

# Composition, structure and dynamic properties of the Earth and planetary interiors: insights from lab experiments

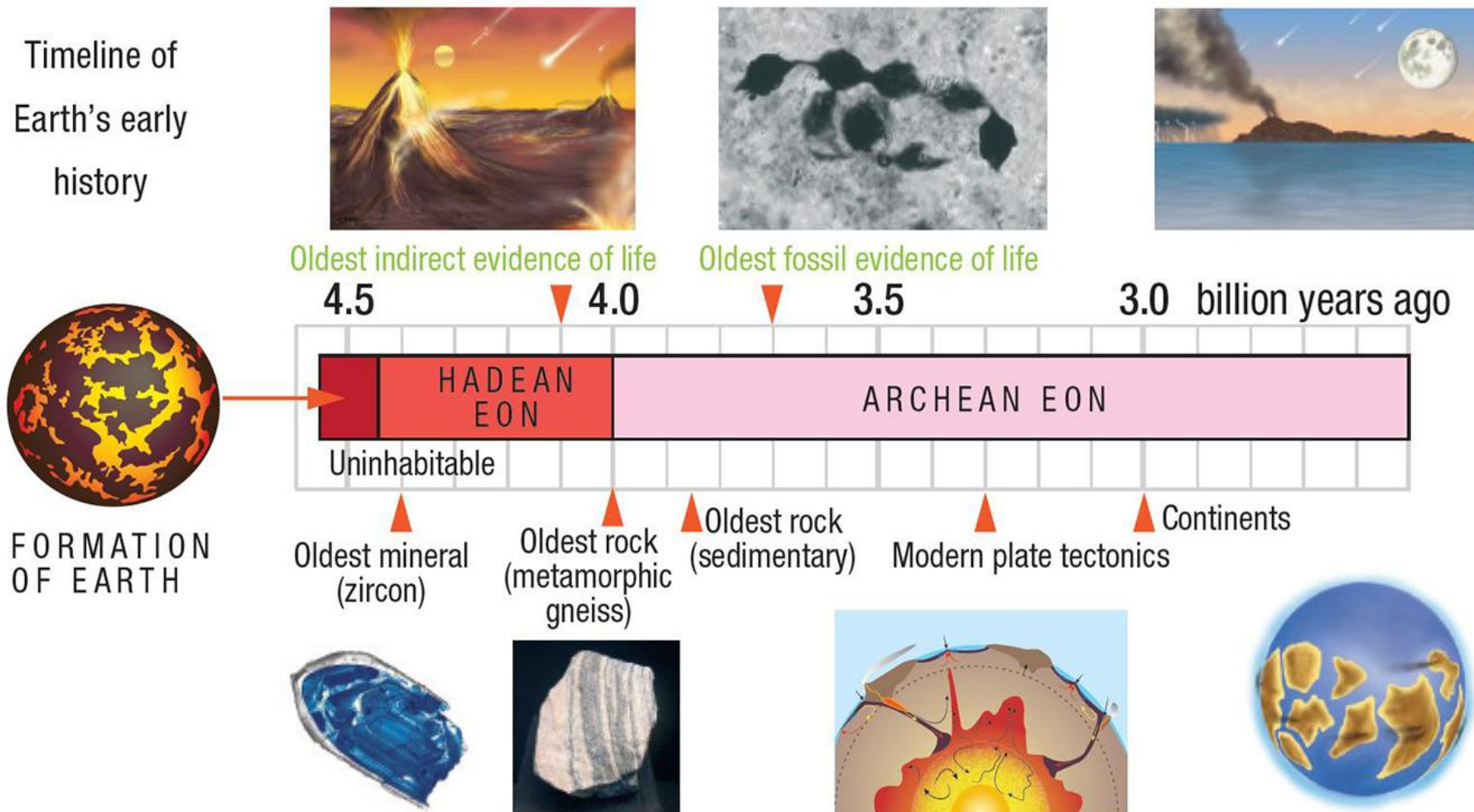
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Clermont Ferrand



Laboratoire Magmas et Volcans  
Clermont Ferrand, France

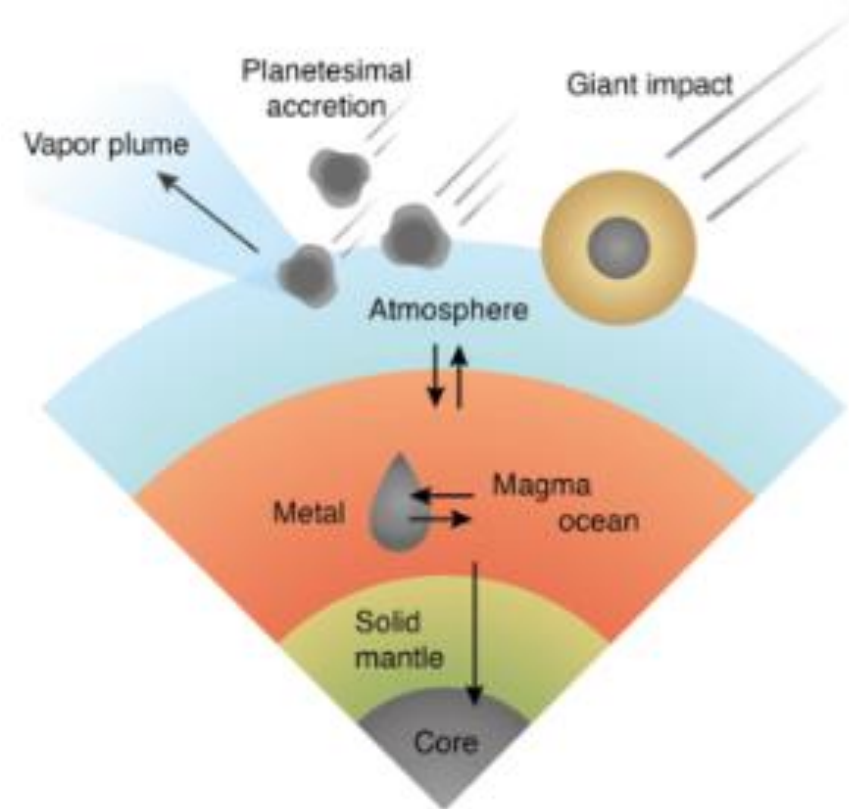
# 1. Introduction

Earth's evolution from a fire-ball state to a habitable planet



# Early Earth

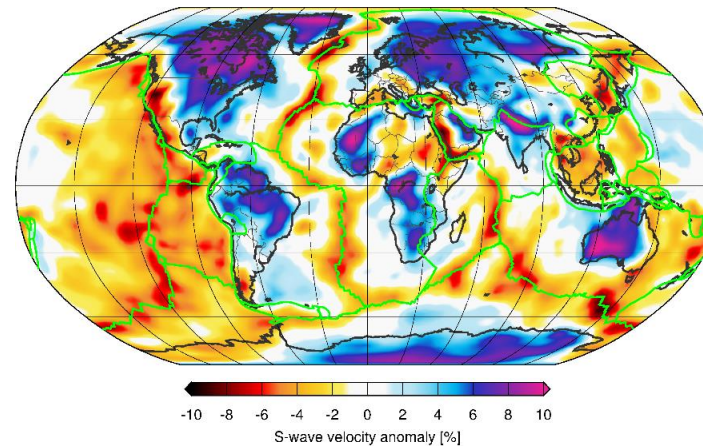
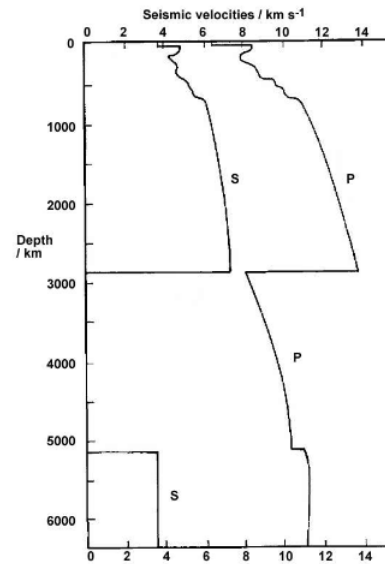
- Formation of the metallic core
- Magma ocean crystallization
- Solidification of the inner core
- Compositional stratification



Source: TITEC

# Present-day Earth

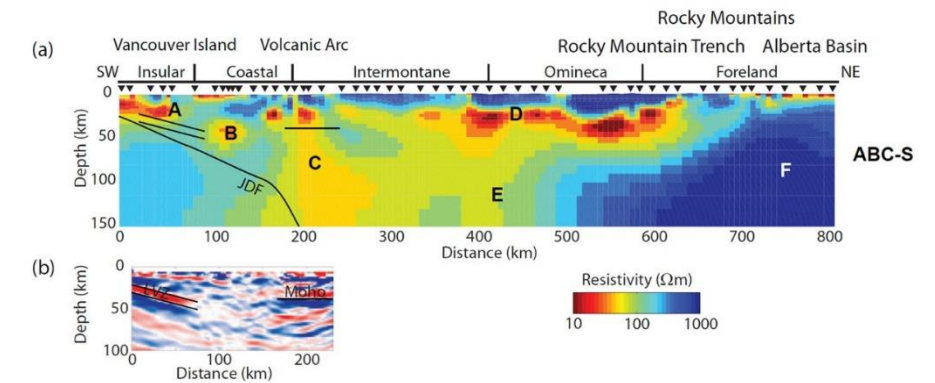
## Seismic wave velocity



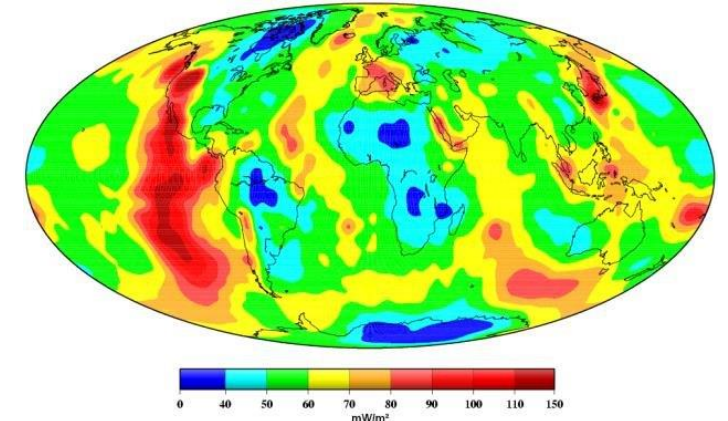
Geophysical profiles represent variation in

- temperature
- composition
- density
- physical state

## Electrical resistivity



## Heat flow



Source: Institute of Physics

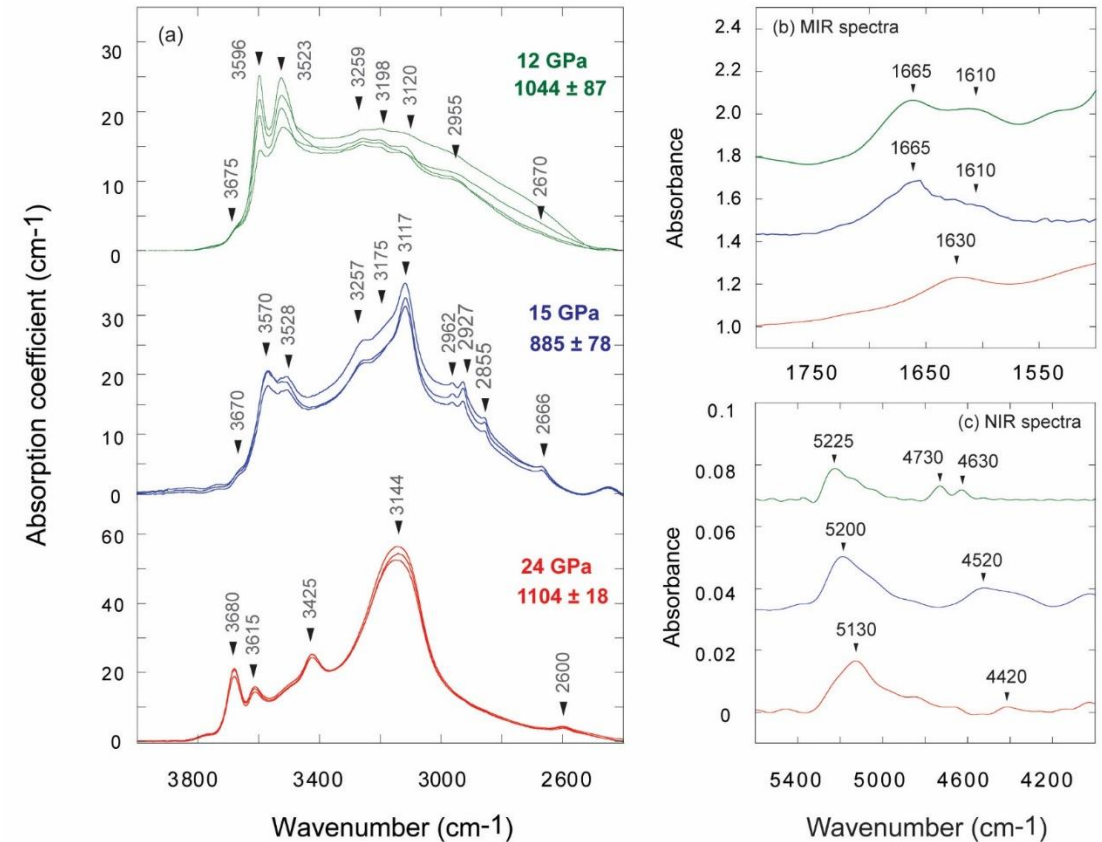
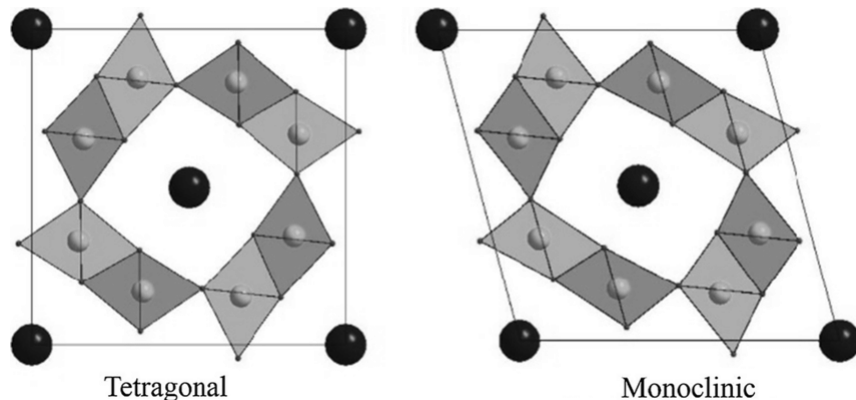
## 2. Applications

- Material circulation
- Mantle heterogeneities
- Heat transport in the Earth and Planetary bodies

# • Transport of water into the lower mantle

- High water contents (2 wt. %) in OIBs fed by plumes rising from the LM (Deschamps et al. 2011)
- Continental sediment components in OIBs (Hart 1988) - EM-II

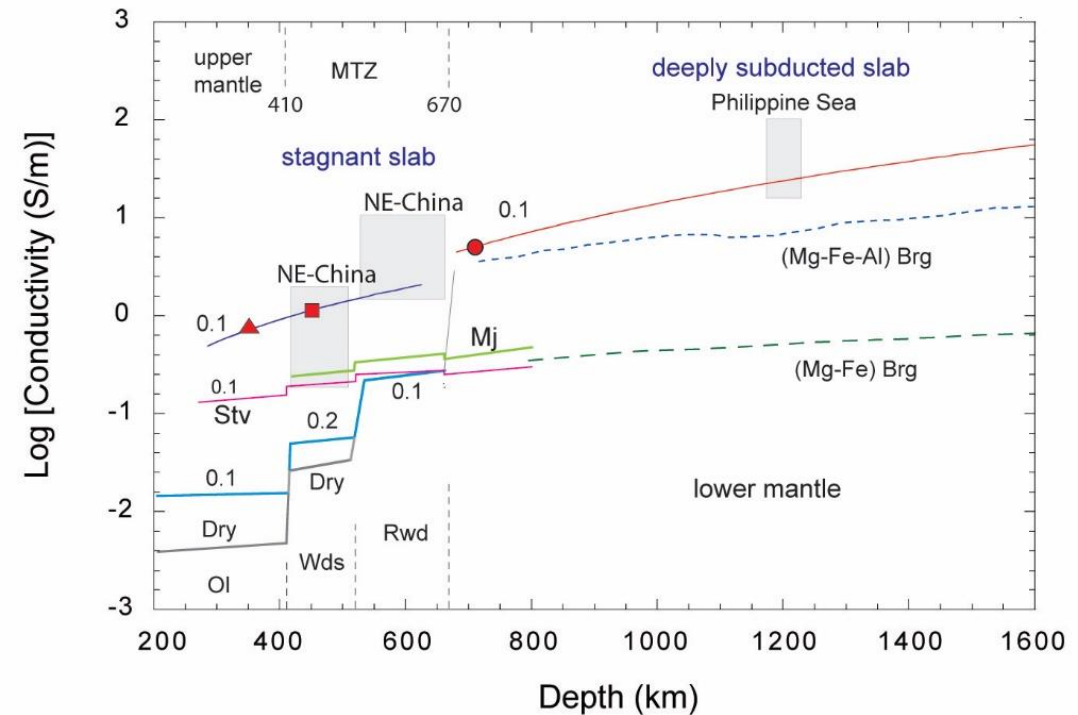
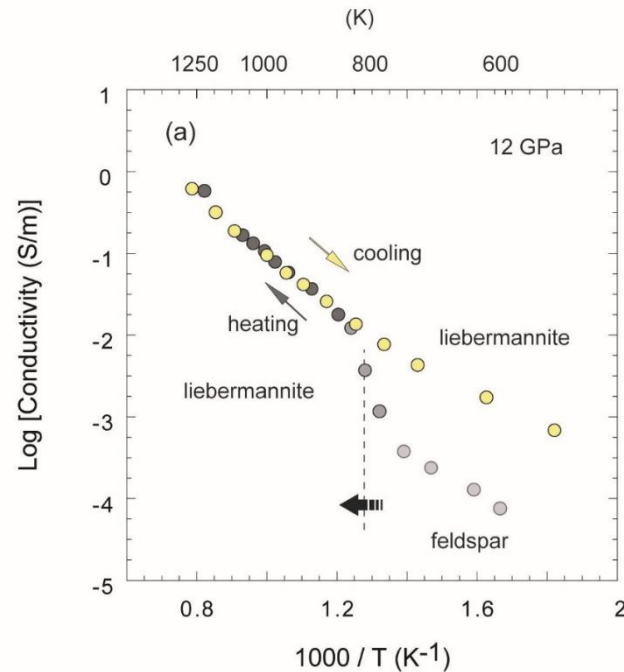
- K-hollandite (liebermannite):
  - principal phase in continental sediments (30 vol. %),
  - K-rich MORB
- Stable from the crust to core
  - K-feld – K-Hol (low P) – K-Hol (High P)
- tunnel structure – a repository for incompatible cations
- possible water carrier





# K-hollandite- A super ionic conductor

Superionic conduction : fast motion of  $K^+$  or  $H^+$  in the tunnel

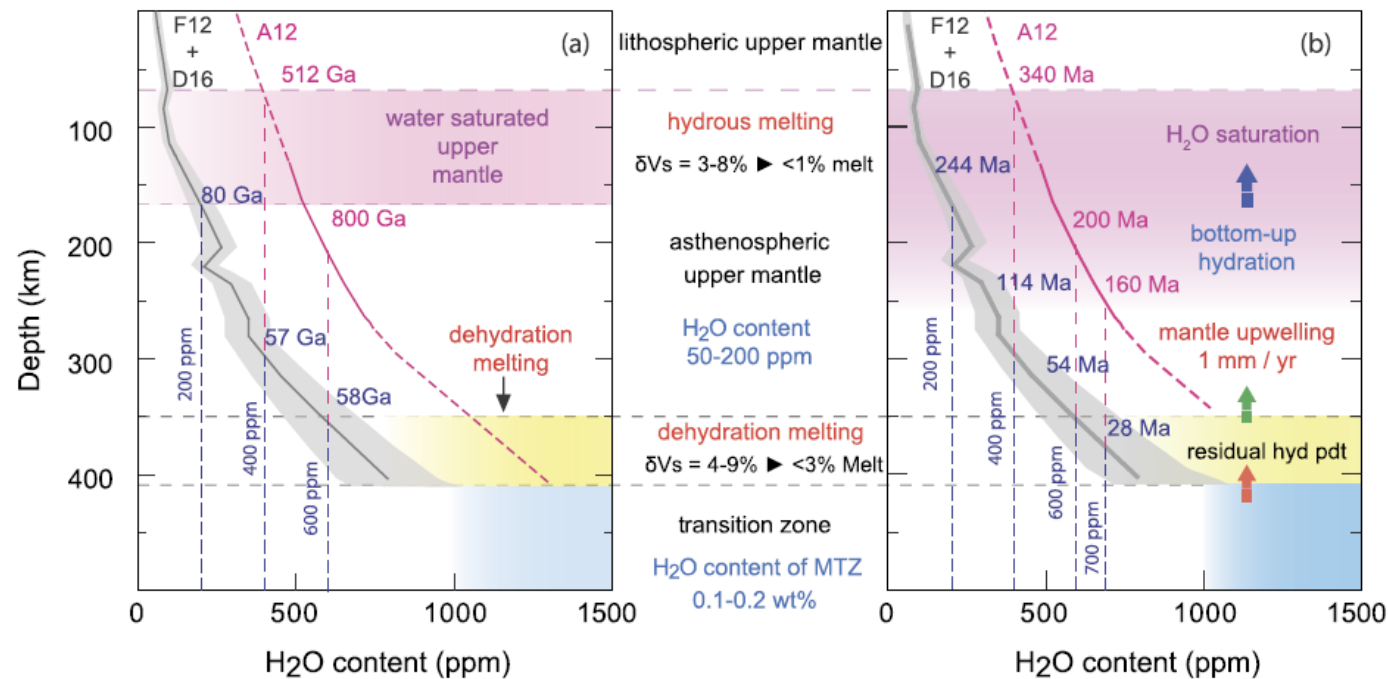


Tracer mineral for subduction pathways of continental sediments

# Water in the Earth's upper mantle

- The earth's upper mantle contains 50-200 et. ppm water

diffusion



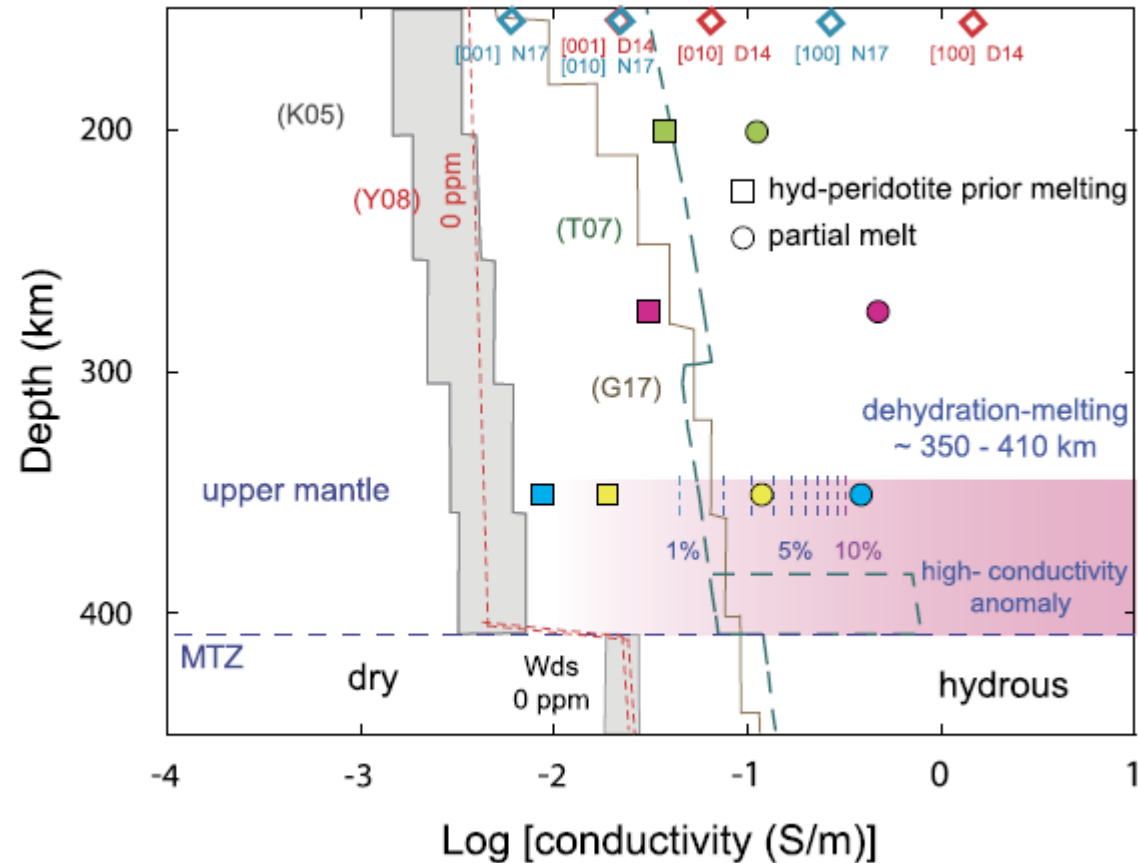
Mantle upwelling

$$\frac{c_x - c_0}{c_s - c_0} = 1 - \operatorname{erf}\left(\frac{x}{2\sqrt{Dt}}\right)$$

$c_s$  = Concentration of source  
 $c_x$  = Concentration at distance x  
 $c_0$  = Concentration of host



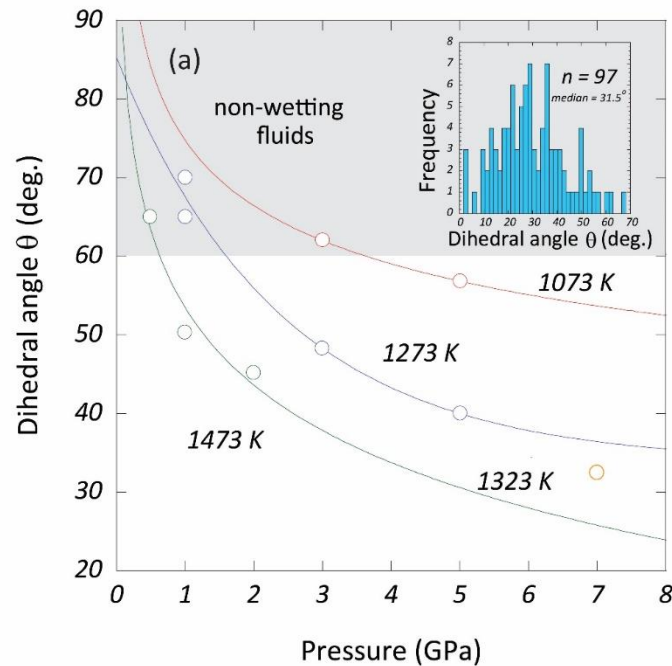
# Conductivity profile of the upper mantle



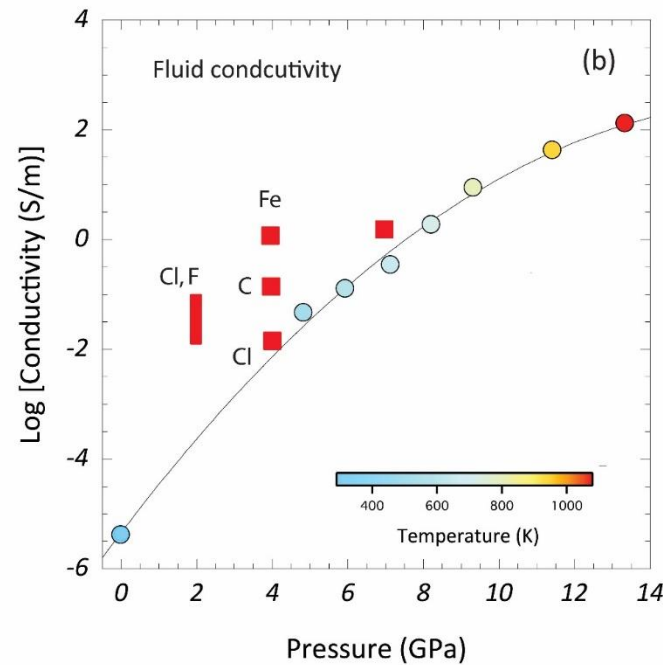
# Physical properties of fluids

- Dehydration of hydrous phases releases aqueous fluids into the mantle wedge
  - Affect the oxidation conditions
  - Mantle flow
  - carrier for the trace elements

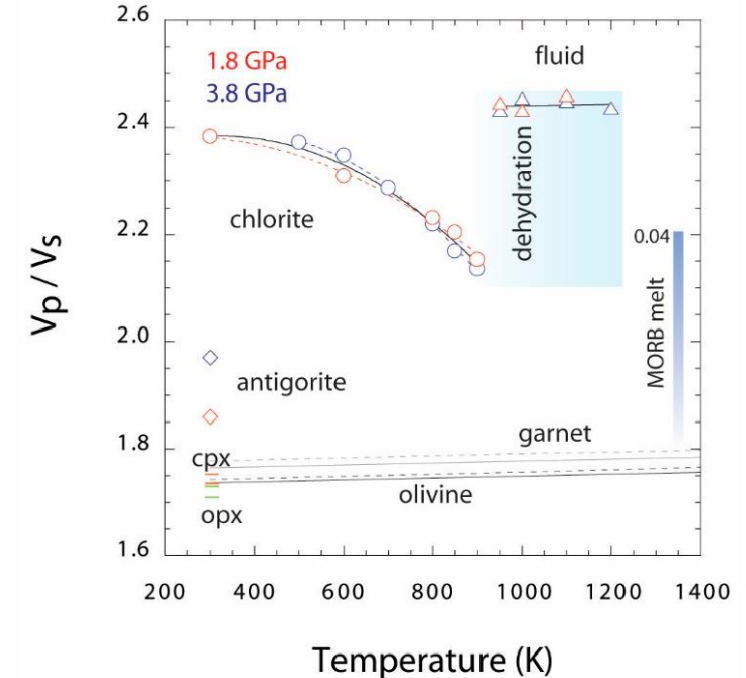
## Wetting properties



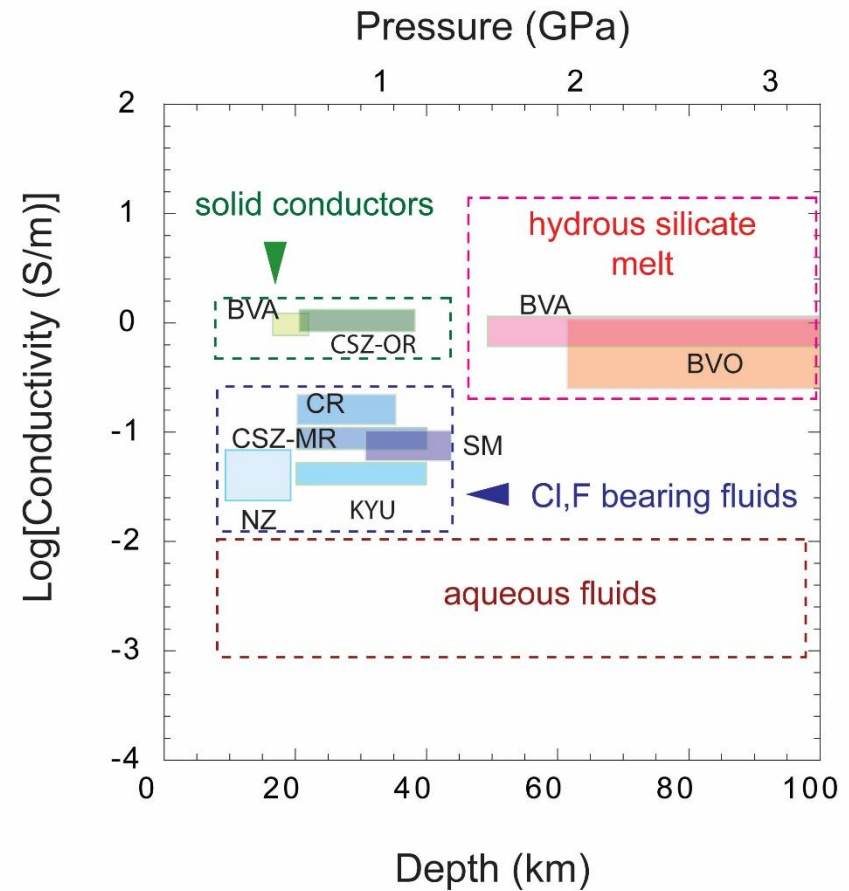
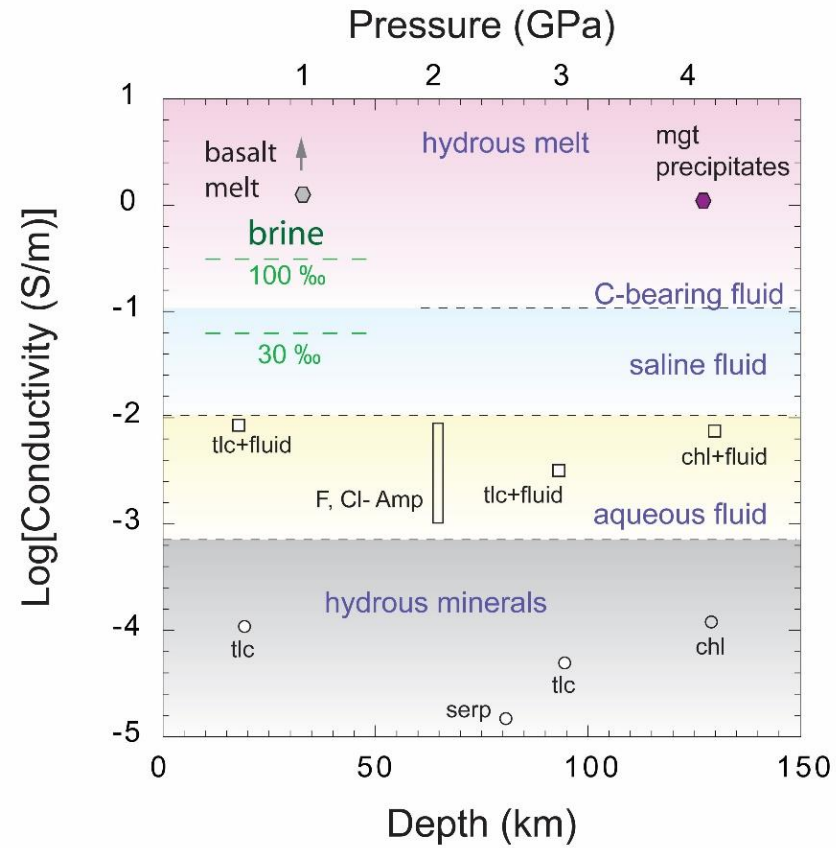
## Electrical properties



## Seismic properties

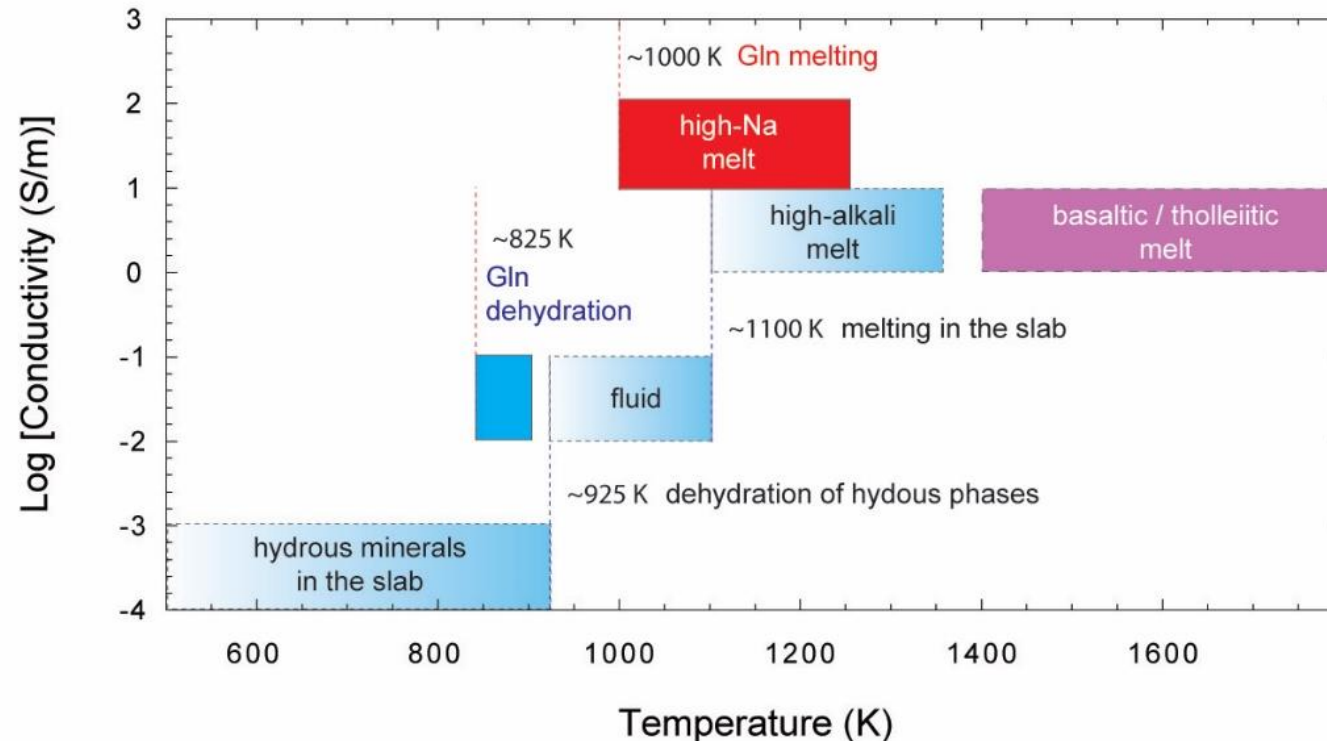


# Fluid in subduction zones

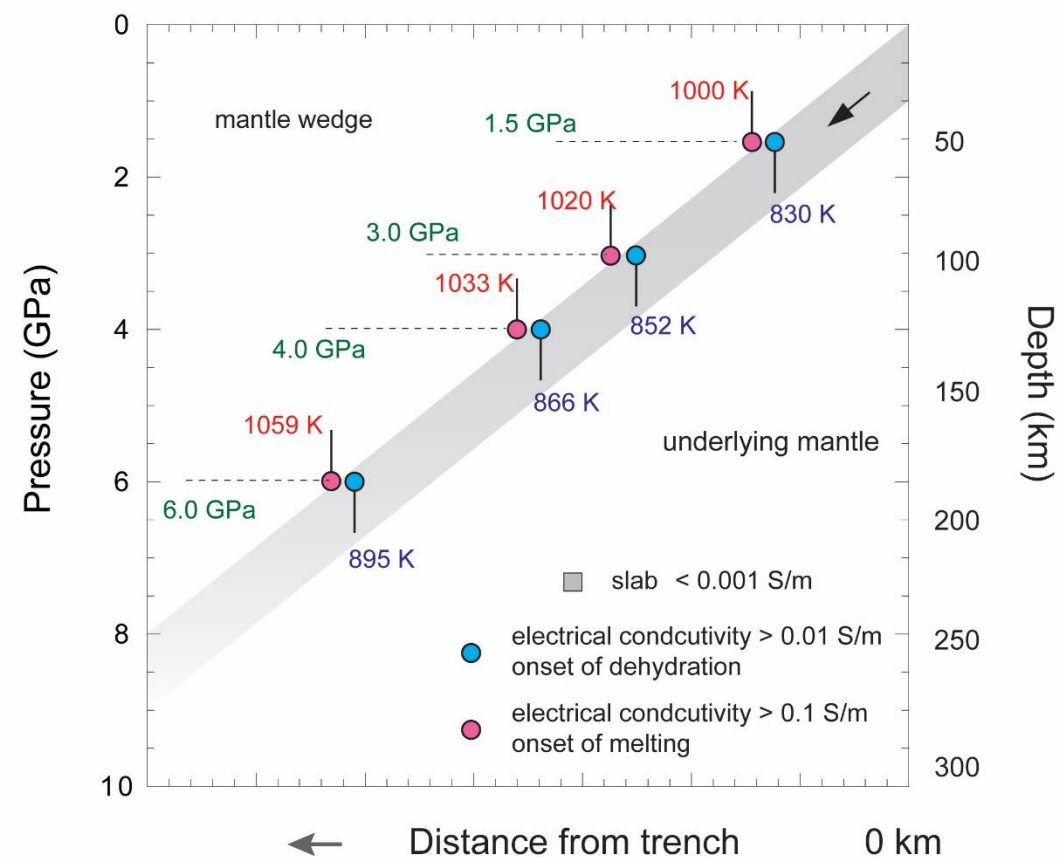


# Electrical conductivity as a geothermometer

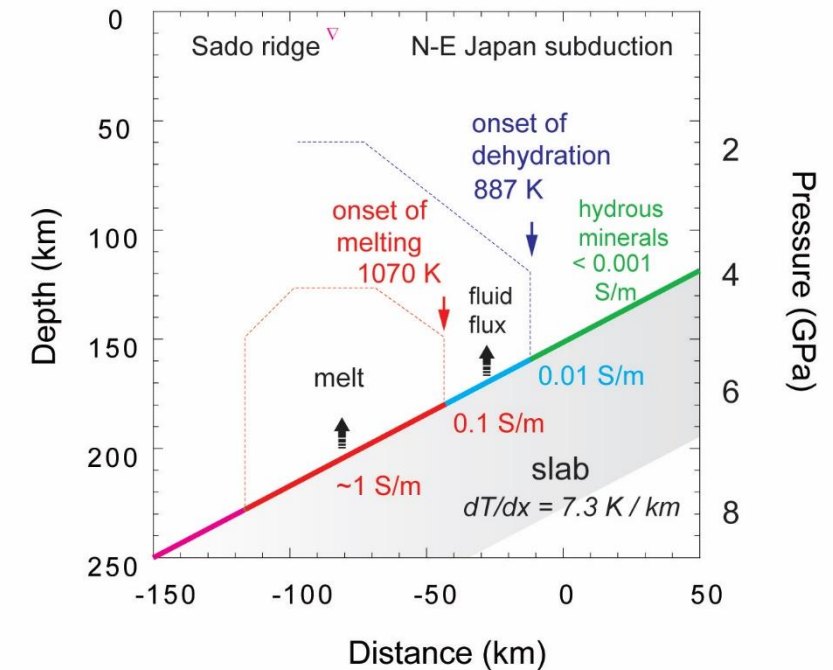
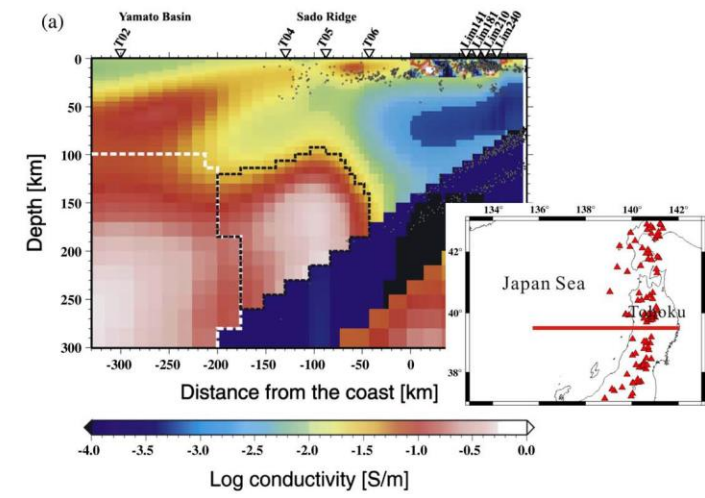
- Slab surface temperature – stability of minerals, melting, deeply subducting slab
- A numerical model based on finite element analyses (e.g. Syracuse et. al 2011)



# Application to NE-Japan subduction system



Manthilake et al. 2021

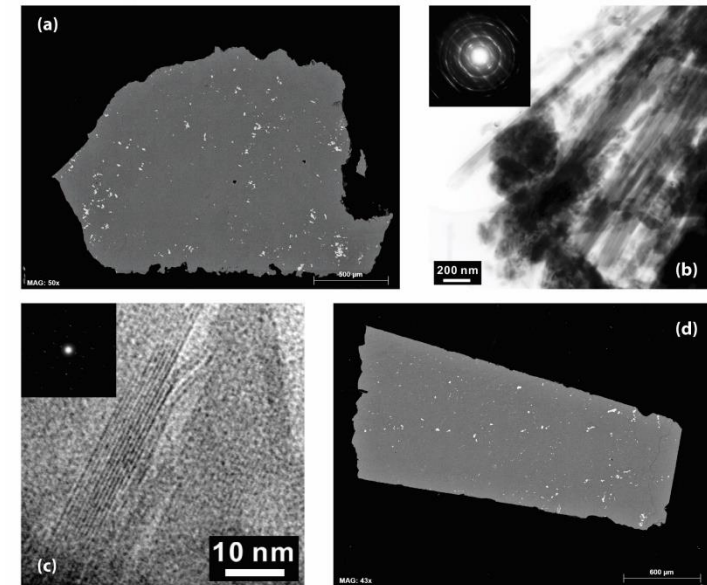


# Recycling of carbon through subduction

- The decarbonation and dehydration in subduction zones are intricately related.  
Most studies evaluate them separately
- Physical properties of CHO fluids
- carbon circulation in the mantle wedge
- the mechanisms of deep carbon subduction

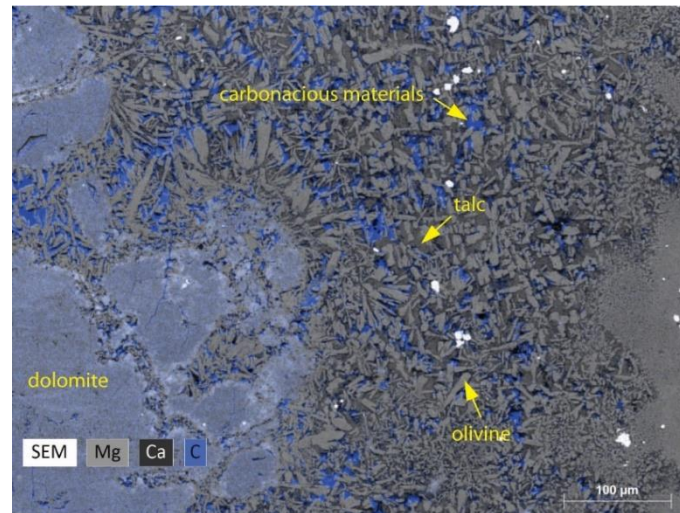
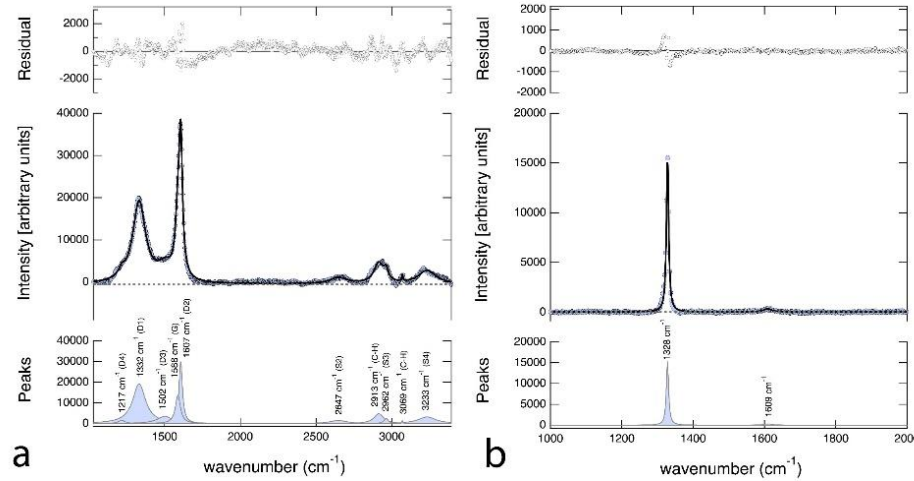
Starting materials:  
Carbonate-bearing chrysotile

HP-HT Experiments :  
4 GPa and up to 975 K





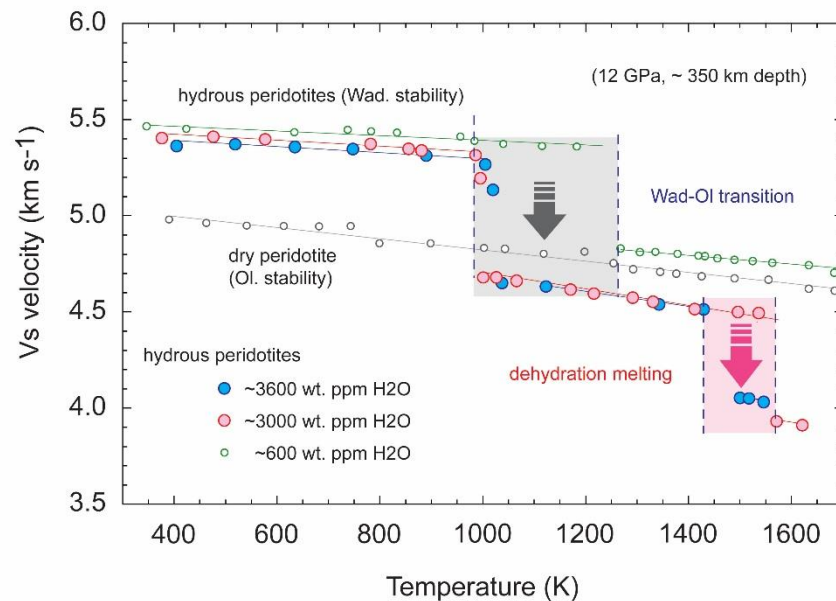
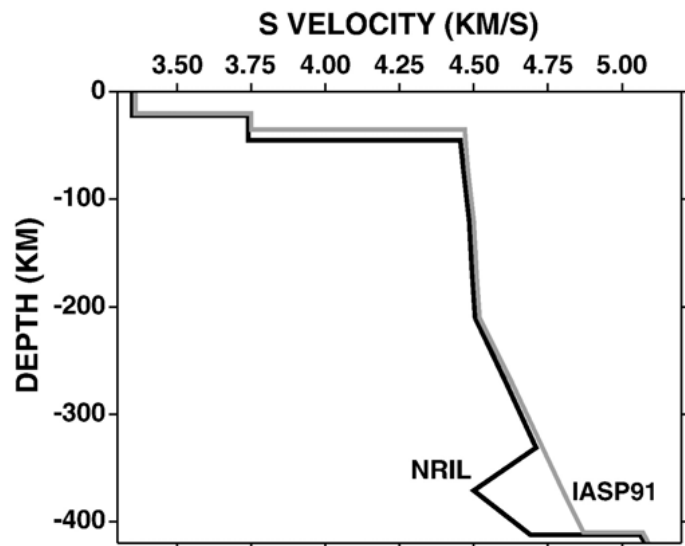
# Recycling of carbon through subduction



- The presence of carbonaceous materials suggests the following reactions  
 $\text{CO}_2 = \text{C} + \text{O}_2$ ,  
 $\text{CO}_2 + 2\text{H}_2\text{O} = \text{CH}_4 + 2\text{O}_2$
- The presence of diamonds at low T suggests the reactions:  
 $\text{CH}_4 = \text{C} + 2\text{H}_2$   
 $\text{CH}_4 + \text{O}_2 = \text{C} + 2\text{H}_2\text{O}$
- Supports the formation of subduction-related micro-diamonds in ultrahigh-pressure metamorphic terrains
- Crystallization of dolomite could be a key process responsible for transporting carbon into the deep Earth

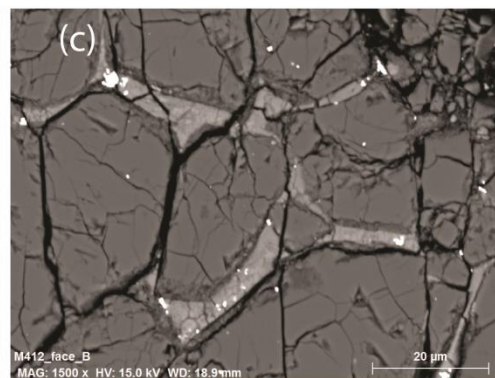
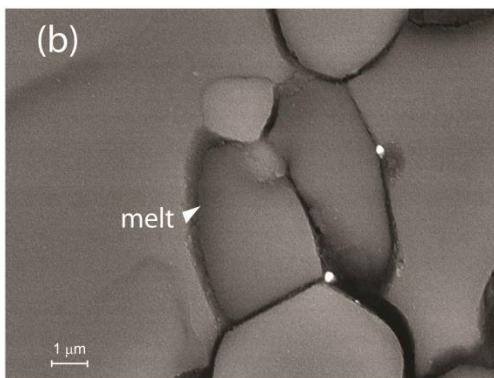
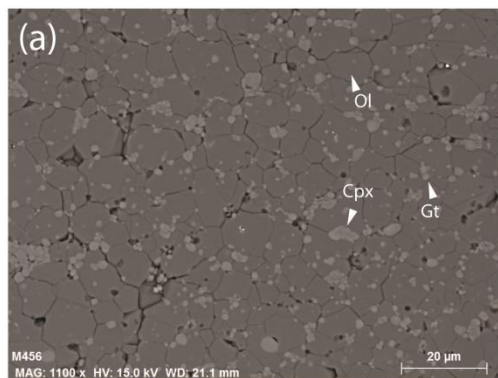
# Large-scale mantle structures

- Partial melt layer atop the mantle transition zone



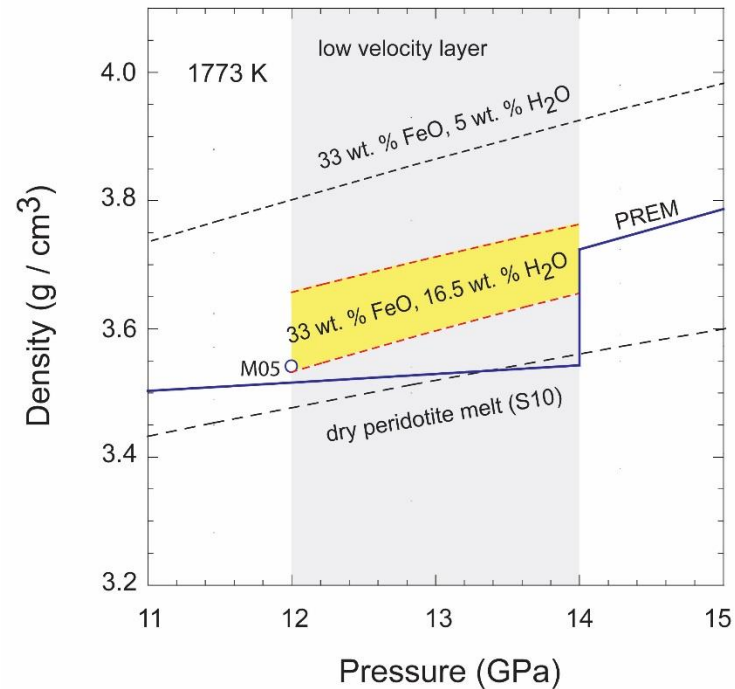
1<sup>st</sup> Vs drop - Wad to Ol transition

2<sup>nd</sup> Vs drop - melting



# Large-scale mantle structures

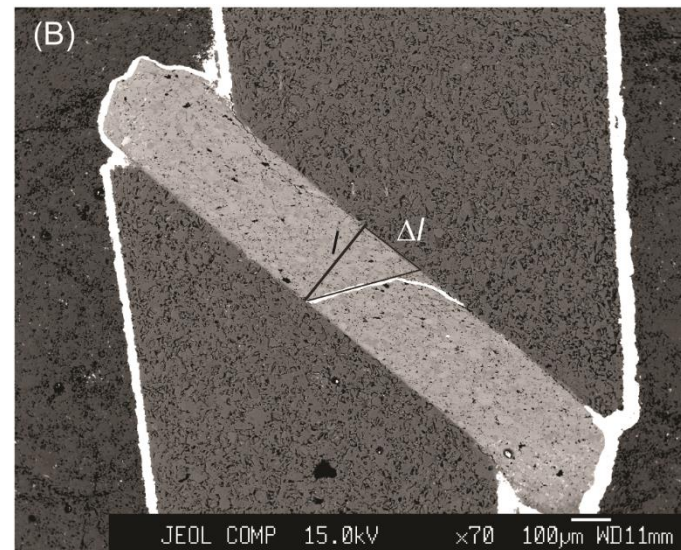
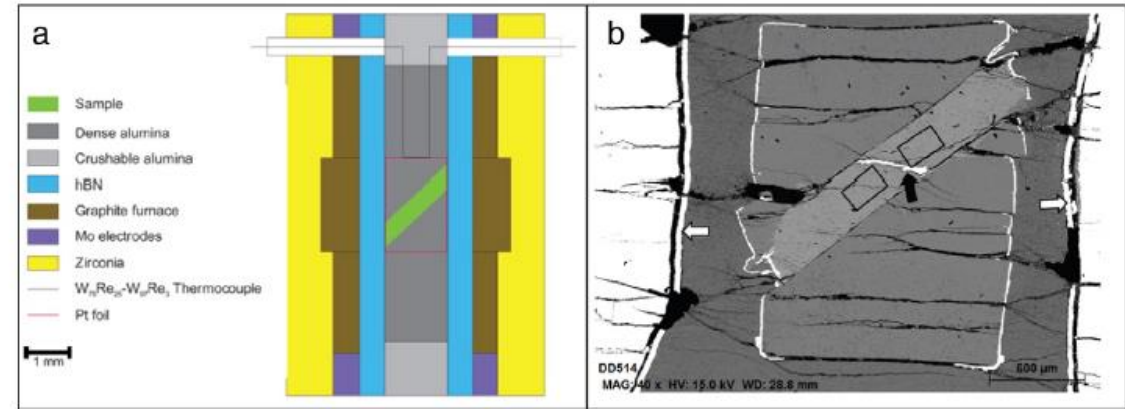
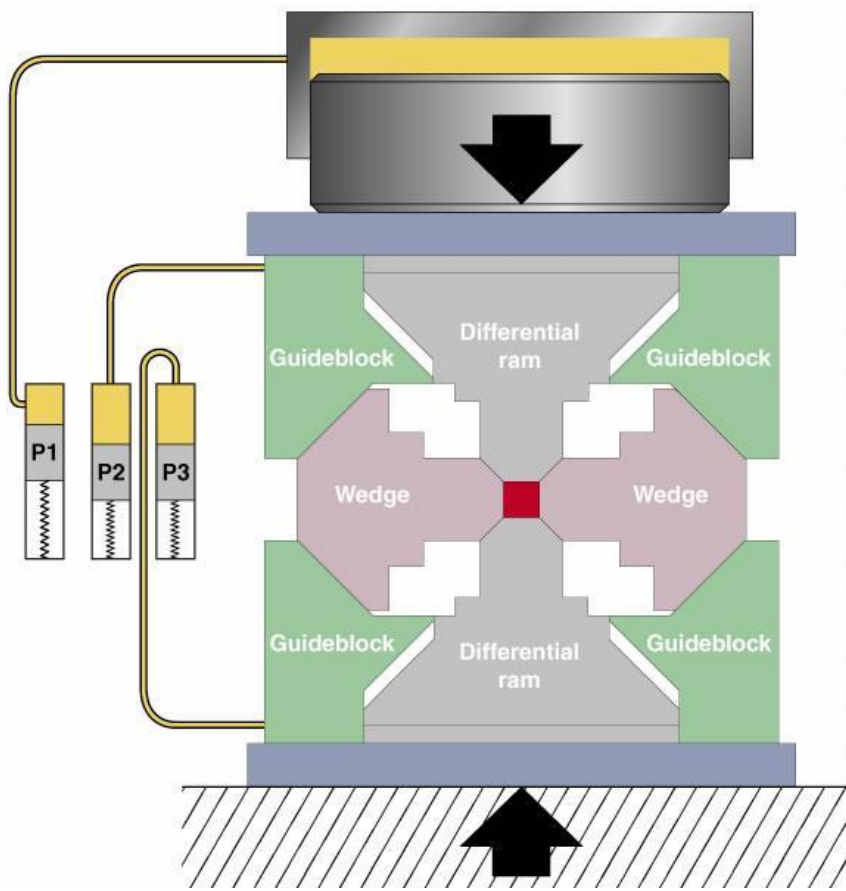
- Stable melt layer atop the mantle transition zone



- The density of melt 3.56- 3.74 g cm<sup>-1</sup>
- Neutrally buoyant at 350-410 km depth
- 4 % shear velocity drop: compatible with 0.7 - 1.0% melt
- MTZ water contents  $\sim 0.22 \pm 0.02$  wt. %
- MTZ is water undersaturated

# Large-scale mantle structures

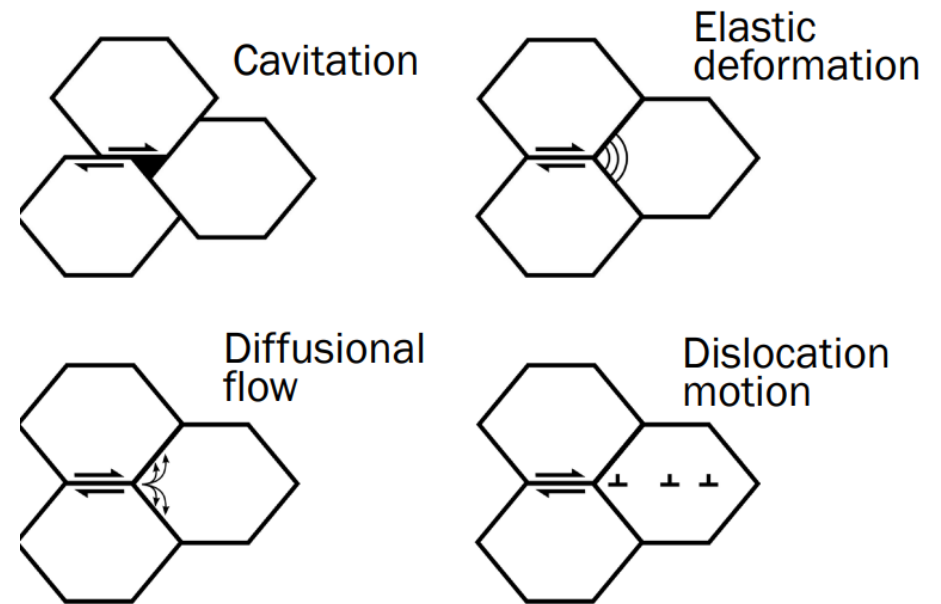
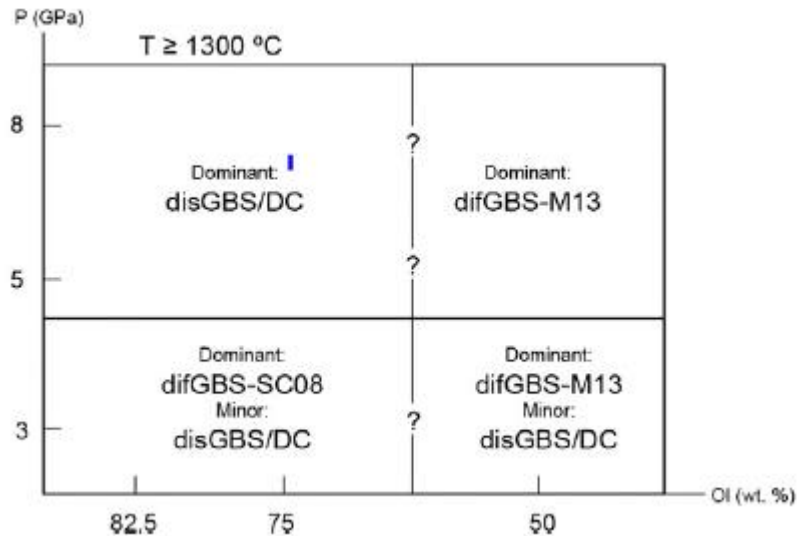
- Deformation of Ol+Opx mixtures in upper mantle conditions



Soustelle and Manthilake 2017

# Large-scale mantle structures

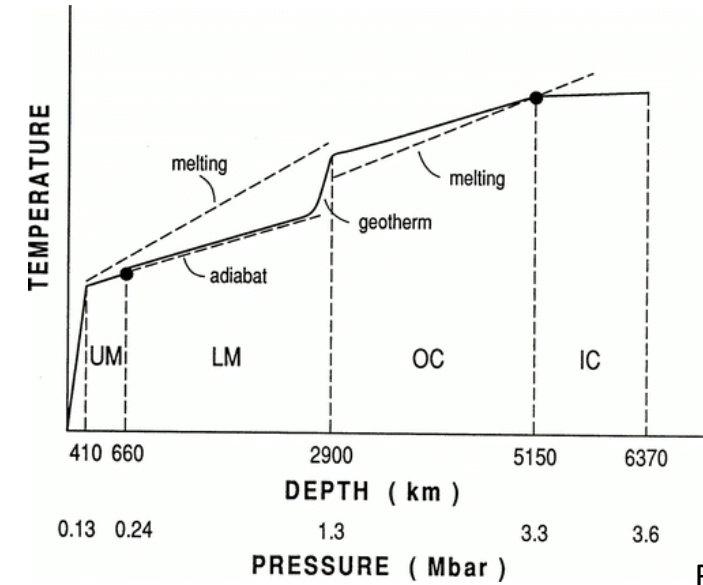
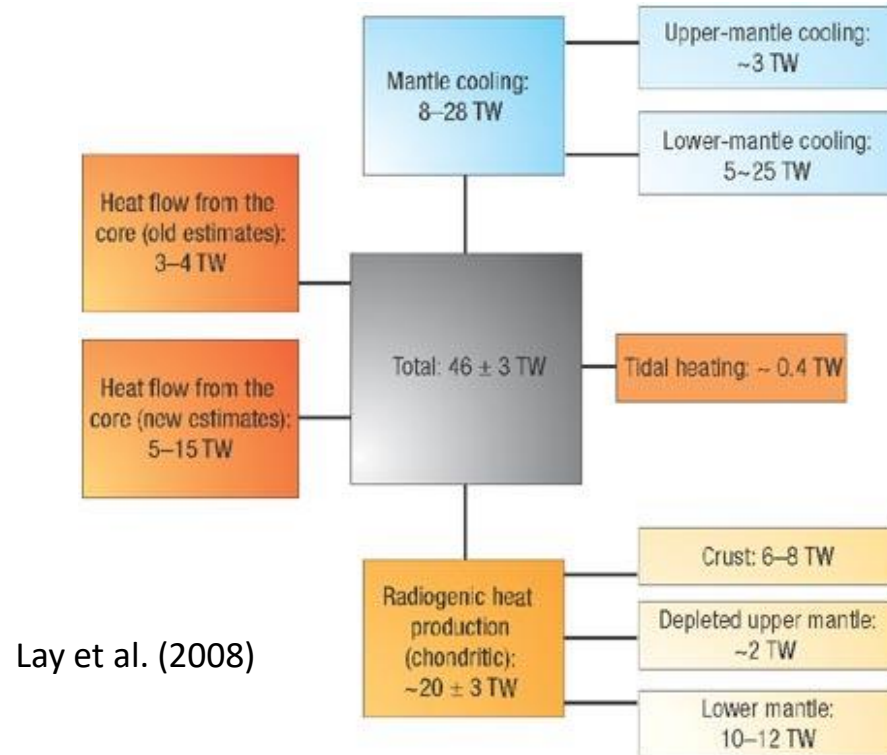
- Deformation of Ol+Opx mixtures



disGBS – dislocation-accommodated grain boundary sliding  
difGBS – diffusion-accommodated grain boundary sliding



# Planetary heat transport



Boehler (1996)

- Heat flows through TBL via conduction

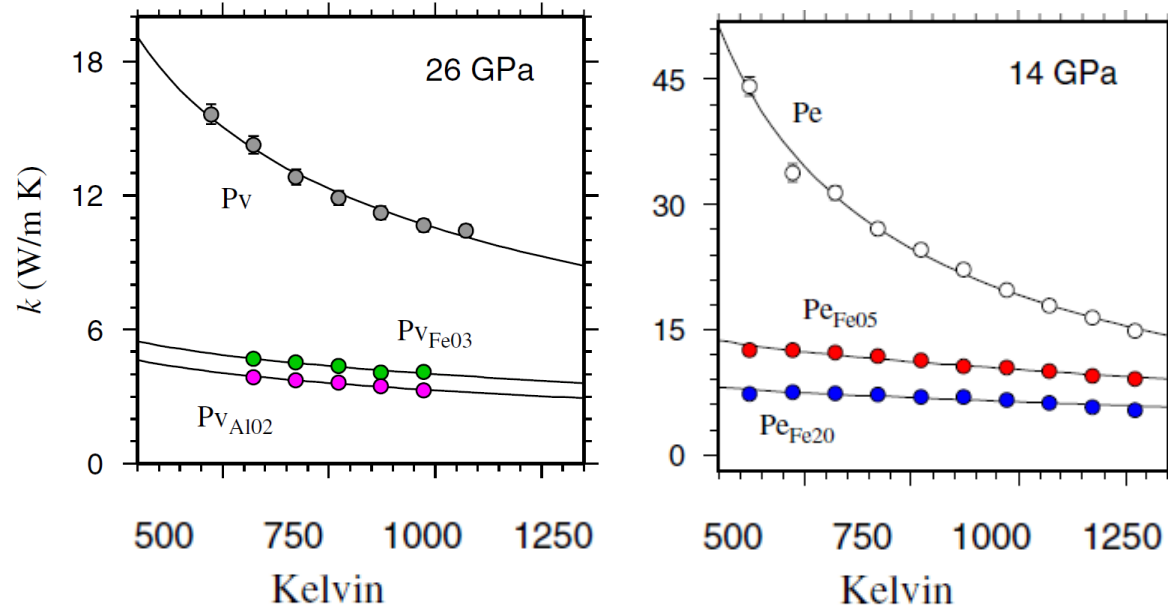
Key parameters

- Thermal conductivity
- Temperature
- Thickness of TBL

$$\frac{\Delta Q}{\Delta t} = -\kappa A \frac{\Delta T}{\Delta x}$$



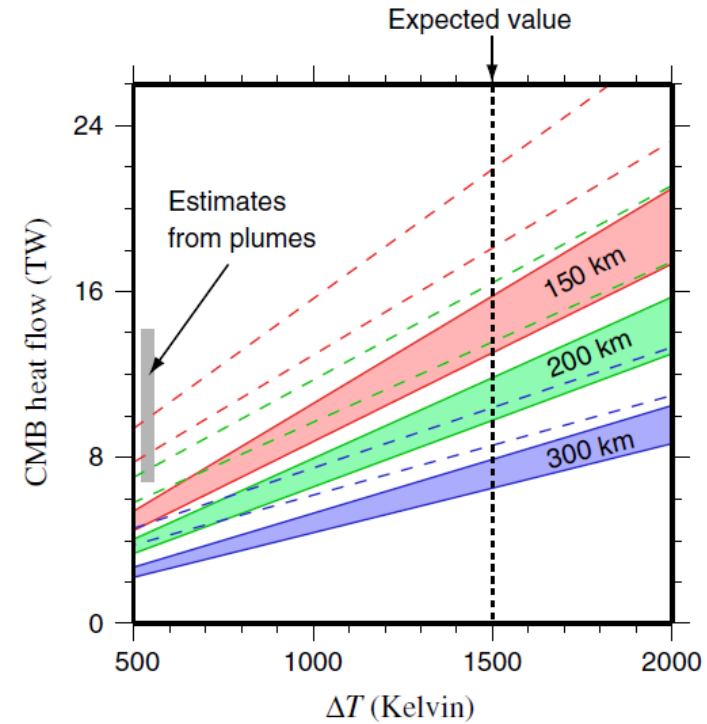
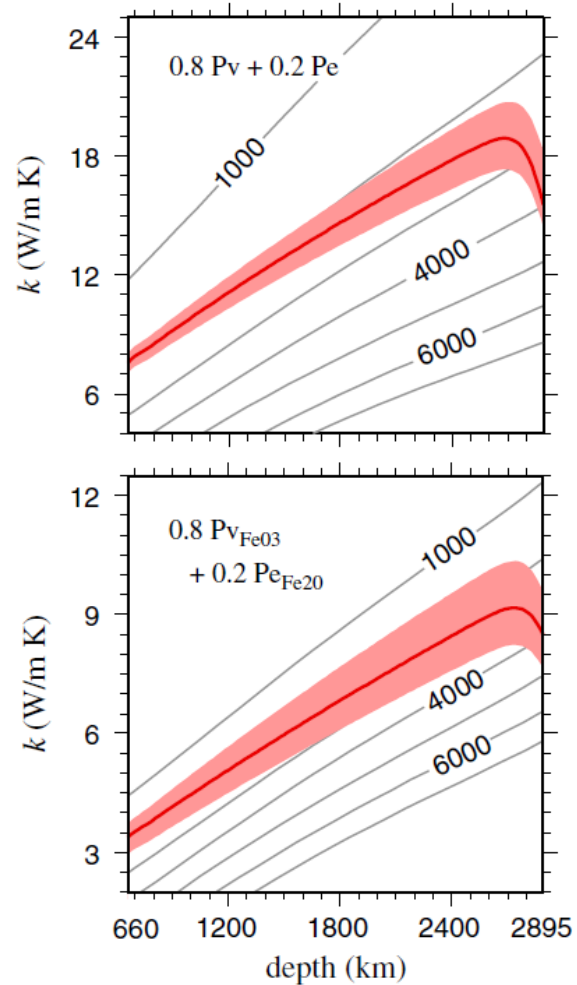
# Pressure and temperature dependence TC of brighmanite



$$\kappa = \kappa_{ref} \left( \frac{T_{ref}}{T} \right)^a \left( \frac{\rho}{\rho_{ref}} \right)^g$$

Effect of impurities – incorporation of Fe, Al decrease the  $\kappa$  by 50 %

# Core-mantle boundary heat flux



pure MgO-MgSiO<sub>3</sub> aggregate  
of 18.9-15.4 W/ m K

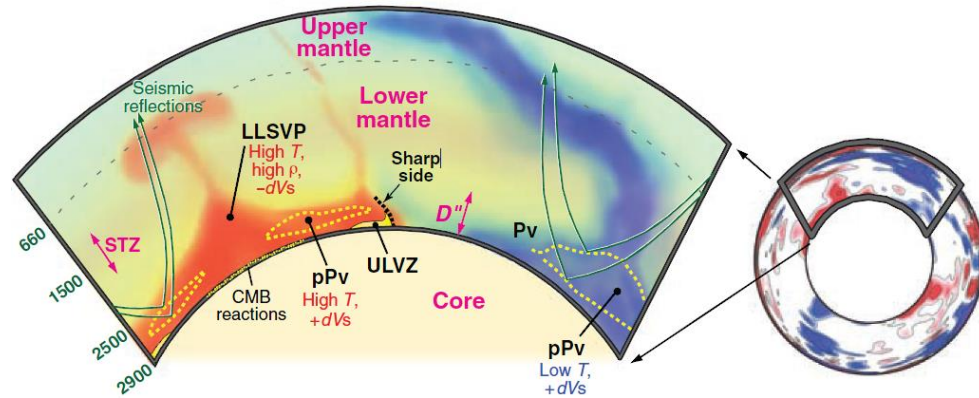
lattice thermal conductivity at TBL  
8-9 W/ m K

CMB heat flow  $11 \pm 1.4$  TW

Manthilake et al. 2011a

## Towards a comprehensive core-mantle boundary heat flow model

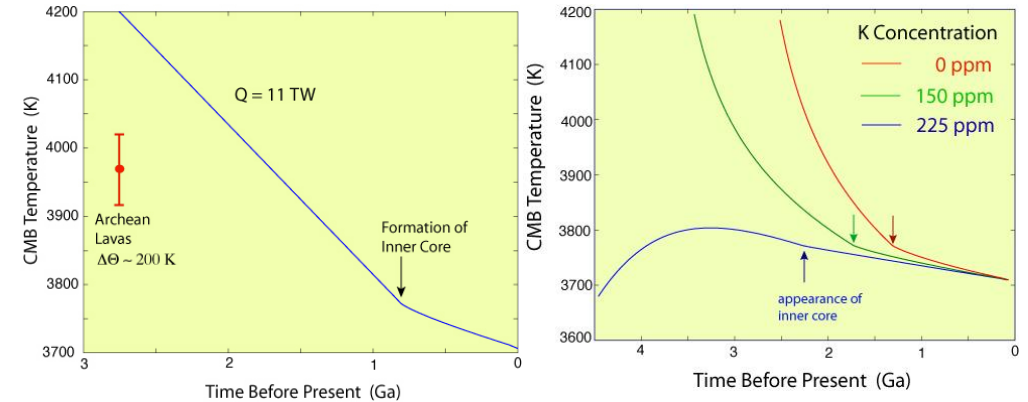
- Thermal conductivity of constituent phases along and across the TBL



Garnero and McNamara (2008)

According to seismology, CMB may consist of a mixture of

- Brg, Fp and PPv
- Subducted materials (MORB)
- Core reactant products,
- Partial melt, metallic alloys



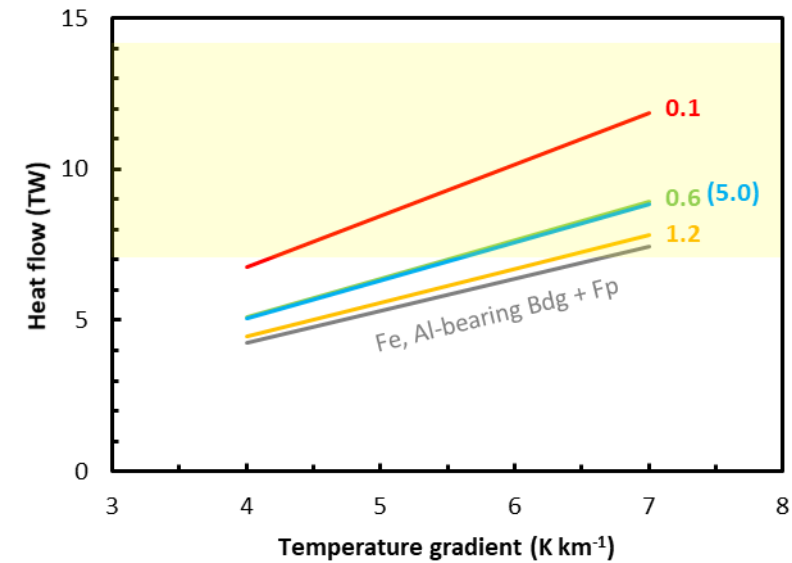
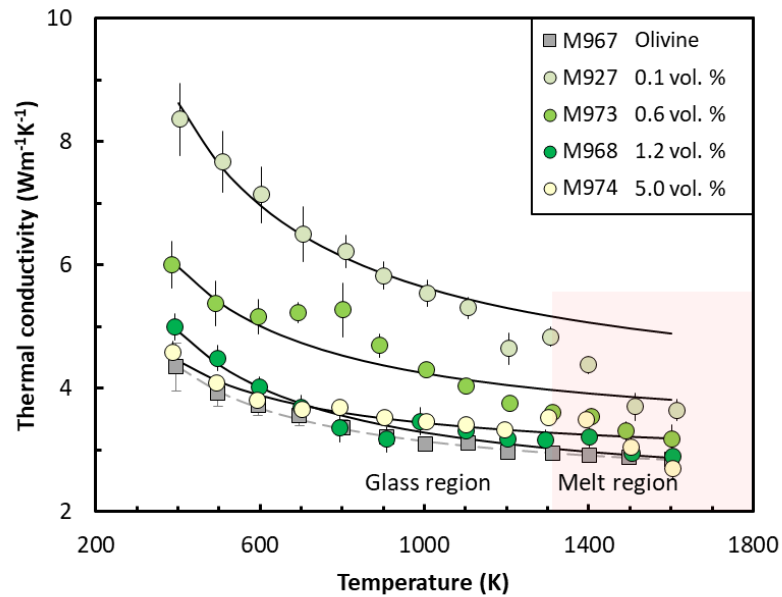
Buffett 1994

- Thermal instabilities at the CMB
  - Formation of mantle plumes
- Evolution of the TBL/heat flow with time
  - Sets the age of the inner core
- constrains the radiogenic isotopes budget in the core/CMB

# A comprehensive core-mantle boundary heat flow model

- The work in progress

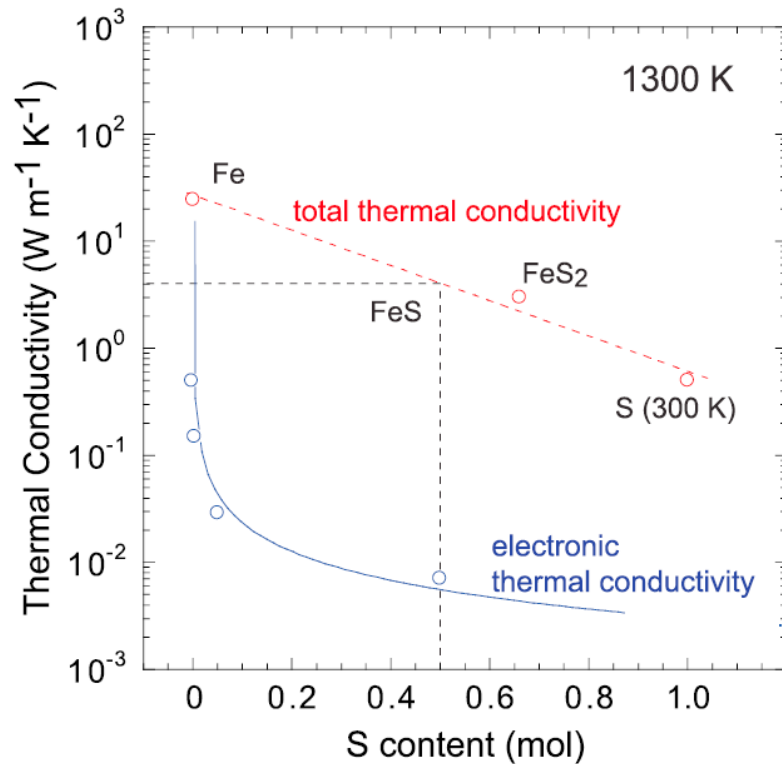
## Characterization of thermal conductivity of a melt bearing system



Zhao et al. in prep

# Thermal evolution of the planet Mercury

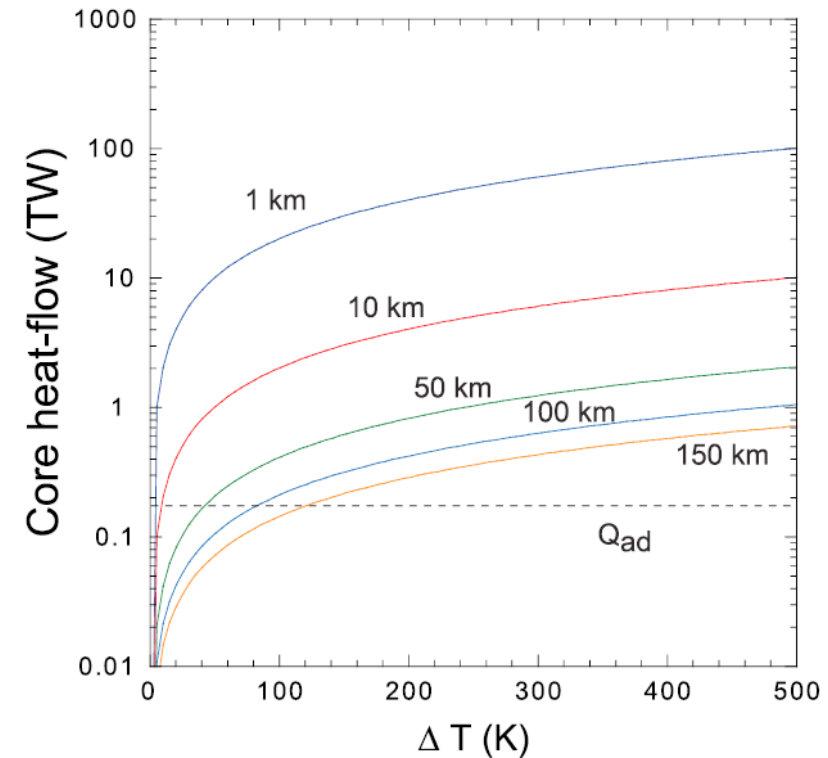
- Thermal conductivity of Fe-S alloys



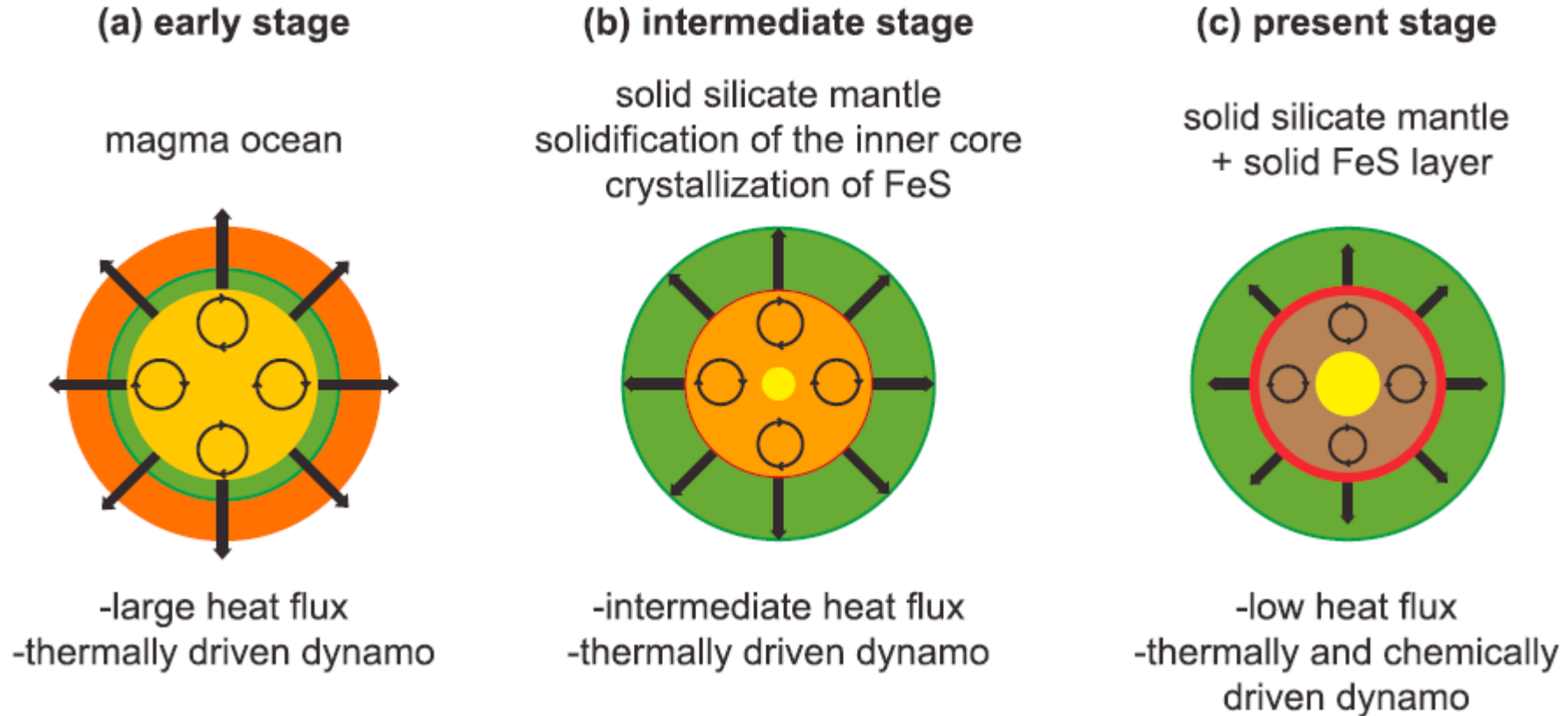
Wiedemann–Franz law

$$\frac{\kappa}{\sigma} = LT$$

Calculated heat flow at Mercury's CMB



# Thermal evolution of the planet Mercury



Heat flow is slightly higher than the  
adiabatic heat flow



## Les devises Shadok



IL VAUT MIEUX POMPER MÊME S'IL NE SE PASSE  
RIEN QUE RISQUER QU'IL SE PASSE QUELQUE CHOSE  
DE PIÈRE EN NE POMPANT PAS.