

# Volatilization Alternative explanation

Transport processes, including volatilization, leaching, lateral mass flux are considered sequentially before degradation and directly after application. Sorption is considered within each process.

## Volatilization

Volatilization is assumed only to take place on the application day, and follows Leistra et al. [2001] by considering mass flux across two boundary layers. Mass flux across an air boundary  $J_{v,a}$  ( $\mu\text{g}/\text{m}^2\text{d}$ ) and mass flux across the top soil layer  $J_{v,s}$  ( $\mu\text{g}/\text{m}^2\text{d}$ ) are given by:

$$J_{v,air} = \frac{-(C_{g,0(t)} - C_{air})}{r_{air(t)}}$$

$$J_{v,soil} = \frac{-(C_{g,z0} - C_{g,0})}{r_{soil(t)}}$$

where,  $r_a$  ( $\text{d}/\text{m}$ ) is the air boundary resistance and  $r_s$  ( $\text{d}/\text{m}$ ) the top soil layer resistance before the pesticide reaches the soil surface from the center of the cell layer ( $0.5 \cdot D_{z0}$ ). Concentration gradients ( $\mu\text{g}/\text{m}^3$ ) are considered based on the concentration at the soil surface  $C_{g,0}$ , the air boundary layer  $C_{air}$  and at the center of top soil layer.

Resistance across the boundary air and soil layers are given respectively by:

$$r_{air(t)} = \frac{D_{z,air}}{D_{a(t)}}$$

$$r_{soil(t)} = \frac{0.5 \cdot D_{z0}}{D_{diff,g}}$$

where  $D_{z,air}$  ( $\text{m}$ ) is the boundary air layer assumed to be equal to the first soil layer ( $k = z0$ ) thickness and the pesticide diffusion coefficient  $D_a$  ( $\text{m}^2/\text{d}$ ), which is dependent on the daily air temperature  $T$  (in  $K$ ) as:

$$D_{a(t)} = \left(\frac{T}{T_r}\right)^{1.75} D_{a,r}$$

and where  $D_{a,r}$  ( $\text{m}^2/\text{d}$ ), is the diffusion coefficient in air at reference temperature  $T_r$ . Two options are available to compute the relative diffusion coefficient  $D_{diff,g}$  ( $\text{m}^2/\text{d}$ ), which is adjusted to the available pore air space  $\theta_{z0,air(t)} = \theta_{z0,sat} - \theta_{z0(t)}$ . The first option is given by [Millington and Quirk, 1960, Jin and Jury, 1996]:

$$D_{diff,g(t)} = \frac{D_{a(t)} \cdot (\theta_{z0,air(t)})^2}{(\theta_{z0,sat})^{2/3}}$$

The second option is given by [Currie, 1960, Bakker et al., 1987]:

$$D_{diff,g(t)} = 2.5 \cdot D_{a(t)} \cdot (\theta_{z0,air(t)})^3$$

During each process, sorption extent is computed (as shown above) to obtain the concentration in aqueous  $C_{z0,aq}$  ( $\mu\text{g}/\text{L}$ ) and gas  $C_{z0,g}$  ( $\mu\text{g}/\text{L}$ ) solutions. It is assumed that the diffusion flux from the top soil layer is equal to the diffusion flux across the boundary air layer and that concentration in air  $C_{air} = 0$ , yielding:

$$C_{g,0} = \frac{r_{air(t)}}{r_{air(t)} + r_{soil(t)}} \cdot C_{g,z0} \quad (1)$$

Substituting eq. 1 into ?? ...

The volatilized flux out of the top layer is given by:

... this has been documented in BEACH Methodology!!

## References

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