

Lithospheric plates : a dynamical perspective

Fanny Garel

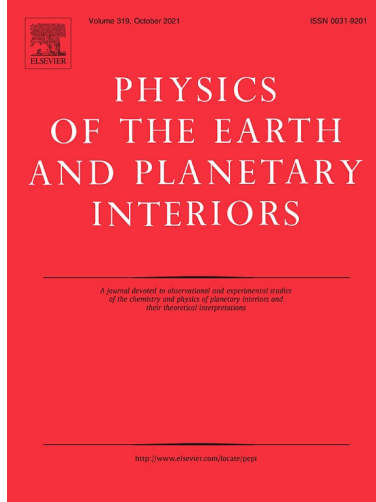


University of Montpellier, France

Deep Earth Doctoral School - Les Houches - Autumn 2022



Dynamical definition of surface plates and subducting slabs



PEPI special issue 20-21

Physical properties and observations of the lithosphere-asthenosphere system

Editors : Rick Aster, Saskia Goes, Derek Schutt

- seismology, heat flow
- electrical conductivity
- gravimetry
- mineral physics
- geotherms
- rheology, melting
- dynamics

...

- **Thermal structures of slabs from gravimetry ?**
- **Breaking plates from below with dislocation creep ?**

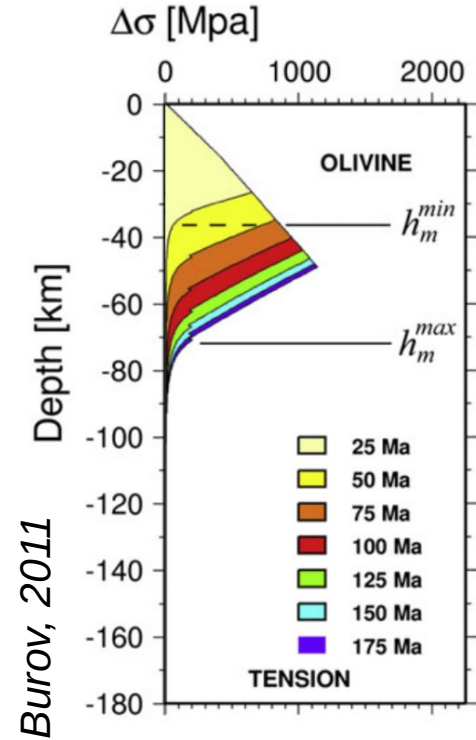


2022-2026

ANR-21-CE49-0009



Concepts and proxies of plates & LAB

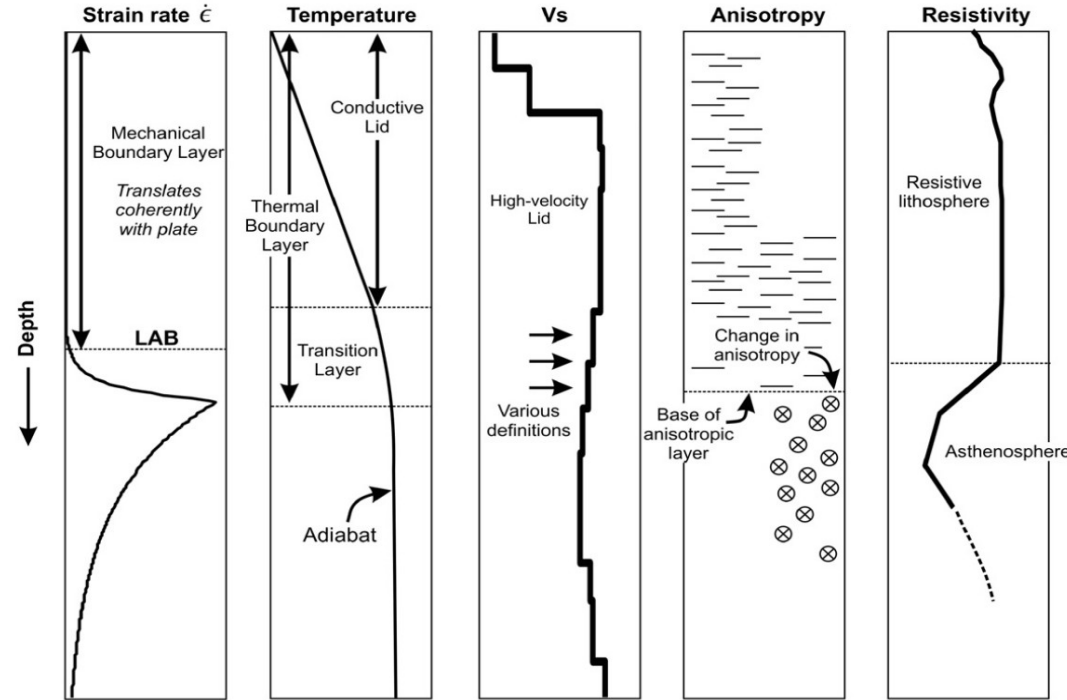


cold, strong,
boundary-layer
lithosphere

vs.

hot, soft, convective
asthenosphere

bathymetry, heat
flow, flexure, shear
wave velocity, radial
anisotropy...



Eaton et al., 2009

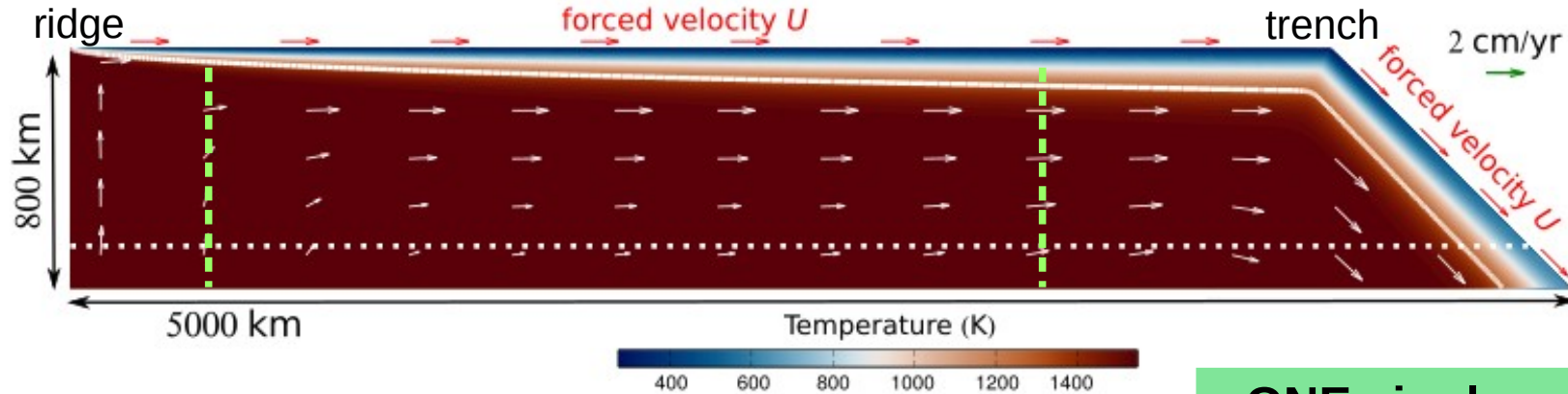
→ velocity profile with depth ? evolution with time ?



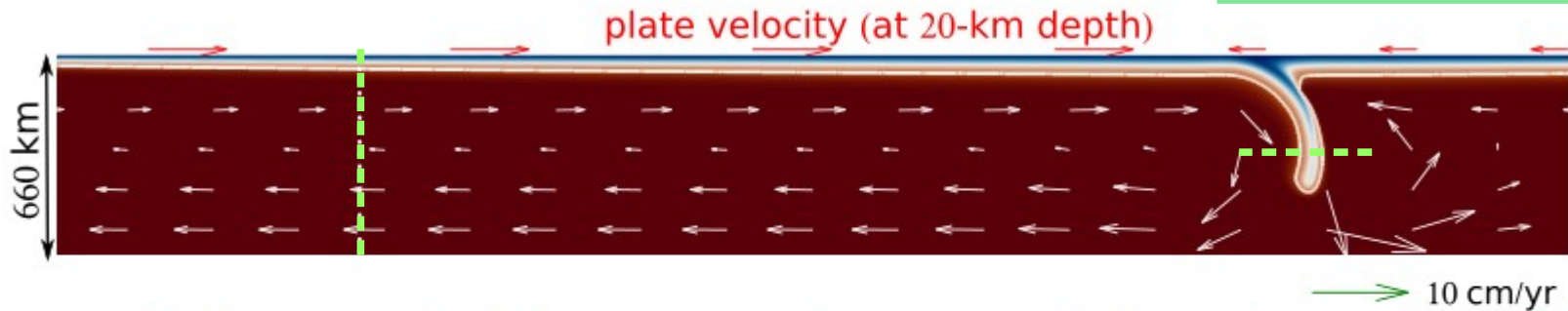
Finite-element thermo-mechanical models



> plate-driven set-up : steady-state



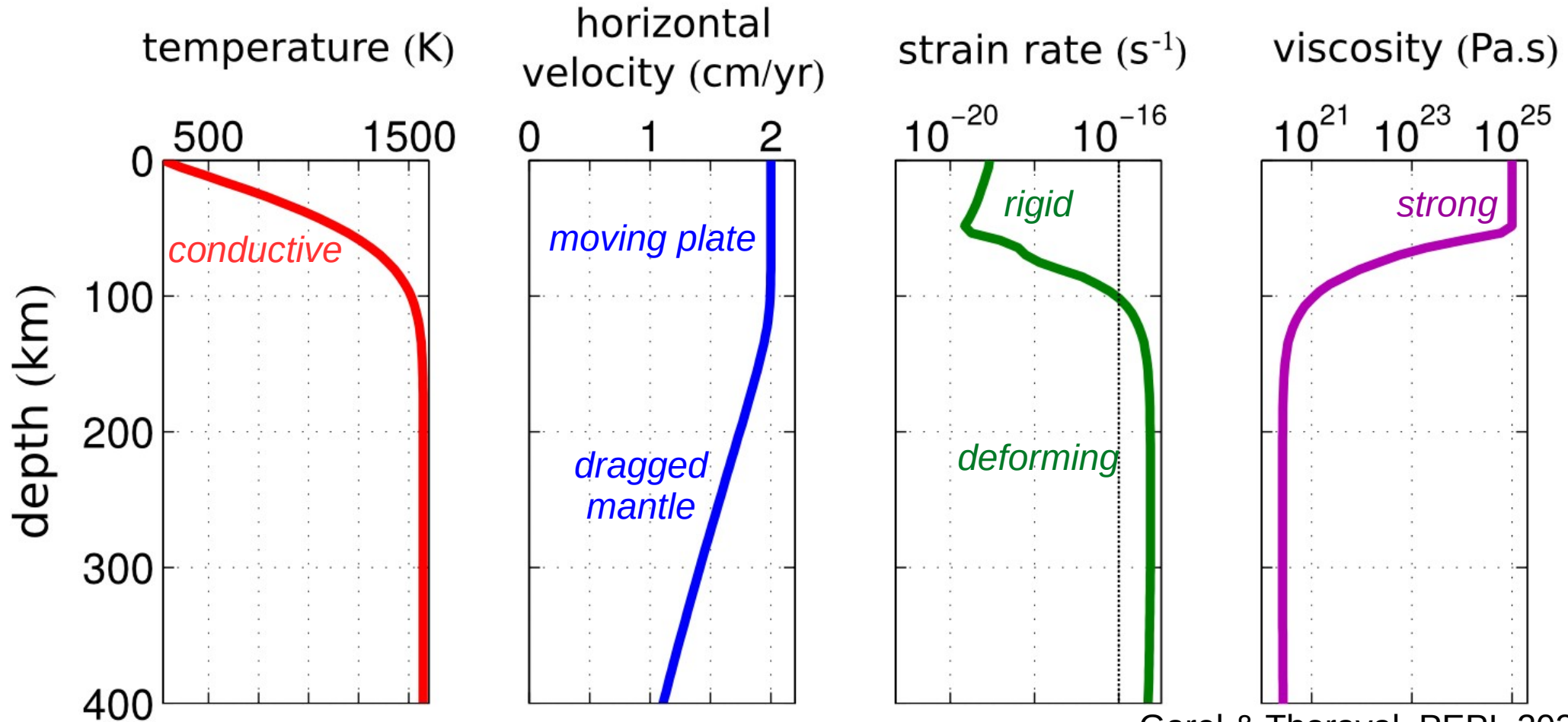
> free subduction set-up : time-evolving



ONE single material for both lithosphere and asthenosphere
(no pre-imposed discontinuity)

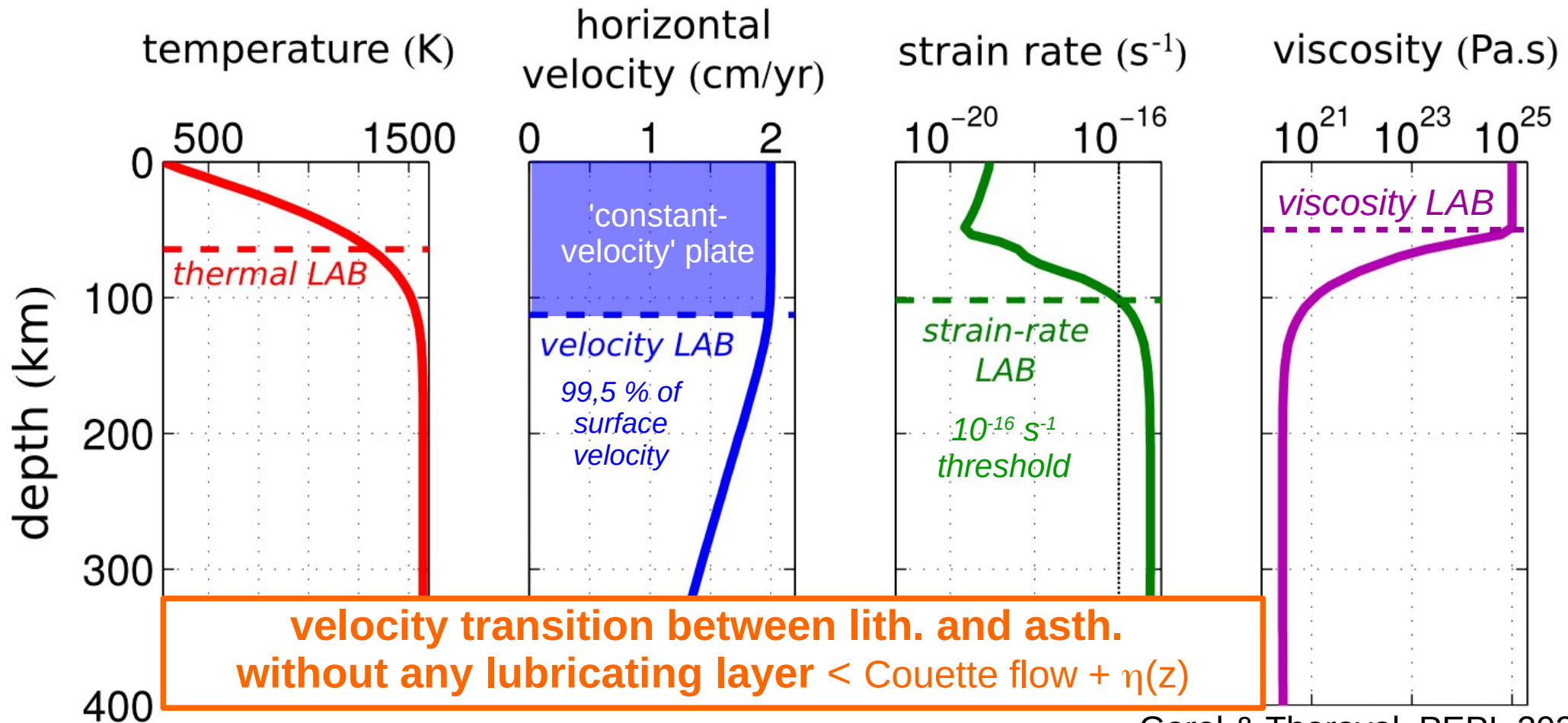
Vertical profiles below a moving plate (2 cm/yr)

PLATE-DRIVEN



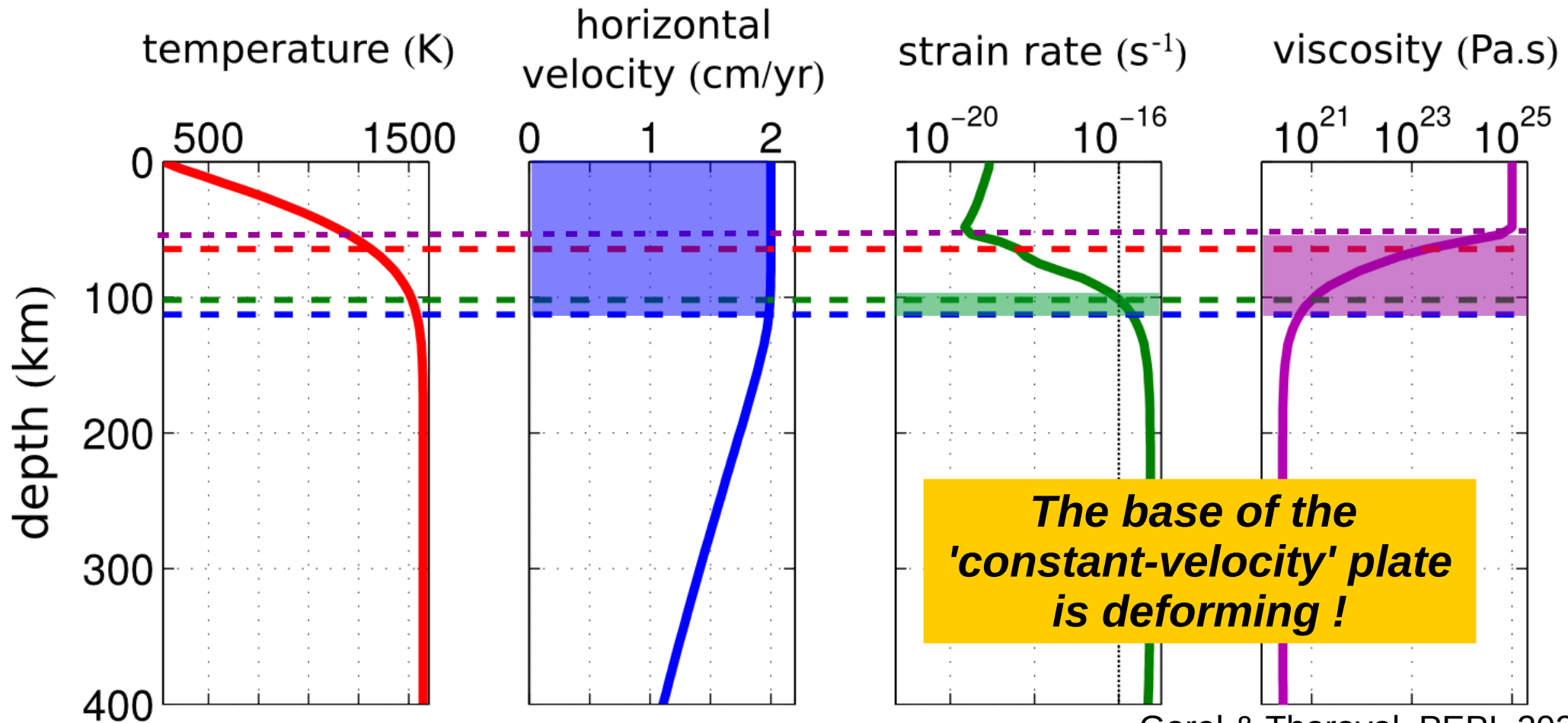
Vertical profiles below a moving plate (2 cm/yr)

PLATE-DRIVEN

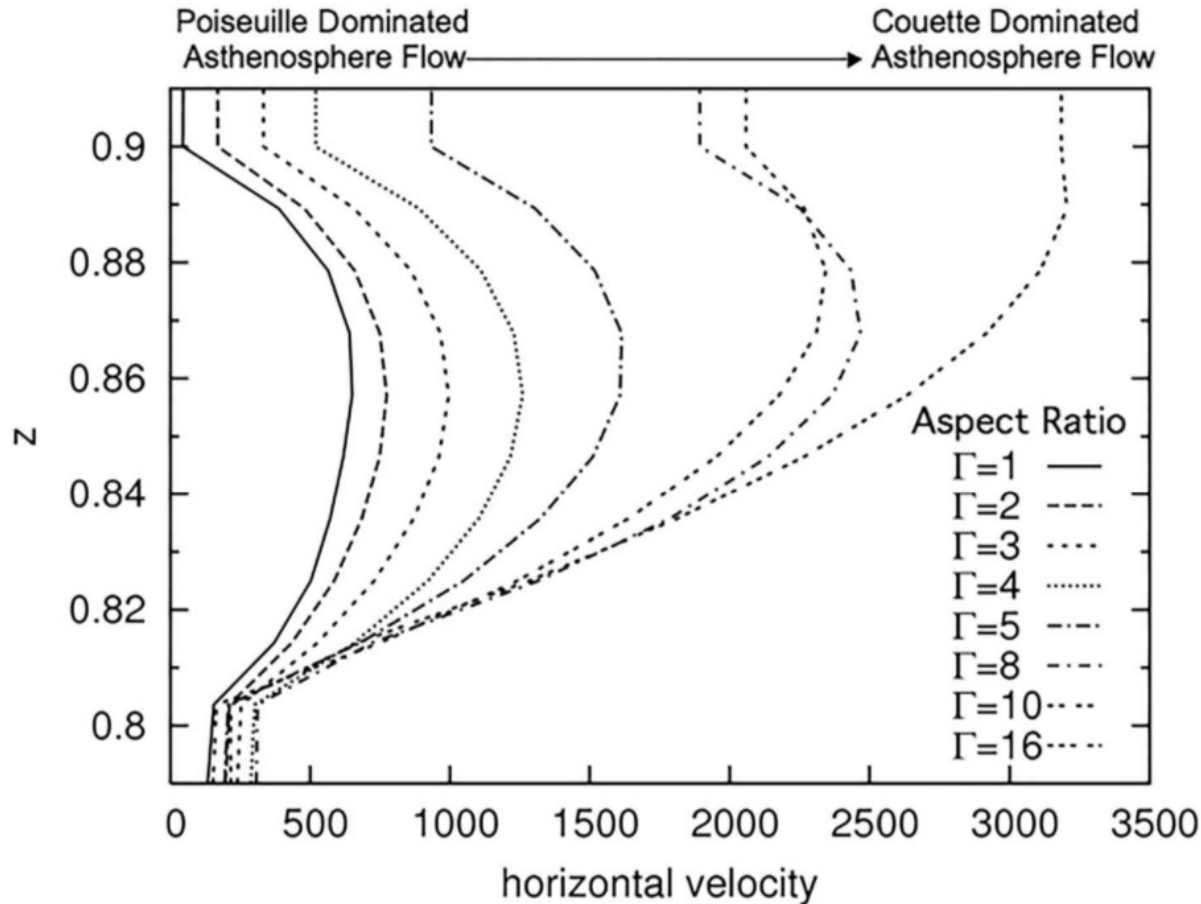


Vertical profiles below a moving plate (2 cm/yr)

PLATE-DRIVEN



Dynamical plate in Couette or Poiseuille flows



**constant-velocity plate
also observed
for 'active'
asthenosphere flow**

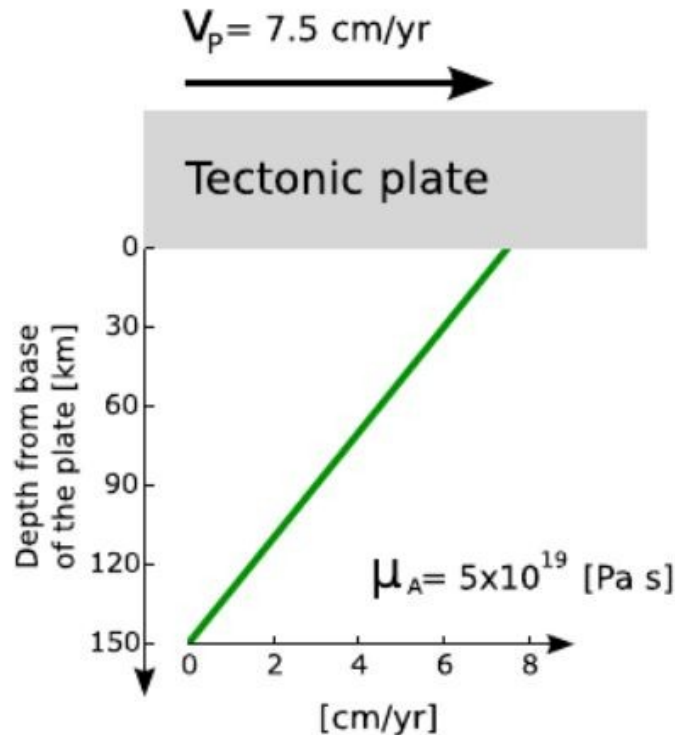
Hoink and Lenardic, 2010
Richards and Lenardic, 2018



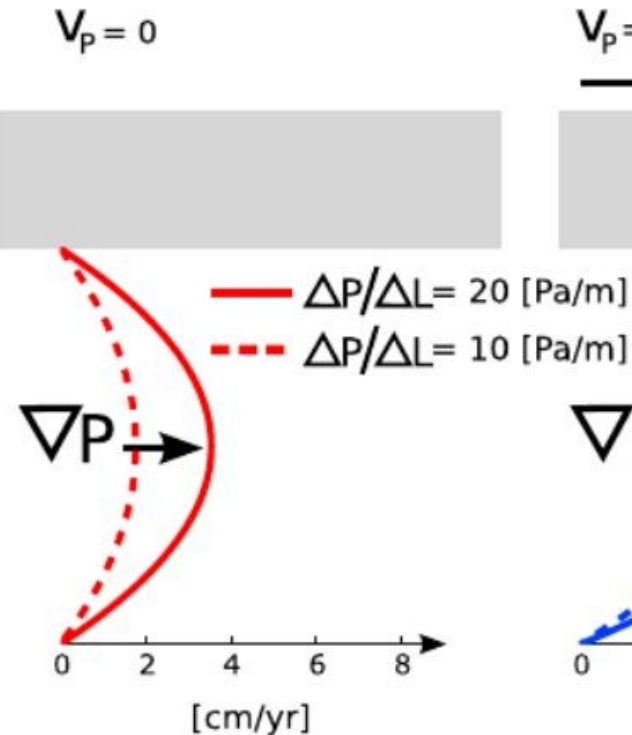
drag of asth.
by Pacific plate
(itself pulled by subduction)

asthenospheric
“push” by Hawaiian
plume

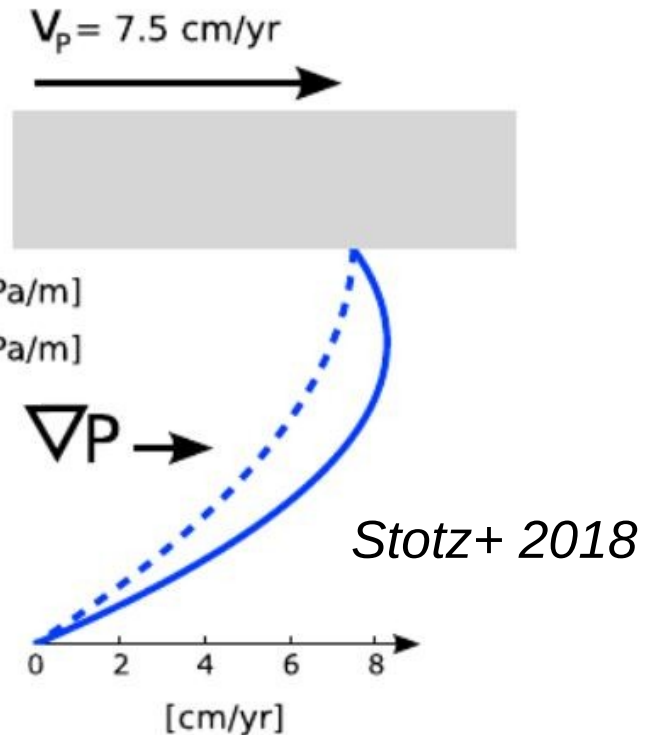
A) Couette flow



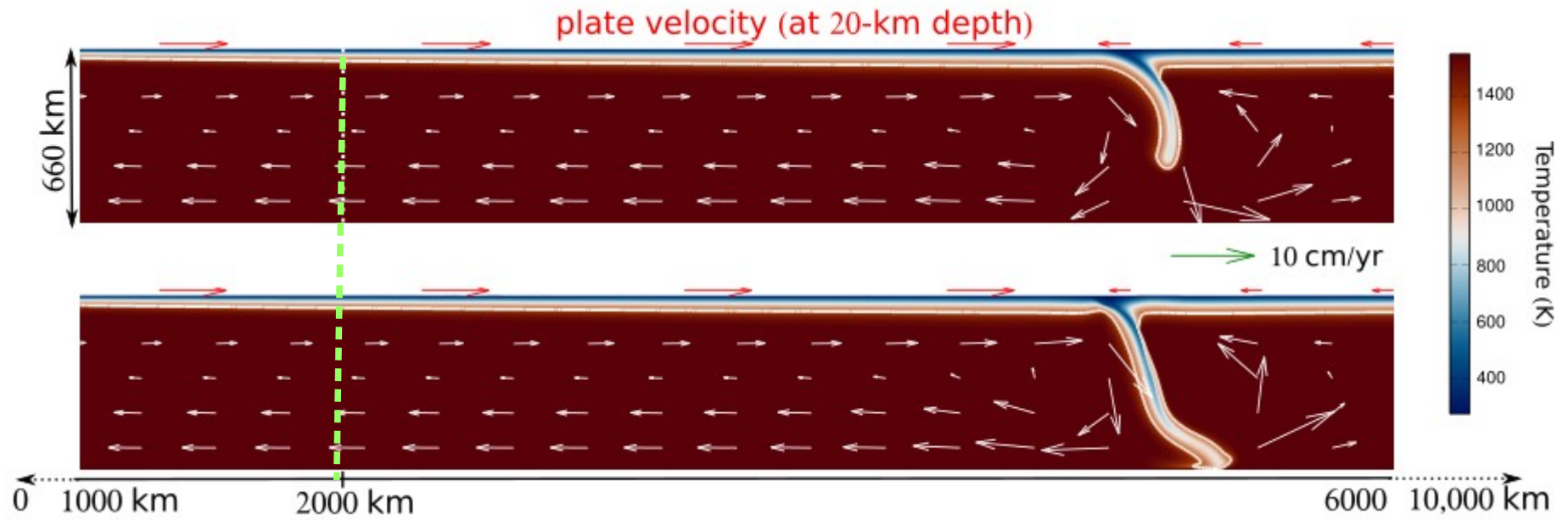
B) Poiseuille flow



C) Combined flow
(Couette and Poiseuille)



Transient asth. flow driven by subduction

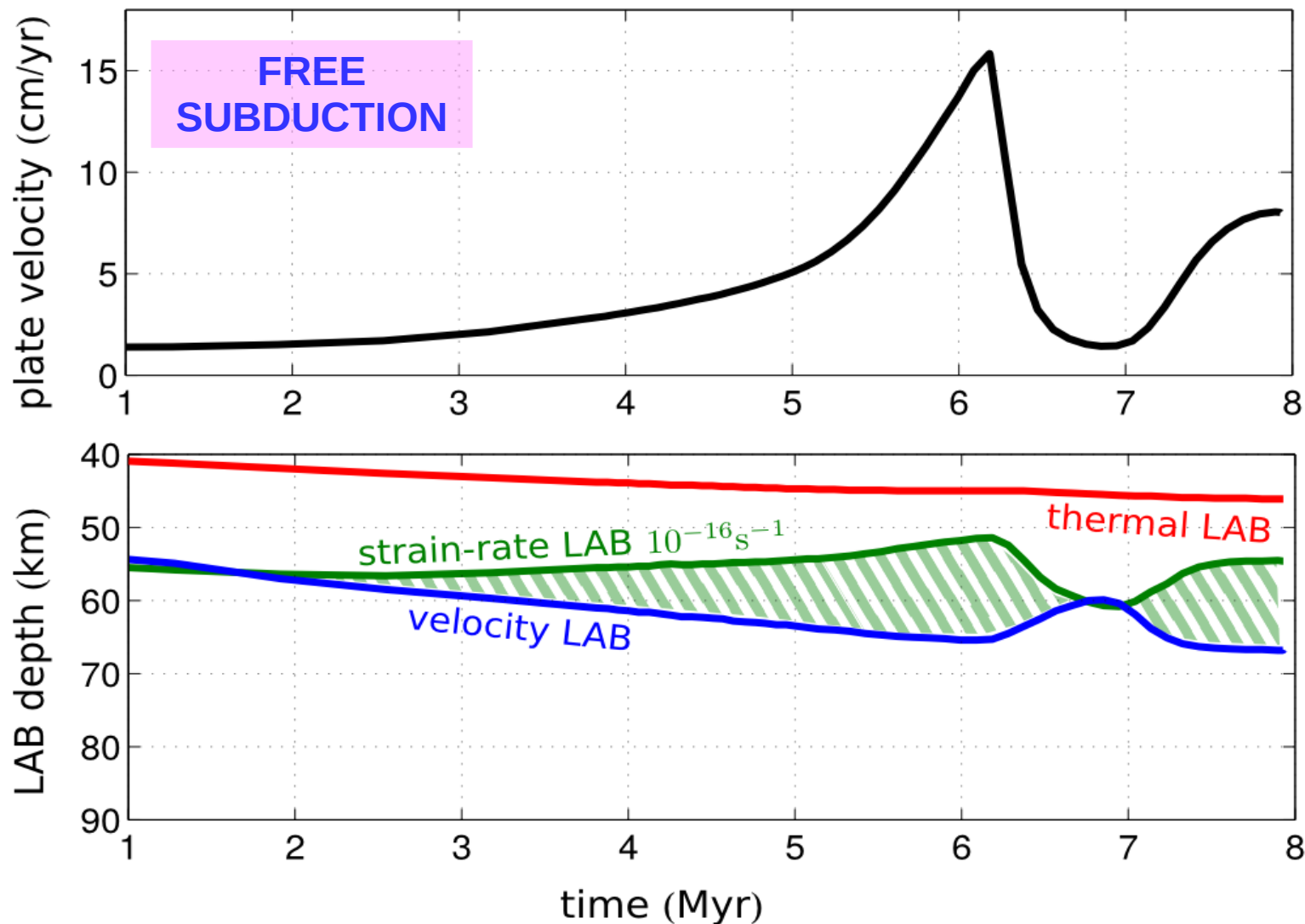


> during plate acceleration, the deformed base of the 'constant-velocity' plate thickens

> after plate deceleration, small sub-plate strain rate may not be enough to "erase" past CPO

- agrees with anisotropy vs. current plate motions (Debayle & Ricard, 2013)

- it takes ~ 30 Myr to accumulate 100 % strain at a rate of 10^{-15} s^{-1}

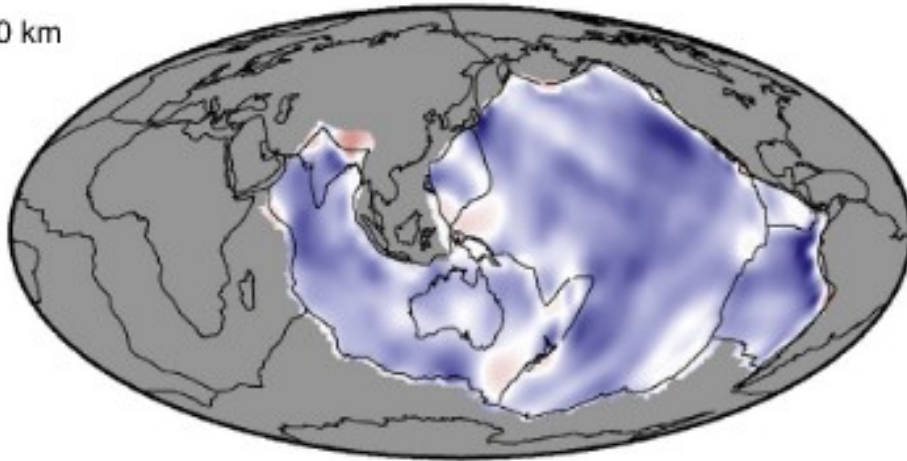


Agreement between fast direction of S_v waves and present-day absolute plate motion (from NUVEL-1A)

blue = parallelism

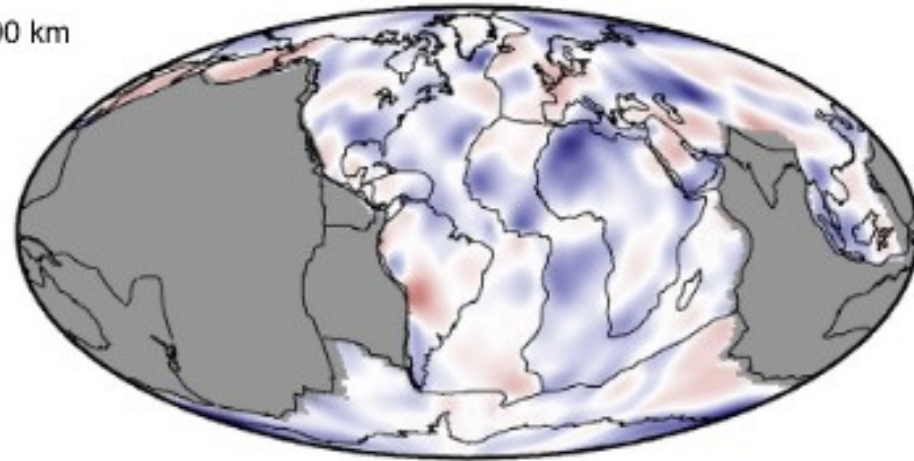
red = orthogonality

200 km



Fast-moving plates

200 km

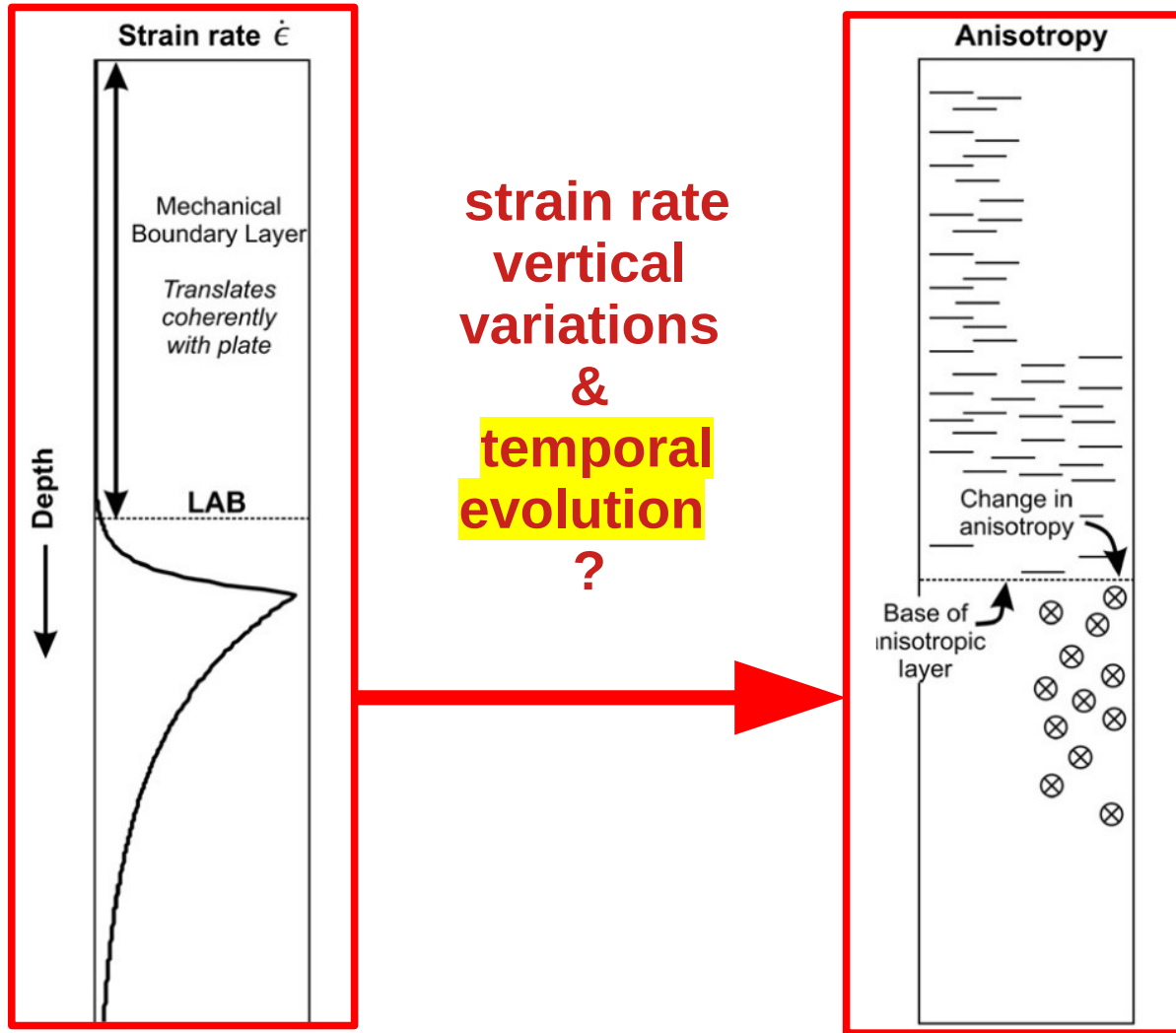


Slow-moving plates

insufficient strain rates in the asthenosphere to generate CPO aligned with plate motion in less than 30 Myr

Debaille & Ricard, 2013

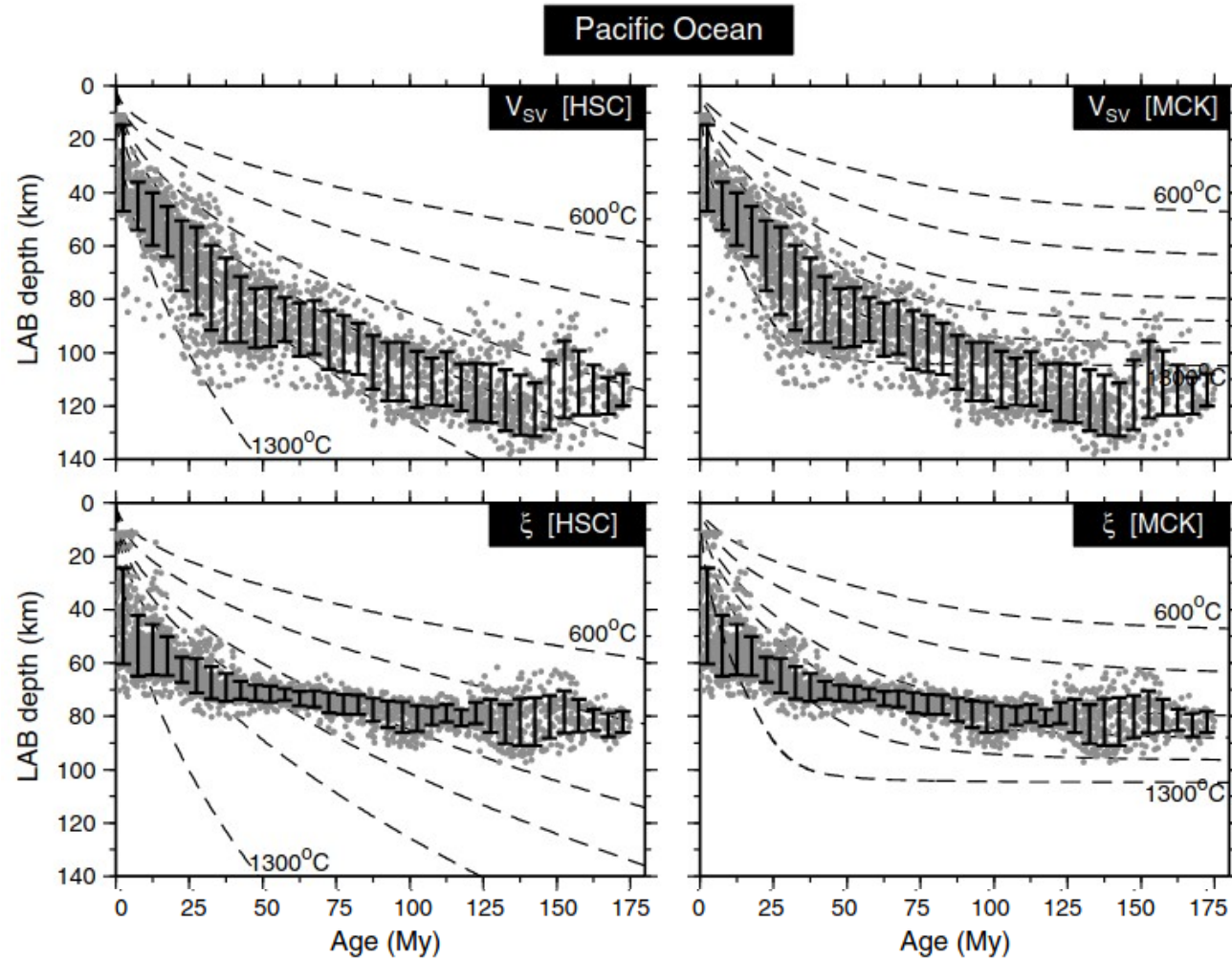
Seismic anisotropy as a proxy for the lith-asth transition ?



Eaton et al.,
2009

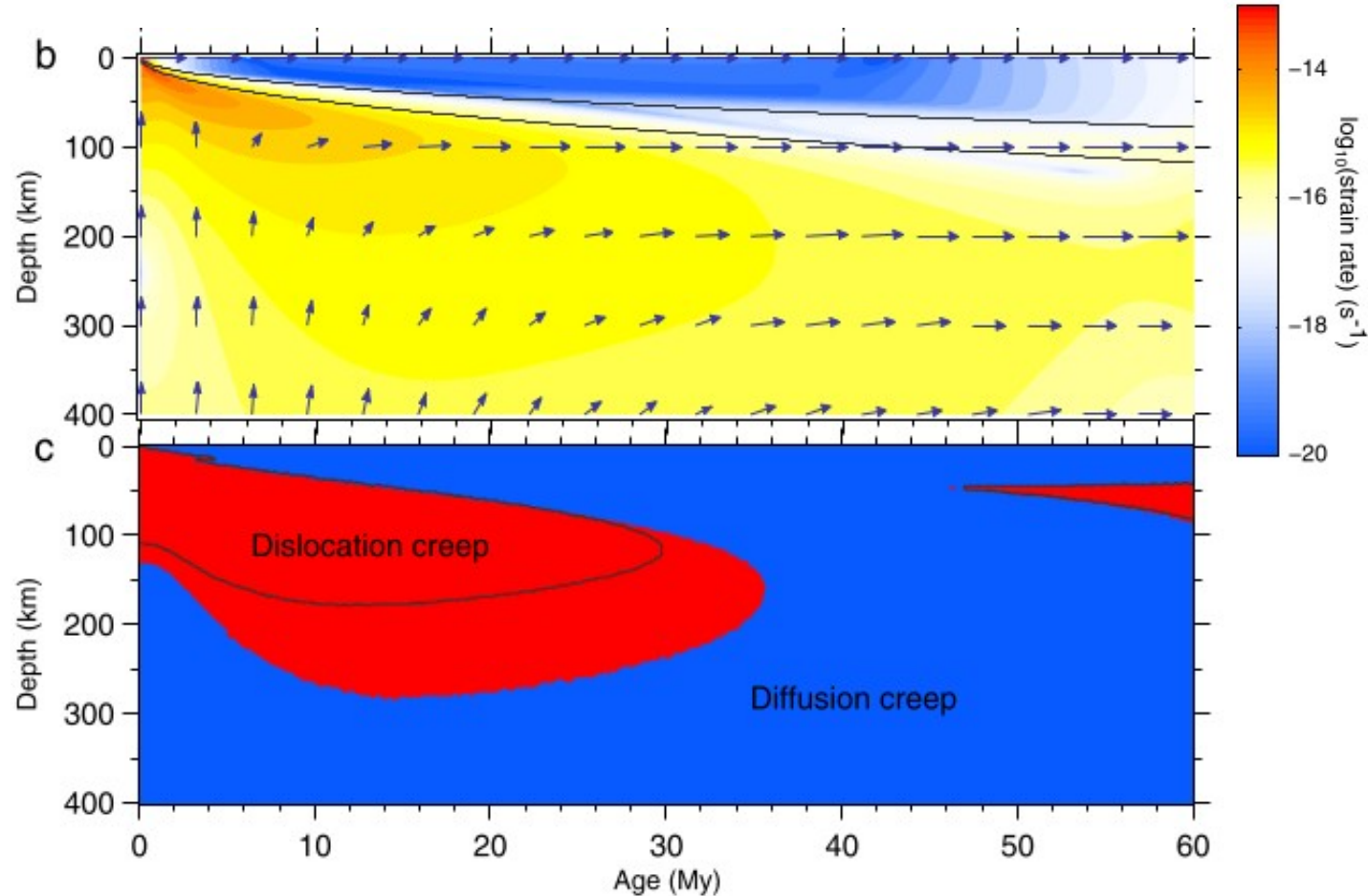


Age-independent radial seismic anisotropy



Burgos et al., 2014

Dislocation & diffusion creep regimes below an oceanic plate

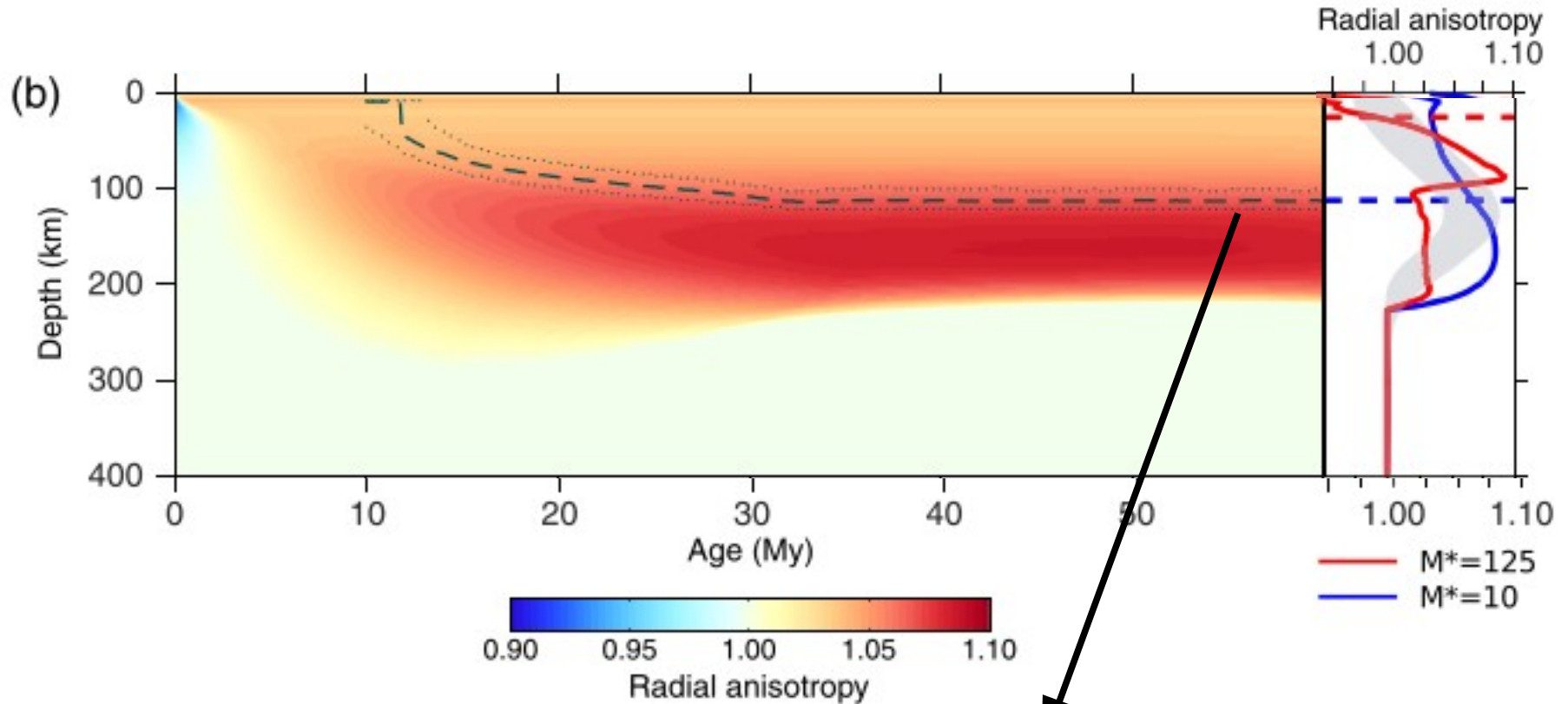


- steady-state **plate-driven flow** of upper mantle
- composite rheology
diffusion
+ **dislocation creep**
- D-Rex calculation of anisotropy

Hedjazian, Garel+.
EPSL 2017



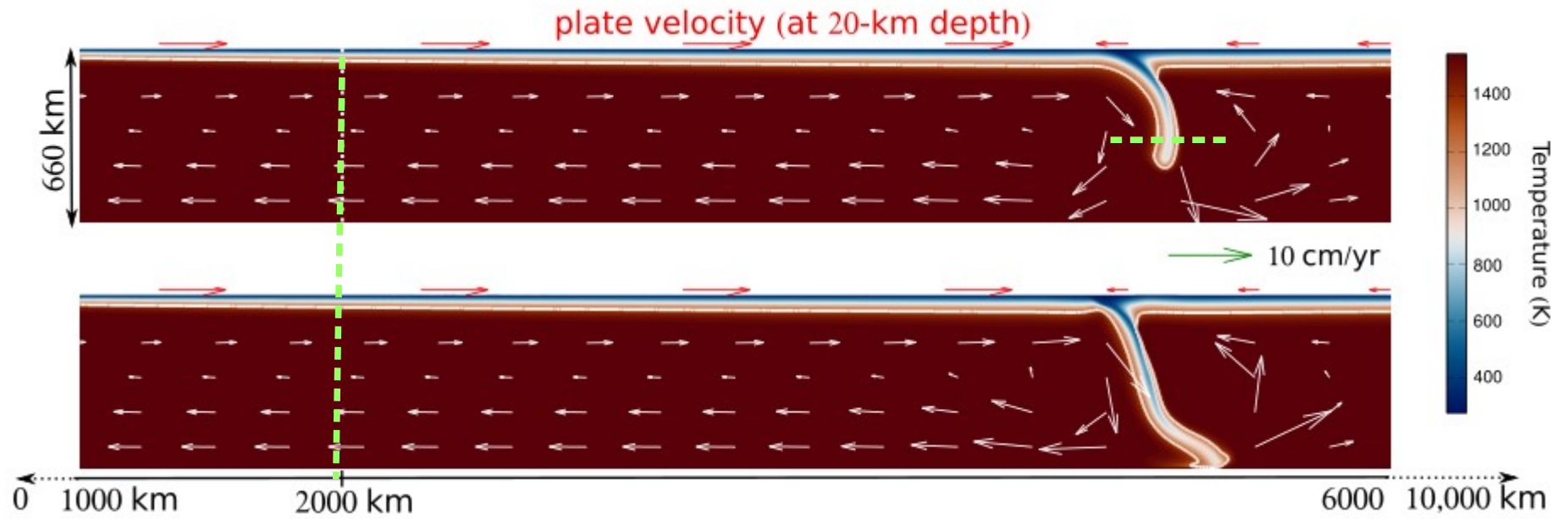
Dislocation & diffusion creep regimes below an oceanic plate

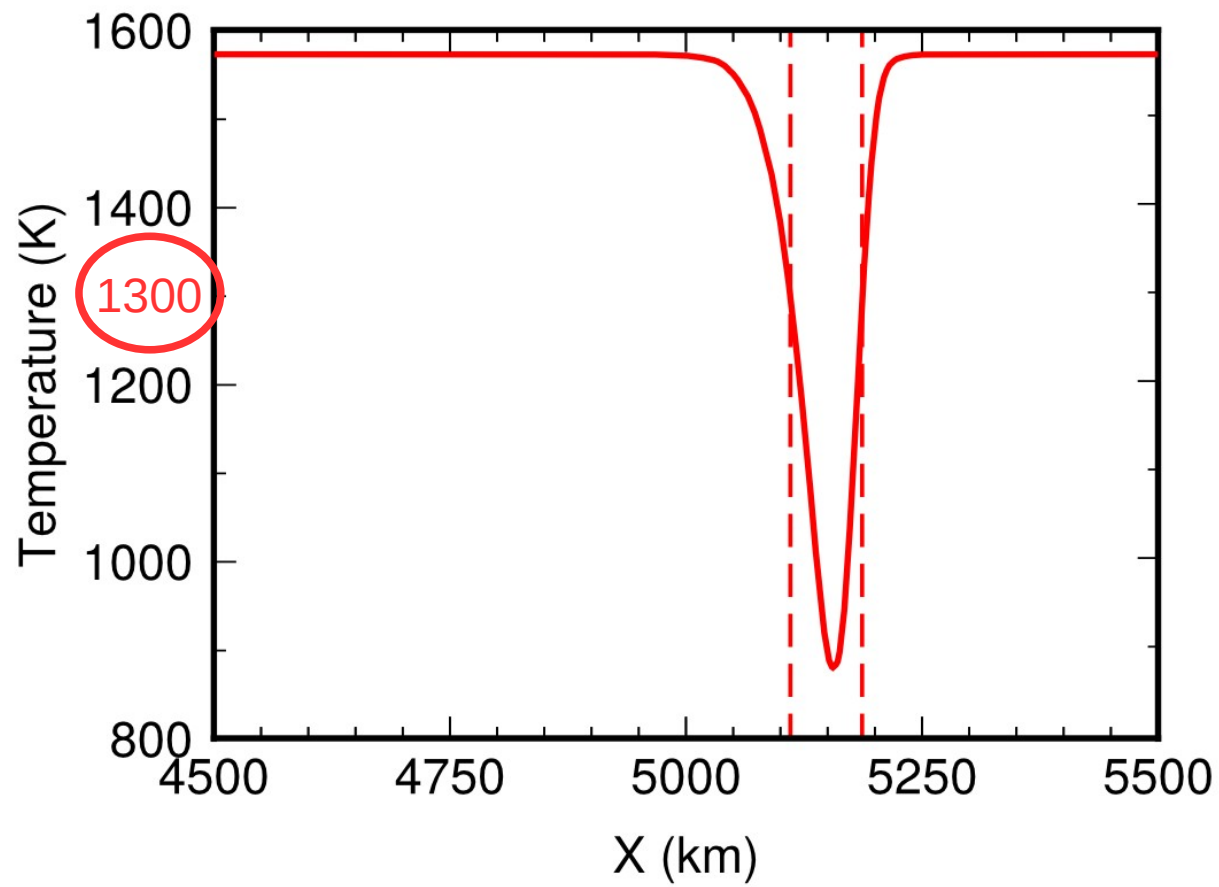


age-independent radial seismic anisotropy !

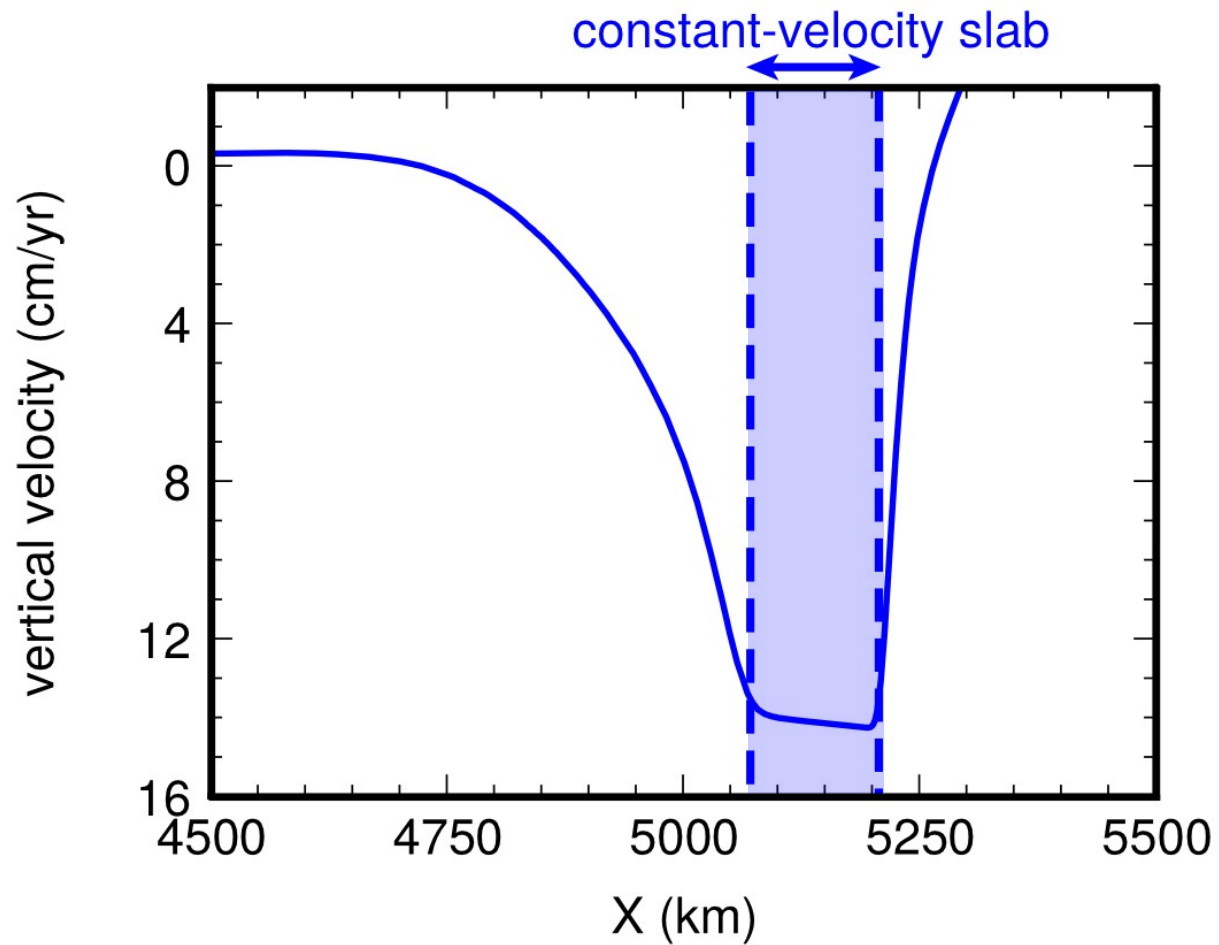
Hedjazian et al.,
EPSL 2017

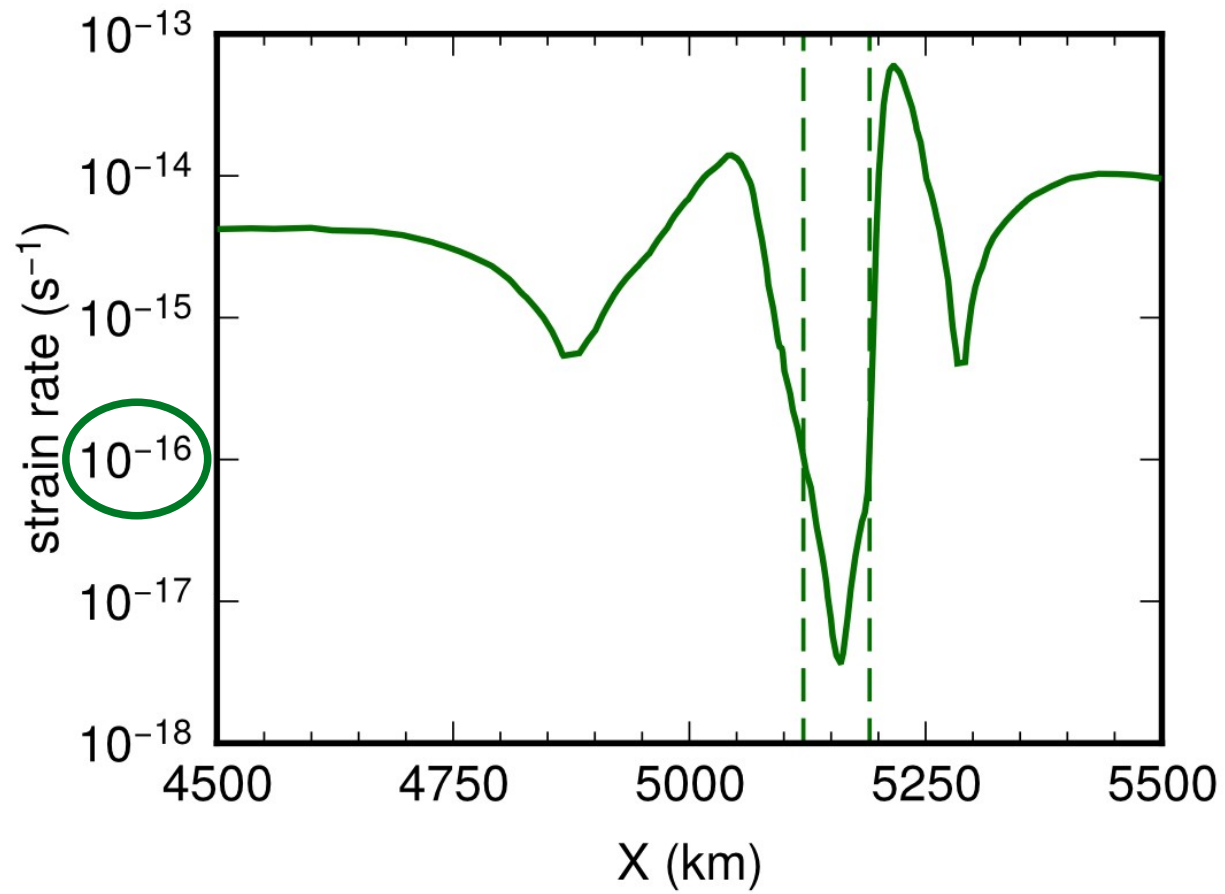
LAB in fast sinking slabs ? (upper mantle)



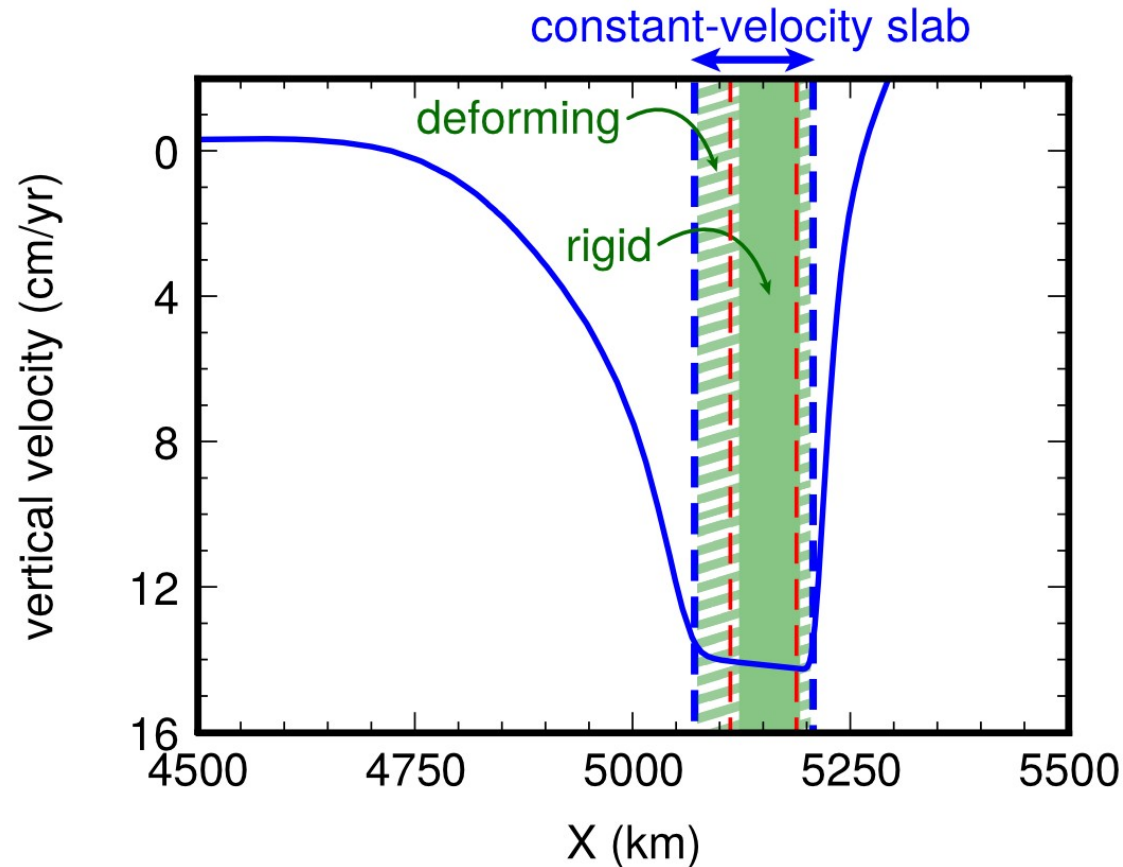


Garel & Thoraval, PEPI, 2021





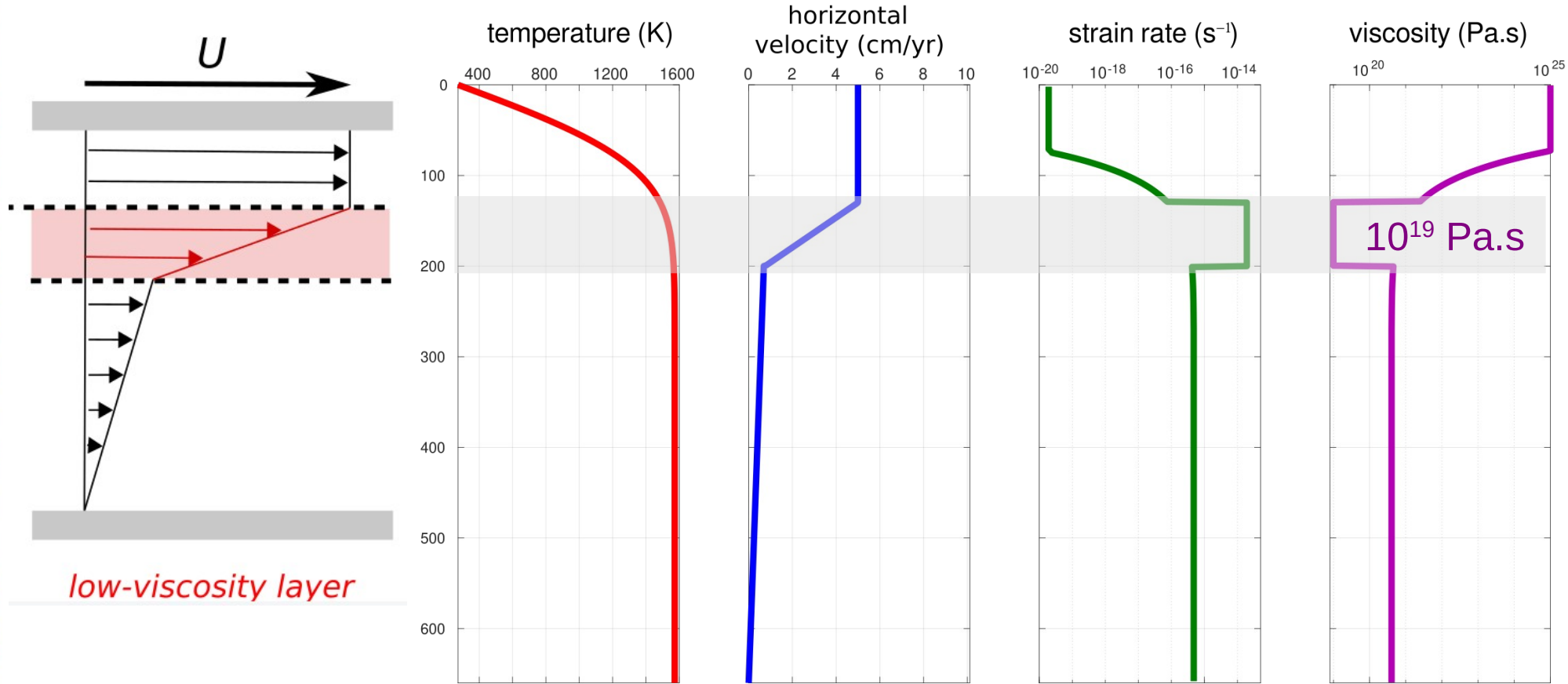
Horizontal profile across a fast-sinking slab



volume of downdip-dragged asthenosphere
dependent on

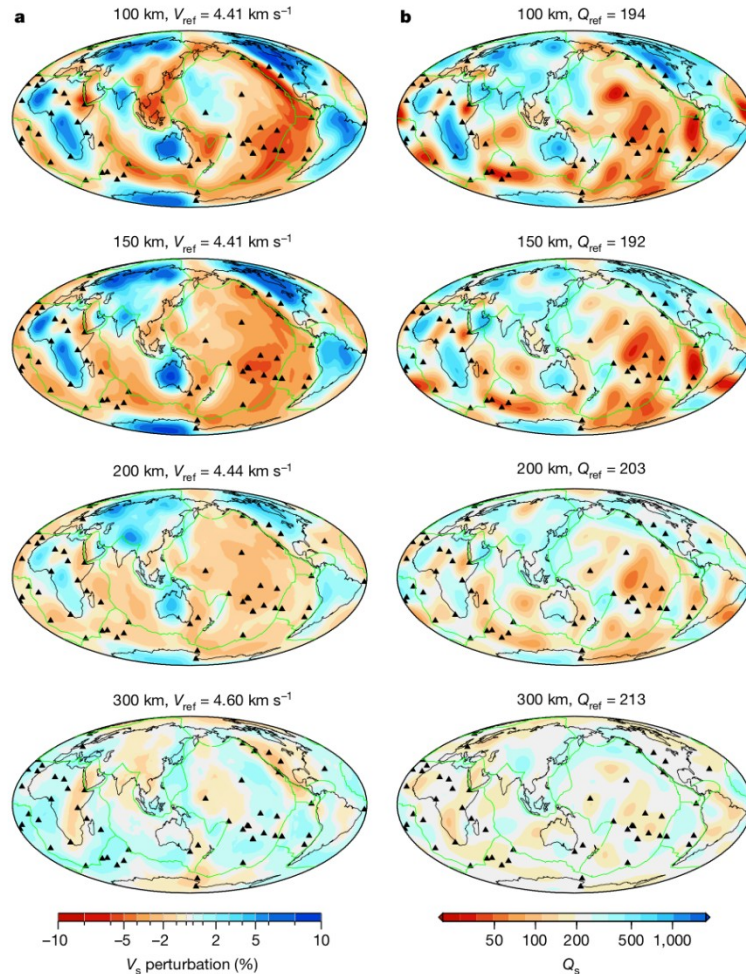
- slab sinking velocity
- **viscosity profile**
(dependent on rheological parameterizations)

A decoupling low-viscosity layer ?



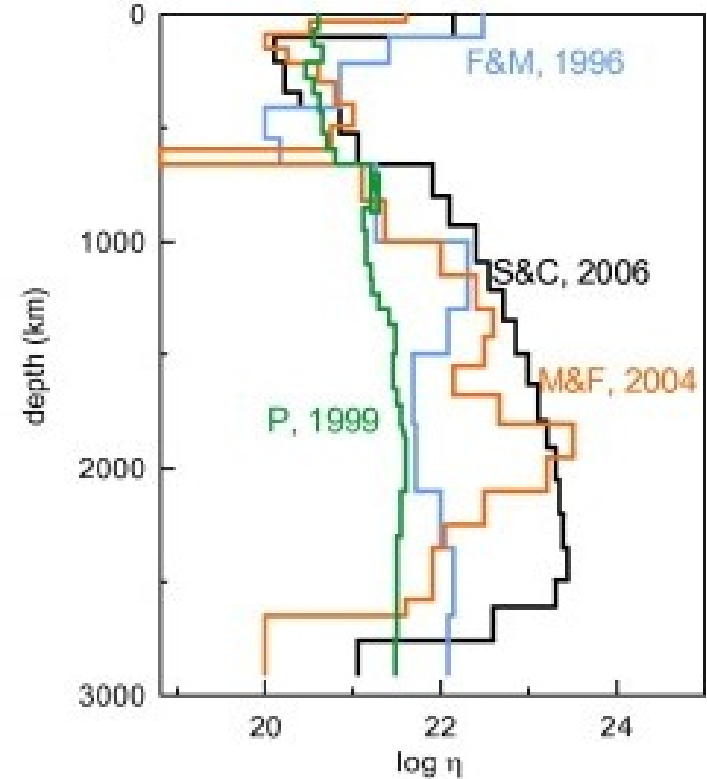
A low-viscosity layer below the plates ?

Shear velocity, attenuation



Debayle et al.,
Nature, 2020

joint inversion of geoid and postglacial rebound data

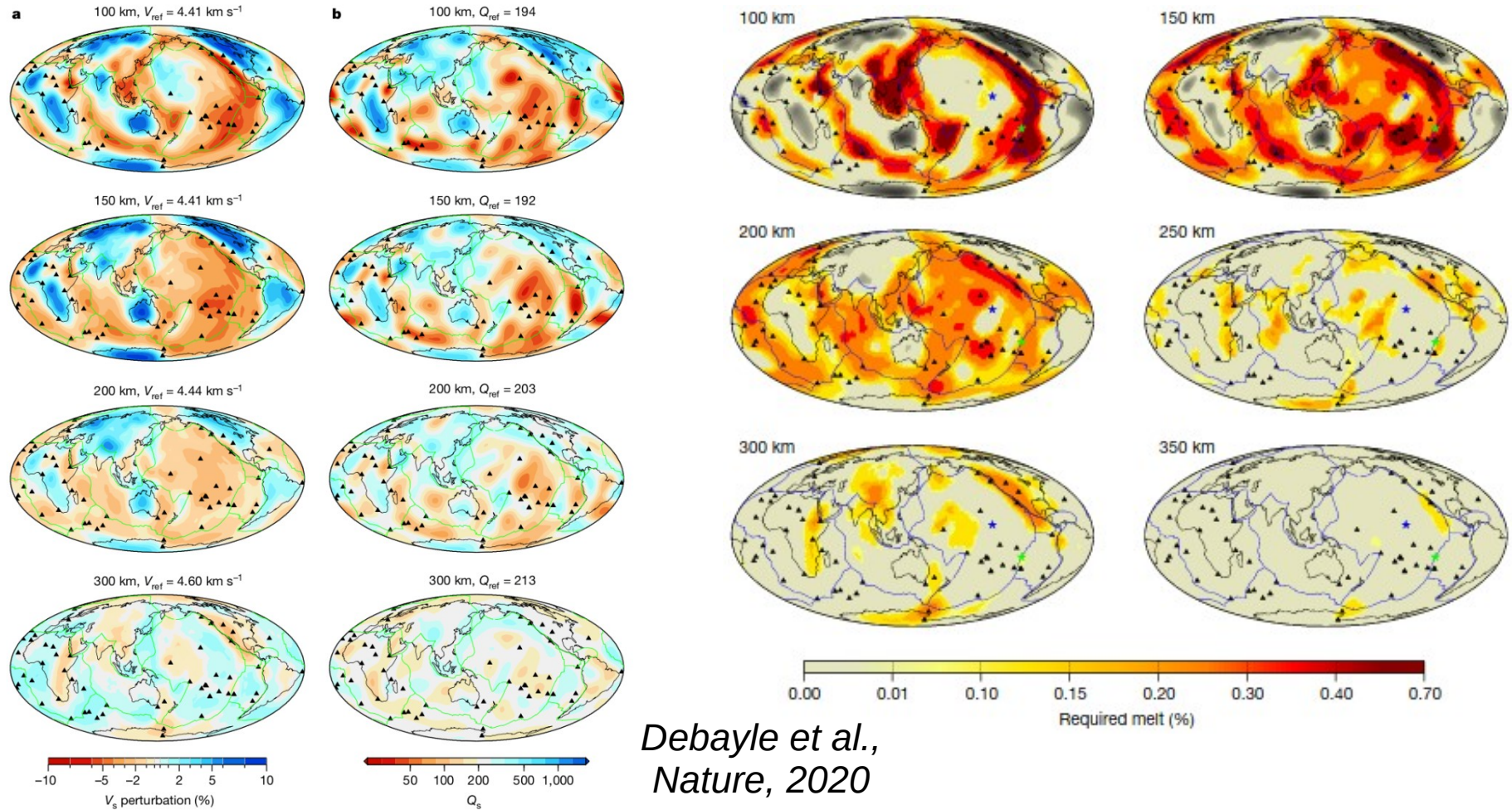


Compilation in Cizkova
et al., *PEPI*, 2012



A low-viscosity layer below the plates ?

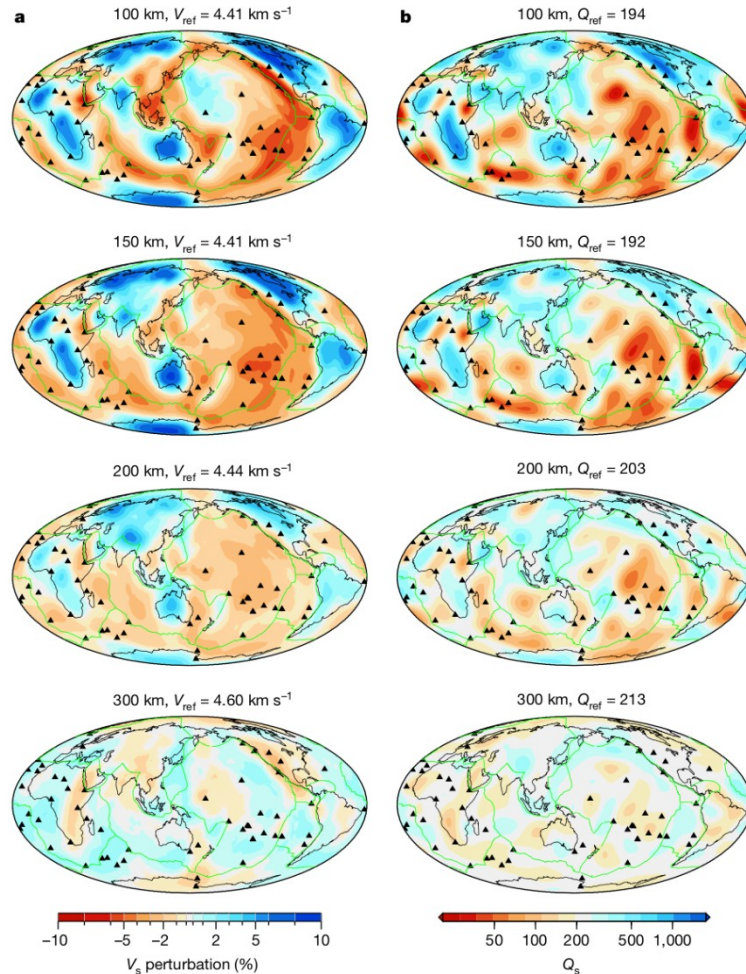
Shear velocity, attenuation -----► partial melt ?



Debayle et al.,
Nature, 2020

A low-viscosity layer below the plates ?

Shear velocity, attenuation -----► partial melt ?



-----► grain size variations ?

creep laws → low-viscosity layer possible with certain rheological parameters

how weak?

10^{19} Pa.s ?

how much weaker?

$10\text{-}100 \times ?$

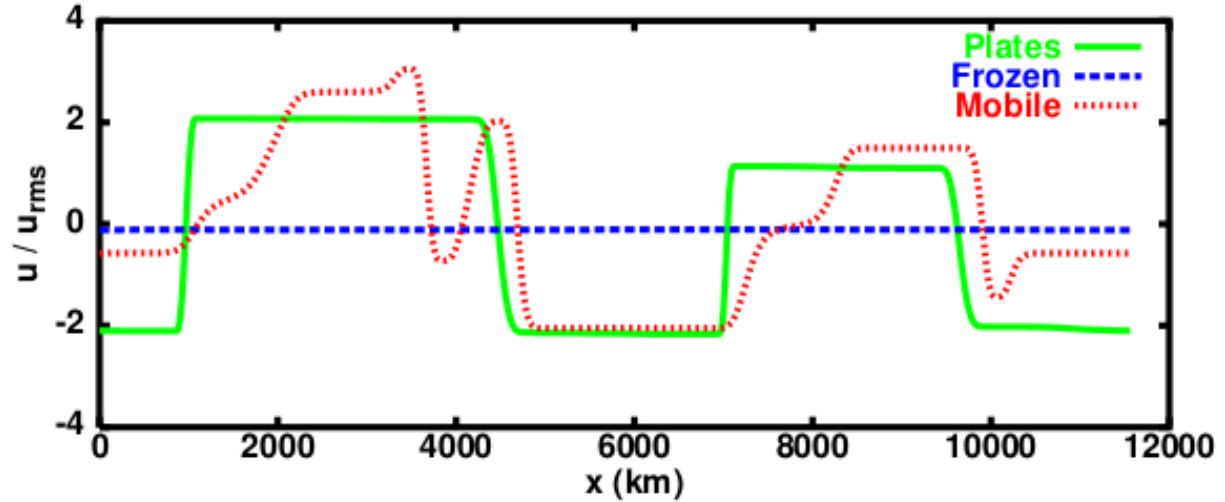
how thick?

$< 250 \text{ km depth ?}$

Debayle et al.,
Nature, 2020

A low-viscosity layer favoring plate tectonics regime

localizing
deformation at
plate boundary



η^* / σ_y (MPa)	Inf	150	100	80	60	50	30	10
1	●	●	◐	◐	◐	●	●	●
100	●	●	●	◐	◐	◐	●	●
300	●	●	●	■	■	◐	●	●
500	●	●	◐	■	■	◐	●	●
1000	●	●	■	■	■	◐	●	●
10000	●	■	■	■	◐	◐	●	●

● Frozen ■ Plates ● Mobile

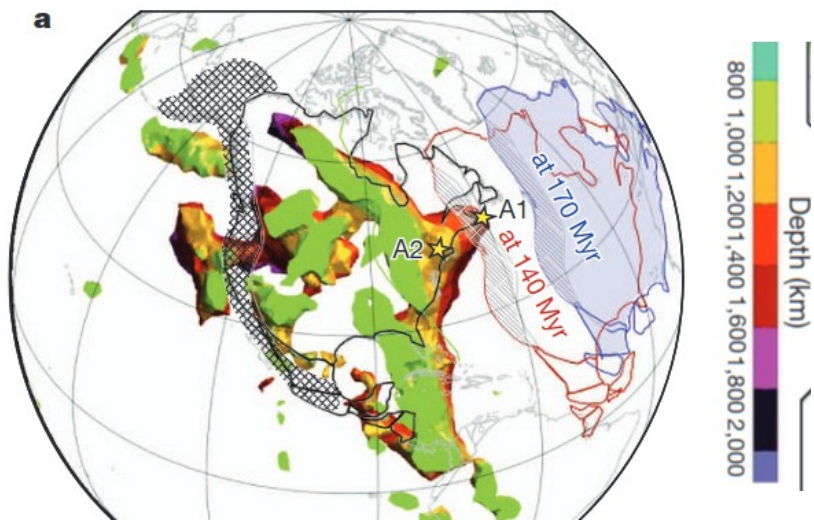
lithosphere mobility as a
function of **yield stress** σ_y
and viscosity contrast η^* for a
sub-plate low-viscosity layer

Richards G^3 2001



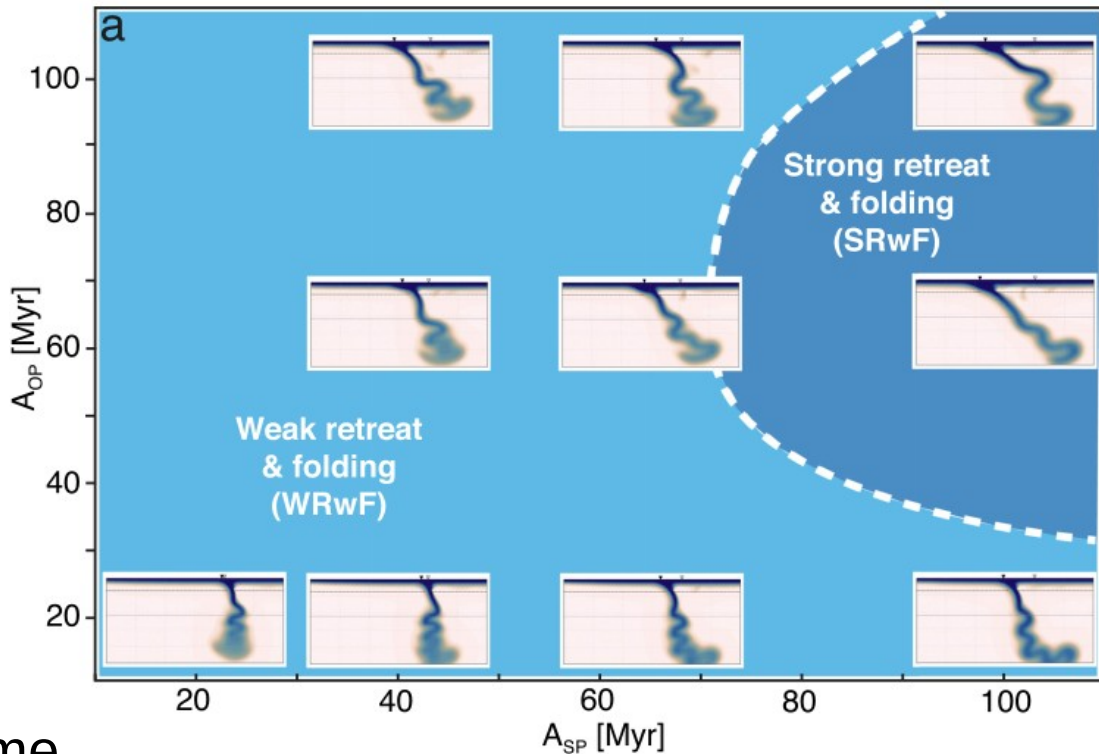
A low-viscosity layer reconciling plate vs. slab velocities

- 5 cmr/yr surface plates vs. 1 cm/yr sinking slabs > folding !?
- imaged slab “walls” = piled & folded ?



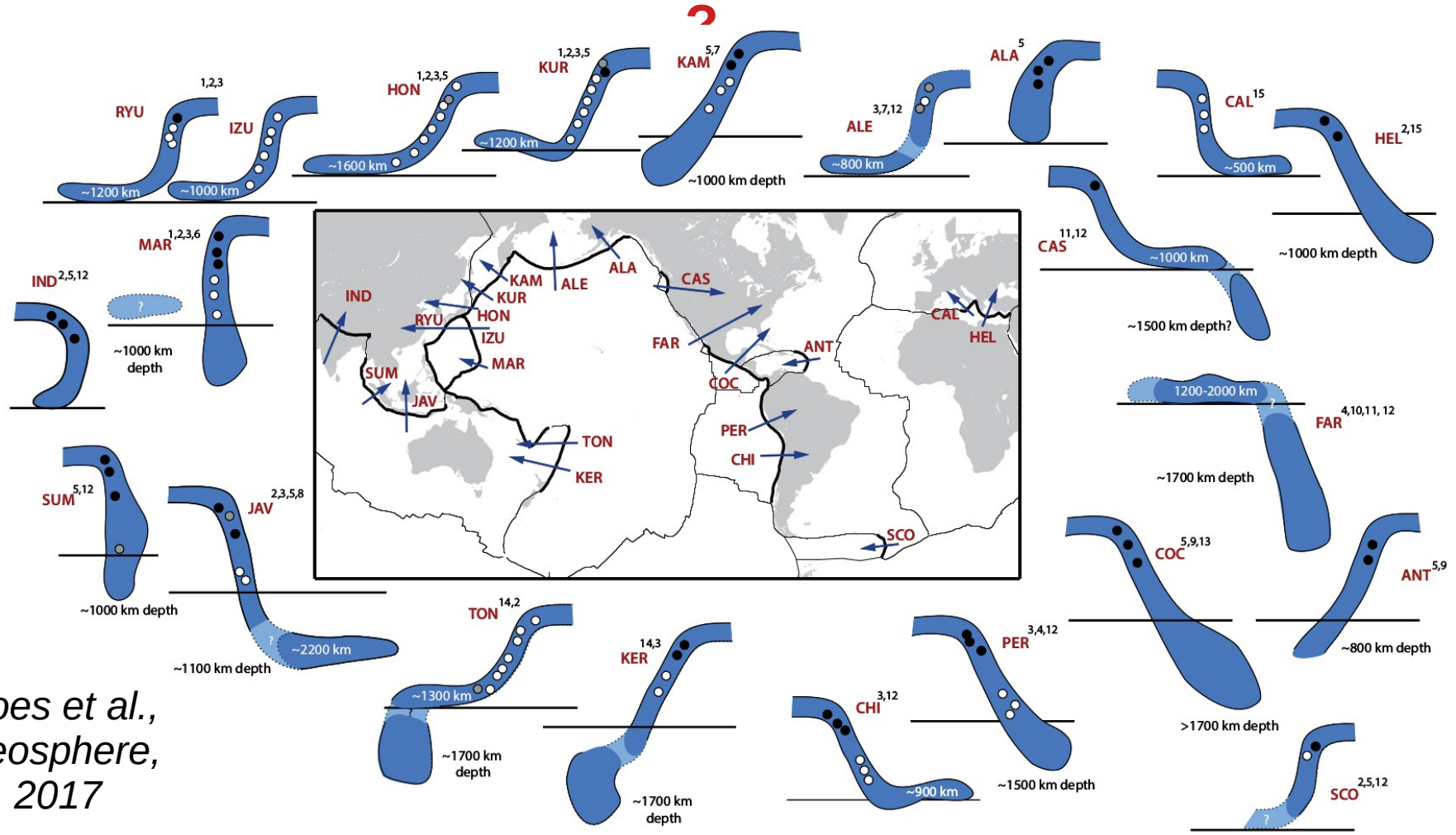
but most subduction models shows inclined/retreating slabs...

> dominance of vertical-folding regime if addition of **subplate 2-5x weak layer**



Cerpa, Sigloch, Garel+ JGR 2022

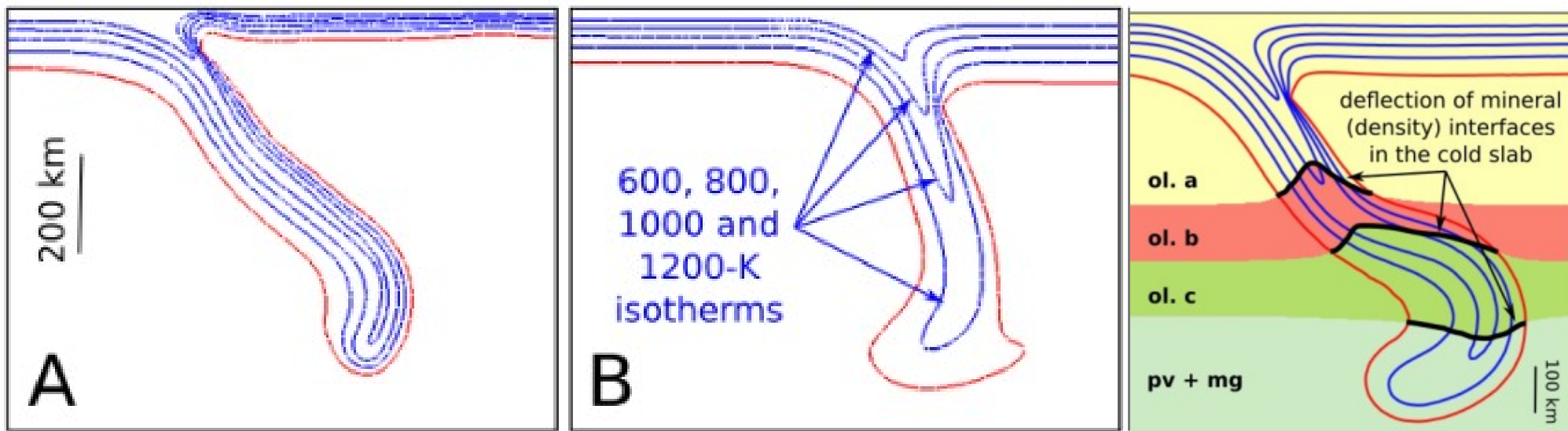
Mass transfers between upper & lower mantle : asthenosphere dragged and trapped by cold deforming slabs



Goes et al.,
Geosphere,
2017

Thermal structure of slabs imaged by spatial gravimetry & gradiometry to retrieve their history of sinking and deformation

starting PhD of **Xavier Vergeron**



CNES GraviSSym
project
(PI Cécilia Cadio)



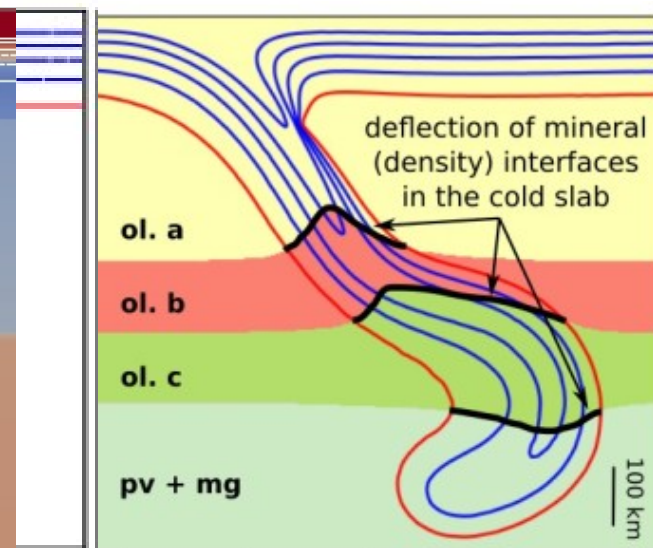
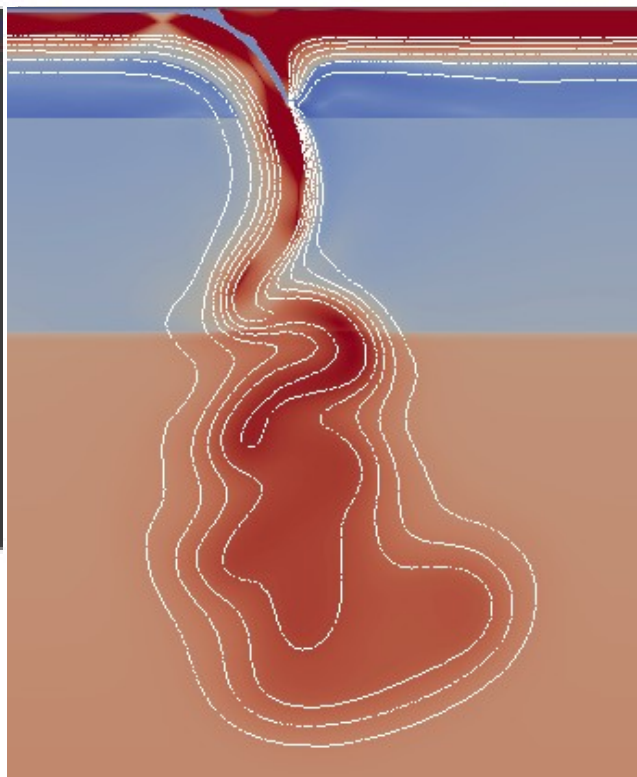
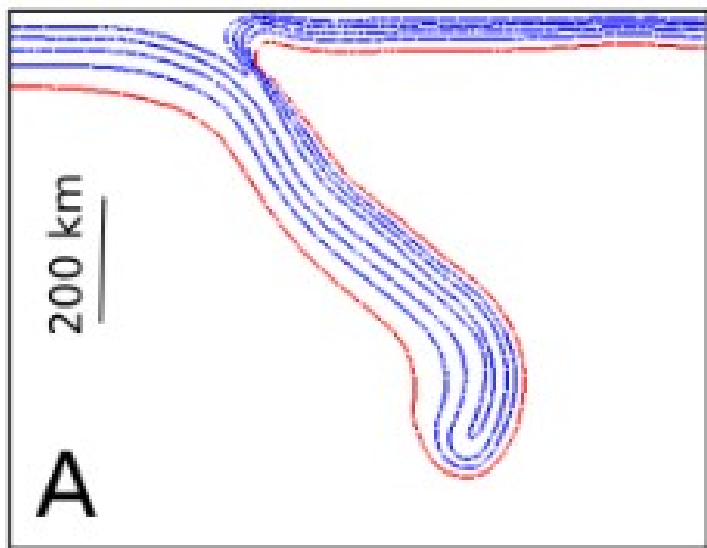
anr

RheoBreak

ANR-21-CE49-0009

Thermal structure of slabs imaged by spatial gravimetry & gradiometry to retrieve their history of sinking and deformation

starting PhD of **Xavier Vergeron**



CNES GraviSSym
project
(PI Cécilia Cadio)



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Breaking plates from below with dislocation creep ?



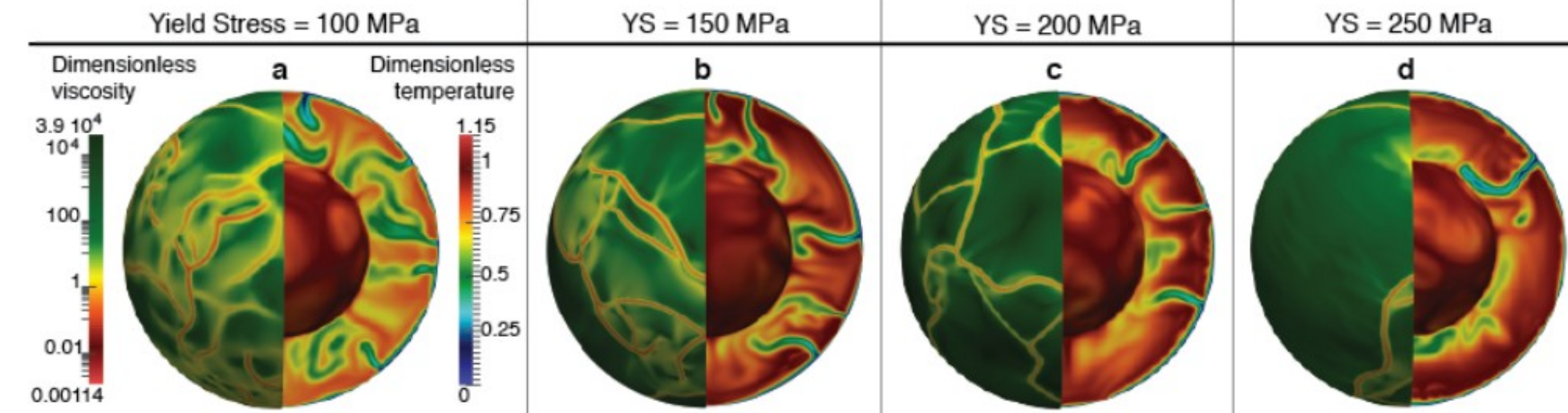
ANR-21-CE49-0009

“Memory” processes proposed for deformation localization feedbacks

- grain size growth vs. recrystallization under stress
- strain weakening
- shear heating

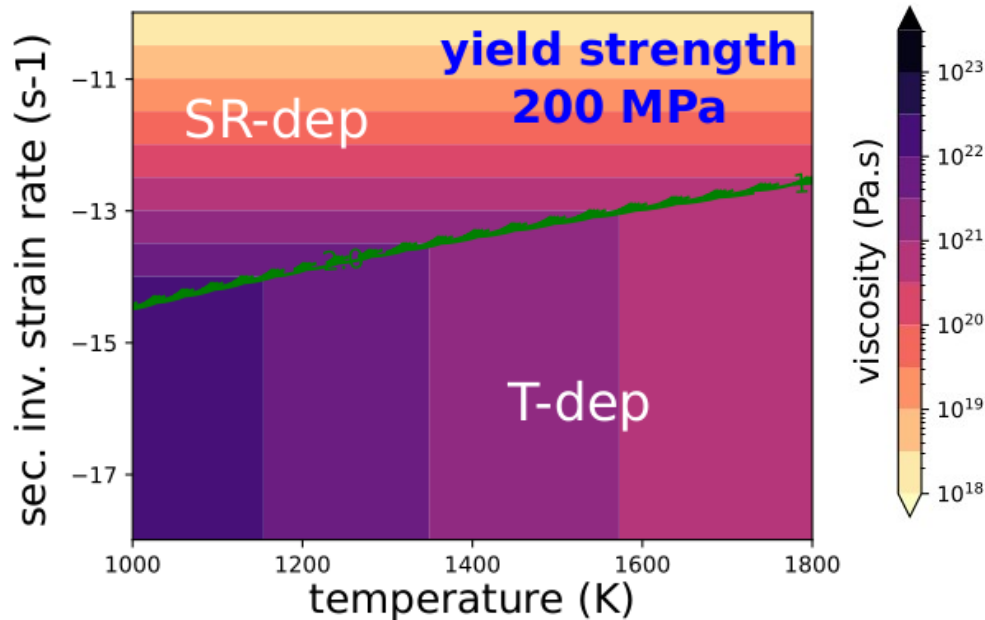
**No memory effect :
shear thinning**

strain-rate dependent viscosity
from dislocation creep ($n > 1$)



Mallard, Coltice+ Science 2016

yield-stress rheology successful in localizing “viscous” plate boundaries



$$\eta_{\text{eff}} = \min \left[\eta(z, T), \frac{\sigma_y(z)}{2\dot{\epsilon}} \right]$$

mantle + cold plates (**T-dep**) weak plate boundaries (**SR-dep**)

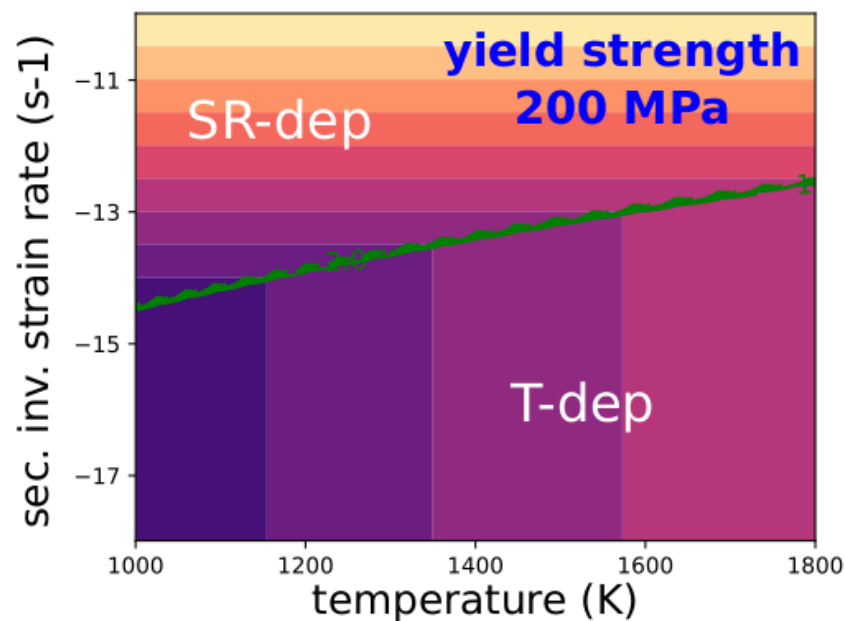
Tackley 1998, 2000 ; Mallard+ 2016 ;
Coltice+, 2017, Arnould+ 2018....



composite dislocation + diffusion creep

compatible with deformation experiments :
**effective viscosity dependent on both
temperature and strain rate**

Gouriet+ EPSL 2019
Garel+ EPSL 2020



**yield-stress rheology successful in
localizing “viscous” plate boundaries**

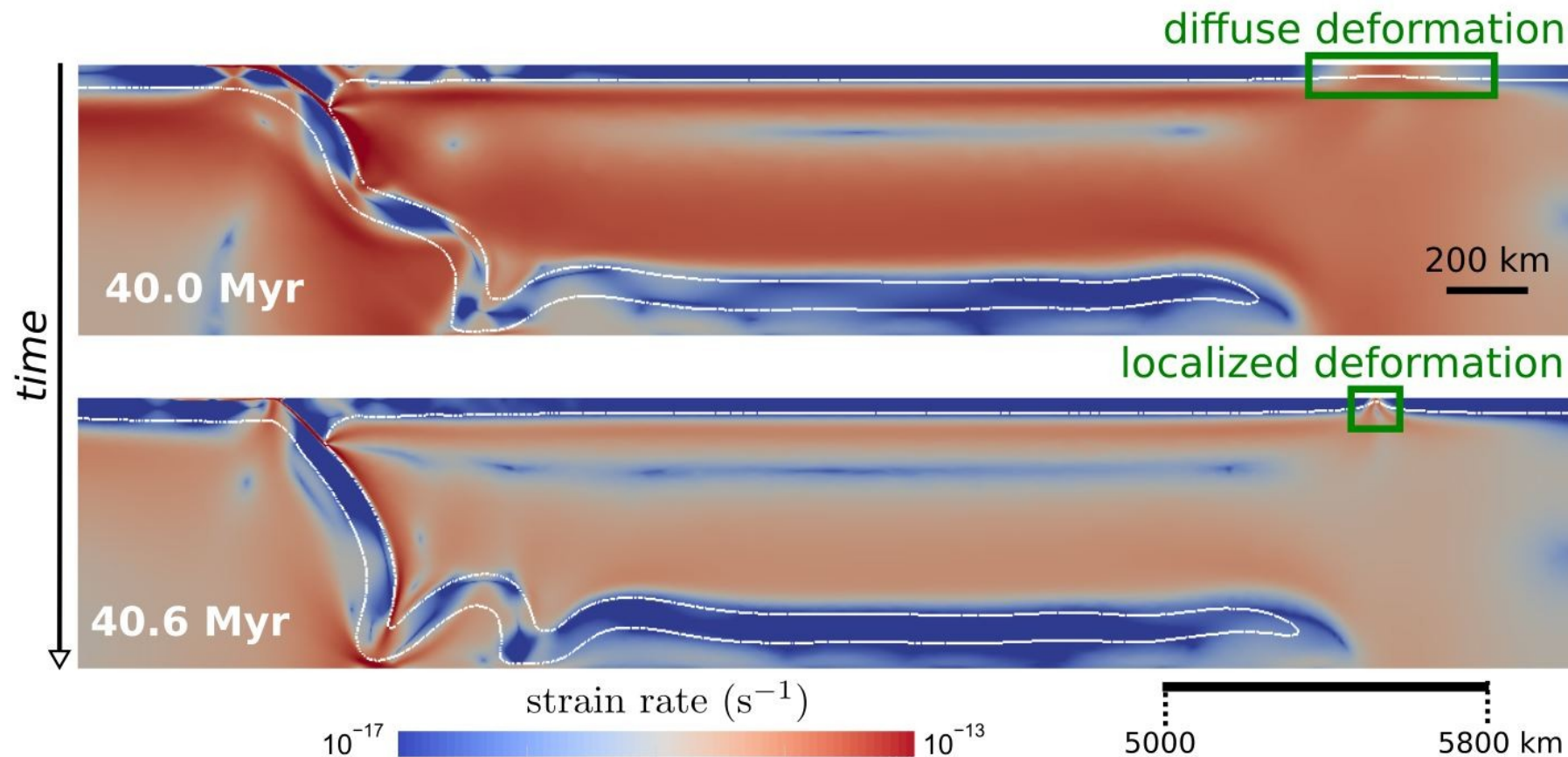
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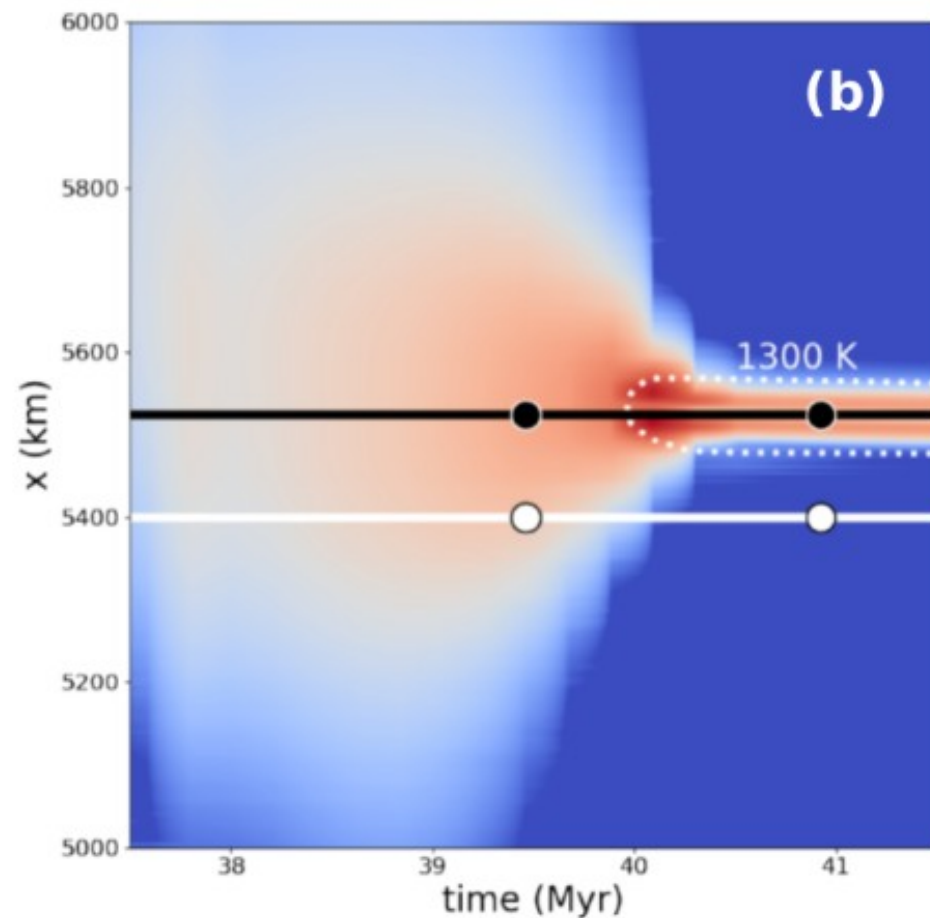
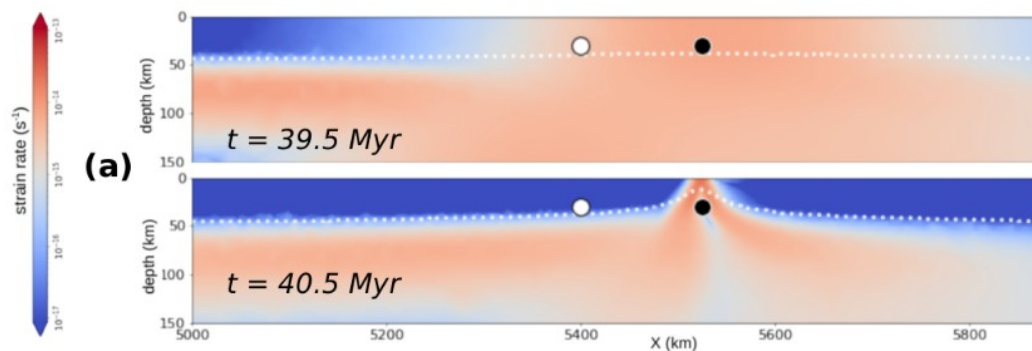
Tackley 1998, 2000 ; Mallard+ 2016 ;
Coltice+, 2017, Arnould+ 2018....

Model set-up enabling intraplate deformation localization

- accelerated subduction dynamics with a fixed upper plate - *Alsaif + EPSL 2020*
- olivine composite rheology *HT and LT disl. creep* + *diff. creep.* + *brittle*



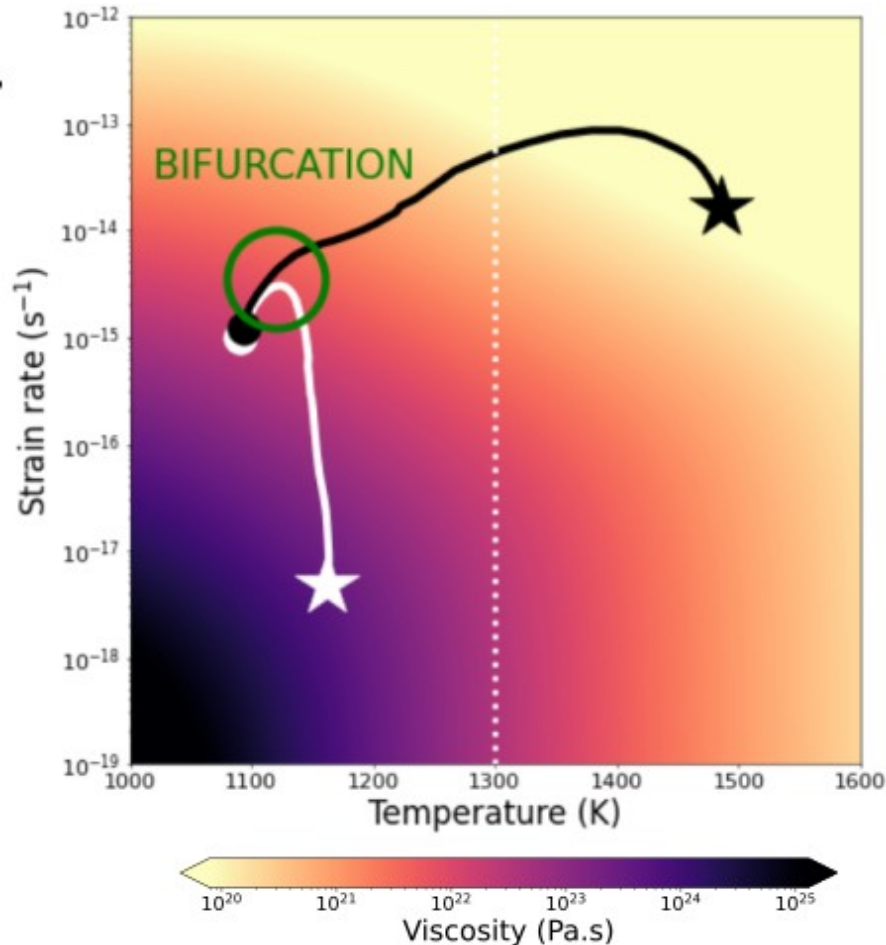
Model set-up enabling intraplate deformation localization



> follow the properties of two sample points through time

deformation-Temperature-time paths (*def-T-t paths*)

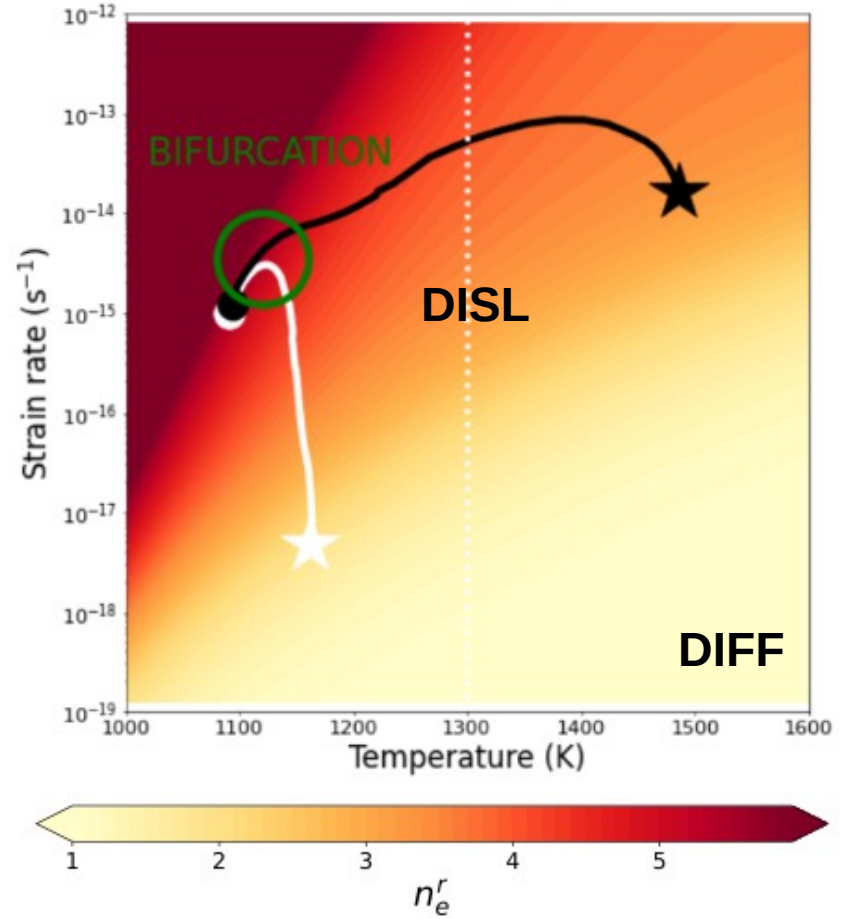
onto a background of viscosity



rheological effective stress exponent for a composite viscosity (diff + disl)

$$\dot{\epsilon} \propto \sigma^n$$

$$n_e^r(T, \dot{\epsilon}) = \frac{\partial \log(\dot{\epsilon})}{\partial \log \sigma}$$



time-dependent "dynamic" effective stress exponent

Montesi & Zuber, 2002

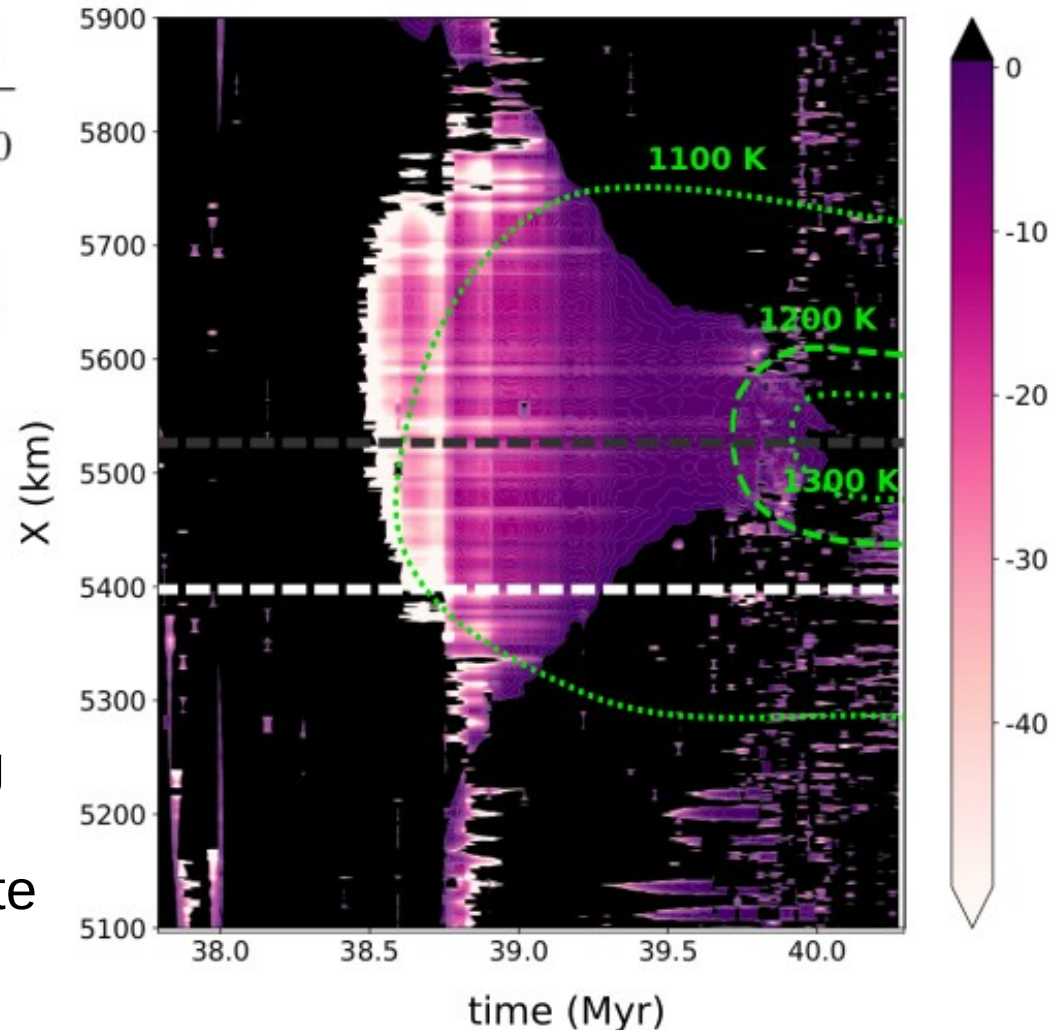
$$\frac{1}{n_e} \equiv \frac{\chi_0}{\sigma} \frac{d\sigma}{d\chi_0}$$

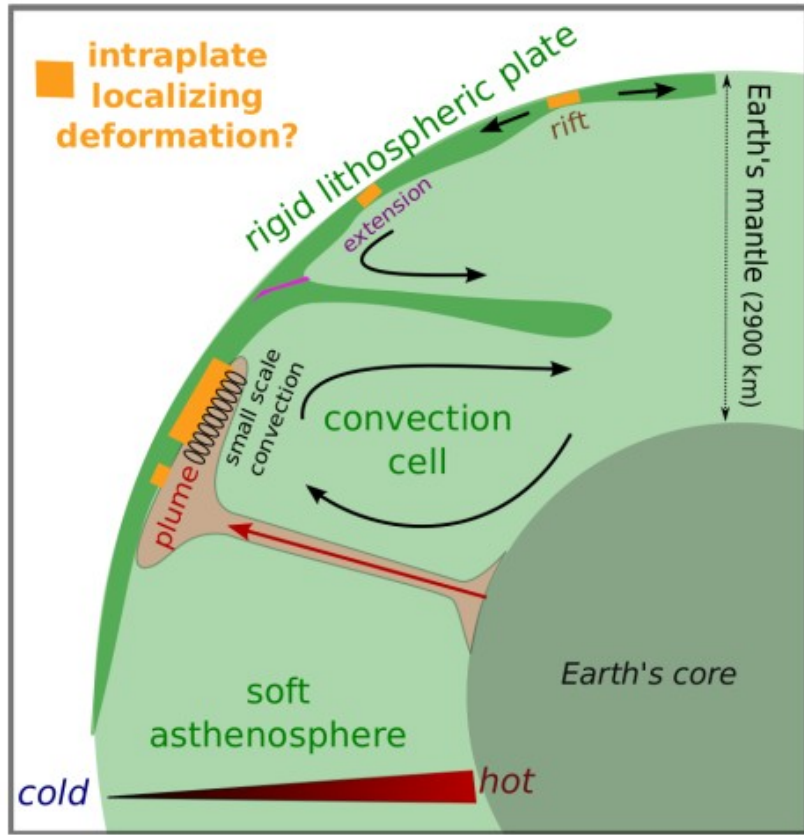
$$\frac{1}{n_e^d} = \frac{d \log(\sigma)}{d \log(\dot{\varepsilon}_{II})} = \frac{d \log(\sigma)}{dt} \frac{dt}{d \log(\dot{\varepsilon}_{II})}$$

*dynamic localization occurs
for negative n_e*

and is strongest for more negative $1/n_e$

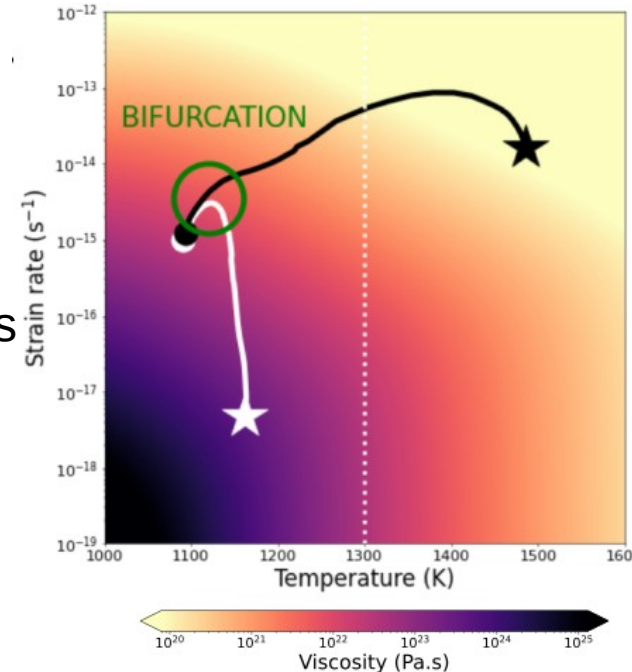
deformation localization
may be caused by a hot upwelling
but hot regions then return to
non-localizing, low-deformation state





- deformation-temperature-viscosity-time paths calculated in models of **various geodynamical contexts**

- rheological param. compatible with def. Exp. but with different gradients of viscosity / exponent n as a function of T and SR



- > why some rheology localize deformation but others do NOT
- > whole-mantle convection models → frequency of plate break-up ?
→ Early Earth ?