

# Process Rate Estimator

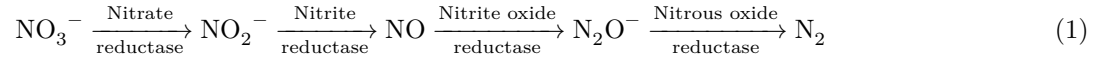
A modeling side-hustle for the ETH group sustainable agroecosystems

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## 1 Introduction

Denitrification is the natural process by which nitrate ( $\text{NO}_3^-$ ) in the soil are converted by bacteria into nitrous oxide ( $\text{N}_2\text{O}$ ) or pure nitrogen ( $\text{N}_2$ ). The latter is called *total denitrification* — the full process described in Equation 1 takes place.



Denitrification occurs in conditions where oxygen is limited, such as waterlogged soils. It is part of the nitrogen cycle, where nitrogen is circulated between the atmosphere, organisms and the earth.

## 2 Formal model description

### 2.1 Model parameters

Table 1: Overview of the parameters used in the model.

Symbol	Code	Name	Value	Unit
$BD$	BD	Bulk density (mass of the many particles of the material divided by the bulk volume)	1.686	g cm <sup>-3</sup>
$\theta_w$	theta_w	Soil volumetric water content		
$\theta_a$	theta_a	Air-filled porosity	$1 - \frac{\theta_w}{\theta_t}$	
$\theta_t$	theta_t	Total soil porosity	$1 - \frac{BD}{2.65}$	
T	temperature	Soil temperature	298	K
$D_s$	D_s	Gas diffusion coefficient	Equation 3	m <sup>2</sup> s <sup>-1</sup>
$D_{fw}$	D_fw	Diffusivity of $\text{N}_2\text{O}$ in water	Equation 5	
$D_{fa}$	D_fa	Diffusivity of $\text{N}_2\text{O}$ in air	Equation 6	
$D_{fa,NTP}$		Free air diffusion coefficient under standard conditions	Equation 6	
$n$	n	Empirical parameter (1)	1.81	
$H$	H	Dimensionless Henry's solubility constant	Equation 4	
$\rho$	rho	Gas density of $\text{N}_2\text{O}$	$1.26 \times 10^6$	mg m <sub>-3</sub>

The diffusion fluxes between soil increments are described by Frick's law (Equation 2).

$$F_{\text{calc}} = \frac{dC}{dZ} D_s \rho \quad (2)$$

Here,  $D_s$  is the gas diffusion coefficient,  $\rho$  is the gas density of  $\text{N}_2\text{O}$ , and  $\frac{dC}{dZ}$  is the  $\text{N}_2\text{O}$  concentration gradient from lower to upper depth. The fluxes are calculated based on  $\text{N}_2\text{O}$  concentration gradients between 105-135 cm, 75-105 cm, 45-75 cm, 15-45 cm, and 0-15 cm depth layers, and ambient air above the soil surface.

$\theta_w$  is the soil volumetric water content,  $\theta_a$  the air-filled porosity, and  $\theta_T$  is the total soil porosity.

The gas diffusion coefficient  $D_s$  was calculated according Equation 3 as established by Millington and Quirk in 1961 (2).

$$D_s = \left( \frac{\theta_w^{\frac{10}{3}} + D_{fw}}{H} + \theta_a^{\frac{10}{3}} \times D_{fa} \right) \times \theta_T^{-2} \quad (3)$$

Here,  $H$  represents a dimensionless form of Henry's solubility constant ( $H'$ ) for  $\text{N}_2\text{O}$  in water at a given temperature. The constant  $H$  for  $\text{N}_2\text{O}$  is calculated as follows:

$$H = \frac{8.5470 \times 10^5 \times \exp \frac{-2284}{T}}{R \times T} \quad (4)$$

Here,  $R$  is the gas constant, and  $T$  is the temperature ( $T = 298 \text{ K}$ ).

$D_{fw}$  was calculated according to Equation 5 as documented by Versteeg and Van Swaaij (1988) (3).

$$D_{fw} = 5.07 \times 10^{-6} \times \exp \frac{-2371}{T} \quad (5)$$

$$D_{fa} = D_{fa, \text{NTP}} \times \left( \frac{T}{273.15} \right)^n \times \left( \frac{101'325}{P} \right) \quad (6)$$

## 2.2 Smoothing curves

The  $\text{N}_2\text{O}$  concentration, site preference as well as  $\delta^{18}\text{O}$  are estimated as a function of time for every depth and every column, separately. To achieve this function approximation, Kernel Regression as implemented in `npreg` is used (4).

## 2.3 State function set

Still to do.

## 3 The data

The study uses data collected from a mesocosm experiment – i.e. an outdoor experiment that examines the natural environment under controlled conditions. The experiment was set up as a randomized complete block design, with 4 varieties and 3 replicates, using 12 non-weighted lysimeters. A non-weighted lysimeter is a device to measure the amount of water that drains through soil, and to determine the types and amounts of dissolved nutrients or contaminants in the water. Each lysimeter had five sampling ports with soil moisture probes and custom-built pore gas sample, at depths of 7.5, 30, 60, 90 and 120 cm below soil surface.

$$4 \times 3 \times 5 \times 161 = 9660 \quad (7)$$

Equation 7 shows how many observations we should expect to have. In reality, some observations are missing.

Code	Name	Description
day_column_depth	Combination	
date_R	Weird date	Year + DOY
column	Column	
depth	Measurement depth	
increment	?	
variety	Wheat variety	
moisture	Soil moisture	
concN03N	NO <sub>3</sub> -N concentration	Nitrate nitrogen concentration ([NO <sub>3</sub> ] = [NO <sub>3</sub> -N] * 4.43).
N03N_ha		
corrected.N2O		
corrected.CO2		
mgN2ONm3		
gN2ONha		
gCO2Cha		
CN		
d15Nbulk		
d15Nalpha		
d15Nbeta		
SP	Site preference	
d18O		Ratio of stable isotopes oxygen-18 ( <sup>18</sup> O) and oxygen-16 ( <sup>16</sup> O).

## References

1. Massman, W. A review of the molecular diffusivities of H<sub>2</sub>O, CO<sub>2</sub>, CH<sub>4</sub>, CO, O<sub>3</sub>, SO<sub>2</sub>, NH<sub>3</sub>, N<sub>2</sub>O, NO, and NO<sub>2</sub> in air, O<sub>2</sub> and N<sub>2</sub> near STP. *Atmospheric environment* **32**, 1111–1127 (1998).
2. Millington, R. & Quirk, J. Permeability of porous solids. *Transactions of the Faraday Society* **57**, 1200–1207 (1961).
3. Versteeg, G. F. & Van Swaaij, W. P. Solubility and diffusivity of acid gases (carbon dioxide, nitrous oxide) in aqueous alkanolamine solutions. *Journal of Chemical & Engineering Data* **33**, 29–34 (1988).
4. Hayfield, T. & Racine, J. S. [Nonparametric econometrics: The np package](#). *Journal of Statistical Software* **27**, 1–32 (2008).