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https://github.com/julianmak/academic-notes

The repository principally contains the compiled products rather than the source for size reasons.

- Associated Python code (as Jupyter notebooks mostly) will be held on the same repository. The source data however might be big, so I am going to be naughty and possibly just refer you to where you might get the data if that is the case (e.g. JRA-55 data). I know I should make properly reproducible binders etc., but I didn't...
- ▶ I do not claim the compiled products and/or code are completely mistake free (e.g. I know I don't write Pythonic code). Use the material however you like, but use it at your own risk.
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# OCES 2003 : Descriptive Physical Oceanography

(a.k.a. physical oceanography by drawing pictures)

Lecture 5: Sea water properties (temperature and salinity)

#### Outline

- thermodynamic vs. mechanical forcing of ocean
  - $\rightarrow$  two lectures on former, four lectures on latter
- ultimately care about density  $\rho$  (sort of, see  $\rho_{\theta}$  and  $\gamma$  later...)
  - $\rightarrow$  temperature *T* and salinity *S* contributions
  - → horizontal and vertical structures
  - → sea surface temperature/salinity (SST, SSS)
  - $\rightarrow$  links to forcing

**Key terms**: density, buoyancy, temperature, salinity, SST, SSS, thermo/halocline, watermass property

### Recap: forces

- sea water moves around since there are forces acting on it
  - $\rightarrow$  e.g. wind  $\Rightarrow$  "pushing" water
  - $\rightarrow$  e.g. uneven heating  $\Rightarrow$  movement to redistribute heat

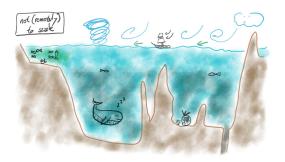


Figure: Schematic of ocean forcing.

to understand/predict circulation  $\sim$  how forces act



# Recap: equations of motion

Denoting u = (u, v) and  $u_3 = (u, v, w)$ , to <u>numerous</u> approximations (!!!) (see OCES 3203) ocean dynamics is governed by

$$\rho_0 \left( \frac{\partial u}{\partial t} + u \cdot \nabla u + 2\Omega \times u \right) = -\nabla p + F_u + D_u \tag{1}$$

$$\frac{\partial p}{\partial z} = -\rho g \tag{2}$$

$$\nabla \cdot \boldsymbol{u}_3 = 0 \tag{3}$$

$$\left(\frac{\partial \mathbf{T}}{\partial t} + \mathbf{u}_3 \cdot \nabla \mathbf{T}\right) = \mathbf{F}_T + \mathbf{D}_T \tag{4}$$

$$\left(\frac{\partial S}{\partial t} + u_3 \cdot \nabla S\right) = F_S + D_S \tag{5}$$

$$\rho = \rho(T, S, p) \tag{6}$$

Respectively, (1) momentum equation, (2) hydrostatic balance, (3) incompressibility, (4) temperature equation, (5) salinity equation, and (6) equation of state (EOS)



### Buoyancy: motivating example

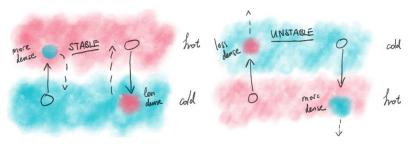
Q. which one is unstable?



Figure: (Un)stable temperature configurations.

### Buoyancy: motivating example

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 $\textbf{Figure:} \ (Un) stable \ temperature \ configurations.$ 

- density (!)  $\rho$  (units: kg m<sup>-3</sup>)
- ▶ buoyancy  $b = -(\delta \rho/\rho_0)g$  (so units of...?)
  - → how "floaty" something is (e.g. warm water, lighter density, **more** buoyant)

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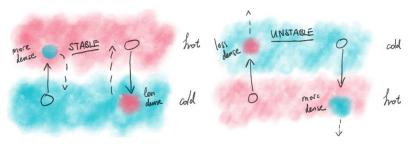


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ultimately it's density/buoyancy we care about



### Sea water properties

Q: What are the distinguishing features of sea water?

Main focus for here: what controls density of seawater?

Another hint: how is it **different** to air?

### Sea water properties: temperature

#### Denote temperature (!) by T

(other temperatures/density next lecture)

- units of °C or K (0°C = 273.15 K)
- ▶ isotherm = lines/surfaces of constant temperature
- ▶ above around  $4^{\circ}$ C (!!!) warmer  $\sim$  less dense, i.e.

$$\rho \nearrow \text{ as } T \searrow$$

- $\rightarrow$  water actually **most dense** around 4°C (EOS next lec.)
- $\rightarrow$  (otherwise consequence for ice?)
- measured by thermometer, sound speed etc.

# Sea water properties: salinity

#### Denote salinity (!) by S

- ▶ given in g kg<sup>-1</sup> or sometimes PSU (practical salinity unit)
   → dimensionless really, use of PSU strongly discouraged
- isohaline = lines/surfaces of constant salinity
- higher salinity = more things dissolved in water, so

$$\rho \nearrow as S \nearrow$$

chemical measure through chlorinity (e.g. titration)

$$S_A \approx 1.80655 \times \text{Chlorinity}$$

- $\rightarrow$  absolute salinity  $S_A$
- usually now done through electrical conductivity (see lec. 19)
  - $\rightarrow$  practical salinity  $S_P$



### Sea water properties: opacity

Sea water is dense and has things dissolved in it

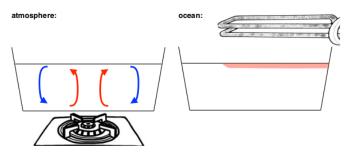
stops light going through (either diffusing or absorbing), opaque
 → light doesn't penetrate very deep, consequence for primary production



Figure: Picture of the sea. CCO Public Domain, taken from phys.org

### Sea water properties: opacity

One very important and interesting difference between atmosphere and ocean:

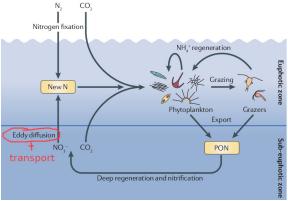


**Figure:** Schematic of principal sources of thermal forcing between atmosphere and ocean. Courtesy of David Marshall (Oxford).

- actually quite hard to drive up/down motion + deep circulation! (cf. Sandströrm's theorem)
  - $\rightarrow$  but there is one, so why? (see Lec 11 18)



### Sea water properties (sort of): other things



- things like nutrients, carbon, oxygen are dissolved in seawater
- negligible (?) impact for the physics, but important for biogeochemistry (some in Lec. 21 + 22) or tracking watermasses (see

### Warning!

There are different types of temperature...

- ightharpoonup in-situ T vs . potential  $\theta$  vs. conservative  $\Theta$
- ▶ subtle but !!!<u>VERY IMPORTANT</u>!!! differences
  - $\rightarrow$  for dynamics usually we don't care about T

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- ...and different types of salinity!
  - ightharpoonup practical  $S_P$  vs. absolute  $S_A$ 
    - $\rightarrow$  absolute salinity (sometimes  $S_A$ ) part of **TEOS-10**

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...and we have different densities  $\rho/\rho_{\theta}/\gamma$  too!

just be aware that subtleties <u>do exist</u> and <u>do matter</u>...

▶ sun + radiation

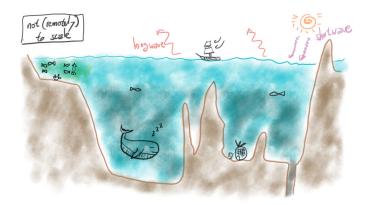


Figure: Schematic of temperature forcing at the ocean surface. Only showing shortwave and longwave radiation (latent heat + sensible heat can be gain or loss; not shown here).

Sea Surface Temperature (SST) mostly(!) by the sun

▶ shortwave radiation  $Q_{sw}$  here (units: W m<sup>-2</sup>) → not shown are longwave radiation (seen as a heat loss), latent + sensible heat (can be gain or loss) see homework? (look up any physical oceanography book under "heat budgets")

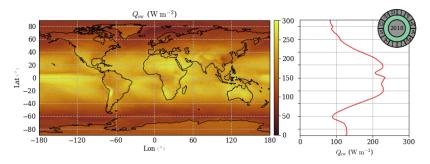


Figure: Year-averaged (left) and year and zonal averaged (right) shortwave radiation, from the JRA55 dataset (Kobayashi et al. 2015, J. Meteor. Soc. Japan). See plot\_jra55\_sample.ipynb

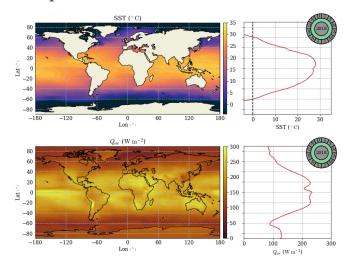


Figure: As previous figure but for SST (top) and shortwave radiation (bot).

ightharpoonup notice correlation between SST and  $Q_{\rm sw}$ 

time-varying data with seasonal cycle (movie here)

▶ river runoff, evaporation, precipitation (see Lec. 5)

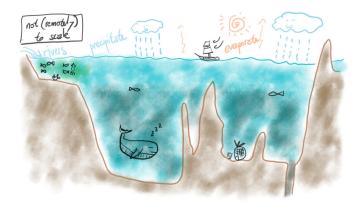


Figure: Schematic of ocean forcing.

#### Sea Surface Salinity (SSS) by

- decrease in salinity by precipitation (units: depends...)
- ▶ **increase** in salinity by evaporation (units: depends...)
- decrease in salinity by runoff (units: depends...)
  - $\rightarrow$  river, ice melt

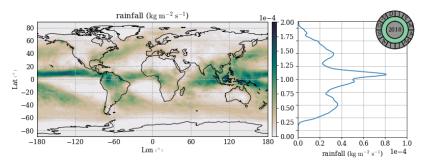


Figure: Year-averaged (left) and year and zonal averaged (right) rainfall (so **decrease** in salinity), from the JRA55 dataset (Kobayashi *et al.* 2015, *J. Meteor. Soc. Japan*). See plot\_jra55\_sample.ipynb

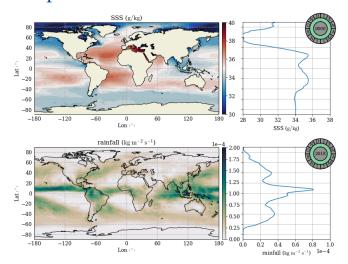


Figure: As previous figure but for SSS (top) and rainfall (bot). Notice how the Atlantic is more salty than the Pacific.

Q. why is SSS high when rainfall is also high?



▶ time-varying data with seasonal cycle (cheating here with SSS!)

#### Observed *T* profiles: sections e.g. (x,z) or (y,z)

ightharpoonup meridional section (i.e. (y, z)) of T

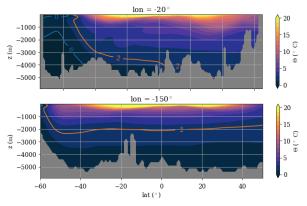


Figure: Meridional section of yearly-averaged conservative temperature in the Atlantic (top) and Pacific (bot), based on World Ocean Atlas 2013 data. Meridional range chosen to roughly correspond to Talley *et al.* (2011) Fig. 4.11 and 4.12. See plot.WOAl3.sample.ipynb

- intrusion of bottom cold waters in Atlantic
- marked temperature contour is "higher" in Pacific



#### Observed *S* profiles: sections e.g. (x, z) or (y, z)

ightharpoonup meridional section (i.e. (y, z)) of S

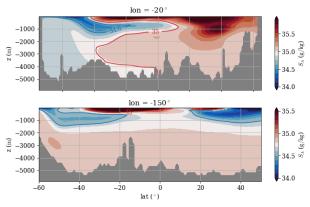


Figure: Meridional section of yearly-averaged absolute salinity in the Atlantic (top) and Pacific (bot), based on World Ocean Atlas 2013 data. Meridional range chosen to roughly correspond to Talley et al. (2011) Fig. 4.11 and 4.12. See plot.WOA13.sample.ipvnb

- distinct salinity signature in salinity in Atlantic
- notice a fresh intermediate layer in Pacific



#### Observed *T* profiles: vertical usually (depth) or (z)

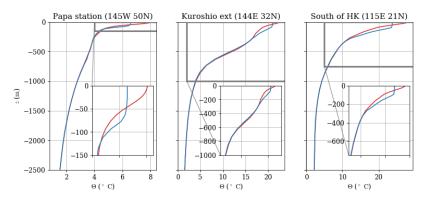


Figure: Vertical variation of Θ at some designated locations, based on WOA13 data. Red and blue line denote summer and winter climatology. See plot\_WOA13\_sample.ipynb

notice a rapidly changing top part ("sharp", think large gradient) and slower varying ("smooth") bottom part



### Observed *T* profiles: vertical usually (depth) or (z)

mixed layer (usually (!) O(100 m))

- very top bit where stratification is weak (see previous panel (c) in winter)
  - $\rightarrow$  strong vertical mixing (see Lec. 17)

### Observed T profiles: vertical usually (depth) or (z)

- mixed layer (usually (!) O(100 m))
  - very top bit where stratification is weak (see previous panel (c) in winter)
    - $\rightarrow$  strong vertical mixing (see Lec. 17)
- thermocline (usually O(200 1000 m))
  - the transition region between the "sharp" and "smooth" part in temperature
    - $\rightarrow$  hence the **thermo** part
    - $\rightarrow$  it's gradients we care about, and thermocline is where temperature gradient is largest **below** the mixed layer
    - $ightarrow 
      ho \sim T$  (or  $\Theta$  which is used here), so it indicates changes in density

#### Observed *S* profiles: vertical usually (depth) or (z)

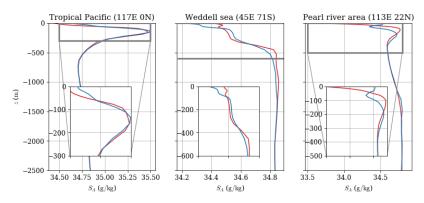


Figure: Vertical variation of  $S_A$  at some designated locations, based on WOA13 data. Red and blue line denote summer and winter climatology. See plot\_WOA13\_sample.ipynb

- ▶ similarly for salt, halocline
  - → large evaporation, ice melt and river runoff here



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$$\rho = \rho(T, S, p)$$

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- !!! complication: different types of T and  $\rho$  relating to concept of work done related to pressure... (next lec)

