

The computation aspects of the equivalent-layer technique: review and perspective

Diego Takahashi^{1,*}, André L. A. Reis², Vanderlei C. Oliveira Jr.¹ and Valéria C. F. Barbosa¹

¹*Observatório Nacional, Department of Geophysics, Rio de Janeiro, Brasil*

²*Universidade do Estado do Rio de Janeiro, Department of Applied Geology, Rio de Janeiro, Brasil*

Correspondence*:
Valéria C.F. Barbosa
valcris@on.br

1 FUNDAMENTALS

- 2 d_i^o
 3 (x_i, y_i, z_i)
 4 $i \in \{1 : D\}$
 5 approximate d_i^o by a harmonic function

$$d_i^p = \sum_{j=1}^P g_{ij} p_j, \quad i \in \{1 : D\}, \quad (1)$$

6 where,

- 7 p_j
 8 (x_j, y_j, z_j)
 9 $j \in \{1 : P\}$

$$g_{ij} \equiv g(x_i - x_j, y_i - y_j, z_i - z_j), \quad z_i < \min\{z_j\}, \quad \forall i \in \{1 : D\}, \quad (2)$$

10 where $\min\{z_j\}$ denotes the minimum z_j (or the vertical coordinate of the shallowest equivalent source).

$$\mathbf{d}_p = \mathbf{G}\mathbf{p}, \quad (3)$$

11 where \mathbf{p} is a $P \times 1$ vector with j -th element p_j representing the scalar physical property of the j -th
 12 equivalent source and \mathbf{G} is a $D \times P$ matrix with element g_{ij} given by equation 2.

13 The equivalent-layer technique consists in solving a linear inverse problem to determine a parameter
 14 vector \mathbf{p} leading to a predicted data vector \mathbf{d}_p (equation 3) *sufficiently close to* the observed data vector \mathbf{d}_o ,
 15 whose i -th element d_i^o is the observed potential field at (x_i, y_i, z_i) . The notion of *closeness* is intrinsically
 16 related to the concept of *vector norm* (e.g., Golub and Loan, 2013, p. 68) or *measure of length* (e.g., Menke,
 17 2018, p. 41). Because of that, almost all methods for determining \mathbf{p} actually estimate a parameter vector $\tilde{\mathbf{p}}$
 18 minimizing a length measure of the difference between \mathbf{d}_p and \mathbf{d}_o (see subsection 1.3). Given an estimate
 19 $\tilde{\mathbf{p}}$, it is then possible to compute a potential field transformation

$$\mathbf{d}_t = \mathbf{T}\tilde{\mathbf{p}}, \quad (4)$$

20 where \mathbf{d}_t is a $T \times 1$ vector with k -th element d_k^t representing the transformed potential field at the position
 21 (x_k, y_k, z_k) , $k \in \{1 : T\}$, and

$$t_{kj} \equiv t(x_k - x_j, y_k - y_j, z_k - z_j), \quad z_k < \min\{z_j\}, \quad \forall k \in \{1 : T\}, \quad (5)$$

22 is a harmonic function representing the kj -th element of the $T \times P$ matrix \mathbf{T} .

23 1.1 Spatial distribution and total number of equivalent sources

24 There is no well-established criteria to define the optimum number P or the spatial distribution of the
 25 equivalent sources. We know that setting an equivalent layer with more (less) sources than potential-field

data usually leads to an underdetermined (overdetermined) inverse problem (e.g., Menke, 2018, p. 52–53). Concerning the spatial distribution of the equivalent sources, the only condition is that they must rely on a surface that is located below and does not cross that containing the potential field data. Soler and Uieda (2021) present a practical discussion about this topic.

From a theoretical point of view, the equivalent layer reproducing a given potential field data set cannot cross the true gravity or magnetic sources. This condition is a consequence of recognizing that the equivalent layer is essentially an indirect solution of a boundary value problem of potential theory (e.g., Roy, 1962; Zidarov, 1965; Dampney, 1969; Li et al., 2014; Reis et al., 2020). In practical applications, however, there is no guarantee that this condition is satisfied. Actually, it is widely known from practical experience (e.g., Gonzalez et al., 2022) that the equivalent-layer technique works even for the case in which the layer cross the true sources.

CRITÉRIOS PARA DEFINIR A PROFUNDIDADE DA CAMADA: DAMPNEY (ESPAÇAMENTO DO GRID) E REIS (ESPAÇAMENTO DAS LINHAS)

1.2 Sensitivity matrix A

Generally, the harmonic function g_{ij} (equation 2) is defined in terms of the inverse distance between the observation point (x_i, y_i, z_i) and the j -th equivalent source at (x_j, y_j, z_j) ,

$$\frac{1}{r_{ij}} \equiv \frac{1}{\sqrt{(x_i - x_j)^2 + (y_i - y_j)^2 + (z_i - z_j)^2}}, \quad (6)$$

or by its partial derivatives of first and second orders, respectively given by

$$\partial_\alpha \frac{1}{r_{ij}} \equiv \frac{-(\alpha_i - \alpha_j)}{r_{ij}^3}, \quad \alpha \in \{x, y, z\}, \quad (7)$$

and

$$\partial_{\alpha\beta} \frac{1}{r_{ij}} \equiv \begin{cases} \frac{3(\alpha_i - \alpha_j)^2}{r_{ij}^5}, & \alpha = \beta, \\ \frac{3(\alpha_i - \alpha_j)(\beta_i - \beta_j)}{r_{ij}^5} - \frac{1}{r_{ij}^3}, & \alpha \neq \beta, \end{cases} \quad \alpha, \beta \in \{x, y, z\}. \quad (8)$$

In this case, the equivalent layer is formed by punctual sources representing monopoles or dipoles (e.g., Dampney, 1969; Emilia, 1973; Leão and Silva, 1989; Cordell, 1992; Oliveira Jr. et al., 2013; Siqueira et al., 2017; Reis et al., 2020; Takahashi et al., 2020; Soler and Uieda, 2021; Takahashi et al., 2022). Another common approach consists in not defining g_{ij} by using equations 6–8, but other harmonic functions obtained by integrating them over the volume of regular prisms (e.g., Li and Oldenburg, 2010; Barnes and Lumley, 2011; Li et al., 2014; Jiriglatu and Ebbing, 2019). There are also some less common approaches defining the harmonic function g_{ij} (equation 2) as the potential field due to plane faces with constant physical property (Hansen and Miyazaki, 1984), doublets (Silva, 1986) or by computing the double integration of the inverse distance function with respect to z (Guspí and Novara, 2009).

A common assumption for most of the equivalent-layer methods is that the harmonic function g_{ij} (equation 2) is independent on the actual physical relationship between the observed potential field and their true sources (e.g., Cordell, 1992; Guspí and Novara, 2009; Li et al., 2014). Hence, g_{ij} can be defined according to the problem. The only condition imposed to this function is that it decays to zero as the observation point (x_i, y_i, z_i) goes away from the position (x_j, y_j, z_j) of the j -th equivalent source.

However, several methods use a function g_{ij} that preserves the physical relationship between the observed potential field and their true sources. For the case in which the observed potential field is gravity data, g_{ij} is commonly defined as a component of the gravitational field produced at (x_i, y_i, z_i) by a point mass or prism located at (x_j, y_j, z_j) , with unit density. On the other hand, g_{ij} is commonly defined as a component of the magnetic induction field produced at (x_i, y_i, z_i) by a dipole or prism located at (x_j, y_j, z_j) , with unit magnetization intensity, when the observed potential field is magnetic data.

For all harmonic functions discussed above, the sensitivity matrix \mathbf{G} (equation 3) is always dense. For scattered potential-field data, \mathbf{G} does not have a well-defined structure, regardless of whether the spatial distribution of the equivalent sources is set. Nevertheless, for the particular case in which (i) there is a single equivalent source right below each potential-field datum and (ii) both data and sources rely on planar and regularly spaced grids, Takahashi et al. (2020, 2022) show that \mathbf{G} assumes a block-Toeplitz Toeplitz-block (BTTB) structure. In this case, the product of \mathbf{G} and an arbitrary vector can be efficiently computed via 2D fast Fourier transform as a discrete convolution.

1.3 General formulation

A general formulation for almost all equivalent-layer methods can be achieved by first considering that the $P \times 1$ parameter vector \mathbf{p} (equation 3) can be reparameterized into a $Q \times 1$ vector \mathbf{q} according to:

$$\mathbf{p} = \mathbf{H} \mathbf{q}, \quad (9)$$

where \mathbf{H} is a $P \times Q$ matrix. This reparameterization is usually defined with $Q \ll P$ to reduce the original number of parameters. The predicted data vector \mathbf{d}_p (equation 3) can then be rewritten as follows:

$$\mathbf{d}_p = \mathbf{G} \mathbf{H} \mathbf{q}. \quad (10)$$

Then, the problem of estimating a parameter vector $\tilde{\mathbf{p}}$ minimizing a length measure of the difference between \mathbf{d}_p (equation 3) and \mathbf{d}_o is replaced by that of estimating an auxiliary vector $\tilde{\mathbf{q}}$ minimizing the goal function

$$\Gamma(\mathbf{q}) = \Phi(\mathbf{q}) + \mu \Theta(\mathbf{q}), \quad (11)$$

which is a combination of particular measures of length given by

$$\Phi(\mathbf{q}) = \|\mathbf{W}_d (\mathbf{d}_o - \mathbf{G} \mathbf{H} \mathbf{q})\|_2^2, \quad (12)$$

and

$$\Theta(\mathbf{q}) = \|\mathbf{W}_q (\mathbf{q} - \mathbf{q}_a)\|_2^2, \quad (13)$$

where μ is a positive scalar controlling the trade-off between $\Phi(\mathbf{q})$ and $\Theta(\mathbf{q})$; $\|\cdot\|_2$ is the 2-norm (or Euclidean norm); \mathbf{W}_q is a matrix imposing prior information on \mathbf{q} given by

$$\mathbf{W}_q = \mathbf{W}_p \mathbf{H}, \quad (14)$$

with \mathbf{W}_p being a matrix imposing prior information on \mathbf{p} ; \mathbf{q}_a is a $Q \times 1$ vector of reference values for \mathbf{q} satisfying

$$\mathbf{p}_a = \mathbf{H} \mathbf{q}_a, \quad (15)$$

with \mathbf{p}_a being a $P \times 1$ vector containing reference values for the original parameter vector \mathbf{p} and \mathbf{W}_d is an $D \times D$ matrix defining the relative importance of each observed datum d_i^o . After obtaining an estimate $\tilde{\mathbf{q}}$ for the reparameterized parameter vector \mathbf{q} (equation 9) minimizing $\Gamma(\mathbf{q})$ (equation 11), the estimate $\tilde{\mathbf{p}}$ for the original parameter vector (equation 3) is computed by

$$\tilde{\mathbf{p}} = \mathbf{H} \tilde{\mathbf{q}}. \quad (16)$$

The reparameterized vector $\tilde{\mathbf{q}}$ is obtained by first computing the gradient of $\Gamma(\mathbf{q})$,

$$\nabla \Gamma(\mathbf{q}) = -2 \mathbf{H}^\top \mathbf{G}^\top \mathbf{W}_d^\top \mathbf{W}_d (\mathbf{d}_o - \mathbf{d}_p) + 2 \mu \mathbf{H}^\top \mathbf{W}_p^\top \mathbf{W}_p \mathbf{H} (\mathbf{q} - \mathbf{q}_a). \quad (17)$$

Then, by considering that $\nabla \Gamma(\tilde{\mathbf{q}}) = \mathbf{0}$ (equation 17), where $\mathbf{0}$ is a vector of zeros, as well as adding and subtracting the term $(\mathbf{H}^\top \mathbf{G}^\top \mathbf{W}_d^\top \mathbf{W}_d \mathbf{G} \mathbf{H}) \mathbf{q}_a$, we obtain

$$\tilde{\mathbf{q}} = \mathbf{q}_a + \mathbf{B} (\mathbf{d}_o - \mathbf{G} \mathbf{H} \mathbf{q}_a), \quad (18)$$

where

$$\mathbf{B} = \left(\mathbf{H}^\top \mathbf{G}^\top \mathbf{W}_d^\top \mathbf{W}_d \mathbf{G} \mathbf{H} + \mu \mathbf{H}^\top \mathbf{W}_p^\top \mathbf{W}_p \mathbf{H} \right)^{-1} \mathbf{H}^\top \mathbf{G}^\top \mathbf{W}_d^\top \mathbf{W}_d, \quad (19)$$

or, equivalently (Menke, 2018, p. 62),

$$\mathbf{B} = \left(\mathbf{H}^\top \mathbf{W}_p^\top \mathbf{W}_p \mathbf{H} \right)^{-1} \mathbf{H}^\top \mathbf{G}^\top \left[\mathbf{G} \mathbf{H} \left(\mathbf{H}^\top \mathbf{W}_p^\top \mathbf{W}_p \mathbf{H} \right)^{-1} \mathbf{H}^\top \mathbf{G}^\top + \mu \left(\mathbf{W}_d^\top \mathbf{W}_d \right)^{-1} \right]^{-1}. \quad (20)$$

Evidently, we have considered that all inverses exist in equations 19 and 20. Matrix \mathbf{B} defined by equation 19 is commonly used for the cases in which $D > P$, i.e., when there are more data d_i^o than parameters p_j (overdetermined problems). On the other hand, for the cases in which there are more parameters than data (underdetermined problems), matrix \mathbf{B} is usually defined according to equation 20.

PAREI AQUÍ

The great majority of equivalent-layer methods use non-null \mathbf{p}_r , μ , \mathbf{P} and \mathbf{D} (equations ??-??) to regularize the inverse problem, impose some constraint on the estimated parameter vector $\tilde{\mathbf{p}}$ and define weights for the observed data according to their (presumed) experimental error. Nonetheless, we limit our analysis to the cases in which \mathbf{p}_r have null elements, matrices \mathbf{D} and \mathbf{P} are equal to the identity and $\mu = 0$ because these terms do not significantly impact the computational performance. Hence, we consider that the estimate $\tilde{\mathbf{p}}$ for overdetermined problems ($I > J$) is obtained by solving the linear system

$$\left(\mathbf{A}^\top \mathbf{A} \right) \tilde{\mathbf{p}} = \mathbf{A}^\top \mathbf{d}, \quad (21)$$

whereas the solution for underdetermined problems ($I < J$) is obtained in two steps by first solving a linear system and then computing a matrix-vector product as follows

$$\begin{aligned} \left(\mathbf{A} \mathbf{A}^\top \right) \mathbf{u} &= \mathbf{d} \\ \tilde{\mathbf{p}} &= \mathbf{A}^\top \mathbf{u} \end{aligned} \quad (22)$$

where \mathbf{u} is a dummy vector.

Let us first consider the following generic linear system

$$\mathbf{H}\mathbf{v} = \mathbf{h}, \quad (23)$$

where \mathbf{H} is a square matrix and \mathbf{v} is a vector of unknowns. Here, we set $\mathbf{H} = \mathbf{A}^\top \mathbf{A}$ and $\mathbf{h} = \mathbf{A}^\top \mathbf{d}$ for solving the overdetermined system (equation 21) or $\mathbf{H} = \mathbf{A} \mathbf{A}^\top$ and $\mathbf{h} = \mathbf{d}$ for solving the underdetermined system (equation 22).

An estimate $\tilde{\mathbf{v}}$ for the unknown vector \mathbf{v} can be obtained by using direct methods, such as LU, Cholesky or QR factorization (Golub and Loan, 2013, p. 111, 163, 246), or iterative methods, such as Gauss-Seidel, Successive Over-Relaxation (SOR) or Conjugate Gradient (Golub and Loan, 2013, p. 612, 619, 625), for example. Here, we briefly discuss the Cholesky factorization and some general aspects of iterative methods.

The Cholesky factorization (Golub and Loan, 2013, p. 163) presumes the existence of an upper triangular matrix \mathbf{G} (the Cholesky factor) such that $\mathbf{H} = \mathbf{G} \mathbf{G}^\top$. Given \mathbf{G} , an estimate $\tilde{\mathbf{v}}$ for the vector of unknowns \mathbf{v} (equation 23) can be obtained by first solving the triangular system

$$\mathbf{G} \mathbf{w} = \mathbf{h}, \quad (24)$$

where \mathbf{w} is a dummy vector, and then computing

$$\mathbf{G}^\top \mathbf{w} = \tilde{\mathbf{v}}. \quad (25)$$

Iterative methods can be roughly defined as a process that starts with an initial approximation $\tilde{\mathbf{v}}_0$ and then produces ever-better approximations $\tilde{\mathbf{v}}_\ell$ for the unknown vector \mathbf{v} (equation 23). At each iteration, it is necessary to evaluate a predefined convergence criterion (CC). If the CC is satisfied, the current approximation is considered as the solution $\tilde{\mathbf{v}}$ of the linear system and the algorithm stops. Otherwise, the current approximation $\tilde{\mathbf{v}}_\ell$ is used to compute a new approximation $\tilde{\mathbf{v}}_{\ell+1}$ and the CC is evaluated again. A pseudo-code for this generic iterative method can be defined as follows:

Algorithm 1: Generic pseudo-code for iterative methods.

Initialization :

```

1  $\ell = 0$  ;
2  $\tilde{\mathbf{v}}_\ell \leftarrow \tilde{\mathbf{v}}_0$  ;
3 Compute auxiliary variables ;
4 while CC is not satisfied do
5   Compute  $\tilde{\mathbf{v}}_{\ell+1}$  ;
6   Compute auxiliary variables ;
7    $\tilde{\mathbf{v}}_\ell \leftarrow \tilde{\mathbf{v}}_{\ell+1}$  ;
8    $\ell \leftarrow \ell + 1$  ;
9 end
10  $\tilde{\mathbf{v}} \leftarrow \tilde{\mathbf{v}}_\ell$ 

```

2 COMPUTATIONAL STRATEGIES

Two important factors affecting the efficiency of a given matrix algorithm are the storage and amount of required arithmetic. Here, we quantify this last factor by counting flops. A flop is a floating point addition, subtraction, multiplication or division (Golub and Loan, 2013, p. 12–14).

NÃO SEI SE TÁ BOM AQUI: To investigate the efficiency of equivalent-layer methods, we consider how they (i) set up and (ii) solve the linear inverse problem to estimate the physical property distribution on the equivalent layer, as well as (iii) perform potential field transformations (equation 4).

We focus on the overall strategies used by the selected methods

2.1 Moving-data windows

Split the observed data $d_i, i \in \{1, \dots, I\}$, into M overlapping subsets (or data windows) formed by I^m data each, $m \in \{1, \dots, M\}$.

The number of data I^m forming the data windows are not necessarily equal to each other.

The data forming a given window are usually adjacent to each other.

Each window has an $I^m \times 1$ observed data vector \mathbf{d}^m .

Let each data window be approximated by a local equivalent layer composed of J^m sources, so that its predicted data vector is given by

$$\mathbf{f}^m = \mathbf{A}^m \mathbf{p}^m, \quad (26)$$

where \mathbf{p}^m is a $J^m \times 1$ vector containing the scalar physical properties of the equivalent sources in the m -th subset and \mathbf{A}^m is an $I^m \times J^m$ matrix whose elements are computed with equation ?? by using only the coordinates of the observed data and equivalent sources in the m -th subset.

The main advantage of this approach is that the estimated parameter vector $\tilde{\mathbf{p}}$ is not obtained by solving the full linear system, but several smaller ones.

Leão and Silva (1989) presented a pioneer work using this approach.

Their method requires a regularly-spaced grid of observed data on a horizontal plane z_0 .

The data windows are defined by square local grids of $\sqrt{I'} \times \sqrt{I'}$ adjacent points, all of them having the same number of points I' .

The equivalent sources in the m -th data window are located below the observation plane, at a constant vertical distance Δz_0 . They are arranged on a regular grid of $\sqrt{J'} \times \sqrt{J'}$ adjacent points following the same grid pattern of the observed data. The local grid of sources for all data windows have the same number of elements J' . Besides, they are vertically aligned, but expands the limits of their corresponding data windows, so that $J' > I'$.

Because of this spatial configuration of observed data and equivalent sources, we have that $\mathbf{A}^m = \mathbf{A}'$ (equation 26) for all data windows, where \mathbf{A}' is a constant matrix.

By omitting the regularization and normalization strategies used by Leão and Silva (1989), their method consists in combining equations 22 and 4 to directly compute the transformed potential field t_c^m at the central point of each data window as follows:

$$t_c^m = (\mathbf{A}' \mathbf{b}')^\top \left[\mathbf{A}' (\mathbf{A}')^\top \right]^{-1} \mathbf{d}^m, \quad m \in \{1, \dots, M\}, \quad (27)$$

where \mathbf{b}' is a $J' \times 1$ vector with elements computed by equation ?? by using all equivalent sources in the m -th subset and only the coordinate of the central point in the m -th data window. Due to the presumed spatial configuration of the observed data and equivalent sources, \mathbf{b}' is the same for all data windows.

Their method can be outlined by the following pseudo-code:

Algorithm 2: Generic pseudo-code for the method proposed by Leão and Silva (1989).

Initialization :

```

1 Set the  $\mathbf{d}^m$  for each data window,  $m \in \{1, \dots, M\}$  ;
2 Set the constant depth  $z_0 + \Delta z_0$  for all equivalent sources ;
3 Compute the vector  $\mathbf{b}'$  associated with the desired potential-field transformation ;
4 Compute the matrix  $\mathbf{A}'$  ;
5 Compute  $(\mathbf{A}'\mathbf{b}')^\top [\mathbf{A}' (\mathbf{A}')^\top]^{-1}$  ;
6  $\ell = 1$  ;
7 while  $\ell < M$  do
8   | Compute  $t_c^m$  (equation 27) ;
9   |  $\ell \leftarrow \ell + 1$  ;
10 end

```

163

164 Note that Leão and Silva (1989) directly compute the transformed potential t_c^m at the central point of each
 165 data window without explicitly computing and storing \mathbf{p}^m (equation 26). It means that their method allows
 166 computing a single potential-field transformation. A different transformation or the same one evaluated at
 167 different points require running their moving-data window method again.

168 Soler and Uieda (2021) generalized the method proposed by Leão and Silva (1989) for irregularly spaced
 169 data on an undulating surface.

The overall steps of their method are defined by the following pseudo-code:

Algorithm 3: Generic pseudo-code for the method proposed by Soler and Uieda (2021).

Initialization :

```

1 Set the  $\mathbf{d}^m$  for each data window,  $m \in \{1, \dots, M\}$  ;
2 Set the depth scheme for all equivalent sources ;
3 Compute the vector  $\mathbf{b}'$  associated with the desired potential-field transformation ;
4 Compute the matrix  $\mathbf{A}'$  ;
5 Compute  $(\mathbf{A}'\mathbf{b}')^\top [\mathbf{A}' (\mathbf{A}')^\top]^{-1}$  ;
6  $\ell = 1$  ;
7 while  $\ell < M$  do
8   | Compute  $t_c^m$  (equation 27) ;
9   |  $\ell \leftarrow \ell + 1$  ;
10 end

```

170

171 PAREI AQUI

172 2.2 Column update

173 Cordell (1992)

174 Guspí and Novara (2009)

175 **2.3 Row update**

176 Algebraic reconstruction techniques (ART) van der Sluis and van der Vorst (2004)

177 Mendonça and Silva (1994)

178 **2.4 Reparameterization**

179 Barnes and Lumley (2011)

180 Oliveira Jr. et al. (2013)

181 Mendonça (2020)

182 **2.5 Wavelet compression**

183 Li and Oldenburg (2010)

184 **2.6 Iterative methods using the original A**

185 Xia and Sprowl (1991)

186 Xia et al. (1993)

187 Siqueira et al. (2017)

188 Jirigalatu and Ebbing (2019)

189 **2.7 Discrete convolution**

190 Takahashi et al. (2020)

191 Takahashi et al. (2022)

3 TEXTO ANTIGO

192 Leão and Silva (1989) reduced the total processing time and memory usage of equivalent-layer technique
 193 by means of a moving data-window scheme. A small moving data window with N_w observations and
 194 a small equivalent layer with M_w equivalent sources ($M_w > N_w$) located below the observations are
 195 established. For each position of a moving-data window, Leão and Silva (1989) estimate a stable solution
 196 \mathbf{p}_w^* by using a data-space approach with the zeroth-order Tikhonov regularization (?), i.e.,

$$\left(\mathbf{A}_w \mathbf{A}_w^\top + \mu \mathbf{I}\right) \mathbf{w} = \mathbf{d}_w^o, \quad (28a)$$

$$\mathbf{A}_w^\top \mathbf{w} = \mathbf{p}_w^*, \quad (28b)$$

197 where \mathbf{w} is a dummy vector, μ is a regularizing parameter, \mathbf{d}_w^o is an N_w -dimensional vector containing
 198 the observed potential-field data, \mathbf{A}_w is an $N_w \times M_w$ sensitivity matrix related to a moving-data window, \mathbf{I}
 199 is an identity matrix of order N_w and the superscript \top stands for a transpose. After estimating an $M_w \times 1$
 200 parameter vector \mathbf{p}_w^* (equation 28b) the desired transformation of the data is only calculated at the central

point of each moving-data window, i.e.:

$$\hat{t}_k = \mathbf{t}_k^\top \mathbf{p}_w^*, \quad (29)$$

where \hat{t}_k is the transformed data calculated at the central point k of the data window and \mathbf{t}_k is an $M_w \times 1$ vector whose elements form the k th row of the $N_w \times N_w$ matrix of Green's functions \mathbf{T} (equation ??) of the desired linear transformation of the data.

By shifting the moving-data window with a shift size of one data spacing, a new position of a data window is set up. Next, the aforementioned process (equations 28b and 29) is repeated for each position of a moving-data window, until the entire data have been processed. Hence, instead of solving a large inverse problem, Leão and Silva (1989) solve several much smaller ones.

To reduce the size of the linear system to be solved, Soler and Uieda (2021) adopted the same strategy proposed, originally, by Leão and Silva (1989) of using a small moving-data window sweeping the whole data. In Leão and Silva (1989), a moving-data window slides to the next adjacent data window following a sequential movement, the predicted data is calculated inside the data window and the desired transformation are only calculated at the center of the moving-data window. Unlike Leão and Silva (1989), Soler and Uieda (2021) do not adopt a sequential order of the data windows; rather, they adopt a randomized order of windows in the iterations of the gradient-boosting algorithm (?, ? and ?). The gradient-boosting algorithm in Soler and Uieda (2021) estimates a stable solution using the data and the equivalent sources that fall within a moving-data window; however, it calculates the predicted data and the residual data in the whole survey data. Next, the residual data that fall within a new position of the data window is used as input data to estimate a new stable solution within the data window which in turn is used to calculate a new predicted data and a new residual data in the whole survey data. Finally, unlike Leão and Silva (1989), in Soler and Uieda (2021) neither the data nor the equivalent sources need to be distributed in regular grids. Indeed, Leão and Silva (1989) built their method using regular grids, but in fact regular grids are not necessary. Regarding the equivalent-source layout, Soler and Uieda (2021) proposed the block-averaged sources locations in which the survey area is divided into horizontal blocks and one single equivalent source is assigned to each block. Each single source per block is placed over the layer with its horizontal coordinates given by the average horizontal positions of observation points. According to Soler and Uieda (2021), the block-averaged sources layout reduces the number of equivalent sources significantly and the gradient-boosting algorithm provides even greater efficiency in terms of data fitting.

3.0.1 The equivalent-data concept

To reduced the total processing time and memory usage of equivalent-layer technique, Mendonça and Silva (1994) proposed a strategy called 'equivalent data concept'. The equivalent data concept is grounded on the principle that there is a subset of redundant data that does not contribute to the final solution and thus can be dispensed. Conversely, there is a subset of observations, called equivalent data, that contributes effectively to the final solution and fits the remaining observations (redundant data). Iteratively, Mendonça and Silva (1994) selected the subset of equivalent data that is substantially smaller than the original dataset. This selection is carried out by incorporating one data point at a time.

According to Mendonça and Silva (1994), the number of equivalent data is about one-tenth of the total number of observations. These authors used the equivalent data concept to carry out an interpolation of gravity data. They showed a reduction of the total processing time and memory usage by, at least, two orders of magnitude as opposed to using all observations in the interpolation process via the classical equivalent-layer technique.

3.0.2 The wavelet compression and lower-dimensional subspace

For large data sets, the sensitivity matrix \mathbf{A} (equation 3) is a drawback in applying the equivalent-layer technique because it is a large and dense matrix.

? transformed a large and full sensitivity matrix into a sparse one by using fast wavelet transforms. In the wavelet domain, ? applied a 2D wavelet transform to each row and column of the original sensitivity matrix \mathbf{A} to expand it in the wavelet bases. This operation can be done by premultiplying the original sensitivity matrix \mathbf{A} by a matrix representing the 2D wavelet transform \mathbf{W}_2 and then the resulting is postmultiplied by the transpose of \mathbf{W}_2 (i.e., \mathbf{W}_2^\top).

$$\tilde{\mathbf{A}} = \mathbf{W}_2 \mathbf{A} \mathbf{W}_2^\top, \quad (30)$$

where $\tilde{\mathbf{A}}$ is the expanded original sensitivity matrix in the wavelet bases with many elements zero or close to zero. Next, the matrix $\tilde{\mathbf{A}}$ is replaced by its sparse version $\tilde{\mathbf{A}}_s$ in the wavelet domain which in turn is obtained by retaining only the large elements of the $\tilde{\mathbf{A}}$. Thus, the elements of $\tilde{\mathbf{A}}$ whose amplitudes fall below a relative threshold are discarded. In ?, the original sensitivity matrix \mathbf{A} is high compressed resulting in a sparse matrix $\tilde{\mathbf{A}}_s$ with a few percent of nonzero elements and the the inverse problem is solved in the wavelet domain by using $\tilde{\mathbf{A}}_s$ and a incomplete conjugate gradient least squares, without an explicit regularization parameter and a limited number of iterations. The solution is obtained by solving the following linear system

$$\tilde{\mathbf{A}}_L^\top \tilde{\mathbf{A}}_L \tilde{\mathbf{p}}_L^* = \tilde{\mathbf{A}}_L^\top \tilde{\mathbf{d}}^o, \quad (31)$$

where $\tilde{\mathbf{p}}_L^*$ is obtained by solving the linear system given by equation 31,

$$\tilde{\mathbf{A}}_L = \tilde{\mathbf{A}}_s \tilde{\mathbf{L}}^{-1}, \quad (32a)$$

$$\tilde{\mathbf{p}}_L = \tilde{\mathbf{L}} \tilde{\mathbf{p}}, \quad (32b)$$

$$\tilde{\mathbf{d}}^o = \mathbf{W}_2 \mathbf{d}^o, \quad (32c)$$

where $\tilde{\mathbf{L}}$ is a diagonal and invertible weighting matrix representing the finite-difference approximation in the wavelet domain. Finally, the distribution over the equivalent layer in the space domain \mathbf{p} is obtained by applying an inverse wavelet transform in two steps, i.e.:

$$\tilde{\mathbf{p}} = \tilde{\mathbf{L}}^{-1} \tilde{\mathbf{p}}_L^*, \quad (33)$$

and

$$\mathbf{p} = \mathbf{W}_2 \tilde{\mathbf{p}}. \quad (34)$$

Although the data misfit quantifying the difference between the observed and predicted data by the equivalent source is calculated in the wavelet domain, we understand that the desired transformation is calculated via equation ?? which uses a full matrix of Green's functions \mathbf{T} .

? used the equivalent-layer technique with a wavelet compression to perform an upward continuation of total-field anomaly between uneven surfaces. For regularly spaced grid of data, ? reported that high compression ratios are achieved with insignificant loss of accuracy. As compared to the upward-continued total-field anomaly by equivalent layer using the dense matrix, ?'s (?) approach, using the Daubechies wavelet, decreased CPU (central processing unit) time by up to two orders of magnitude.

? overcame the solution of intractable large-scale equivalent-layer problem by using the subspace method (e.g., ?, ?; ?, ?; ?, ?; ?, ?). The subspace method reduces the dimension of the linear system of equations to be solved. Given a higher-dimensional space (e.g., M -dimensional model space, \mathbb{R}^M), there exists many lower-dimensional subspaces (e.g., Q -dimensional subspace) of \mathbb{R}^M . The linear inverse problem related to the equivalent-layer technique consists in finding an M -dimension parameter vector $\mathbf{p} \in \mathbb{R}^M$ which adequately fits the potential-field data. The subspace method looks for a parameter vector who lies in a Q -dimensional subspace of \mathbb{R}^M which, in turn, is spanned by a set of Q vectors $\mathbf{v}_i = 1, \dots, Q$, where $\mathbf{v}_i \in \mathbb{R}^M$. In matrix notation, the parameter vector in the subspace method can be written as

$$\mathbf{p} = \mathbf{V} \boldsymbol{\alpha}, \quad (35)$$

where \mathbf{V} is an $M \times Q$ matrix whose columns $\mathbf{v}_i = 1, \dots, Q$ form a basis vectors for a subspace Q of \mathbb{R}^M . In equation 35, the parameter vector \mathbf{p} is defined as a linear combination in the space spanned by Q basis vectors $\mathbf{v}_i = 1, \dots, Q$ and $\boldsymbol{\alpha}$ is a Q -dimensional unknown vector to be determined. The main advantage of the subspace method is that the linear system of M equations in M unknowns to be originally solved is reduced to a new linear system of Q equations in Q unknowns which requires much less computational effort since $Q \ll M$, i.e.:

$$\mathbf{V}^T \mathbf{A}^T \mathbf{A} \mathbf{V} \boldsymbol{\alpha}^* = \mathbf{V}^T \mathbf{d}^o. \quad (36)$$

To avoid the storage of matrices \mathbf{A} and \mathbf{V} , ? evaluates an element of the matrix $\mathbf{A}\mathbf{V}$ by calculating the dot product between the row of matrix \mathbf{A} and the column of the matrix \mathbf{B} . After estimating $\boldsymbol{\alpha}^*$ (equation 36) belonging to a Q -dimensional subspace of \mathbb{R}^M , the distribution over the equivalent layer \mathbf{p} in the \mathbb{R}^M is obtained by applying equation 35. The choice of the Q basis vectors $\mathbf{v}_i = 1, \dots, Q$ (equation 35) in the subspace method is not strict. ?, for example, chose the eigenvectors yielded by applying the singular value decomposition of the matrix containing the gridded data set. The number of eigenvectors used to form basis vectors will depend on the singular values.

The proposed subspace method for solving large-scale equivalent-layer problem by ? was applied to estimate the mass excess or deficiency caused by causative gravity sources.

3.0.3 The quadtree discretization

To make the equivalent-layer technique tractable, ? also transformed the dense sensitivity matrix \mathbf{A} (equation 3) into a sparse matrix. In ?, a sparse version of the sensitivity matrix is achieved by grouping equivalent sources (e.g., they used prisms) distant from an observation point together to form a larger prism or larger block. Each larger block has averaged physical properties and averaged top- and bottom-surfaces of the grouped smaller prisms (equivalent sources) that are encompassed by the larger block. The authors called it the 'larger averaged block' and the essence of their method is the reduction in the number of equivalent sources, which means a reduction in the number of parameters to be estimated implying in model dimension reduction.

The key of the ?'s (?) method is the algorithm for deciding how to group the smaller prisms. In practice, these authors used a recursive bisection process that results in a quadtree discretization of the equivalent-layer model.

By using the quadtree discretization, ? were able to jointly process multiple components of airborne gravity-gradient data using a single layer of equivalent sources. To our knowledge, ? are the pioneers on processing full-tensor gravity-gradient data jointly. In addition to computational feasibility, ?'s (?) method reduces low-frequency noise and can also remove the drift in time-domain from the survey data. Those

310 authors stressed that the G_{zz} -component calculated through the single estimated equivalent-layer model
 311 projected on a grid at a constant elevation by inverting full gravity-gradient data has the low-frequency error
 312 reduced by a factor of 2.4 as compared to the inversion of an individual component of the gravity-gradient
 313 data.

314 3.0.4 The reparametrization of the equivalent layer

315 Oliveira Jr. et al. (2013) reparametrized the whole equivalent-layer model by a piecewise bivariate-
 316 polynomial function defined on a set of Q equivalent-source windows. In Oliveira Jr. et al.'s (2013)
 317 approach, named polynomial equivalent layer (PEL), the parameter vector within the k th equivalent-source
 318 window \mathbf{p}^k can be written in matrix notation as

$$\mathbf{p}^k = \mathbf{B}^k \mathbf{c}^k, \quad k = 1 \dots Q, \quad (37)$$

319 where \mathbf{p}^k is an M_w -dimensional vector containing the physical-property distribution within the k th
 320 equivalent-source window, \mathbf{c}^k is a P -dimensional vector whose l th element is the l th coefficient of the
 321 α th-order polynomial function and \mathbf{B}^k is an $M_w \times P$ matrix containing the first-order derivative of the
 322 α th-order polynomial function with respect to one of the P coefficients.

323 By using a regularized potential-field inversion, Oliveira Jr. et al. (2013) estimates the polynomial
 324 coefficients for each equivalent-source window by solving the following linear system

$$\left(\mathbf{B}^\top \mathbf{A}^\top \mathbf{A} \mathbf{B} + \mu \mathbf{I} \right) \mathbf{c}^* = \mathbf{B}^\top \mathbf{A}^\top \mathbf{d}^o, \quad (38)$$

325 where μ is a regularizing parameter, \mathbf{c}^* is an estimated H -dimensional vector containing all coefficients
 326 describing all polynomial functions within all equivalent-source windows which compose the entire
 327 equivalent layer, \mathbf{I} is an identity matrix of order H ($H = PQ$) and \mathbf{B} is an $M \times H$ block diagonal matrix
 328 such that the main-diagonal blocks are \mathbf{B}^k matrices (equation 37) and all off-diagonal blocks are zero
 329 matrices. For ease of the explanation of equation 38, we keep only the zeroth-order Tikhonov regularization
 330 and omitting the first-order Tikhonov regularization (?) which was also used by Oliveira Jr. et al. (2013).

331 The main advantage of the PEL is solve H -dimensional system of equations (equation 38), where H
 332 totalizes the number of polynomial coefficients composing all equivalent-source windows, requiring a
 333 lower computational effort since $H \ll N$. To avoid the storage of matrices \mathbf{A} and \mathbf{B} , Oliveira Jr. et al.
 334 (2013) evaluate an element of the matrix $\mathbf{A}\mathbf{B}$ by calculating the dot product between the row of matrix \mathbf{A}
 335 and the column of the matrix \mathbf{B} . After estimating all polynomial coefficients of all windows, the estimated
 336 coefficients (\mathbf{c}^* in equation 38) are transformed into a single physical-property distribution encompassing
 337 the entire equivalent layer.

338 As stated by Oliveira Jr. et al. (2013), the computational efficiency of PEL approach stems from the fact
 339 that the total number of polynomial coefficients H required to depict the physical-property distribution
 340 within the equivalent layer is generally much smaller than the number of equivalent sources. Consequently,
 341 this leads to a considerably smaller linear system that needs to be solved. Hence, the main strategy of
 342 polynomial equivalent layer is the model dimension reduction.

343 The polynomial equivalent layer was applied to perform upward continuations of gravity and magnetic
 344 data and reduction to the pole of magnetic data.

3.0.5 The iterative scheme without solving a linear system

There exists a class of methods that iteratively estimate the distribution of physical properties within an equivalent layer without the need to solve linear systems. The method initially introduced by Cordell (1992) and later expanded upon by Guspí and Novara (2009) updates the physical property of sources, located beneath each potential-field data, by removing the maximum residual between the observed and fitted data. In addition, Xia and Sprowl (1991) and Xia et al. (1993) have developed efficient iterative algorithms for updating the distribution of physical properties within the equivalent layer in the wavenumber and space domains, respectively. Specifically, in Xia and Sprowl's (1991) method the physical-property distribution is updated by using the ratio between the squared depth to the equivalent source and the gravitational constant multiplied by the residual between the observed and predicted observation at the measurement station. Neither of these methods solve linear systems.

Following this class of methods of iterative equivalent-layer technique that does not solve linear systems, Siqueira et al. (2017) developed a fast iterative equivalent-layer technique for processing gravity data in which the sensitivity matrix \mathbf{A} (equation 3) is replaced by a diagonal matrix $N \times N$, i.e.:

$$\tilde{\tilde{\mathbf{A}}} = 2 \pi \gamma \Delta \mathbf{S}^{-1}, \quad (39)$$

where γ is Newton's gravitational constant and $\Delta \mathbf{S}^{-1}$ is a diagonal matrix of order N whose diagonal elements Δs_i , $i = 1, \dots, N$ are the element of area centered at the i th horizontal coordinates of the i th observation point. The physical foundations of Siqueira et al.'s (2017) method rely on two constraints: i) the excess of mass; and ii) the positive correlation between the gravity observations and the mass distribution over the equivalent layer.

Although Siqueira et al.'s (2017) method does not solve any linear system of equations, it can be theoretically explained by solving the following linear system at the k th iteration:

$$\tilde{\tilde{\mathbf{A}}}^T \tilde{\tilde{\mathbf{A}}} \Delta \hat{\mathbf{p}}^k = \tilde{\tilde{\mathbf{A}}}^T \mathbf{r}^k, \quad (40)$$

where \mathbf{r}^k is an N -dimensional residual vector whose i th element is calculated by subtracting the i th observed data d_i^o from the i th fitted data d_i^k at the k th iteration, i.e.,

$$r_i^k = d_i^o - d_i^k. \quad (41)$$

and $\Delta \hat{\mathbf{p}}^k$ is an estimated N -dimensional vector of parameter correction.

Because $\tilde{\tilde{\mathbf{A}}}$, in equation 40, is a diagonal matrix (equation 39), the parameter correction estimate is directly calculated without solving system of linear equations, and thus, an i th element of $\Delta \hat{\mathbf{p}}^k$ is directly calculated by

$$\Delta \hat{p}_i^k = \frac{\Delta s_i r_i^k}{2 \pi \gamma}. \quad (42)$$

The mass distribution over the equivalent layer is updated by:

$$\hat{p}_i^{k+1} = \hat{p}_i^k + \Delta \hat{p}_i^k. \quad (43)$$

Siqueira et al.'s (2017) method starts from a mass distribution on the equivalent layer, whose i th mass p_i^o is proportional to the i th observed data d_i^o , i.e.,

$$p_i^o = \frac{\Delta s_i d_i^o}{2 \pi \gamma}. \quad (44)$$

Siqueira et al. (2017) applied their fast iterative equivalent-layer technique to interpolate, calculate the horizontal components, and continue upward (or downward) gravity data.

For jointly process two gravity gradient components, Jiriglatu and Ebbing (2019) used the Gauss-FFT for forward calculation of potential fields in the wavenumber domain combined with Landweber's iteration coupled with a mask matrix \mathbf{M} to reduce the edge effects without increasing the computation cost. The mask matrix \mathbf{M} is defined in the following way: if the corresponding pixel does not contain the original data, the element of \mathbf{M} is set to zero; otherwise, it is set to one. The k th Landweber iteration is given by

$$\mathbf{p}_{k+1} = \mathbf{p}_k + \omega \left[\mathbf{A}_1^\top (\mathbf{d}_1 - \mathbf{M} \mathbf{A}_1 \mathbf{p}_k) + \mathbf{A}_2^\top (\mathbf{d}_2 - \mathbf{M} \mathbf{A}_2 \mathbf{p}_k) \right], \quad (45)$$

where ω is a relaxation factor, \mathbf{d}_1 and \mathbf{d}_2 are the two gravity gradient components and \mathbf{A}_1 and \mathbf{A}_2 are the corresponding gravity gradient kernels. Jiriglatu and Ebbing (2019) applied their method for processing two horizontal curvature components of Falcon airborne gravity gradient.

3.0.6 The convolutional equivalent layer with BTTB matrices

? (? , ?) introduced the convolutional equivalent layer for gravimetric and magnetic data processing, respectively.

? demonstrated that the sensitivity matrix \mathbf{A} (equation 3) associated with a planar equivalent layer formed by a set of point masses, each one directly beneath each observation point and considering a regular grid of observation points at a constant height has a symmetric block-Toeplitz Toeplitz-block (BTTB) structure. A symmetric BTTB matrix has, at least, two attractive properties. The first one is that it can be defined by using only the elements forming its first column (or row). The second attractive property is that any BTTB matrix can be embedded into a symmetric Block-Circulant Circulant-Block (BCCB) matrix. This means that the full sensitivity matrix \mathbf{A} (equation 3) can be completely reconstruct by using the first column of the BCCB matrix only. In what follows, ? computed the forward modeling by using only a single equivalent source. Specifically, it is done by calculating the eigenvalues of the BCCB matrix that can be efficiently computed by using only the first column of the BCCB matrix via 2D fast Fourier transform (2D FFT). By comparing with the classic approach in the Fourier domain, the convolutional equivalent layer for gravimetric data processing proposed by ? performed upward- and downward-continue gravity data with a very small border effects and noise amplification.

By using the original idea of the convolutional equivalent layer proposed by ? for gravimetric data processing, ? developed the convolutional equivalent layer for magnetic data processing. By assuming a regularly spaced grid of magnetic data at a constant height and a planar equivalent layer of dipoles, ? proved that the sensitivity matrix linked with this layer possess a BTTB structure in the specific scenario where each dipole is exactly beneath each observed magnetic data point. ? used a conjugate gradient least-squares (CGLS) algorithm which does not require an inverse matrix or matrix-matrix multiplication. Rather, it only requires matrix-vector multiplications per iteration, which can be effectively computed using the 2D FFT as a discrete convolution. The matrix-vector product only uses the elements that constitute the

first column of the associated BTTB matrix, resulting in computational time and memory savings. ? (?) showed the robustness of the convolutional equivalent layer in processing magnetic survey that violates the requirement of regular grids in the horizontal directions and flat observation surfaces.

The matrix-vector product in ? (? , ?) (e.g., $\mathbf{d} = \mathbf{A}\mathbf{p}$, such as in equation 3) is the main issue to be solved. To solve it efficiently, these authors invoked the auxiliary linear system

$$\mathbf{w} = \mathbf{C}\mathbf{v} , \quad (46)$$

where \mathbf{w} and \mathbf{v} are, respectively, vectors of data and parameters completed by zeros and \mathbf{C} is a BCCB matrix formed by $2Q \times 2Q$ blocks, where each block \mathbf{C}_q , $q = 0, \dots, Q - 1$, is a $2P \times 2P$ circulant matrix. The first column of \mathbf{C} is obtained by rearranging the first column of the sensitivity matrix \mathbf{A} (equation 3). Because a BCCB matrix is diagonalized by the 2D unitary discrete Fourier transform (DFT), \mathbf{C} can be written as

$$\mathbf{C} = (\mathbf{F}_{2Q} \otimes \mathbf{F}_{2P})^* \mathbf{\Lambda} (\mathbf{F}_{2Q} \otimes \mathbf{F}_{2P}) , \quad (47)$$

where the symbol “ \otimes ” denotes the Kronecker product (?), \mathbf{F}_{2Q} and \mathbf{F}_{2P} are the $2Q \times 2Q$ and $2P \times 2P$ unitary DFT matrices (? , p. 31), respectively, the superscript “ $*$ ” denotes the complex conjugate and $\mathbf{\Lambda}$ is a $4QP \times 4QP$ diagonal matrix containing the eigenvalues of \mathbf{C} . Due to the diagonalization of the matrix \mathbf{C} , the auxiliary system (equation 46) can be rewritten by using equation 47 and premultiplying both sides of the result by $(\mathbf{F}_{2Q} \otimes \mathbf{F}_{2P})$, i.e.,

$$\mathbf{\Lambda} (\mathbf{F}_{2Q} \otimes \mathbf{F}_{2P}) \mathbf{v} = (\mathbf{F}_{2Q} \otimes \mathbf{F}_{2P}) \mathbf{w} . \quad (48)$$

By applying the vec-operator (?) to both sides of equation 48, by premultiplying both sides of the result by \mathbf{F}_{2Q}^* and then postmultiplying both sides of the result by \mathbf{F}_{2P}^*

$$\mathbf{F}_{2Q}^* [\mathbf{L} \circ (\mathbf{F}_{2Q} \mathbf{V} \mathbf{F}_{2P})] \mathbf{F}_{2P}^* = \mathbf{W} , \quad (49)$$

where “ \circ ” denotes the Hadamard product (? , p. 298) and \mathbf{L} , \mathbf{V} and \mathbf{W} are $2Q \times 2P$ matrices obtained by rearranging, along their rows, the elements forming the diagonal of matrix $\mathbf{\Lambda}$, vector \mathbf{v} and vector \mathbf{w} , respectively. The left side of equation 49 contains the 2D Inverse Discrete Fourier Transform (IDFT) of the term in brackets, which in turn represents the Hadamard product of matrix \mathbf{L} and the 2D DFT of matrix \mathbf{V} . Matrix \mathbf{L} contains the eigenvalues of $\mathbf{\Lambda}$ (equation 47) and can be efficiently computed by using only the first column of the BCCB matrix \mathbf{C} (equation 46).

Actually, in ? (? , ?) a fast 2D discrete circular convolution (?) is used to process very large gravity and magnetic datasets efficiently. The convolutional equivalent layer was applied to perform upward continuation of large magnetic datasets. Compared to the classical Fourier approach, ?’s (?) method produces smaller border effects without using any padding scheme.

Without taking advantage of the symmetric BTTB structure of the sensitivity matrix (? , ?) that arises when gravimetric observations are measured on a horizontally regular grid, on a flat surface and considering a regular grid of equivalent sources within a horizontal layer, ? explored the symmetry of the gravity kernel to reduce the number of forward model evaluations. By exploiting the symmetries of the gravity kernels and redundancies in the forward model evaluations on a regular grid and combining the subspace solution based on eigenvectors of the gridded dataset, ? estimated the mass excess or deficiency produced by anomalous sources with positive or negative density contrast.

3.0.7 The deconvolutional equivalent layer with BTTB matrices

To avoid the iterations of the conjugate gradient method in ?, we can employ the deconvolution process. Equation 49 shows that estimate the matrix \mathbf{V} , containing the elements of parameter vector \mathbf{p} , is a inverse problem that could be solved by deconvolution. From equation 49, the matrix \mathbf{V} can be obtain by deconvolution, i.e.

$$\mathbf{V} = \mathbf{F}_{2Q}^* \left[\frac{(\mathbf{F}_{2Q} \mathbf{W} \mathbf{F}_{2P})}{\mathbf{L}} \right] \mathbf{F}_{2P}^* . \quad (50)$$

Equation 50 shows that the parameter vector (in matrix \mathbf{V}) can be theoretically obtain by dividing each potential-field observations (in matrix \mathbf{W}) by each eigenvalues (in matrix \mathbf{L}). Hence, the parameter vector is constructed by element-by-element division of data by eigenvalues.

However, the deconvolution often is extremely unstable. This means that a small change in data can lead to an enormous change in the estimated parameter. Hence, equation 50 requires regularization to be useful. We used wiener deconvolution to obtain a stable solution, i.e.,

$$\mathbf{V} = \mathbf{F}_{2Q}^* \left[(\mathbf{F}_{2Q} \mathbf{W} \mathbf{F}_{2P}) \frac{\mathbf{L}^*}{(\mathbf{L} \mathbf{L}^* + \mu)} \right] \mathbf{F}_{2P}^* , \quad (51)$$

where the matrix \mathbf{L}^* contains the complex conjugate eigenvalues and μ is a parameter that controls the degree of stabilization.

3.1 Solution stability

The solution stability of the equivalent-layer methods is rarely addressed. Here, we follow the numerical stability analysis presented in Siqueira et al. (2017).

Let us assume noise-free potential-field data \mathbf{d} , we estimate a physical-property distribution \mathbf{p} (estimated solution) within the equivalent layer. Then, the noise-free data \mathbf{d} are contaminated with additive D different sequences of pseudorandom Gaussian noise, creating different noise-corrupted potential-field data \mathbf{d}_ℓ^o , $\ell = 1, \dots, D$. From each \mathbf{d}_ℓ^o , we estimate a physical-property distribution $\hat{\mathbf{p}}_\ell$ within the equivalent layer.

Next, for each noise-corrupted data \mathbf{d}_ℓ^o and estimated solution $\hat{\mathbf{p}}_\ell$, the ℓ th model perturbation δp_ℓ and the ℓ th data perturbation δd_ℓ are, respectively, evaluated by

$$\delta p_\ell = \frac{\|\hat{\mathbf{p}}_\ell - \mathbf{p}\|_2}{\|\mathbf{p}\|_2}, \quad \ell = 1, \dots, D, \quad (52)$$

and

$$\delta d_\ell = \frac{\|\mathbf{d}_\ell^o - \mathbf{d}\|_2}{\|\mathbf{d}\|_2}, \quad \ell = 1, \dots, D. \quad (53)$$

Regardless of the particular method used, the following inequality (?, p. 66) is applicable:

$$\delta p_\ell \leq \kappa \delta d_\ell, \quad \ell = 1, \dots, D, \quad (54)$$

where κ is the constant of proportionality between the model perturbation δp_ℓ (equation 52) and the data perturbation δd_ℓ (equation 53). The constant κ acts as the condition number of an invertible matrix in a given inversion, and thus measures the instability of the solution. The larger (smaller) the value of κ the more unstable (stable) is the estimated solution.

471 Equation 54 shows a linear relationship between the model perturbation and the data perturbation. By
472 plotting δp_ℓ (equation 52) against δd_ℓ (equation 53) produced by a set of D estimated solution obtained by
473 applying a given equivalent-layer method, we obtain a straight line behaviour described by equation 54.
474 By applying a linear regression, we obtain a fitted straight line whose estimated slope (κ in equation 54)
475 quantifies the solution stability.

476 Here, the analysis of solution stability is numerically conducted by applying the classical equivalent-
477 layer technique with zeroth-order Tikhonov regularization, the convolutional method for gravimetric and
478 magnetic data, the deconvolutional method (equation 50) and the deconvolutional method with different
479 values for the Wiener stabilization (equation 51).

4 NUMERICAL SIMULATIONS

We investigated different computational algorithms for inverting gravity disturbances and total-field anomalies. To test the capability of the fast equivalent-layer technique for processing that potential field data we measure of the computational effort by counting the number of floating-point operations (*flops*), such as additions, subtractions, multiplications, and divisions (Golub and Loan, 2013) for different number of observation points, ranging from 10, 000 up to 1, 000, 000. The results generated when using iterative methods are set to $it = 50$ for the number of iterations.

4.1 Floating-point operations calculation

To measure the computational effort of the different algorithms to solve the equivalent layer linear system, a non-hardware dependent method can be useful because allow us to do direct comparison between them. Counting the floating-point operations (*flops*), i.e., additions, subtractions, multiplications and divisions is a good way to quantify the amount of work of a given algorithm (Golub and Loan, 2013). For example, the number of *flops* necessary to multiply two vectors \mathbb{R}^N is $2N$. A common matrix-vector multiplication with dimension $\mathbb{R}^{N \times N}$ and \mathbb{R}^N , respectively, is $2N^2$ and a multiplication of two matrices $\mathbb{R}^{N \times N}$ is $2N^3$. Figure ?? shows the total *flops* count for the different methods presented in this review with a crescent number of data, ranging from 10, 000 to 1, 000, 000 for the gravity equivalent layer and figure ?? for magnetic data.

4.1.1 Normal equations using Cholesky decomposition

The equivalent sources can be estimated directly from solving the normal equations 3. In this work we will use the Cholesky decompositions method to calculate the necessary *flops*. In this method it is calculated the lower triangle of $\mathbf{A}^T \mathbf{A}$ ($1/2N^3$), the Cholesky factor ($1/3N^3$), a matrix-vector multiplication ($2N^2$) and finally solving the triangular system ($2N^2$), totalizing

$$f_{classical} = \frac{5}{6}N^3 + 4N^2 \quad (55)$$

4.1.2 Window method (Leão and Silva, 1989)

The moving data-window scheme (Leão and Silva, 1989) solve N linear systems with much smaller sizes (equation 28b). For our results we are considering a data-window of the same size of wich the authors presented in theirs work ($N_w = 49$) and the same number of equivalent sources ($M_w = 225$). We are doing this process for all the other techniques to standardize the resolution of our problem. Using the Cholesky decomposition with this method the *flops* are

$$f_{window} = N \frac{5}{6} M_w N_w^2 + 4N_w M_w \quad (56)$$

4.1.3 PEL method (Oliveira Jr. et al., 2013)

The polynomial equivalent layer uses a simliar approach od moving windows from Leão and Silva (1989). For this operations calculation (equation 38) we used a first degree polynomial (two variables) and each window contains $N_s = 1, 000$ observed data and $M_s = 1, 000$ equivalent sources. Following the steps given in (Oliveira Jr. et al., 2013) the total *flops* becomes

$$f_{pel} = \frac{1}{3}H^3 + 2H^2 + 2NM_sH + H^2N + 2HN + 2NP \quad (57)$$

where H is the number of constant coefficients for the first degree polynomial ($P = 3$) times the number of windows ($P \times N/N_s$).

4.1.4 Conjugate gradient least square (CGLS)

The CGLS method is a very stable and fast algorithm for solving linear systems iteratively. Its computational complexity involves a matrix-vector product outside the loop ($2N^2$), two matrix-vector products inside the loop ($4N^2$) and six vector products inside the loop ($12N$) (?)

$$f_{cglS} = 2N^2 + it(4N^2 + 12N) \quad (58)$$

4.1.5 Wavelet compression method with CGLS (?)

For the wavelet method (equation 31) we have calculated a coompression rate of 98% ($C_r = 0.02$) for the threshold as the authors used in ? and the wavelet transformation requiring $\log_2(N)$ flops each (equations 30 and 32c), with its inverse also using the same number of operations (equation 34). Combined with the conjugate gradient least square necessary steps and iterations, the number of flops are

$$f_{wavelet} = 2NC_r + 4N \log_2(N) + it(4N \log_2(N) + 4NC_r + 12C_r) \quad (59)$$

4.1.6 Fast equivalent layer for gravity data (Siqueira et al., 2017)

The fast equivalent layer from Siqueira et al. (2017) solves the linear system in it iterations. The main cost of this method (equations 40,41, 42 and 43)is the matrix-vector multiplication to asses the predicted data ($2N^2$) and three simply element by element vector sum, subtraction and division ($3N$ total)

$$f_{siqueira} = it(3N + 2N^2) \quad (60)$$

4.1.7 Convolutional equivalent layer for gravity data (?)

This methods replaces the matrix-vector multiplication of the iterative fast-equivalent technique (Siqueira et al., 2017) by three steps, involving a Fourier transform, an inverse Fourier transform, and a Hadamard product of matrices (equation 49). Considering that the first column of our BCCB matrix has $4N$ elements, the flops count of this method is

$$f_{convgrav} = \kappa 4N \log_2(4N) + it(27N + \kappa 8N \log_2(4N)) \quad (61)$$

In the resultant count we considered a *radix*-2 algorithm for the fast Fourier transform and its inverse, which has a κ equals to 5 and requires $\kappa 4N \log_2(4N)$ flops each. The Hadarmard product of two matrices of $4N$ elements with complex numbers takes $24N$ flops. Note that equation 61 is different from the one presented in ? because we also added the flops necessary to calculate the eigenvalues in this form. It does not differentiate much in order of magnitude because the iterative part is the most costful.

536 4.1.8 Convolutional equivalent layer for magnetic data (?)

537 The convolutional equivalent layer for magnetic data uses the same flops count of the main operations as
 538 in the gravimetric case (equation 49), the difference is the use of the conjugate gradient algorithm to solve
 539 the inverse problem. It requires a Hadamard product outside of the iterative loop and the matrix-vector and
 540 vector-vector multiplications inside the loop as seen in equation 58.

$$f_{convmag} = \kappa 16N \log_2(4N) + 24N + it(\kappa 16N \log_2(4N) + 60N) \quad (62)$$

541 4.1.9 Deconvolutional method

542 The deconvolution method does not require an iterative algorithm, rather it solves the estimative of the
 543 physical properties in a single step using the $4N$ eigenvalues of the BCCB matrix as in the convolutional
 544 method. From equation 50 it is possible to deduce this method requires two fast Fourier transform
 545 ($\kappa 4N \log_2(4N)$), one for the eigenvalues and another for the data transformation, a element by element
 546 division ($24N$) and finally, a fast inverse Fourier transform for the final estimative ($\kappa 4N \log_2(4N)$).

$$f_{deconv} = \kappa 12N \log_2(4N) + 24N \quad (63)$$

547 Using the deconvolutional method with a Wiener stabilization adds two multiplications of complex
 548 elements of the conjugates eigenvalues ($24N$ each) and the sum of $4N$ elements with the stabilization
 549 parameter μ as shown in equation 51

$$f_{deconvwiener} = \kappa 12N \log_2(4N) + 76N \quad (64)$$

CONFLICT OF INTEREST STATEMENT

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

AUTHOR CONTRIBUTIONS

The Author Contributions section is mandatory for all articles, including articles by sole authors. If an appropriate statement is not provided on submission, a standard one will be inserted during the production process. The Author Contributions statement must describe the contributions of individual authors referred to by their initials and, in doing so, all authors agree to be accountable for the content of the work. Please see here for full authorship criteria.

FUNDING

Diego Takahashi was supported by a Post-doctoral scholarship from CNPq (grant 300809/2022-0) Valéria C.F. Barbosa was supported by fellowships from CNPq (grant 309624/2021-5) and FAPERJ (grant 26/202.582/2019). Vanderlei C. Oliveira Jr. was supported by fellowships from CNPq (grant 315768/2020-7) and FAPERJ (grant E-26/202.729/2018).

ACKNOWLEDGMENTS

We thank the Brazilian federal agencies CAPES, CNPq, state agency FAPERJ and Observatório Nacional research institute and Universidade do Estado do Rio de Janeiro.

DATA AVAILABILITY STATEMENT

The datasets generated for this study can be found in the frontiers-paper Github repository link: <https://github.com/DiegoTaka/frontiers-paper>.

REFERENCES

- Aster, R. C., Borchers, B., and Thurber, C. H. (2019). *Parameter Estimation and Inverse Problems* (Elsevier), 3 edn.
- Barnes, G. and Lumley, J. (2011). Processing gravity gradient data. *GEOPHYSICS* 76, I33–I47. doi:10.1190/1.3548548
- Cordell, L. (1992). A scattered equivalent-source method for interpolation and gridding of potential-field data in three dimensions. *Geophysics* 57, 629–636
- Dampney, C. N. G. (1969). The equivalent source technique. *GEOPHYSICS* 34, 39–53. doi:10.1190/1.1439996
- Emilia, D. A. (1973). Equivalent sources used as an analytic base for processing total magnetic field profiles. *GEOPHYSICS* 38, 339–348. doi:10.1190/1.1440344
- Golub, G. H. and Loan, C. F. V. (2013). *Matrix Computations (Johns Hopkins Studies in the Mathematical Sciences)* (Johns Hopkins University Press), 4 edn.
- Gonzalez, S. P., Barbosa, V. C. F., and Oliveira Jr., V. C. (2022). Analyzing the ambiguity of the remanent-magnetization direction separated into induced and remanent magnetic sources. *Journal of Geophysical Research: Solid Earth* 127, 1–24. doi:10.1029/2022JB024151

- Guspi, F. and Novara, I. (2009). Reduction to the pole and transformations of scattered magnetic data using Newtonian equivalent sources. *GEOPHYSICS* 74, L67–L73. doi:10.1190/1.3170690
- Hansen, R. O. and Miyazaki, Y. (1984). Continuation of potential fields between arbitrary surfaces. *GEOPHYSICS* 49, 787–795. doi:10.1190/1.1441707
- Jirigalatu, J. and Ebbing (2019). A fast equivalent source method for airborne gravity gradient data. *Geophysics* 84, G75–G82. doi:10.1190/GEO2018-0366.1
- Leão, J. W. D. and Silva, J. B. C. (1989). Discrete linear transformations of potential field data. *Geophysics* 54, 497–507. doi:10.1190/1.1442676
- Li, Y., Nabighian, M., and Oldenburg, D. W. (2014). Using an equivalent source with positivity for low-latitude reduction to the pole without striation. *GEOPHYSICS* 79, J81–J90. doi:10.1190/geo2014-0134.1
- Li, Y. and Oldenburg, D. W. (2010). Rapid construction of equivalent sources using wavelets. *GEOPHYSICS* 75, L51–L59. doi:10.1190/1.3378764
- Mendonça, C. A. (2020). Subspace method for solving large-scale equivalent layer and density mapping problems. *GEOPHYSICS* 85, G57–G68. doi:10.1190/geo2019-0302.1
- Mendonça, C. A. and Silva, J. B. C. (1994). The equivalent data concept applied to the interpolation of potential field data. *Geophysics* 59, 722–732. doi:10.1190/1.1443630
- Menke, W. (2018). *Geophysical data analysis: Discrete inverse theory* (Elsevier), 4 edn.
- Oliveira Jr., V. C., Barbosa, V. C. F., and Uieda, L. (2013). Polynomial equivalent layer. *GEOPHYSICS* 78, G1–G13. doi:10.1190/geo2012-0196.1
- Reis, A. L. A., Oliveira Jr., V. C., and Barbosa, V. C. F. (2020). Generalized positivity constraint on magnetic equivalent layers. *Geophysics* 85, 1–45. doi:10.1190/geo2019-0706.1
- Roy, A. (1962). Ambiguity in geophysical interpretation. *GEOPHYSICS* 27, 90–99. doi:10.1190/1.1438985
- Silva, J. B. C. (1986). Reduction to the pole as an inverse problem and its application to low-latitude anomalies. *GEOPHYSICS* 51, 369–382. doi:10.1190/1.1442096
- Siqueira, F., Oliveira Jr., V. C., and Barbosa, V. C. F. (2017). Fast iterative equivalent-layer technique for gravity data processing: A method grounded on excess mass constraint. *GEOPHYSICS* 82, G57–G69. doi:10.1190/GEO2016-0332.1
- Soler, S. R. and Uieda, L. (2021). Gradient-boosted equivalent sources. *Geophysical Journal International* 227, 1768–1783. doi:10.1093/gji/ggab297
- Takahashi, D., Oliveira Jr., V. C., and Barbosa, V. C. (2022). Convolutional equivalent layer for magnetic data processing. *Geophysics* 87, 1–59
- Takahashi, D., Oliveira Jr., V. C., and Barbosa, V. C. F. (2020). Convolutional equivalent layer for gravity data processing. *GEOPHYSICS* 85, G129–G141. doi:10.1190/geo2019-0826.1
- van der Sluis, A. and van der Vorst, H. A. (2004). Innovation and intellectual property rights. In *The Oxford Handbook of Innovation*, eds. J. Fagerberg, D. C. Mowery, and R. R. Nelson (Oxford: Oxford University Press), chap. 10. 266–290
- Xia, J. and Sprowl, D. R. (1991). Correction of topographic distortion in gravity data. *Geophysics* 56, 537–541
- Xia, J., Sprowl, D. R., and Adkins-Heljeson, D. (1993). Correction of topographic distortions in potential-field data; a fast and accurate approach. *Geophysics* 58, 515–523. doi:10.1190/1.1443434
- Zidarov, D. (1965). Solution of some inverse problems of applied geophysics. *Geophysical Prospecting* 13, 240–246. doi:10.1111/j.1365-2478.1965.tb01932.x