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<https://github.com/julianmak/academic-notes>

The repository principally contains the compiled products rather than the source for size reasons.

- ▶ Associated Python code (as Jupyter notebooks mostly) will be held on the same repository. The source data however might be big, so I am going to be naughty and possibly just refer you to where you might get the data if that is the case (e.g. JRA-55 data). I know I should make properly reproducible binders etc., but I didn't...
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OCES 2003 : Descriptive Physical Oceanography

(a.k.a. physical oceanography by drawing pictures)

Lecture 12: Gyres 2 (Western intensification)

Thur 11th Mar

Outline

- ▶ (wind driven?) gyres
- ▶ depth-independent model with no topography: vorticity balance
 - boundary acting as sink of vorticity
 - **Western intensification** and WBCs
- ▶ complications (extras, or things of interest)
 - topography
 - depth-dependence
 - buoyancy forcing

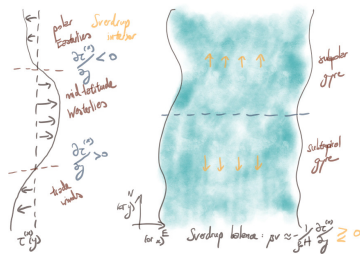
Key terms: Western Boundary Currents (WBC), vorticity balance, western intensification

Recap: Sverdrup balance

► Sverdrup balance (in meridional)

$$\beta v \approx F_{\tau}(x, y) = -\frac{1}{\rho_0 H} \frac{\partial \tau^{(x)}}{\partial y}$$

→ Coriolis balancing **wind stress curl**



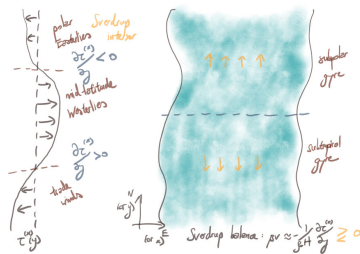
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What happens at boundaries?

- **friction** important over boundary layer, Sverdrup balance breaks down
- **boundary condition** fixes (more or less) direction of flow
→ intensification has to be on the **West**

Single (subtropical) gyre: vorticity balance

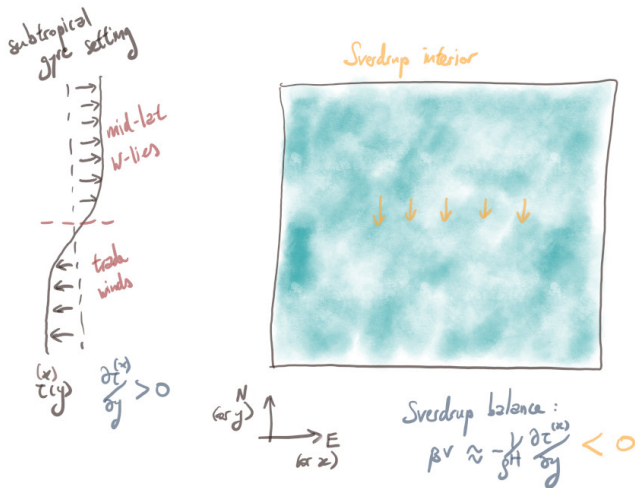


Figure: Schematic for NH subtropical gyre.

Single (subtropical) gyre: vorticity balance

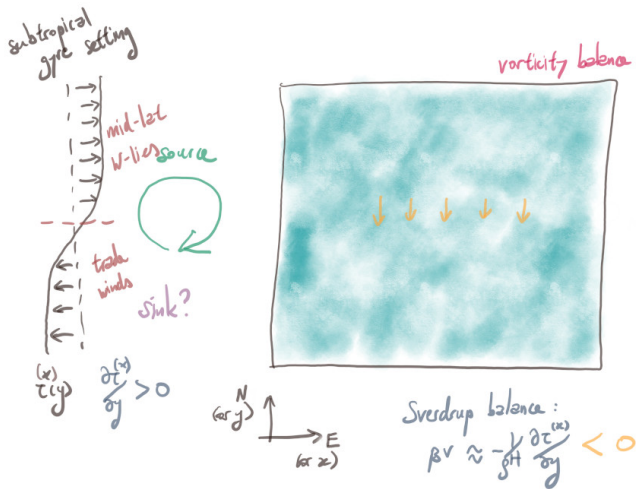


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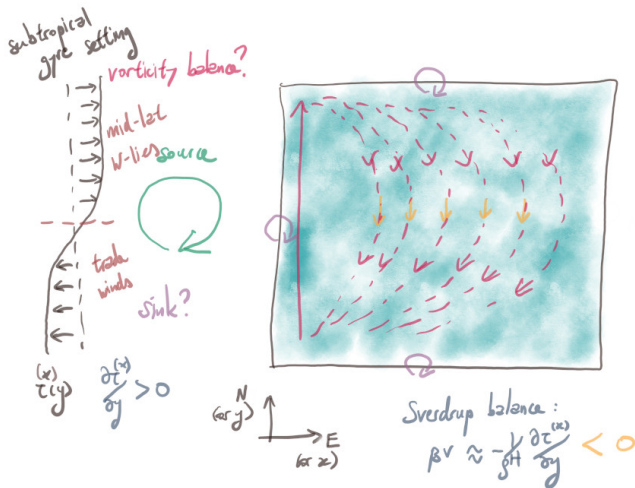


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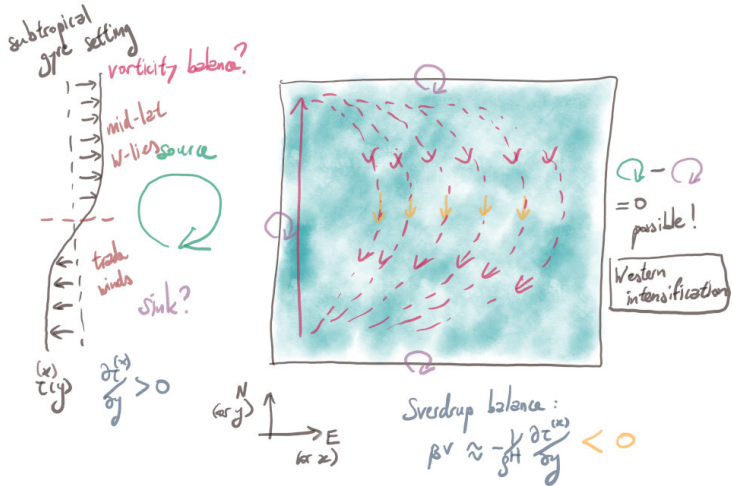


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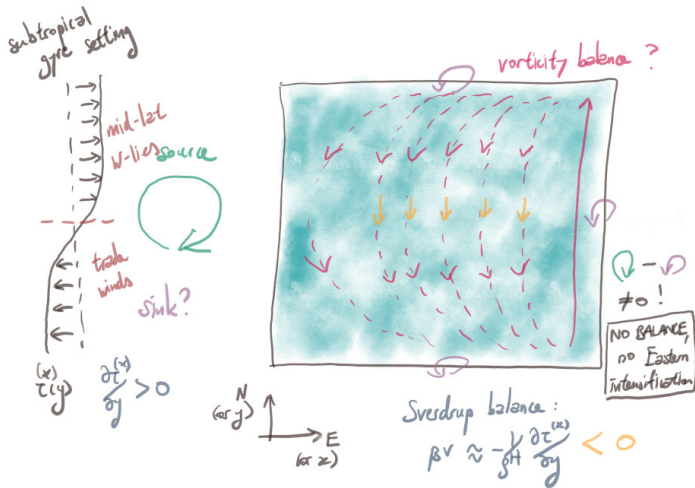


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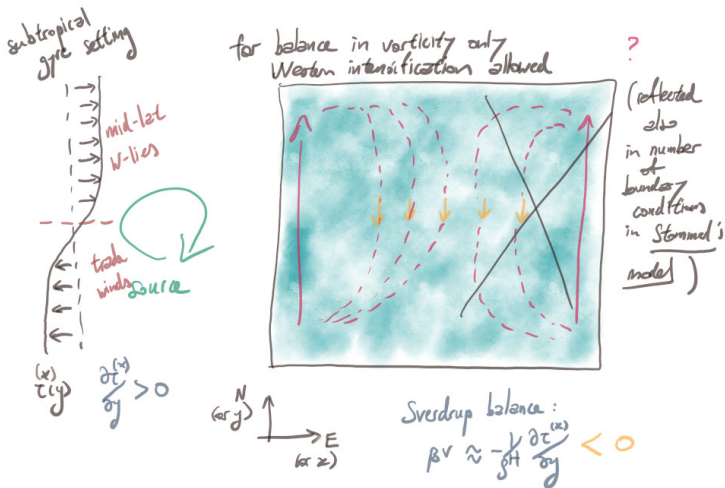


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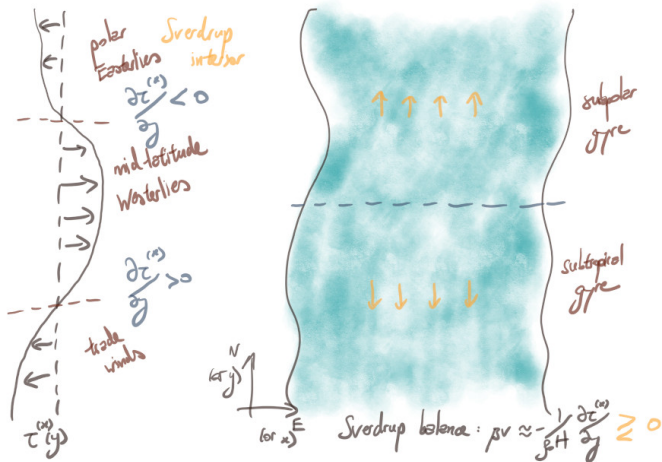


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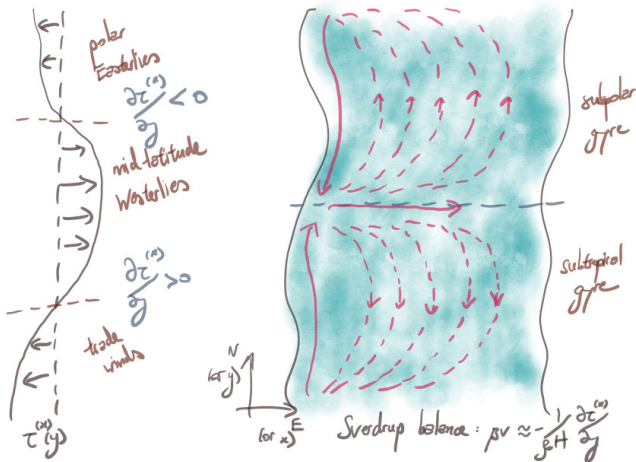


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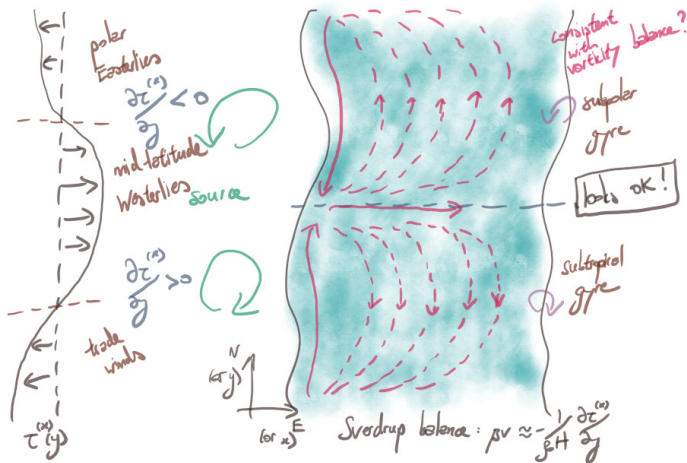


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Summary of arguments

In this depth-independent (cf. **barotropic**) model of the gyre:

- ▶ **Sverdrup balance** (wind stress curl \sim Coriolis) tells you what interior flow you have
 - links also with **con/divergence** (up/downwelling)
- ▶ boundary conditions fixes **flow orientation** and has to be **Western intensification**
 - energetically unfavourable (wind blowing against the detached current)
 - only the Western Boundary can be a sink of the input vorticity

Aside:

- ▶ Stommel's original model only allows for one boundary condition to put on
 - have to put on the West, asymptotic matching for solution (see maybe OCES 3203)
 - see Ch. 14 of Vallis (2006), *Atmospheric and Ocean Fluid Dynamics* for more

Some notes

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→ model here made **assumptions** (e.g. depth-independence) to give a **deep** theory

Some notes

Useful lesson: if your argument misses on some aspects, either

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Value of the model/theory to be measured by:

- ▶ does it work?
- ▶ did you learn something from it?
- ▶ does it tell you what you might want to consider next?

Extras

Other discrepancies:

- ▶ no asymmetry between subtropical and subpolar gyres
→ subtropical gyre more prominent in observations
- ▶ WBCs and resulting near-surface transport goes **polewards**
→ return EQ-ward flow in the east?

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Model assumptions:

1. no bathymetry
2. depth-independence (cf. **barotropic** but I try not to use that term)
3. wind-driven

Extras: topography

The ocean has bathymetric features

- ▶ go around or go over? depends on **energetics!**

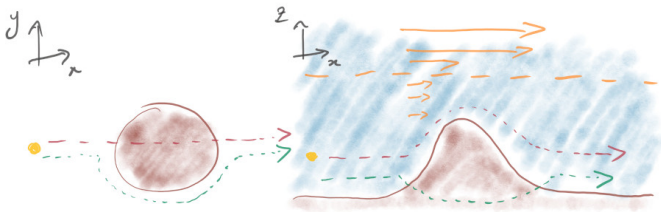


Figure: Possible paths when facing an obstacle: go over or go around. Depending on H as well one might be more preferable than the other. Note also that the strong flow above the obstacle is intuitively not strongly influenced by the obstacle (despite measures such as **JEBAR** stating otherwise; see e.g. Cane, Kamenkovich & Krupitsky, 1998, *J. Phys. Oceanogr.*).

- ▶ influence not throughout depth because **baroclinicity**
→ despite things like **JEBAR** stating otherwise...

Extras: topography

- ▶ fluid wants to try and conserve **angular momentum**
 - contrast this with **linear** momentum p
 - related to **vorticity**, but as $q = (\omega + f)/H$ (**potential vorticity**, PV)

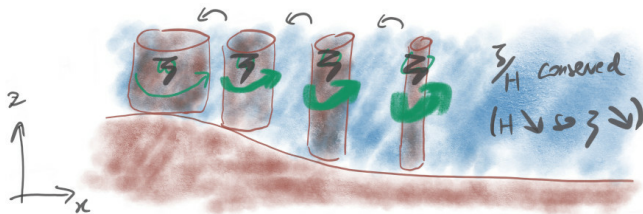


Figure: Conservation of $q = \omega/H$ (assuming $|\omega| \gg |f|$ for illustration). As H decreases, the spinning gets slower to compensate so that q is conserved.

- ▶ cf. ballerina effect (or the spinning-chair-with-leg-sticking out effect)

Extras: topography

over slopes we now have more things coming into play

(e.g. Salmon, 1998; Hughes & de Cuevas, 2001, *J. Phys. Oceanogr.*)

- ▶ **bottom pressure torque** (related to PV conservation)
- ▶ **bottom drag** (instead of friction as before)

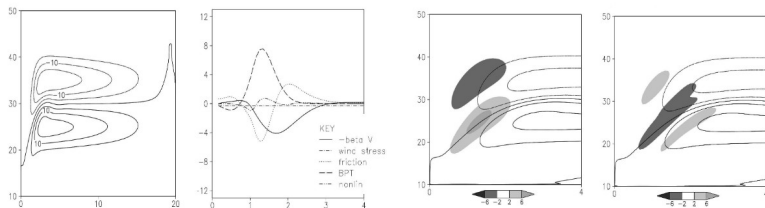


Figure: Results from a depth-independent gyre model with a slope on the west, showing (a) streamfunction, (b) balances between terms, (c) bottom pressure torque forcing, and (d) frictional forcing. Adapted from Jackson, Hughes & Williams, 2006, *J. Phys. Oceanogr.* (their Figs. 1–4). Also see Williams & Follows (2011), Ch. 8.3.

Extras: vertical variations + baroclinicity

- ▶ there are of course vertical gradients in general ocean flow
→ related to stratification through **thermal wind shear relation** (see next Lec.)
→ allows for **baroclinic instability** (more in Lec 13 + 17)

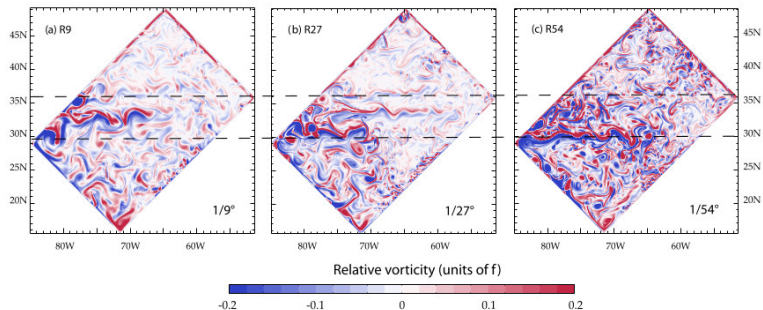


Figure: Snapshots of surface relative vorticity of a double gyre model at different resolutions. From Lévy *et al.*, (2010), *Ocean Model*. (modified from their Fig. 3).

Extras: vertical variations + baroclinicity

- deep WBCs but going the other way (see again in Lec. 14)
 - similar type of argument above (e.g. Stommel & Arons, 1960, *Deep-Sea Res.*)
 - modifications by **topography** (steering), **mixing** (re-circulations) etc.

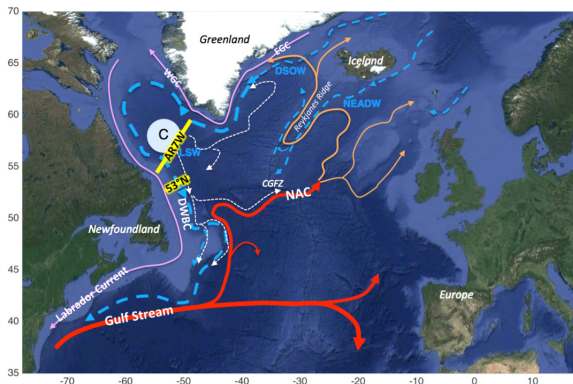


Figure: Schematic of WBC (Gulf stream, the red line) and the deep WBC (the blue-dashed line). From Handmann *et al.*, 2018, *J. Geophys. Res: Oceans* (their Fig. 1).

Extras: vertical variations + baroclinicity

► deep EBCs? (Yang, Tziperman & Speer, 2020, *Geophys. Res. Lett.*)

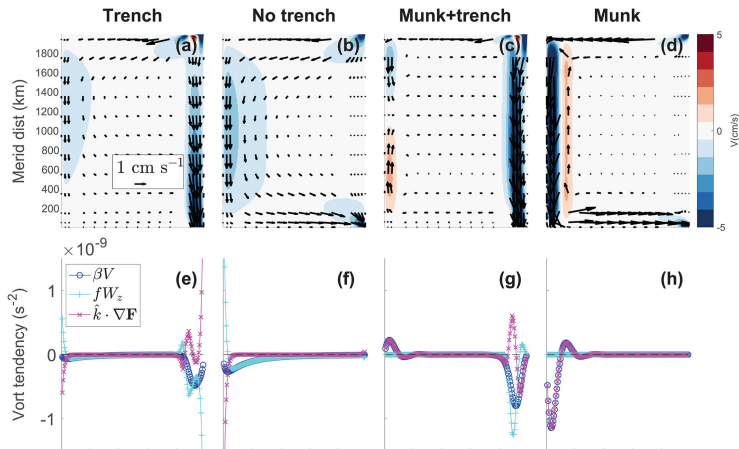


Figure: Meridional flow and vorticity budget and bathymetry (rows) from Yang, Tziperman & Speer, 2020, *Geophys. Res. Lett.* (modified from their Fig. 3).

Extras: buoyancy forcing

- ▶ classical theory of wind-driven gyres, but what about **thermodynamic/buoyancy** forcing?

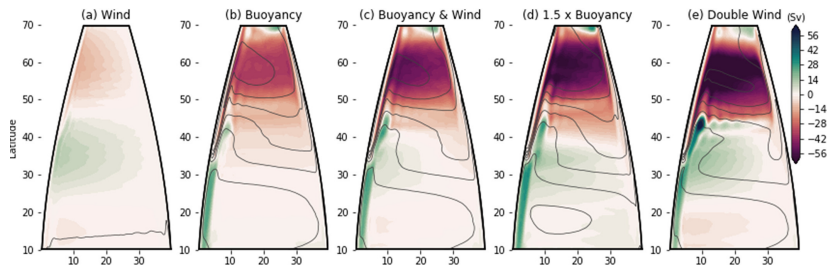


Figure: Barotropic streamfunction (rows) for a few experiments, showing the barotropic streamfunction (as shading) and the SST (as gray lines). From Hogg & Gayen (2020), *Geophys. Res. Lett.* (modified from their Fig. 3).

- ▶ **subpolar** gyre particularly affected by buoyancy forcing
- ▶ **eddies** seem to be needed (cf. Colin de Verdière, 1988, *J. Mar. Res.*)

Summary

Simple (but by no means **complete**) theory for **gyre circulation**

- ▶ depth-independent, wind-driven with no topography
- ▶ **Sverdrup interior** fixes interior flow
- ▶ **vorticity balance** fixes orientation and gives **Western intensification**

→ vorticity input by wind can **only** (!) be removed at the West (in this model...)

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→ vorticity input by wind can **only** (!) be removed at the West (in this model...)

- ▶ start from some assumptions, just roll with it (carefully and logically) and see where it takes you...
→ just pictures and considering whether things are pos/negative above (wasn't too bad right?)

Theories are not static and do evolve!

Summary

If things in the **simple** theory doesn't work, either

- ▶ the logic was wrong (!)
- ▶ the theory is too simple and needs more details
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“All models are wrong, but some are useful”

– attributed to George Box