

# The computation aspects of the equivalent-layer technique: review and perspective

Diego Takahashi  $^{1,*}$ , André L. A. Reis  $^2$ , Vanderlei C. Oliveira Jr.  $^1$  and Valéria C. F. Barbosa  $^1$ 

Correspondence\*: Valéria C.F. Barbosa valcris@on.br

<sup>&</sup>lt;sup>1</sup>Observatório Nacional, Department of Geophysics, Rio de Janeiro, Brasil

<sup>&</sup>lt;sup>2</sup> Universidade do Estado do Rio de Janeiro, Department of Applied Geology, Rio de Janeiro, Brasil

# 1 FUNDAMENTALS

- 2 Let d be a  $D \times 1$  vector, whose i-th element  $d_i$  is the observed potential field at the position  $(x_i, y_i, z_i)$ ,
- 3  $i \in \{1:D\}$ , of a topocentric Cartesian system with x, y and z axes pointing to north, east and down,
- 4 respectively. Consider that  $d_i$  can be satisfactorily approximated by a harmonic function

$$f_i = \sum_{j=1}^{P} g_{ij} p_j , \quad i \in \{1 : D\} ,$$
 (1)

- 5 where,  $p_i$  represents the scalar physical property of a virtual source (i.e., monopole, dipole, prism) located
- 6 at  $(x_j, y_j, z_j), j \in \{1 : P\}$  and

$$g_{ij} \equiv g(x_i - x_j, y_i - y_j, z_i - z_j), \quad z_i < \min\{z_j\}, \quad \forall i \in \{1 : D\},$$
 (2)

- 7 is a harmonic function, where  $\min\{z_j\}$  denotes the minimum  $z_j$ , or the vertical coordinate of the shallowest
- 8 virtual source. These virtual sources are called equivalent sources and they form an equivalent layer. In
- 9 matrix notation, the potential field produced by all equivalent sources at all points  $(x_i, y_i, z_i)$ ,  $i \in \{1 : D\}$ ,
- 10 is given by:

$$\mathbf{f} = \mathbf{G}\mathbf{p} \,, \tag{3}$$

- 11 where  $\mathbf{p}$  is a  $P \times 1$  vector with j-th element  $p_j$  representing the scalar physical property of the j-th
- 12 equivalent source and G is a  $D \times P$  matrix with element  $g_{ij}$  given by equation 2.
- 13 The equivalent-layer technique consists in solving a linear inverse problem to determine a parameter
- 14 vector p leading to a predicted data vector f (equation 3) sufficiently close to the observed data vector d,
- whose i-th element  $d_i$  is the observed potential field at  $(x_i, y_i, z_i)$ . The notion of *closeness* is intrinsically
- related to the concept of vector norm (e.g., Golub and Van Loan, 2013, p. 68) or measure of length (e.g.,
- 17 Menke, 2018, p. 41). Because of that, almost all methods for determining p actually estimate a parameter
- vector  $\tilde{\mathbf{p}}$  minimizing a length measure of the difference between  $\mathbf{f}$  and  $\mathbf{d}$  (see subsection 1.3). Given an
- 19 estimate  $\tilde{p}$ , it is then possible to compute a potential field transformation

$$\mathbf{t} = \mathbf{A}\tilde{\mathbf{p}} \,, \tag{4}$$

where t is a  $T \times 1$  vector with k-th element  $t_k$  representing the transformed potential field at the position

21 
$$(x_k, y_y, z_k), k \in \{1:T\}$$
, and

$$a_{kj} \equiv a(x_k - x_j, y_k - y_j, z_k - z_j), \quad z_k < \min\{z_j\}, \quad \forall k \in \{1:T\},$$
 (5)

22 is a harmonic function representing the kj-th element of the  $T \times P$  matrix A.

# 23 1.1 Spatial distribution and total number of equivalent sources

- There is no well-established criteria to define the optimum number P or the spatial distribution of the
- 25 equivalent sources. We know that setting an equivalent layer with more (less) sources than potential-field
- 26 data usually leads to an underdetermined (overdetermined) inverse problem (e.g., Menke, 2018, p. 52–53).
- 27 Concerning the spatial distribution of the equivalent sources, the only condition is that they must rely on a
- 28 surface that is located below and does not cross that containing the potential field data. Soler and Uieda
- 29 (2021) present a practical discussion about this topic.

From a theoretical point of view, the equivalent layer reproducing a given potential field data set cannot cross the true gravity or magnetic sources. This condition is a consequence of recognizing that the equivalent layer is essentially an indirect solution of a boundary value problem of potential theory (e.g., Roy, 1962; Zidarov, 1965; Dampney, 1969; Li et al., 2014; Reis et al., 2020). In practical applications, however, there is no guarantee that this condition is satisfied. Actually, its is widely known from practical experience (e.g., Gonzalez et al., 2022) that the equivalent-layer technique works even for the case in which the layer cross the true sources.

Regarding the depth of the equivalent layer, Dampney (1969) proposed a criterion based on horizontal data sampling, suggesting that the equivalent-layer depth should be between two and six times the horizontal grid spacing, considering evenly spaced data. However, when dealing with a survey pattern that has unevenly spaced data, Reis et al. (2020) adopted an alternative empirical criterion. According to their proposal, the depth of the equivalent layer should range from two to three times the spacing between adjacent flight lines. The criteria of Dampney (1969) and Reis et al. (2020) are valid for planar equivalent layers. Cordell (1992) have proposed and an alternative criterion for scattered data that leads to an undulating equivalent layer. This criterion have been slightly modified by Guspí et al. (2004), Guspí and Novara (2009) and Soler and Uieda (2021), for example, and consists in setting one equivalent source below each datum at a depth proportional to the horizontal distance to the nearest neighboring data points. Soler and Uieda (2021) have compared different strategies for defining the equivalent sources depth for the specific problem of interpolating gravity data, but they have not found significant differences between them.

#### 49 **1.2 Matrix** G

37

38

39 40

41

42 43

44

45

46

47

48

Generally, the harmonic function  $g_{ij}$  (equation 2) is defined in terms of the inverse distance between the observation point  $(x_i, y_i, z_i)$  and the j-th equivalent source at  $(x_j, y_j, z_j)$ ,

$$\frac{1}{r_{ij}} \equiv \frac{1}{\sqrt{(x_i - x_j)^2 + (y_i - y_j)^2 + (z_i - z_j)^2}},$$
(6)

52 or by its partial derivatives of first and second orders, respectively given by

$$\partial_{\alpha} \frac{1}{r_{ij}} \equiv \frac{-(\alpha_i - \alpha_j)}{r_{ij}^3} , \quad \alpha \in \{x, y, z\} , \tag{7}$$

53 and

$$\partial_{\alpha\beta} \frac{1}{r_{ij}} \equiv \begin{cases} \frac{3(\alpha_i - \alpha_j)^2}{r_{ij}^5} , & \alpha = \beta ,\\ \frac{3(\alpha_i - \alpha_j)(\beta_i - \beta_j)}{r_{ij}^5} - \frac{1}{r_{ij}^3} , & \alpha \neq \beta , \end{cases} \quad \alpha, \beta \in \{x, y, z\} . \tag{8}$$

In this case, the equivalent layer is formed by punctual sources representing monopoles or dipoles (e.g., Dampney, 1969; Emilia, 1973; Leão and Silva, 1989; Cordell, 1992; Oliveira Jr. et al., 2013; Siqueira et al., 55 2017; Reis et al., 2020; Takahashi et al., 2020; Soler and Uieda, 2021; Takahashi et al., 2022). Another 56 common approach consists in not defining  $g_{ij}$  by using equations 6–8, but other harmonic functions 57 58 obtained by integrating them over the volume of regular prisms (e.g., Li and Oldenburg, 2010; Barnes and Lumley, 2011; Li et al., 2014; Jirigalatu and Ebbing, 2019). There are also some less common approaches 59 defining the harmonic function  $g_{ij}$  (equation 2) as the potential field due to plane faces with constant physical property (Hansen and Miyazaki, 1984), doublets (Silva, 1986) or by computing the double 61 62 integration of the inverse distance function with respect to z (Guspí and Novara, 2009).

A common assumption for most of the equivalent-layer methods is that the harmonic function  $g_{ij}$ 63 (equation 2) is independent on the actual physical relationship between the observed potential field and 64 their true sources (e.g., Cordell, 1992; Guspí and Novara, 2009; Li et al., 2014). Hence,  $g_{ij}$  can be 65 defined according to the problem. The only condition imposed to this function is that it decays to zero 66 as the observation point  $(x_i, y_i, z_i)$  goes away from the position  $(x_j, y_j, z_j)$  of the j-th equivalent source. 67 However, several methods use a function  $g_{ij}$  that preserves the physical relationship between the observed 68 potential field and their true sources. For the case in which the observed potential field is gravity data,  $g_{ij}$ 69 is commonly defined as a component of the gravitational field produced at  $(x_i, y_i, z_i)$  by a point mass or 70 prism located at  $(x_j, y_j, z_j)$ , with unit density. On the other hand,  $g_{ij}$  is commonly defined as a component 71 of the magnetic induction field produced at  $(x_i, y_i, z_i)$  by a dipole or prism located at  $(x_i, y_i, z_i)$ , with unit magnetization intensity, when the observed potential field is magnetic data. 73

74 The main challenge in the equivalent-layer technique is the computational complexity associated with 75 handling large datasets. This complexity arises because the sensitivity matrix G (equation 3) is dense regardless of the harmonic function  $g_{ij}$  (equation 2) employed. In the case of scattered potential-field data, the structure of G is not well-defined, regardless of the spatial distribution of the equivalent sources. 77 However, in a specific scenario where (i) each potential-field datum is directly associated with a single 78 equivalent source located directly below it, and (ii) both the data and sources are based on planar and 79 regularly spaced grids, Takahashi et al. (2020, 2022) demonstrate that G exhibits a block-Toeplitz Toeplitz-80 block (BTTB) structure. In such cases, the product of G and an arbitrary vector can be efficiently computed 81 using a 2D fast Fourier transform as a discrete convolution.

# 83 1.3 General formulation

A general formulation for almost all equivalent-layer methods can be achieved by first considering that the  $P \times 1$  parameter vector  $\mathbf{p}$  (equation 3) can be reparameterized into a  $Q \times 1$  vector  $\mathbf{q}$  according to:

$$\mathbf{p} = \mathbf{H} \, \mathbf{q} \,, \tag{9}$$

where **H** is a  $P \times Q$  matrix. The predicted data vector **f** (equation 3) can then be rewritten as follows:

$$\mathbf{f} = \mathbf{G} \mathbf{H} \mathbf{q} . \tag{10}$$

Note that the original parameter vector  $\mathbf{p}$  is defined in a P-dimensional space whereas the reparameterized parameter vector  $\mathbf{q}$  (equation 9) lies in a Q-dimensional space. For convenience, we use the terms P-space and Q-space to designate them.

In this case, the problem of estimating a parameter vector  $\tilde{\mathbf{p}}$  minimizing a length measure of the difference between  $\mathbf{f}$  (equation 3) and  $\mathbf{d}$  is replaced by that of estimating an auxiliary vector  $\tilde{\mathbf{q}}$  minimizing the goal function

$$\Gamma(\mathbf{q}) = \Phi(\mathbf{q}) + \mu \Theta(\mathbf{q}), \qquad (11)$$

93 which is a combination of particular measures of length given by

$$\Phi(\mathbf{q}) = (\mathbf{d} - \mathbf{f})^{\top} \mathbf{W}_d (\mathbf{d} - \mathbf{f}) , \qquad (12)$$

94 and

$$\Theta(\mathbf{q}) = (\mathbf{q} - \bar{\mathbf{q}})^{\top} \mathbf{W}_q (\mathbf{q} - \bar{\mathbf{q}}) , \qquad (13)$$

- where the regularization parameter  $\mu$  is a positive scalar controlling the trade-off between the data-misfit
- 96 function  $\Phi(\mathbf{q})$  and the regularization function  $\Theta(\mathbf{q})$ ;  $\mathbf{W}_d$  is a  $D \times D$  symmetric matrix defining the relative
- 97 importance of each observed datum  $d_i$ ;  $\mathbf{W}_q$  is a  $Q \times Q$  symmetric matrix imposing prior information on  $\mathbf{q}$ ;
- 98 and  $\bar{\mathbf{q}}$  is a  $Q \times 1$  vector of reference values for  $\mathbf{q}$  that satisfies

$$\bar{\mathbf{p}} = \mathbf{H}\,\bar{\mathbf{q}}\,,\tag{14}$$

- 99 where  $\bar{\mathbf{p}}$  is a  $P \times 1$  vector containing reference values for the original parameter vector  $\mathbf{p}$ .
- After obtaining an estimate  $\tilde{\mathbf{q}}$  for the reparameterized parameter vector  $\mathbf{q}$  (equation 9), the estimate  $\tilde{\mathbf{p}}$  for the original parameter vector (equation 3) is computed by

$$\tilde{\mathbf{p}} = \mathbf{H}\,\tilde{\mathbf{q}}\,. \tag{15}$$

The reparameterized vector  $\tilde{\mathbf{q}}$  is obtained by first computing the gradient of  $\Gamma(\mathbf{q})$ ,

$$\nabla \Gamma(\mathbf{q}) = -2\mathbf{H}^{\mathsf{T}} \mathbf{G}^{\mathsf{T}} \mathbf{W}_d (\mathbf{d} - \mathbf{f}) + 2\mu \mathbf{W}_q (\mathbf{q} - \bar{\mathbf{q}}) . \tag{16}$$

103 Then, by considering that  $\nabla\Gamma(\tilde{\mathbf{q}}) = \mathbf{0}$  (equation 16), where  $\mathbf{0}$  is a vector of zeros, as well as adding and subtracting the term  $(\mathbf{H}^{\top}\mathbf{G}^{\top}\mathbf{W}_{d}\mathbf{G}\mathbf{H})\bar{\mathbf{q}}$ , we obtain

$$\tilde{\boldsymbol{\delta}}_q = \mathbf{B}\,\boldsymbol{\delta}_d\,,\tag{17}$$

105 where

$$\tilde{\boldsymbol{\delta}}_{q} = \tilde{\mathbf{q}} - \bar{\mathbf{q}} \,, \tag{18}$$

$$\delta_d = \mathbf{d} - \mathbf{G} \mathbf{H} \,\bar{\mathbf{q}} \,, \tag{19}$$

107

106

$$\mathbf{B} = \left(\mathbf{H}^{\top} \mathbf{G}^{\top} \mathbf{W}_{d} \mathbf{G} \mathbf{H} + \mu \mathbf{W}_{q}\right)^{-1} \mathbf{H}^{\top} \mathbf{G}^{\top} \mathbf{W}_{d},$$
(20)

108 or, equivalently (Menke, 2018, p. 62),

$$\mathbf{B} = \mathbf{W}_q^{-1} \mathbf{H}^{\top} \mathbf{G}^{\top} \left( \mathbf{G} \mathbf{H} \mathbf{W}_q^{-1} \mathbf{H}^{\top} \mathbf{G}^{\top} + \mu \mathbf{W}_d^{-1} \right)^{-1} . \tag{21}$$

- 109 Evidently, we have considered that all inverses exist in equations 20 and 21.
- The  $Q \times D$  matrix B defined by equation 20 is commonly used for the case in which D > Q, i.e., when
- there are more data than parameters (overdetermined problems). In this case, we consider that the estimate
- 112  $\tilde{q}$  is obtained by solving the following linear system for  $\delta_q$  (equation 18):

$$\left(\mathbf{H}^{\top}\mathbf{G}^{\top}\mathbf{W}_{d}\mathbf{G}\mathbf{H} + \mu\mathbf{W}_{q}\right)\tilde{\boldsymbol{\delta}}_{q} = \mathbf{H}^{\top}\mathbf{G}^{\top}\mathbf{W}_{d}\boldsymbol{\delta}_{d}.$$
(22)

- 113 On the other hand, for the cases in which D < Q (underdetermined problems), matrix B is usually defined
- 114 according to equation 21. In this case, the general approach involves estimating  $\tilde{q}$  in two steps. The first
- 115 consists in solving a linear system for a dummy vector, which is subsequently used to compute  $\tilde{q}$  by a

116 matrix-vector product as follows:

$$\left(\mathbf{G}\,\mathbf{H}\,\mathbf{W}_{q}^{-1}\,\mathbf{H}^{\top}\mathbf{G}^{\top} + \mu\mathbf{W}_{d}^{-1}\right)\mathbf{u} = \boldsymbol{\delta}_{d}$$

$$\tilde{\boldsymbol{\delta}}_{q} = \mathbf{W}_{q}^{-1}\,\mathbf{H}^{\top}\mathbf{G}^{\top}\mathbf{u}$$
(23)

- 117 where **u** is a dummy vector. After obtaining  $\tilde{\delta}_q$  (equations 22 and 23), the estimate  $\tilde{\mathbf{q}}$  is computed with
- 118 equation 18.
- 119 1.3.1 Optional normalization strategy
- Setting the regularization parameter  $\mu$  (equation 11) can be very difficult due to scale differences between
- 121 G and p (equation 3) or GH and q (equation 9. When faced with this scenario, a popular strategy (e.g., Li
- and Oldenburg, 2010; Soler and Uieda, 2021) involves creating the linear system (equations 22 and 23) by
- 123 substituting GH and q with

$$\mathbf{G}_n = \mathbf{G} \,\mathbf{H} \,\mathbf{N} \,, \quad \mathbf{q}_n = \mathbf{N}^{-1} \mathbf{q} \,, \tag{24}$$

- 124 and then finding the solution  $\tilde{\mathbf{q}}_n$  for the normalized parameter vector  $\mathbf{q}_n$ . The estimate  $\tilde{\mathbf{q}}$  for the
- reparameterized parameter vector q (equation 9) is subsequently obtained by removing the normalization
- 126 as follows:

$$\tilde{\mathbf{q}} = \mathbf{N}\,\tilde{\mathbf{q}}_n\,,\tag{25}$$

- where N is an invertible normalization matrix. This strategy usually constrains the practical range of the
- 128 regularization parameter  $\mu$  (equation 11).
- 129 1.3.2 Classical approach
- The classical approach in the equivalent-layer technique consists in using  $\mathbf{H} = \mathbf{I}_P$ , so that P = Q,  $\mathbf{p} = \mathbf{q}$
- 131 (equation 9),  $\bar{\mathbf{p}} = \bar{\mathbf{q}}$  (equation 14),  $\tilde{\mathbf{p}} = \tilde{\mathbf{q}}$  (equation 15),  $\mathbf{p}_n = \mathbf{q}_n$  (equation 24) and  $\tilde{\mathbf{p}}_n = \tilde{\mathbf{q}}_n$  (equation
- 132 25). In this case, the linear system (equations 22 and 23) is directly solved for

$$\tilde{\boldsymbol{\delta}}_p = \tilde{\mathbf{p}} - \bar{\mathbf{p}} \,, \tag{26}$$

instead of  $\tilde{\delta}_q$  (equation 18).

#### 2 NOTATION FOR SUBVECTORS AND SUBMATRICES

Here, we use a notation inspired on that presented by (Van Loan, 1992, p. 4) to represent subvectors and submatrices. Subvectors of d, for example, are specified by d[i], where i is a list of integer numbers that "pick out" the elements of d forming the subvector d[i]. For example, i = (1, 6, 4, 6) gives the subvector d[i] =  $[d_1 \ d_6 \ d_4 \ d_6]^{\top}$ . Note that the list i of indices may be sorted or not and it may also have repeated indices. For the particular case in which the list has a single element i = (i), then it can be used to extract the *i*-th element  $d_i \equiv d[i]$  of d. Regular lists can be represented by using the colon notation. For example,

$$\mathbf{i} = (3:8) \Leftrightarrow \mathbf{d}[\mathbf{i}] = [d_3 \ d_4 \ \dots \ d_8]^{\top}$$
 $\mathbf{i} = (:8) \Leftrightarrow \mathbf{d}[\mathbf{i}] = [d_1 \ d_2 \ \dots \ d_8]^{\top}$ ,
 $\mathbf{i} = (3:) \Leftrightarrow \mathbf{d}[\mathbf{i}] = [d_3 \ d_4 \ \dots \ d_D]^{\top}$ 

34 where D is the number of elements forming d.

The notation above can also be used to define submatrices of the  $D \times P$  matrix G. For example,  $\mathbf{i} = (2, 7, 4, 6)$  and  $\mathbf{j} = (1, 3, 8)$  lead to the submatrix

$$\mathbf{G[i,j]} = \begin{bmatrix} g_{21} & g_{23} & g_{28} \\ g_{71} & g_{73} & g_{78} \\ g_{41} & g_{43} & g_{48} \\ g_{61} & g_{63} & g_{68} \end{bmatrix}.$$

Note that, in this case, the lists i and j "pick out", respectively, the rows and columns of G that form the submatrix G[i, j]. The i-th row of G is given by the  $1 \times P$  vector G[i, j]. Similarly, the  $D \times 1$  vector G[i, j] represents the j-th column. Finally, we may use the colon notation to define the following submatrix:

$$\mathbf{i} = (2:5), \mathbf{j} = (3:7) \Leftrightarrow \mathbf{G}[\mathbf{i}, \mathbf{j}] = \begin{bmatrix} g_{23} & g_{24} & g_{25} & g_{26} & g_{27} \\ g_{33} & g_{34} & g_{35} & g_{36} & g_{37} \\ g_{43} & g_{44} & g_{45} & g_{46} & g_{47} \\ g_{53} & g_{54} & g_{55} & g_{56} & g_{57} \end{bmatrix}$$

which contains the contiguous elements of G from rows 2 to 5 and from columns 3 to 7.

# 3 COMPUTATIONAL STRATEGIES

- 136 Here, we review some strategies for reducing the computational cost of equivalent-layer technique.
- Typically, estimating a parameter vector  $\tilde{\mathbf{p}}$  or  $\tilde{\mathbf{q}}$  requires to solve a large-scale linear inversion (equations
- 138 22 and 23). This, in turn, means to deal with some obstacles concerning large computational cost:
- 139 (i) the large computer memory to store large and full matrices;
- 140 (ii) the long computation time to multiply a matrix by a vector; and
- 141 (iii) the long computation time to solve a large linear system of equations.
- 142 Two important factors affecting the efficiency of a given matrix algorithm are the storage and amount of
- 143 required arithmetic. Here, we quantify this last factor by counting *flops*, which are floating point additions,
- subtractions, multiplications or divisions (Golub and Van Loan, 2013, p. 12–14). We focus on the overall
- 145 strategies used by the selected methods.

#### 146 3.1 Moving window

- 147 The initial approach to enhance the computational efficiency of the equivalent-layer technique is
- 148 commonly denoted moving window and involves first splitting the observed data  $d_i$ ,  $i \in \{1 : D\}$ , into
- 149 M overlapping subsets (or data windows) formed by  $D^m$  data each,  $m \in \{1: M\}$ . The data inside the
- 150 m-th window are usually adjacent to each other and have indices defined by an integer list  $i^m$  having
- 151  $D^m$  elements. The number of data  $D^m$  forming the data windows are not necessarily equal to each other.
- 152 Each data window has a  $D^m \times 1$  observed data vector  $\mathbf{d}^m \equiv \mathbf{d}[\mathbf{i}^m]$ . The second step consists in defining
- 153 a set of P equivalent sources with scalar physical property  $p_i$ ,  $j \in \{1 : P\}$ , and also split them into M
- overlapping subsets (or source windows) formed by  $P^m$  data each,  $m \in \{1: M\}$ . The sources inside the
- 155 m-th window have indices defined by an integer list  $j^m$  having  $P^m$  elements. Each source window has a
- with white white state of the larger has a state of the s
- 156  $P^m \times 1$  parameter vector  $\mathbf{p}^m$  and is located right below the corresponding m-th data window. Then, each

157 
$$\mathbf{d}^m \equiv \mathbf{d}[\mathbf{i}^m]$$
 is approximated by

$$\mathbf{f}^m = \mathbf{G}^m \mathbf{p}^m \,, \tag{27}$$

where  $G^m \equiv G[i^m, j^m]$  is a submatrix of G (equation 3) formed by the elements computed with equation 2 using only the data and equivalent sources located inside the window m-th. The main idea of the moving-window approach is using the  $\tilde{\mathbf{p}}^m$  estimated for each window to obtain (i) an estimate  $\tilde{\mathbf{p}}$  of the parameter vector for the entire equivalent layer or (ii) a given potential-field transformation  $\mathbf{t}$  (equation 4). The main advantages of this approach is that (i) the estimated parameter vector  $\tilde{\mathbf{p}}$  or transformed potential field are not obtained by solving the full, but smaller linear systems and (ii) the full matrix G (equation 3) is never stored.

Leão and Silva (1989) presented a pioneer work using the moving-window approach. Their method 165 requires a regularly-spaced grid of observed data on a horizontal plane  $z_0$ . The data windows are defined by 166 square local grids of  $\sqrt{D'} \times \sqrt{D'}$  adjacent points, all of them having the same number of points D'. The 167 equivalent sources in the m-th data window are located below the observation plane, at a constant vertical 168 distance  $\Delta z_0$ . They are arranged on a regular grid of  $\sqrt{P'} \times \sqrt{P'}$  adjacent points following the same 169 grid pattern of the observed data. The local grid of sources for all data windows have the same number 170 of elements P'. Besides, they are vertically aligned, but expands the limits of their corresponding data 171 windows, so that D' < P'. Because of this spatial configuration of observed data and equivalent sources, 172 we have that  $G^m = G'$  (equation 27) for all data windows (i.e.,  $\forall m \in \{1 : M\}$ ), where G' is a  $D' \times P'$ 173 constant matrix. 174

By omitting the normalization strategy used by Leão and Silva (1989), their method consists in directly computing the transformed potential field  $t_c^m$  at the central point  $(x_c^m, y_c^m, z_0 + \Delta z_0)$  of each data window as follows:

$$t_c^m = \left(\mathbf{G}'\mathbf{a}'\right)^{\top} \left[\mathbf{G}' \left(\mathbf{G}'\right)^{\top} + \mu \mathbf{I}_{D'}\right]^{-1} \mathbf{d}^m, \quad m \in \{1:M\},$$
(28)

where  $I_{D'}$  is the identity matrix of order D' and a' is a  $P' \times 1$  vector with elements computed by equation 178 5 by using all equivalent sources in the m-th subset and only the coordinate of the central point in the m-th 179 data window. Due to the presumed spatial configuration of the observed data and equivalent sources, a' and 180 G' are the same for all data windows. Hence, only the data vector  $d^m$  is modified according to the position 181 of the data window. Note that equation 28 combines the potential-field transformation (equation 4) with 182 the solution of the undetermined problem (equation 23) for the particular case in which  $\mathbf{H} = \mathbf{W}_q = \mathbf{I}_{P'}$ 183 (equations 9 and 13),  $W_d = I_{D'}$  (equation 12),  $\bar{p} = 0$  (equation 14), where  $I_{P'}$  and  $I_{D'}$  are identity 184 matrices of order P' and D', respectively, and 0 is a vector of zeros. 185

The method proposed by Leão and Silva (1989) can be outlined by the Algorithm 1. Note that Leão and Silva (1989) directly compute the transformed potential  $t_c^m$  at the central point of each data window without explicitly computing and storing an estimated for  $\mathbf{p}^m$  (equation 27). It means that their method allows computing a single potential-field transformation. A different transformation or the same one evaluated at different points require running their moving-data window method again.

Soler and Uieda (2021) generalized the method proposed by Leão and Silva (1989) for irregularly spaced data on an undulating surface. A direct consequence of this generalization is that a different submatrix  $G^m \equiv G[i^m, j^m]$  (equation 27) must be computed for each window. Differently from Leão and Silva (1989), Soler and Uieda (2021) store the computed  $\tilde{p}^m$  for all windows and subsequently use them to obtain a desired potential-field transformation (equation 4) as the superposed effect of all windows. The estimated  $\tilde{p}^m$  for all windows are combined to form a single  $P \times 1$  vector  $\tilde{p}$ , which is an estimate for original

# Algorithm 1: Generic pseudo-code for the method proposed by Leão and Silva (1989).

```
Initialization:

1 Set the indices \mathbf{i}^m for each data window, m \in \{1:M\};

2 Set the indices \mathbf{j}^m for each source window, m \in \{1:M\};

3 Set the constant depth z_0 + \Delta z_0 for all equivalent sources;

4 Compute the vector \mathbf{a}' associated with the desired potential-field transformation;

5 Compute the matrix \mathbf{G}';

6 Compute (\mathbf{G}'\mathbf{a}')^{\top} \left[\mathbf{G}' (\mathbf{G}')^{\top} + \mu \mathbf{I}_{D'}\right]^{-1};

7 m = 1;

8 while m < M do

9 | Compute t_c^m (equation 28);

10 | m \leftarrow m + 1;

11 end
```

parameter vector  $\mathbf{p}$  (equation 3). For each data window, Soler and Uieda (2021) solve an overdetermined problem (equation 22) for  $\tilde{\mathbf{p}}^m$  by using  $\mathbf{H} = \mathbf{W}_q = \mathbf{I}_{P^m}$  (equations 9 and 13),  $\mathbf{W}_d^m$  (equation 12) equal to a diagonal matrix of weights for the data inside the m-th window and  $\bar{p} = \mathbf{0}$  (equation 14), so that

$$\left[ \left( \mathbf{G}^m \right)^{\top} \mathbf{W}_d^m \mathbf{G}^m + \mu \mathbf{I}_{P'} \right] \tilde{\mathbf{p}}^m = \left( \mathbf{G}^m \right)^{\top} \mathbf{W}_d^m \mathbf{d}^m.$$
 (29)

It is important to stress here that Soler and Uieda (2021) used the normalization strategy described in section 1.3.1, but we have conveniently omitted here. The overall steps of their method are defined by the Algorithm 2. Note that this algorithm starts with a residuals vector **r** that is iteratively updated. At each iteration, the potential field predicted a source window is computed at all observation points and removed from the residuals vector **r**.

# Algorithm 2: Generic pseudo-code for the method proposed by Soler and Uieda (2021).

```
Initialization:
 1 Set the indices i^m for each data window, m \in \{1 : M\};
2 Set the indices \mathbf{j}^m for each source window, m \in \{1:M\};
   Set the depth of all equivalent sources;
 4 Set a D \times 1 residuals vector \mathbf{r} = \mathbf{d};
 5 Set a P \times 1 vector \tilde{\mathbf{p}} = \mathbf{0};
 6 m = 1:
    while m < M do
 7
          Set the matrix \mathbf{W}_d^m;
 8
          Compute the matrix G^m;
          Compute \tilde{\mathbf{p}}^m (equation 29);
10
          \tilde{\mathbf{p}}[\mathbf{j}^m] \leftarrow \tilde{\mathbf{p}}[\mathbf{j}^m] + \tilde{\mathbf{p}}^m;
11
          \mathbf{r} \leftarrow \mathbf{r} - \mathbf{G}[:, \mathbf{j}^m] \, \tilde{\mathbf{p}}^m \; ;
12
          m \leftarrow m + 1;
13
14 end
```

#### 205 3.2 Column update

Cordell (1992) proposed a computational strategy that was later used by Guspí and Novara (2009) and relies on first defining one equivalent source located right below each observed data  $d_i$ ,  $i \in \{1 : D\}$ , at

a vertical coordinate  $z_i + \Delta z_i$ , where  $\Delta z_i$  is proportional to the distance from the *i*-th observation point 209  $(x_i, y_i, z_i)$  to its closest neighbor. The second step consists in updating the physical property  $p_i$  of a given equivalent source,  $j \in \{1 : D\}$  and remove its predicted potential field from the observed data vector d, 210 producing a residuals vector r. Then, the same procedure is repeated for other sources with the purpose of 211 iteratively updating r and the  $D \times 1$  parameter vector p containing the physical property of all equivalent 212 sources. At the end, the algorithm produces an estimate  $\tilde{p}$  for the parameter vector yielding a predicted 213 potential field f (equation 3) satisfactorily fitting the observed data d according to a given criterion. Note 214 that the method proposed by Cordell (1992) iteratively solves the linear  $G\tilde{p} \approx d$  with a  $D \times D$  matrix G. 215 At each iteration, only a single column of G (equation 3) is used. An advantage of this *column-update* 216 approach is that the full matrix G is never stored. 217

Algorithm 3 delineates the Cordell's method. Note that a single column  $G[:, i_{max}]$  of the  $D \times D$  matrix G (equation 3) is used per iteration, where  $i_{max}$  is the index of the maximum absolute value in r. As pointed out by Cordell (1992), the method does not necessarily decrease monotonically along the iterations. Besides, the method may not converge depending on how the vertical distances  $\Delta z_i$ ,  $i \in \{1:D\}$ , controlling the depths of the equivalent sources are set. According to Cordell (1992), the maximum absolute value  $r_{max}$  in r decreases robustly at the beginning and oscillates within a narrowing envelope for the subsequent iterations.

# Algorithm 3: Generic pseudo-code for the method proposed by Cordell (1992).

```
Initialization:
```

```
1 Compute a D \times 1 vector \Delta z whose i-th element \Delta z_i is a vertical distance controlling the depth of
      the i-th equivalent source, i \in \{1 : D\};
 2 Set a tolerance \epsilon;
    Set a maximum number of iteration ITMAX;
 4 Set a D \times 1 residuals vector \mathbf{r} = \mathbf{d};
 5 Set a D \times 1 vector \tilde{\mathbf{p}} = \mathbf{0};
 6 Define the maximum absolute value r_{\text{max}} in \mathbf{r};
 7 \ m = 1;
    while (r_{\max} > \epsilon) and (m < \text{ITMAX}) do
          Define the coordinates (x_{\text{max}}, y_{\text{max}}, z_{\text{max}}) and index i_{\text{max}} of the observation point associated with
          \tilde{\mathbf{p}}[i_{\max}] \leftarrow \tilde{\mathbf{p}}[i_{\max}] + (a_{\max} \Delta \mathbf{z}[i_{\max}]);
10
          \mathbf{r} \leftarrow \mathbf{r} - (\mathbf{G}[:, i_{\text{max}}] \, \tilde{\mathbf{p}}[i_{\text{max}}]) ;
11
          Define the new r_{\text{max}} in r;
12
13
         m \leftarrow m + 1;
14 end
```

#### 3.3 Row update

225

226

227228

229

230

231

232

Mendonça and Silva (1994) proposes an algebraic reconstruction technique (ART) (e.g., van der Sluis and van der Vorst, 1987, p. 58) to estimate a parameter vector  $\tilde{\mathbf{p}}$  for a regular grid of P equivalent sources on a horizontal plane  $z_0$ . Such methods iterate on the linear system rows to estimate corrections for the parameter vector, which may substantially save computer time and memory required to compute and store the full linear system matrix along the iterations. The convergence of such *row-update methods* depends on the linear system condition. The main advantage of such methods is not computing and storing the full linear system matrix, but iteratively using its rows. The particular ART method proposed by Mendonça and

Silva (1994) considers that 233

$$\mathbf{d} = \begin{bmatrix} \mathbf{d}_e \\ \mathbf{d}_r \end{bmatrix} , \quad \mathbf{G} = \begin{bmatrix} \mathbf{G}_e \\ \mathbf{R}_r \end{bmatrix} , \tag{30}$$

where  $\mathbf{d}_e$  and  $\mathbf{d}_r$  are  $D_e \times 1$  and  $D_r \times 1$  vectors and  $\mathbf{G}_e$  and  $\mathbf{G}_r$  are  $D_e \times P$  and  $D_r \times P$  matrices, 234 235 respectively. Mendonça and Silva (1994) designate  $d_e$  and  $d_r$  as, respectively, equivalent and redundant data. With the exception of a normalization strategy, Mendonça and Silva (1994) calculate a  $P \times 1$  estimated 236 parameter vector  $\tilde{\mathbf{p}}$  by solving an underdetermined problem (equation 23) involving only the equivalent 237 238

data  $\mathbf{d}_e$  (equation 30) for the particular case in which  $\mathbf{H} = \mathbf{W}_p = \mathbf{I}_P$  (equations 9 and 13),  $\mathbf{W}_d = \mathbf{I}_{D_e}$ 

(equation 12) and  $\bar{p} = 0$  (equation 14), which results in 239

$$(\mathbf{F} + \mu \, \mathbf{I}_{D_e}) \, \mathbf{u} = \mathbf{d}_e \tilde{\mathbf{p}} = \mathbf{G}_e^{\top} \mathbf{u} ,$$
(31)

where  $\mathbf{F}$  is a  $P \times P$  matrix that replaces  $\mathbf{G}_e \mathbf{G}_e^{\top}$ . Mendonça and Silva (1994) presume that the estimated parameter vector  $\tilde{\mathbf{p}}$  obtained from equation 31 leads to a  $D_r \times 1$  residuals vector

$$\mathbf{r} = \mathbf{d}_r - \mathbf{G}_r \tilde{\mathbf{p}} \tag{32}$$

having a maximum absolute value  $r_{\max} \leq \epsilon$ , where  $\epsilon$  is a predefined tolerance. 242

The overall method of Mendonça and Silva (1994) is defined by Algorithm 4. It is important noting 243 that the number  $D_e$  of equivalent data in  $\mathbf{d}_e$  increases by one per iteration, which means that the order 244 of the linear system in equation 31 also increases by one at each iteration. Those authors also propose 245 a computational strategy based on Cholesky factorization (e.g., Golub and Van Loan, 2013, p. 163) for 246 efficiently updating  $(\mathbf{F} + \mu \mathbf{I}_{D_e})$  at a given iteration (line 16 in Algorithm 4) by computing only its new 247 elements with respect to those computed in the previous iteration. 248

#### 3.4 Reparameterization

249

250

251

252

253

254

255

256

257

258

259

260

261

262

263

264 265

266

Another approach for improving the computational performance of equivalent-layer technique consists in setting a  $P \times Q$  reparameterization matrix **H** (equation 9) with  $Q \ll P$ . This strategy has been used in applied geophysics for decades (e.g., Skilling and Bryan, 1984; Kennett et al., 1988; Oldenburg et al., 1993; Barbosa et al., 1997) and is known as *subspace method*. The main idea relies in reducing the linear system dimension from the original P-space to a lower-dimensional subspace (the Q-space). An estimate  $\tilde{\mathbf{q}}$  for the reparameterized parameter vector  $\mathbf{q}$  is obtained in the Q-space and subsequently used to obtain an estimate  $\tilde{p}$  for the parameter vector p (equation 3) in the P-space by using equation 9. Hence, the key aspect of this reparameterization approach is solving an appreciably smaller linear inverse problem for  $\tilde{\mathbf{q}}$ than that for the original parameter vector  $\tilde{\mathbf{p}}$  (equation 3).

Oliveira Jr. et al. (2013) have used this approach to describe the physical property distribution on the equivalent layer in terms of piecewise bivariate polynomials. Specifically, their method consists in splitting a regular grid of equivalent sources into source windows inside which the physical-property distribution is described by bivariate polynomial functions. The key aspect of their method relies on the fact that the total number of coefficients required to define the bivariate polynomials is considerably smaller than the original number of equivalent sources. Hence, they formulate a linear inverse problem for estimating the polynomial coefficients and use them later to compute the physical property distribution on the equivalent layer.

280

281

282

283

284

285

286

Algorithm 4: Generic pseudo-code for the method proposed by Mendonça and Silva (1994).

```
Initialization:
 1 Set a regular grid of P equivalent sources at a horizontal plane z_0;
 2 Set a tolerance \epsilon;
    Set a D \times 1 residuals vector \mathbf{r} = \mathbf{d};
 4 Define the maximum absolute value r_{max} in r;
 5 Define the index i_{max} of r_{max};
6 Define the list of indices i_r of the remaining data in r;
    Define \mathbf{d}_e = \mathbf{d}[i_{\text{max}}];
    Compute (\mathbf{F} + \mu \mathbf{I}_{D_e}) and \mathbf{G}_e;
    Compute \tilde{\mathbf{p}} (equation 31);
10 Compute \mathbf{r} = \mathbf{d}[\mathbf{i}_r] - \mathbf{G}[\mathbf{i}_r,:] \tilde{\mathbf{p}};
11 Define the maximum absolute value r_{\text{max}} in r;
    13
          Define the list of indices i_r of the remaining elements in r;
14
         \mathbf{d}_e \leftarrow egin{bmatrix} \mathbf{d}_e \ \mathbf{d}[i_{\mathtt{max}}] \end{bmatrix}
15
          Update (\mathbf{F} + \mu \mathbf{I}_{D_e}) and \mathbf{G}_e;
16
          Update p (equation 31);
17
          Update \mathbf{r} = \mathbf{d}[\mathbf{i}_r] - \mathbf{G}[\mathbf{i}_r,:] \tilde{\mathbf{p}};
18
          Define the maximum absolute value r_{max} in r;
19
20 end
```

The method proposed by Oliveira Jr. et al. (2013) consists in solving an overdetermined problem (equation 22) for estimating the polynomial coefficients  $\tilde{\bf q}$  with  ${\bf W}_d={\bf I}_D$  (equation 12) and  $\bar{q}={\bf 0}$  (equation 14), so that

$$\left(\mathbf{H}^{\top}\mathbf{G}^{\top}\mathbf{G}\,\mathbf{H} + \mu\,\mathbf{W}_{q}\right)\tilde{\mathbf{q}} = \mathbf{H}^{\top}\mathbf{G}^{\top}\,\mathbf{d}\,,\tag{33}$$

where  $\mathbf{W}_q = \mathbf{H}^{\top} \mathbf{W}_p \mathbf{H}$  is defined by a matrix  $\mathbf{W}_p$  representing the zeroth- and first-order Tikhonov regularization (e.g., Aster et al., 2019, p. 103). Note that, in this case, the prior information is defined in the P-space for the original parameter vector  $\mathbf{p}$  and then transformed to the Q-space. Another characteristic of their method is that it is valid for processing irregularly-spaced data on an undulating surface.

Mendonça (2020) also proposed a reparameterization approach for the equivalent-layer technique. Their approach, however, consists in setting  $\bf H$  as a truncated singular value decomposition (SVD) (e.g., Aster et al., 2019, p. 55) of the observed potential field. Differently from Oliveira Jr. et al. (2013), however, the method of Mendonça (2020) requires a regular grid of potential-field data on horizontal plane. Another difference is that these authors uses  ${\bf W}_q = {\bf I}_Q$  (equation 13), which means that the regularization is defined directly in the Q-space.

Before Oliveira Jr. et al. (2013) and Mendonça (2020), Barnes and Lumley (2011) also proposed a computationally efficient method for equivalent-layer technique based on reparameterization. A key difference, however, is that Barnes and Lumley (2011) did not set a  $P \times Q$  reparameterization matrix **H** (equation 9) with Q << P. Instead, they used a matrix **H** with  $Q \approx 1.7 P$ . Their central idea is setting a reparameterization scheme that groups distant equivalent sources into blocks by using a bisection process. This scheme leads to a quadtree representation of the physical-property distribution on the equivalent layer, so that matrix **GH** (equation 10) is notably sparse. Barnes and Lumley (2011) explore this sparsity in

solving the overdetermined problem for  $\tilde{\bf q}$  (equation 33) via conjugate-gradient method (e.g., Golub and Van Loan, 2013, sec. 11.3).

# 3.5 Wavelet compression

290 Previously to Barnes and Lumley (2011), the idea of transforming the dense matrix G (equation 3) into a sparse one has already been used in the context of equivalent-layer technique. Li and Oldenburg (2010) 291 292 proposed a method that applies the discrete wavelet transform to introduce sparsity into the original dense 293 matrix G. Those authors approximate a planar grid of potential-field data by a regularly-spaced grid of equivalent sources, so that the number of data D and sources P is the same, i.e., D = P. Specifically, Li 294 295 and Oldenburg (2010) proposed a method that applies the wavelet transform to the original dense matrix G296 and sets to zero the small coefficients that are below a given threshold, which results in an approximating sparse representation of G in the wavelet domain. They first consider the following approximation 297

$$\mathbf{d}_w \approx \mathbf{G}_s \, \mathbf{p}_w \,, \tag{34}$$

298 where

289

$$\mathbf{d}_w = \mathbf{W} \, \mathbf{d} \,, \quad \mathbf{p}_w = \mathbf{W} \, \mathbf{p} \,, \tag{35}$$

are the observed data and parameter vector in the wavelet domain; W is a  $D \times D$  orthogonal matrix defining a discrete wavelet transform; and  $G_s$  is a sparse matrix obtained by setting to zero the elements of

$$G_w = \mathcal{W} G \mathcal{W} \tag{36}$$

301 with absolute value smaller than a given threshold.

Li and Oldenburg (2010) apply the normalization strategy defined in section 1.3.1 to equation 34 by using a diagonal normalization matrix N (equation 24) and then formulate an overdetermined problem (equation 304 22) with  $H = I_P$  (equations 9),  $\mu = 0$  (equation 11),  $W_d = I_D$  (equation 12) and  $\bar{p} = 0$  (equation 14) so

305 that

$$\left(\mathbf{G}_n^{\top}\mathbf{G}_n\right)\tilde{\mathbf{p}}_n = \mathbf{G}_n^{\top}\mathbf{d}_w \,, \tag{37}$$

306 where  $G_n$  and  $\tilde{p}_n$  are defined according to equation 24 in terms of the sparse matrix  $G_s$  (equation 34).

307 They solve this linear system (equation 37) with the conjugate-gradient method (e.g., Golub and Van Loan,

308 2013, sec. 11.3) and use it to obtain an estimate  $\tilde{p}$  for the parameter vector given by

$$\tilde{\mathbf{p}} = \boldsymbol{\mathcal{W}}^{\top} \left( \mathbf{N} \, \tilde{\mathbf{p}}_n \right) \,, \tag{38}$$

309 where the term within parentheses removes the normalization (equation 25) and matrix  $\boldsymbol{\mathcal{W}}^{\top}$  applies an 310 inverse wavelet transform.

#### 311 3.6 Iterative methods using the full matrix G

In the context of equivalent-layer technique, there are methods (e.g., Li and Oldenburg, 2010; Barnes and Lumley, 2011) that modify the original linear system and then iteratively solve this modified system by using the well-established conjugate gradient method (e.g., Golub and Van Loan, 2013, sec. 11.3). Xia and Sprowl (1991) also proposed an iterative method for estimating a parameter vector  $\tilde{\mathbf{p}}$  (equation 3). Their method, however, uses the original matrix  $\mathbf{G}$  without previously computing a compression or reparameterization, for example. More than two decades later, Siqueira et al. (2017) have deduced

322

323

324

325 326

327

328 329

330

332 333

334

335

336

337 338

339

340

341

essentially the same method presented by Xia and Sprowl (1991), but by following a more theoretical approach based on the Gauss' theorem (e.g., Kellogg, 1967, p. 43) and the total excess of mass (e.g., Blakely, 1996, p. 60). Besides, Siqueira et al. (2017) have shown that this method produces very stable

321 solutions, even for noise-corrupted potential-field data.

The iterative method proposed by Xia and Sprowl (1991) and Siqueira et al. (2017) is outlined in Algorithm 5, presumes an equivalent layer formed by monopoles (point masses) and can be applied to irregularly-spaced data on an undulating surface. Note that the residuals r are used to compute a correction Δp for the parameter vector at each iteration (line 11), which requires a matrix-vector product involving the full matrix G. Interestingly, this approach for estimating the physical property distribution on an equivalent layer is the same originally proposed by Bott (1960) for estimating the basement relief under sedimentary basins. The method of Xia and Sprowl (1991) and Siqueira et al. (2017) was originally proposed for processing gravity data, but can be potentially applied to any harmonic function because it actually represents an iterative solution of the classical *Dirichlet's problem* or the *first boundary value problem of potential theory* (Kellogg, 1967, p. 236) on a plane.

Recently, Jirigalatu and Ebbing (2019) presented another iterative method for estimating a parameter vector  $\tilde{\mathbf{p}}$  (equation 3). With the purpose of combining different potential-field data, their method basically modifies that shown in Algorithm 5 by changing the initial approximation and the iterative correction for the parameter vector. Specifically, Jirigalatu and Ebbing (2019) replace line 3 by  $\tilde{\mathbf{p}} = \mathbf{0}$ , where  $\mathbf{0}$  is a vector of zeros, and line 8 by  $\Delta \mathbf{p} = \omega \mathbf{G}^{\top} \mathbf{r}$ , where  $\omega$  is a positive scalar defined by trial and error. Note that this modified approach requires two matrix-vector products involving the full matrix  $\mathbf{G}$  per iteration. To overcome the high computational cost of these two products, Jirigalatu and Ebbing (2019) set an equivalent layer formed by prisms and compute their predicted potential field in the wavenumber domain by using the Gauss-FFT technique Zhao et al. (2018).

INCLUIMOS Xia et al. (1993) - FOURIER - OU NÃO?

**Algorithm 5:** Generic pseudo-code for the iterative methods proposed by Xia and Sprowl (1991) and Siqueira et al. (2017). The symbol " $\circ$ " denotes the entrywise or Hadamard product (e.g., Horn and Johnson, 1991, p. 298) and s is a  $P \times 1$  vector whose j-th element is a predefined element of area centered at the j-th equivalent source.

```
Initialization:
     Set P equivalent sources on a horizontal plane z_0;
     Set the auxiliary vector s;
     Set a tolerance \epsilon;
 4 Compute \tilde{\mathbf{p}} = \mathbf{s} \circ \mathbf{d};
 5 Compute G;
    Compute \mathbf{r} = \mathbf{d} - \mathbf{G}\,\tilde{\mathbf{p}};
     Compute \|\mathbf{r}\|;
     while (\|\mathbf{r}\| > \epsilon) do
            \Delta \mathbf{p} = \mathbf{s} \circ \mathbf{r};
            \tilde{\mathbf{p}} \leftarrow \tilde{\mathbf{p}} + \Delta \mathbf{p};
10
            \mathbf{r} \leftarrow \mathbf{r} - \mathbf{G} \, \mathbf{\Delta} \mathbf{p};
11
            Compute \|\mathbf{r}\|;
12
13 end
```

#### 342 3.7 Discrete convolution

- Recently, Takahashi et al. (2020, 2022) have shown that G assumes a doubly block Toeplitz (Jain, 1989,
- p. 28) or block-Toeplitz Toeplitz-block (BTTB) (Chan and Jin, 2007, p. 67) structure for the particular case
- 345 in which (i) there is a single equivalent source right below each potential-field datum and (ii) both data and
- 346 sources rely on planar and regularly spaced grids. This structure allows formulating the product of G and
- an arbitrary vector as a fast discrete convolution (Van Loan, 1992, section 4.2).
- Takahashi et al. (2020, 2022) consider a regular grid of D potential-field data at points  $(x_i, y_i, z_0)$ ,
- 349  $i \in \{1:D\}$ , on a horizontal plane  $z_0$ . The data indices i may be ordered along the x- or y-direction,
- 350 which results in an x- or y-oriented grid, respectively. They also consider a single equivalent source
- 351 located right below each datum, at a constant vertical coordinate  $z_0 + \Delta z$ ,  $\Delta z > 0$ . In this case, the number
- 352 of data and equivalent sources are equal to each other (i.e., D = P) and G (equation 3) is a  $D \times D$  matrix
- 353 formed by  $D_B \times D_B$  blocks, where each block has  $D_b \times D_b$  elements, with  $D = D_B D_b$ .
- Let  $\mathcal{D}$  be a  $2D_B \times 2D_b$  matrix formed by four blocks of  $D_B \times D_b$  elements. All blocks have null
- 355 elements, except the first block in the first quadrant, which has the observed data  $d_i$ ,  $i \in \{1 : D\}$  arranged
- 356 along its rows or columns if the data grid has a x- or y- oriented pattern, respectively. Let us also consider
- 357 a  $2D_B \times 2D_b$  matrix  $\mathcal{P}$  having the same structure of  $\mathcal{D}$ , but containing the elements  $p_j$  of the parameter
- 358 vector  $\mathbf{p}, j \in \{1 : P\}.$
- 359 The first block (in the first quadrant) contains
- 360 The conjugate gradient (CG) is a very popular iterative method for solving linear systems. This method
- 361 was originally developed to solve systems having a square and positive definite matrix. There are two
- 362 adapted versions of the CG method. The first is called *conjugate gradient normal equation residual* (CGNR)
- 363 Golub and Van Loan (2013, sec. 11.3) or conjugate gradient least squares (CGLS) (Aster et al., 2019, p.
- 364 165) and is used to solve overdetermined problem (equation 22). The second is called *conjugate gradient*
- 365 normal equation error (CGNE) method Golub and Van Loan (2013, sec. 11.3) and is used to solve the
- 366 underdetermined problem (equation 23).
- 367 Takahashi et al. (2020)
- 368 Takahashi et al. (2022)

#### 4 TEXTO ANTIGO

- 369 4.0.1 The wavelet compression and lower-dimensional subspace
- For large data sets, the sensitivity matrix A (equation 3) is a drawback in applying the equivalent-layer
- 371 technique because it is a large and dense matrix.
- ? transformed a large and full sensitivity matrix into a sparse one by using fast wavelet transforms. In
- 373 the wavelet domain, ? applyied a 2D wavelet transform to each row and column of the original sensitivity
- and matrix A to expand it in the wavelet bases. This operation can be done by premultiplying the original
- 375 sensitivity matrix A by a matrix representing the 2D wavelet transform  $W_2$  and then the resulting is
- 376 postmultiplied by the transpose of  $W_2$  (i.e.,  $W_2^{\perp}$ ).

$$\tilde{\mathbf{A}} = \mathbf{W_2} \, \mathbf{A} \, \mathbf{W_2}^{\top} \,, \tag{39}$$

where  $\tilde{\bf A}$  is the expanded original sensitivity matrix in the wavelet bases with many elements zero or close to zero. Next, the matrix  $\tilde{\mathbf{A}}$  is replaced by its sparse version  $\tilde{\mathbf{A}}_{\mathbf{s}}$  in the wavelet domain which in turn 378 is obtained by retaining only the large elements of the A. Thus, the elements of A whose amplitudes 379 fall below a relative threshold are discarded. In ?, the original sensitivity matrix A is high compressed 380 resulting in a sparce matrix As with a few percent of nonzero elements and the the inverse problem is 381 382 solved in the wavelet domain by using  $A_s$  and a incomplete conjugate gradient least squares, without an explicit regularization parameter and a limited number of iterations. The solution is obtained by solving the 383 following linear system 384

$$\tilde{\mathbf{A}}_{\mathbf{L}}^{\top} \, \tilde{\mathbf{A}}_{\mathbf{L}} \, \tilde{\mathbf{p}}_{\mathbf{L}}^{*} = \tilde{\mathbf{A}}_{\mathbf{L}}^{\top} \, \tilde{\mathbf{d}}^{o} \,, \tag{40}$$

where  $\tilde{\mathbf{p}}_{\mathbf{L}}^*$  is obtained by solving the linear system given by equation 40,

$$\tilde{\mathbf{A}}_{\mathbf{L}} = \tilde{\mathbf{A}}_{\mathbf{s}} \tilde{\mathbf{L}}^{-1}, \tag{41a}$$

$$\tilde{\mathbf{p}}_{\mathbf{L}} = \tilde{\mathbf{L}}\tilde{\mathbf{p}}, \tag{41b}$$

$$\tilde{\mathbf{d}}^{o} = \mathbf{W}_{2} \mathbf{d}^{o}, \tag{41c}$$

$$\tilde{\mathbf{d}}^o = \mathbf{W_2} \, \mathbf{d}^o, \tag{41c}$$

where  $\hat{\mathbf{L}}$  is a diagonal and invertible weighting matrix representing the finite-difference approximation in 386 the wavelet domain. Finally, the distribution over the equivalent layer in the space domain p is obtained by 387 applying an inverse wavelet transform in two steps, i.e.: 388

$$\tilde{\mathbf{p}} = \tilde{\mathbf{L}}^{-1} \, \tilde{\mathbf{p}}_{\mathbf{L}}^* \,, \tag{42}$$

389 and

393

394

395

396

397

$$\mathbf{p} = \mathbf{W_2} \, \tilde{\mathbf{p}} \,. \tag{43}$$

Although the data misfit quantifying the difference between the observed and predicted data by the 390 equivalent source is calculated in the wavelet domain, we understand that the desired transformation is 391 calculated via equation ?? which uses a full matrix of Green's functions T. 392

? used the equivalent-layer technique with a wavelet compression to perform an upward continuation of total-field anomaly between uneven surfaces. For regularly spaced grid of data, ? reported that high compression ratios are achived with insignificant loss of accuracy. As compared to the upward-continued total-field anomaly by equivalent layer using the dense matrix, ?'s (?) approach, using the Daubechies wavelet, decreased CPU (central processing unit) time by up to two orders of magnitude.

398 ? overcame the solution of intractable large-scale equivalent-layer problem by using the subspace method (e.g., Skilling and Bryan, 1984; Kennett et al., 1988; Oldenburg et al., 1993; Barbosa et al., 1997). The 399 subspace method reduces the dimension of the linear system of equations to be solved. Given a higher-400 dimensional space (e.g., M-dimensional model space,  $\mathbb{R}^M$ ), there exists many lower-dimensional subspaces 401 (e.g., Q-dimensional subspace) of  $\mathbb{R}^M$ . The linear inverse problem related to the equivalent-layer technique 402 consists in finding an M-dimension parameter vector  $\mathbf{p} \in \mathbb{R}^M$  which adequately fits the potential-field 403 data. The subspace method looks for a parameter vector who lies in a Q-dimensional subspace of  $\mathbb{R}^M$ 404 which, in turn, is spanned by a set of Q vectors  $\mathbf{v}_i = 1,...,Q$ , where  $\mathbf{v}_i \in \mathbb{R}^M$  In matrix notation, the 405 parameter vector in the subspace method can be written as

$$p = V \alpha$$
, (44)

407 where V is an  $M \times Q$  matrix whose columns  $\mathbf{v}_i = 1, ..., Q$  form a basis vectors for a subspace Q of  $\mathbb{R}^M$ .

- 408 In equation 44, the parameter vector **p** is defined as a linear combination in the space spanned by Q basis
- 409 vectors  $\mathbf{v}_i = 1, ..., Q$  and  $\boldsymbol{\alpha}$  is a Q-dimensional unknown vector to be determined. The main advantage of
- 410 the subspace method is that the linear system of M equations in M unknowns to be originally solved is
- 411 reduced to a new linear system of Q equations in Q unknowns which requires much less computational
- 412 effort since  $Q \ll M$ , i.e.:

$$\mathbf{V}^{\top} \mathbf{A}^{\top} \mathbf{A} \mathbf{V} \alpha^* = \mathbf{V}^{\top} \mathbf{d}^o. \tag{45}$$

- 413 To avoid the storage of matrices A and V,? evaluates an element of the matrix AV by calculating the dot
- 414 product between the row of matrix A and the column of the matrix B. After estimating  $\alpha^*$  (equation 45)
- 415 belonging to a Q-dimensional subspace of  $\mathbb{R}^M$ , the distribution over the equivalent layer **p** in the  $\mathbb{R}^M$  is
- 416 obtained by applying equation 44. The choice of the Q basis vectors  $\mathbf{v}_i = 1, ..., Q$  (equation 44) in the
- 417 subspace method is not strict. ?, for example, chose the eigenvectors yielded by applying the singular value
- 418 decomposition of the matrix containing the gridded data set. The number of eigenvectors used to form
- 419 basis vectors will depend on the singular values.
- The proposed subspace method for solving large-scale equivalent-layer problem by ? was applied to
- 421 estimate the mass excess or deficiency caused by causative gravity sources.

# 422 4.0.2 The quadtree discretization

- To make the equivalent-layer technique tractable, ? also transformed the dense sensitivity matrix A
- 424 (equation 3) into a sparse matrix. In ?, a sparce version of the sensitivity matrix is achived by grouping
- 425 equivalent sources (e.g., they used prisms) distant from an observation point together to form a larger prism
- 426 or larger block. Each larger block has averaged physical properties and averaged top- and bottom-surfaces
- 427 of the grouped smaller prisms (equivalent sources) that are encompassed by the larger block. The authors
- 428 called it the 'larger averaged block' and the essence of their method is the reduction in the number of
- 429 equivalent sources, which means a reduction in the number of parameters to be estimated implying in
- 430 model dimension reduction.
- The key of the ?'s (?) method is the algorithm for deciding how to group the smaller prisms. In
- 432 practice, these authors used a recursive bisection process that results in a quadtree discretization of the
- 433 equivalent-layer model.
- By using the quadtree discretization, ? were able to jointly process multiple components of airborne
- 435 gravity-gradient data using a single layer of equivalent sources. To our knowledge, ? are the pioneers on
- 436 processing full-tensor gravity-gradient data jointly. In addition to computational feasibility, ?'s (?) method
- 437 reduces low-frequency noise and can also remove the drift in time-domain from the survey data. Those
- 438 authors stressed that the  $G_{zz}$ -component calculated through the single estimated equivalent-layer model
- 439 projected on a grid at a constant elevation by inverting full gravity-gradient data has the low-frequency error
- 440 reduced by a factor of 2.4 as compared to the inversion of an individual component of the gravity-gradient
- 441 data.

#### 442 4.0.3 The reparametrization of the equivalent layer

- Oliveira Jr. et al. (2013) reparametrized the whole equivalent-layer model by a piecewise bivariate-
- 444 polynomial function defined on a set of Q equivalent-source windows. In Oliveira Jr. et al.'s (2013)
- 445 approach, named polynomial equivalent layer (PEL), the parameter vector within the kth equivalent-source

446 window  $\mathbf{p}^k$  can be written in matrix notation as

$$\mathbf{p}^k = \mathbf{B}^k \, \mathbf{c}^k \,, \qquad k = 1 \dots Q \,, \tag{46}$$

where  $\mathbf{p}^k$  is an  $M_w$ -dimensional vector containing the physical-property distribution within the kth equivalent-source window,  $\mathbf{c}^k$  is a P-dimensional vector whose lth element is the lth coefficient of the  $\alpha$ th-order polynomial function and  $\mathbf{B}^k$  is an  $M_w \times P$  matrix containing the first-order derivative of the  $\alpha$ th-order polynomial function with respect to one of the P coefficients.

By using a regularized potential-field inversion, Oliveira Jr. et al. (2013) estimates the polynomial coefficients for each equivalent-source window by solving the following linear system

$$\left(\mathbf{B}^{\top} \mathbf{A}^{\top} \mathbf{A} \mathbf{B} + \mu \mathbf{I}\right) \mathbf{c}^{*} = \mathbf{B}^{\top} \mathbf{A}^{\top} \mathbf{d}^{o}, \tag{47}$$

where  $\mu$  is a regularizing parameter,  $\mathbf{c}^*$  is an estimated H-dimensional vector containing all coefficients describing all polynomial functions within all equivalent-source windows which compose the entire equivalent layer,  $\mathbf{I}$  is an identity matrix of order  $H(H=P\dot{Q})$  and  $\mathbf{B}$  is an  $M\times H$  block diagonal matrix such that the main-diagonal blocks are  $\mathbf{B}^k$  matrices (equation 46) and all off-diagonal blocks are zero matrices. For ease of the explanation of equation 47, we keep only the zeroth-order Tikhonov regularization and omitting the first-order Tikhonov regularization (?) which was also used by Oliveira Jr. et al. (2013).

The main advantage of the PEL is solve H-dimensional system of equations (equation 47), where H totalizes the number of polynomial coefficients composing all equivalent-source windows, requiring a lower computational effort since H <<< N. To avoid the storage of matrices A and B, Oliveira Jr. et al. (2013) evaluate an element of the matrix AB by calculating the dot product between the row of matrix AB and the column of the matrix BB. After estimating all polynomial coefficients of all windows, the estimated coefficients ( $C^*$  in equation 47) are transformed into a single physical-property distribution encompassing the entire equivalent layer.

As stated by Oliveira Jr. et al. (2013), the computational efficiency of PEL approach stems from the fact that the total number of polynomial coefficients *H* required to depict the physical-property distribution within the equivalent layer is generally much smaller than the number of equivalent sources. Consequently, this leads to a considerably smaller linear system that needs to be solved. Hence, the main strategy of polynomial equivalent layer is the model dimension reduction.

The polynomial equivalent layer was applied to perform upward continuations of gravity and magnetic data and reduction to the pole of magnetic data.

#### 473 4.0.4 The iterative scheme without solving a linear system

There exists a class of methods that iteratively estimate the distribution of physical properties within an 474 equivalent layer without the need to solve linear systems. The method initially introduced by Cordell (1992) 475 and later expanded upon by Guspí and Novara (2009) updates the physical property of sources, located 476 beneath each potential-field data, by removing the maximum residual between the observed and fitted data. 477 In addition, Xia and Sprowl (1991) and Xia et al. (1993) have developed efficient iterative algorithms for 478 updating the distribution of physical properties within the equivalent layer in the wavenumber and space 479 domains, respectively. Specifically, in Xia and Sprowl's (1991) method the physical-property distribution is 480 updated by using the ratio between the squared depth to the equivalent source and the gravitational constant 481

- 482 multiplied by the residual between the observed and predicted observation at the measurement station.
- 483 Neither of these methods solve linear systems.
- Following this class of methods of iterative equivalent-layer technique that does not solve linear systems,
- 485 Siqueira et al. (2017) developed a fast iterative equivalent-layer technique for processing gravity data in
- 486 which the sensitivity matrix A (equation 3) is replaced by a diagonal matrix  $N \times N$ , i.e.:

$$\tilde{\tilde{\mathbf{A}}} = 2 \,\pi \,\gamma \,\Delta \mathbf{S}^{-1} \,, \tag{48}$$

- 487 where  $\gamma$  is Newton's gravitational constant and  $\Delta S^{-1}$  is a diagonal matrix of order N whose diagonal
- 488 elements  $\Delta s_i$ , i=1,...,N are the element of area centered at the ith horizontal coordinates of the ith
- 489 observation point. The physical foundations of Siqueira et al.'s (2017) method rely on two constraints: i) the
- 490 excess of mass; and ii) the positive correlation between the gravity observations and the mass distribution
- 491 over the equivalent layer.
- 492 Although Siqueira et al.'s (2017) method does not solve any linear system of equations, it can be
- 493 theoretically explained by solving the following linear system at the kth iteration:

$$\tilde{\tilde{\mathbf{A}}}^{\top} \tilde{\tilde{\mathbf{A}}} \Delta \hat{\mathbf{p}}^k = \tilde{\tilde{\mathbf{A}}}^{\top} \mathbf{r}^k, \tag{49}$$

494 where  $\mathbf{r}^k$  is an N-dimensional residual vector whose ith element is calculated by subtracting the ith

495 observed data  $d_i^o$  from the *i*th fitted data  $d_i^k$  at the *k*th iteration, i.e.,

$$r_i^k = d_i^o - d_i^k \,. \tag{50}$$

496 and  $\Delta \hat{\mathbf{p}}^k$  is an estimated N-dimensional vector of parameter correction.

- Because  $\tilde{\tilde{A}}$ , in equation 49, is a diagonal matrix (equation 48), the parameter correction estimate is
- 498 directly calculated without solving system of linear equations, and thus, an ith element of  $\Delta \hat{\mathbf{p}}^k$  is directly
- 499 calculated by

$$\Delta \hat{p}^k_{\ i} = \frac{\Delta s_i \ r_i^k}{2 \ \pi \ \gamma} \ . \tag{51}$$

The mass distribution over the equivalent layer is updated by:

$$\hat{p}_i^{k+1} = \hat{p}_i^k + \Delta \hat{p}_i^k. \tag{52}$$

Siqueira et al.'s (2017) method starts from a mass distribution on the equivalent layer, whose ith mass  $p_i^o$  is

502 proportional to the *i*th observed data  $d_i^o$ , i.e.,

$$p_i^o = \frac{\Delta s_i \, d_i^o}{2 \, \pi \, \gamma} \,. \tag{53}$$

503 Siqueira et al. (2017) applied their fast iterative equivalent-layer technique to interpolate, calculate the

- 504 horizontal components, and continue upward (or downward) gravity data.
- For jointly process two gravity gradient components, Jirigalatu and Ebbing (2019) used the Gauss-FFT
- 506 for forward calculation of potential fields in the wavenumber domain combined with Landweber's iteration
- 507 coupled with a mask matrix M to reduce the edge effects without increasing the computation cost. The

mask matrix M is defined in the following way: if the corresponding pixel does not contain the original data, the element of M is set to zero; otherwise, it is set to one. The kth Landweber iteration is given by

$$\mathbf{p}_{k+1} = \mathbf{p}_k + \omega \left[ \mathbf{A_1}^{\top} (\mathbf{d_1} - \mathbf{M} \mathbf{A_1} \mathbf{p}_k) + \mathbf{A_2}^{\top} (\mathbf{d_2} - \mathbf{M} \mathbf{A_2} \mathbf{p}_k) \right],$$
 (54)

- where  $\omega$  is a relaxation factor,  $d_1$  and  $d_2$  are the two gravity gradient components and  $A_1$  and  $A_2$  are the
- 511 corresponding gravity gradient kernels. Jirigalatu and Ebbing (2019) applied their method for processing
- 512 two horizontal curvature components of Falcon airborne gravity gradient.
- 513 4.0.5 The convolutional equivalent layer with BTTB matrices
- ? (?, ?) introduced the convolutional equivalent layer for gravimetric and magnetic data processing, respectively.
- 516 ? demonstrated that the sensitivity matrix A (equation 3) associated with a planar equivalent layer formed by a set of point masses, each one directly beneath each observation point and considering a regular grid of 517 observation points at a constant height has a symmetric block-Toeplitz Toeplitz-block (BTTB) structure. A 518 519 symmetric BTTB matrix has, at least, two attractive properties. The first one is that it can be defined by using only the elements forming its first column (or row). The second attractive property is that any BTTB 520 matrix can be embedded into a symmetric Block-Circulat Circulant-Block (BCCB) matrix. This means that 521 the full sensitivity matrix A (equation 3) can be completely reconstruct by using the first column of the 522 BCCB matrix only. In what follows, ? computed the forward modeling by using only a single equivalent 523 source. Specifically, it is done by calculating the eigenvalues of the BCCB matrix that can be efficiently 524 computed by using only the first column of the BCCB matrix via 2D fast Fourier transform (2D FFT). 525
- 526 By comparing with the classic approach in the Fourier domain, the convolutional equivalent layer for
- 527 gravimetric data processing proposed by ? performed upward- and downward-continue gravity data with a
- 528 very small border effects and noise amplification.
- By using the original idea of the convolutional equivalent layer proposed by ? for gravimetric data
- processing, ? developed the convolutional equivalent layer for magnetic data processing. By assuming a regularly spaced grid of magnetic data at a constant height and a planar equivalent layer of dipoles, ?
- 532 proved that the sensitivity matrix linked with this layer possess a BTTB structure in the specific scenario
- 533 where each dipole is exactly beneath each observed magnetic data point. ? used a conjugate gradient
- least-squares (CGLS) algorithm which does not require an inverse matrix or matrix-matrix multiplication.
- 535 Rather, it only requires matrix-vector multiplications per iteration, which can be effectively computed using
- 536 the 2D FFT as a discrete convolution. The matrix-vector product only uses the elements that constitute the
- 537 first column of the associated BTTB matrix, resulting in computational time and memory savings. ? (?)
- 538 showed the robustness of the convolutional equivalent layer in processing magnetic survey that violates the
- 539 requirement of regular grids in the horizontal directions and flat observation surfaces.
- The matrix-vector product in ? (?, ?) (e.g., d = Ap, such as in equation 3) is the main issue to be solved.
- 541 To solve it efficiently, these authors involked the auxiliary linear system

$$\mathbf{w} = \mathbf{C}\mathbf{v} \,, \tag{55}$$

- 542 where w and v are, respectively, vectors of data and parameters completed by zeros and C is a BCCB
- matrix formed by  $2Q \times 2Q$  blocks, where each block  $\mathbf{C}_q$ ,  $q = 0, \dots, Q 1$ , is a  $2P \times 2P$  circulant matrix.
- The first column of C is obtained by rearranging the first column of the sensitivity matrix A (equation 3).

Because a BCCB matrix is diagonalized by the 2D unitary discrete Fourier transform (DFT), C can be

546 written as

$$\mathbf{C} = (\mathbf{F}_{2O} \otimes \mathbf{F}_{2P})^* \Lambda (\mathbf{F}_{2O} \otimes \mathbf{F}_{2P}) , \qquad (56)$$

547 where the symbol " $\otimes$ " denotes the Kronecker product (?),  $\mathbf{F}_{2Q}$  and  $\mathbf{F}_{2P}$  are the  $2Q \times 2Q$  and  $2P \times 2P$ 

- unitary DFT matrices (?, p. 31), respectively, the superscritpt "\*" denotes the complex conjugate and  $\Lambda$  is
- 549 a  $4QP \times 4QP$  diagonal matrix containing the eigenvalues of C. Due to the diagonalization of the matrix
- 550 C, the auxiliary system (equation 55) can be rewritten by using equation 56 and premultiplying both sides
- 551 of the result by  $(\mathbf{F}_{2O} \otimes \mathbf{F}_{2P})$ , i.e.,

$$\mathbf{\Lambda} \left( \mathbf{F}_{2Q} \otimes \mathbf{F}_{2P} \right) \mathbf{v} = \left( \mathbf{F}_{2Q} \otimes \mathbf{F}_{2P} \right) \mathbf{w} . \tag{57}$$

By applying the vec-operator (?) to both sides of equation 57, by premultiplying both sides of the result by

53  $\mathbf{F}_{2Q}^*$  and then postmultiplying both sides of the result by  $\mathbf{F}_{2P}^*$ 

$$\mathbf{F}_{2Q}^{*} \left[ \mathbf{L} \circ \left( \mathbf{F}_{2Q} \mathbf{V} \mathbf{F}_{2P} \right) \right] \mathbf{F}_{2P}^{*} = \mathbf{W} , \qquad (58)$$

- where " $\circ$ " denotes the Hadamard product (?, p. 298) and L, V and W are  $2Q \times 2P$  matrices obtained
- by rearranging, along their rows, the elements forming the diagonal of matrix  $\Lambda$ , vector  $\mathbf{v}$  and vector  $\mathbf{w}$ ,
- 556 respectively. The left side of equation 58 contains the 2D Inverse Discrete Fourier Transform (IDFT) of the
- 557 term in brackets, which in turn represents the Hadamard product of matrix L and the 2D DFT of matrix V.
- Matrix L contains the eigenvalues of  $\Lambda$  (equation 56) and can be efficiently computed by using only the
- 559 first column of the BCCB matrix C (equation 55).
- Actually, in ? (?, ?) a fast 2D discrete circular convolution (Van Loan, 1992) is used to process very
- large gravity and magnetic datasets efficiently. The convolutional equivalent layer was applied to perform
- 562 upward continuation of large magnetic datasets. Compared to the classical Fourier approach, ?'s (?) method
- 563 produces smaller border effects without using any padding scheme.
- Without taking advantage of the symmetric BTTB structure of the sensitivity matrix (?, ?) that arises
- 565 when gravimetric observations are measured on a horizontally regular grid, on a flat surface and considering
- 566 a regular grid of equivalent sources whithin a horizontal layer, ? explored the symmetry of the gravity
- 567 kernel to reduce the number of forward model evaluations. By exploting the symmetries of the gravity
- 568 kernels and redundancies in the forward model evaluations on a regular grid and combining the subspace
- kernels and redundances in the forward model evaluations on a regular grid and combining the subspace
- solution based on eigenvectors of the gridded dataset, ? estimated the mass excess or deficiency produced
- 570 by anomalous sources with positive or negative density contrast.

#### 571 4.0.6 The deconvolutional equivalent layer with BTTB matrices

- To avoid the iterations of the conjugate gradient method in ?, we can employ the deconvolution process.
- 573 Equation 58 shows that estimate the matrix V, containing the elements of parameter vector p, is a
- 574 inverse problem that could be solved by deconvolution. From equation 58, the matrix V can be obtain by
- 575 deconvolution, i.e.

$$\mathbf{V} = \mathbf{F}_{2Q}^* \left[ \frac{\left( \mathbf{F}_{2Q} \mathbf{W} \, \mathbf{F}_{2P} \right)}{\mathbf{L}} \right] \mathbf{F}_{2P}^* \,. \tag{59}$$

- Equation 59 shows that the parameter vector (in matrix V) can be theoretically obtain by dividing each
- 577 potential-field observations (in matrix W) by each eigenvalues (in matrix L). Hence, the parameter vector
- 578 is constructed by element-by-element division of data by eigenvalues.

However, the deconvolution often is extremely unstable. This means that a small change in data can lead to an enormous change in the estimated parameter. Hence, equation 59 requires regularization to be useful.

581 We usede wiener deconvolution to obtain a stable solution, i.e.,

$$\mathbf{V} = \mathbf{F}_{2Q}^* \left[ \left( \mathbf{F}_{2Q} \mathbf{W} \, \mathbf{F}_{2P} \right) \frac{\mathbf{L}^*}{(\mathbf{L} \, \mathbf{L}^* + \mu)} \right] \mathbf{F}_{2P}^* \,, \tag{60}$$

where the matrix  $L^*$  contains the complex conjugate eigenvalues and  $\mu$  is a parameter that controls the degree of stabilization.

# 584 4.1 Solution stability

The solution stability of the equivalent-layer methods is rarely addressed. Here, we follow the numerical stability analysis presented in Siqueira et al. (2017).

Let us assume noise-free potential-field data  $\mathbf{d}$ , we estimate a physical-property distribution  $\mathbf{p}$  (estimated solution) within the equivalent layer. Then, the noise-free data  $\mathbf{d}$  are contaminated with additive D different sequences of pseudorandom Gaussian noise, creating different noise-corrupted potential-field data  $\mathbf{d}_{\ell}^{\mathbf{o}}$ ,  $\ell = 1, ..., D$ . From each  $\mathbf{d}_{\ell}^{\mathbf{o}}$ , we estimate a physical-property distribution  $\hat{\mathbf{p}}_{\ell}$  within the equivalent layer.

Next, for each noise-corrupted data  $\mathbf{d}_{\ell}^{\mathbf{o}}$  and estimated solution  $\hat{\mathbf{p}}_{\ell}$ , the  $\ell$ th model perturbation  $\delta p_{\ell}$  and the  $\ell$ th data perturbation  $\delta d_{\ell}$  are, respectively, evaluated by

$$\delta p_{\ell} = \frac{\parallel \hat{\mathbf{p}}_{\ell} - \mathbf{p} \parallel_2}{\parallel \mathbf{p} \parallel_2}, \quad \ell = 1, ..., D,$$
 (61)

593 and

$$\delta d_{\ell} = \frac{\parallel \mathbf{d}_{\ell}^{\mathbf{o}} - \mathbf{d} \parallel_2}{\parallel \mathbf{d} \parallel_2}, \quad \ell = 1, ..., D.$$

$$(62)$$

Regardless of the particular method used, the following inequality (?, p. 66) is applicable:

$$\delta p_{\ell} \le \kappa \, \delta d_{\ell}, \quad \ell = 1, ..., D,$$
 (63)

where  $\kappa$  is the constant of proportionality between the model perturbation  $\delta p_{\ell}$  (equation 61) and the data perturbation  $\delta d_{\ell}$  (equation 62). The constant  $\kappa$  acts as the condition number of an invertible matrix in a given inversion, and thus measures the instability of the solution. The larger (smaller) the value of  $\kappa$  the more unstable (stable) is the estimated solution.

Equation 63 shows a linear relationship between the model perturbation and the data perturbation. By plotting  $\delta p_{\ell}$  (equation 61) against  $\delta d_{\ell}$  (equation 62) produced by a set of D estimated solution obtained by applying a given equivalent-layer method, we obtain a straight line behaviour described by equation 63. By applying a linear regression, we obtain a fitted straight line whose estimated slope ( $\kappa$  in equation 63) quantifies the solution stability.

Here, the analysis of solution stability is numerically conducted by applying the classical equivalent-layer technique with zeroth-order Tikhonov regularization, the convolutional method for gravimetric and magnetic data, the deconvolutional method (equation 59) and the deconvolutional method with different values for the Wiener stabilization (equation 60).

varides for the whener stabilization (equation 60)

# 5 NUMERICAL SIMULATIONS

- 608 We investigated different computational algorithms for inverting gravity disturbances and total-field
- anomalies. To test the capability of the fast equivalent-layer technique for processing that potential field
- 610 data we measure of the computational effort by counting the number of floating-point operations (*flops*),
- such as additions, subtractions, multiplications, and divisions (Golub and Van Loan, 2013) for different
- 612 number of observation points, ranging from 10,000 up to 1,000,000. The results generated when using
- iterative methods are set to it = 50 for the number of iterations.

# 614 5.1 Floating-point operations calculation

- To measure the computational effort of the different algorithms to solve the equivalent layer linear system,
- a non-hardware dependent method can be useful because allow us to do direct comparison between them.
- 617 Counting the floating-point operations (flops), i.e., additions, subtractions, multiplications and divisions
- 618 is a good way to quantify the amount of work of a given algorithm (Golub and Van Loan, 2013). For
- example, the number of flops necessary to multiply two vectors  $\mathbb{R}^N$  is 2N. A common matrix-vector
- 620 multiplication with dimension  $\mathbb{R}^{N\times N}$  and  $\mathbb{R}^N$ , respectively, is  $2N^2$  and a multiplication of two matrices
- 621  $\mathbb{R}^{N\times N}$  is  $2N^3$ . Figure ?? shows the total flops count for the different methods presented in this review
- 622 with a crescent number of data, ranging from 10,000 to 1,000,000 for the gravity equivalent layer and
- 623 figure ?? for magnetic data.

# 624 5.1.1 Normal equations using Cholesky decomposition

- The equivalent sources can be estimated directly from solving the normal equations 3. In this work we
- 626 will use the Cholesky decompositions method to calculate the necessary flops. In this method it is calculated
- 627 the lower triangule of  $A^TA$  (1/2N<sup>3</sup>), the Cholesky factor (1/3N<sup>3</sup>), a matrix-vector multiplication (2N<sup>2</sup>)
- 628 and finally solving the triangular system  $(2N^2)$ , totalizing

$$f_{classical} = \frac{5}{6}N^3 + 4N^2 \tag{64}$$

# 629 5.1.2 Window method (Leão and Silva, 1989)

- The moving data-window scheme (Leão and Silva, 1989) solve N linear systems with much smaller
- 631 sizes (equation ??). For our results we are considering a data-window of the same size of wich the authors
- 632 presented in theirs work ( $N_w = 49$ ) and the same number of equivalent sources ( $M_w = 225$ ). We are doing
- 633 this process for all the other techniques to standardize the resolution of our problem. Using the Cholesky
- 634 decomposition with this method the *flops* are

$$f_{window} = N\frac{5}{6}M_w N_w^2 + 4N_w M_w (65)$$

# 635 5.1.3 PEL method (Oliveira Jr. et al., 2013)

- The polynomial equivalent layer uses a simliar approach od moving windows from Leão and Silva (1989).
- 637 For this operations calculation (equation 47) we used a first degree polynomial (two variables) and each
- 638 window contains  $N_s = 1,000$  observed data and  $M_s = 1,000$  equivalent sources. Following the steps
- 639 given in (Oliveira Jr. et al., 2013) the total flops becomes

$$f_{pel} = \frac{1}{3}H^3 + 2H^2 + 2NM_sH + H^2N + 2HN + 2NP$$
 (66)

- 640 where H is the number of constant coefficients for the first degree polynomial (P=3) times the number 641 of windows  $(P \times N/N_s)$ .
- 642 5.1.4 Conjugate gradient least square (CGLS)
- The CGLS method is a very stable and fast algorithm for solving linear systems iteratively. Its computational complexity envolves a matrix-vector product outside the loop  $(2N^2)$ , two matrix-vector products inside the loop  $(4N^2)$  and six vector products inside the loop (12N) (?)

$$f_{cals} = 2N^2 + it(4N^2 + 12N) (67)$$

- 646 5.1.5 Wavelet compression method with CGLS (?)
- For the wavelet method (equation 40) we have calculated a coompression rate of 98% (  $C_r = 0.02$  )
- for the threshold as the authors used in ? and the wavelet transformation requiring  $\log_2(N)$  flops each
- 649 (equations 39 and 41c), with its inverse also using the same number of operations (equation 43). Combined
- 650 with the conjugate gradient least square necessary steps and iterations, the number of *flops* are

$$f_{wavelet} = 2NC_r + 4N\log_2(N) + it(4N\log_2(N) + 4NC_r + 12C_r)$$
(68)

- 651 5.1.6 Fast equivalent layer for gravity data (Siqueira et al., 2017)
- The fast equivalent layer from Siqueira et al. (2017) solves the linear system in *it* iterations. The main
- 653 cost of this method (equations 49,50, 51 and 52)is the matrix-vector multiplication to asses the predicted
- data  $(2N^2)$  and three simply element by element vector sum, subtraction and division (3N total)

$$f_{siqueira} = it(3N + 2N^2) \tag{69}$$

- 655 5.1.7 Convolutional equivalent layer for gravity data (?)
- This methods replaces the matrix-vector multiplication of the iterative fast-equivalent technique (Siqueira
- et al., 2017) by three steps, involving a Fourier transform, an inverse Fourier transform, and a Hadamard
- product of matrices (equation 58). Considering that the first column of our BCCB matrix has 4N elements,
- 659 the flops count of this method is

$$f_{convgrav} = \kappa 4N \log_2(4N) + it(27N + \kappa 8N \log_2(4N)) \tag{70}$$

In the resultant count we considered a radix-2 algorithm for the fast Fourier transform and its inverse, which has a  $\kappa$  equals to 5 and requires  $\kappa 4N \log_2(4N)$  flops each. The Hadarmard product of two matrices of 4N elements with complex numbers takes 24N flops. Note that equation 70 is different from the one presented in ? because we also added the flops necessary to calculate the eigenvalues in this form. It does not differentiate much in order of magnitude because the iterative part is the most costful.

#### 665 5.1.8 Convolutional equivalent layer for magnetic data (?)

The convolutional equivalent layer for magnetic data uses the same flops count of the main operations as in the gravimetric case (equation 58), the difference is the use of the conjugate gradient algorithm to solve the inverse problem. It requires a Hadamard product outside of the iterative loop and the matrix-vector and vector-vector multiplications inside the loop as seem in equation 67.

$$f_{convmag} = \kappa 16N \log_2(4N) + 24N + it(\kappa 16N \log_2(4N) + 60N)$$
(71)

# 670 5.1.9 Deconvolutional method

The deconvolution method does not require an iterative algorithm, rather it solves the estimative of the physical properties in a single step using the 4N eigenvalues of the BCCB matrix as in the convolutional method. From equation 59 it is possible to deduce this method requires two fast Fourier transform ( $\kappa 4N \log_2(4N)$ ), one for the eigenvalues and another for the data transformation, a element by element division (24N) and finally, a fast inverse Fourier transform for the final estimative ( $\kappa 4N \log_2(4N)$ ).

$$f_{deconv} = \kappa 12N \log_2(4N) + 24N \tag{72}$$

Using the deconvolutional method with a Wiener stabilization adds two multiplications of complex elements of the conjugates eigenvalues (24N each) and the sum of 4N elements with the stabilization parameter  $\mu$  as shown in equation 60

$$f_{deconvviener} = \kappa 12N \log_2(4N) + 76N \tag{73}$$

# **CONFLICT OF INTEREST STATEMENT**

- 679 The authors declare that the research was conducted in the absence of any commercial or financial
- 680 relationships that could be construed as a potential conflict of interest.

#### **AUTHOR CONTRIBUTIONS**

- 681 The Author Contributions section is mandatory for all articles, including articles by sole authors. If an
- 682 appropriate statement is not provided on submission, a standard one will be inserted during the production
- 683 process. The Author Contributions statement must describe the contributions of individual authors referred
- 684 to by their initials and, in doing so, all authors agree to be accountable for the content of the work. Please
- see here for full authorship criteria.

#### **FUNDING**

- 686 Diego Takahashi was supported by a Post-doctoral scholarship from CNPq (grant 300809/2022-0) Valéria
- 687 C.F. Barbosa was supported by fellowships from CNPq (grant 309624/2021-5) and FAPERJ (grant
- 688 26/202.582/2019). Vanderlei C. Oliveira Jr. was supported by fellowships from CNPq (grant 315768/2020-
- 689 7) and FAPERJ (grant E-26/202.729/2018).

#### **ACKNOWLEDGMENTS**

- 690 We thank the brazillian federal agencies CAPES, CNPq, state agency FAPERJ and Observatório Nacional
- 691 research institute and Universidade do Estado do Rio de Janeiro.

# **DATA AVAILABILITY STATEMENT**

- 692 The datasets generated for this study can be found in the frontiers-paper Github repository link:
- 693 https://github.com/DiegoTaka/frontiers-paper.

#### **REFERENCES**

- 694 Aster, R. C., Borchers, B., and Thurber, C. H. (2019). Parameter Estimation and Inverse Problems
- 695 (Elsevier), 3 edn.
- 696 Barbosa, V. C. F., Silva, J. B., and Medeiros, W. E. (1997). Gravity inversion of basement relief using
- approximate equality constraints on depths. *Geophysics* 62, 1745–1757
- 698 Barnes, G. and Lumley, J. (2011). Processing gravity gradient data. GEOPHYSICS 76, I33–I47. doi:10.
- 699 1190/1.3548548
- 700 Blakely, R. J. (1996). Potential Theory in Gravity and Magnetic Applications (Cambridge University
- 701 press)
- 702 Bott, M. H. P. (1960). The use of Rapid Digital Computing Methods for Direct Gravity Interpretation
- of Sedimentary Basins. Geophysical Journal International 3, 63–67. doi:10.1111/j.1365-246X.1960.
- 704 tb00065.x
- 705 Chan, R. H.-F. and Jin, X.-Q. (2007). An introduction to iterative Toeplitz solvers, vol. 5 (SIAM)
- 706 Cordell, L. (1992). A scattered equivalent-source method for interpolation and gridding of potential-field
- data in three dimensions. *Geophysics* 57, 629–636

- 708 Dampney, C. N. G. (1969). The equivalent source technique. GEOPHYSICS 34, 39–53. doi:10.1190/1.
- 709 1439996
- 710 Emilia, D. A. (1973). Equivalent sources used as an analytic base for processing total magnetic field
- 711 profiles. *GEOPHYSICS* 38, 339–348. doi:10.1190/1.1440344
- Golub, G. H. and Van Loan, C. F. (2013). *Matrix Computations*. Johns Hopkins Studies in the Mathematical
- 713 Sciences (Johns Hopkins University Press), 4 edn.
- 714 Gonzalez, S. P., Barbosa, V. C. F., and Oliveira Jr., V. C. (2022). Analyzing the ambiguity of the remanent-
- 715 magnetization direction separated into induced and remanent magnetic sources. *Journal of Geophysical*
- 716 Research: Solid Earth 127, 1–24. doi:10.1029/2022JB024151
- 717 Guspí, F., Introcaso, A., and Introcaso, B. (2004). Gravity-enhanced representation of measured geoid
- undulations using equivalent sources. Geophysical Journal International 159, 1-8. doi:10.1111/j.
- 719 1365-246X.2004.02364.x
- 720 Guspí, F. and Novara, I. (2009). Reduction to the pole and transformations of scattered magnetic data using
- Newtonian equivalent sources. GEOPHYSICS 74, L67–L73. doi:10.1190/1.3170690
- 722 Hansen, R. O. and Miyazaki, Y. (1984). Continuation of potential fields between arbitrary surfaces.
- 723 *GEOPHYSICS* 49, 787–795. doi:10.1190/1.1441707
- 724 Horn, R. A. and Johnson, C. R. (1991). *Topics in Matrix Analysis* (Cambridge University Press), 1 edn.
- 725 Jain, A. K. (1989). Fundamentals of Digital Image Processing (Pearson), 1 edn.
- 726 Jirigalatu, J. and Ebbing (2019). A fast equivalent source method for airborne gravity gradient data.
- 727 Geophysics 84, G75–G82. doi:10.1190/GEO2018-0366.1
- 728 Kellogg, O. D. (1967). Foundations of Potential Theory (Springer-Verlag), reprint from the first edition of
- 729 1929 edn.
- 730 Kennett, B., Sambridge, M., and Williamson, P. (1988). Subspace methods for large inverse problems with
- multiple parameter classes. *Geophysical Journal International* 94, 237–247
- 732 Leão, J. W. D. and Silva, J. B. C. (1989). Discrete linear transformations of potential field data. Geophysics
- 733 54, 497–507. doi:10.1190/1.1442676
- 734 Li, Y., Nabighian, M., and Oldenburg, D. W. (2014). Using an equivalent source with positivity for low-
- latitude reduction to the pole without striation. *GEOPHYSICS* 79, J81–J90. doi:10.1190/geo2014-0134.
- 736 1
- 737 Li, Y. and Oldenburg, D. W. (2010). Rapid construction of equivalent sources using wavelets.
- 738 *GEOPHYSICS* 75, L51–L59. doi:10.1190/1.3378764
- 739 Mendonça, C. A. (2020). Subspace method for solving large-scale equivalent layer and density mapping
- 740 problems. *GEOPHYSICS* 85, G57–G68. doi:10.1190/geo2019-0302.1
- 741 Mendonça, C. A. and Silva, J. B. C. (1994). The equivalent data concept applied to the interpolation of
- 742 potential field data. *Geophysics* 59, 722–732. doi:10.1190/1.1443630
- 743 Menke, W. (2018). Geophysical data analysis: Discrete inverse theory (Elsevier), 4 edn.
- 744 Oldenburg, D., McGillivray, P., and Ellis, R. (1993). Generalized subspace methods for large-scale inverse
- 745 problems. *Geophysical Journal International* 114, 12–20
- 746 Oliveira Jr., V. C., Barbosa, V. C. F., and Uieda, L. (2013). Polynomial equivalent layer. GEOPHYSICS 78,
- 747 G1-G13. doi:10.1190/geo2012-0196.1
- 748 Reis, A. L. A., Oliveira Jr., V. C., and Barbosa, V. C. F. (2020). Generalized positivity constraint on
- 749 magnetic equivalent layers. *Geophysics* 85, 1–45. doi:10.1190/geo2019-0706.1
- 750 Roy, A. (1962). Ambiguity in geophysical interpretation. GEOPHYSICS 27, 90-99. doi:10.1190/1.
- 751 1438985

- 752 Silva, J. B. C. (1986). Reduction to the pole as an inverse problem and its application to low-latitude anomalies. *GEOPHYSICS* 51, 369–382. doi:10.1190/1.1442096
- 754 Siqueira, F., Oliveira Jr., V. C., and Barbosa, V. C. F. (2017). Fast iterative equivalent-layer technique for
- gravity data processing: A method grounded on excess mass constraint. *GEOPHYSICS* 82, G57–G69.
- 756 doi:10.1190/GEO2016-0332.1
- 757 Skilling, J. and Bryan, R. (1984). Maximum entropy image reconstruction-general algorithm. *Monthly*758 *Notices of the Royal Astronomical Society, Vol. 211, NO. 1, P. 111, 1984* 211, 111
- Soler, S. R. and Uieda, L. (2021). Gradient-boosted equivalent sources. *Geophysical Journal International* 227, 1768–1783. doi:10.1093/gji/ggab297
- Takahashi, D., Oliveira Jr., V. C., and Barbosa, V. C. (2022). Convolutional equivalent layer for magnetic
   data processing. *Geophysics* 87, 1–59
- Takahashi, D., Oliveira Jr., V. C., and Barbosa, V. C. F. (2020). Convolutional equivalent layer for gravity data processing. *GEOPHYSICS* 85, G129–G141. doi:10.1190/geo2019-0826.1
- van der Sluis, A. and van der Vorst, H. A. (1987). Numerical solution of large, sparse linear algebraic
- systems arising from tomographic problems. In Seismic tomography with applications in global
- seismology and exploration geophysics, ed. G. Nolet (D. Reidel Publishing Company), chap. 3. 49–83
- Van Loan, C. F. (1992). Computational Frameworks for the fast Fourier transform. Frontiers in Applied
   Mathematics (SIAM)
- Xia, J. and Sprowl, D. R. (1991). Correction of topographic distortion in gravity data. *Geophysics* 56,
   537–541
- Xia, J., Sprowl, D. R., and Adkins-Heljeson, D. (1993). Correction of topographic distortions in potential-field data; a fast and accurate approach. *Geophysics* 58, 515–523. doi:10.1190/1.1443434
- 774 Zhao, G., Chen, B., Chen, L., Liu, J., and Ren, Z. (2018). High-accuracy 3D Fourier forward modeling
- of gravity field based on the Gauss-FFT technique. *Journal of Applied Geophysics* 150, 294–303.
- 776 doi:10.1016/j.jappgeo.2018.01.002
- Zidarov, D. (1965). Solution of some inverse problems of applied geophysics. *Geophysical Prospecting* 13, 240–246. doi:10.1111/j.1365-2478.1965.tb01932.x