

Outline

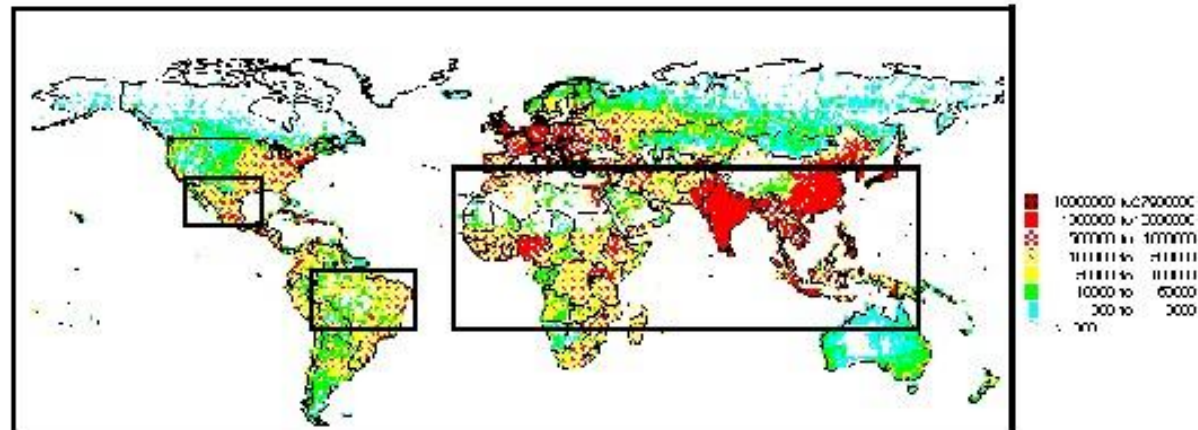
- ▶ Monsoon
- ▶ Radiative forcing
- ▶ Climate Feedbacks

Monsoon

Monsoon

- Monsoon is a climatological feature covering roughly half the tropics (1/4 of the global surface)
- Strictly, a system where the winds and precipitation reverses (summer rain, winter dry)
- Host 65% of the world's population
- Small changes in year-to-year climate can be catastrophic

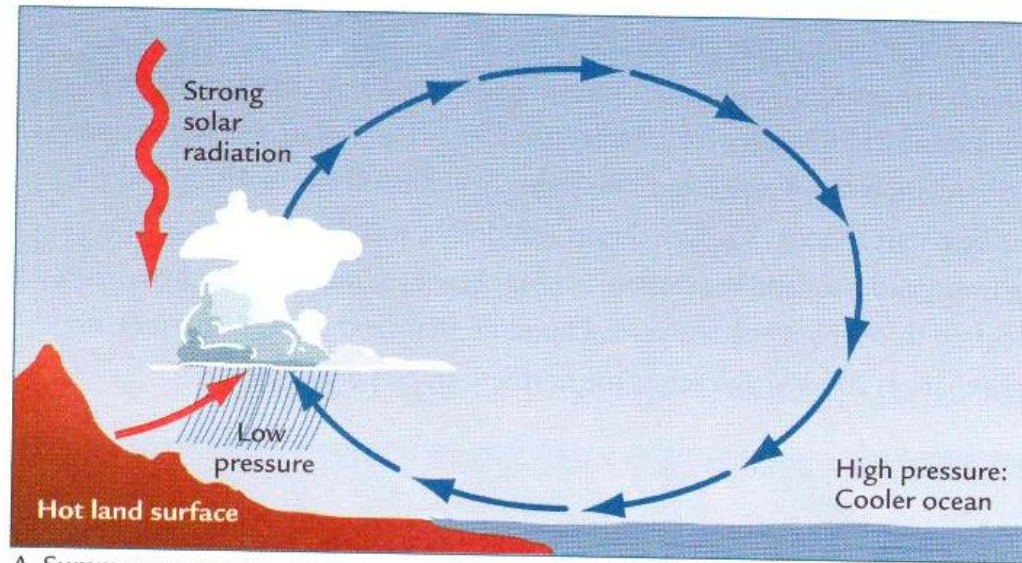
DEMOGRAPHICS OF THE MONSOON REGIONS



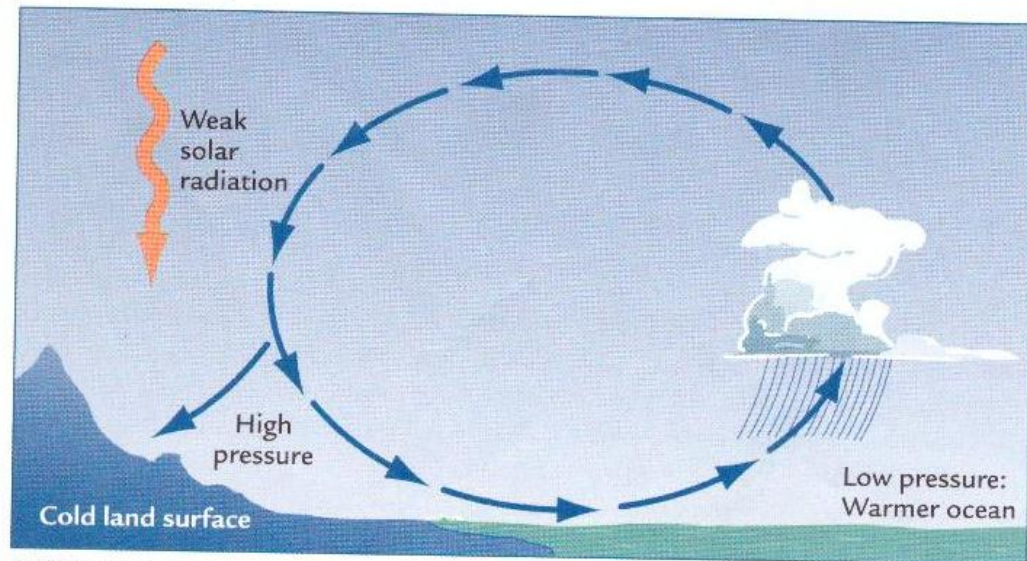
1950		1995		2025	
Rank	Country	Rank	Country	Rank	Country
1	China	1	China	1	India
2	India	2	India	2	China
3	USA	3	USA	3	Pakistan
4	Russia	4	Indonesia	4	USA
5	Japan	5	Brazil	5	Nigeria
6	Indonesia	6	Russia	6	Indonesia
7	Germany	7	Pakistan	7	Brazil
8	Brazil	8	Japan	8	Bangladesh
9	UK	9	Bangladesh	9	Ethiopia
10	Italy	10	Nigeria	10	Iran

Data from World Population Estimate (1995).

Seasonal monsoon circulations

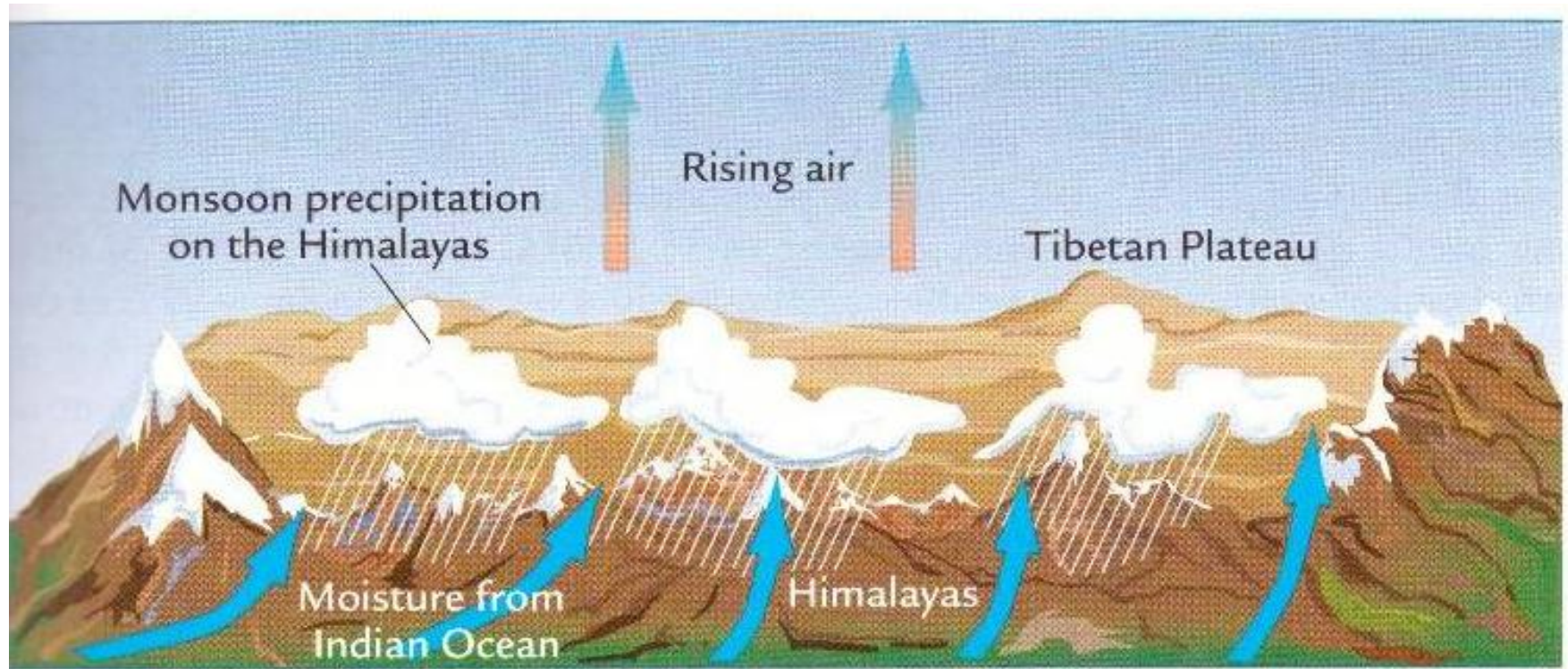


A Summer monsoon

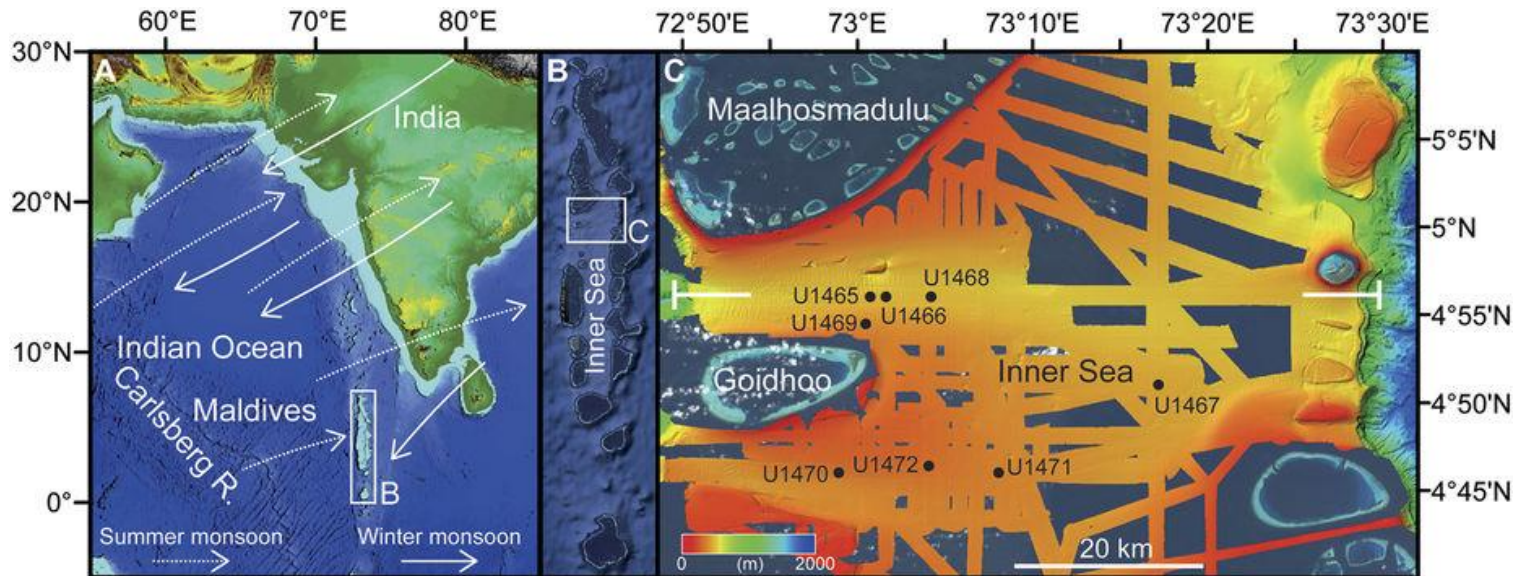


B Winter monsoon

Seasonal monsoon circulations



The abrupt onset of the modern South Asian Monsoon winds



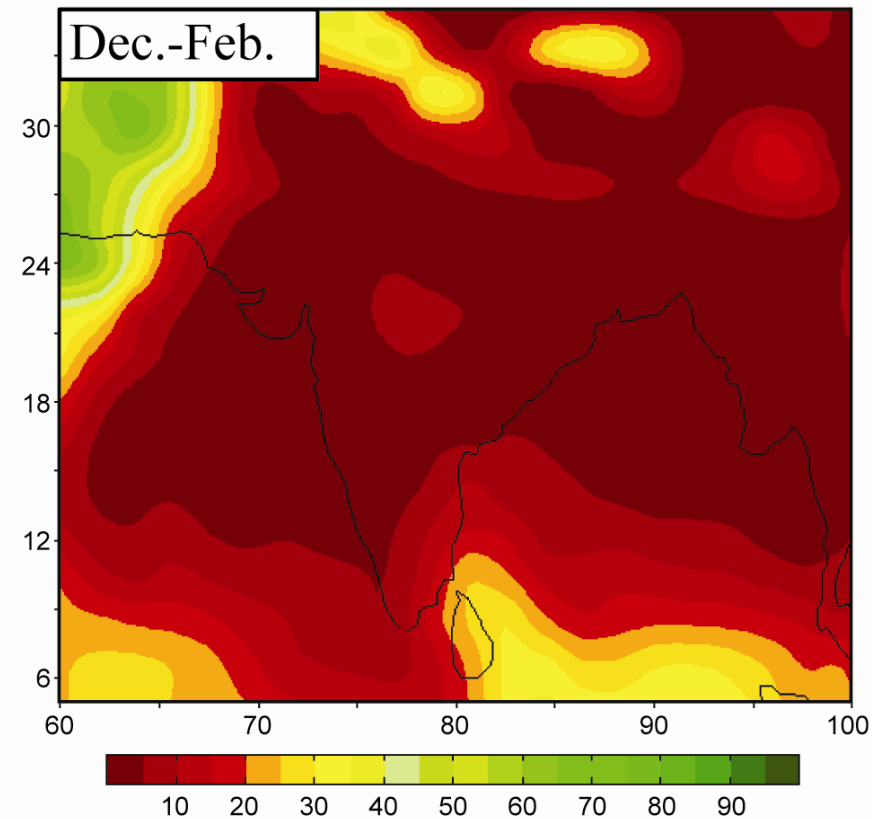
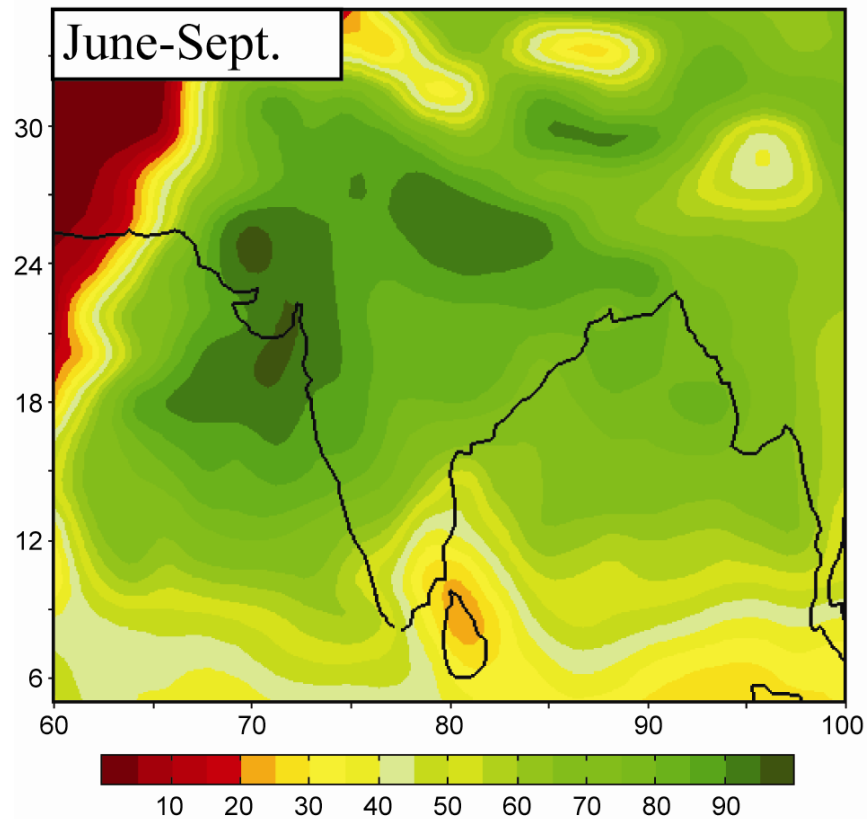
Dating the deposits of SAM wind-driven currents in IODP cores from the Maldives yields an age of 12.9 Ma indicating an abrupt Monsoon onset

The Winter Monsoon

- During autumn and winter, air above land cools faster than over the water.
- The winds are reversed from the summer monsoon flow—at the surface from land to sea
- Winter monsoon is a dry season

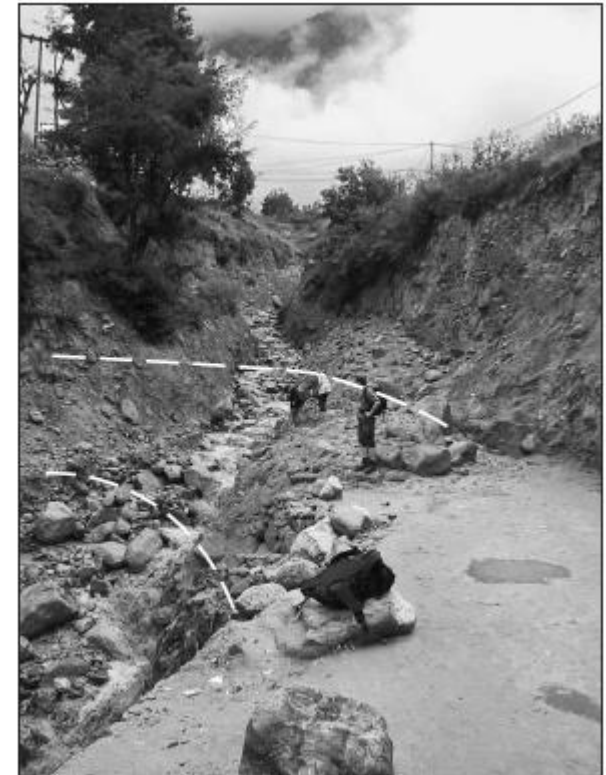
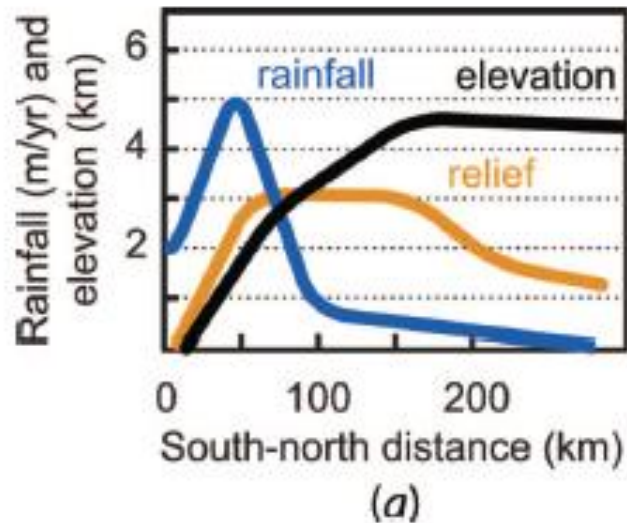
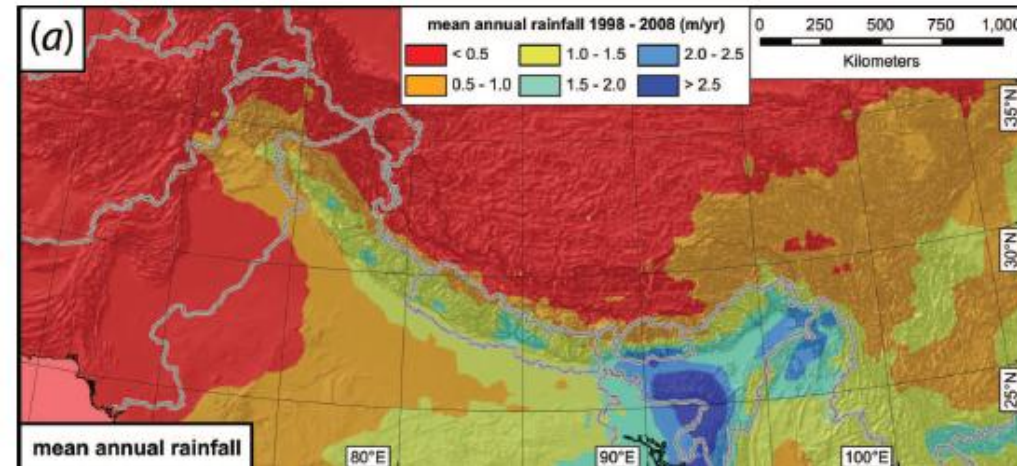
Modern Precipitation

Rainfall 1958-2007 (% of annual precipitation)

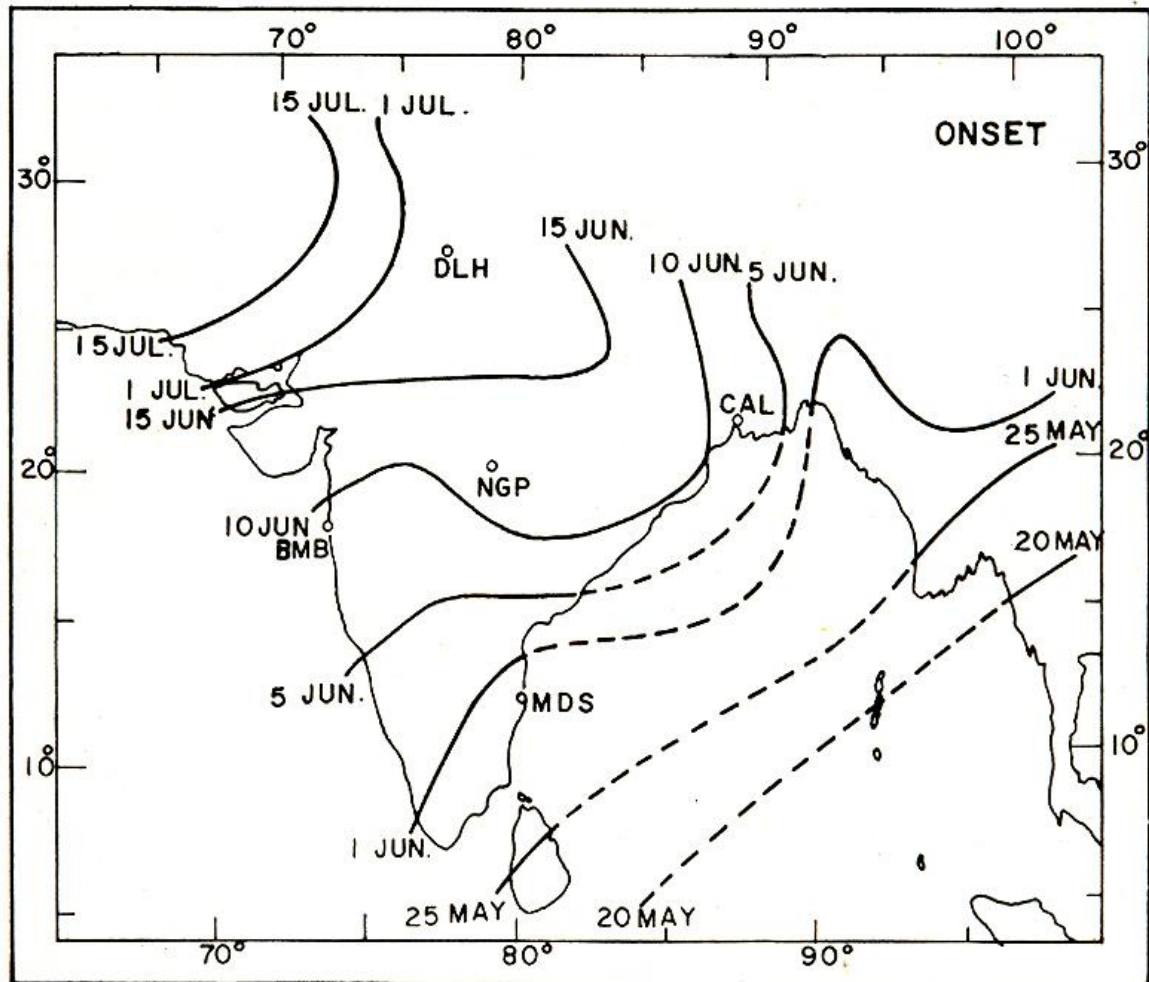


Modern Precipitation

Abnormal monsoon years



Monsoon-onset phase



Monsoon precipitation

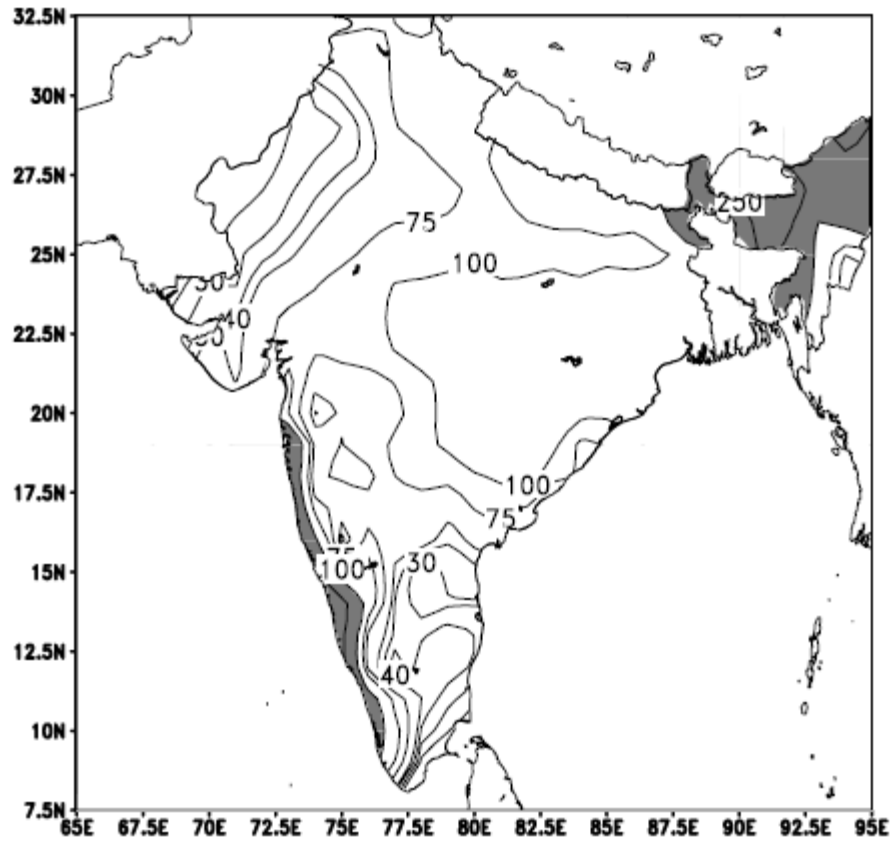
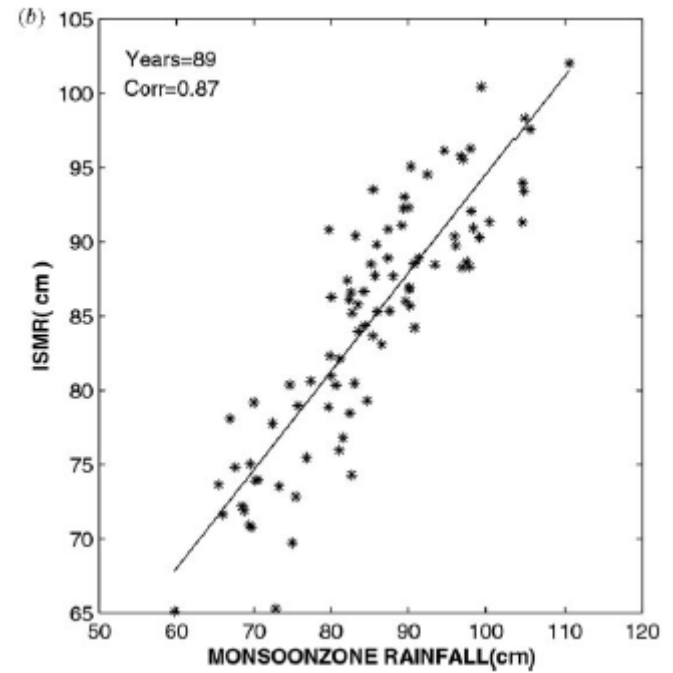
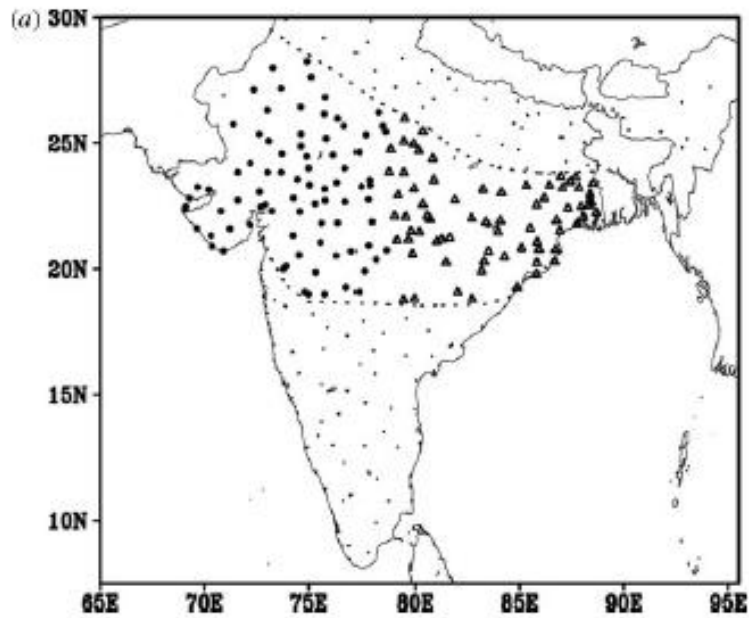


Figure 2 Mean June–September rainfall over the Indian region south of 30° N.

Core monsoon zone

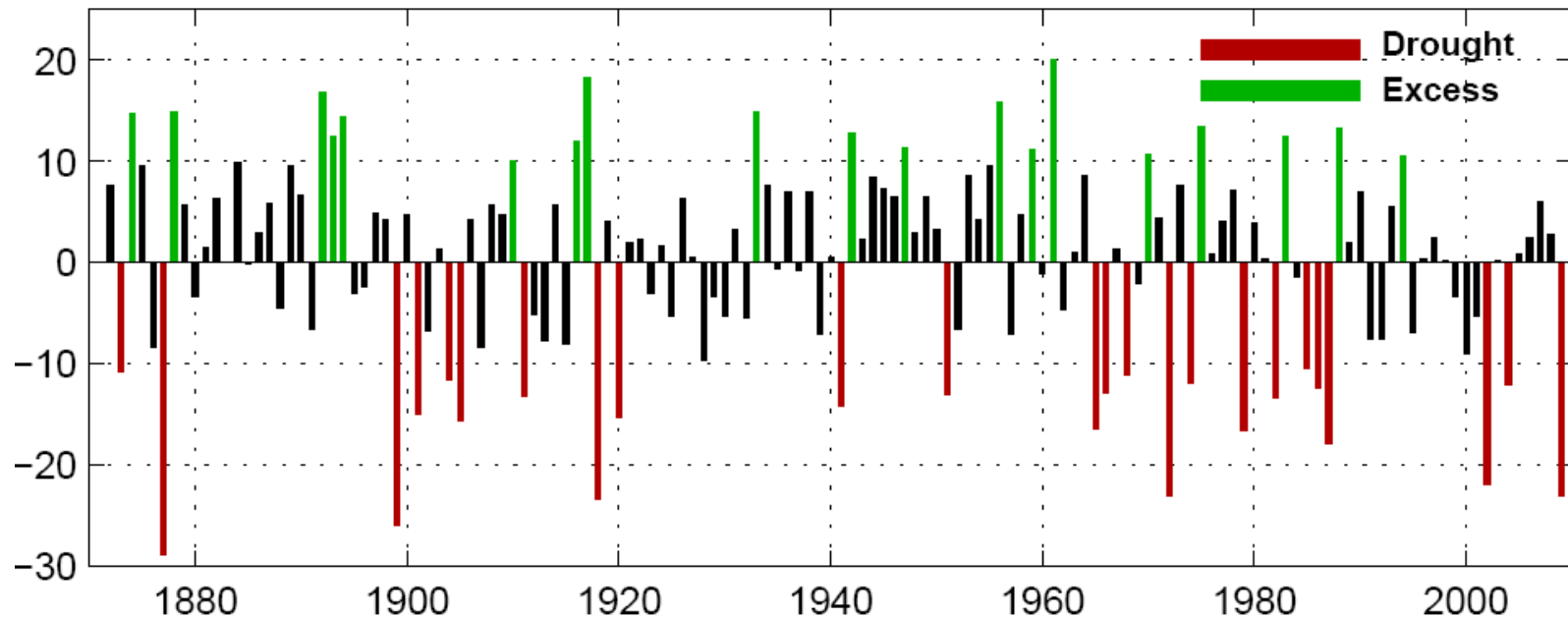


Interannual and Interdecadal variability

Interannual variability

- The long term mean of the Indian summer monsoon rainfall (ISMR) is about 85cm.
- Consider the interannual (year to year) variation of the anomaly (defined as the actual value in any year minus the mean value) of ISMR as a percentage of the mean value during 1876-2010.

Interannual Variation of the anomaly of ISMR (as % of the mean) during 1876-2010



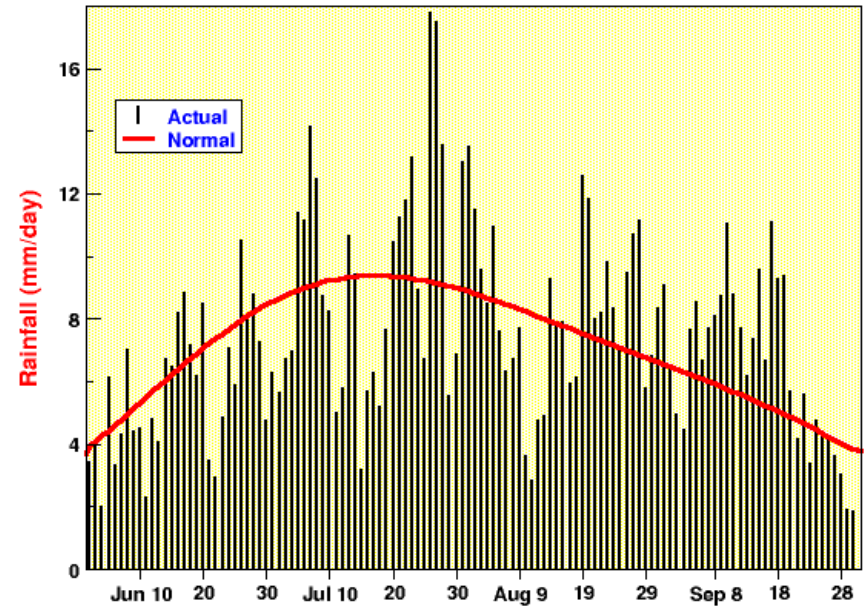
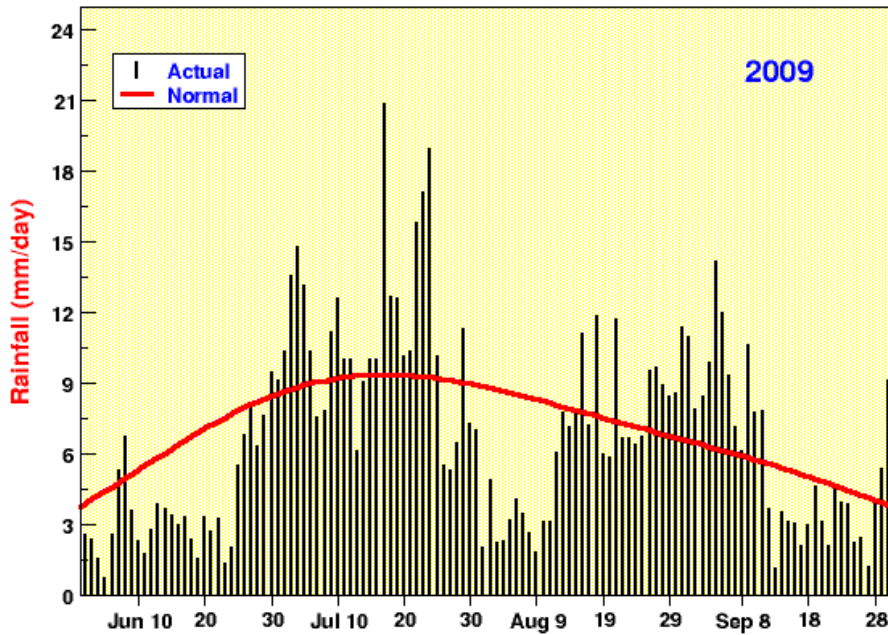
Drought: ISMR anomaly $< -10\%$ of the mean

Excess rainfall seasons: ISMR anomaly $> 10\%$ of the mean

Frequent droughts: 1899-1920 (7 in 21 years) ; 1965-87 (10 in 28 years) and again since 2002

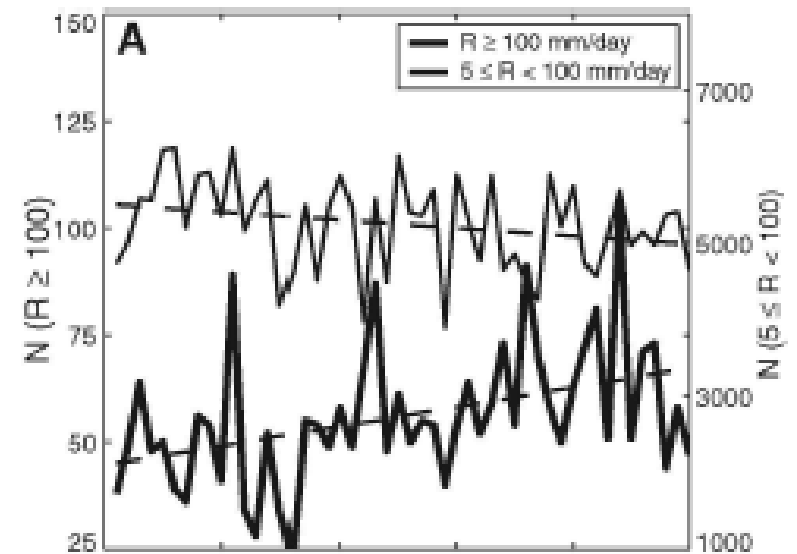
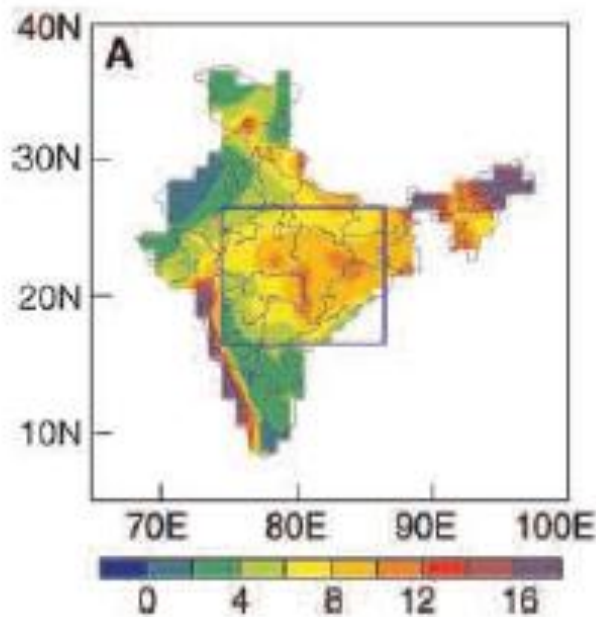
Less frequent during 1878-98 (0 in 21 years); 1921-64 (2 in 44 years)

Variation of the all-India daily rainfall during the summer monsoon seasons of 2009 and 2010



Increasing Trend of Extreme Rain Events Over India in a Warming Environment

B. N. Goswami,^{1*} V. Venugopal,² D. Sengupta,² M. S. Madhusoodanan,² Prince K. Xavier²



Forcing factors of Indian monsoon

Snow cover

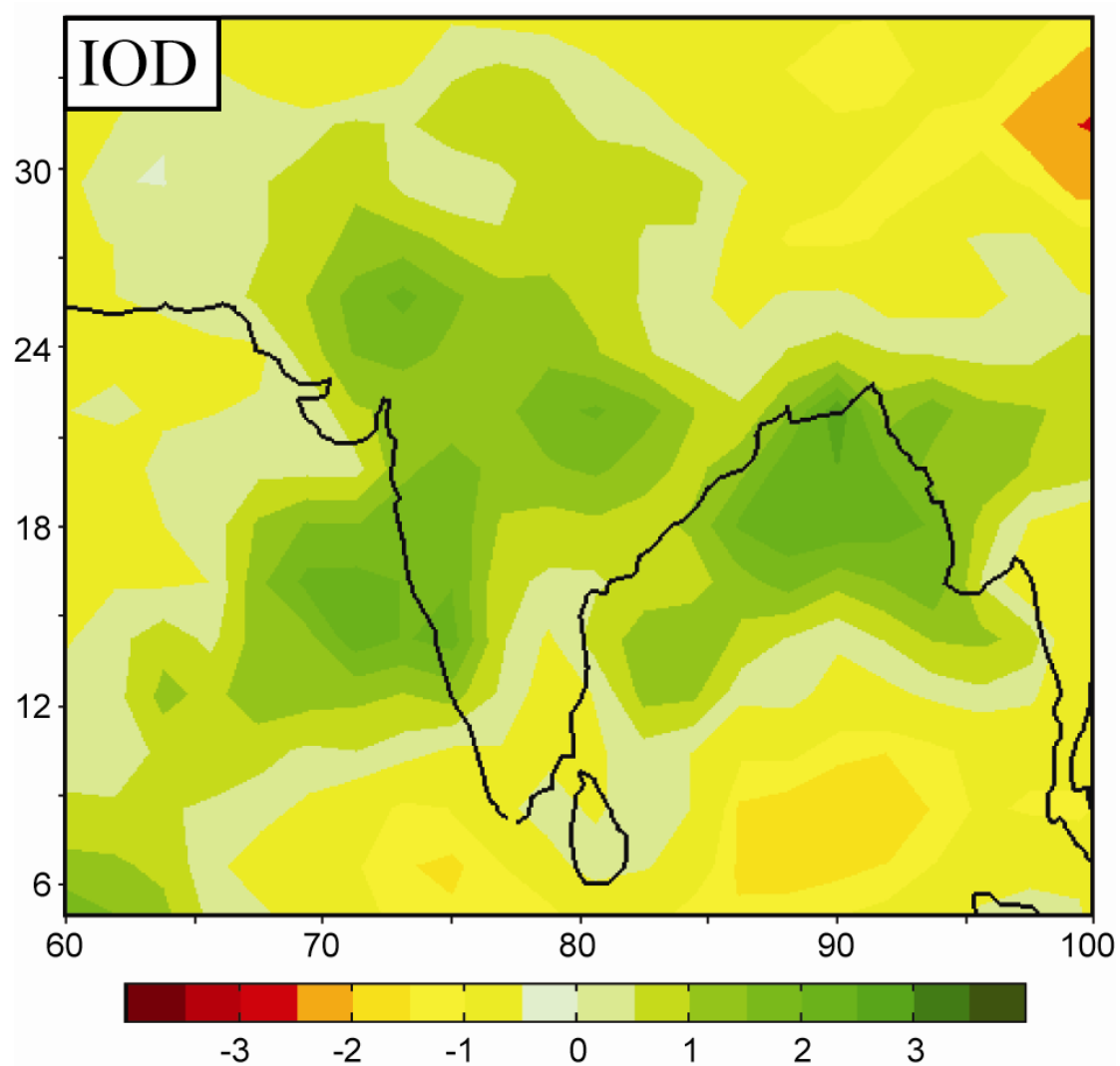
- Indian Monsoon
 - North American Monsoon
-
- Connection with neighboring snow cover of the preceding winter.
 - *Blanford Hypothesis*: Winter/Spring snow cover in the Himalayas mountain region has an inverse relationship with June-September Indian rainfall.

Snow cover

Contradictory results

- In recent years there have been contradictions against the Blanford Hypothesis.
- Fusallo (2004), Bamzai & Shukla (1999) and Ye & Bao (2001):
 - Positive correlation between snow cover and monsoonal rainfall anomalies
 - These results are at odds with the Blanford Hypothesis for the spatial distribution of snow cover.

Indian Ocean Dipole



What is IOD?

- First described by Dr. Saji et al. in 1999
- Dipole Mode Index (DMI)

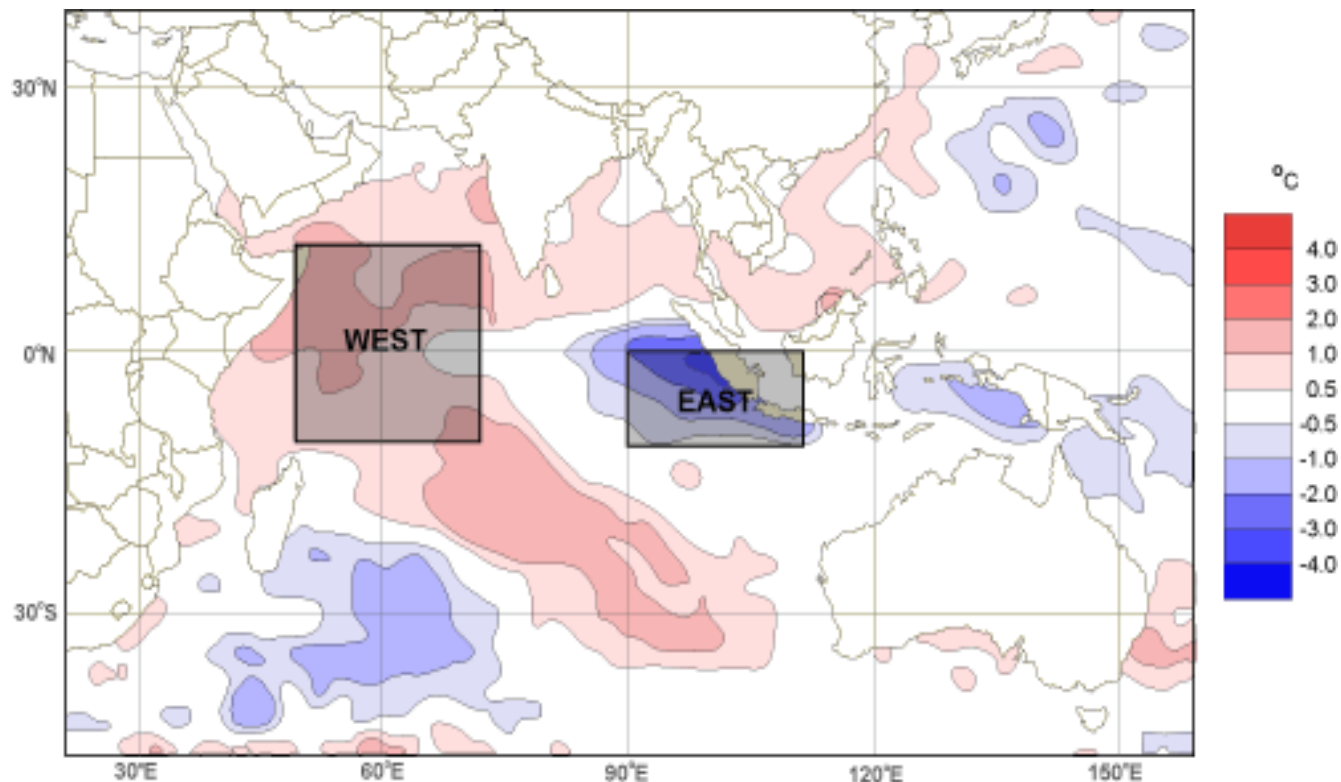
A dipole mode in the tropical Indian Ocean

R. M. Saji^{*}, B. M. Goswami[†], P. M. Unnikrishnan[‡] & T. Yamagata[‡]

^{*}Institute for Global Change Research, SEAMUS N 7F, 1-2-1 Shibaaura, Minato-ku, Tokyo 105 6791, Japan

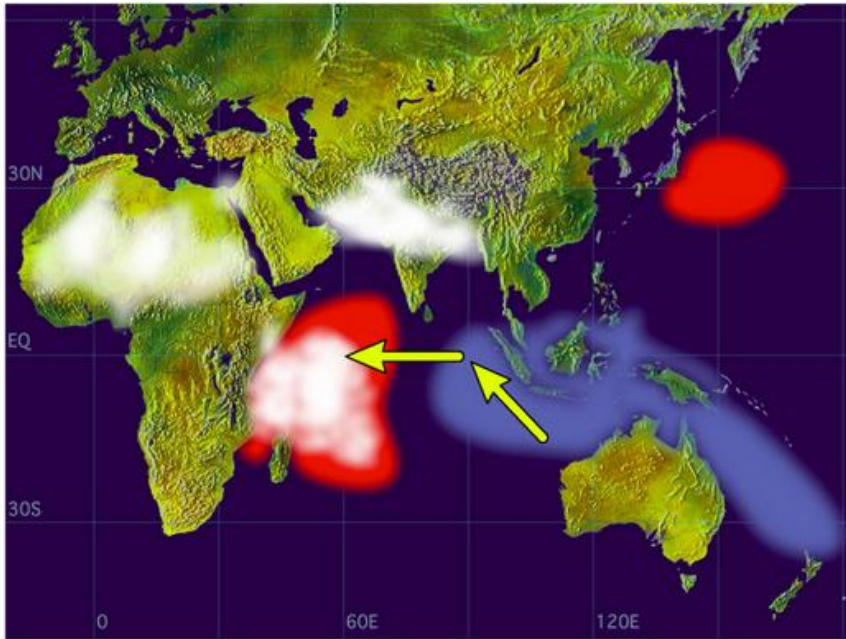
[†]Center for Atmospheric and Oceanic Science, Indian Institute of Science, Bangalore 560 012, India

[‡]Department of Earth and Planetary Physics, Graduate School of Science, The University of Tokyo, Tokyo 113 0033, Japan

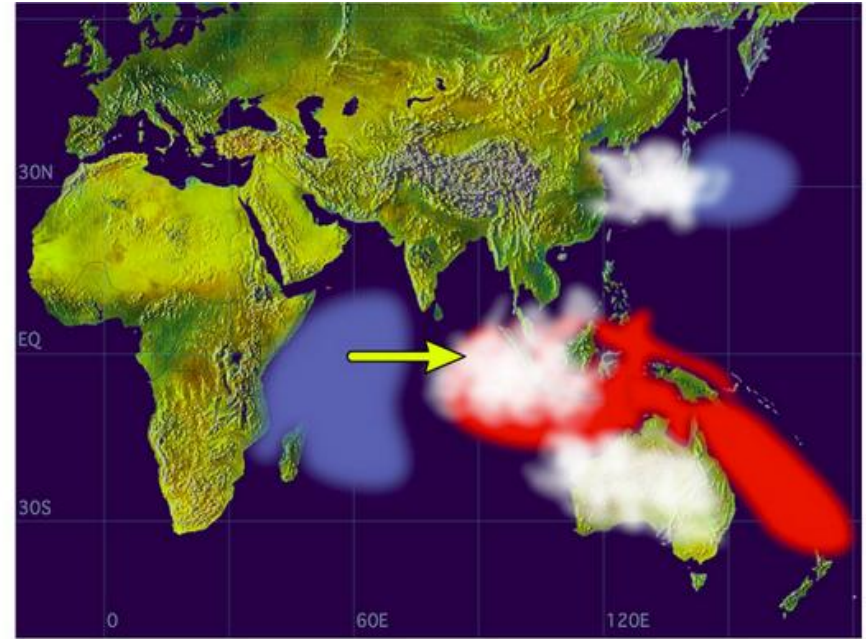


Positive & Negative Phases of IOD

- Positive: **cooler** in the tropical eastern Indian Ocean and **warmer** in the tropical western Indian Ocean
- Negative: **warmer** in the tropical eastern Indian Ocean and **cooler** in the tropical western Indian Ocean



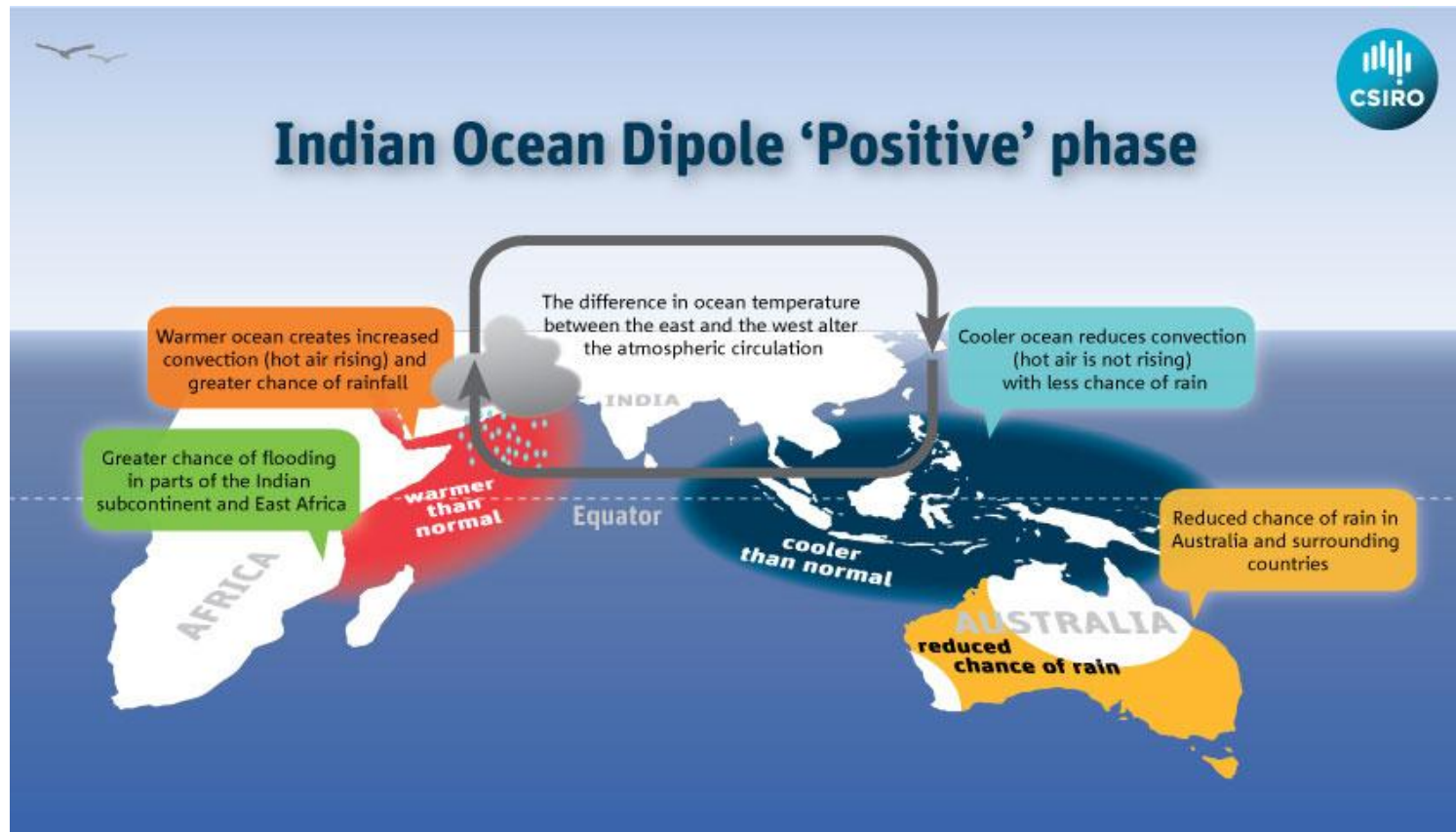
Positive IOD Phase



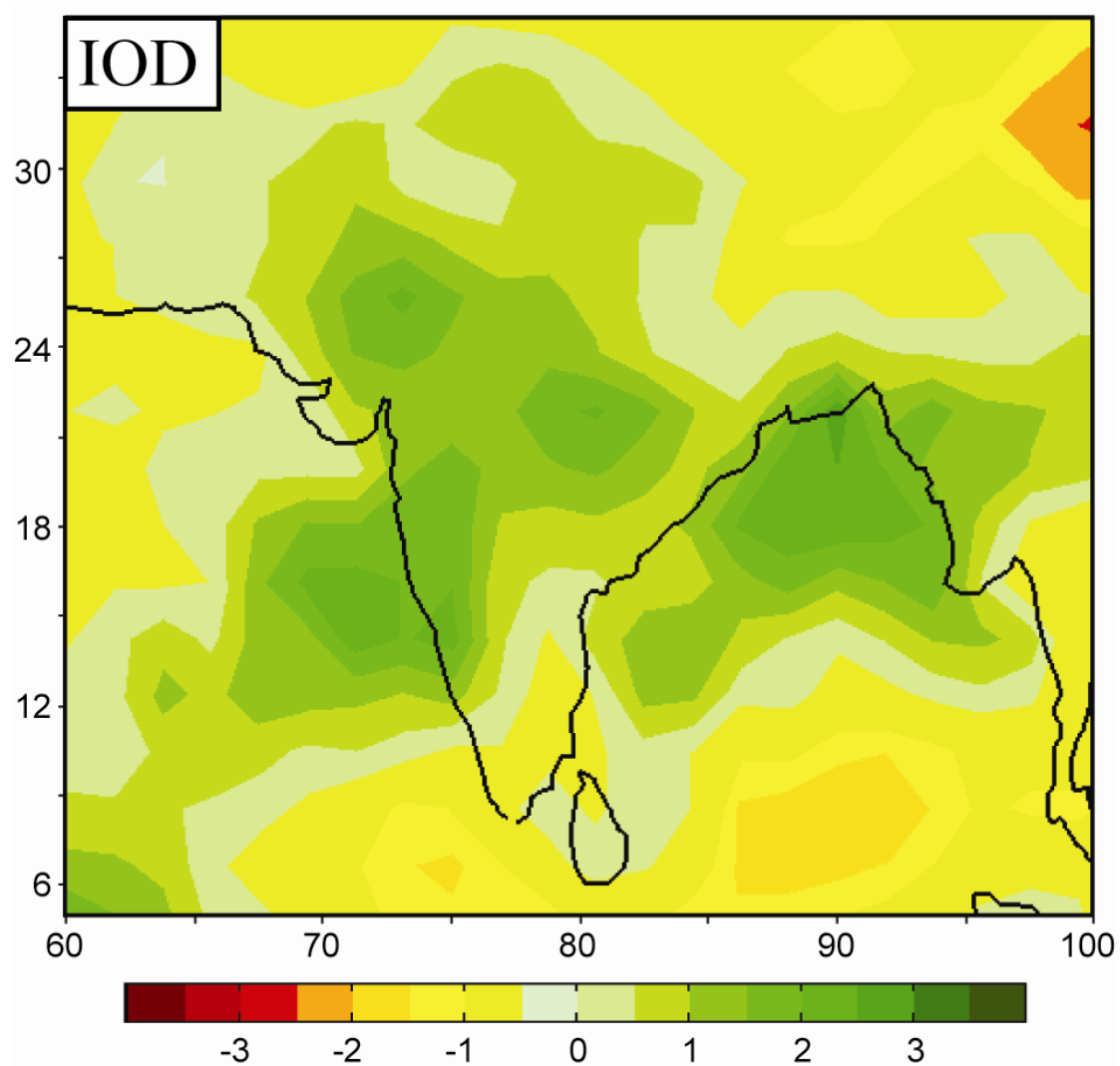
Negative IOD Phase

IOD Influence on Surrounding Climate

- A positive phase of the IOD tends to cause droughts in East Asia and Australia, and flooding in parts of the Indian subcontinent and East Africa.



Indian Ocean Dipole

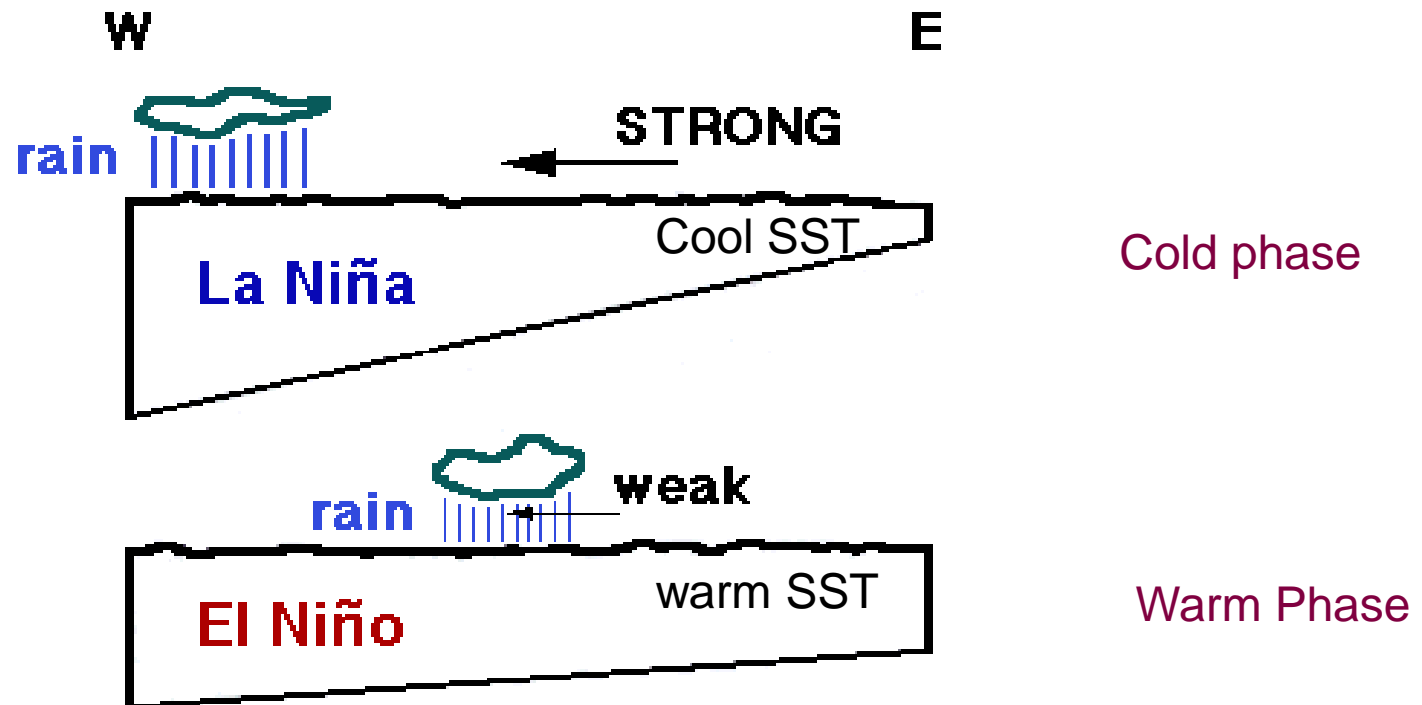


Questions still unsolved?

A better understanding of the relationship between the Indian Ocean Dipole and extreme weather events

- How IOD formed
- Relationship between ENSO and IOD
- How they affect the climate together

ENSO-monsoon teleconnections



Surface Layer of Low Latitude
Pacific Ocean (Warm Layer)

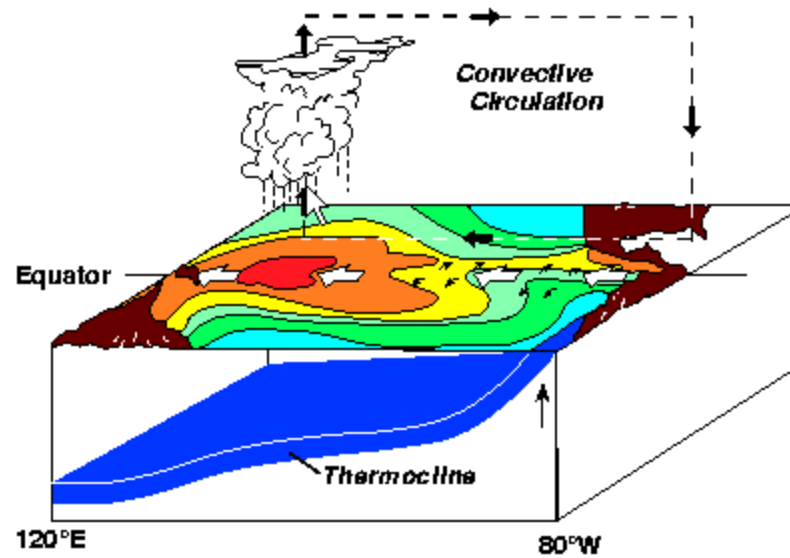
τ of oscillation ~ 4 years

Historically Speaking

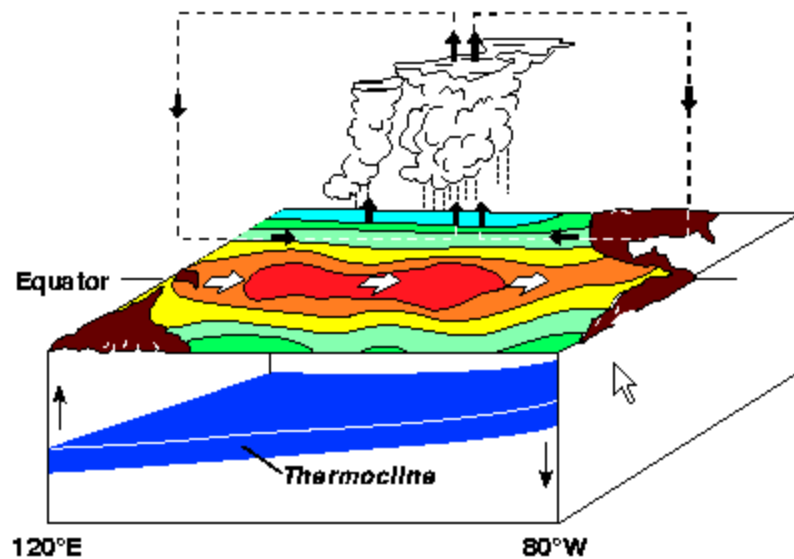
El Niño and La Niña events tend to develop during the period Apr-Jun and they:

- *Tend to reach their maximum strength during Dec-Feb*
- *Typically persist for 9-12 months, though occasionally persisting for up to 2 years*
- *Typically recur every 2 to 7 years*

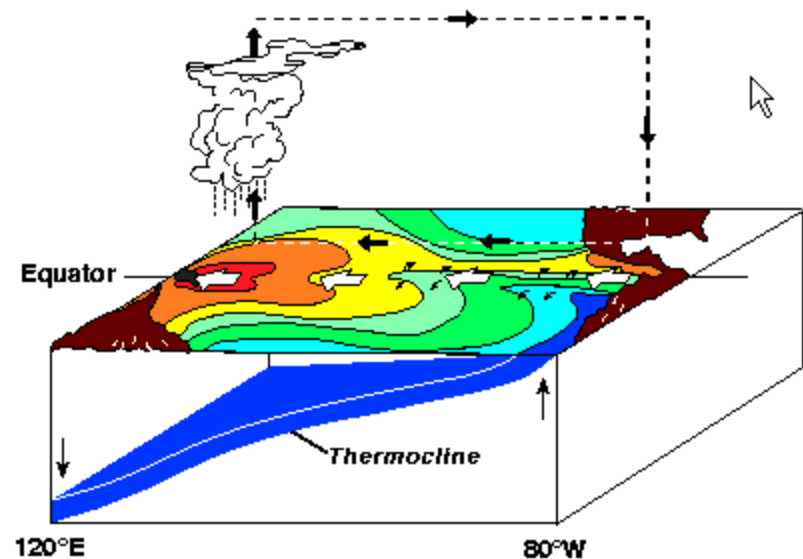
Normal Conditions



El Niño Conditions



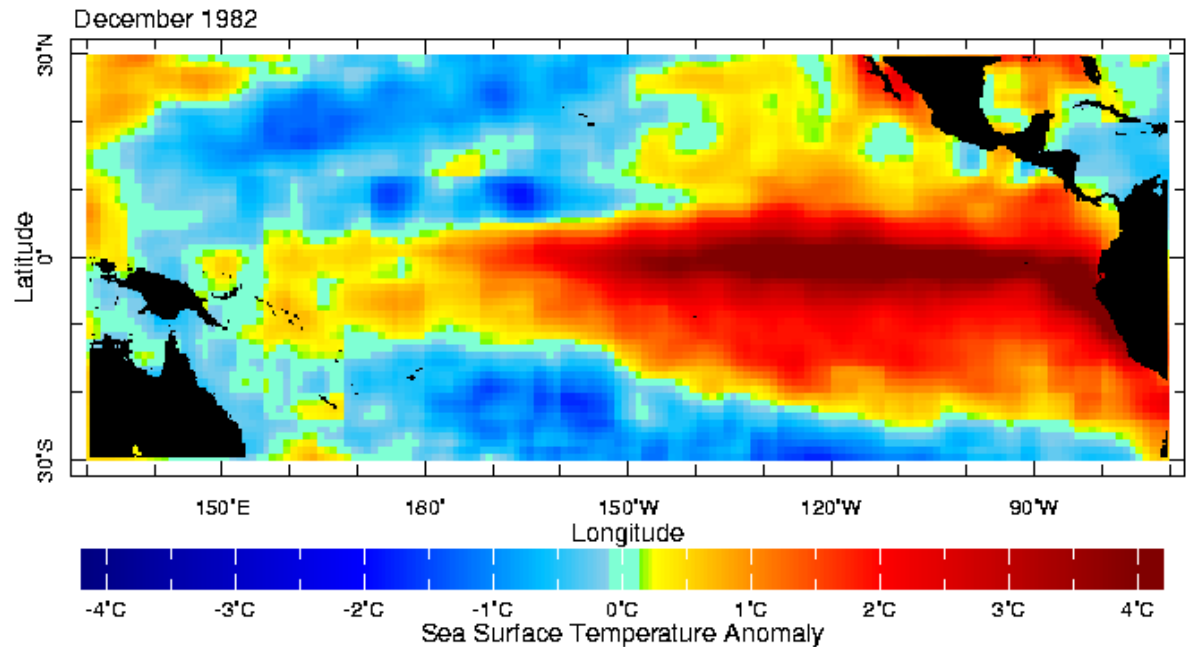
La Niña Conditions



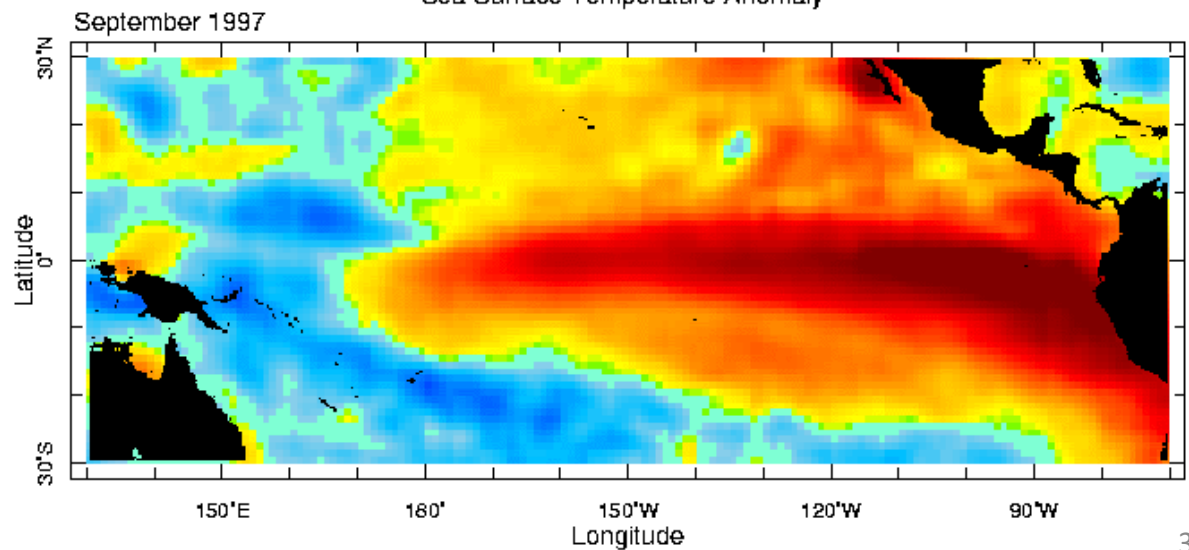
Largest El Niños of 20th century

1982

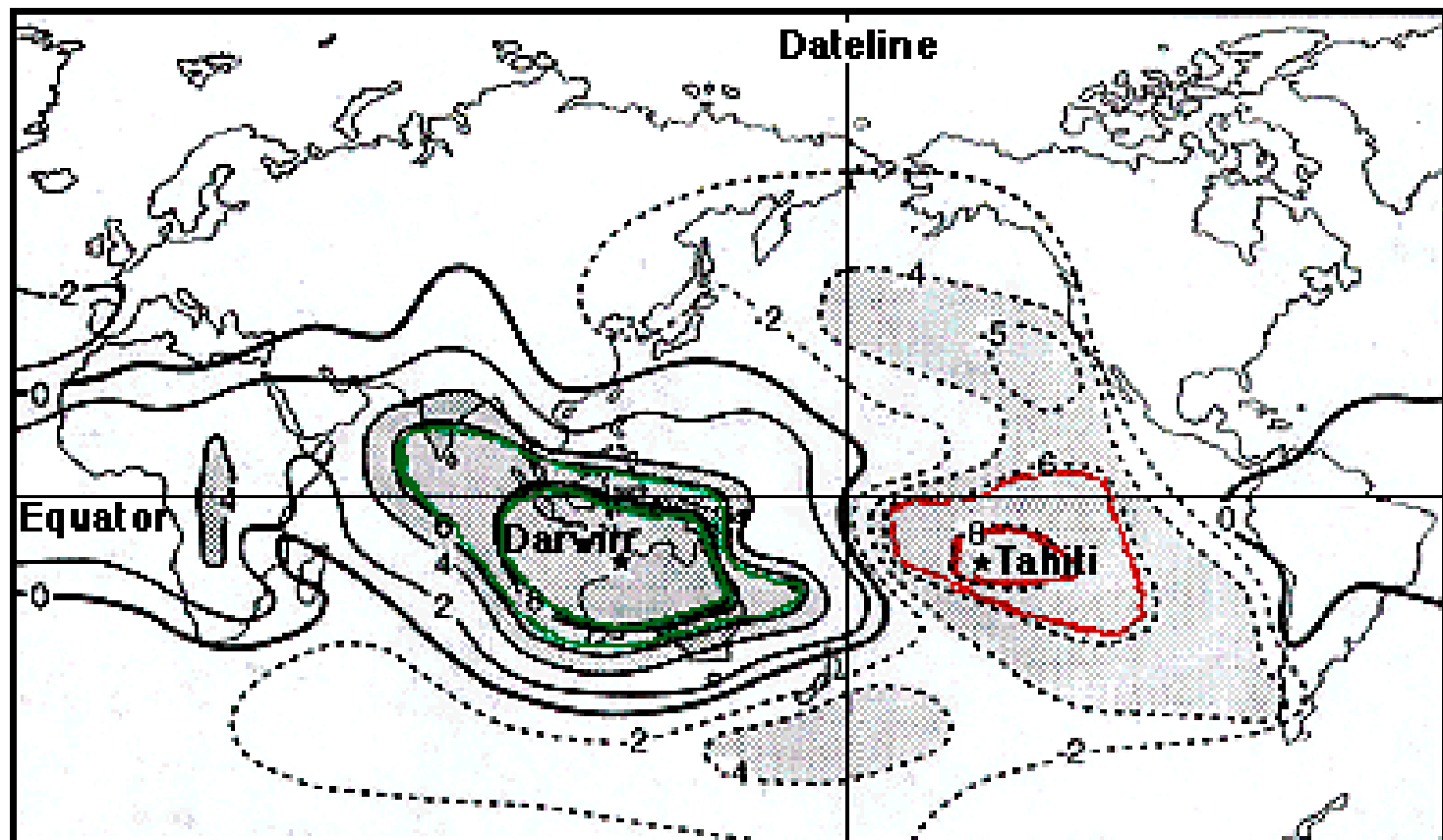
SST Anomaly (relative to
the average state)



1997



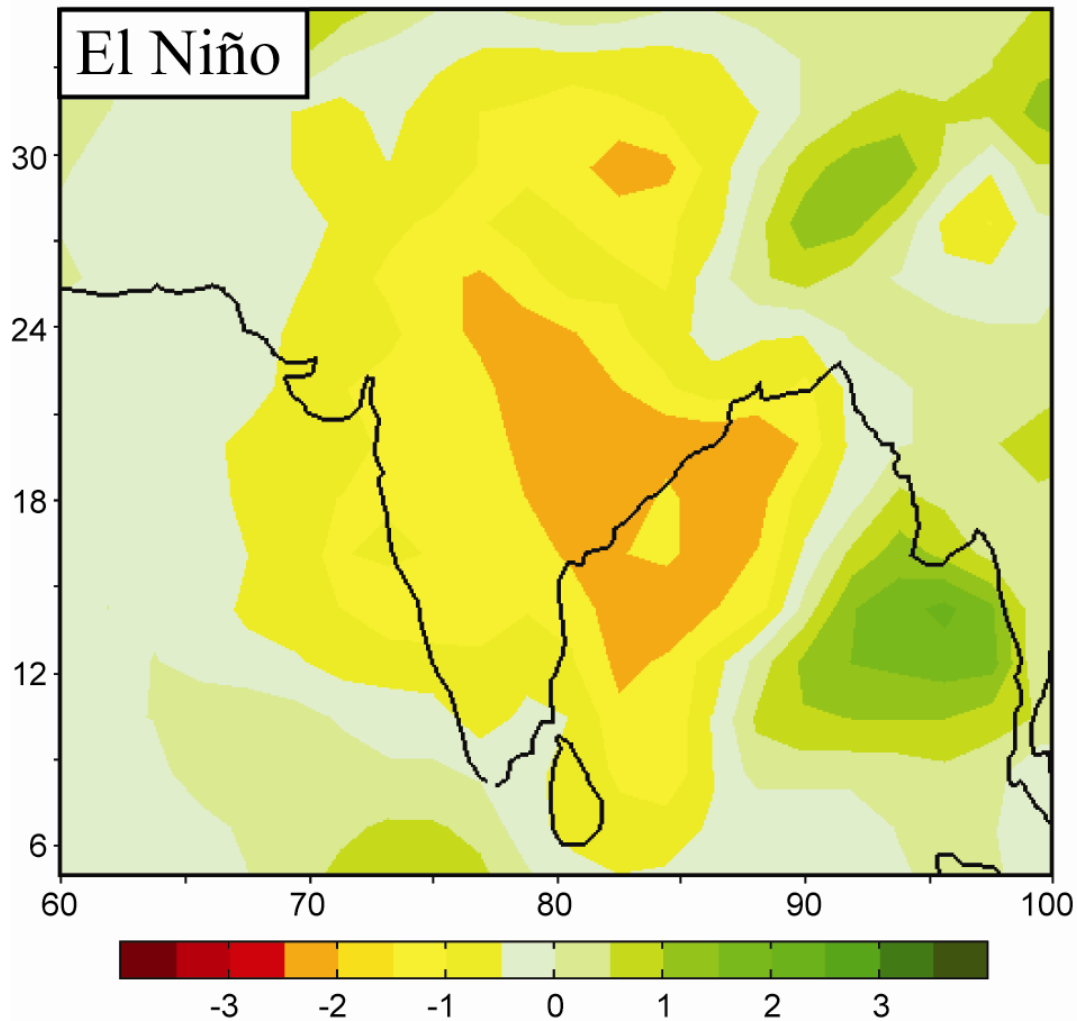
**SOI: Tahiti and Darwin as "centers of action",
mslp correlations between two locations**



Tahiti and Darwin are at opposite ends of the Southern Oscillation's seesaw,

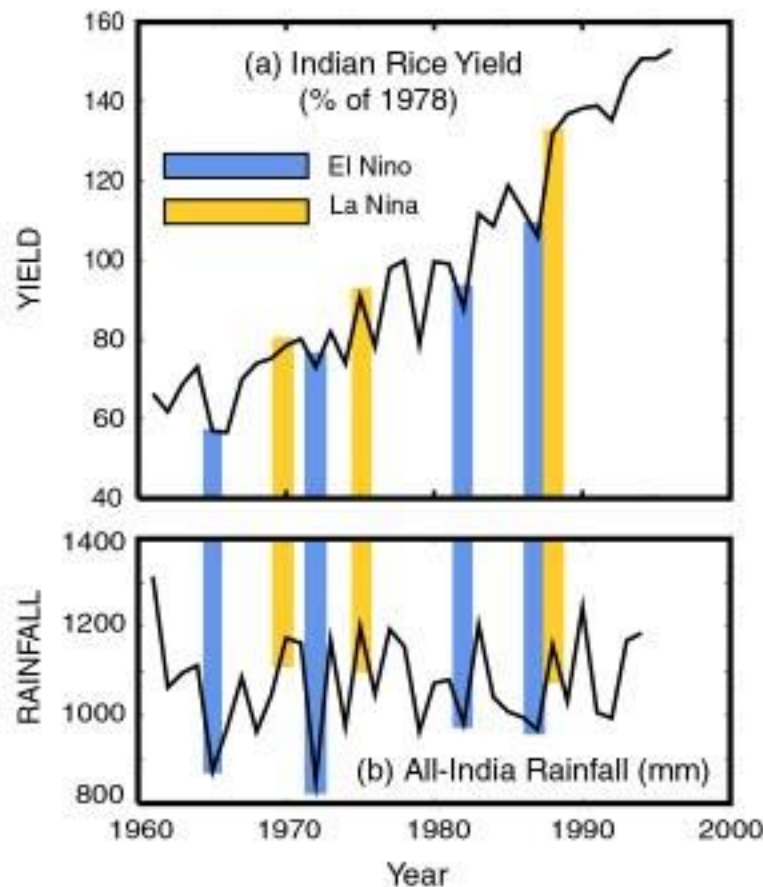
$$\text{SOI} = \text{Tahiti SLP} - \text{Darwin SLP}$$

Global teleconnections : ENSO



Spatial inhomogeneity in monsoon precipitation associated with teleconnections

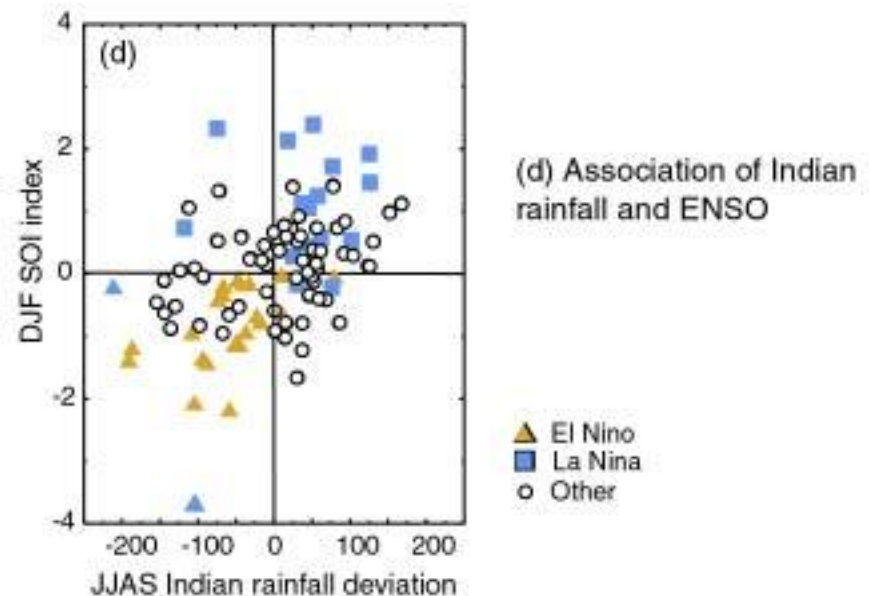
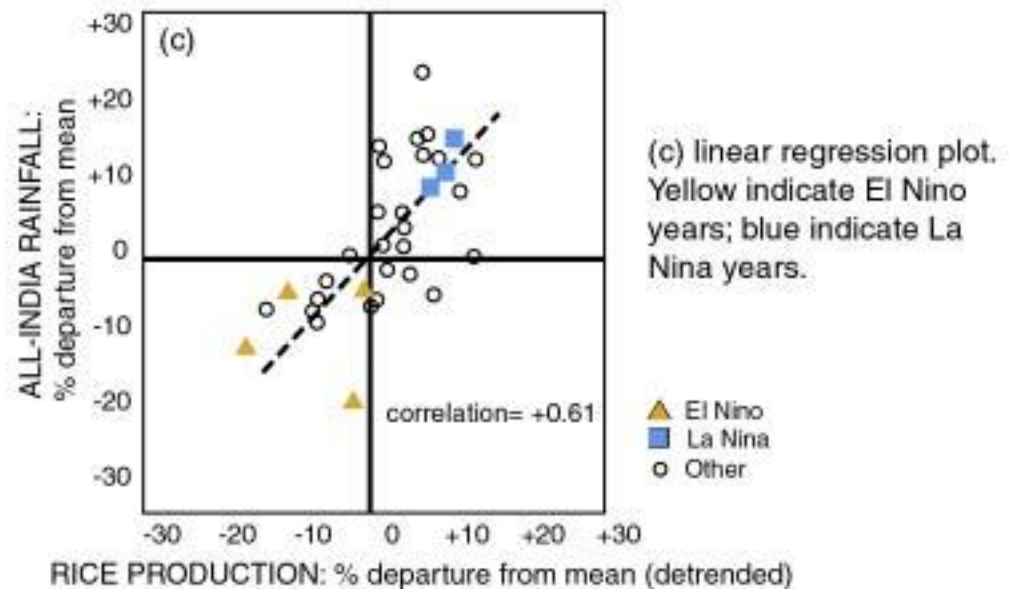
Association of monsoon rainfall and agriculture



From Webster et al. (1998)

(a) The relationship between Indian rainfall and rice production from 1960 to 1996 relative to 1978 production.

(b) the All-India rainfall Index for the corresponding years;

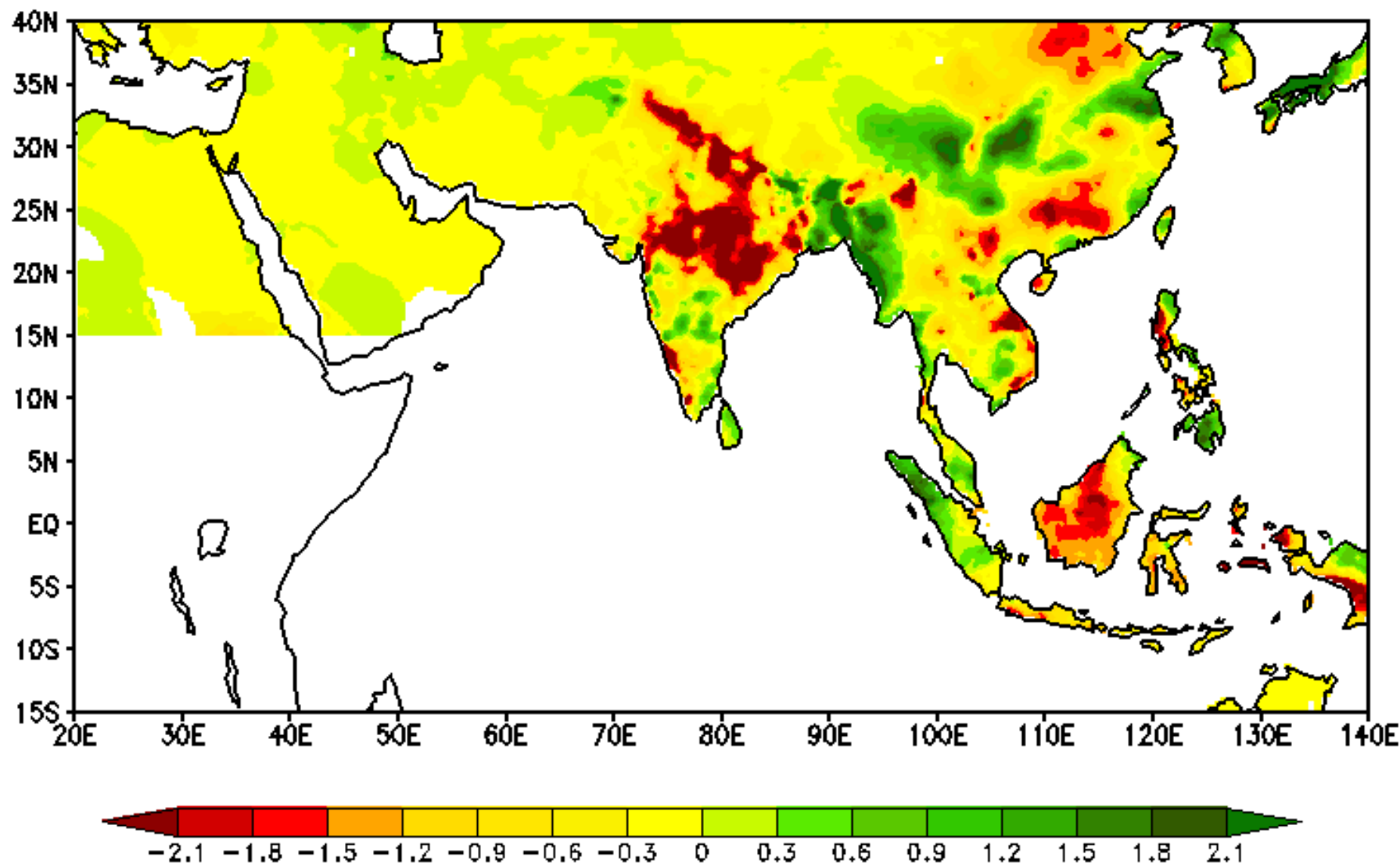


On the Weakening Relationship Between the Indian Monsoon and ENSO

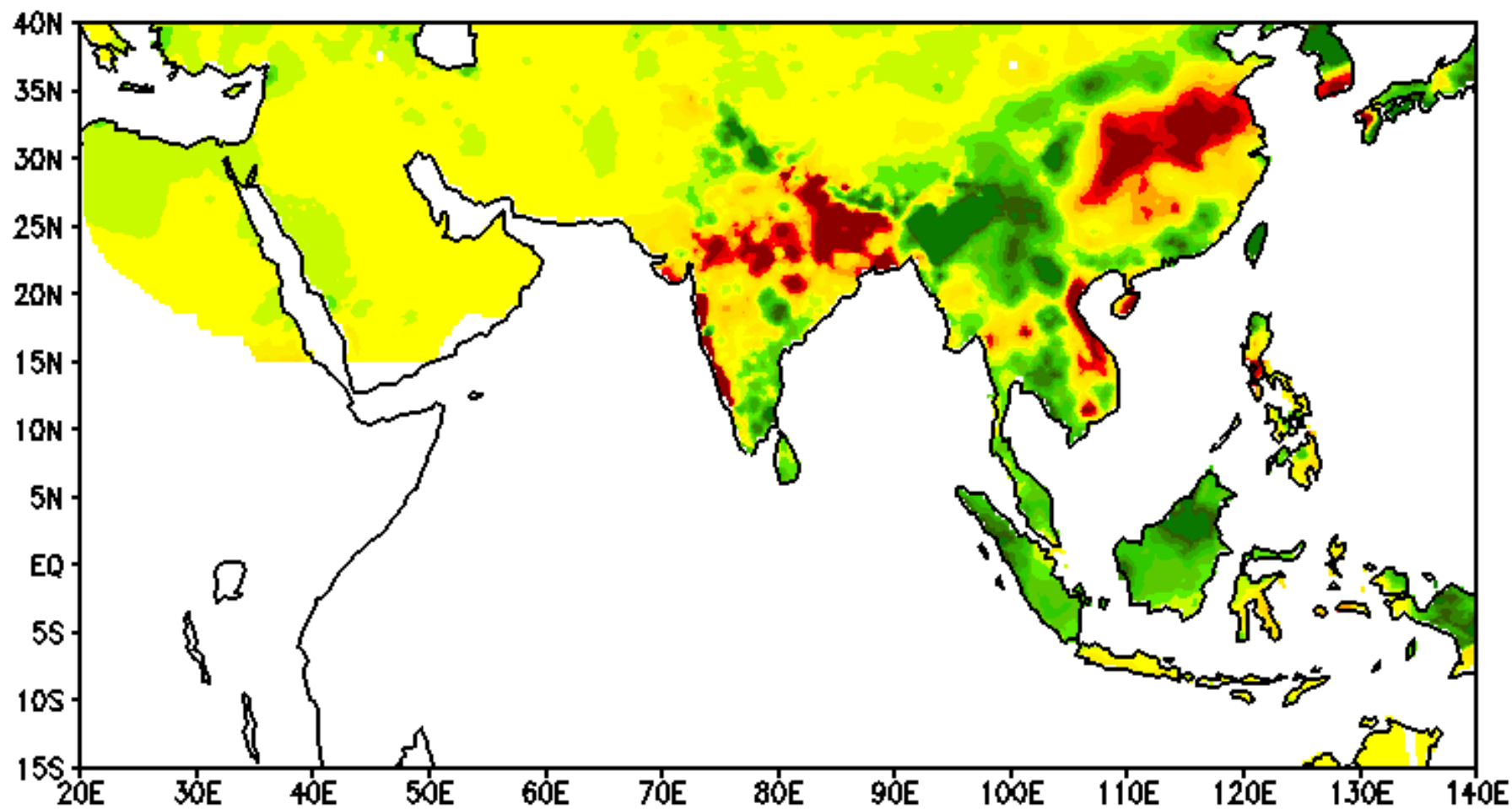
K. Krishna Kumar,^{1*}† Babji Rajagopalan,² Mark A. Cane²

Analysis of the 140-year historical record suggests that the inverse relationship between the El Niño–Southern Oscillation (ENSO) and the Indian summer monsoon (weak monsoon arising from warm ENSO event) has broken down in recent decades. Two possible reasons emerge from the analysis. A southeastward shift in the Walker circulation anomalies associated with ENSO events may lead to a reduced subsidence over the Indian region, thus favoring normal monsoon conditions. Additionally, increased surface temperatures over Eurasia in winter and spring, which are a part of the midlatitude continental warming trend, may favor the enhanced land-ocean thermal gradient conducive to a strong monsoon. These observations raise the possibility that the Eurasian warming in recent decades helps to sustain the monsoon rainfall at a normal level despite strong ENSO events.

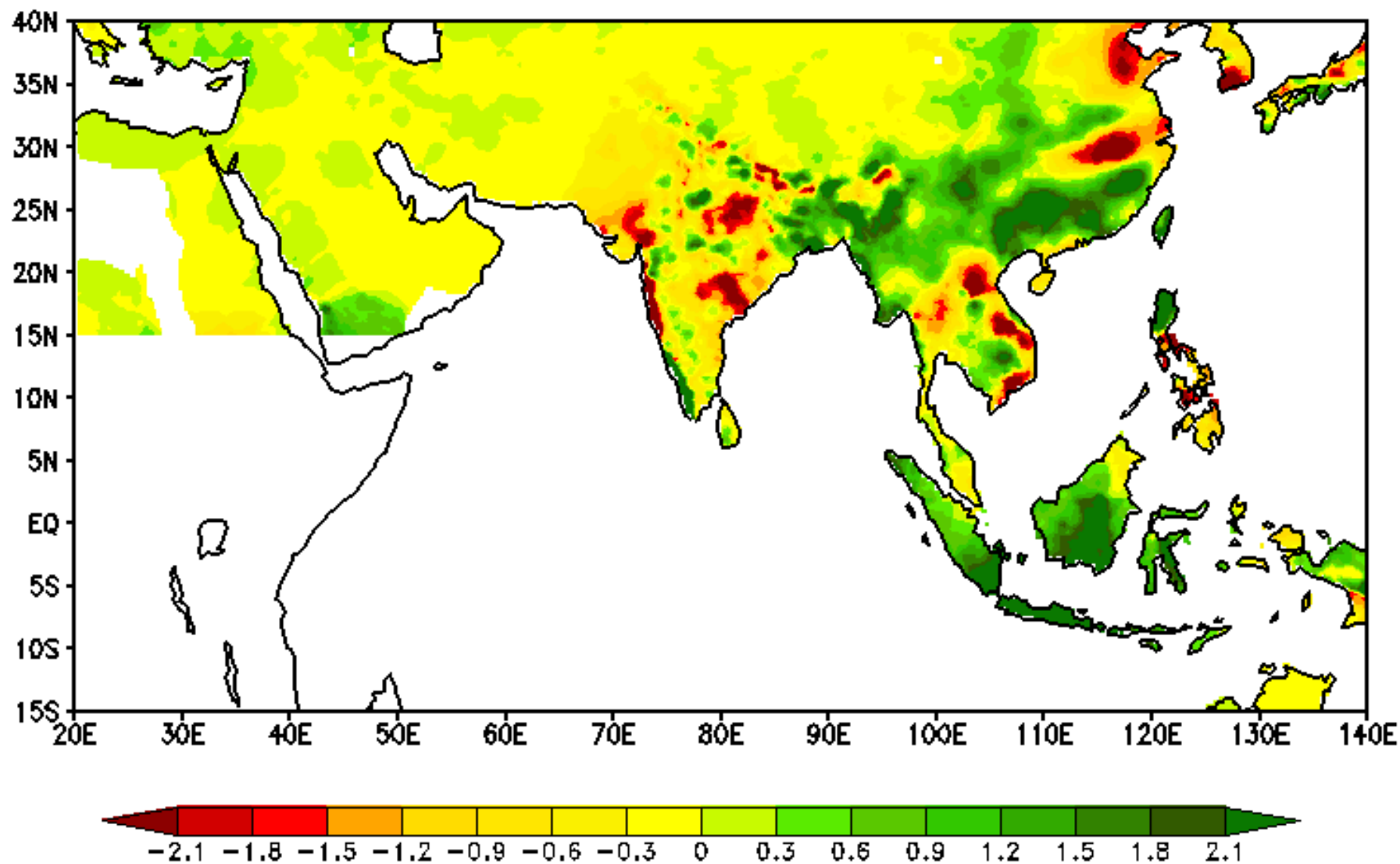
1965



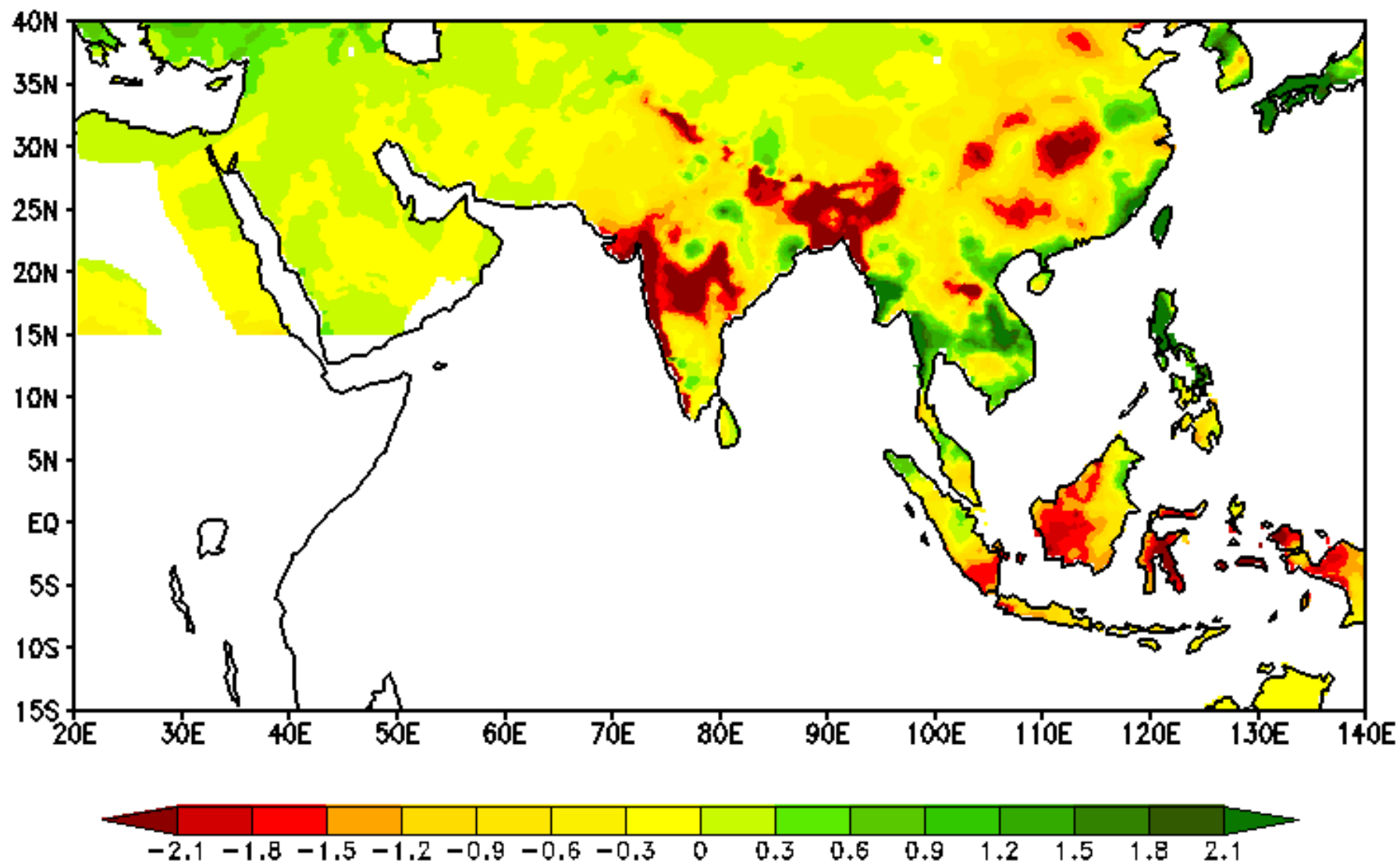
1966



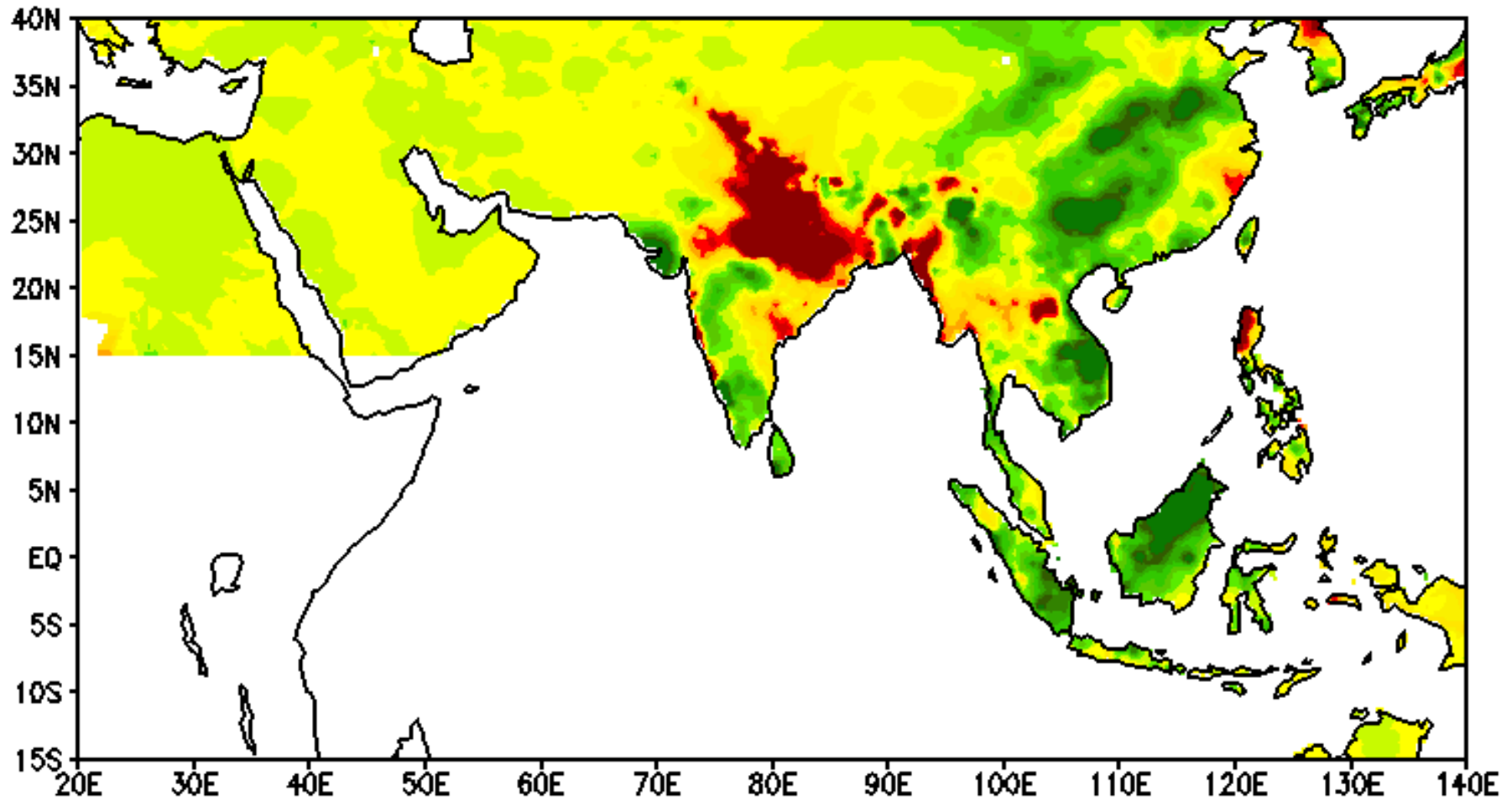
1968



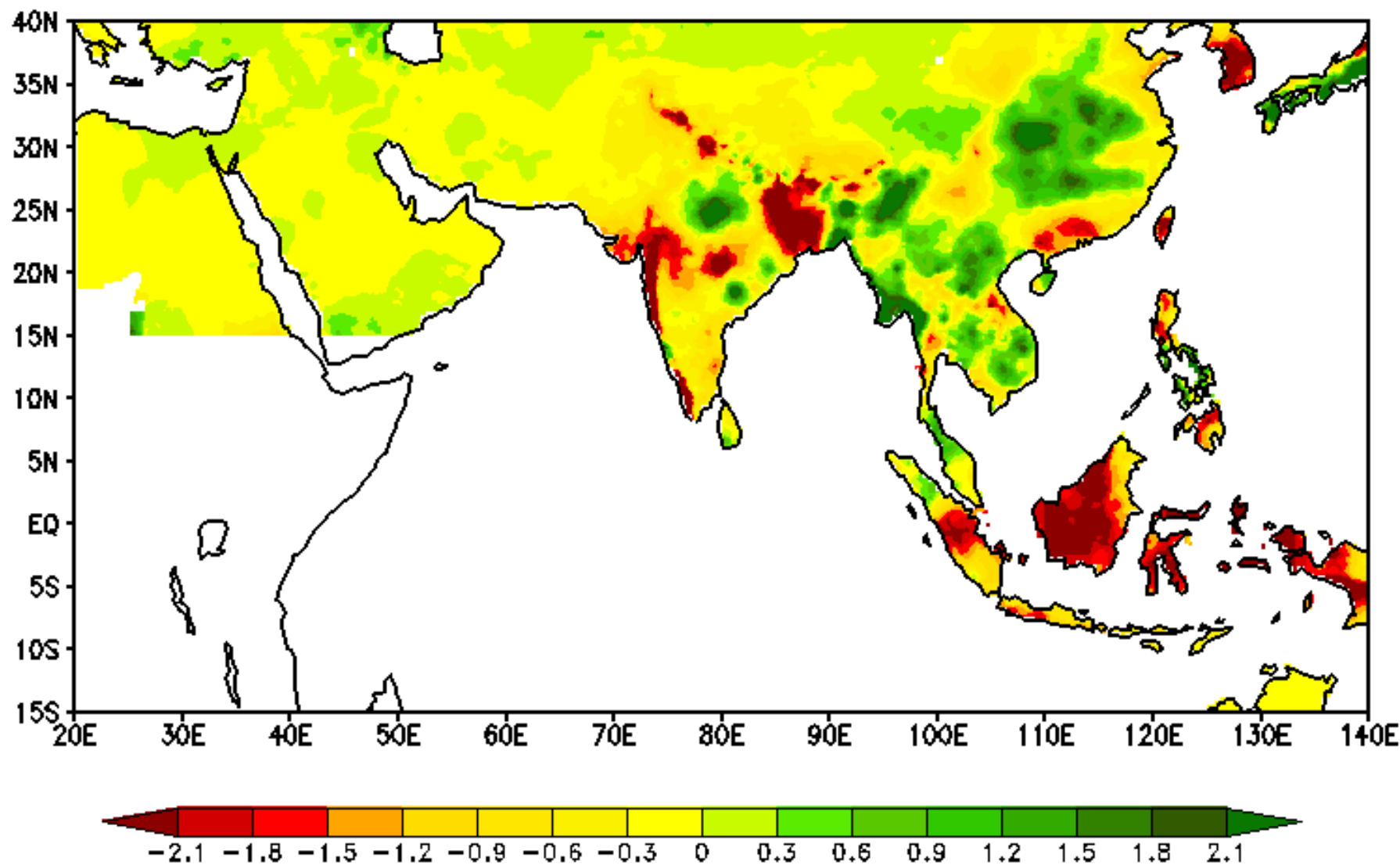
1972



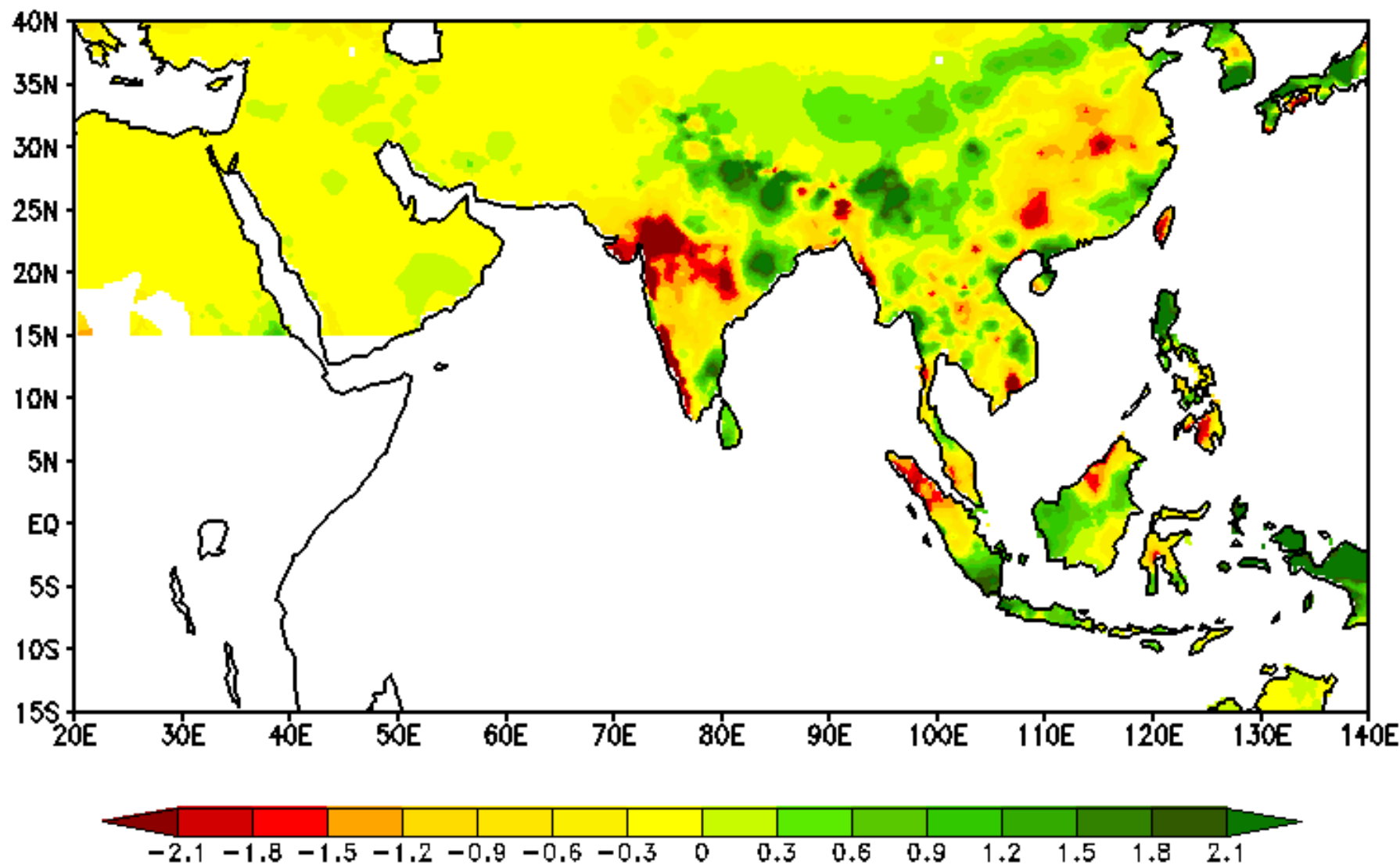
1979



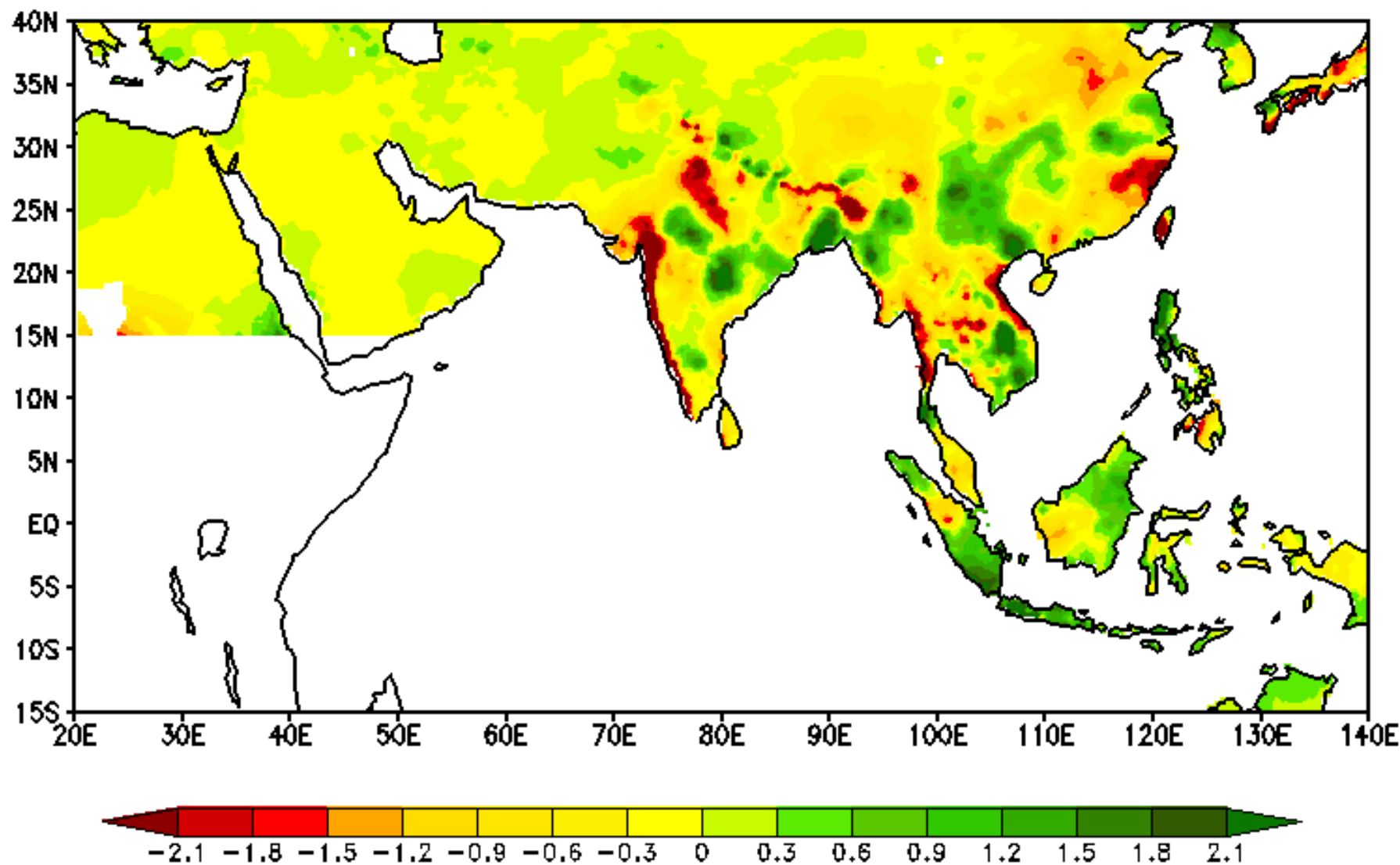
1982



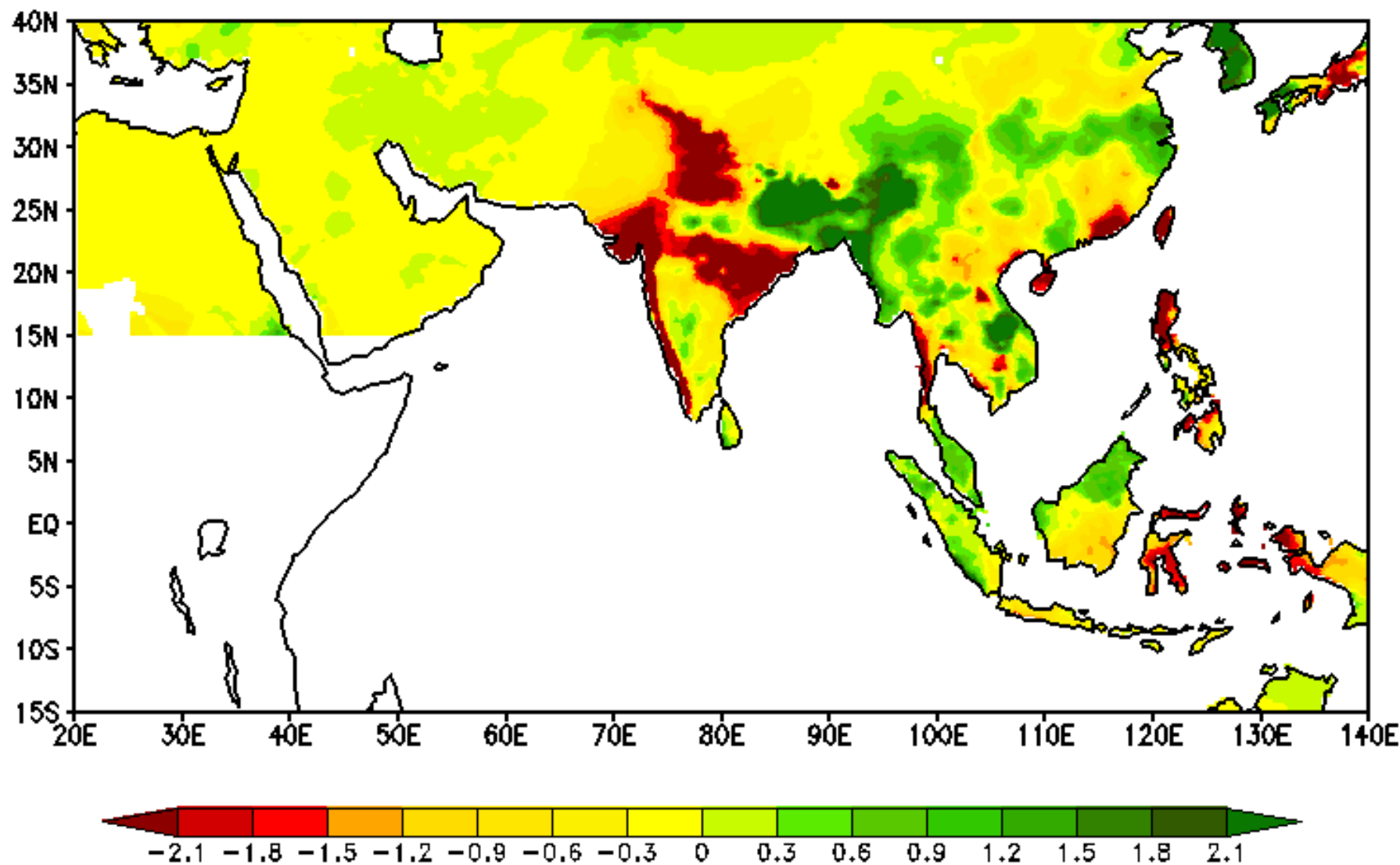
1985



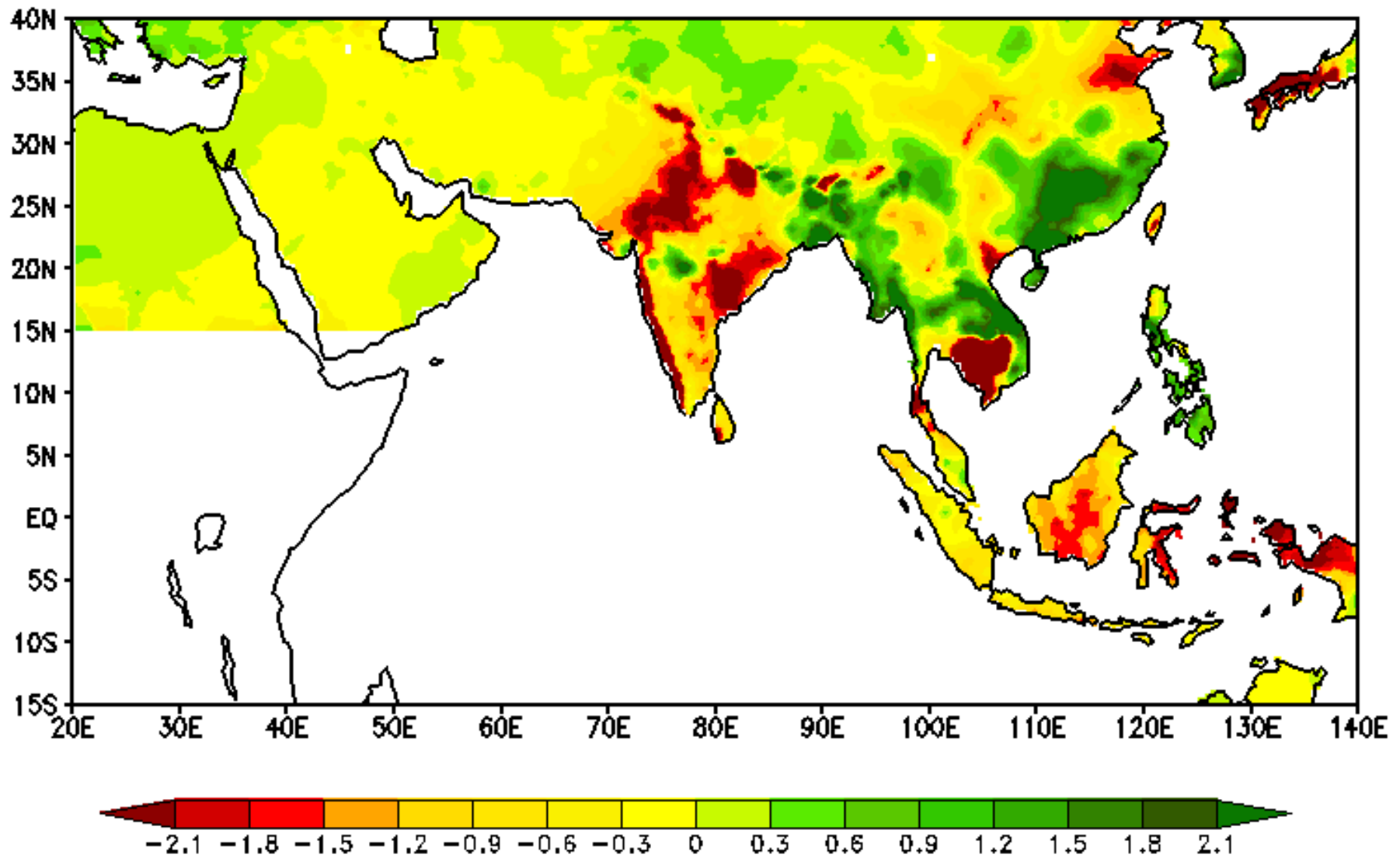
1986



1987

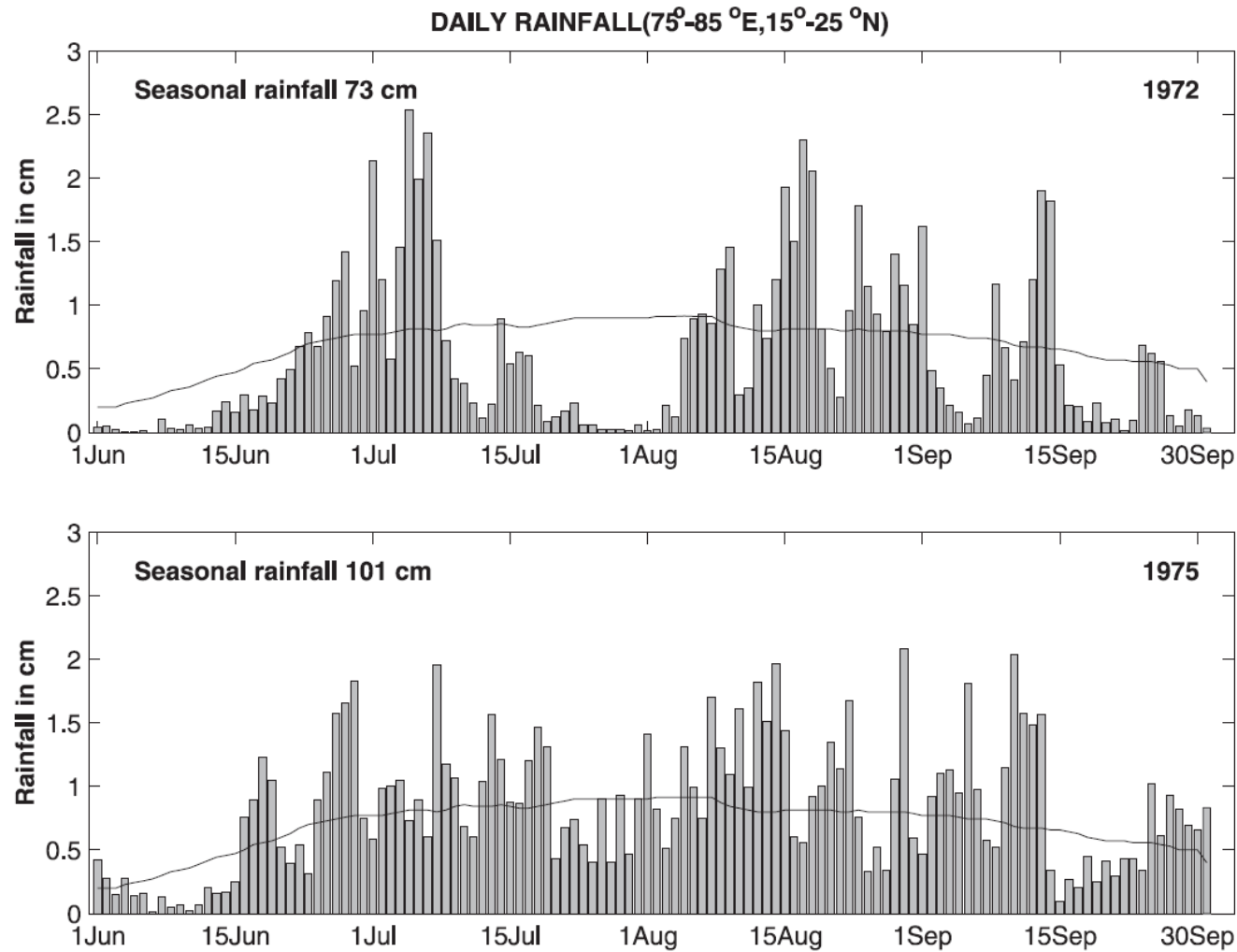


2002



Break Monsoon

Breaks in Indian summer monsoon



Breaks in Indian summer monsoon

