

## Boring but important disclaimers:

- ▶ If you are not getting this from the GitHub repository or the associated Canvas page (e.g. CourseHero, Chegg etc.), you are probably getting the substandard version of these slides Don't pay money for those, because you can get the most updated version for free at

<https://github.com/julianmak/academic-notes>

The repository principally contains the compiled products rather than the source for size reasons.

- ▶ Associated Python code (as Jupyter notebooks mostly) will be held on the same repository. The source data however might be big, so I am going to be naughty and possibly just refer you to where you might get the data if that is the case (e.g. JRA-55 data). I know I should make properly reproducible binders etc., but I didn't...
- ▶ I do not claim the compiled products and/or code are completely mistake free (e.g. I know I don't write Pythonic code). Use the material however you like, but use it at your own risk.
- ▶ As said on the repository, I have tried to honestly use content that is self made, open source or explicitly open for fair use, and citations should be there. If however you are the copyright holder and you want the material taken down, please flag up the issue accordingly and I will happily try and swap out the relevant material.

# OCES 2003 : Descriptive Physical Oceanography

(a.k.a. physical oceanography by drawing pictures)

## Lecture 13: Southern Ocean and ACC

# Outline

- ▶ Southern Ocean and the Antarctic Circumpolar Current (ACC)
  - largest/strongest current in the world
- ▶ beyond the homogeneous gyre example
  - wind forcing and Ekman overturning cell (+ existence of counter overturning cell)
  - thermal wind shear relation (cf. Lec. 7 + 8)
  - stratification + form stress
  - baroclinic instability (see Lec. 17)
  - influence on MOC (see next Lec.)

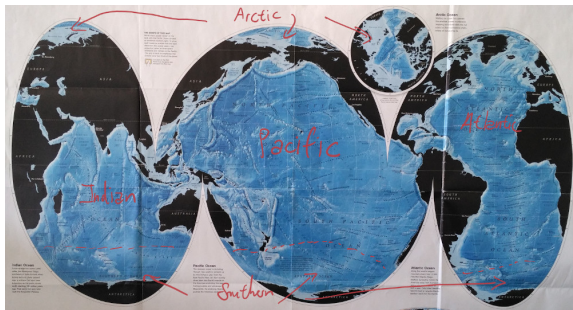
**Key terms:** ACC, Ekman + eddy overturning cell, thermal wind, (interfacial/topographic) form stress, baroclinicity

## Recap: Southern ocean (slide from Lec 2)

Oceans separated horizontally by continental land masses

- constraints on dynamics + circulation (contrast this to atmosphere)

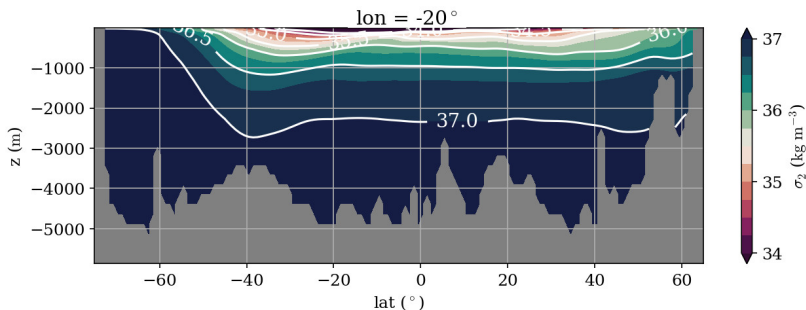
→ Southern ocean slightly different...







## Recap: stratification



**Figure:** Meridional section in the Atlantic of  $\sigma_2$ . See `plot_eos.ipynb`. Strongly tilting isopycnals present in the Southern Ocean, in contrast to relatively flat isopycnals in the basins.

- ▶ meridional section plot of  $\sigma_2$  (see Lec. 6)
- ▶ strongly tilting isopycnals (see Lec. 6) in Southern Ocean
- ▶ wind coming out of page
- ▶ → but sign + profile of wind stress curl? (see Lec. 9)

# Forcings around the Southern Ocean

buoyancy/thermodynamic forcing:

- ▶  $T_{\text{air}}$  **cold** (high latitudes), heat **loss** from ocean
- ▶  $\Rightarrow$  buoyancy **loss** (water getting denser)



# Forcings around the Southern Ocean

buoyancy/thermodynamic forcing:

- ▶  $T_{\text{air}}$  **cold** (high latitudes), heat **loss** from ocean
- ▶  $\Rightarrow$  buoyancy **loss** (water getting denser)
- ▶ if ice **melting**, release freshwater  
 $\Rightarrow$  buoyancy **gain** (water getting lighter)  
 $\rightarrow$  vice-versa if forming ice (**brine rejection**)

# Forcings around the Southern Ocean

buoyancy/thermodynamic forcing:

- ▶  $T_{\text{air}}$  **cold** (high latitudes), heat **loss** from ocean
- ▶  $\Rightarrow$  buoyancy **loss** (water getting denser)
- ▶ if ice **melting**, release freshwater  
 $\Rightarrow$  buoyancy **gain** (water getting lighter)  
 $\rightarrow$  vice-versa if forming ice (**brine rejection**)

mechanical forcing:

- ▶ mid-latitude prevailing Eastward winds (Westerlies)  
 $\Rightarrow$  **E-ward** momentum injection

# Forcings around the Southern Ocean

buoyancy/thermodynamic forcing:

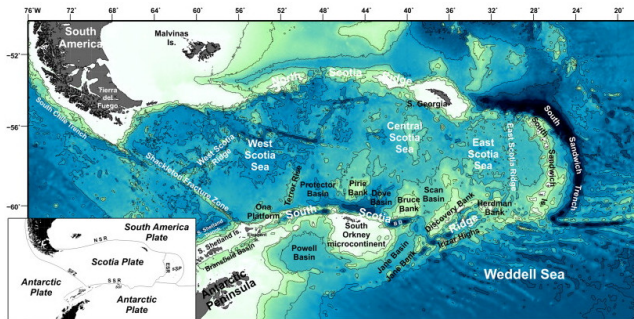
- ▶  $T_{\text{air}}$  **cold** (high latitudes), heat **loss** from ocean
- ▶  $\Rightarrow$  buoyancy **loss** (water getting denser)
- ▶ if ice **melting**, release freshwater  
 $\Rightarrow$  buoyancy **gain** (water getting lighter)  
 $\rightarrow$  vice-versa if forming ice (**brine rejection**)

mechanical forcing:

- ▶ mid-latitude prevailing Eastward winds (Westerlies)  
 $\Rightarrow$  **E-ward** momentum injection
- ▶ bathymetric features  
 $\Rightarrow$  **take out** momentum (see Lec 10) via **topographic form stress**

(see later)

# Forcings around the Southern Ocean



**Figure:** Bathymetry around the Drake passage. Figure modified from Civile *et al.* (2012), *Tectonophysics* (top half of their Fig. 1)

Some notable bathymetric features:

- ▶ **Drake passage**, a choke point for the ACC
- ▶ **Kerguelen plateau**, a wide ridge (not shown here)  
→ water depth can vary from 4000 to 1000 m (recall **PV conservation** Lec. 12)

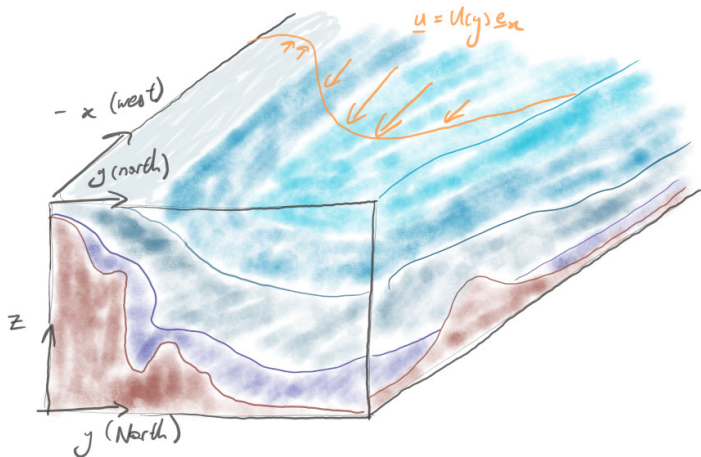
# gyres vs. ACC

gyres	ACC
bounded latitudes	unblocked latitudes
WBC: intense narrow current $\approx 30$ Sv transport depth-independent theory ok?	cf. <b>atmosphere</b> ACC: reasonably “fast”, but broad $\approx 130$ Sv transport depth-independent theory “fails” e.g. Gill (1968) <i>J. Fluid Mech.</i> but see Marshall <i>et al.</i> (2016) <i>Ocean Modell.</i>
Sverdrup balance OK	<b>eddies</b> important

- ▶ despite differences, **dynamical** concepts shared between the two

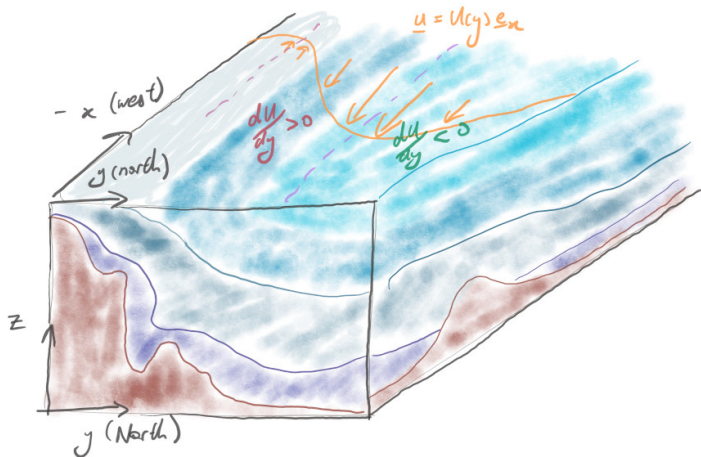
again, **dynamics** important!

# Ekman driven circulation



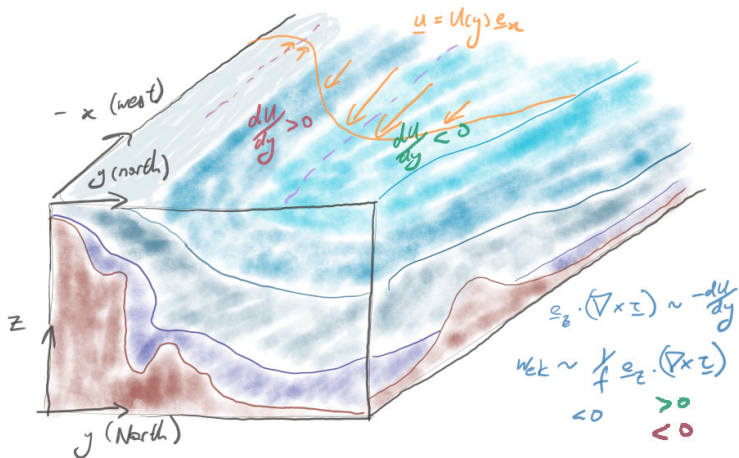
**Figure:** Schematic of wind forcing over Southern Ocean and associated Ekman circulation (recall Lec. 9).  $f < 0$  because we are in the Southern Hemisphere. Diagram based on Olbers [ref here](#)

# Ekman driven circulation



**Figure:** Schematic of wind forcing over Southern Ocean and associated Ekman circulation (recall Lec. 9).  $f < 0$  because we are in the Southern Hemisphere. Diagram based on Olbers [ref here](#)

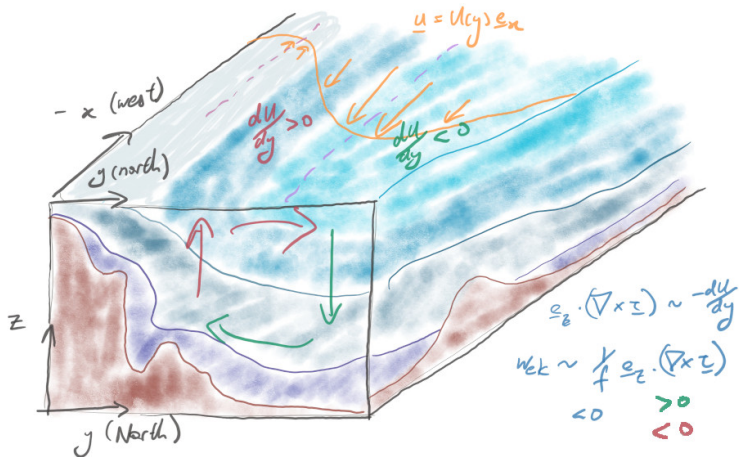
# Ekman driven circulation



**Figure:** Schematic of wind forcing over Southern Ocean and associated Ekman circulation (recall Lec. 9).  $f < 0$  because we are in the Southern Hemisphere. Diagram based on Olbers [ref here](#)



# Ekman driven circulation



**Figure:** Schematic of wind forcing over Southern Ocean and associated Ekman circulation (recall Lec. 9).  $f < 0$  because we are in the Southern Hemisphere. Diagram based on Olbers [ref here](#)

# Ekman driven circulation

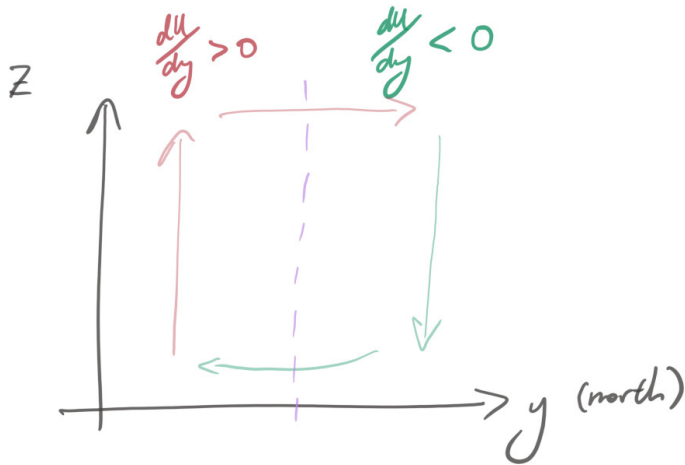


Figure: Ekman overturning and its consequences.

# Ekman driven circulation

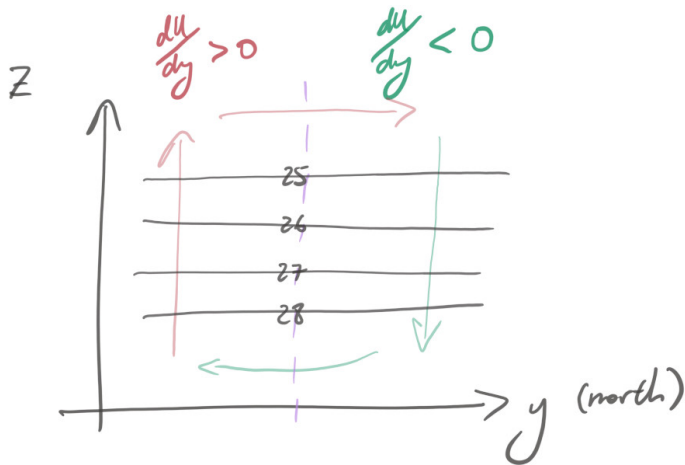


Figure: Ekman overturning and its consequences.

# Ekman driven circulation

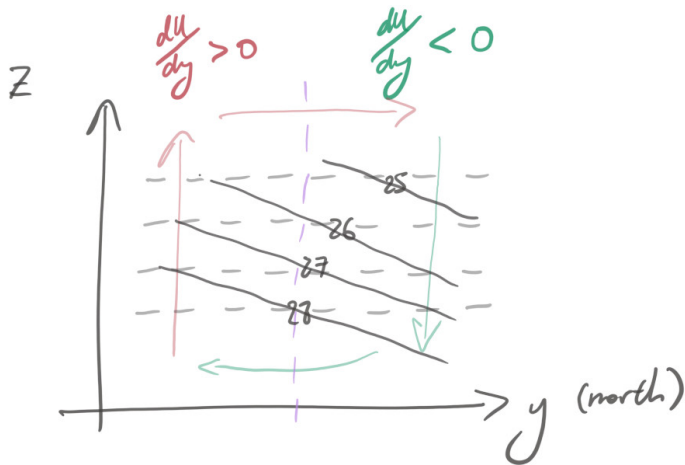


Figure: Ekman overturning and its consequences.

# Ekman driven circulation

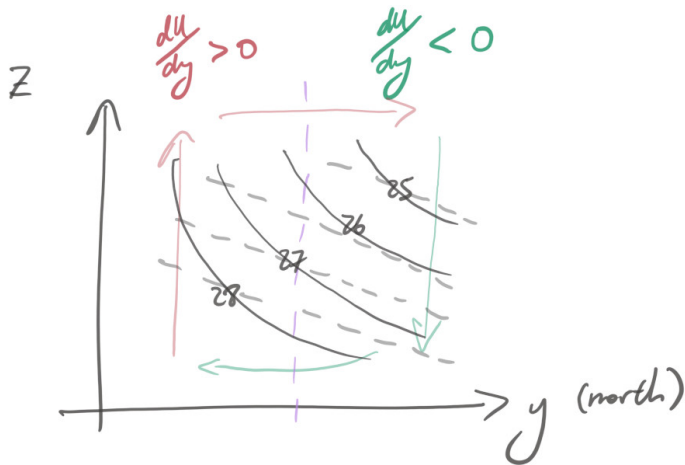


Figure: Ekman overturning and its consequences.

# Ekman driven circulation

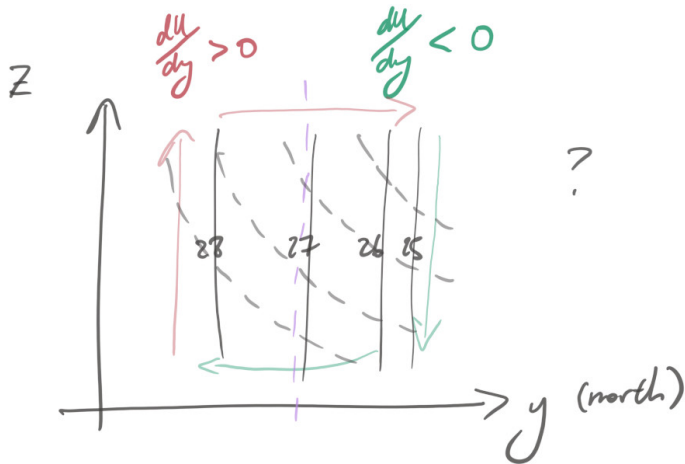


Figure: Ekman overturning and its consequences.

# Ekman driven circulation

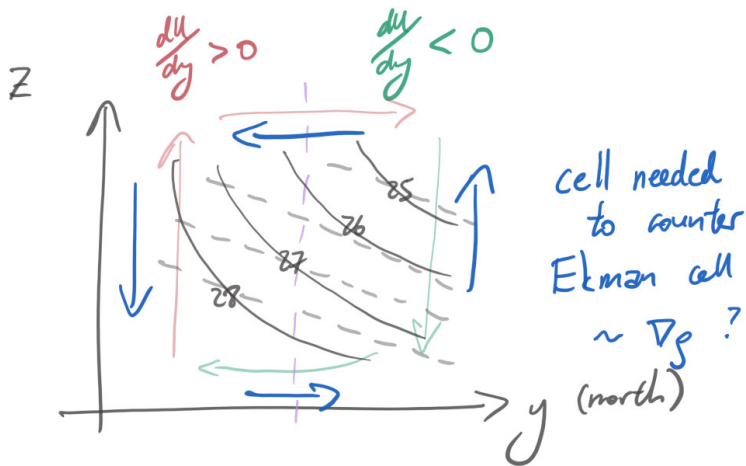


Figure: Ekman overturning and its consequences.

# Thermal wind relation

SO stratification control, how is it related to ACC?



# Thermal wind relation

SO stratification control, how is it related to ACC? Recall **hydrostatic balance** (Lec 7) and **geostrophic balance** (Lec 8):

$$\frac{\partial p}{\partial z} = -\rho g, \quad f \mathbf{e}_z \times \mathbf{u}_g = -\frac{1}{\rho_0} \nabla p$$

# Thermal wind relation

SO stratification control, how is it related to ACC? Recall **hydrostatic balance** (Lec 7) and **geostrophic balance** (Lec 8):

$$\frac{\partial p}{\partial z} = -\rho g, \quad f \mathbf{e}_z \times \mathbf{u}_g = -\frac{1}{\rho_0} \nabla p$$

Combine the two?

# Thermal wind relation

SO stratification control, how is it related to ACC? Recall **hydrostatic balance** (Lec 7) and **geostrophic balance** (Lec 8):

$$\frac{\partial p}{\partial z} = -\rho g, \quad f \mathbf{e}_z \times \mathbf{u}_g = -\frac{1}{\rho_0} \nabla p$$

Combine the two?

- eliminate  $p$ , so take  $\nabla$  of hydrostatic balance and  $\partial/\partial z$  of geostrophic balance (shenanigans here!):

$$\frac{\partial \nabla p}{\partial z} = -g \nabla \rho, \quad f \mathbf{e}_z \times \frac{\partial \mathbf{u}_g}{\partial z} \stackrel{!}{=} -\frac{1}{\rho_0} \frac{\partial}{\partial z} \nabla p$$

# Thermal wind relation

SO stratification control, how is it related to ACC? Recall **hydrostatic balance** (Lec 7) and **geostrophic balance** (Lec 8):

$$\frac{\partial p}{\partial z} = -\rho g, \quad f \mathbf{e}_z \times \mathbf{u}_g = -\frac{1}{\rho_0} \nabla p$$

Combine the two?

- ▶ eliminate  $p$ , so take  $\nabla$  of hydrostatic balance and  $\partial/\partial z$  of geostrophic balance (shenanigans here!):

$$\frac{\partial \nabla p}{\partial z} = -g \nabla \rho, \quad f \mathbf{e}_z \times \frac{\partial \mathbf{u}_g}{\partial z} \stackrel{!}{=} -\frac{1}{\rho_0} \frac{\partial}{\partial z} \nabla p$$

- ▶ eliminate pressure to get

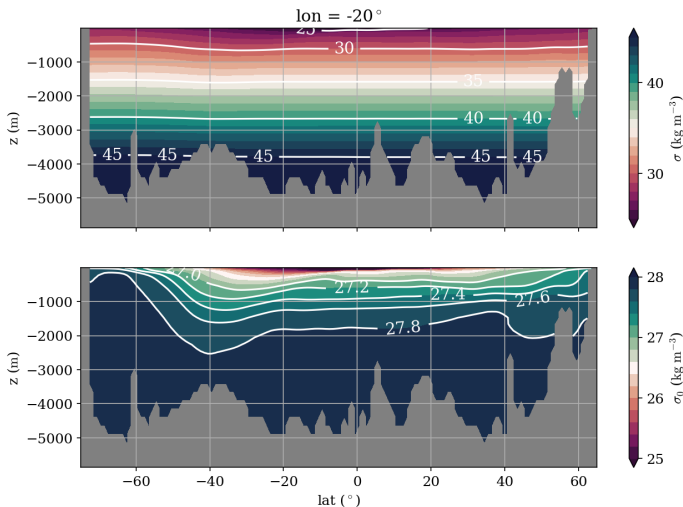
$$f \mathbf{e}_z \times \frac{\partial \mathbf{u}_g}{\partial z} \stackrel{!}{=} -\frac{g}{\rho_0} \nabla \rho$$

# Thermal wind relation

$$f \mathbf{e}_z \times \frac{\partial \mathbf{u}_g}{\partial z} \stackrel{!}{=} -\frac{g}{\rho} \nabla \rho$$

- ▶ **thermal wind relation** (ignoring some shenanigans for now) says that  
**horizontal gradients in  $\rho \sim$  vertical gradients in  $u_g$**
- ▶ tilting isopycnals implies there a geostrophic flow
  - system needs to be rotating
  - more tilt = stronger flow
  - direction depends on tilt and hemisphere (because  $f$  changes sign)
- ? combine with SSH to get **vertical profile** of geostrophic flow? (see Lec. 20, OCES 3203, maybe OCES 3301)

# Thermal wind relation



**Figure:** Meridional section in the Atlantic of (top) in-situ density and (bot)  $\sigma_0$ . See `plot_eos.ipynb`.  $\sigma_0$  implies there is a thermal wind coming out of the page in the Southern Ocean (because  $f < 0$ ), consistent with what we know. On the other hand, in-situ density  $\sigma$  implies basically nothing is going on, which we know is not true.

## 2/3 way point

- ▶ wind puts momentum in + induces an **overturning** via **Ekman suction/pumping** (Lec. 9)
  - **steepens** isopycnals (Lec. 5 + 6)
  - cannot continue indefinitely (otherwise **convectively unstable**), existence of **counter overturning cell** (see later)
- ▶ **thermal wind shear relation**
  - hydrostatic + geostrophic balance (Lec. 7 + 8)
  - tilting isopycnals = geostrophic flow
  - consistent with E-ward wind momentum input

## 2/3 way point

- ▶ wind puts momentum in + induces an **overturning** via **Ekman suction/pumping** (Lec. 9)
  - **steepens** isopycnals (Lec. 5 + 6)
  - cannot continue indefinitely (otherwise **convectively unstable**), existence of **counter overturning cell** (see later)
- ▶ **thermal wind shear relation**
  - hydrostatic + geostrophic balance (Lec. 7 + 8)
  - tilting isopycnals = geostrophic flow
  - consistent with E-ward wind momentum input
- ▶ **baroclinic** theory here, **vertical structure** + **stratification** involved
  - cf. homogeneous gyre theory, but ideas here also apply somewhat there



# Form stress

## How is momentum removed?

- ▶ removal at bottom (momentum input by wind  $\sim$  momentum removal at ocean floor <sub>loss to land</sub>)
  - but how? (**topographic form stress**)
  - role of **PV conservation**? (Lec. 12)
- ▶ how is it **transferred** vertically? (**interfacial form stress**)

# Form stress

## How is momentum removed?

- ▶ removal at bottom (momentum input by wind  $\sim$  momentum removal at ocean floor <sub>loss to land</sub>)
  - but how? (**topographic form stress**)
  - role of **PV conservation**? (Lec. 12)
- ▶ how is it **transferred** vertically? (**interfacial form stress**)

in essence **pressure gradients** (Lec. 7)

# Form stress

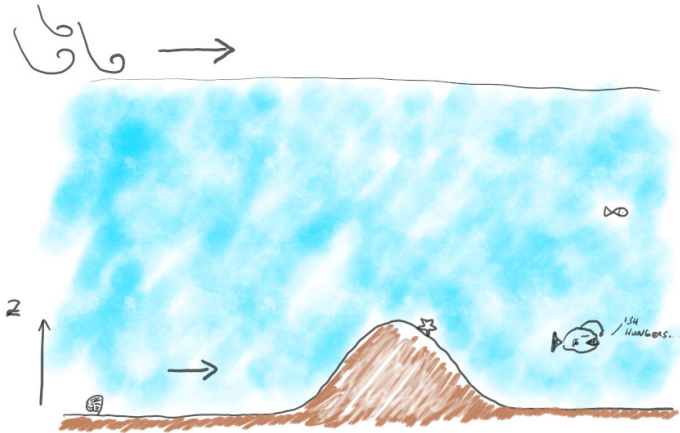


Figure: Schematic of (interfacial + topographic) form stress.

# Form stress

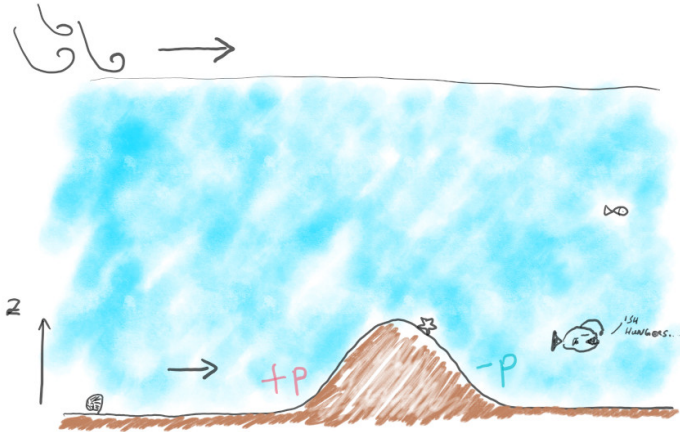


Figure: Schematic of (interfacial + topographic) form stress.

# Form stress

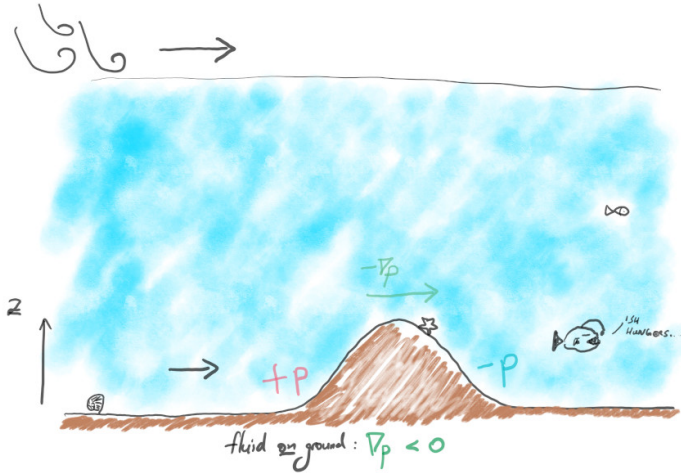


Figure: Schematic of (interfacial + topographic) form stress.

# Form stress

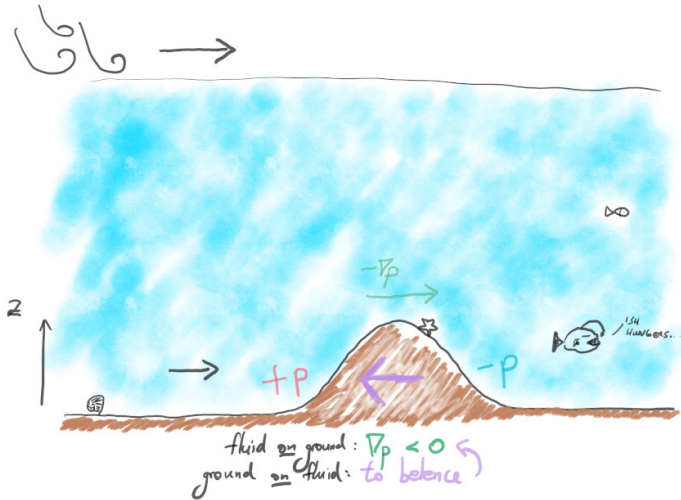


Figure: Schematic of (interfacial + topographic) form stress.

# Form stress

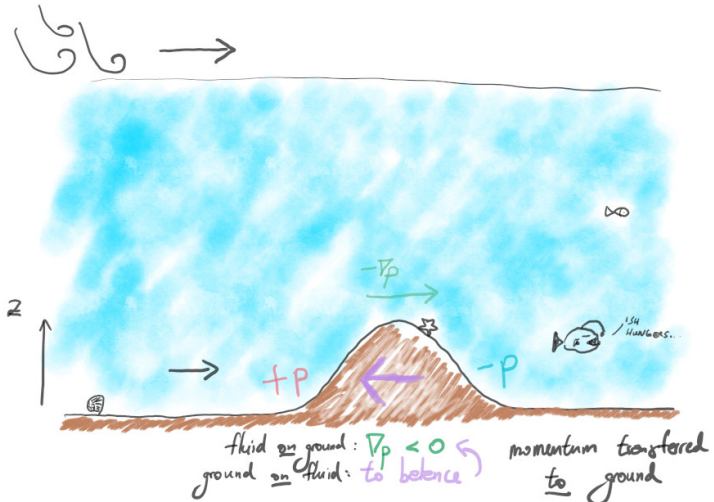


Figure: Schematic of (interfacial + topographic) form stress.

# Form stress

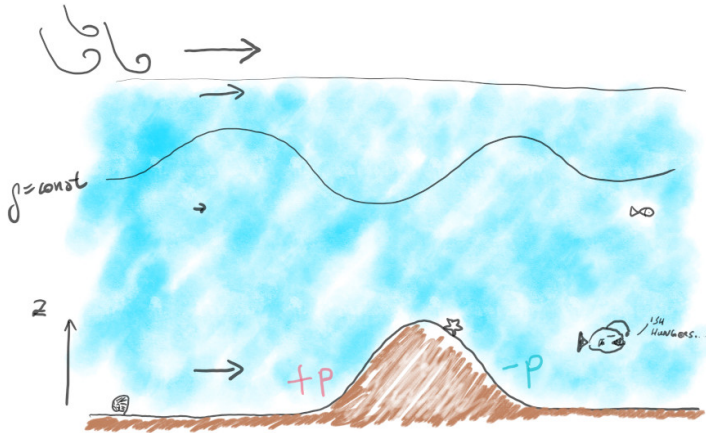


Figure: Schematic of (interfacial + topographic) form stress.



# Form stress

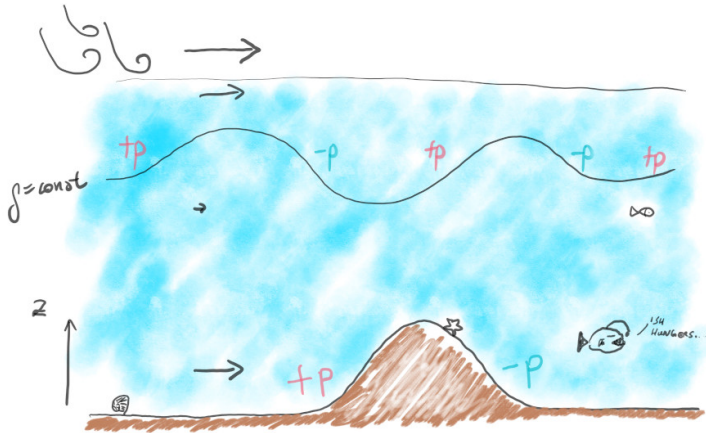


Figure: Schematic of (interfacial + topographic) form stress.

# Form stress

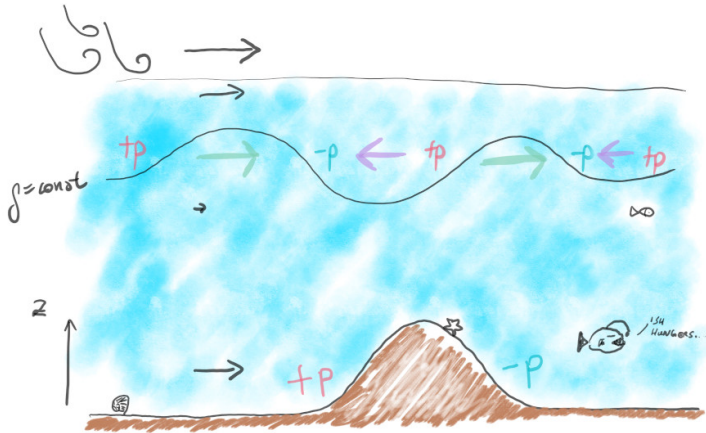


Figure: Schematic of (interfacial + topographic) form stress.

# Form stress

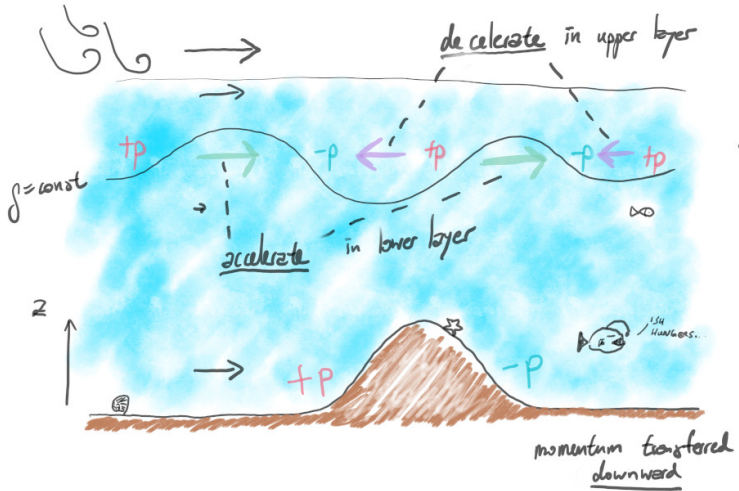


Figure: Schematic of (interfacial + topographic) form stress.



## $f/H$ contours

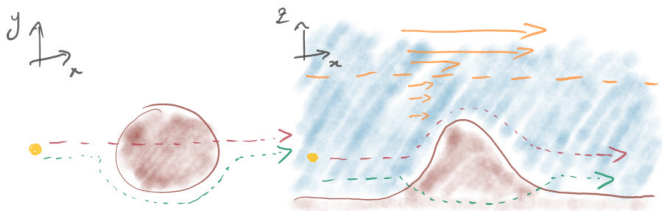


Figure: Some consequences of  $f/H$  contours.

- ▶ if bathymetric feature not that tall ( $H \approx H_{\text{ref}}$ ), just go **over**
  - not that much “pressing” onto bathymetry
  - weaker  $\nabla p$  so smaller topographic form stress
  - **weak** deceleration, weak topographic influence
- ▶ theories neglecting baroclinicity + topography results in ACC transport being ridiculously large

## $f/H$ contours

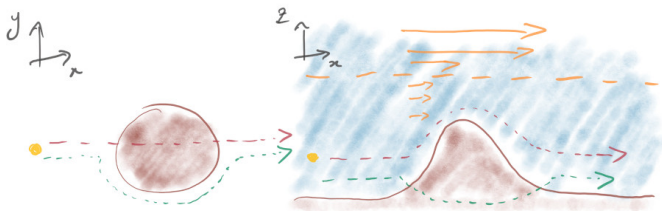
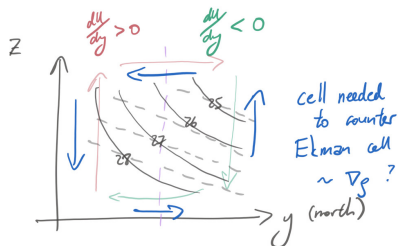


Figure: Some consequences of  $f/H$  contours.

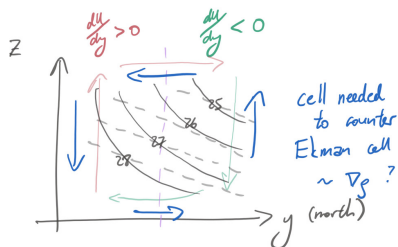
- ▶ if bathymetric feature tall enough ( $H < H_{\text{ref}}$ ),  $f/H$  contours can be **blocked**
  - significant “pressing” onto bathymetry
  - strong  $\nabla p$  so stronger topographic form stress
  - **strong** deceleration, topographic influence significant
  - e.g. Drake passage, Kerguelen plateau

## Quick brief on baroclinic instability (more in Lec. 17)



- ▶ what is the source of the counter overturning cell?
  - strength probably needs to be related to **isopycnal slopes**

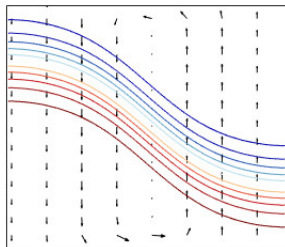
# Quick brief on baroclinic instability (more in Lec. 17)



- ▶ what is the source of the counter overturning cell?  
→ strength probably needs to be related to **isopycnal slopes**

## baroclinic instability

- ▶ sloping isopycnals  $\sim$  vertically sheared flow (thermal wind)
- ▶ sheared flow  $\Rightarrow$  instability (see Lec. 17)  
→ **reduce** vertical shear  $\sim$  **flatten** isopycnals

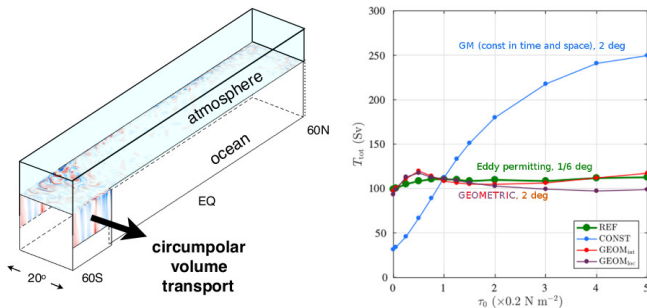




# Quick brief on baroclinic instability (more in Lec. 17)

eye candy

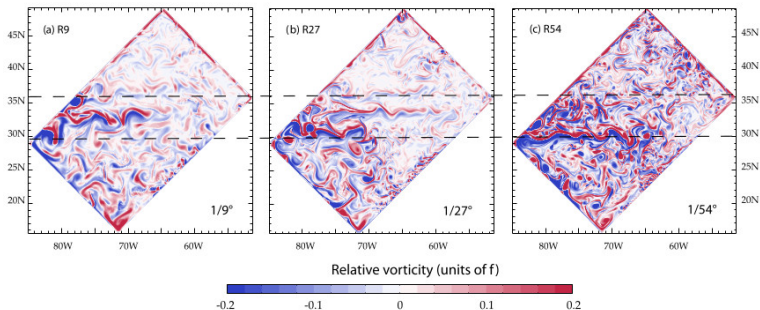
# Quick brief on baroclinic instability (more in Lec. 17)



**Figure:** Idealised sector model from Munday, Johnson & Marshall (2013), *J. Phys. Oceanogr.* and results on ACC transport (related to Southern Ocean overturning) sensitivity with changes of wind depending on **mesoscale parameterisation**, from Mak *et al.* (2018), *J. Phys. Oceanogr.*.

- **residual** of Ekman and eddy cell affects SO stratification
  - affects ACC transport through thermal wind
  - can have **global** effect via isopycnal connectivity to all ocean basins (see Lec. 14)

# Quick brief on baroclinic instability (more in Lec. 17)



**Figure:** Snapshots of surface relative vorticity of a double gyre model at different resolutions. From Lévy *et al.*, (2010), *Ocean Model*. (modified from their Fig. 3).

- important in gyres too (see Lec. 12)
  - shaping the WBC, **bio-physical interaction**, momentum transfer etc.

# Summary

- ▶ Southern Ocean the “center” of the global ocean
  - unblocked latitudes
  - stratification here can influence **global** stratification (and in turn global **MOC**, ocean heat content etc.) (see Lec. 14)
- ▶ ACC largest current in the world
  - tilting isopycnals, **thermal wind relation** (geostrophic + hydrostatic balance) (more in Lec. 20)
  - **Ekman** vs. **eddy** overturning cell, **residual** affecting overall stratification
  - role of **form stress** and  **$f/H$  contours**
  - **baroclinic eddies** a source of form stress + flattens isopycnals (more in Lec. 17)

**Key role of **dynamics**, applicable to ACC + gyres!**