comparing the results shown in table 4 and table 3 to their daily counterpart it shows similar values showing that the model proposed in Plauborg can be extended to hourly timeseries.

### 4.3 Deep learning models

If not got-  
ten to  
work by  
6.May,  
then  
delete. Try  
to do the  
stat (with-  
out model)  
to have re-  
sults.

## 5 Discussion

### 5.1 The Fall discrepancy

A phenomenon that arose during performance evaluations was that the linear models struggles with the Fall season. The difference graphs shows a clear over or under estimation that are larger than  $10\sigma$ . When investigating the coefficients to the model this discrepancy can be contributed to the intercept that during low temperature ( $< 5^\circ\text{C}$ ) giving either an over estimation or an under estimation. Further more when removing the calculation of the intercept the same phenomenon is still present.

Write a  
discussion

mention  
anomalies  
and reasons

Show table  
of parameters

### 5.2 Basis for results

The inclusion of previous temperatures gives an improved estimation, even on hourly basis. The coefficients for both daily and hourly are observed to be  $< 1$  making it a mean temperature and the fourier terms would estimate the function[20]

$$e^{z/D} \sin(\omega t - z/D + \phi)$$

Since the term  $\exp(z/D)$  is constant we would be estimating  $C \sin(\omega t - Q) = \sin(\omega t) \cos(Q) - \sin(Q) \cos(\omega t)$ , where  $Q$  is  $z/D - \phi$  and is constant. This will be extrapolated to a simple sum of sines and cosines as the model does. Together the Plauborg model would estimate

$$E_{\text{year}}(T) + e^{z/D} \sin(\omega t - z/D + \phi) \approx E_{\text{period}}(T) + \sum e^{z/D} \sin(Q_i) \cos(i\omega t_i) + \sum e^{z/D} \cos(Q_i) \sin(i\omega t_i)$$

A possible reason for the residuals could be the soil types at the stations making so a universal, high accuracy model would not be feasible unless including other metrological metrics (air pressure, humidity, soil type, soil texture, etc...) or including other non-linear features ( $\sqrt{\text{temperature}}$ , 2D temperature change in depth and time, etc...).

### 5.3 Future work

The models chosen in this study is not a representative sample of current knowledge of soil temperature modelling, and this study did not aim for optimizing the models beyond what the original authors have already done with the exception for base models used for comparison purposes. A more comprehensive is needed of more complex models that utilizes cutting edge technologies, techniques, and theory. One of which is logic based models, for instance ASPER[10] that tries to incorporate logical descriptions of the problem and limits the model for better or equal results based on fewer samples[11]. Another approach is to incorporate randomness into the deterministic models to explain the variation in the data, for instance fractional Brownian motion[21].

move some  
of this to  
introduction

## 6 Conclusion

Everything is okay

## 7 Bibliography

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## Glossary

D | H | M | S

D

### DataFrame

A table of values. The name is from the python library Pandas used in this study.. 8

H

### Hashmap

A list of items where their unique placmnt in the list is detemend by their unique refrence key using a function that maps the key to a placement in the list.. 8

M

### Multiple Seasonal-Trend decomposition using LOESS

Based in the traditional Seasonal-Trend decomposition using LOESS it decomposes a time-series into several seasons, trend, and residual[18].. 9, 18

S

### Seasonal-Trend decomposition using LOESS

Takes a timeseries and decomposing it into a trend component, season component, and residual component using local regression for smoothing[17].. 9, 18

## Acronyms

K | L | M | N | S

K

### Kilden

Norwegian Institute of Bioeconomy Research Kilden. 6

L

### LMT

Norwegian Institute of Bioeconomy Research LandbruksMeteorologisk service. 6, 9

M

### MET

The Norwegian Meteorological Institute. 6, 9

### MSTL

Multiple Seasonal-Trend decomposition using LOESS. 9

**N**

**NMBU**

The Norwegian University of Life Science. 12

**S**

**STL**

Seasonal-Trend decomposition using LOESS. 9

## A Plots

## B Tables

scope	spesific scope	RMSE °C	MAE°C	bias °C	$\log(\kappa(\text{model}))$	digit sensitivity	R <sup>2</sup>
global	—	2.677	2.061	0.527	-0.322	-1	0.756
region	Østfold	2.564	2	0.176	-0.33	-1	0.8
region	Vestfold	2.565	1.958	0.785	-0.33	-1	0.81
region	Trøndelag	2.937	2.279	0.753	-0.32	-1	0.484
region	Innlandet	2.62	2.001	0.374	-0.324	-1	0.798
local	52	2.504	1.976	-1.2	-0.328	-1	0.803
local	41	2.135	1.665	0.13	-0.321	-1	0.872
local	37	2.513	1.938	0.067	-0.326	-1	0.83
local	118	3.029	2.422	1.722	-0.323	-1	0.656
local	50	2.176	1.7	0.815	-0.318	-1	0.836
local	42	2.739	2.099	0.746	-0.323	-1	0.807
local	38	2.983	2.323	1.149	-0.323	-1	0.736
local	30	2.276	1.708	0.428	-0.327	-1	0.859
local	57	3.079	2.419	0.744	-0.33	-1	0.617
local	39	2.79	2.186	0.633	-0.316	-1	0.615
local	34	3.158	2.424	0.811	-0.325	-1	-0.342
local	15	2.706	2.094	0.827	-0.323	-1	0.484
local	27	2.455	1.885	0.335	-0.33	-1	0.839
local	26	2.757	2.105	0.892	-0.324	-1	0.801
local	17	3.063	2.29	0.102	-0.329	-1	0.754
local	11	2.346	1.847	-0.023	-0.325	-1	0.755

Table 3: Results from hourly version of the Plauborg model for 20cm depth.

scope	specific scope	RMSE °C	MAE °C	bias °C	$\log(\kappa(\text{model}))$	digit sensitivity	R <sup>2</sup>
global	—	2.531	1.927	0.596	-0.439	-1	0.794
region	Østfold	2.448	1.894	0.512	-0.447	-1	0.816
region	Vestfold	2.412	1.81	0.733	-0.445	-1	0.846
region	Trøndelag	2.821	2.176	0.784	-0.442	-1	0.554
region	Innlandet	2.391	1.809	0.306	-0.449	-1	0.845
local	52	2.514	1.964	-0.349	-0.444	-1	0.636
local	41	1.938	1.519	0.151	-0.444	-1	0.903
local	37	2.344	1.804	0.237	-0.441	-1	0.857
local	118	2.928	2.322	1.639	-0.44	-1	0.706
local	50	1.908	1.472	0.558	-0.444	-1	0.884
local	42	2.501	1.885	0.703	-0.445	-1	0.852
local	38	3.055	2.368	1.363	-0.433	-1	0.754
local	30	2.072	1.555	0.354	-0.448	-1	0.892
local	57	2.906	2.263	0.677	-0.441	-1	0.677
local	39	2.77	2.151	0.701	-0.442	-1	0.633
local	34	3.009	2.303	0.857	-0.439	-1	-0.036
local	15	2.589	1.991	0.903	-0.447	-1	0.562
local	27	2.277	1.724	0.163	-0.445	-1	0.872
local	26	2.532	1.918	0.821	-0.444	-1	0.843
local	17	2.705	2.008	0.027	-0.445	-1	0.82
local	11	2.146	1.666	0.038	-0.439	-1	0.823

Table 4: Results from hourly version of the Plauborg model for 10cm depth.

scope	spesific scope	RMSE °C	MAE°C	bias °C	$\log(\kappa(\text{model}))$	digit sensitivity	R <sup>2</sup>
global	—	1.91	1.536	0.643	-1.909	-2	0.876
region	Østfold	1.94	1.541	-0.073	-1.918	-2	0.885
region	Vestfold	1.71	1.341	0.236	-1.918	-2	0.915
region	Trøndelag	1.843	1.56	1.461	-1.91	-2	0.797
region	Innlandet	2.16	1.735	1.019	-1.915	-2	0.863
local	52	2.33	1.873	-1.402	-1.914	-2	0.83
local	41	1.748	1.409	-0.371	-1.918	-2	0.914
local	37	1.877	1.496	0.353	-1.904	-2	0.905
local	118	1.742	1.384	1.139	-1.914	-2	0.886
local	50	1.251	0.985	0.096	-1.912	-2	0.946
local	42	1.966	1.54	0.346	-1.91	-2	0.901
local	38	1.721	1.367	0.515	-1.915	-2	0.912
local	30	1.817	1.471	-0.014	-1.904	-2	0.91
local	57	1.841	1.538	1.427	-1.91	-2	0.863
local	39	1.729	1.468	1.402	-1.913	-2	0.852
local	34	2.102	1.834	1.815	-1.91	-2	0.406
local	15	1.681	1.411	1.215	-1.912	-2	0.801
local	27	1.924	1.534	0.753	-1.91	-2	0.901
local	26	2.528	2.101	1.578	-1.91	-2	0.833
local	17	2.445	1.907	1.39	-1.905	-2	0.844
local	11	1.735	1.443	0.463	-1.919	-2	0.866

Table 5: Results from daily version of the Plauborg model for 20cm depth.

scope	spesific scope	RMSE °C	MAE°C	bias °C	$\log(\kappa(\text{model}))$	digit sensitivity	R <sup>2</sup>
global	—	2.075	1.622	0.607	-1.269	-2	0.861
region	Østfold	2.168	1.704	0.24	-1.256	-2	0.856
region	Vestfold	2.022	1.564	0.219	-1.265	-2	0.892
region	Trøndelag	1.957	1.528	1.235	-1.26	-2	0.785
region	Innlandet	2.167	1.712	0.711	-1.261	-2	0.873
local	52	2.418	1.837	-0.636	-1.266	-2	0.664
local	41	1.975	1.587	-0.293	-1.26	-2	0.9
local	37	2.206	1.755	0.373	-1.264	-2	0.873
local	118	2.165	1.697	1.137	-1.267	-2	0.839
local	50	1.395	1.105	-0.046	-1.262	-2	0.938
local	42	2.239	1.75	0.333	-1.262	-2	0.881
local	38	2.42	1.908	0.667	-1.26	-2	0.845
local	30	1.914	1.519	-0.046	-1.267	-2	0.908
local	57	1.978	1.547	1.108	-1.263	-2	0.85
local	39	1.896	1.455	1.193	-1.26	-2	0.828
local	34	2.142	1.687	1.535	-1.263	-2	0.475
local	15	1.806	1.428	1.114	-1.269	-2	0.787
local	27	2.063	1.627	0.396	-1.265	-2	0.895
local	26	2.43	1.937	1.251	-1.259	-2	0.855
local	17	2.273	1.789	0.9	-1.259	-2	0.873
local	11	1.879	1.504	0.339	-1.254	-2	0.864

Table 6: Results from daily version of the Plauborg model for 10cm depth.

scope	spesific scope	RMSE °C	MAE°C	bias °C	$\log(\kappa(\text{model}))$	digit sensitivity	R <sup>2</sup>
global	—	4.504	3.473	2.486	-0.796	-1	0.309
region	Østfold	4.348	3.363	1.901	-0.796	-1	0.424
region	Vestfold	4.564	3.47	2.297	-0.796	-1	0.397
region	Trøndelag	4.438	3.508	3.175	-0.796	-1	-0.178
region	Innlandet	4.687	3.567	2.597	-0.796	-1	0.353
local	52	3.556	2.841	0.559	-0.796	-1	0.604
local	41	4.248	3.286	1.677	-0.796	-1	0.491
local	37	4.754	3.675	2.174	-0.796	-1	0.391
local	118	4.726	3.654	3.208	-0.796	-1	0.162
local	50	4.048	3.025	2.207	-0.796	-1	0.434
local	42	4.863	3.741	2.364	-0.796	-1	0.393
local	38	4.832	3.682	2.601	-0.796	-1	0.308
local	30	4.465	3.433	2.015	-0.796	-1	0.456
local	57	4.655	3.636	3.153	-0.796	-1	0.125
local	39	4.31	3.39	3.083	-0.796	-1	0.081
local	34	4.581	3.673	3.469	-0.796	-1	-1.822
local	15	4.198	3.342	3.006	-0.796	-1	-0.241
local	27	4.672	3.547	2.535	-0.796	-1	0.415
local	26	5.17	4.009	3.282	-0.796	-1	0.302
local	17	5.042	3.835	2.913	-0.796	-1	0.335
local	11	3.821	2.924	1.692	-0.796	-1	0.35

Table 7: Results from the linear regression model for 20cm depth.



scope	spesific scope	RMSE °C	MAE°C	bias °C	$\log(\kappa(\text{model}))$	digit sensitivity	R <sup>2</sup>
global	—	4.23	3.267	2.302	-0.638	-1	0.424
region	Østfold	4.236	3.28	2.015	-0.638	-1	0.45
region	Vestfold	4.277	3.26	2.019	-0.638	-1	0.517
region	Trøndelag	4.133	3.274	2.893	-0.638	-1	0.043
region	Innlandet	4.282	3.255	2.242	-0.638	-1	0.504
local	52	3.679	2.889	1.226	-0.638	-1	0.221
local	41	3.976	3.07	1.494	-0.638	-1	0.593
local	37	4.501	3.503	2.07	-0.638	-1	0.473
local	118	4.5	3.486	2.93	-0.638	-1	0.306
local	50	3.611	2.702	1.766	-0.638	-1	0.584
local	42	4.571	3.525	2.109	-0.638	-1	0.506
local	38	4.815	3.741	2.502	-0.638	-1	0.388
local	30	4.053	3.106	1.733	-0.638	-1	0.588
local	57	4.293	3.356	2.775	-0.638	-1	0.295
local	39	4.057	3.193	2.835	-0.638	-1	0.213
local	34	4.259	3.414	3.175	-0.638	-1	-1.076
local	15	3.918	3.141	2.799	-0.638	-1	-0.003
local	27	4.272	3.236	2.078	-0.638	-1	0.551
local	26	4.714	3.651	2.902	-0.638	-1	0.456
local	17	4.518	3.435	2.478	-0.638	-1	0.499
local	11	3.567	2.713	1.529	-0.638	-1	0.51

Table 8: Results from the linear regression model for 10cm depth.



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