



"Relax. Pot temperatures have been going up and down for centuries."

Heatwave, in the ocean



Claudia Deal, Lucy Courtenay and Sophie McClellan enjoy the amazingly warm water at Bronte Beach. Picture: Sam Ruttyn

ANTON ROSE

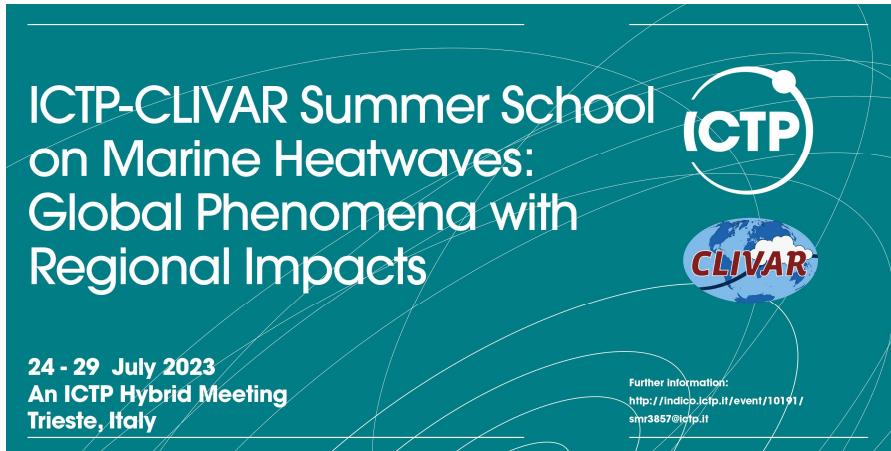
the maximum recorded is 25.2°C.

exceptionally high sea water tem-

23.8°C," the website said.

towards New Zealand. But this year





Marine Heatwaves and Impacts on the Atmosphere



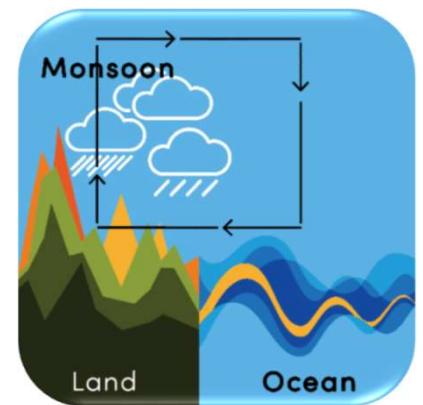
Roxy Mathew Koll

Indian Institute of Tropical Meteorology

Tropical Cyclone



Monsoon



Terrestrial Heatwaves



Why the Indo-Pacific?

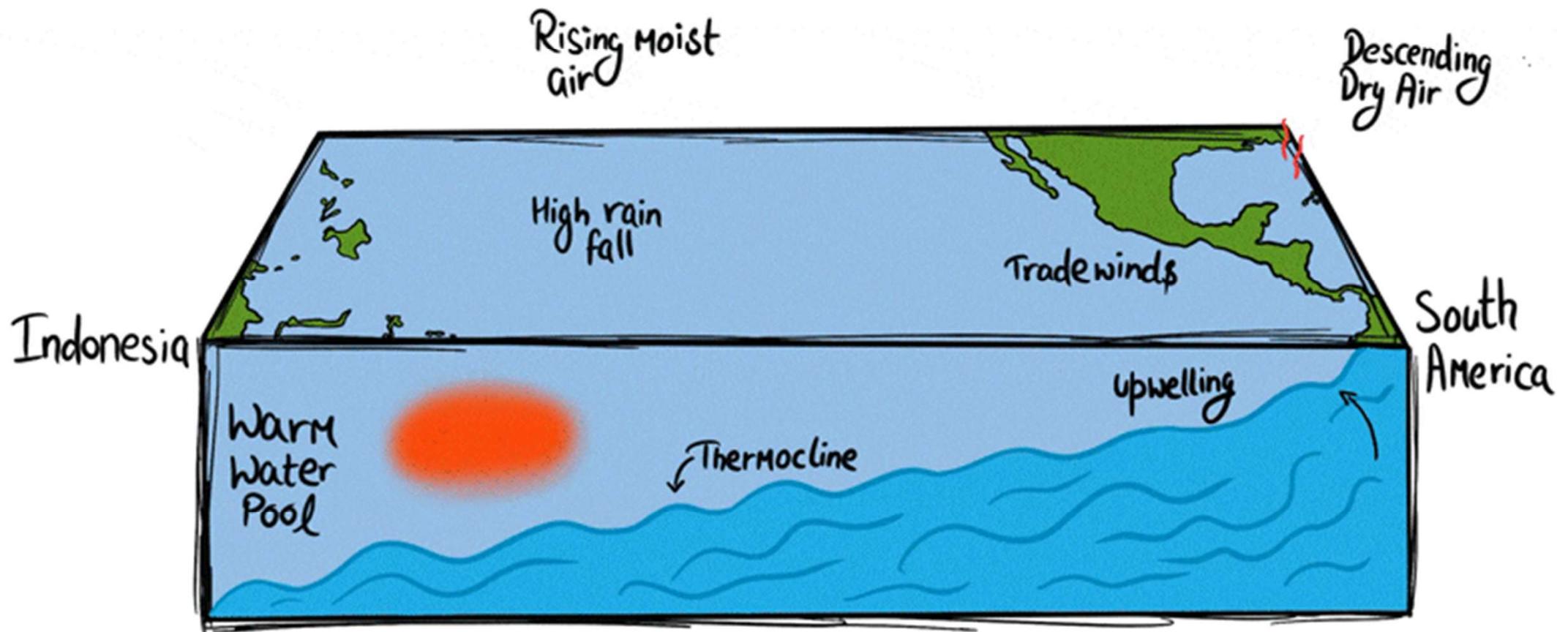
Map based on Population



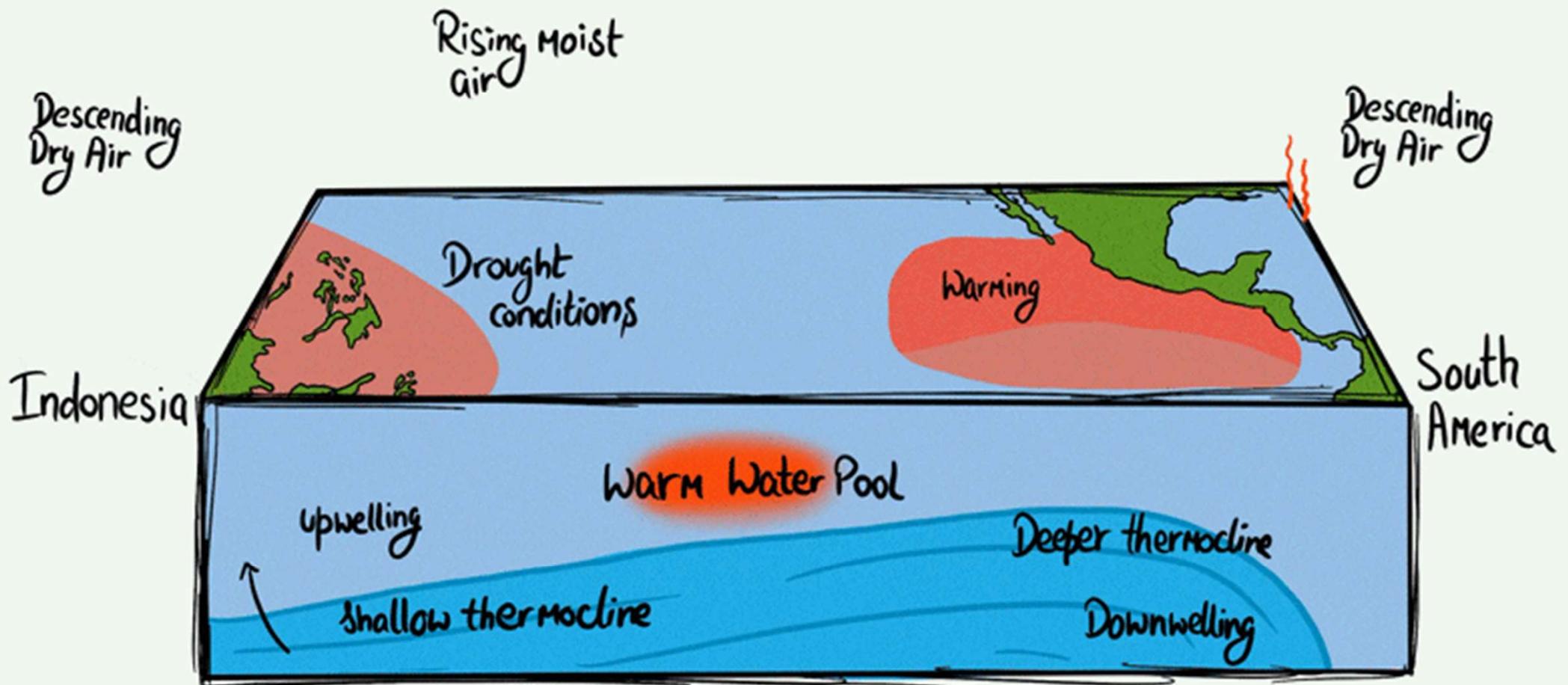
Ocean talking to the Atmosphere

The oceans impact weather and climate by heating (and cooling) the lower atmosphere. In particular, as seawater evaporates, the ocean surface cools; and when the moisture later condenses into cloud droplets, this heat is released, warming the atmosphere. This moistening, and then warming, makes the air buoyant, driving low-level baroclinicity and atmospheric convection, causing wind convergence at the surface and divergence aloft. At the equator, ocean heating of the atmosphere can result in towering convective clouds that reach the top of the troposphere. These disturbances in turn drive teleconnections in the atmosphere, affecting weather and climate remotely.

El Niño Southern Oscillation — ENSO neutral

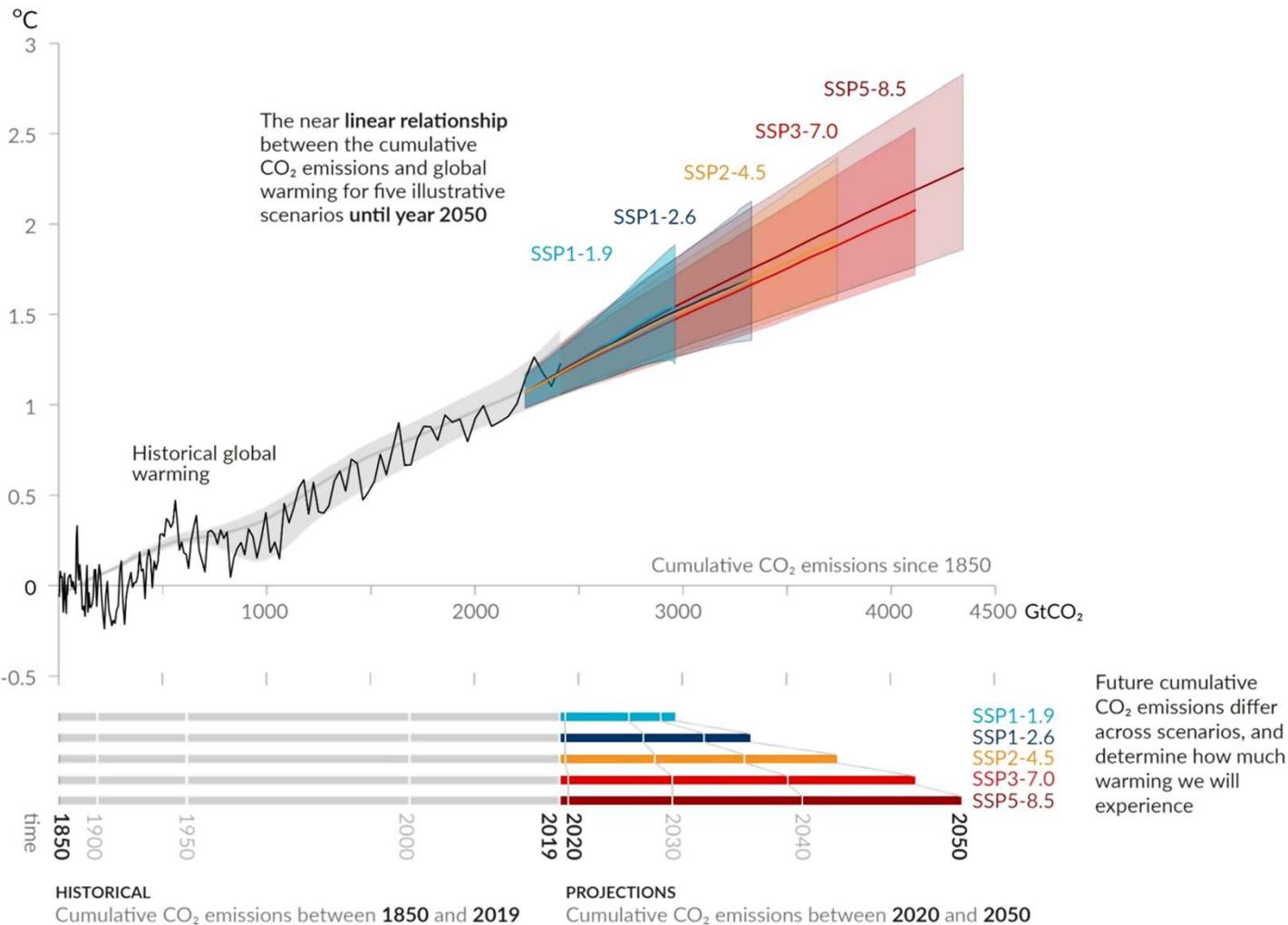


The El Niño



Global Warming

Global surface temperature increase since 1850-1900 ($^{\circ}\text{C}$) as a function of cumulative CO_2 emissions (Gt CO_2)

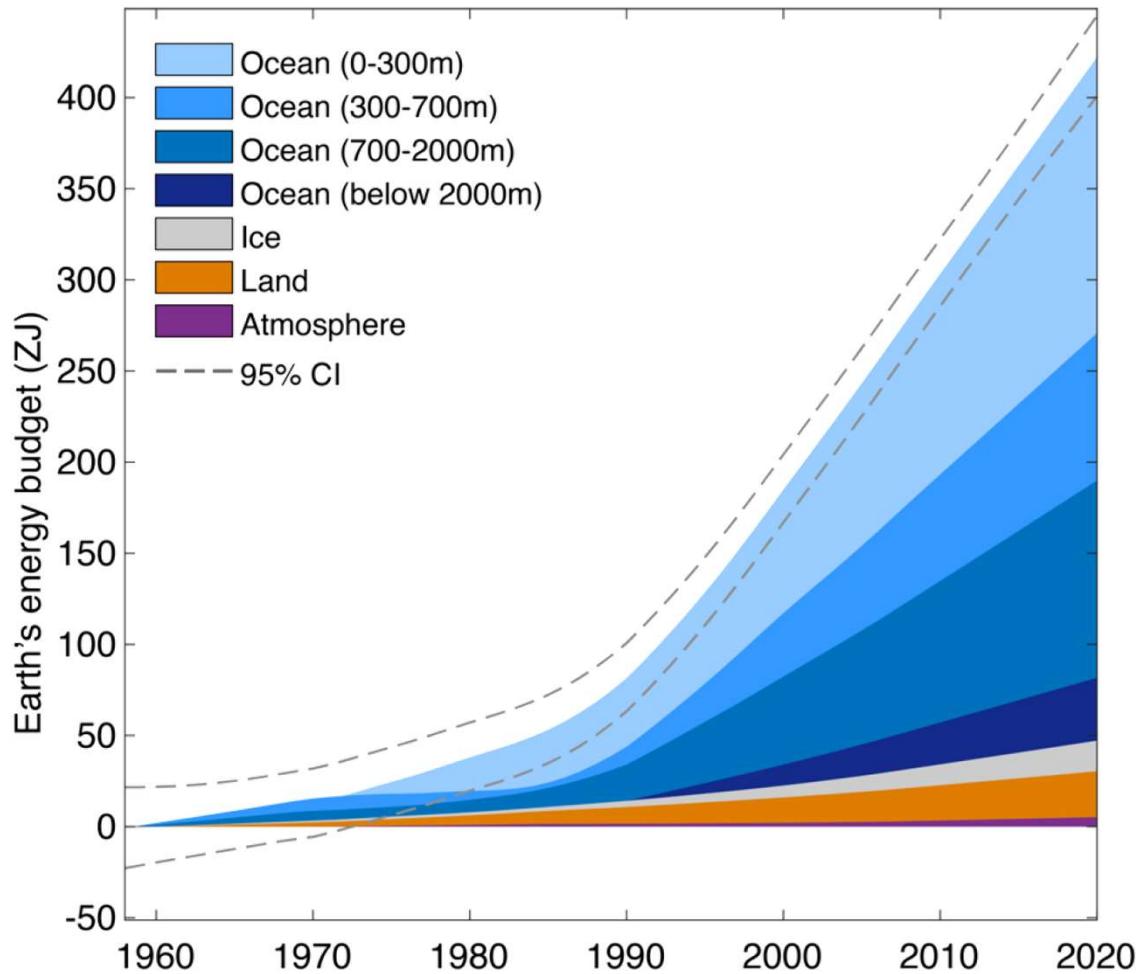


The global mean temperature has reached 1.1°C as of now.

It will cross 1.5°C in the current decade or next, and 2°C during 2040–2060.

This is because the nationally determined contributions (NDCs) submitted by nations via Paris Agreement are insufficient to flatten the curve.

Where does the Heat from Global Warming go?



Oceans take up 93% of the additional heat from global warming.

Atmosphere + Land + Ice accounts for less than 7% of the heat gain.

Extra heat intake rate by Oceans is
= 6 Hiroshima atomic bomb detonations per second

Heat Capacity of Ocean is higher than Land or Atmosphere

Heat capacity of soil/rocks and water,

$$C_p(\text{water}) = 4000 \text{ J kg}^{-1} \text{ } ^\circ\text{C}^{-1}$$

$$C_p(\text{rock/land}) = 800 \text{ J kg}^{-1} \text{ } ^\circ\text{C}^{-1}$$

Top 2 ½ meters of ocean can store the same heat as the entire troposphere

The volume of water which exchanges heat with the atmosphere per sq.meter of surface (depth of 100m) is 100 m³. The density of water is 1000 kg/m³.

mass = density × volume = $m_{\text{water}} = 10^5 \text{ kg}$.

Seasonal heat storage for ocean

$$\begin{aligned}\Delta E_{\text{oceans}} &= C_p(\text{water}) m_{\text{water}} \Delta T \text{ (} \Delta T = 10^\circ\text{C, is the typical change in temperature from winter to summer)} \\ &= (\mathbf{X})(\mathbf{X})(\mathbf{X}) \text{ Joules} = \mathbf{X} \text{ Joules}\end{aligned}$$

The volume of land which exchanges heat with the atmosphere 1 m³. Suppose the density of rock is 3,000 kg/m³, the mass of the soil and rock in contact with the atmosphere is 3,000 kg.

$$\begin{aligned}\Delta E_{\text{land}} &= C_p(\text{rock}) m_{\text{rock}} \Delta T \text{ (} \Delta T = 20^\circ\text{C)} \\ &= (\mathbf{X})(\mathbf{X})(\mathbf{X}) \text{ Joules} = \mathbf{X} \text{ Joules}\end{aligned}$$

$$\Delta E_{\text{oceans}} / \Delta E_{\text{land}} = \mathbf{X}$$

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$$\begin{aligned}\Delta E_{\text{land}} &= C_p(\text{rock}) m_{\text{rock}} \Delta T \text{ (} \Delta T = 20^\circ\text{C)} \\ &= \mathbf{(800)(3000)(20^\circ) Joules = 4.8 \times 10^7 Joules}\end{aligned}$$

$$\Delta E_{\text{oceans}} / \Delta E_{\text{land}} = 100$$

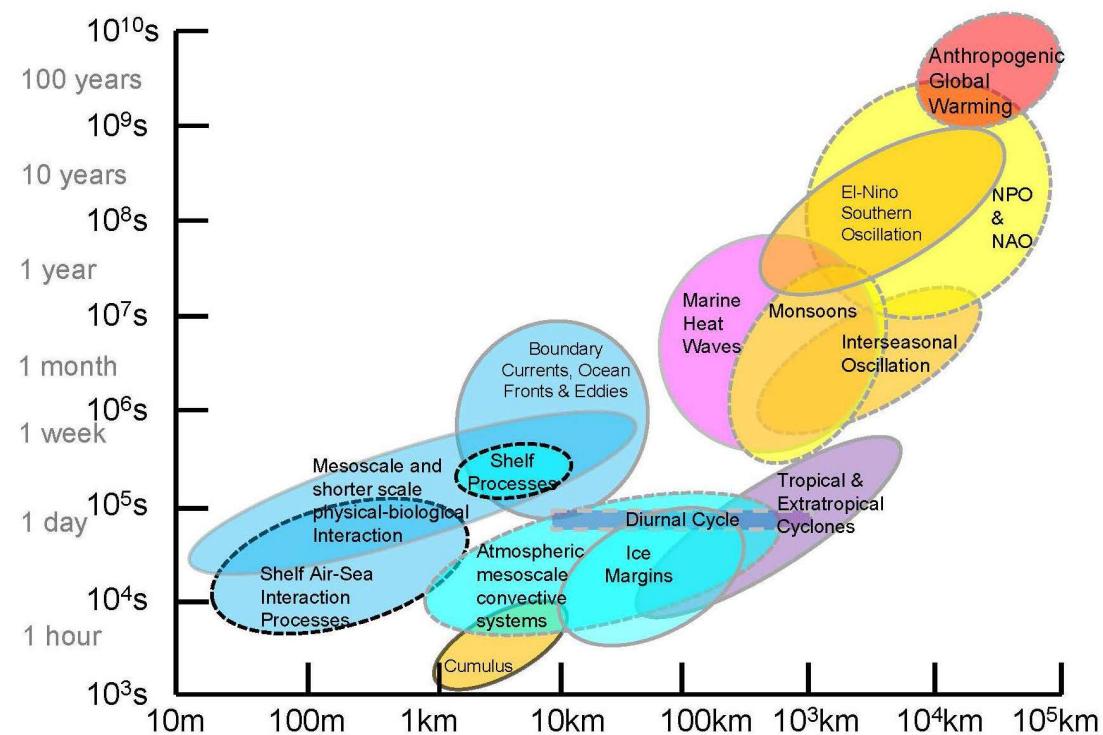
Affects Weather and Climate System

Because the ocean's capacity to store heat is about 1000 times greater than that of the atmosphere, long-term weather and climate predictability has its origins in the oceans.

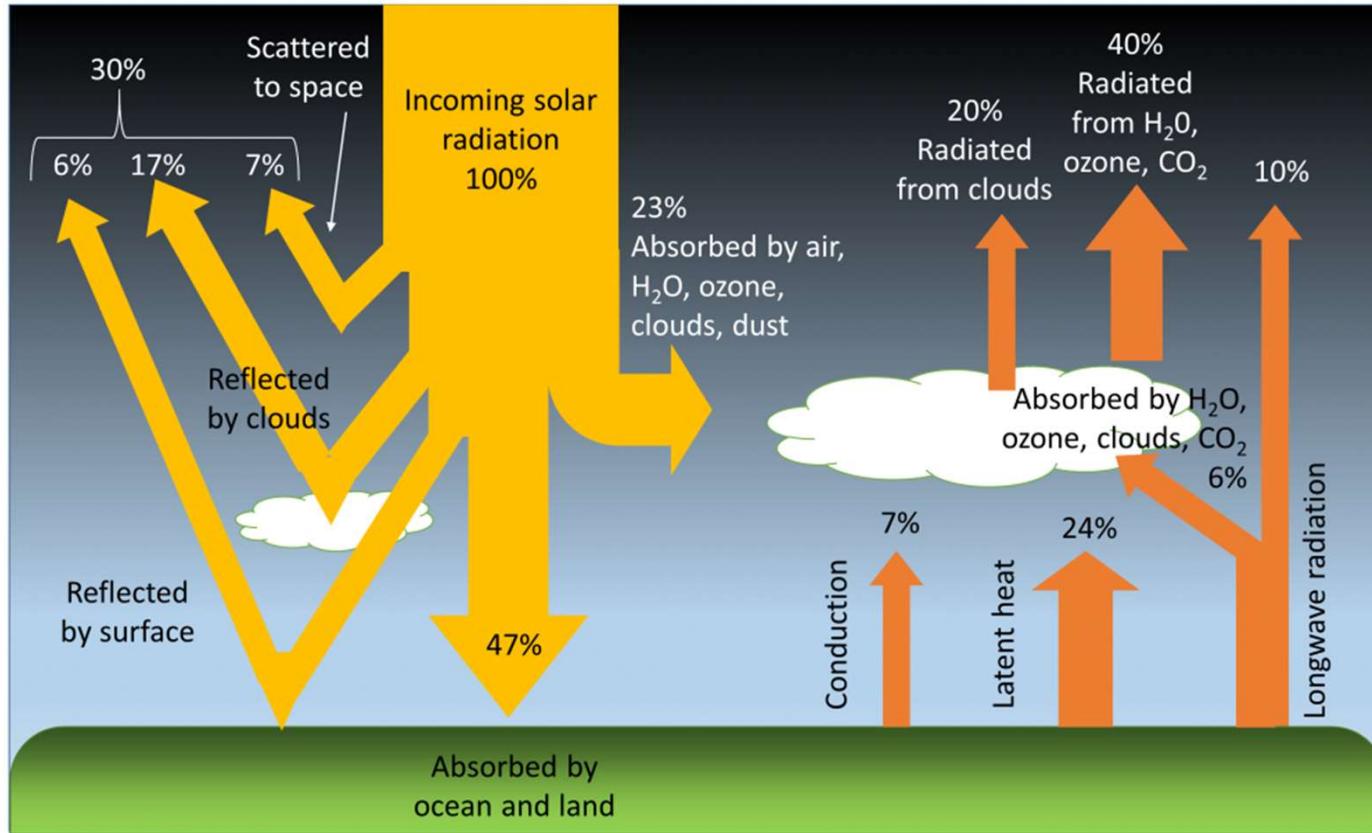
Heat storage and release occurs on a range of time scales and can provide predictability at 10–100 days (e.g., MJO, Monsoon), seasonal-interannual time scales (e.g., ENSO), and decades (e.g., PDO, AMO).

Predictions of weather and climate on these time scales have great economic benefits for agriculture, water resource management, energy management, human and ecosystem health among others.

To achieve useful predictions we must be able to **quantify** where, when and how much heat is released to the atmosphere.



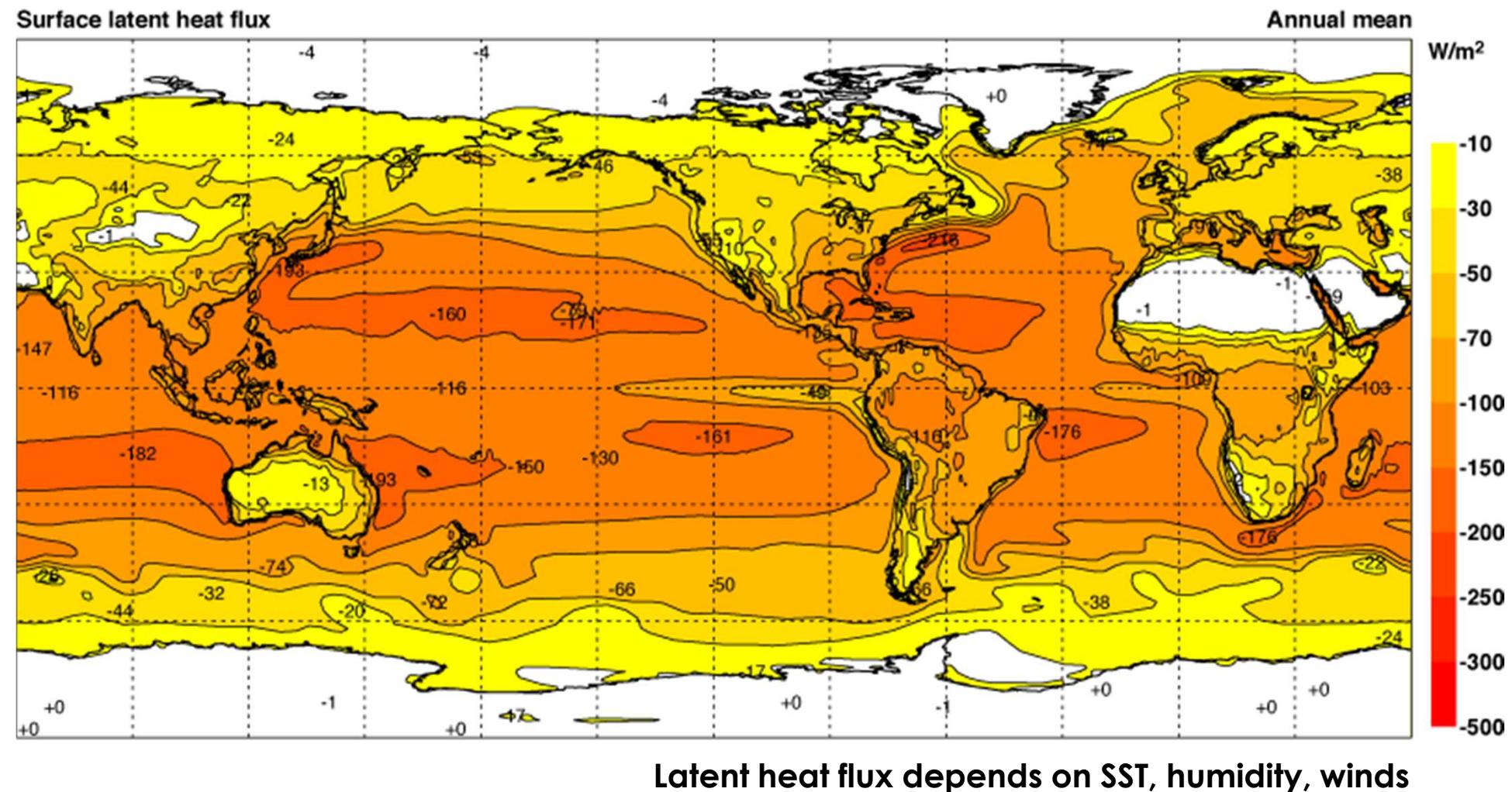
Heat Exchange between Ocean and Atmosphere



Sensible heat flux depends on SST, air temperature, winds

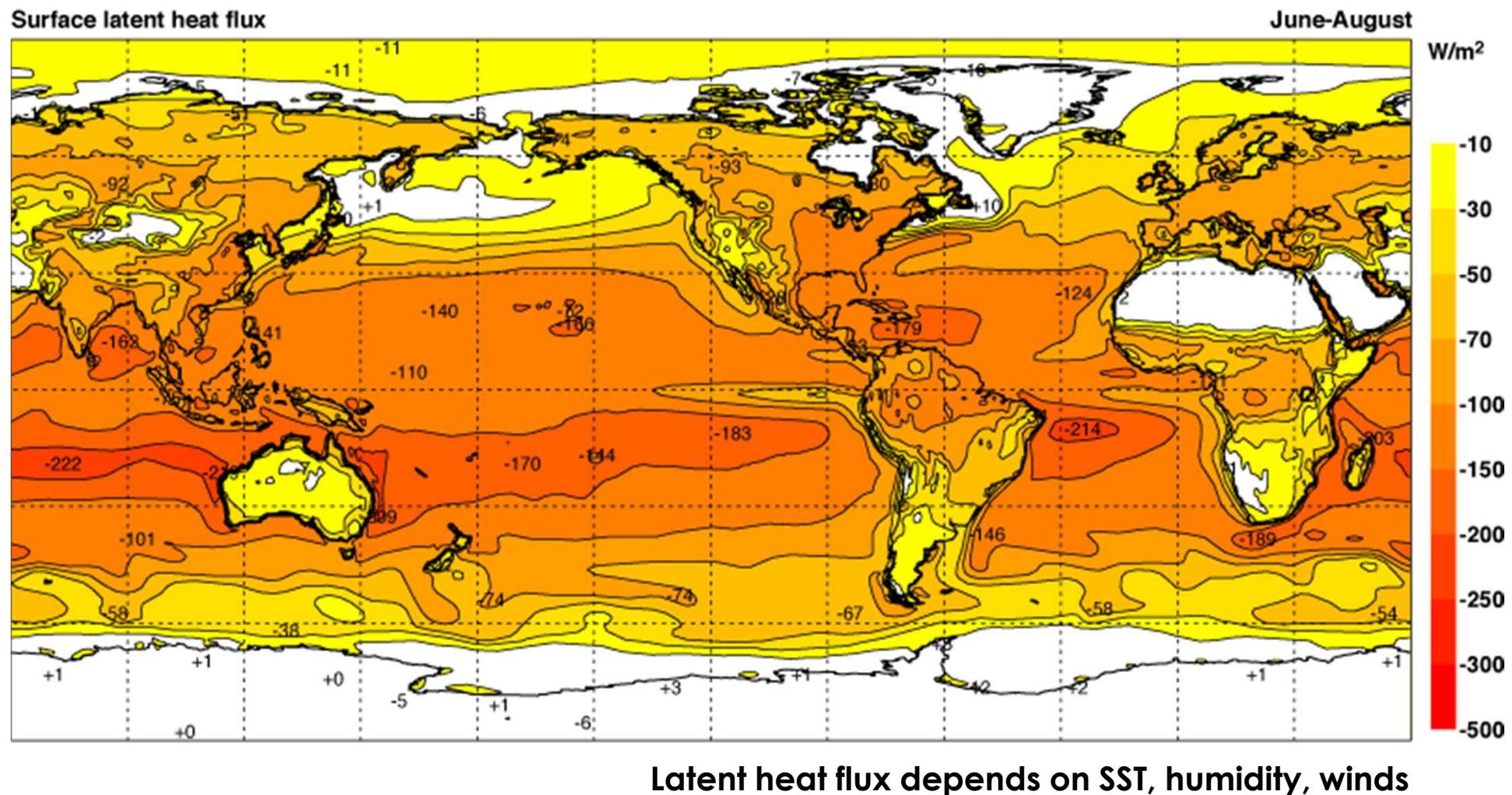
Latent heat flux depends on SST, humidity, winds

Latent Heat Flux at the Surface — Annual



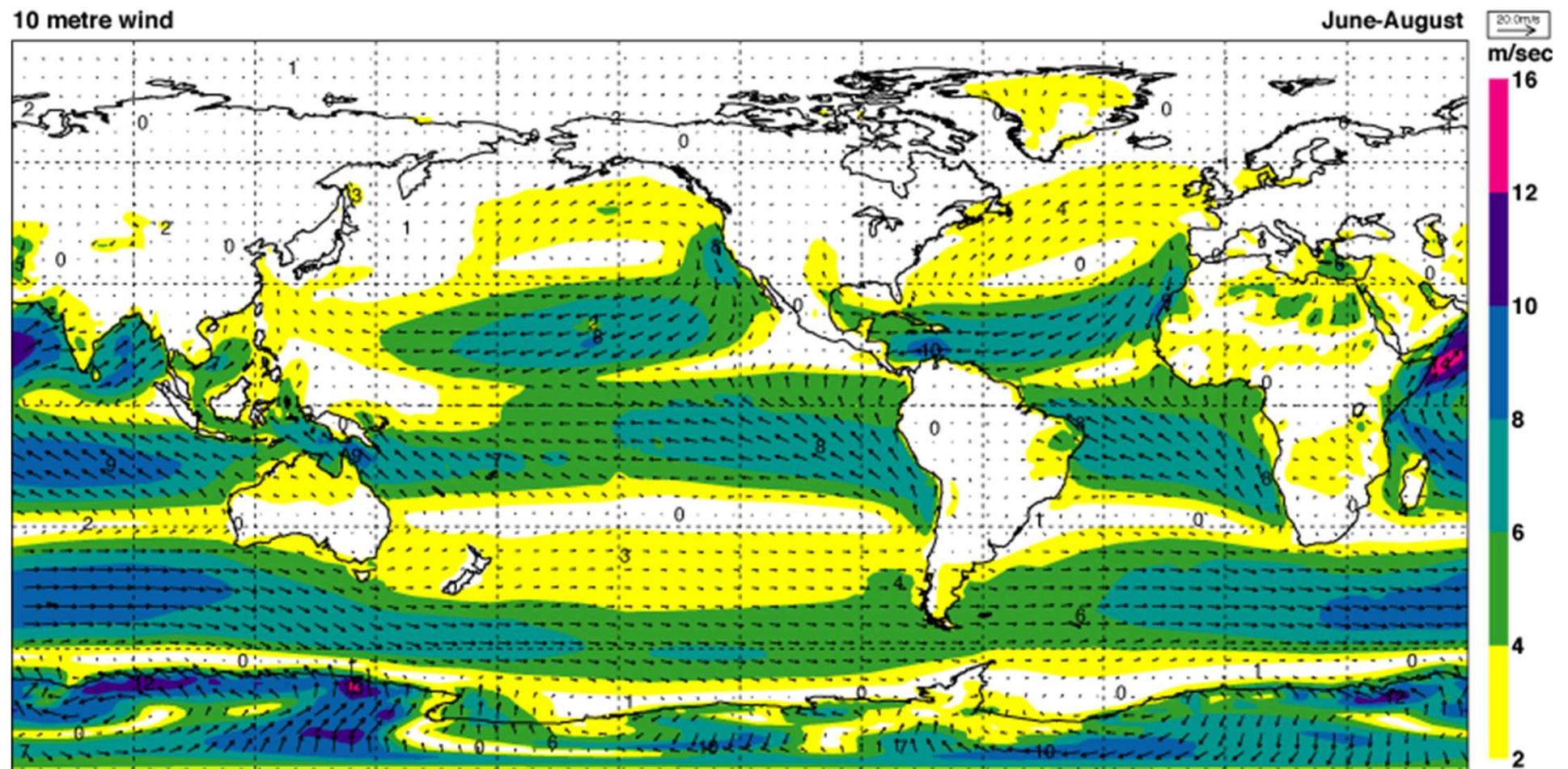
ERA Surface Climatologies

Latent Heat Flux at the Surface — northern Summer



ERA Surface Climatologies

Winds — northern Summer

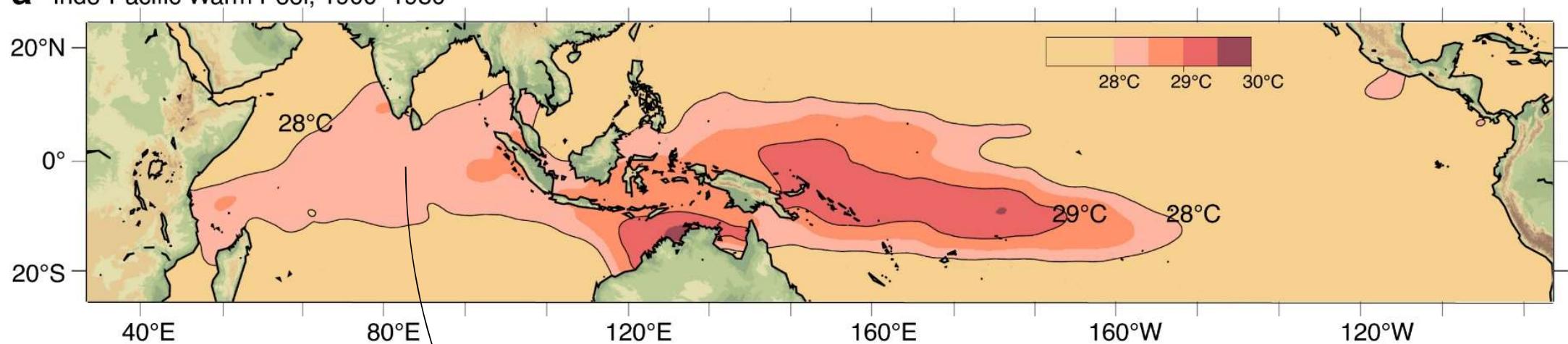


ERA Surface Climatologies

Indo-Pacific Warm Pool, Heat Engine of the Globe

Indo-Pacific warm pool – Sea Surface Temperatures (SST) above 28°C

a Indo-Pacific Warm Pool, 1900–1980

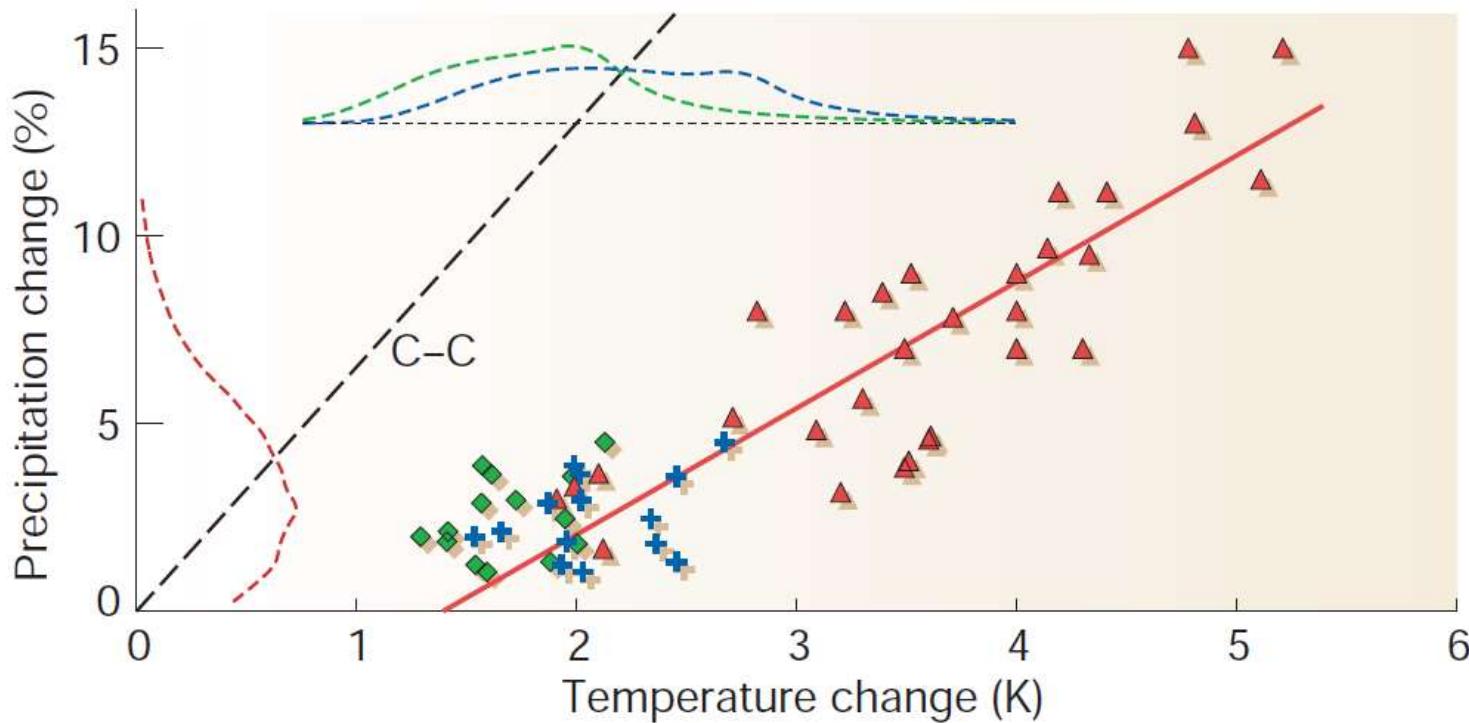


red colors indicate warm pool
— above 28°C

Thermodynamic Response to Temperature Change

The Clausius–Clapeyron relation implies that specific humidity and hence atmospheric moisture would increase roughly exponentially with temperature

7%/°C – substantially smaller than the sensitivity change documented.



Allan and Ingram, *Nature*, 2002; Allan and Soden, *Science*, 2006; Wentz et al., *Science*, 2007

Classical SST-Precipitation relationship

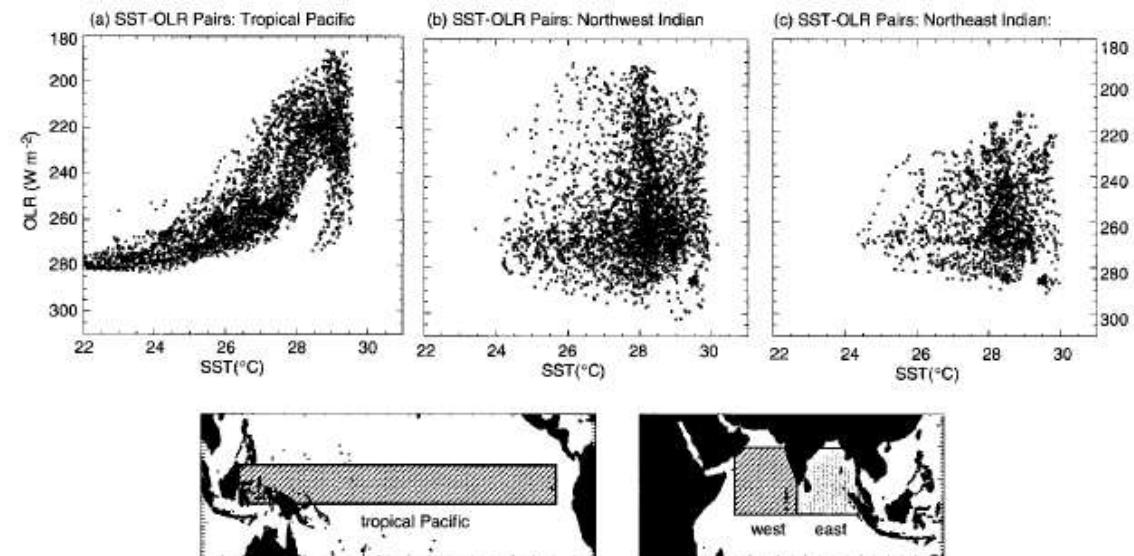
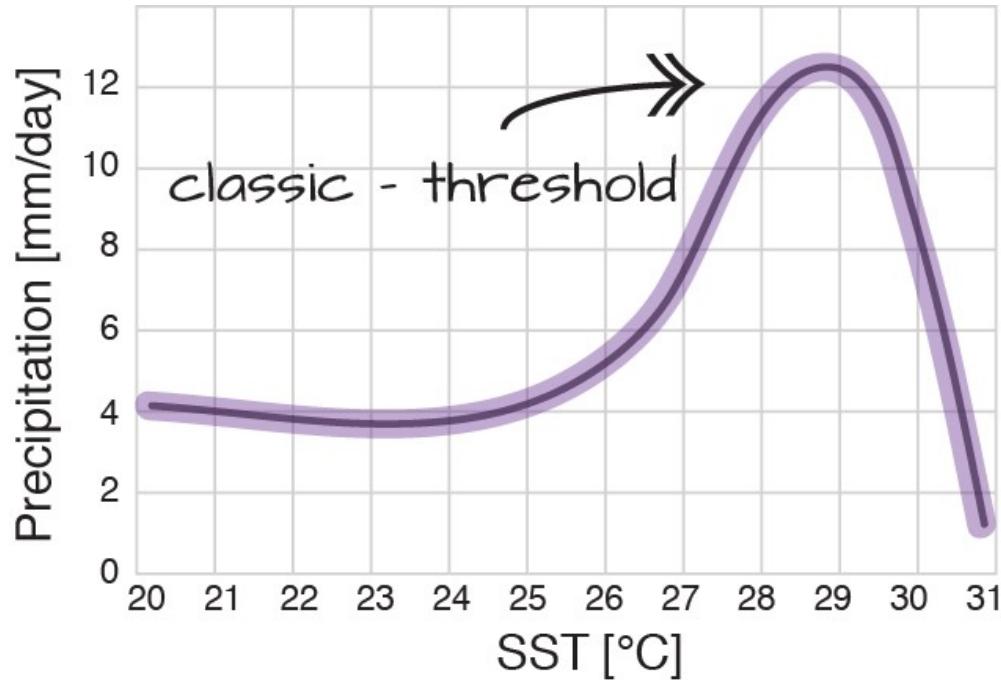
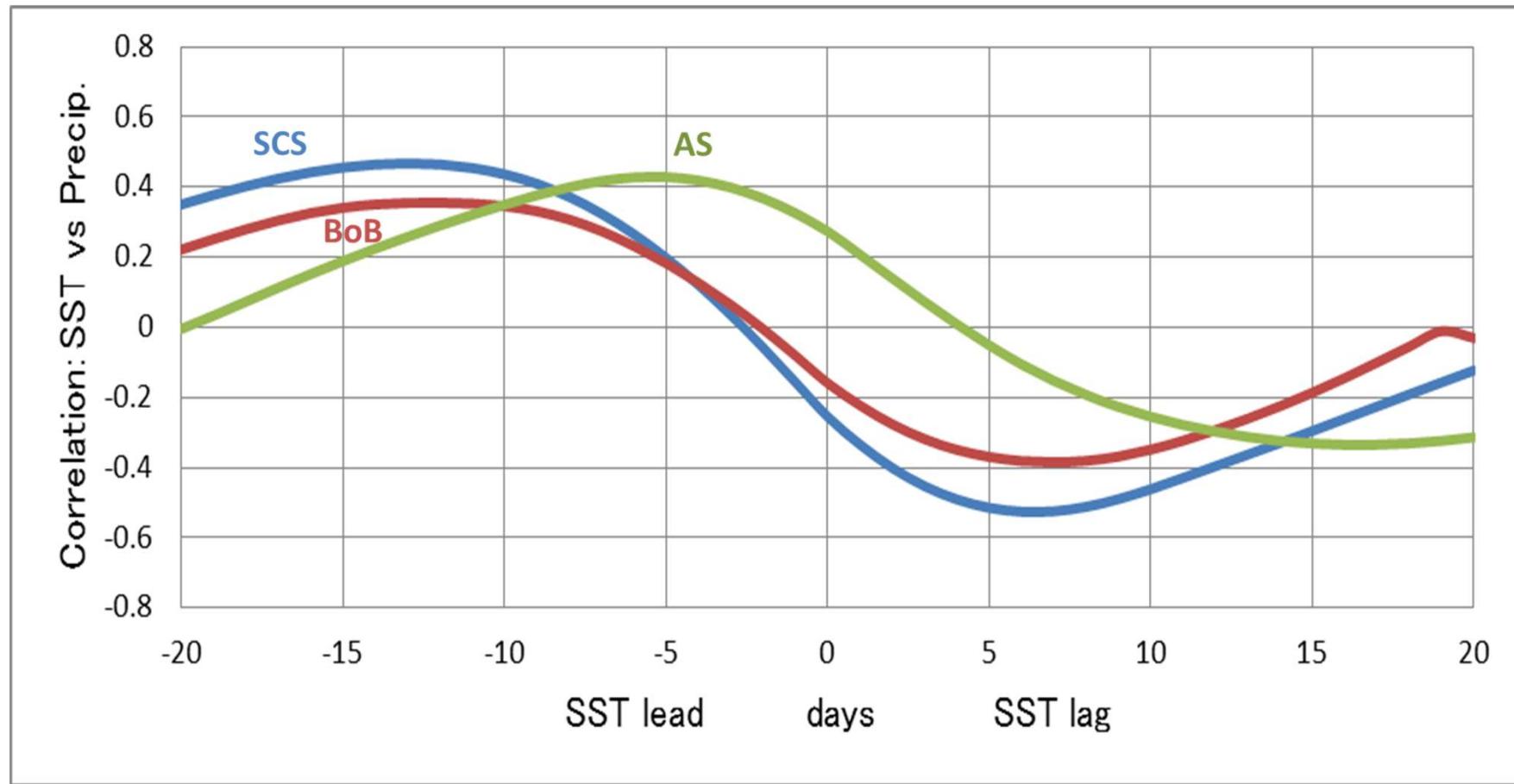


FIG. 3. OLR-SST (W m^{-2} and $^{\circ}\text{C}$ with ordinate scale reversed) pairs for (a) the tropical Pacific Ocean in the region $5^{\circ}\text{S}-5^{\circ}\text{N}$, $120^{\circ}\text{E}-90^{\circ}\text{W}$, (b) the northern Indian Ocean for the region $\text{equator}-20^{\circ}\text{N}$, $55^{\circ}-80^{\circ}\text{E}$, and (c) the northern Indian Ocean, $\text{equator}-20^{\circ}\text{N}$, $80^{\circ}-100^{\circ}\text{E}$. The geographical areas are depicted by the areas within the shaded boxes on map panel. Mean monthly values are plotted for Mar, Apr, and May for all points in the boxes. Data are from 1972-89.

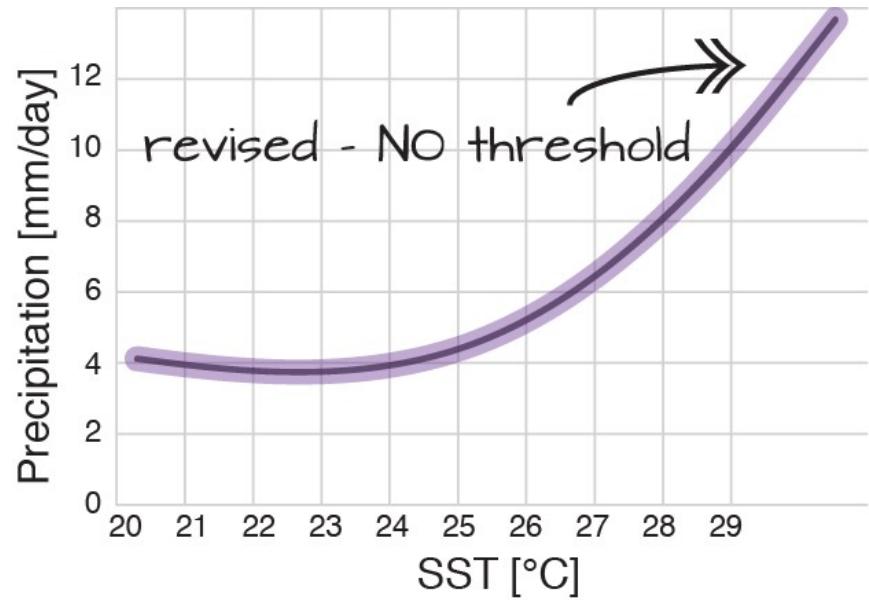
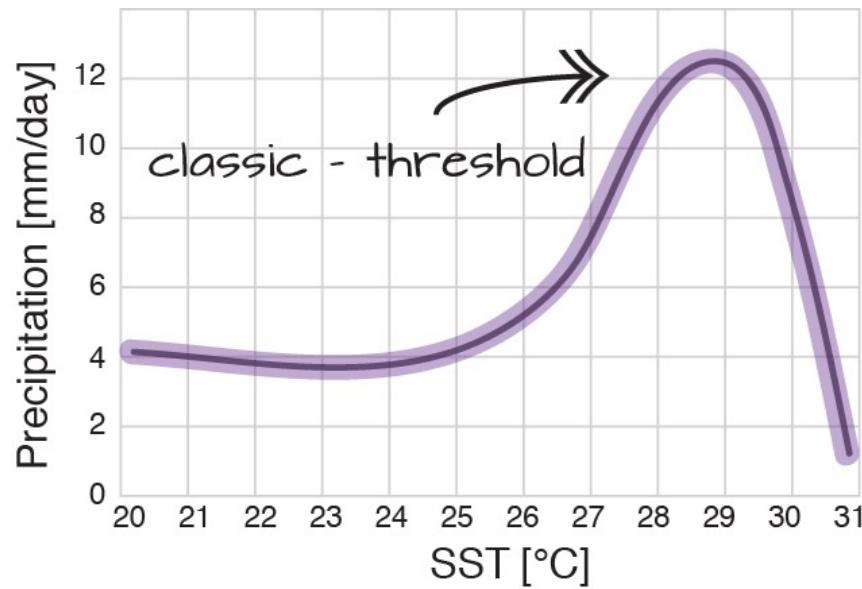
Earlier studies: Precipitation increases monotonously at SSTs beyond **26°C**,
but limited to: **Upper threshold of 28 - 29°C**

Explanation given: Precipitation tends to occur where
positive convective available potential energy (CAPE) exists
-> the occurrence of deep convection will tend to squelch CAPE?

SST-Precipitation relationship leads/lags by several days



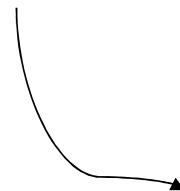
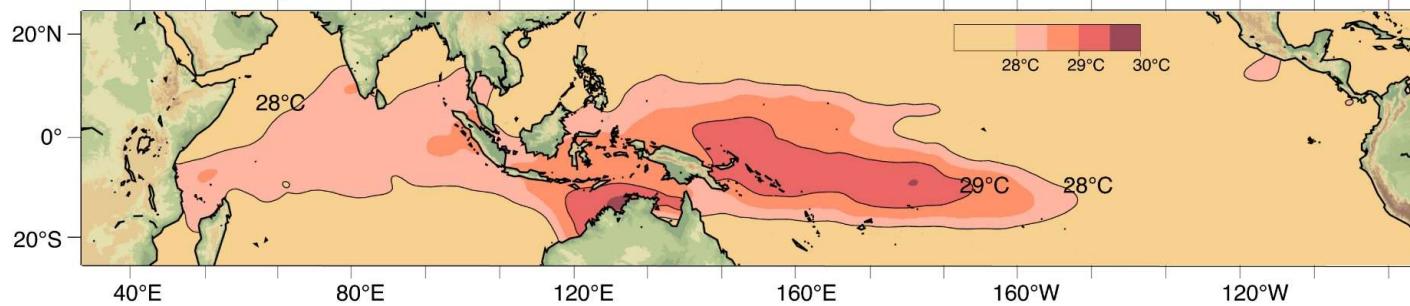
SST-Precipitation relationship considering the lag



Indo-Pacific Warm Pool, Heat Engine of the Globe

Indo-Pacific warm pool – Sea Surface Temperatures (SST) above 28°C

a Indo-Pacific Warm Pool, 1900–1980

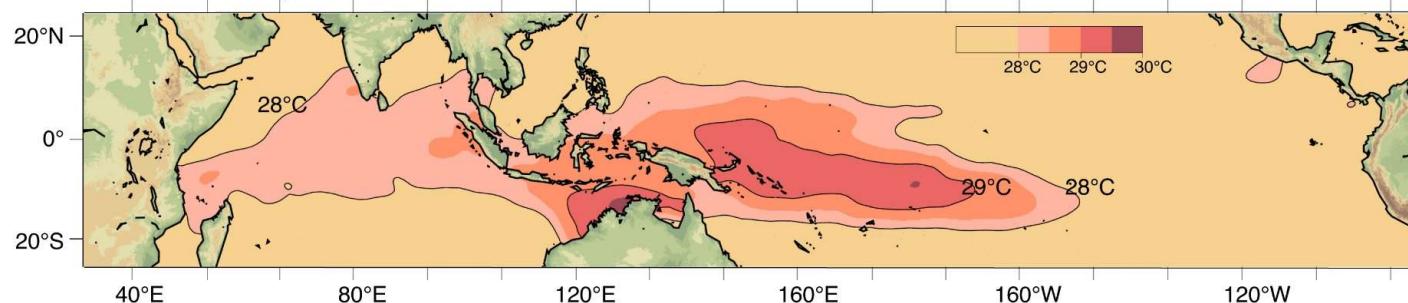


red colors indicate warm pool
—above 28°C

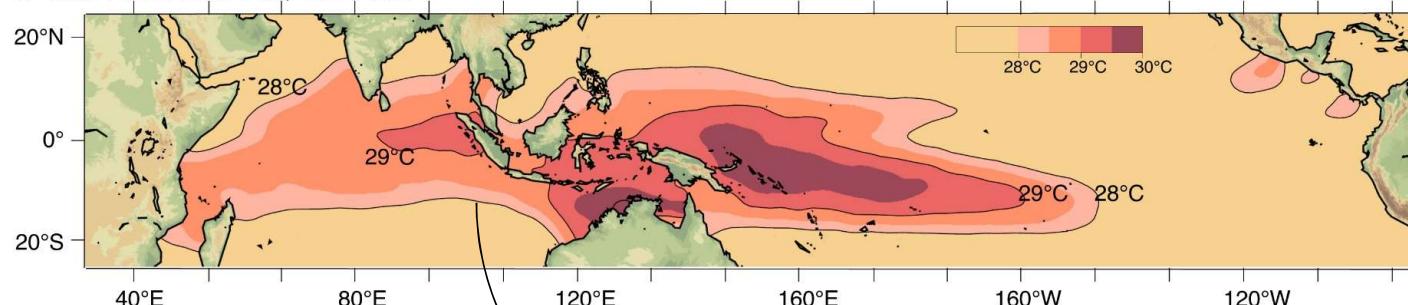
Indo-Pacific Warm Pool, Heat Engine of the Globe

Indo-Pacific warm pool – Sea Surface Temperatures (SST) above 28°C has expanded zonally and meridionally in the recent decades

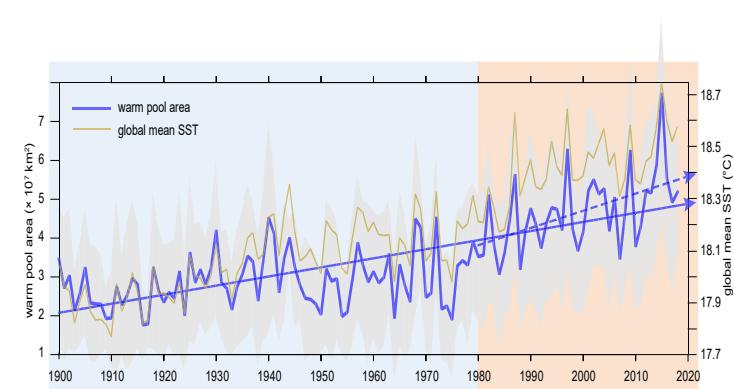
a Indo-Pacific Warm Pool, 1900–1980



b Indo-Pacific Warm Pool, 1981–2018

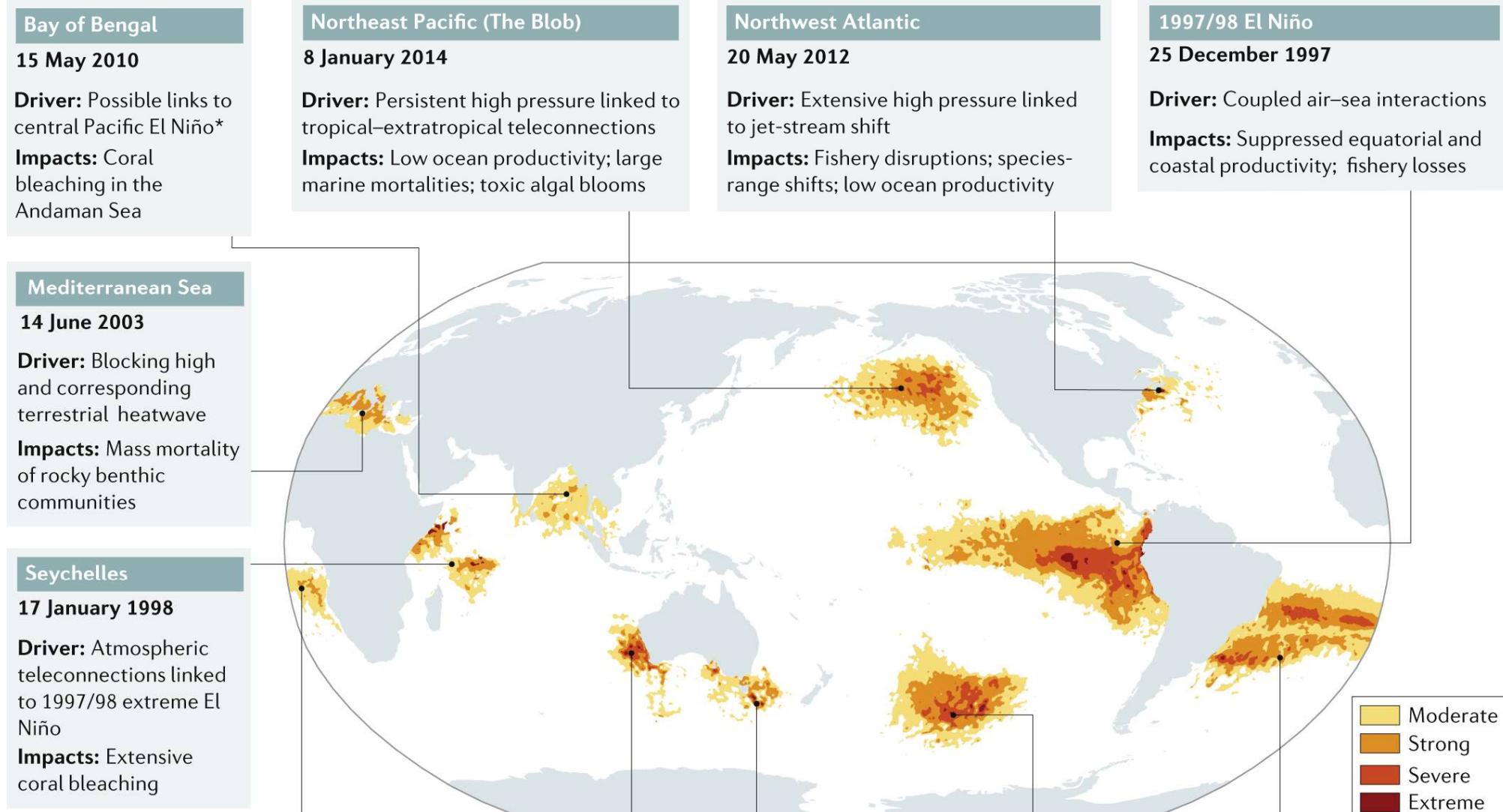


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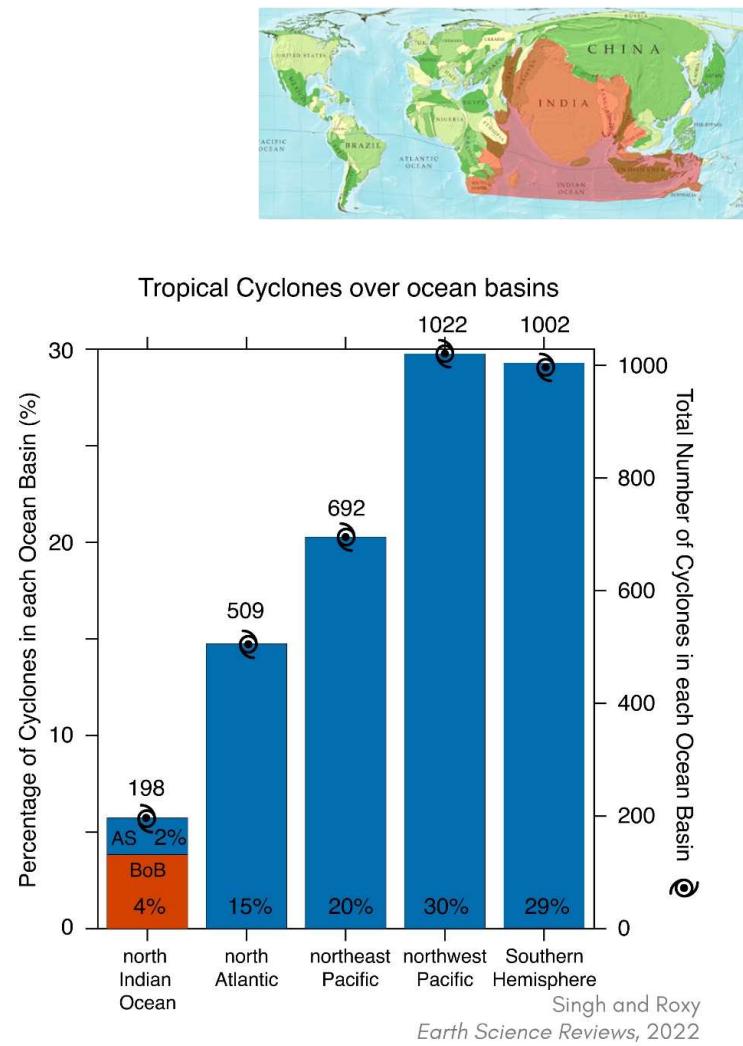
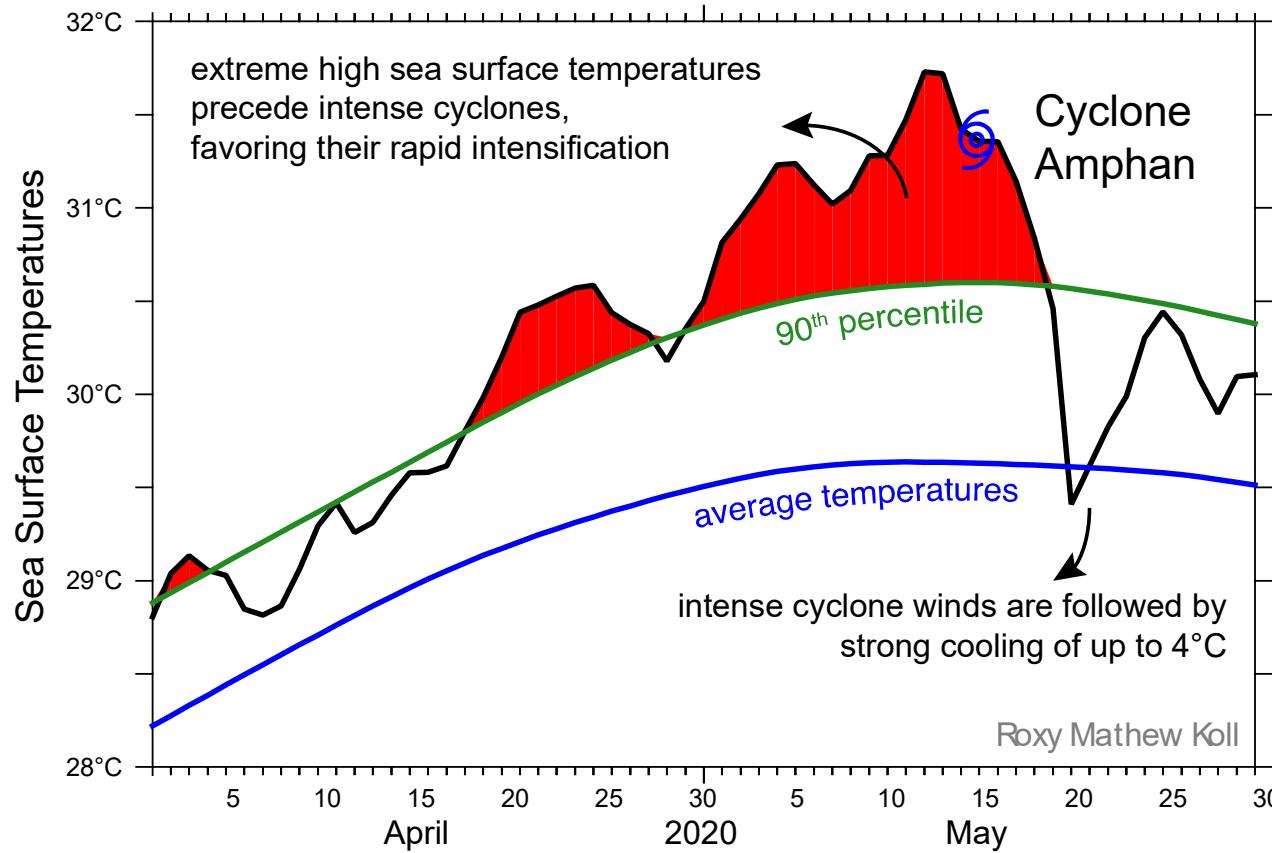


rate of expansion:
40 million sq.km per year

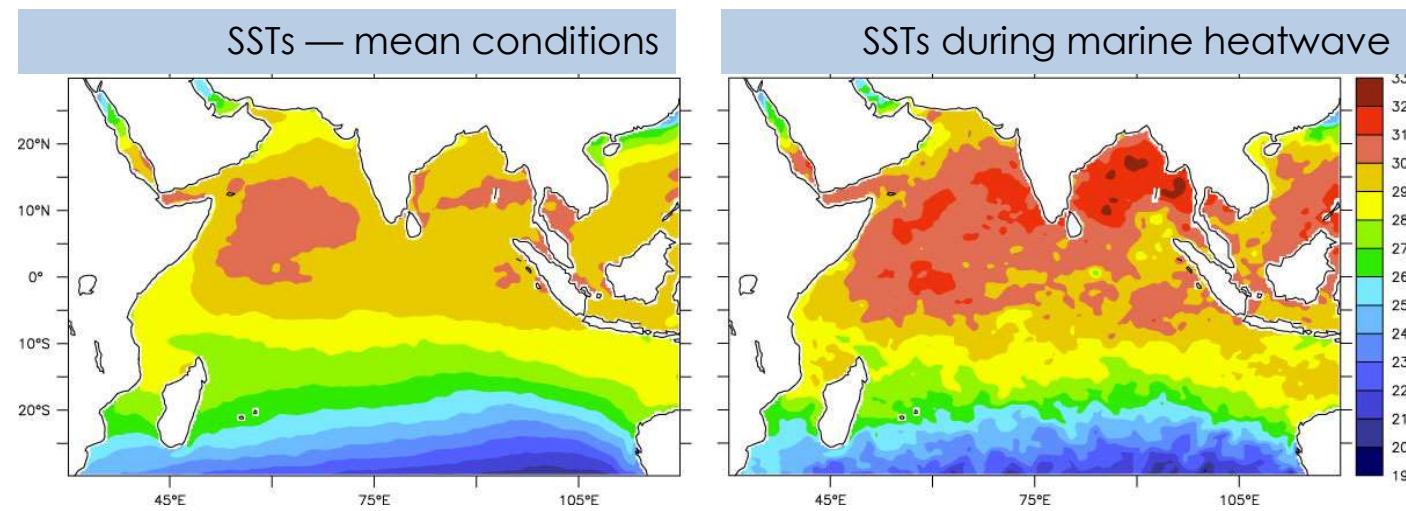
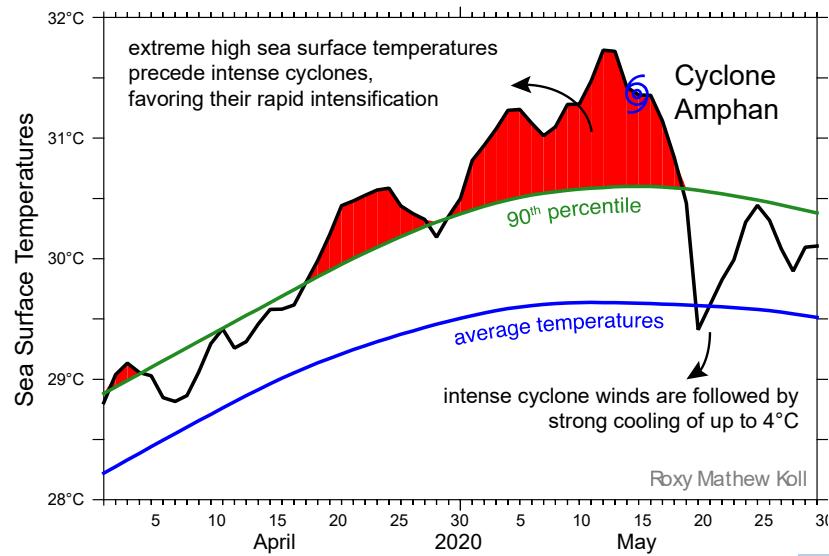
Marine Heatwaves in the Warm Pool



Marine Heatwaves and Cyclones



Marine Heatwaves and Cyclones

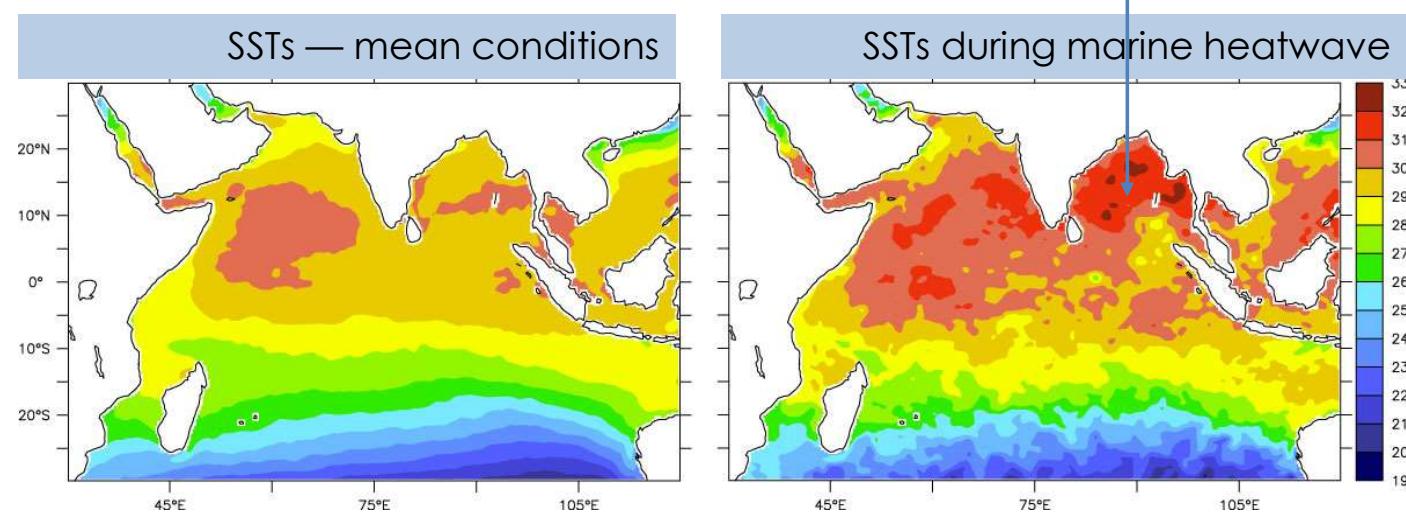


Marine Heatwaves and Corals



Corals have a mucus membrane — known as **zooxanthellae** — which acts as their shield and gives the color. Warm temperatures bleach it away.

These **marine heat waves** led to coral bleaching in Gulf of Mannar.



Marine Heatwaves and Cyclones

In-situ observations show much higher temperatures.

Moored Buoys Data

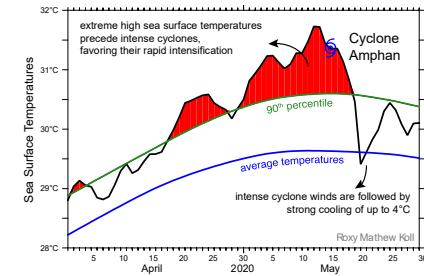
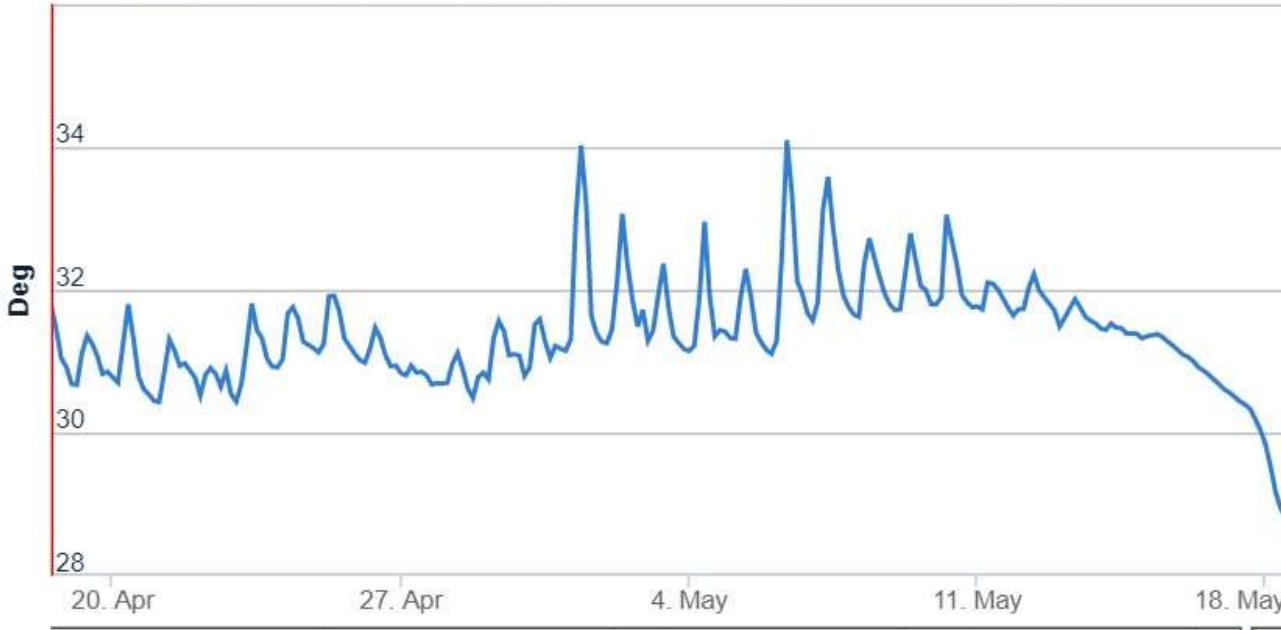
Parameter **SST Skin**

Submit

BD13 NIOT's Moored Buoy

Zoom **1m** **3m** **6m** **YTD** **1y** **All**

From **Apr 18, 2020** To **May 18, 2020**

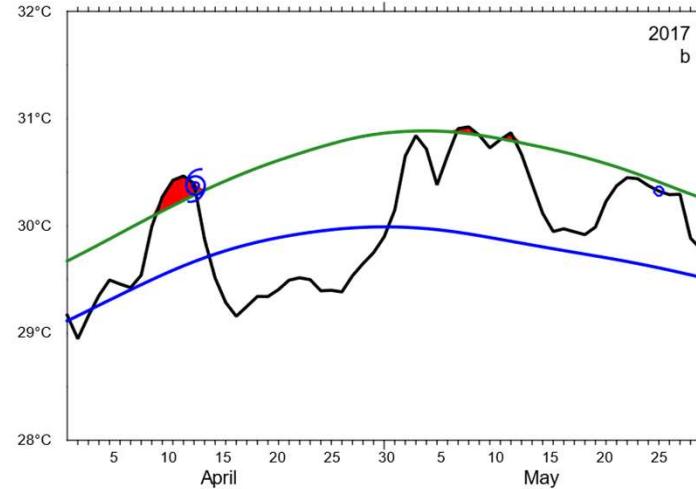
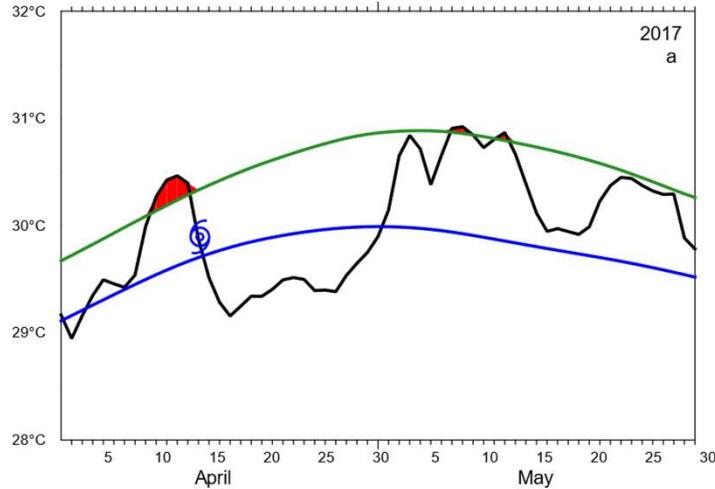


Bay of Bengal recorded surface temperatures of 32-34°C, before Cyclone Amphan.

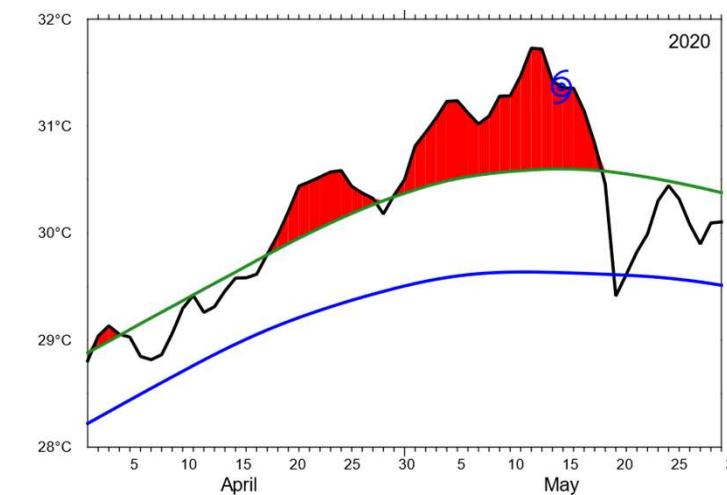
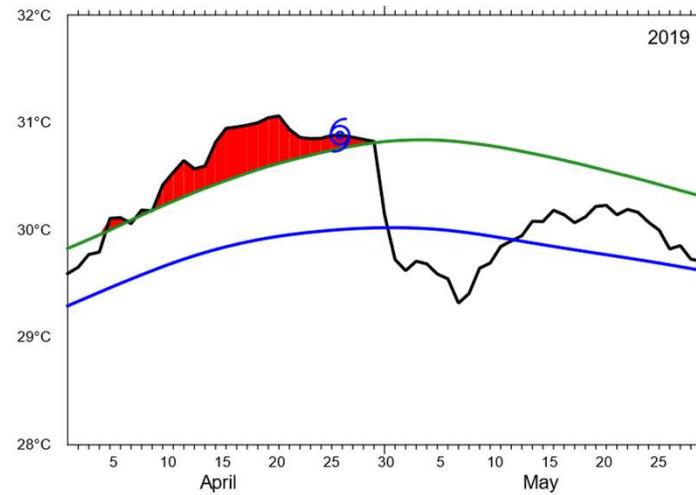
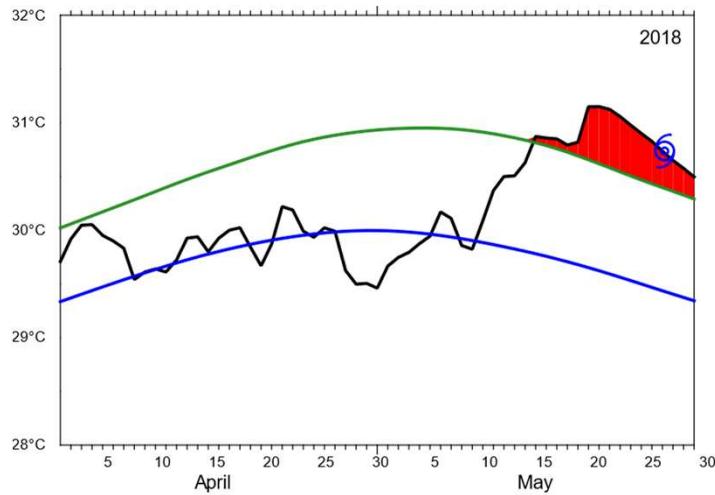
We have never seen such high values until now.

We need better ocean observations — and that needs regional partnership!

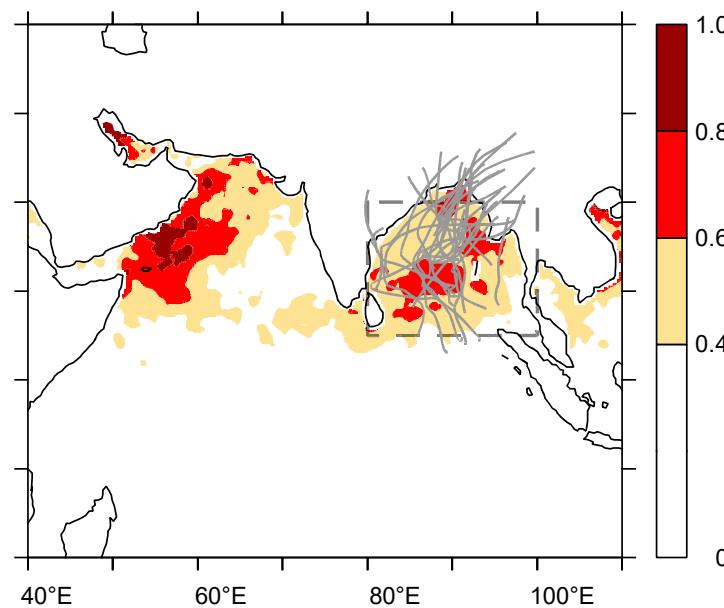
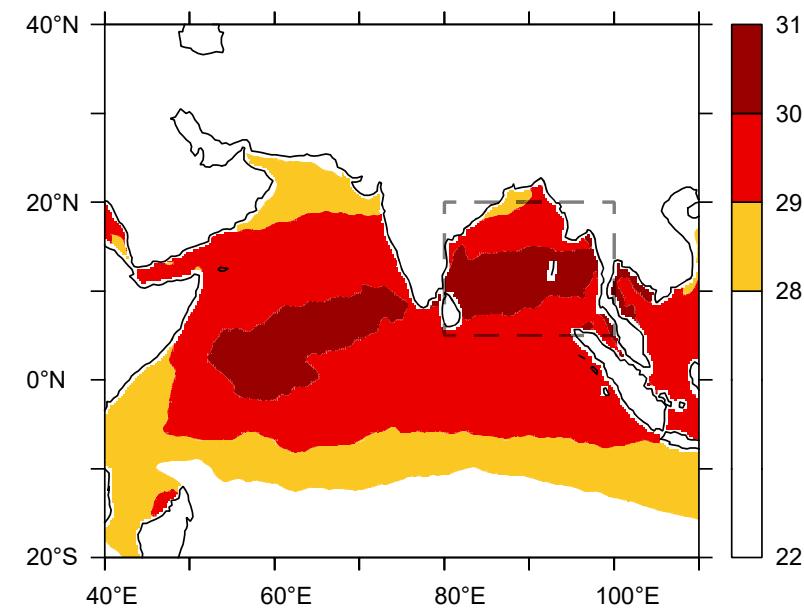
Marine Heatwaves and Cyclones



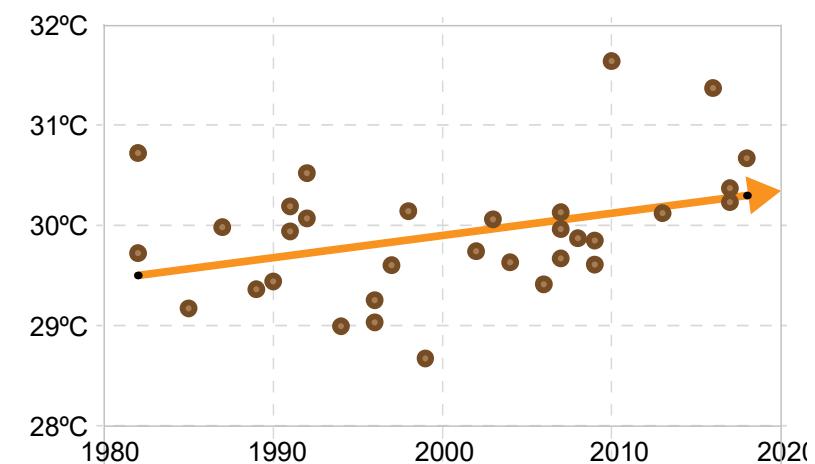
About 90% of the cyclones during pre-monsoon season were preceded by marine heatwaves



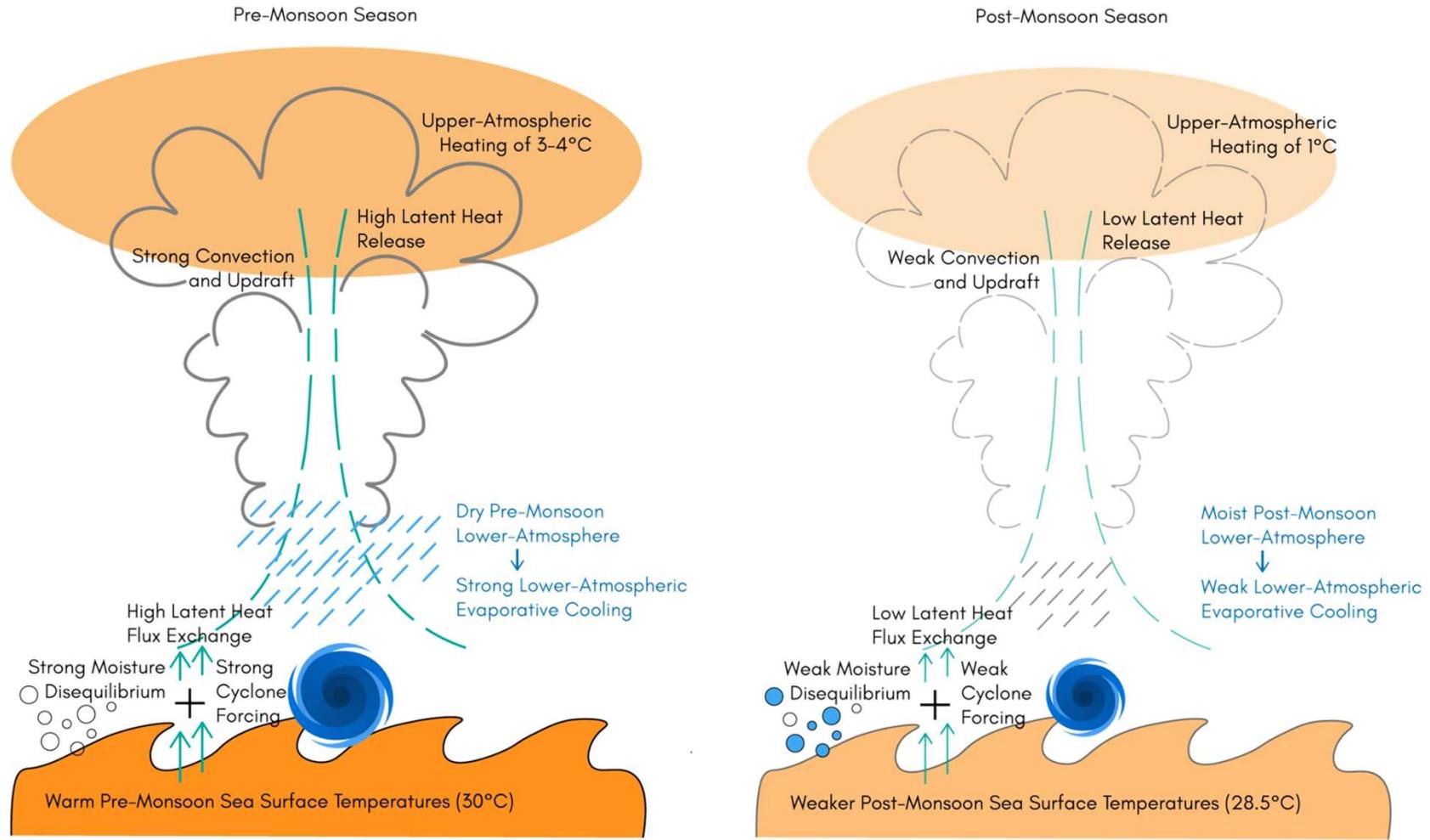
Marine Heatwaves and Cyclones



High SSTs and MHWs precede cyclogenesis, and these SSTs have been increasing in the Bay of Bengal

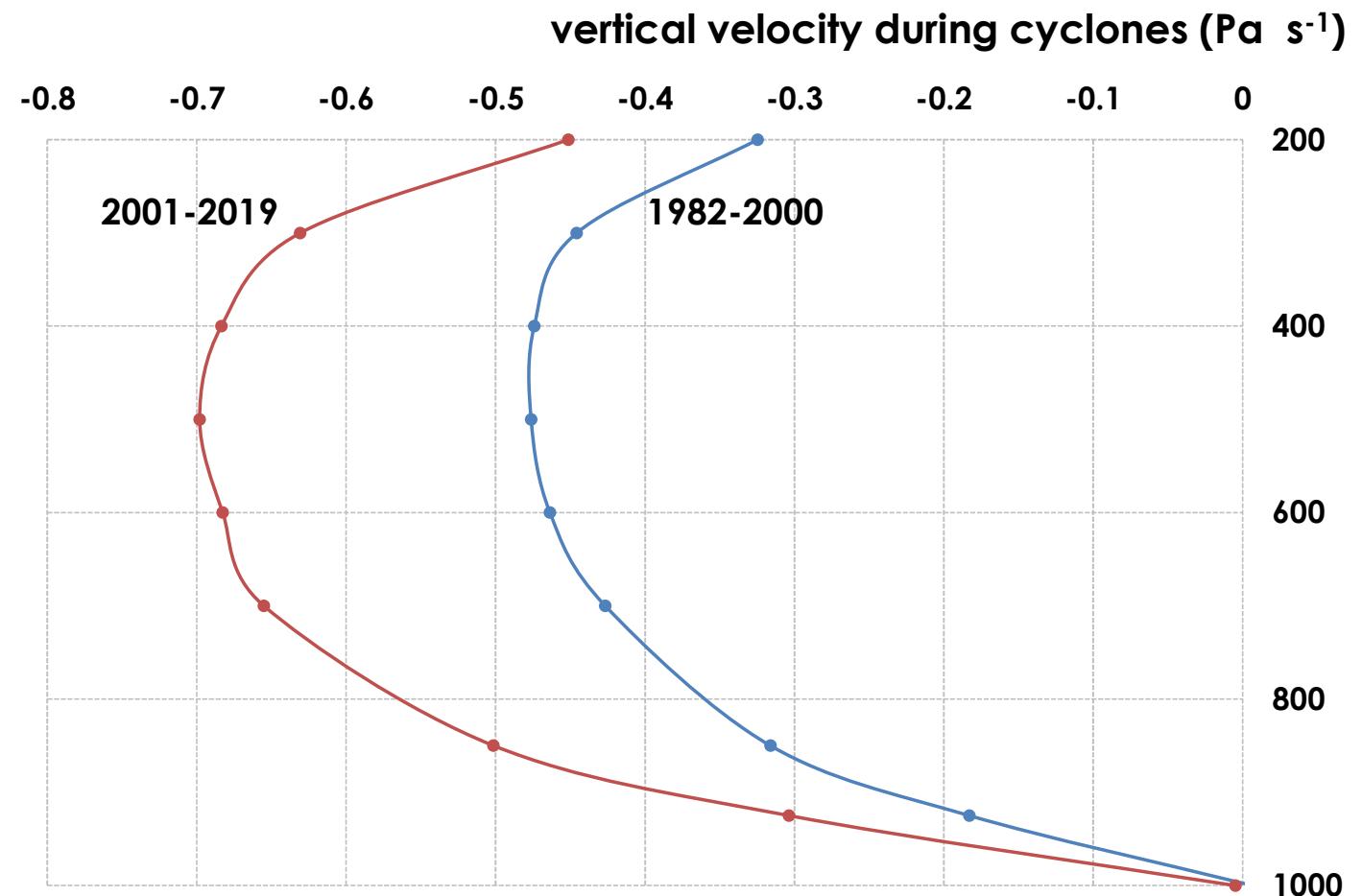


Upper Atmospheric Heating during Cyclone Seasons

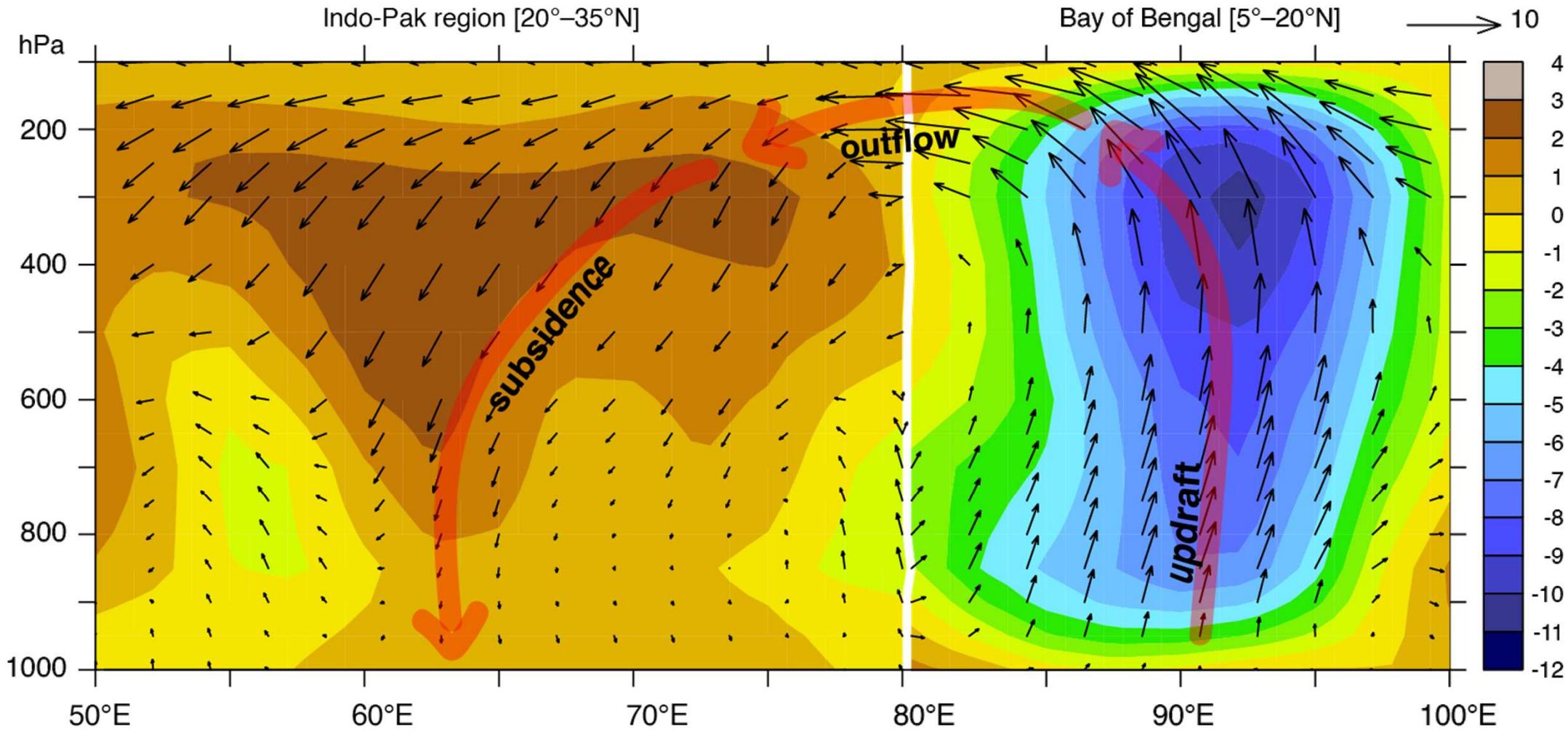


Cyclones and Enhanced Updraft

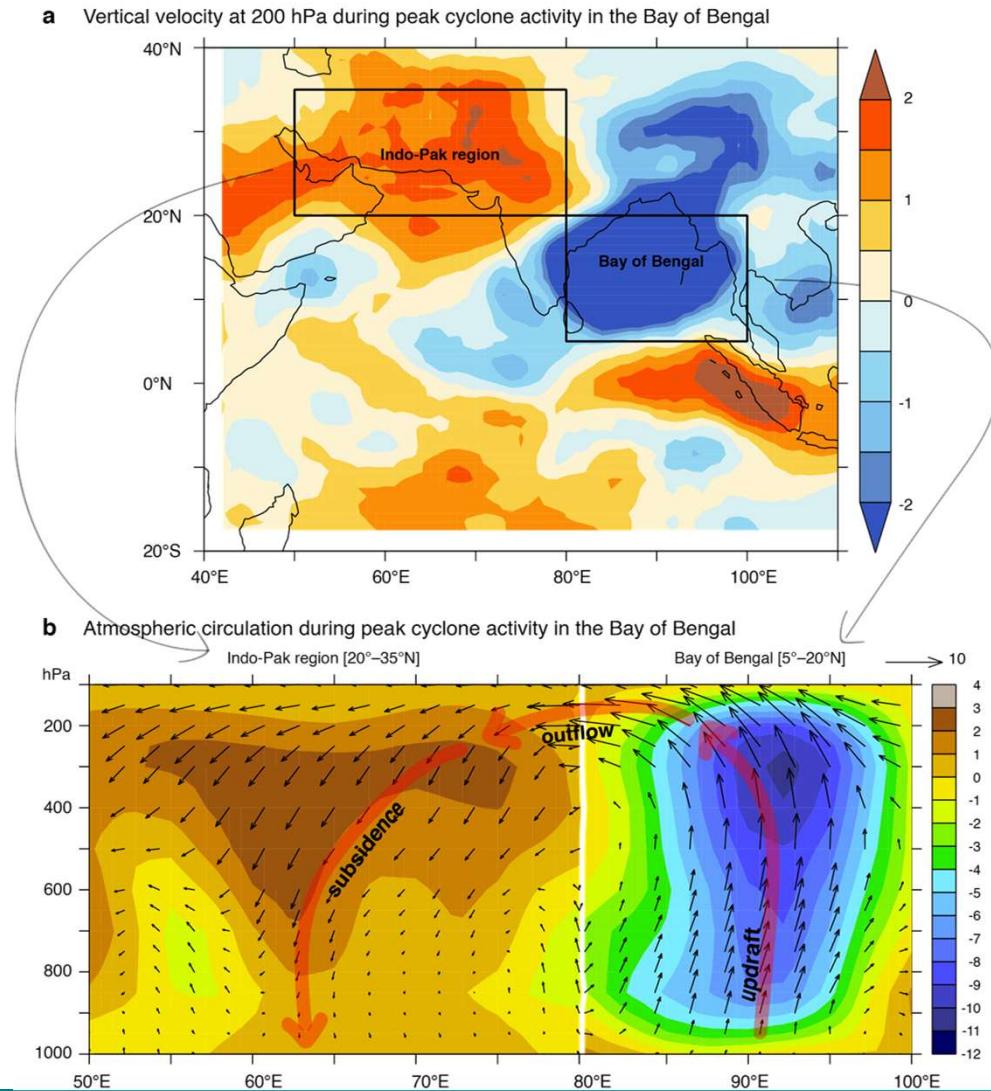
**High SSTs and latent heat flux from cyclones
lead to enhanced updraft (vertical velocity)**



Cyclones and Enhanced Updraft



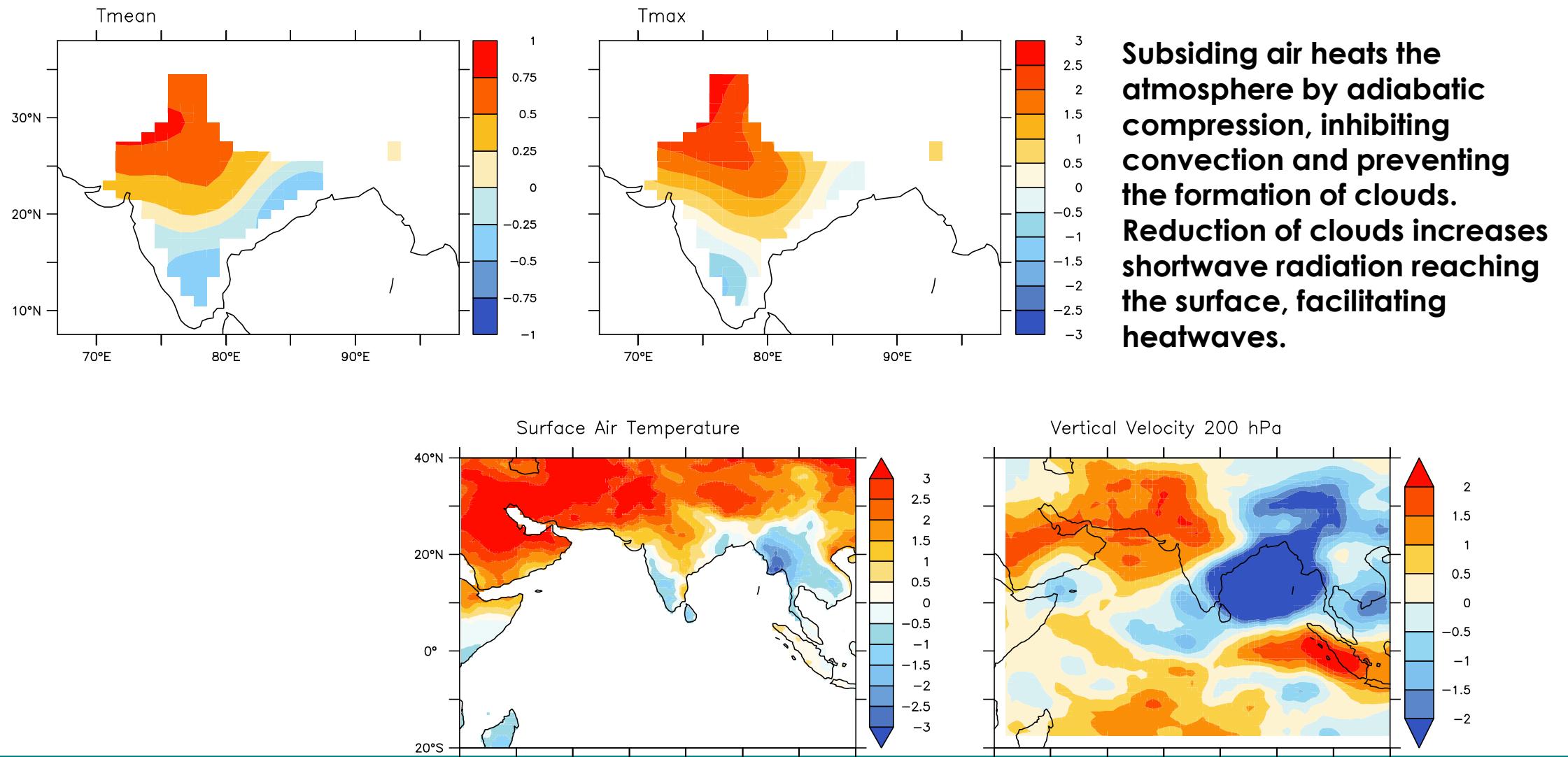
Enhanced updraft over BoB and subsidence over Indo-Pak



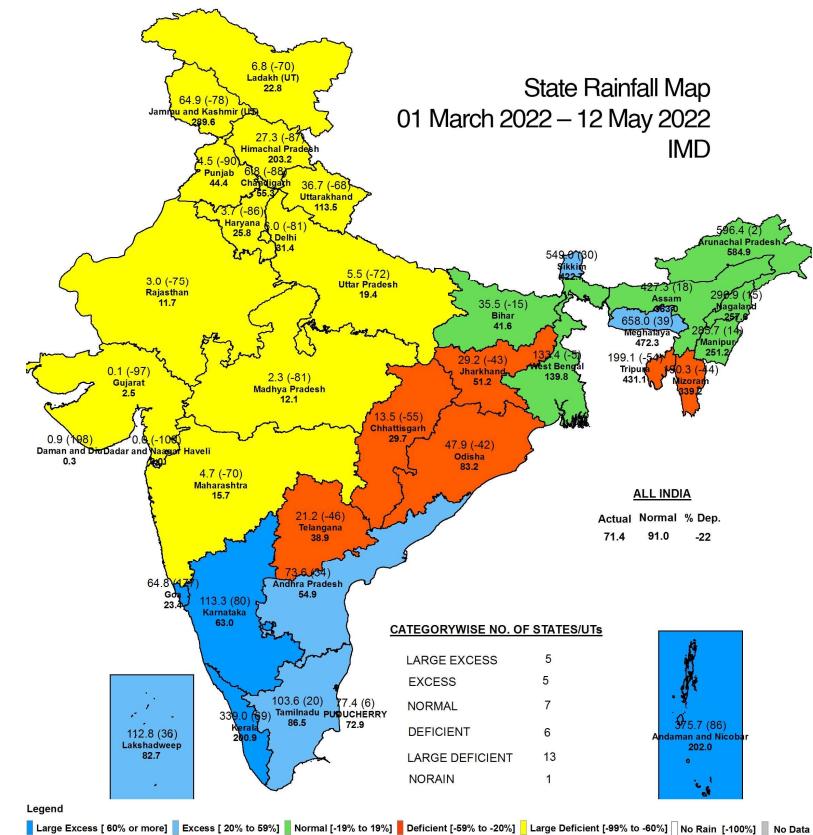
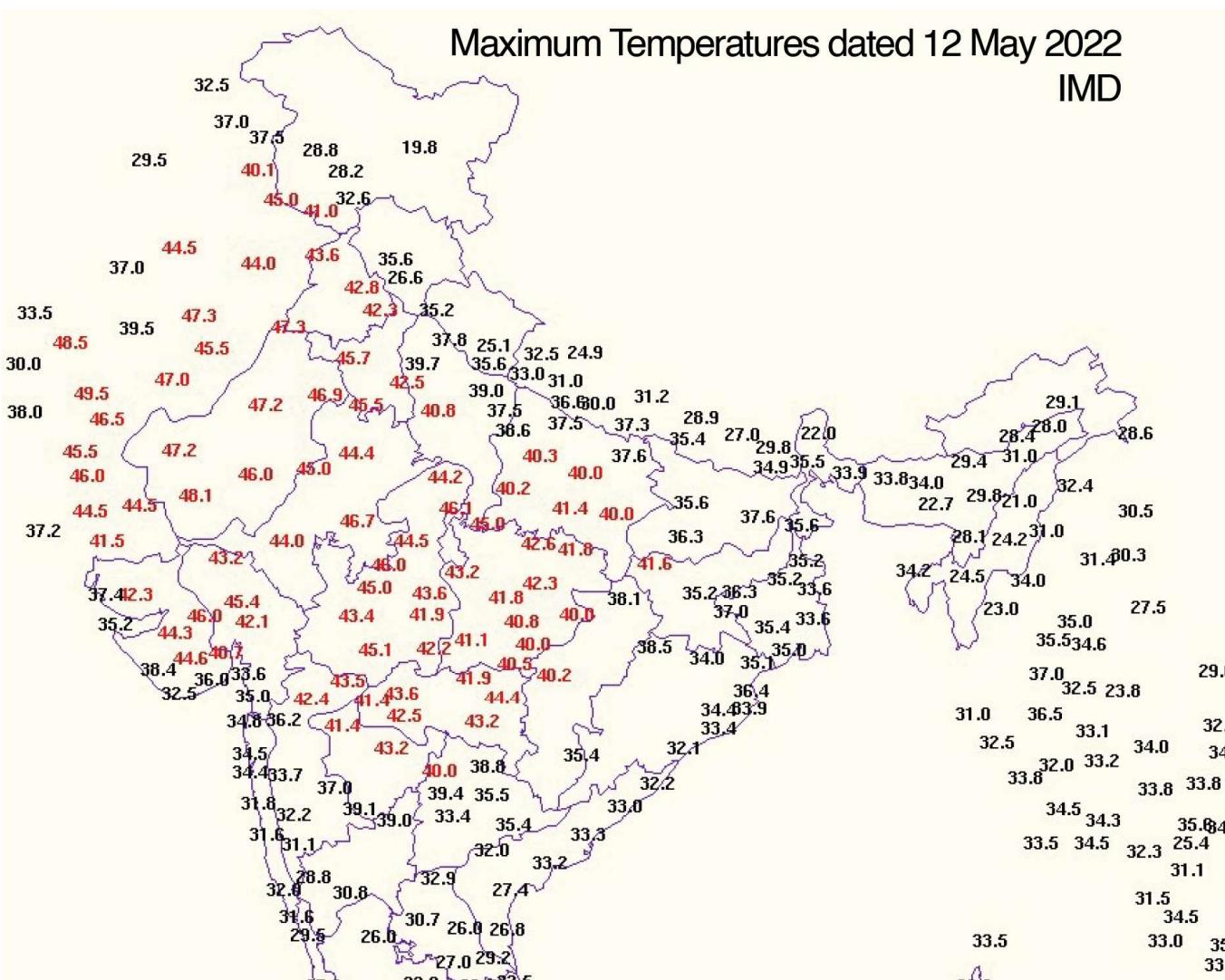
The spatial distribution of the circulation cells show enhanced updraft over the Bay of Bengal and subsidence over the Indo-Pak region

Subsiding air heats the atmosphere by adiabatic compression, inhibiting convection and preventing the formation of clouds. Reduction of clouds increases shortwave radiation reaching the surface, facilitating heatwaves.

Temperature anomalies following BoB Cyclones

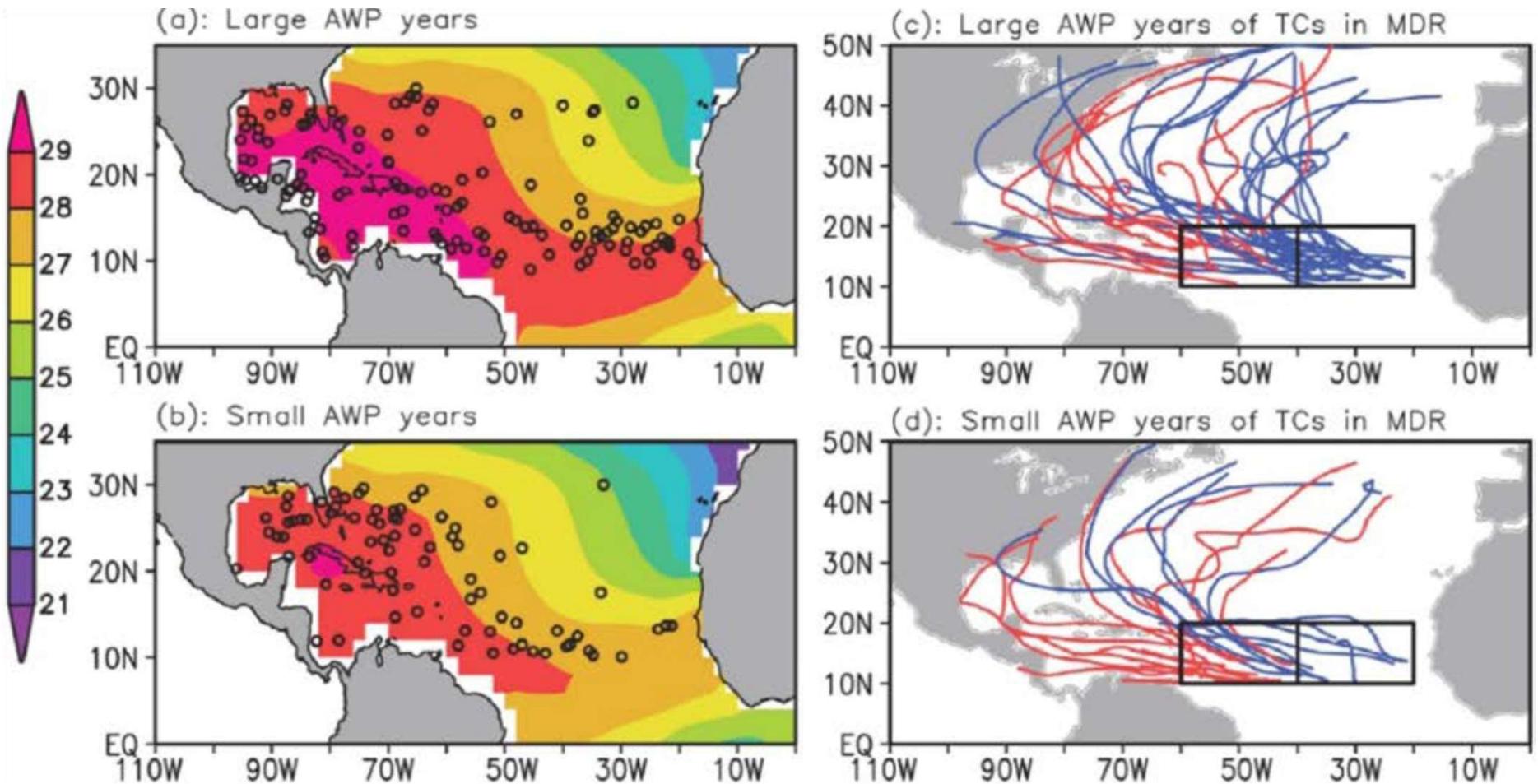


Heat and Rainfall Deficits following BoB Cyclone Asani



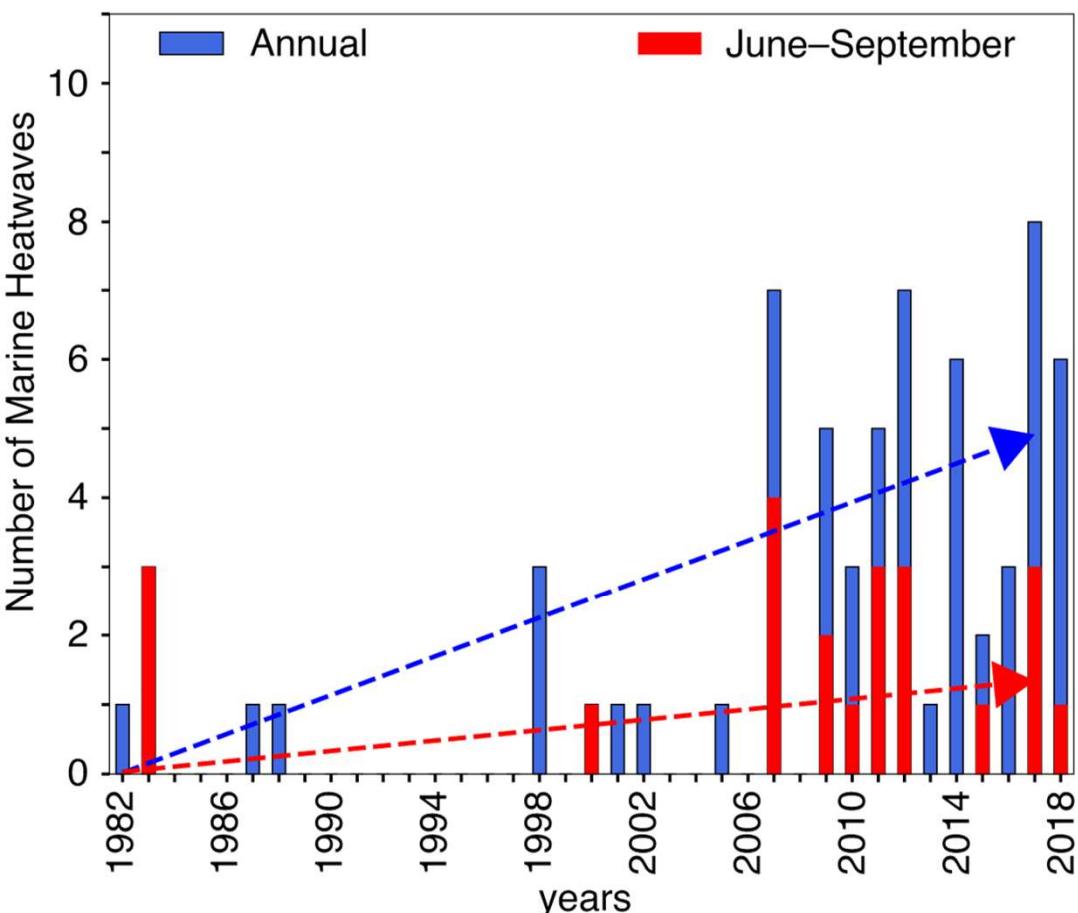


Atlantic Warm Pool and Cyclones

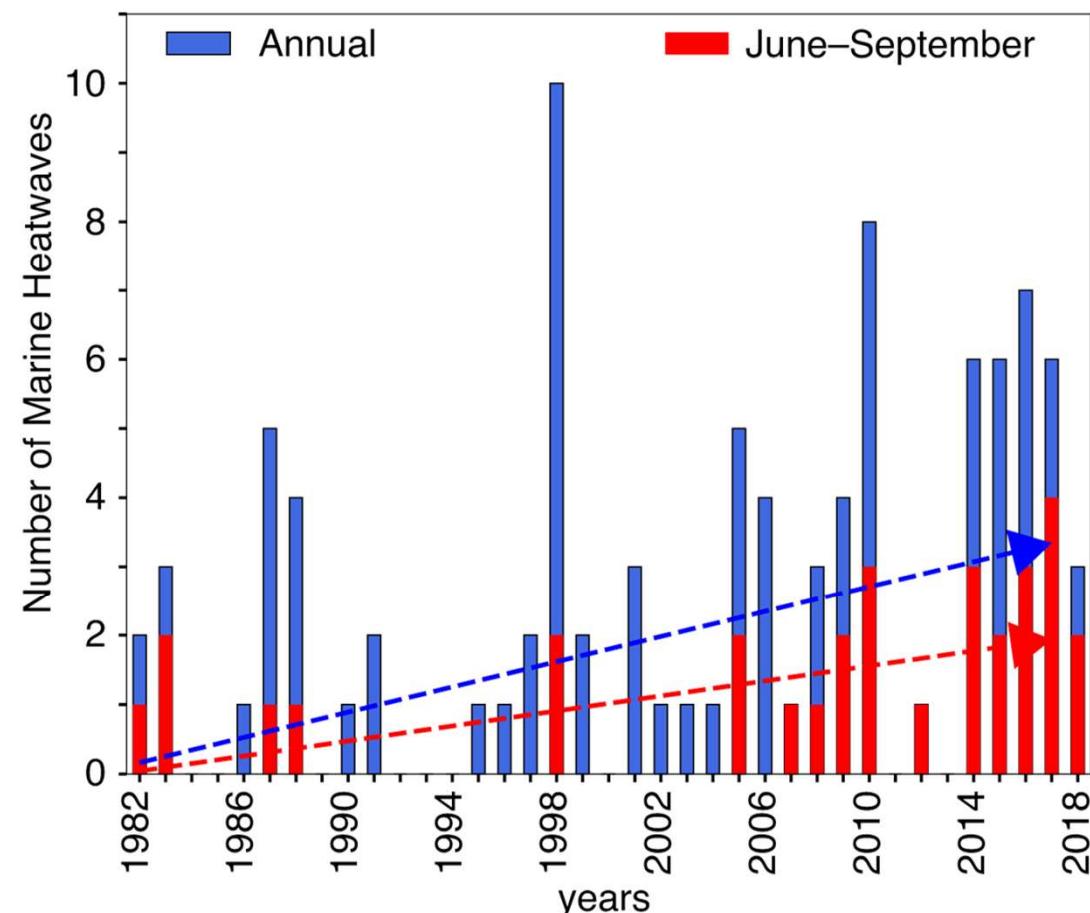


Marine Heatwaves and the Monsoon

a Marine Heatwaves in the western Indian Ocean



b Marine Heatwaves in the north Bay of Bengal

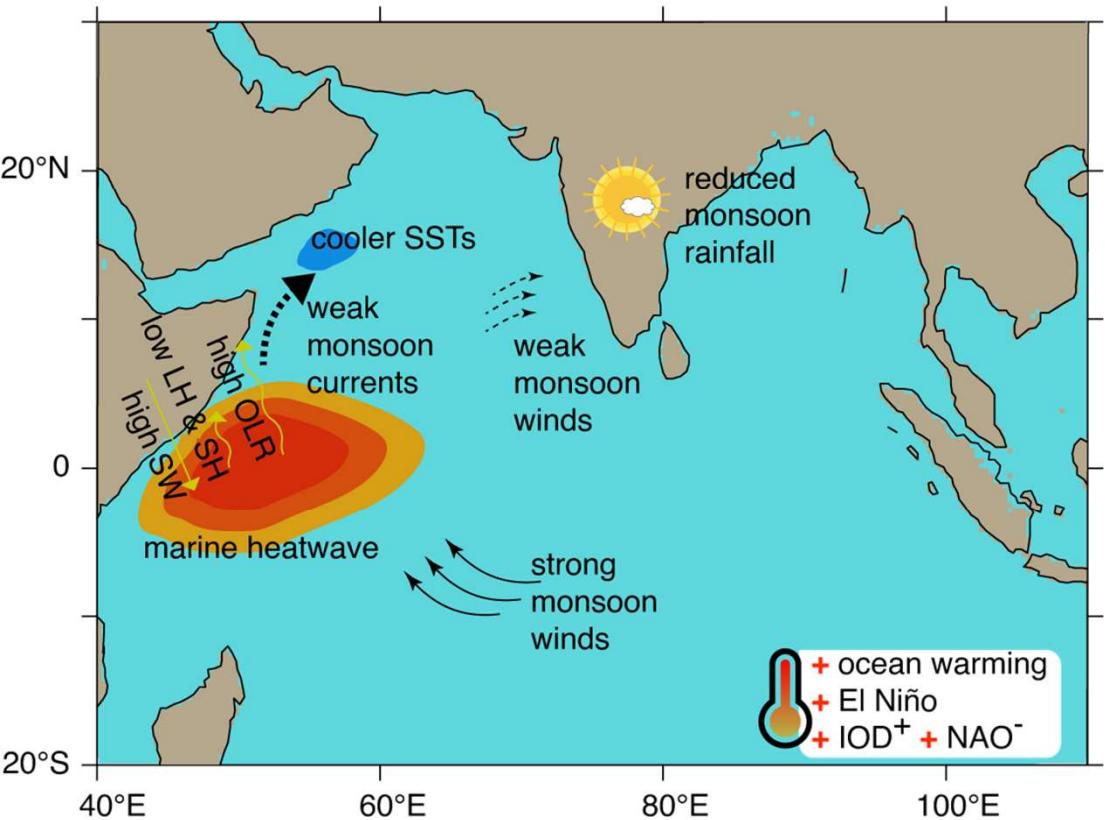


Marine Heatwaves and the Monsoon

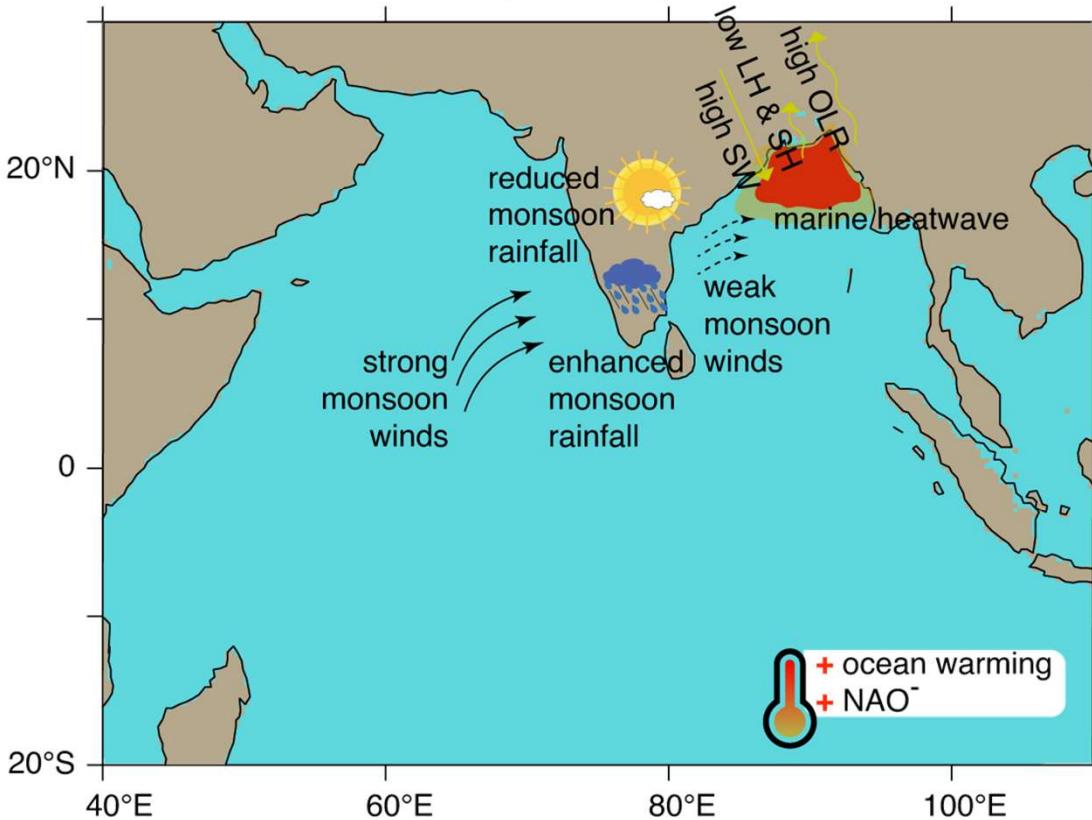
Marine heatwaves in the Indian Ocean and their impact on the monsoon

Saranya et al.
JGR Oceans, 2022

MHWs in the western Indian Ocean



MHWs in the north Bay of Bengal



Indian Ocean gearing up for a near-Permanent MHW state

