

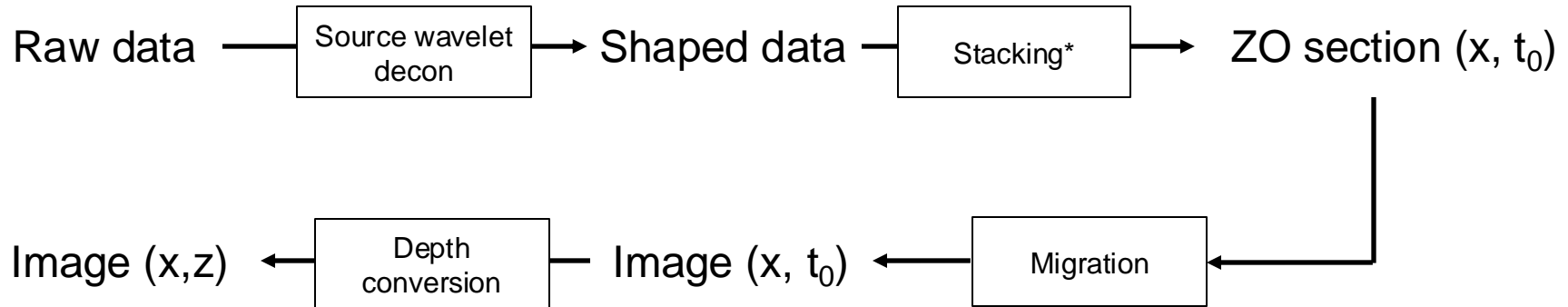
11. Seismic processing

M. Ravasi

ERSE 210 Seismology

Basic Seismic Processing Flow

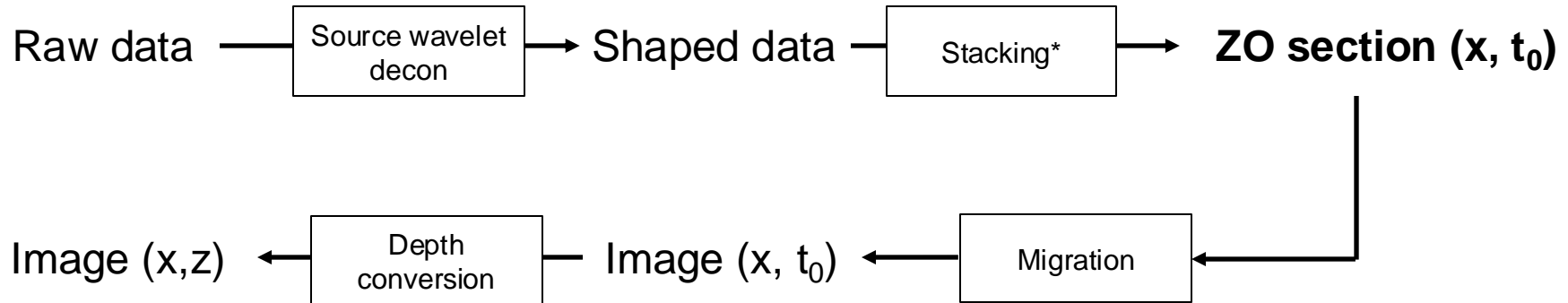
Historically, seismic data processing was composed of a few simple steps, mostly aimed at **stacking** traces to enhance the SNR of the recorded data.



* This also includes a step of velocity estimation

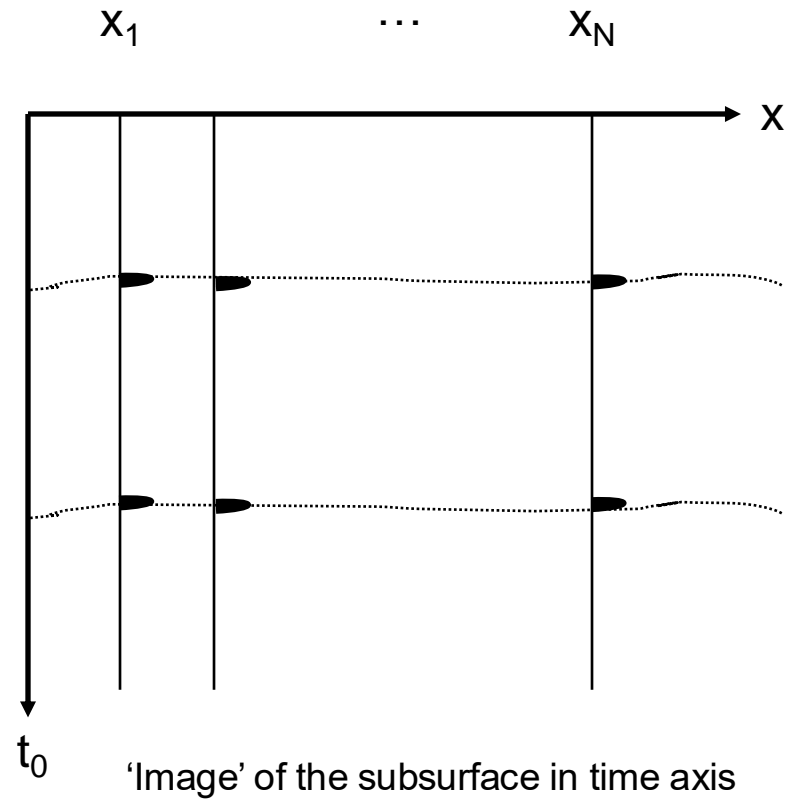
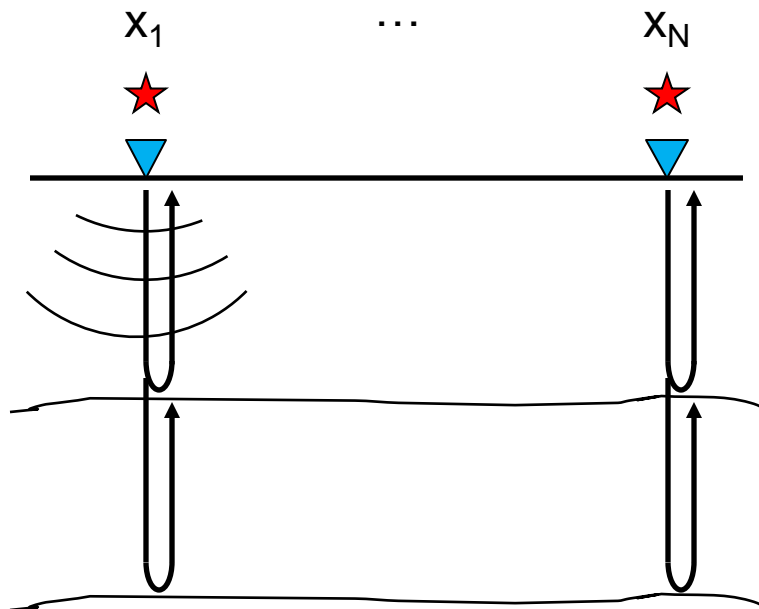
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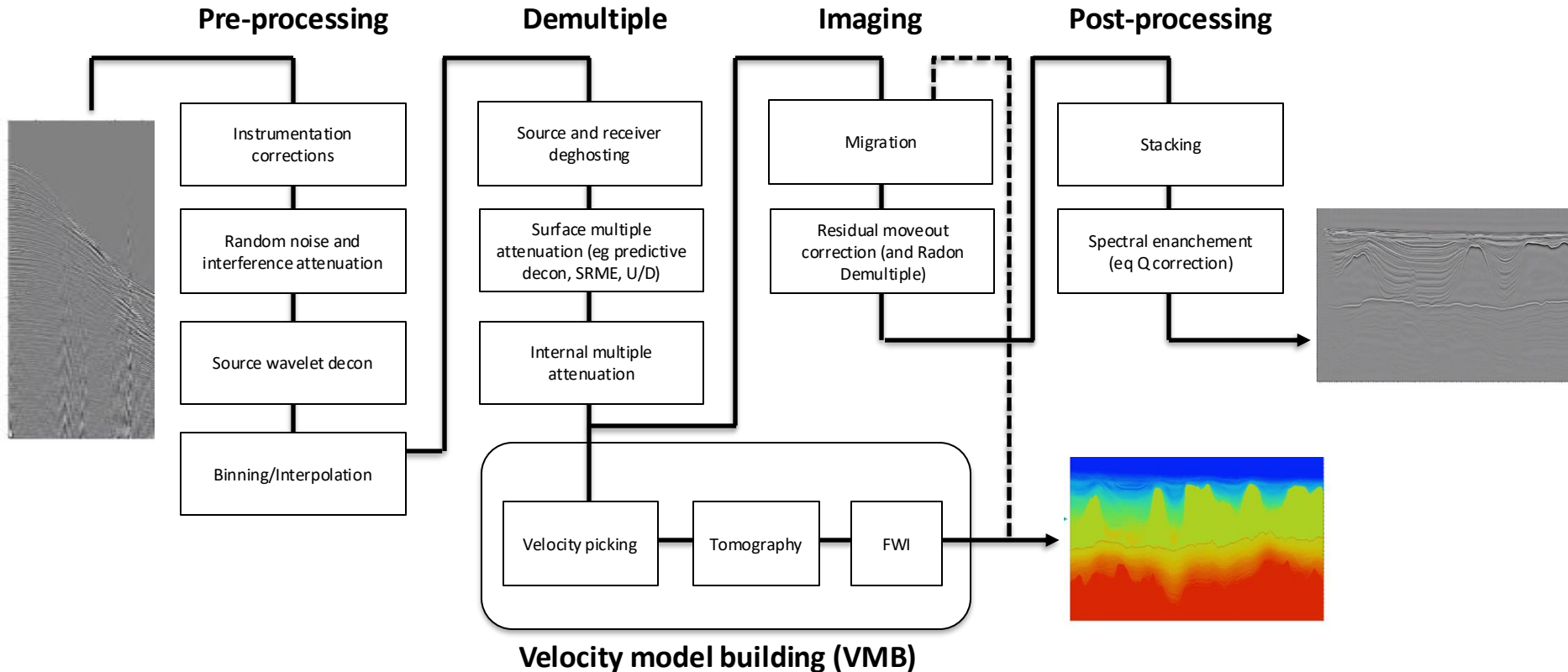


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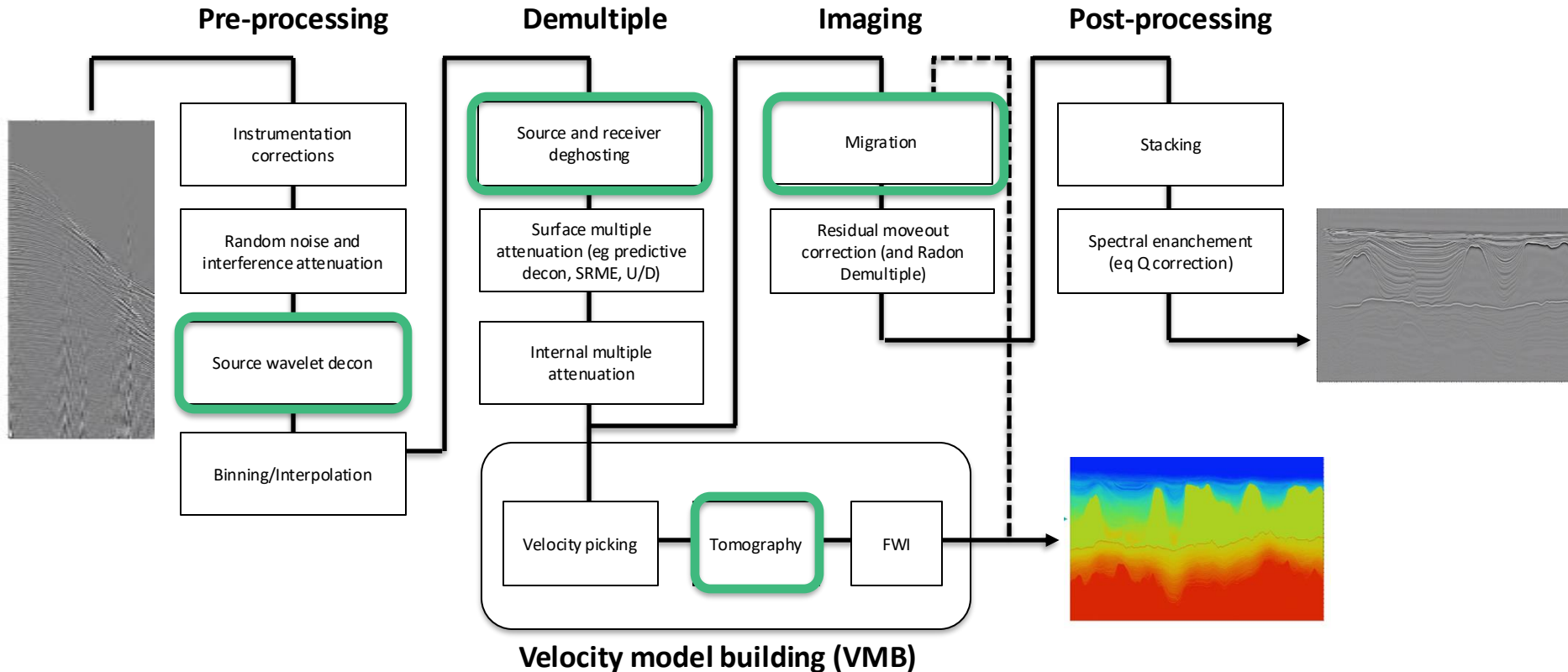
Zero-offset section



Modern Seismic Processing flow

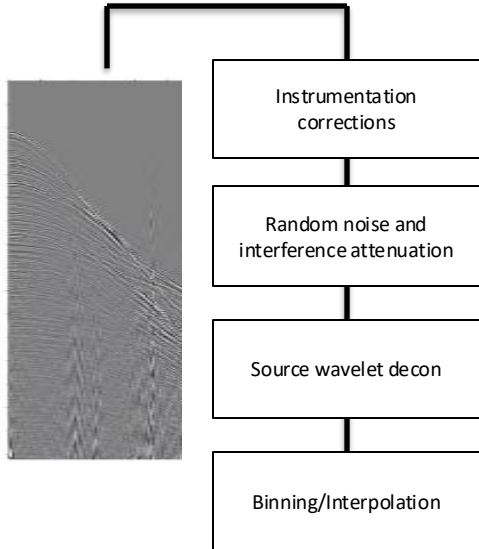


Modern Seismic Processing flow

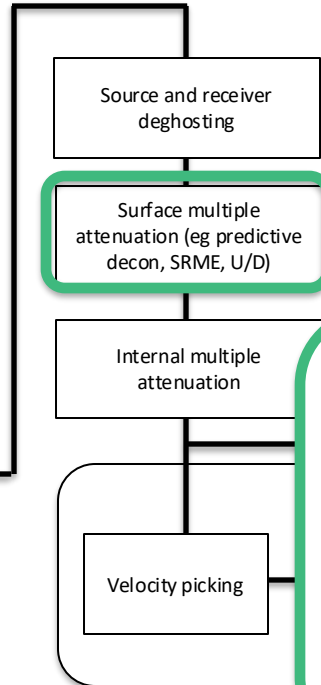


Modern Seismic Processing flow

Pre-processing

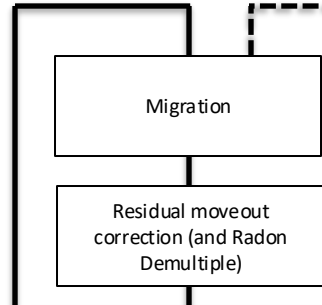


Demultiple

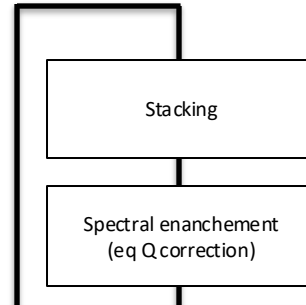


Velocity model

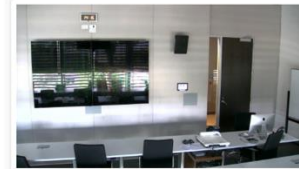
Imaging



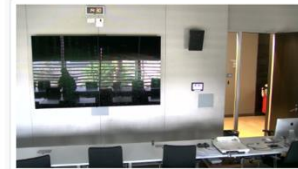
Post-processing



2 Lectures from Dr. Eric Vershuur available in Blackboard

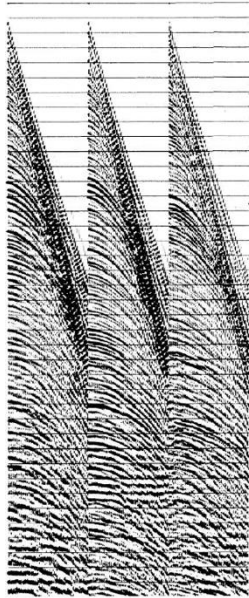


1h 46m
ErSE 210 Seismology - Lecture_15 - 24-2025
Nov 3, 2022 2:16 PM

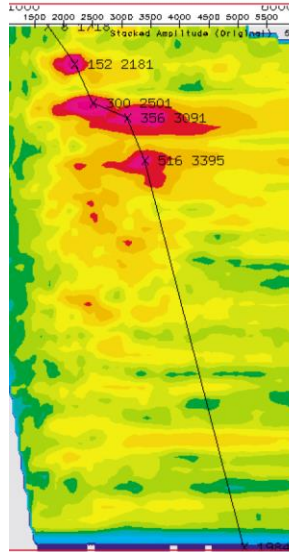


1h 24m
ErSE 210 Seismology - Lecture_14 - 24-2025
Oct 31, 2022 2:30 PM

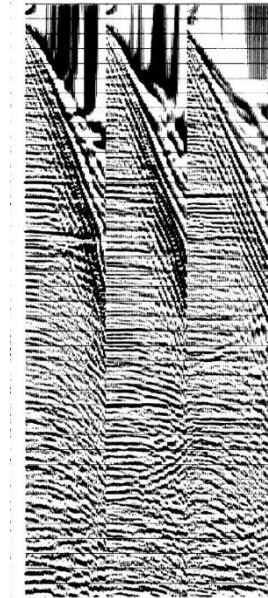
Common midpoint stacking = NMO analysis



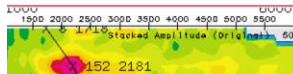
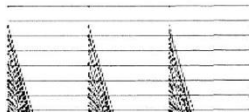
Velocity analysis



NMO Correction

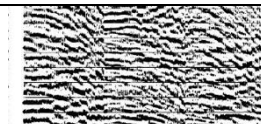
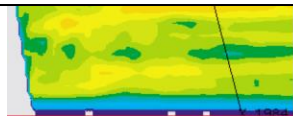
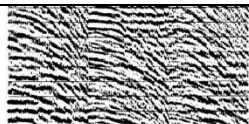


Common midpoint stacking = NMO analysis

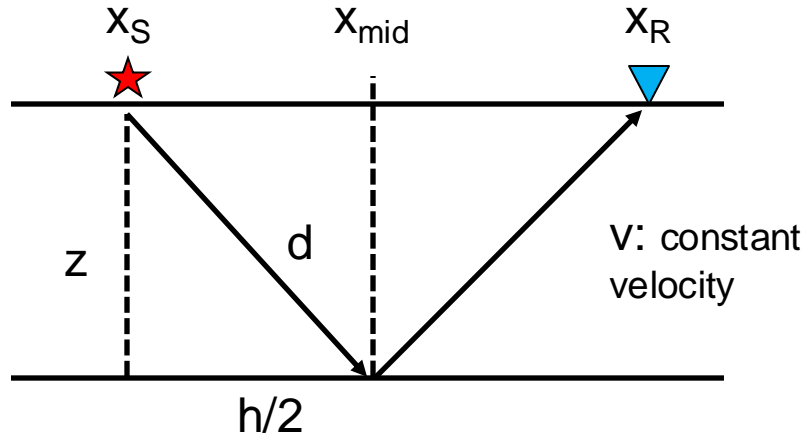


Simple procedure to:

- Increase SNR of seismic data
- Obtain a first estimate of the velocity model of the subsurface (to be later refined by tomographic inversion / FWI methods)



Normal moveout



- If $x_S = x_R$ Two-way traveltimes



$$d_0 = z \rightarrow t_0 = 2z/v$$

- If $x_S \neq x_R$

$$d = \sqrt{z^2 + \frac{h^2}{4}} \rightarrow t = 2 \sqrt{\frac{z^2}{v^2} + \frac{h^2}{4v^2}}$$

$$= \sqrt{t_0^2 + \frac{h^2}{v^2}}$$

Normal moveout

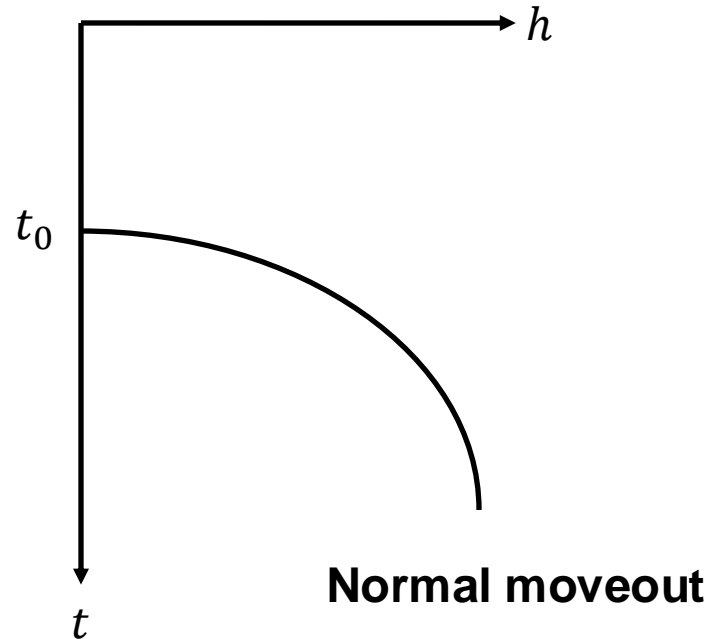
Rearranging the travelttime equation:

$$t^2 = t_0^2 + \frac{h^2}{v^2}$$



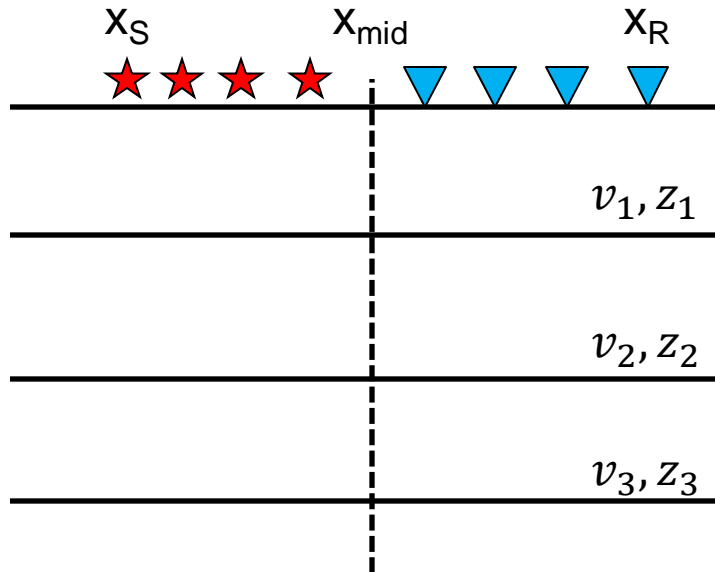
$$t^2 - \frac{h^2}{v^2} = t_0^2$$

Hyperbola in h-t axes



Normal moveout

This equation is not only valid for a single layer of constant velocity; a simple extension exists for a stack of N layers:



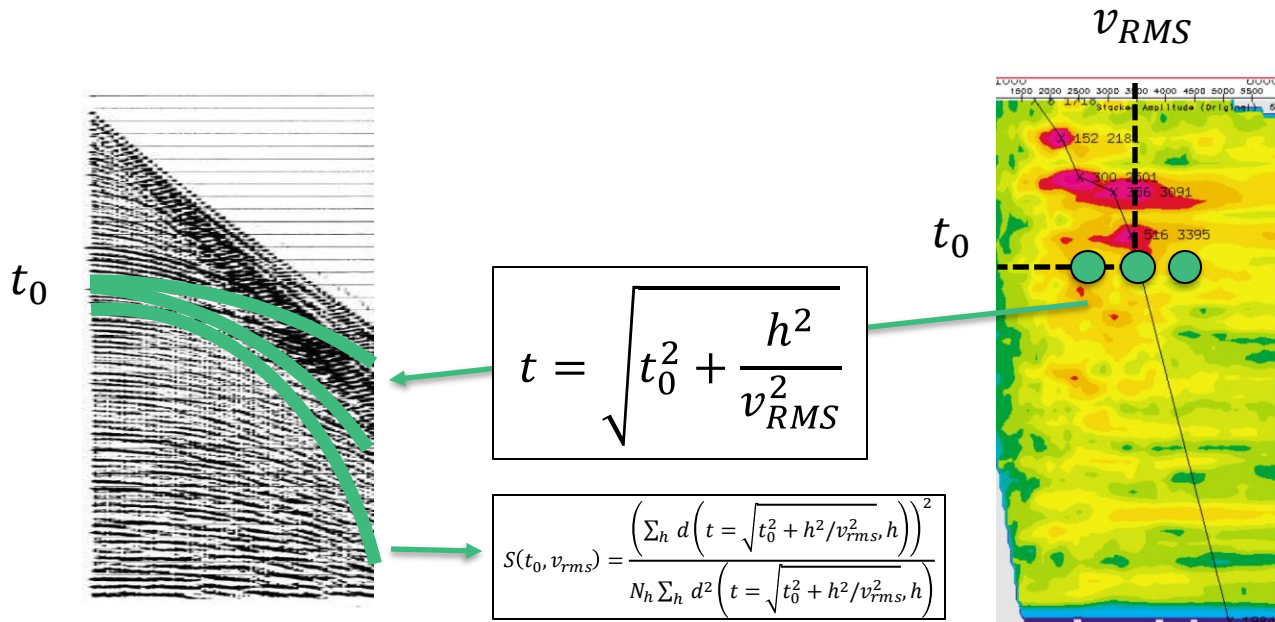
$$t^2 = t_0^2 + \frac{h^2}{v_{RMS}^2}$$

Root-Mean Squared Velocity

$$v_{RMS}^2 = \frac{1}{t_0} \sum_i v_i^2 \Delta\tau_i$$

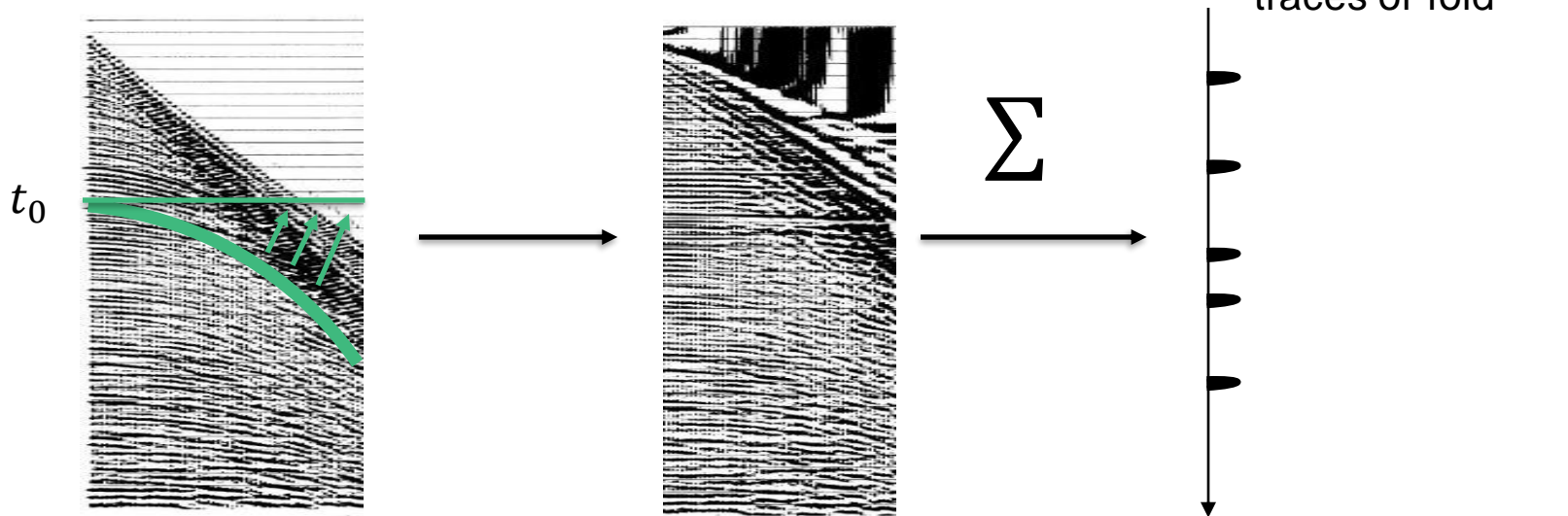
$$\Delta\tau_i = \frac{2z_i}{v_i}, t_0 = \sum_i \Delta\tau_i$$

Semblance analysis



Fit all possible parametric curves with (t_0, v_{RMS}) and 'sum'

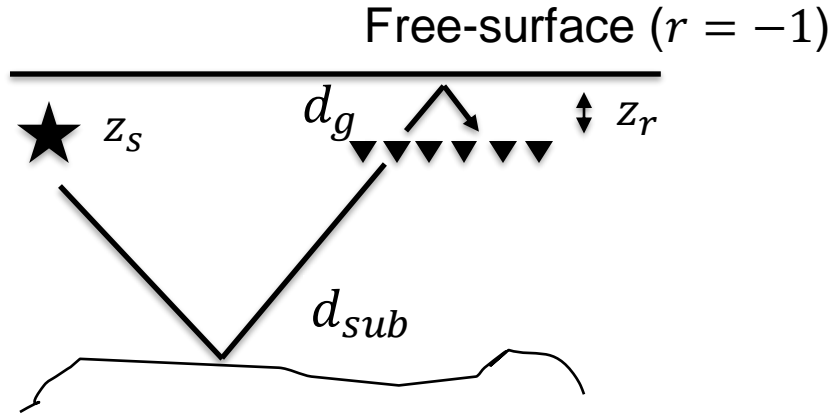
NMO correction and CMP stacking



Flatten CMP gather and sum along best fitting line $\tilde{t} = \sqrt{\tilde{t}_0^2 + \frac{h^2}{\tilde{v}_{RMS}^2}}$

Deghosting

Marine seismic data are acquired placing sources and receivers below the water-air interface (= free surface); this creates a **ghost** effect in the data



$$\begin{aligned}d(t) &= d_{sub}(t) + d_{ghost}(t) \\&= d_{sub}(t) + r d_{sub}(t - t_g) \\&= d_{sub}(t) - d_{sub}(t - t_g)\end{aligned}$$

$$t_g \approx 2z_R/v_{water} \text{ at zero incidence}$$

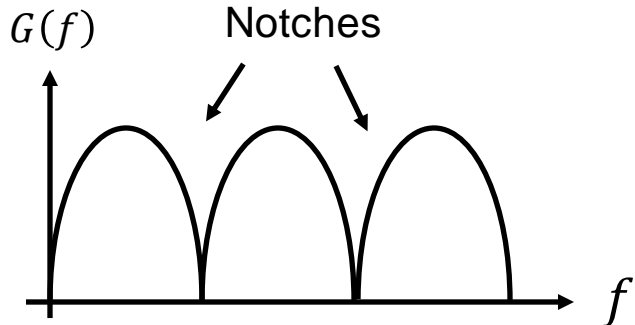
Deghosting

The effect of the ghost arrival on the recorded seismic data is better understood in the frequency domain:

$$d(t) = d_{sub}(t) - d_{sub}(t - t_g) \longleftrightarrow D(f) = D_{sub}(f) - D_{sub}(f)e^{-j2\pi f t_g}$$

$$D(f) = D_{sub}(f) \underbrace{[1 - e^{-j2\pi f t_g}]}_{G(f)}$$

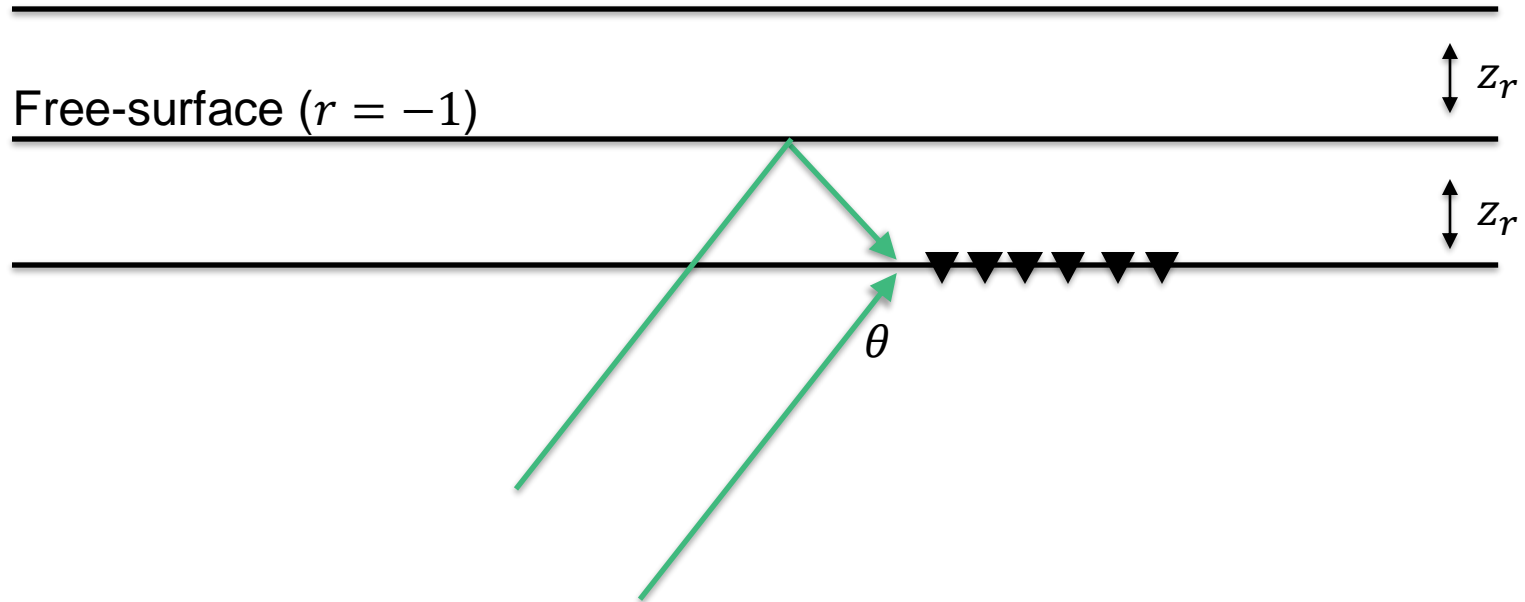
$G(f)$: ghost model



$$f_{notch} = \frac{nv_{water}}{2z_r} \quad \text{where} \quad \frac{2\pi f 2z_r}{v_{water}} = 0 + 2\pi n$$

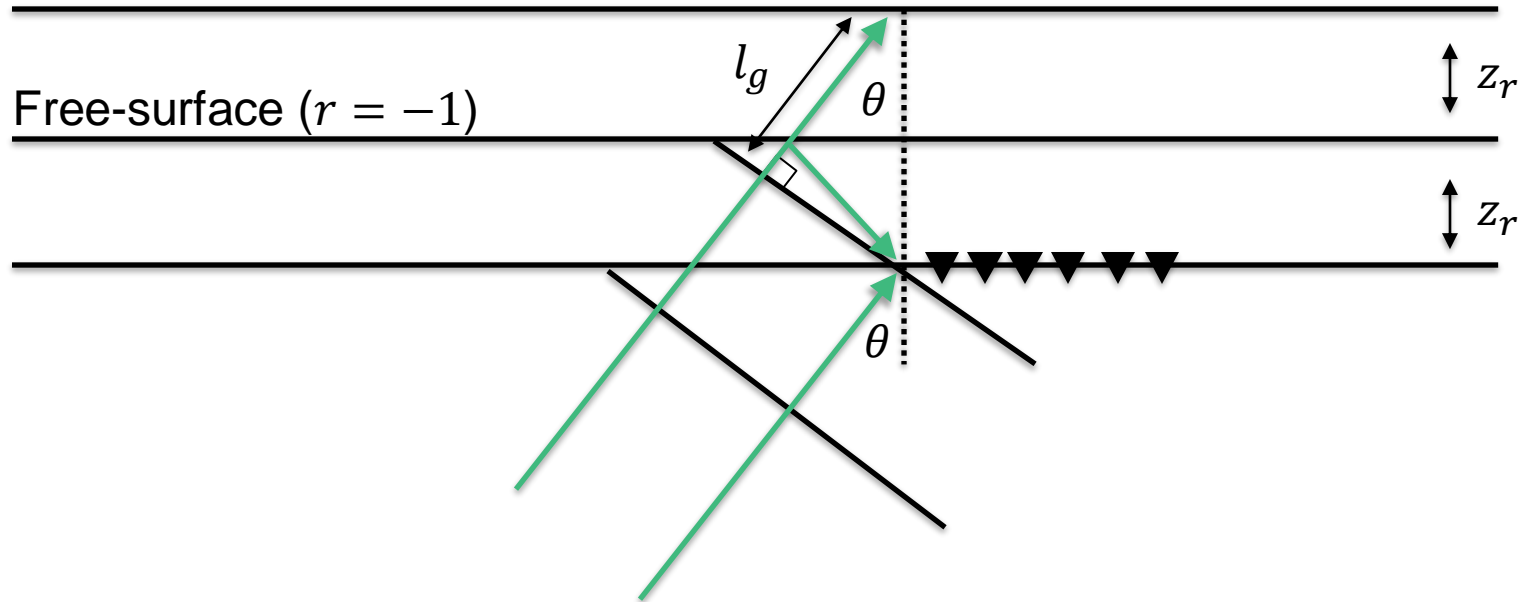
Deghosting

We can model the ghost more accurately by considering the angle of incidence θ



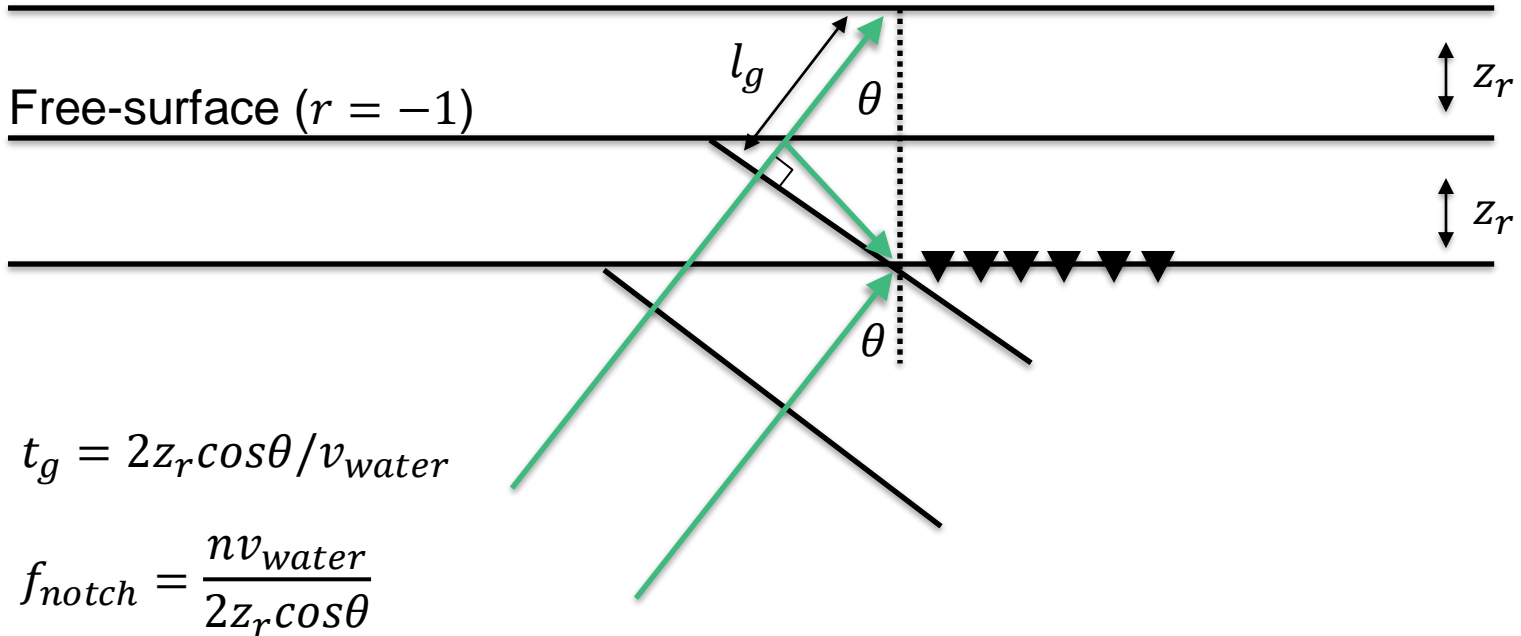
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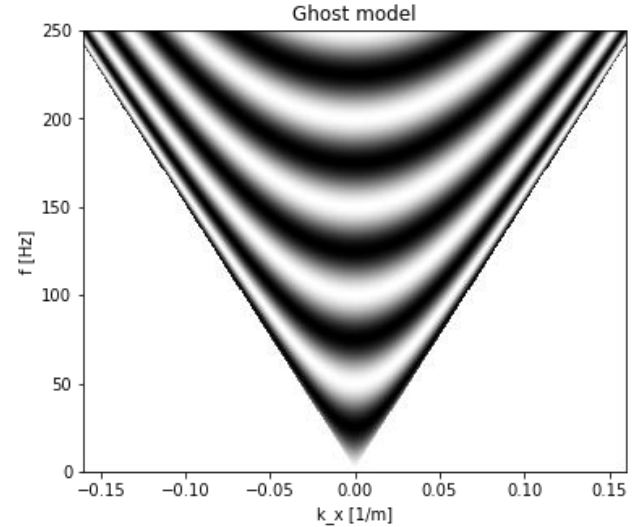


Deghosting

We can model the ghost more accurately by considering the angle of incidence θ

$$f_{notch} = \frac{nv_{water}}{2z_r \cos\theta}$$

$$D(f, k_x) = D_{sub}(f, k_x) [1 - e^{-j2\pi k_z(2z_r)}]$$



Deghosting

Deghosting is the process of ‘deconvolving’ the ghost model from data:]

$$D_{sub}(f, k_x) \approx \frac{D(f, k_x)}{[1 - e^{-j2\pi k_z(2z_r)}]}$$

Since this process allows recovering some frequencies that the free-surface had suppressed (ghost notches), **broadband seismic** refers to the combination of acquisition and processing strategies to remove/attenuate ghost effects.

A similar effect also exist on the source side (source-deghosting).