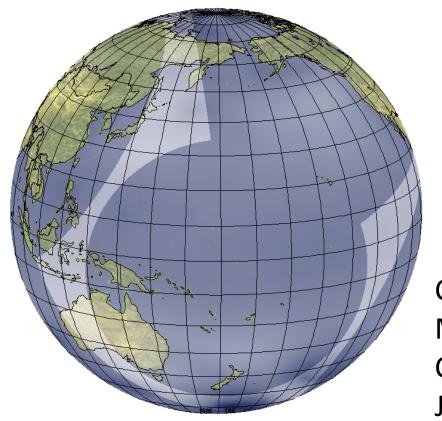
The Global Environmental Multi-scale model (GEM) on the Yin-Yang grid system

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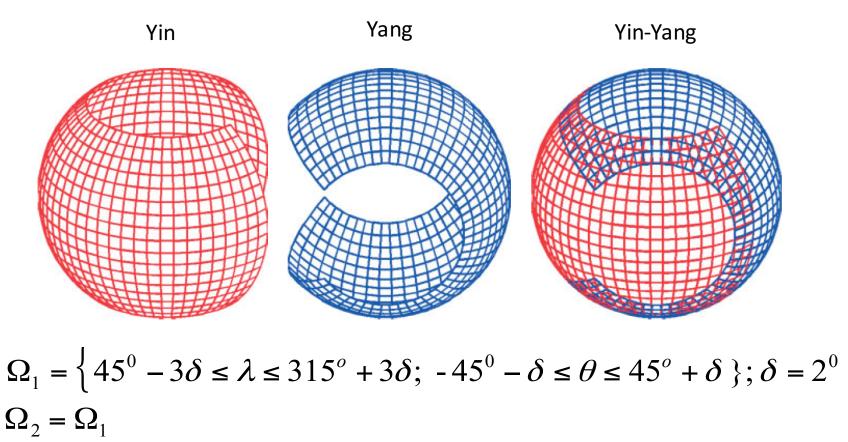


Collaboration with: Monique Tanguay, Claude Girard and Jean de Grandpré

Outline

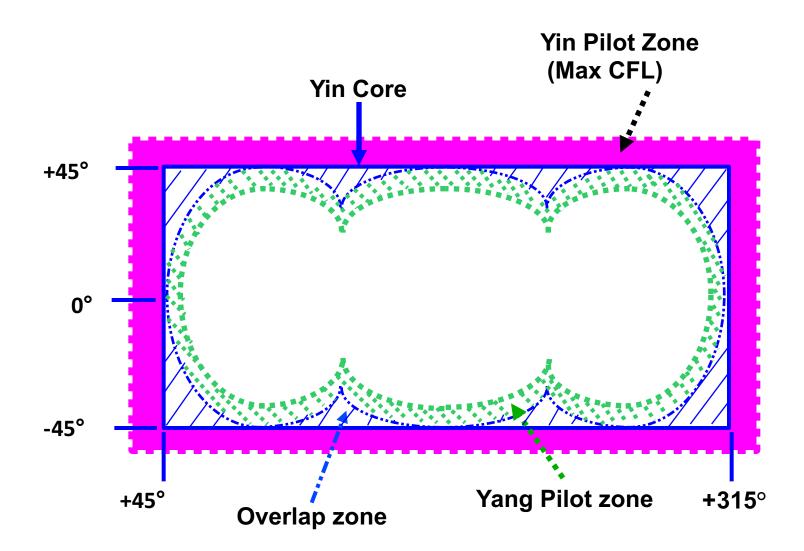
- Yin-Yang grid
- GEM model equations
- Domain decomposition method used to solve GEM model equations on Yin-Yang grid
- Superstorm Sandy

GEM Yin-Yang



The global forecast is based on the two-way nesting method between 2-limited area models.

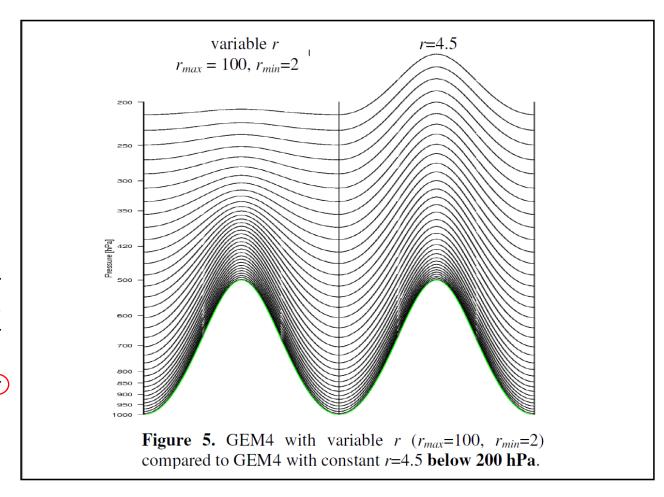
Qaddouri and Lee 2011 Q.J.R.Meteorol.Soc.137:1913-1926



Vertical coordinate ζ : A log-hydrostatic-pressure type

$$\ln(\pi) = \zeta + B(\zeta)s$$
; $B(\zeta) = \left(\frac{\zeta - \zeta_T}{\zeta_S - \zeta_T}\right)^r$; $s = \ln(\frac{\pi_S}{p_0})$; $p_0 = 1000 \,\text{hPa}$

Smooth transitions from terrain following levels near the ground to isobaric surfaces in the upper troposphere with appropriate choice of parameter r



GEM non-hydrostatic equations

Momentum Eqs.

Continuity Eq.

Thermodynamic Eq.

$$\frac{d\mathbf{V}_{h}}{dt} + f\mathbf{k}\mathbf{x}\mathbf{V}_{h} + R_{d}T_{v}\nabla_{\xi} \ln p + (1+\mu)\nabla_{\xi}\phi = 0$$

$$\frac{dw}{dt} - g\mu = 0$$
Non-hydrostatic
$$\frac{d}{dt}\ln\left(\pi\frac{\partial\ln\pi}{\partial\xi}\right) + \nabla_{\xi}\cdot\mathbf{V}_{h} + \frac{\partial\dot{\xi}}{\partial\xi} = 0$$

$$\frac{d\ln T_{v}}{dt} - \kappa\frac{d\ln p}{dt} = 0$$

$$\frac{d\phi}{dt} - gw = 0$$

$$R_{d}T_{v} + \frac{p}{\pi}\frac{\partial\phi}{\partial\ln\pi} = 0$$

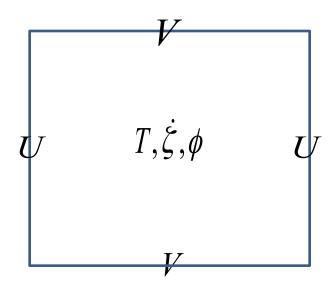
$$1 + \mu - \frac{p}{\pi}\frac{\partial\ln p}{\partial\ln\pi} = 0$$

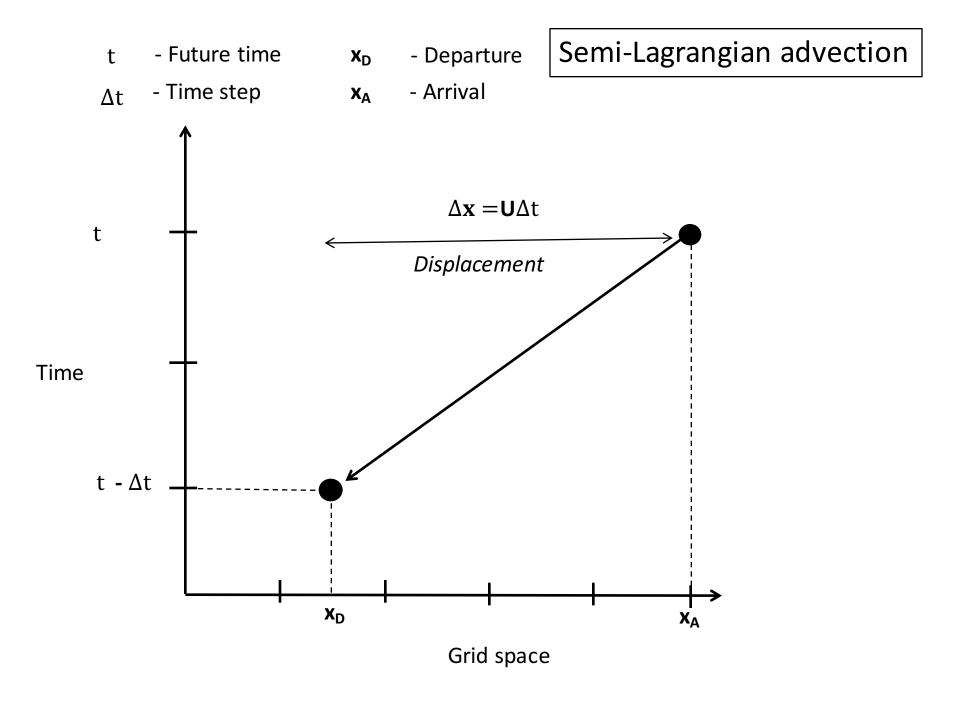
 $\ln \pi = \zeta + Bs$

Vertical discretization: Charney-Phillips

Momentum levels	Charney - Phillips Grid	Thermodynamic levels
0	$\phi_T, q_T - \zeta_T$	1/2
1 -	$T, w, \dot{\zeta}$	3/2
2 -	V_h, ϕ, q	
-	$T, w, \dot{\zeta}$	5/2
	V_h, ϕ, q	
-	$T, w, \dot{\zeta}$	
N-1 -	$T, w, \dot{\zeta}$	
N -	V_h, ϕ, q	 N.1/
N+1	ϕ_s, q_s	$N + \frac{1}{2}$

Horizontal discretization: Arakawa-C grid





Two time-level semi-Lagrangian implicit scheme (1)

$$\Delta \mathbf{r}^{i} = \Delta t \left[b_{A} \mathbf{v}(\mathbf{r}, t) + (1 - b_{A}) \mathbf{v}(\mathbf{r} - \Delta \mathbf{r}^{i-1}, t - \Delta t) \right]$$
Estimate Displacement

Also Crank-Nicholson iterations needed to estimate $\mathbf{v}(\mathbf{r},t)$

Two time-level semi-Lagrangian implicit scheme (2)

$$\frac{dF_i}{dt} + G_i = 0$$

$$\frac{dF_i}{dt} \approx \frac{F_i^A - F_i^D}{\Delta t} \qquad G_i \approx b^A G_i^A + (1 - b^A) G_i^D$$

Semi-Lagrangian Advection

$$G_i \approx b^A G_i^A + (1 - b^A) G_i^D$$

Average

 $0.5 \le b^A \le 0.6$

Off-centering for stability

A:
$$(\mathbf{r},t)$$
 Arrival
 D: $(\mathbf{r} - \Delta \mathbf{r}, t - \Delta t)$ Departure

$$\frac{F_i^A - F_i^D}{\Delta t} + b^A G_i^A + (1 - b^A) G_i^D = 0$$

Two time-level semi-Lagrangian implicit scheme (3)

Separating the time levels

$$\frac{F_i^A}{\tau} + G_i^A = \frac{F_i^D}{\tau} - \beta G_i^D \equiv R_i$$

LEFT-HAND side

Time t (Unknown)

RIGHT-HAND side

Time t- (Known)
Interpolated at D

 $(\tau = \Delta t b^A; \beta = (1 - b^A)/b^A)$

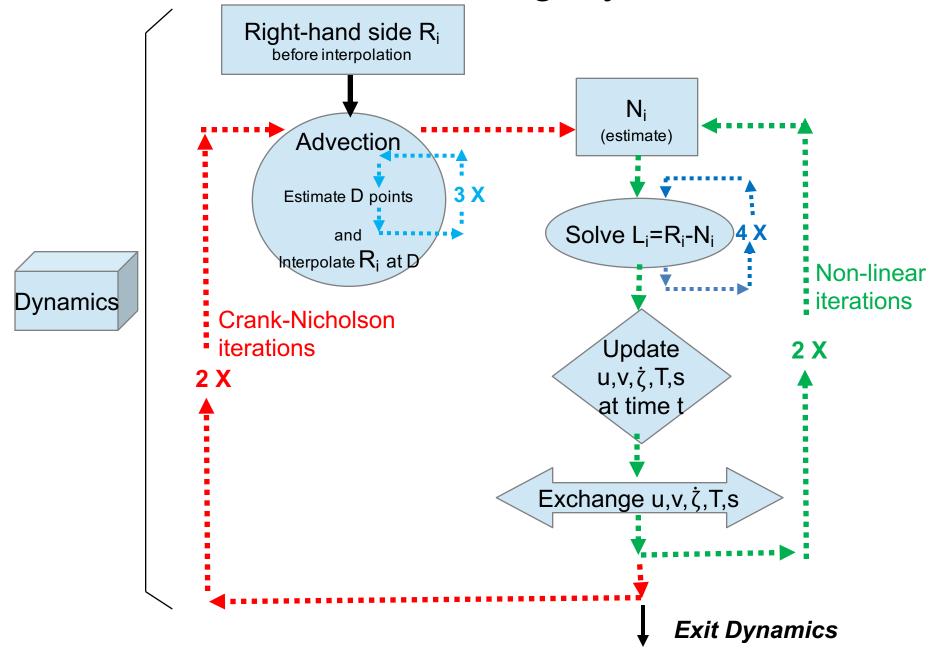
Since LEFT-HAND side is too difficult to solve,

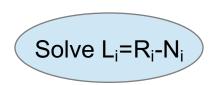
we decompose it into L (Linear) and N (residual Non-linear) parts

$$\frac{F_i^A}{\tau} + G_i^A = L_i + N_i = R_i$$

Solve the LEFT-HAND side $L_i = R_i - N_i$ Non-linear iterations since N_i is estimated

Flow of GEM Yin-Yang Dynamic core





Elliptic problem on Yin-Yang grid with Schwarz method

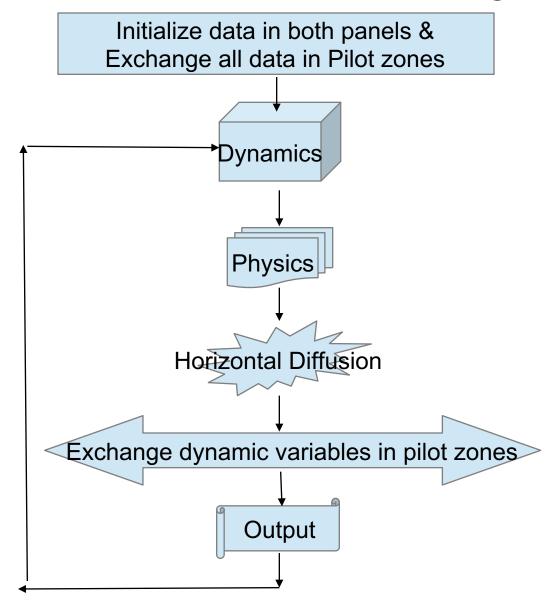
Global solution is obtained by solving iteratively 2 elliptic sub-problems (Yin/Yang)

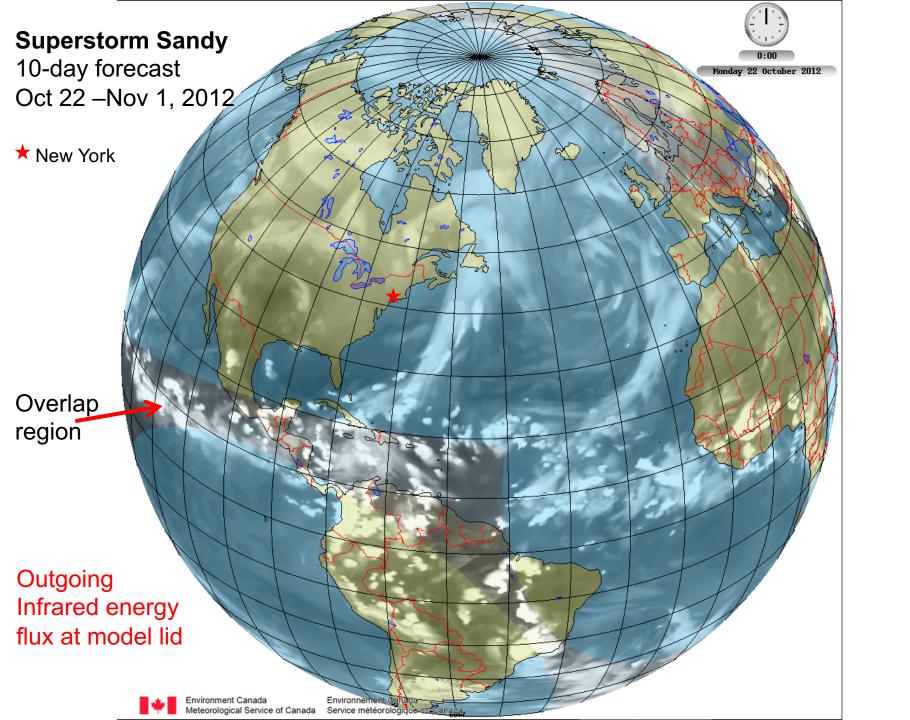
- 1) receive Boundary conditions (BCs) from the other panel; solve local elliptic problem.
- 2) if boundary conditions converge, stop; else send BCs to other panel; goto 1

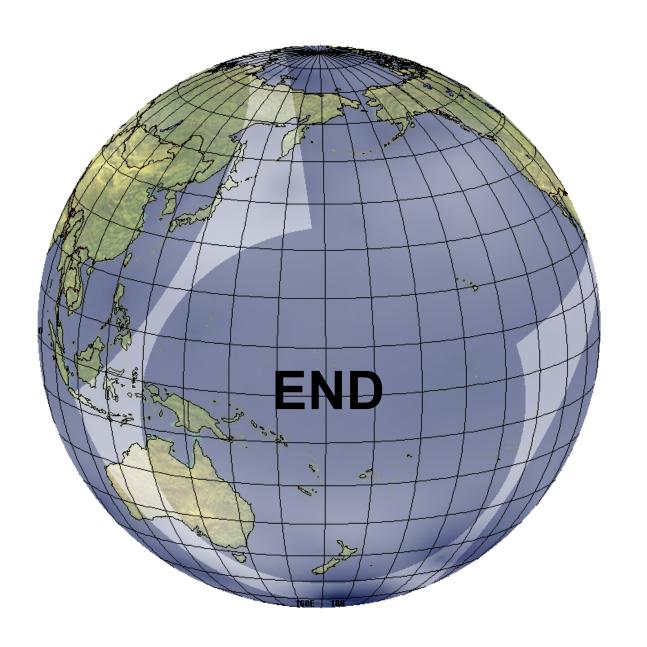
eg: GEM Yin-Yang 25km needs only 4 iterations for convergence with 2 degrees overlap

Qaddouri et al. 2008 Appl.Numerical.Math.,58,4,459-471

Flow of GEM Yin-Yang







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