

# Wind Resource

## B. Wind Motion

### 1. Macroscopic Motions

#### • AT 30° – 60° NORTH/SOUTH LATITUDE

SINKING AIR FROM BOTH  
EQUATOR & HIGHER LATITUDES  
CAUSES HIGH PRESSURE

Polar Easterlies

Prevailing Westerlies

– AIR FLOWS NORTH & EAST

– WESTERLIES (CORIOLIS) Tradewinds

#### • AT POLES, COOL AIR SINKS

Doldrums

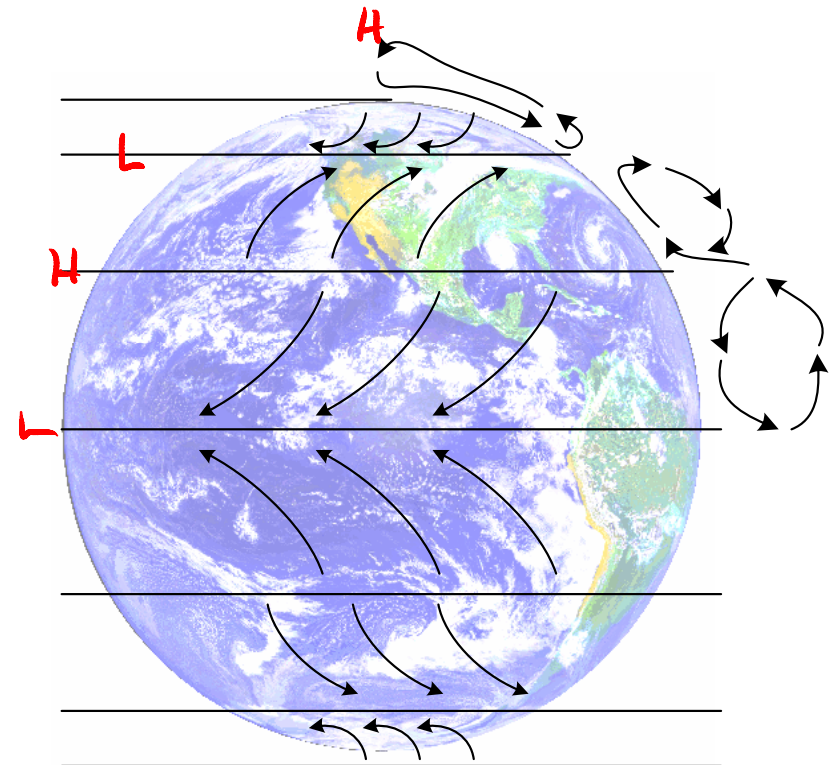
– AIR FLOW SOUTH & WEST Tradewinds

– POLAR EASTERLIES

Prevailing Westerlies

Polar Easterlies

#### • AT 60° NORTH, MEETING OF POLAR EASTERLIES & WESTERLIES



# Wind Resource

## B. Wind Motion

### 2. Microscopic Motions

SEVERAL EXAMPLES

LAND / SEA BREEZES

VALLEY / MOUNTAIN BREEZES

FOHN / SANTA ANA

THUNDERSTORMS

LAND SEA BREEZES

DAY TIME

WARM LAND & COOL OCEAN

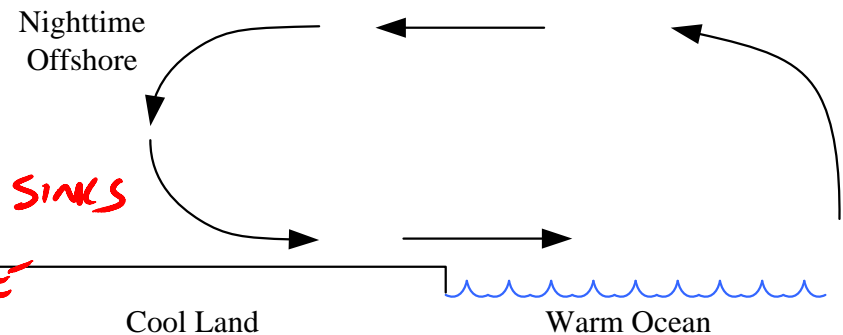
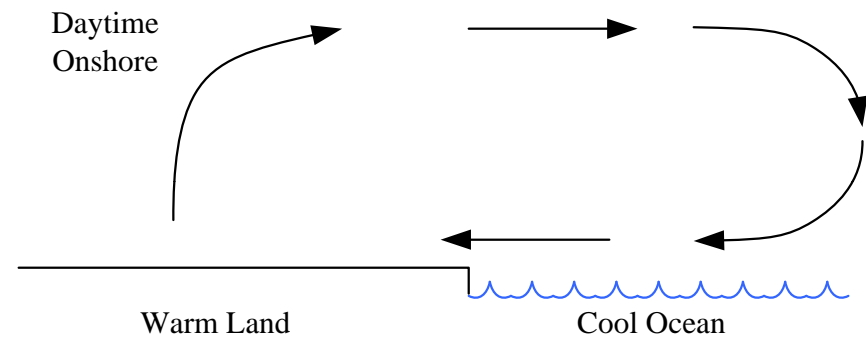
AIR RISES ON LAND, COOLS & SINKS

OVER OCEAN → ONSHORE BREEZE

NIGHT TIME

COOL LAND & WARM OCEAN

OFFSHORE BREEZE



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## B. Wind Motion

### 2. Microscopic Motions

MOUNTAIN/VALLEY BREEZE

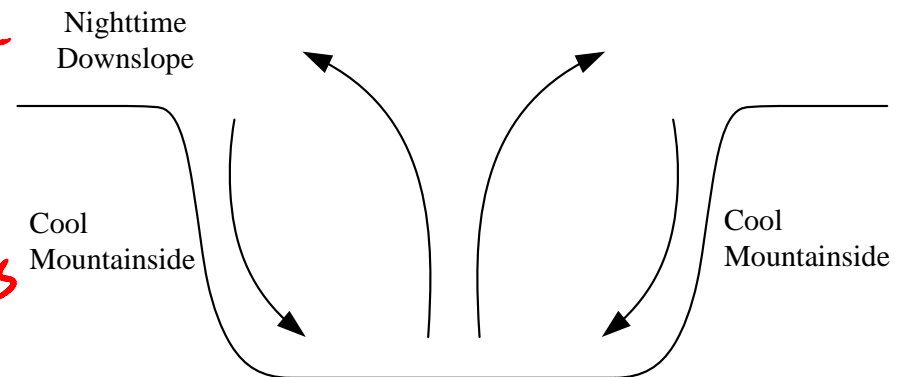
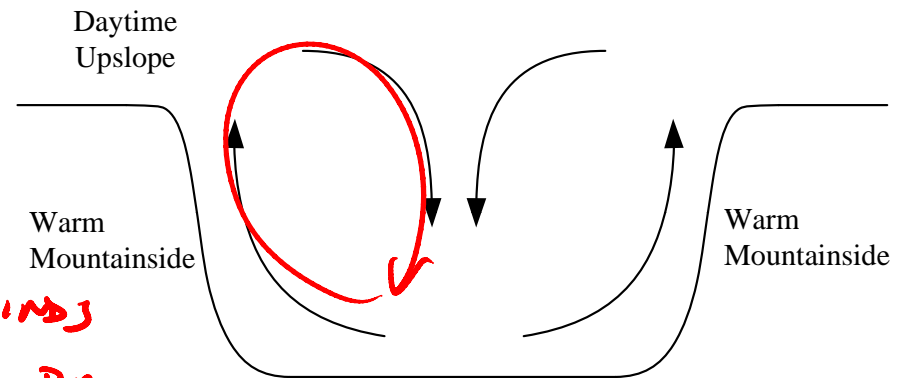
DAYTIME

VALLEY WALL WARMS

→ INDUCES UPSLOPE WINDS

→ AIR COOLS & FALLS BACK

TO VALLEY FLOOR



### 3. INBETWEEN - MESOSCALE

LARGE WEATHER SYSTEMS

WINTER STORMS

HURRICANES & CYCLONES

# Wind Resource

## B. Wind Motion

### 3. Time Variation

WINDS VARY WITH TIME FOR MANY REASONS

TIME-SCALES ASSOCIATED WITH VARIATIONS ARE VERY DIFFERENT

SHORT-TERM VARIATIONS

TURBULENCE & OTHER UNSTEADINESS

CAUSES WIND VARIATIONS WITH TIME SCALES FROM SECONDS TO ~hour

DAILY VARIATIONS

VARIATIONS DUE TO HEATING OF ATMOSPHERE & LAND/SEA DURING DAY & COOLING AT NIGHT

# Wind Resource

## B. Wind Motion

### 3. Time Variation

SEASONAL VARIATIONS

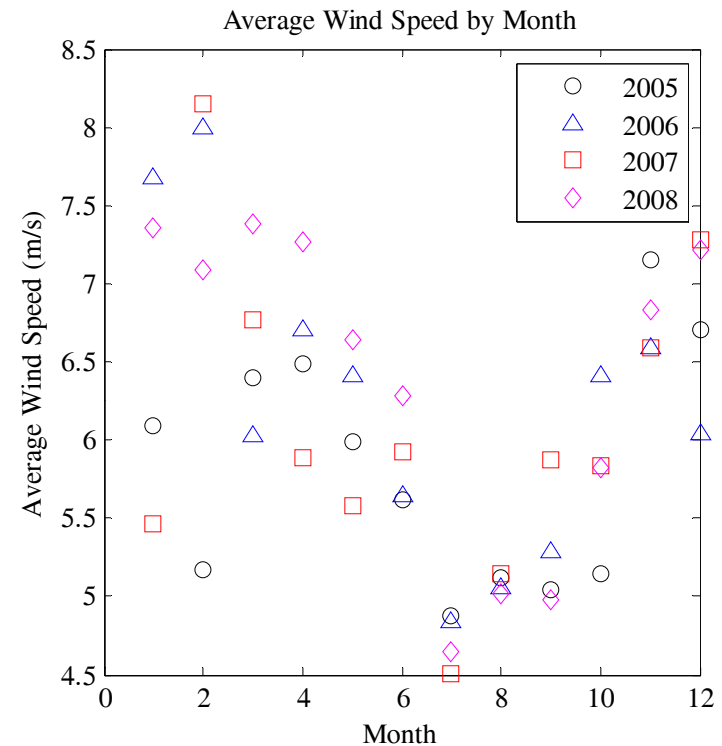
SIGNIFICANT CHANGES  
OVER COURSE OF YEAR  
ARE COMMON

EXACT CASE VALUES

> 1 YEAR VARIATIONS

CLIMATE

SHIFTS OCCUR



# Wind Resource

## B. Wind Motion

### 4. Mechanics

CONSIDER 4 FORCES ACTING ON A FLUID

PRESSURE FORCES

CORIOLIS FORCE (ARISES FROM EARTH'S ROTATION)

INERTIAL FORCES (DUE TO LARGE SCALE WIND PATTERNS)

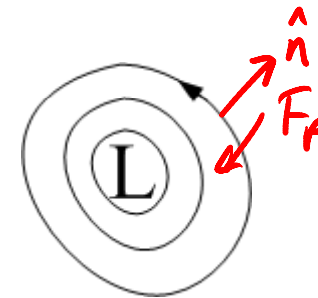
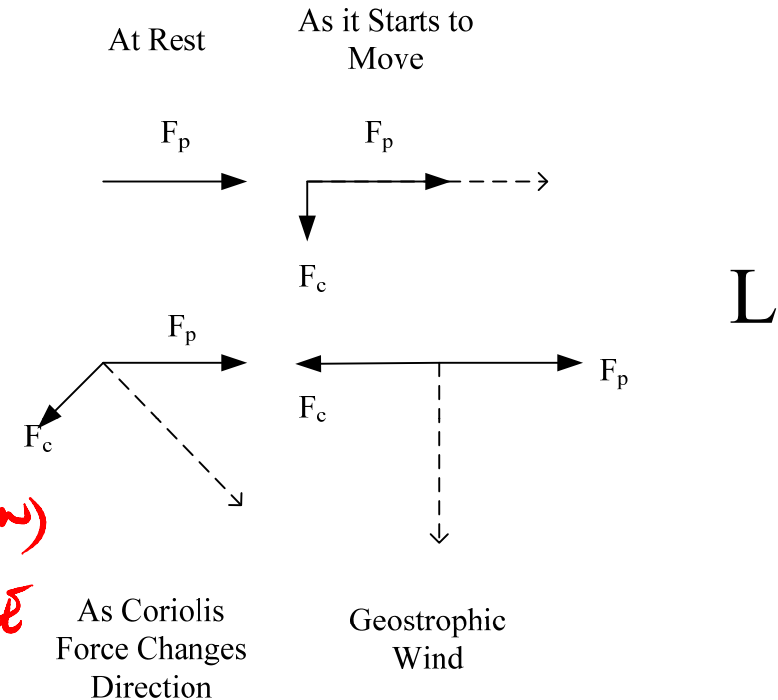
FRICTIONAL FORCES

PRESSURE FORCES

$$\frac{F_p}{m} = -\frac{1}{\rho} \frac{\partial p}{\partial n}$$

$\frac{\partial p}{\partial n} \equiv$  PRESSURE GRADIENT NORMAL TO THE ISO BARS

$\rho \equiv$  DENSITY



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## B. Wind Motion

### 4. Mechanics

CORIOUS FUELLES  
 $\frac{F_c}{m} = f U$  ← NOTE WIND FOR BE PRESENT TO DEVELOP CORIOUS F TO DEVELOP

$f \equiv$  CORIOUS PARAMETER

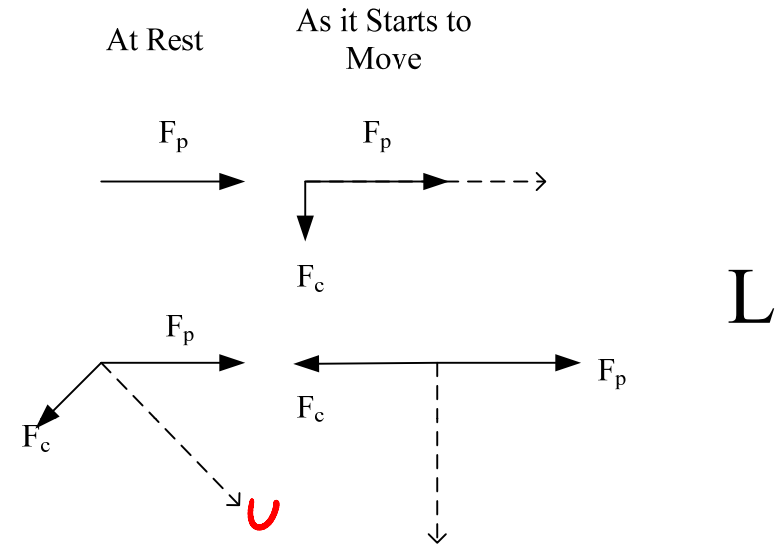
$f = 2\omega \sin(\phi)$   $\phi$  - LATITUDE

ACTS PERPENDICULARLY TO WIND DIRECTION  
 $\omega$  - ANGULAR ROTATION

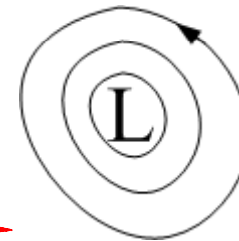
$$0 = \frac{F_p}{m} - \frac{F_c}{m}$$

$$0 = -\frac{1}{\tau} \frac{\partial p}{\partial n} - f U_g$$

$U_g$  IS GEOSTROPHIC WIND



As Coriolis Force Changes Direction



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## B. Wind Motion

### 4. Mechanics

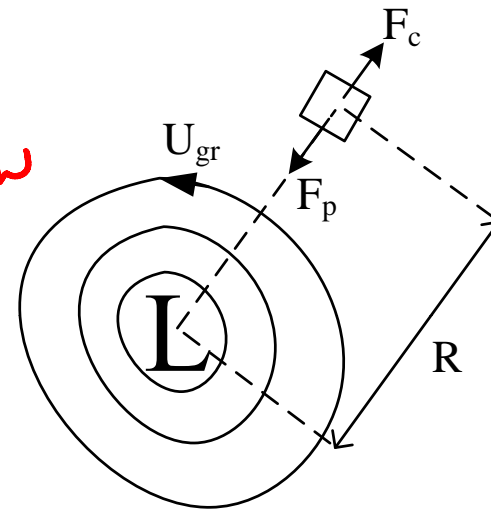
INERTIAL FORCES

DUE TO LARGE SCALE ROTATION  
PRESSURE & CORIOLIS WILL  
HAVE A SLIGHT DIFFERENCE

"INERTIAL  
FORCE"

$$\frac{U_{gr}^2}{R}$$

$$= \frac{F_p}{m} - \frac{F_c}{m}$$



$$\frac{U_{gr}^2}{R} = -\frac{1}{\rho} \frac{\partial p}{\partial n} - f U_{gr}$$

FROM BEFORE

$$0 = -\frac{1}{\rho} \frac{\partial p}{\partial n} - f U_g$$

COMBINES TO SHOW

$$\frac{U_{gr}^2}{fR} + U_{gr} = U_g$$

FRICTIONAL FORCES

IN THE LOWEST PART OF THE ATMOSPHERE, FRICTIONAL FORCES DEVELOP NO-SLIP } => CONFINED TO THE BL VISCOSITY