

# Boussinesq Approximation and Nondimensionalization for Mantle Convection

Ivan Cherkashin

March 23, 2016

The simplest form of equations of the Boussinesq approximation is nondimensionalized in the context of numerical modeling of mantle convection. Physical considerations underlying the appropriate choices of time, velocity, and pressure scales for nondimensionalization of these equations are explained, which is not widely presented in the literature. The physical meaning of all nondimensional numbers arising as a result of nondimensionalization is discussed.

The main goal of this survey is to demonstrate the underlying physical reasoning behind nondimensionalization and to obtain a nondimensional form of the equations, commonly used in numerical modeling, that depends only on a single parameter: the Rayleigh number.

## Contents

<b>1. Equations of Boussinesq Approximation in Dimensional Form</b>	<b>4</b>
1.1. Physical Meaning of the Equations . . . . .	6
<b>2. Applications of the Boussinesq Approximation</b>	<b>6</b>
<b>3. History of the Boussinesq Approximation</b>	<b>7</b>

<b>4. Boussinesq Approximation in Computational Models of Mantle Convection</b>	<b>8</b>
<b>5. Critique of the Boussinesq Approximation</b>	<b>9</b>
<b>6. Nondimensionalization and Choice of Scales</b>	<b>10</b>
6.1. Characteristic Length Scale . . . . .	11
6.2. Characteristic Temperature Scale . . . . .	11
6.3. Characteristic Time Scale . . . . .	12
6.4. Characteristic Velocity Scale . . . . .	13
6.5. Characteristic Pressure Scale . . . . .	13
<b>7. Nondimensional Form of the Equations of Boussinesq Approximation</b>	<b>14</b>
7.1. Physical Meaning of the Rayleigh Number . . . . .	15
<b>A. Buckingham's II Theorem</b>	<b>15</b>
A.1. Content of Buckingham's II-Theorem . . . . .	16
A.2. Physical Dimensions Form a Linear Space . . . . .	16
A.3. Dimensional Matrix . . . . .	17
A.3.1. Dimensionless Quantities and Kernel of the Dimensional Matrix	17
A.3.2. Dimensionally Independent Physical Quantities and the Rank of the Dimensional Matrix . . . . .	18
A.4. Buckingham's Theorem . . . . .	18
A.5. Dimensionality Reduction of a Physical Functional Dependence . . .	19
<b>References</b>	<b>21</b>

## Introduction

Nondimensionalization is a powerful method of investigating complex systems when the exact dynamics is practically impossible to compute yet it is still possible to estimate the mutual importance of various factors on the overall dynamics. For example, although, in general, the exact solution to the Navier-Stokes equations cannot be computed, nondimensionalization of these equations may reveal the relative importance of pressure, viscosity, and other factors on the overall behavior of the flow. Thus, knowing that the dissipation of energy and momentum due to viscosity is negligible compared to the vigor of the inertial motion of the fluid, which manifests itself in a very large value of a nondimensional parameter called Reynolds number, it is possible to neglect the viscous terms in the Navier-Stokes equations, thus turn-

ing them into Euler's equations. Analogously, when Reynolds number is small and inertial motion is negligible compared to viscous effects, Navier-Stokes equations turn into Stokes equations. These examples demonstrate how nondimensionalization may serve as a tool leading to useful approximations and deeper understanding of the physics underlying mathematical equations.

At a first glance, nondimensionalization appears to be a simple scheme: first, a change of variables that would make the physical variables involved in the equations nondimensional, followed by simple algebraic manipulations with the coefficients of the equations and, finally, taking limits with respect to certain coefficients of interest, such as Reynolds number mentioned above. However, such unintelligent approach may lead to mistakes. For example, under incorrect choice of scales, it may lead to disappearance of the pressure term in the Stokes equations when they are derived from Navier-Stokes, which totally contradicts the physical fact that pressure is an important factor even for highly viscous flows. Thus, correct nondimensionalization requires understanding of the physics underlying the decision one makes when a certain scale must be chosen in order to nondimensionalize a particular physical quantity. Otherwise, nondimensionalization becomes a dull set of tricks, which cannot amount to genuine knowledge and understanding of the natural world.

Unfortunately, although nondimensionalization has been successfully employed for a long time as a powerful method of investigation, no universal algorithm was developed for applying this method to all of the possible physical situations. It is natural, since it is impossible to account for all of the possible combinations of physical factors in an abstract mathematical scheme. Thus, when one wishes to nondimensionalize the equations of mantle convection, the first difficulty that arises is an absence of a reliable source devoted to this specific context. Because of that, a researcher is compelled to search through a vast, inhomogeneous array of literature in order to answer questions regarding the nondimensionalization of the equations of mantle convection. This is obviously inefficient and the valuable knowledge is lost if a successful researcher never shares his synthesis with the community.

The present work is attempts to address this very problem. It is a synthesis of the numerous, oftentimes very little but important ideas from the sources the author encountered. It is hoped to be a point of growth in the mantle convection community, providing a source that explains nondimensionalization in the context of mantle convection. The author hopes that this work save precious time for other researches and students, which should be devoted to solving new interesting problems.

The goal of this work is to derive a nondimensional form of the equations of the Boussinesq approximation in the context of mantle convection in a form that depends on a single nondimensional parameter: the Rayleigh number. To do so,

we start nondimensionalize the incompressible Stokes equations and the advection-diffusion equation for temperature, explaining the physical meaning of this procedure and its underlying subtleties. Only the simplest form of the equations is considered in which the viscosity and thermal diffusivity are constant. This simplest case is, nevertheless, important, since it is often used for validating competing computational software that solve the same problem employing various numerical methodologies and algorithms.

In addition to presenting nondimensionalization of the equations commonly used in the mantle convection community, the author hopes that his work is also useful from a pedagogical perspective. Hopefully, it answers the questions that members of the community, especially students, often have, and, even more importantly, that it explains how the answers are actually found.

## Acknowledgements

The author would like to thank Professor Donald L. Turcotte, and Professor Louise Kellogg from Earth and Planetary Sciences, and Professor Elbridge G. Puckett from the Department of Mathematics and the Computational Infrastructure for Geodynamics at UC Davis for their comments, suggestions, and support for this work.

## 1. Equations of Boussinesq Approximation in Dimensional Form

We will consider the simplest form of equations used to model convection in the mantle. Specifically, we will refer to the simplest special case of the equations used by the ASPECT mantle convection numerical simulation software [2, p. 13].

Namely, our equations are based on the following assumptions. Firstly, linear relationship between density and temperature perturbations in the fluid is assumed. In other words, Boussinesq approximation is used for the equation of state of the fluid:

$$\rho(\theta_0 + \theta) - \rho(\theta_0) \approx \rho(\theta_0) \left( \frac{1}{\rho(\theta_0)} \frac{\partial \rho}{\partial \theta} \Big|_{\theta_0} \right) \theta = \rho_0 \beta \theta \quad (1)$$

where  $\rho(\theta_0) = \rho_0$  and  $\theta_0$  are the reference density and temperature of the fluid, respectively, and

$$\beta = \frac{1}{\rho_0} \frac{\partial \rho}{\partial \theta} \Big|_{\theta_0}, \quad [\beta] = \frac{1}{\Theta}$$

is the *volumetric coefficient of thermal expansion*<sup>1</sup> that characterizes the decrease of density with temperature. Its name makes sense if we recall the relationship between the specific volume and density:

$$v = \frac{1}{\rho} \implies \beta = \frac{1}{\rho_0} \frac{\partial \rho}{\partial \theta} \Big|_{\theta_0} = v_0 \frac{\partial v^{-1}}{\partial \theta} \Big|_{\theta_0} = -\frac{v_0}{v_0^2} \frac{\partial v}{\partial \theta} \Big|_{\theta_0} = -\frac{1}{v_0} \frac{\partial v}{\partial \theta} \Big|_{\theta_0}$$

Secondly, the viscosity  $\nu$  of the fluid is also considered constant.

Thirdly, the coefficient of thermal diffusivity  $\kappa$  is considered constant. Finally, it is assumed that no sources of heat are present in the fluid.

The following notation is used to describe the process of mantle convection. The symbols  $L, M, T, \Theta$  represent the dimensions of length, mass, time, and temperature, respectively;  $[X]$  means the dimension of a physical quantity  $X$ .

$[\mathbf{x}] = L$  is the vector describing a location of a particle of the fluid in space.

$[\mathbf{v}] = LT^{-1}$  is the velocity field of the flow.

$[t] = T$  stands for time.

$\mathbf{e}_z$  is the unit normal to the surface of the Earth.

$[\rho] = ML^{-3}$  stands for density.

$[g] = LT^{-2}$  is the acceleration in the uniform gravitational field.

$[\nu] = \left[\frac{\mu}{\rho}\right] = L^2T^{-1}$  is the kinematic viscosity of the fluid.

$[p] = ML^{-1}T^{-2}$  is *dynamic* pressure (i.e., the one due to inertial movement of fluid, which should not be confused with thermodynamic pressure).

$[\theta] = \Theta$  stands for temperature of the fluid.

$[\kappa] = L^2T^{-1}$  is the thermal diffusivity coefficient, which is a measure of the intensity of the diffusion of temperature.

---

<sup>1</sup>Note that  $\beta < 0$  since the density of the fluid generally decreases with the increase of temperature.

Based on these assumptions and notation, the equations simplify to the following form:<sup>2</sup>

$$\nu \Delta \mathbf{v} = \frac{1}{\rho_0} \nabla p - g\theta \mathbf{e}_z \quad (2)$$

$$\nabla \cdot \mathbf{v} = 0 \quad (3)$$

$$\partial_t \theta + \mathbf{v} \cdot \nabla \theta = \kappa \Delta \theta \quad (4)$$

Naturally, these equations must be supplemented by appropriate boundary and initial conditions in order to describe a concrete system.

### 1.1. Physical Meaning of the Equations

Equations (2) and (3) are known as *incompressible Stokes equations*. They accurately approximate highly viscous or slow ("creeping") flows and are different from the *incompressible Navier-Stokes equations* by neglecting the inertial terms  $\partial_t \mathbf{v} + \mathbf{v} \cdot \nabla \mathbf{v}$ .

More specifically, equation (2) describes the conservation of momentum in the flow. The equation (3) is the incompressibility constraint on the flow.

The equation (4) describes propagation of heat in the fluid in terms of the spatial and temporal evolution of the temperature field. The terms  $\partial_t \theta + \mathbf{v} \cdot \nabla \theta$  describe the transport of temperature by the flow.<sup>3</sup> The term  $\kappa \Delta \theta$  describes the transport of heat by conduction or, equivalently, diffusion of temperature.

## 2. Applications of the Boussinesq Approximation

The equations of Boussinesq approximation, both in the simplest form considered here and in more complicated ones, are used extensively in the modeling of mantle convection in Earth and other planets of the solar system. [23] Oceanography is

<sup>2</sup>Note that all the terms in the equation (2) were divided by  $\rho_0$ .

<sup>3</sup>An alternative name is *convection-diffusion* equation. In terms of semantics, advection is preferable to convection in this context. Indeed, *convection* means the flow of the fluid in response to the temperature gradients present in the fluid. *Advection*, however, is a term used to describe a transport of a quantity (such as temperature, concentration of a chemical, etc.) in response to the fluid motion. That is why, in this case, temperature is *advected*. Of course, advection of temperature (i.e. heat transfer) causes convection of a fluid and vice versa, so these processes are interrelated. However, heat transfer by fluid motion is often referred to as *convection*, but, in the light of our previous remarks, it would be more appropriate to call it *advection* and *advective heat transport*. However, this is a terminological inconsistency rooted in history.

another important application of the Boussinesq approximation [17]. Moreover, there are "various thermal, geophysical, astrophysical and magnetohydrodynamic problems in the framework of 'Boussinesquian fluid dynamics'" that are based on the equations discussed in this survey. [26]

Naturally, it is important to thoroughly understand the simplest equations first before proceeding to the more complicated models used in applications. Yet the list of the above mentioned scientific fields in which Boussinesq approximation is used serves as an effective illustration of the necessity to understand this approximation for today's researcher.

### 3. History of the Boussinesq Approximation

The history of scientific study of convection is well summarized in an article by A.V. Getling [9]:

The role of non-uniform heating as the producer of most types of fluid motions in the Universe was first recognised in the mid-eighteenth century, nearly simultaneously by George Hadley and Mikhail Lomonosov. Well-directed studies of convection in horizontal fluid layers heated from below trace back to Benard's experiments [3], in which the instability mechanism was, however, not purely thermal and was closely related to the *thermocapillary effect*. Lord Rayleigh [22] was the first to consider a linear problem of the onset of thermal convection in a horizontal layer, and a more comprehensive analysis of this problem was given by Pellew and Southwell [20]. A highly extensive survey of the linear stability problems, including investigations of the effects of rotation and magnetic field on Rayleigh-Benard convection, was presented in a classical monograph by Chandrasekhar [6]. Subsequent studies mainly dealt with nonlinear convection regimes and related pattern-formation processes. The volume of relevant publications has grown dramatically, and a number of monographs of a more or less wide scope summarize them [a concise review of many results that refer specifically to Rayleigh-Benard convection and were obtained by the end of the 1990s can be found in Getling [10]].

The first theoretical analysis of convection by Lord Rayleigh marks the beginning of the study of the Boussinesq approximation and the accompanying equations. The work [12, p. 545] presents a brief history of this subject matter:

Although these equations are named after Boussinesq [5], they seem to have been first used by Oberbeck [19]. The plausibility argument given by Chandrasekhar [6] is often referenced, but the first attempt at a detailed derivation in a dynamical situation was made by Spiegel and Veronis [25]. They considered a perfect gas of constant properties and used an order of magnitude argument. Similar assumptions and methods were used by Gebhart [7] and Plate [21].

Mihaljan [18] used a mathematically rigorous small parameter expansion technique to derive the Boussinesq equations. He assumed that density was a linear function of temperature only and that the other properties were constant. A generalization of this approach was presented by Malkus [15,16] who considered a perfect gas and allowed thermal diffusivity and viscosity to vary with temperature only.

The theoretical study of the equations of Boussinesq approximation continues and has become especially relevant for understanding the validity of the computational models of mantle convection.

## 4. Boussinesq Approximation in Computational Models of Mantle Convection

We will provide a brief description of the software projects devoted to numerical modeling of mantle convection problems that the author is familiar with. It must be kept in mind that this software generally uses more complicated equations than those considered in this survey, although the latter can be solved as a special case of the more general equations and are often used to validate the results of the computations. More specifically, the more general equations may contain additional terms such as heat sources inside the mantle; the characteristics of the fluid such as viscosity and thermal diffusivity may be non-constant. [2, p. 13].

A variety of computational software projects for modeling mantle convection has been developed and supported by the Computational Infrastructure for Geodynamics (CIG): [1]

**ASPECT:** Finite element parallel code to simulate problems in thermal convection in both 2D and 3D models - currently in alpha testing.

**CitcomCU:** Finite element parallel code capable of modeling thermochemical convection in a three-dimensional domain appropriate for convection within the Earth's mantle.



**CitcomS:** Finite element code designed to solve compressible thermochemical convection problems relevant to Earth’s mantle.

**ConMan:** Finite element program for the solution of the equations of incompressible, infinite-Prandtl number convection in two dimensions, originally written by Scott King, Arthur Raefsky, and Brad Hager.

**Ellipsis3D:** Three-dimensional version of the particle-in-cell finite element code Ellipsis, a solid modeling code for visco-elastoplastic materials. The particle-in-cell method combines the strengths of the Lagrangian and Eulerian formulations of mechanics while bypassing their limitations.

**HC:** Global mantle circulation solver following Hager & O’Connell (1981) which can compute velocities, tractions, and geoid for simple density distributions and plate velocities.

**Taras Gerya’s Finite difference code:** Professor Taras Gerya [8] provides several examples of MATLAB programs that accompany his book. His computational methodology is the finite-difference method, which makes his book valuable since this method generally receives less attention in research papers and software projects than the finite-element method.

## 5. Critique of the Boussinesq Approximation

Following the arguments presented in the work [11] by G.I. Barenblatt, we will show that there is a flaw in the dimensional analysis of these equations that makes them, in fact, inadequate for describing mantle convection. It is important to note, however, that it does not mean that these equations have no value for other purposes, such as validation of numerical algorithms.

The main deficiency of our analysis, besides assuming the characteristics of the fluid constant and the fluid itself incompressible, is the assumption that “the contribution of viscous energy dissipation to the thermal balance of the fluid” is negligible [12, p. 545]:

”In principle, the contribution of viscous energy dissipation to the thermal balance of the fluid should also be taken into account. To do this, one additional parameter must be included, the mechanical equivalent of heat  $J$ .” [11, p. 45].

As one can see in section 1, this parameter was neglected in our discussion.

Through dimensional analysis, the author obtains an important nondimensional parameter [11, p. 46]

$$\Pi_3 = \frac{Jc}{\beta g L}$$

As G.I. Barenblatt notes, neglecting this parameter is inappropriate for modeling mantle convection since the layers of fluid representing the mantle are not thin, unlike in the original experiments with convection in thin layers of fluid by Bernard<sup>4</sup>:

”In what follows, we shall discuss convective motion in thin layers, for which the parameter  $\Pi_3$  is large ( $\Pi_3 \gg 1$ ), so that the effect of this parameter on the similarity conditions may be neglected ... However, *when modelling convection in the Earth’s mantle the parameter  $\Pi_3$  is of order unity and cannot be neglected.*” [11, pp. 46-47]

Thus, based on G.I. Barenblatt’s arguments, a more physically adequate model of mantle convection must include generation of heat due to viscous energy dissipation<sup>5</sup>. This must be kept in mind when making conclusions based on the results of computational simulations obtained by solving the equations (2), (3), and (4) numerically. [23, cf. p. 772 for the discussion of the importance of viscous dissipation in mantle convection]

Nevertheless, as was mentioned before, this critique does not render these equations useless: at the very least, they are useful for numerical analysis and validation of computations, as well as for demonstrating the idea of nondimensionalization.

## 6. Nondimensionalization and Choice of Scales

Given 9 parameters describing mantle convection in our simplified model ( $\mathbf{x}, \mathbf{v}, \rho, \nu, \kappa, g, p, \theta, t$ ) and 4 units of measurement ( $L, M, T, \Theta$ ), by Buckingham’s  $\Pi$ -theorem we can form  $9 - 4 = 5$  nondimensional parameters describing the system. [11]

In order to nondimensionalize the equations (2), (3), and (4), we choose characteristic scales for the length  $L$ , time  $T$ , velocity  $V$ , pressure  $P$ , and temperature  $\Theta$  and perform a change of variables such that these quantities are now measured against these characteristic scales:

$$\mathbf{x}' = \frac{\mathbf{x}}{L}, t' = \frac{t}{T}, \mathbf{v}' = \frac{\mathbf{v}}{V}, p' = \frac{p}{P}, \theta' = \frac{\theta}{\Theta}$$

---

<sup>4</sup>The same assumption was made by Lord Rayleigh in his theoretical analysis of convection in a thin layer of fluid heated from below.

<sup>5</sup>I.e. friction between layers of fluid.

or, equivalently,

$$\mathbf{x} = L\mathbf{x}', t = Tt', \mathbf{v} = V\mathbf{v}', p = Pp', \theta = \Theta\theta'$$

Substituting these new variables in the equations (2), (3), and (4) and dropping the primes for convenience, we obtain the following system of equations:

$$\frac{\nu V}{L^2} \Delta \mathbf{v} = \frac{P}{\rho_0 L} \nabla p - g\Theta\theta \mathbf{e}_z \quad (5)$$

$$\frac{V}{L} \nabla \cdot \mathbf{v} = 0 \quad (6)$$

$$\frac{\Theta}{T} \partial_t \theta + \frac{\Theta V}{L} \mathbf{v} \cdot \nabla \theta = \frac{\kappa \Theta}{L^2} \Delta \theta \quad (7)$$

or, after dividing by the terms both sides of the equations by the appropriate terms such that the coefficients in front of the leftmost terms become 1, we obtain

$$\Delta \mathbf{v} = \frac{PL}{\rho_0 \nu V} \nabla p - \frac{g\Theta L^2}{\nu V} \theta \mathbf{e}_z \quad (8)$$

$$\nabla \cdot \mathbf{v} = 0 \quad (9)$$

$$\partial_t \theta + \frac{VT}{L} \mathbf{v} \cdot \nabla \theta = \frac{\kappa T}{L^2} \Delta \theta \quad (10)$$

The next problem in nondimensionalizing these equations is choosing the appropriate characteristic scales for the length  $L$ , velocity  $V$ , pressure  $P$ , time  $T$ , and temperature  $\Theta$  that adequately reflect the physical processes in the system.

### 6.1. Characteristic Length Scale

The characteristic length scale  $L$  is chosen based on the size of the physical domain where the flow takes place. For example, in case of a flow past a cylinder, it may be the radius of a cylinder. In case of convection in a box, it may be the length of a side of the box.

### 6.2. Characteristic Temperature Scale

The characteristic temperature scale  $\Theta$  is typically defined as either the average temperature  $\Theta = \theta_{avg}$  in the fluid, or as the maximum temperature difference in the fluid  $\Theta = \max \theta - \min \theta$ .

### 6.3. Characteristic Time Scale

There are two natural choices of time scale in the context of highly viscous flows that are inherent to mantle convection. [10, pp. 19-20]

The first choice is based on the viscosity time scale:

$$T_{viscosity} = \frac{L^2}{\nu}$$

The second choice is based on the heat conduction time scale:

$$T_{conduction} = \frac{L^2}{\kappa}$$

In order to choose the appropriate scale for the context of mantle convection, we observe that their ratio is equal to the Prandtl number:

$$\frac{T_{conduction}}{T_{viscosity}} = \frac{\nu}{\kappa} = Pr$$

A commonly used approximation for the dynamics of Earth's mantle is the limit of infinite viscosity, which is equivalent to the limit of infinite Prandtl number since thermal conductivity is usually fixed or bounded: [23, p. 267]

$$\nu \longrightarrow \infty \iff Pr \longrightarrow \infty$$

Physically, this limit can be interpreted as

$$T_{viscosity} \ll T_{conduction}$$

or

$$T_{viscosity} \longrightarrow 0$$

The interpretation of this limit is that information due to viscous interactions propagates infinitely fast. But infinitely fast processes cannot serve as a measure of real physical processes because they are infinitely slow relative to infinitely fast processes. Thus, the viscosity time scale becomes singular in the common mantle convection approximation, which makes it an inappropriate choice for the time scale.

Therefore, the only choice left is the time scale associated with heat conduction,

$$T = T_{conduction} = \frac{L^2}{\kappa}$$

Another argument supporting this choice of scale is the fact that heat conduction takes place regardless of the macroscopic flow of the fluid. Since initially the fluid is at rest in typical problems of mantle convection, measurable viscous dissipation processes cannot start until the fluid has been set in an adequately intense macroscopic flow. At the same time, heat conduction starts immediately because a gradient of temperature is always present in the mantle.

#### 6.4. Characteristic Velocity Scale

In typical problems of convection, the fluid is initially at rest and the researchers are interested in modeling the development of convection from the state of equilibrium. Therefore, there is no natural characteristic velocity in such context since in the early phase of convection the average velocity is close to zero but may be rather large in the developed phase.

Therefore, it is plausible to define characteristic velocity simply as a ratio of characteristic length and time:

$$V = \frac{L}{T} = \frac{L}{\frac{L^2}{\kappa}} = \frac{\kappa}{L}$$

#### 6.5. Characteristic Pressure Scale

An appropriate pressure scale for low Reynolds number flows (highly viscous or "creeping flows") is based on an estimate of viscous shear stress:

$$\tau = \frac{\rho\nu V}{L}$$

Following [14, pp. 433-434], we shall verify that this choice of scale is appropriate by comparing it with another possible choice of scale based on characteristic kinetic energy density of the flow.

There are two physical choices for nondimensionalizing the pressure:

$$P_1 = \tau = \frac{\rho\nu V}{L} = \frac{\rho\nu\kappa}{L^2}$$

based on the characteristic viscous shear stress  $\tau$  in the fluid, and

$$P_2 = KE = \rho V^2 = \frac{\rho\kappa^2}{L^2}$$

which is based on the characteristic kinetic energy density  $KE$  of the flow. Their ratio is equal to the Prandtl number :

$$\frac{P_1}{P_2} = \frac{\tau}{KE} = \frac{\frac{\rho\nu\kappa}{L^2}}{\frac{\rho\kappa^2}{L^2}} = \frac{\frac{\nu}{L^2}}{\frac{\kappa}{L^2}} = \frac{T_{conduction}}{T_{viscosity}} = \frac{\nu}{\kappa} = Pr \iff P_2 = \frac{1}{Pr}P_1$$

A commonly used approximation for the dynamics of Earth's mantle is the limit of infinite viscosity, which is equivalent to the limit of infinite Prandtl number since thermal conductivity is usually fixed or bounded: [23, p. 267]

$$\nu \longrightarrow \infty \iff Pr \longrightarrow \infty \implies P_2 \longrightarrow 0$$

That means that the viscous shear stress dominates the pressure balance in highly viscous flows and that the kinetic energy of such flows is negligible.

In addition, since the scale associated with  $P_2$  vanishes, it implies that the gradient of pressure term  $\nabla p$  in the equation (8) vanishes as well. This contradicts the physical fact that pressure influences highly viscous ("creeping") flows. [14, pp. 433-434]

Thus, the appropriate choice of characteristic pressure scale is the one associated with viscous shear stress:

$$P = P_1 = \tau = \frac{\rho\nu\kappa}{L^2}$$

## 7. Nondimensional Form of the Equations of Boussinesq Approximation

Substituting the scales discussed above into the equations (8), (9), and (10), we obtain the following:

$$\Delta \mathbf{v} = \nabla p - \frac{g\Theta L^3}{\nu\kappa} \theta \mathbf{e}_z \quad (11)$$

$$\nabla \cdot \mathbf{v} = 0 \quad (12)$$

$$\partial_t \theta + \mathbf{v} \cdot \nabla \theta = \Delta \theta \quad (13)$$

Considering the definition of Rayleigh number

$$Ra = \frac{g\Theta\beta L^3}{\nu\kappa}$$

we obtain the following system of equations in nondimensional form:

$$\Delta \mathbf{v} = \nabla p - Ra \theta \mathbf{e}_z \quad (14)$$

$$\nabla \cdot \mathbf{v} = 0 \quad (15)$$

$$\partial_t \theta + \mathbf{v} \cdot \nabla \theta = \Delta \theta \quad (16)$$

As we can see, the only governing parameter of the system is the Rayleigh number, which has made this approximation widely used for numerical experiments and validation of numerical algorithms for solving mantle convection equations.

### 7.1. Physical Meaning of the Rayleigh Number

According to G.I. Barenblatt, "the parameter  $Ra$ , the Rayleigh number, is named after the great English physicist Rayleigh who was the first to study the onset of convection in a horizontal layer theoretically," [11, p. 47] and is a measure of the influence of temperature gradients on the mechanical motion of the fluid.

Since

$$Ra = \frac{g\Theta\beta L^3}{\nu\kappa} = \frac{g\Theta\beta LT}{\nu} = \frac{g\Theta\beta T}{\frac{\nu}{L}} = \frac{\rho V_{buoyancy}}{\rho V_{viscosity}}$$

it can be interpreted as a ratio of momentum gained due to the buoyant force, resulting from thermal expansion of the fluid, and the momentum dissipated by the viscosity.

Another interpretation of Rayleigh number is as a ratio of the work of buoyant force and viscous shear stress:

$$Ra = \frac{g\Theta\beta L^3}{\nu\kappa} = \frac{\rho g\Theta\beta L}{\frac{\rho\nu\kappa}{L^2}} = \frac{Work_{buoyancy}}{Work_{stress}}$$

In other words, Rayleigh number measures how far from the equilibrium a system is, or is a measure of the instability of the system. Indeed, while the buoyant force destabilizes the motion, viscosity stabilizes it through energy dissipation.

## A. Buckingham's $\Pi$ Theorem

The Buckingham  $\Pi$ -Theorem is discussed both from the perspective of revealing the rank of a set of physical quantities, as well as from the perspective of reducing the number of parameters in an equation linking physical quantities. It is shown

that the Buckingham  $\Pi$ -Theorem is simply a consequence of a fundamental theorem of linear algebra relating the dimensions of the image, kernel, and the number of columns of a matrix.

### A.1. Content of Buckingham's $\Pi$ -Theorem

The content of Buckingham's  $\Pi$ -theorem is an answer to the question:

Why out of  $n$  physical quantities, of which  $k$  are dimensionally independent, can we obtain  $n - k$  dimensionless parameters?

The answer is simply a consequence of a fundamental theorem of linear algebra. [4, 13, 24, 27]

### A.2. Physical Dimensions Form a Linear Space

Physical quantities are measured against a conventionally chosen system of units  $\mathbf{X} = ([X_1] \ \dots \ [X_n])$ . Therefore, the dimension of any physical quantity is simply an algebraic combination of the units of measurements:

$$[A] = \left[ \prod_{i=1}^n X_i^{d_A^i} \right] = \sum_{i=1}^n d_A^i [X_i]$$

For example, in the unit system of mass  $M$ , length  $L$ , and time  $T$ , the dimension of  $P$  pressure is

$$[P] = \frac{M}{LT^2} = [M] - [L] - 2[T]$$

We see that we can establish an isomorphism between the dimensions of physical quantities and a linear space of vectors corresponding to these dimensions.

Indeed, consider the dimension of a product of two physical quantities:

$$[AB] = \left[ \prod_{i=1}^n X_i^{d_A^i} \prod_{i=1}^n X_i^{d_B^i} \right] = \left[ \prod_{i=1}^n X_i^{d_A^i} X_i^{d_B^i} \right] = \left[ \prod_{i=1}^n X_i^{d_A^i + d_B^i} \right] = \sum_{i=1}^n (d_A^i + d_B^i) [X_i] = [A] + [B]$$

Evidently, vectors  $[X_i]$  form a basis of this linear space of dimensions. It is convenient to choose a canonical orthonormal basis for the fundamental units of measurements:



$$[X_i] \equiv \begin{pmatrix} 0 \\ \vdots \\ 1 \\ \vdots \\ 0 \end{pmatrix} \text{ where 1 is at the } i\text{-th component and 0 elsewhere.}$$

For example, if we establish the following correspondence between the basis vectors and the dimensions of mass  $M$ , length  $L$ , and time  $T$ :

$$[M] = \begin{pmatrix} 1 \\ 0 \\ 0 \end{pmatrix}, \quad [L] = \begin{pmatrix} 0 \\ 1 \\ 0 \end{pmatrix}, \quad [T] = \begin{pmatrix} 0 \\ 0 \\ 1 \end{pmatrix}$$

the dimension of pressure will be expressed thus:

$$[P] = \frac{M}{LT^2} = [M] - [L] - 2[T] = \begin{pmatrix} 1 \\ -1 \\ -2 \end{pmatrix}$$

### A.3. Dimensional Matrix

Considering these facts and recalling the definition of matrix-vector multiplication, we can see that  $\mathbf{X} = ([X_1] \ \dots \ [X_n])$  forms a *dimensional matrix* of the system of physical quantities. Thus, the dimension vector  $d_A$  of any physical quantity  $[A]$  in the basis  $\mathbf{X} = ([X_1] \ \dots \ [X_n])$  is linearly related to the dimensional matrix  $\mathbf{X}$ :

$$[A] = \sum_{i=1}^n d_A^i [X_i] = \mathbf{X} d_A$$

#### A.3.1. Dimensionless Quantities and Kernel of the Dimensional Matrix

A physical quantity  $[A]$  is *dimensionless* if

$$A = \prod_{i=1}^n X_i^{d_A^i} = 1$$

which means that

$$[A] = \sum_{i=1}^n d_A^i [X_i] = \mathbf{X} d_A = [1] = 0 \implies d_A \in \ker \mathbf{X}$$

Thus, there is a one-to-one correspondence between the kernel, or nullspace, of the dimensional matrix  $\mathbf{X}$  and the nondimensional quantities that can be formed out of the system of physical quantities  $\mathbf{X} = ([X_1] \ \dots \ [X_n])$ . Therefore, the number of distinct nondimensional numbers that it is possible to produce out of a system of dimensional quantities  $(X_1, \dots, X_n)$  is equal to  $\dim \ker \mathbf{X}$ , i.e. the dimension of the nullspace of the dimension matrix  $\mathbf{X}$ .

### A.3.2. Dimensionally Independent Physical Quantities and the Rank of the Dimensional Matrix

Physical quantities  $(X_1, \dots, X_n)$  are dimensionally independent if their only combination that can produce a dimensionless number is a trivial combination. In other words, the dimension vectors of dimensionally independent quantities are linearly independent:

$$\prod_{i=1}^n X_i^{d^i} = 1 \iff \sum_{i=1}^n d^i [X_i] = 0 \implies d^i = 0$$

Thus, the dimension matrix composed of the dimensions of  $n$  dimensionally independent physical quantities has full rank:

$$\text{rank } \mathbf{X} = n$$

## A.4. Buckingham's Theorem

Now it is possible to prove Buckingham's  $\Pi$  theorem. Usually, the number of dimensionally independent physical quantities in the system is known, and it is interesting to know how many dimensionless quantities can be produced out of this system. Such question is of interest since it is preferable to work with as little number of parameters as possible.

Let  $(X_1, \dots, X_k, \dots, X_n)$  be a system of  $n$  physical quantities of which the first  $k$  are dimensionally independent. The dimensional matrix of this system

$$\mathbf{X} = ([X_1], \dots, [X_k], \dots, [X_n])$$

has  $k$  linearly independent columns and, hence, its rank is  $k$ :

$$\dim \text{im } \mathbf{X} = \text{rank } \mathbf{X} = k$$

By a fundamental theorem of linear algebra, the sum of the dimensions of the image and the kernel of a linear operator is equal to the number of columns of its matrix:

$$\dim \ker \mathbf{X} + \dim \text{im } \mathbf{X} = \dim \ker \mathbf{X} + k = n$$

Since it was established that the number of distinct nondimensional numbers characterizing the system  $(X_1, \dots, X_k, \dots, X_n)$  is equal to the dimension of the nullspace of its dimensional matrix, we obtain the result of the  $\Pi$ -theorem:

$$\dim \ker \mathbf{X} = n - k$$

Thus, given  $n$  physical quantities of which  $k$  are dimensionally independent, we can form  $n - k$  nondimensional numbers fully characterizing the system. This allows us to reduce the number of parameters needed to describe the system.

In addition, we established that, given a dimensional matrix of a set of physical quantities, the problem of finding the number of dimensionally independent quantities among them reduces to the problem of computing the rank of the dimensional matrix.

### **A.5. Dimensionality Reduction of a Physical Functional Dependence**

We have now established that given  $n$  physical quantities of which  $k$  are dimensionally independent, we can form  $n - k$  nondimensional numbers fully characterizing the system. However, what does this imply for functional dependencies of one physical quantity on others?

Based on the previous result we intuitively expect that a functional dependence of the form

$$X_0 = f(X_1, \dots, X_k, \dots, X_n)$$

could be described by  $n - k$  nondimensional parameters instead of  $n$  original dimensional parameters.

Indeed, let  $(X_0, X_1, \dots, X_k, \dots, X_n)$  be a system of physical quantities of which  $k$  are dimensionally independent: the dimensionally independent quantities will be used as the units of measurement.

Furthermore,  $X_0$  is a function of other quantities:

$$X_0 = f(X_1, \dots, X_k, \dots, X_n) \quad (17)$$

The dimensions of  $X_0$  and  $(X_{k+1}, \dots, X_n)$  can be expressed in terms of the dimensions of the dimensionally independent quantities  $(X_1, \dots, X_k)$ , since they are our units of measurement:

$$\begin{aligned} [X_0] &= \sum_{i=1}^k d_0^i [X_i] = \left[ \prod_{i=1}^k X_i^{d_0^i} \right] \\ [X_{k+j}] &= \sum_{i=1}^k d_j^i [X_i] = \left[ \prod_{i=1}^k X_i^{d_j^i} \right], \quad j \in \overline{1, n-k} \end{aligned}$$

If we change the scales for the units of measurements, i.e.

$$X_i \mapsto a_i X_i, \quad i \in \overline{1, k} \quad (18)$$

the dimensionally-dependent quantities must transform accordingly:

$$X_0 \mapsto \prod_{i=1}^k (a_i X_i)^{d_0^i} = \left( \prod_{i=1}^k a_i^{d_0^i} \right) \prod_{i=1}^k X_i^{d_0^i} = \left( \prod_{i=1}^k a_i^{d_0^i} \right) X_0 \quad (19)$$

$$X_{k+j} \mapsto \prod_{i=1}^k (a_i X_i)^{d_j^i} = \left( \prod_{i=1}^k a_i^{d_j^i} \right) \prod_{i=1}^k X_i^{d_j^i} = \left( \prod_{i=1}^k a_i^{d_j^i} \right) X_{k+j}, \quad j \in \overline{1, n-k} \quad (20)$$

Since  $(X_1, \dots, X_k)$  are dimensionally independent units of measurement, it is possible to choose scales based on these quantities, i.e.

$$a_i = \frac{1}{X_i}, \quad i \in \overline{1, k}$$

In such case, (18), (19), and (20) become

$$a_i X_i = 1, \quad i \in \overline{1, k} \quad (21)$$

$$\left(\prod_{i=1}^k a_i^{d_0^i}\right) X_0 = \frac{X_0}{\left(\prod_{i=1}^k X_i^{d_0^i}\right)} = \Pi \quad (22)$$

$$\left(\prod_{i=1}^k a_i^{d_j^i}\right) X_{k+j} = \frac{X_{k+j}}{\left(\prod_{i=1}^k X_i^{d_j^i}\right)} = \Pi_j, \quad j \in \overline{1, n-k} \quad (23)$$

Consequently, the physical functional dependence (17) transforms accordingly under the change of scales (21), (22), and (23):

$$X_0 = f(X_1, \dots, X_k, \dots, X_n) \mapsto \Pi = f(1, \dots, 1, \Pi_1, \dots, \Pi_{n-k}) = F(\Pi_1, \dots, \Pi_{n-k})$$

Thus, a dimensional functional dependence involving  $n$  dimensional parameters

$$X_0 = f(X_1, \dots, X_k, \dots, X_n)$$

has been reduced to a nondimensional functional dependence involving only  $n - k$  dimensionless parameters

$$\Pi = F(\Pi_1, \dots, \Pi_{n-k})$$

This relationship is particularly useful, for example, in nondimensionalization of fluid dynamics equations because it reveals the relative influence of various terms of the equations on the overall behavior of the flow.

## References

- [1] *Computational Infrastructure for Geodynamics (CIG)*, <https://geodynamics.org/cig/software/#mc>.
- [2] Wolfgang Bangerth, Martin Kronbichler, and Timo Heister, *High accuracy mantle convection simulation through modern numerical methods*, Geophysical Journal International **191** (2012), 12–29.
- [3] H. Bénard, *Les tourbillons cellulaires dans une nappe liquide*, Rev. Gén. Sciences Pure Appl. **11** (1900), 1261–1271, 1309–1328.

- [4] G.W. Bluman and S. Kumei, *Symmetries and Differential Equations*, Springer-Verlag, 1989.
- [5] J. Boussinesq, *Theorie Analytique de la Chaleur*, vol. 2, Gauthier-Villars, 1903.
- [6] S. Chandrasekhar, *Hydrodynamic and Hydromagnetic Stability*, Oxford University Press, 1961.
- [7] B. Gebhart, *Heat Transfer*, 2 ed., McGraw-Hil, 1971.
- [8] Taras Gerya, *Introduction to Numerical Geodynamic Modelling*, Cambridge University Press, 2009.
- [9] A. V. Getling, *Rayleigh-Benard Convection*, Scholarpedia **7** (2012), no. 7, 7702.
- [10] A.V. Getling, *Rayleigh-Benard Convection: Structures and Dynamics*, World Scientific, 1998.
- [11] Barenblatt G.I., *Scaling*, Cambridge University Press, Cambridge, UK, 2003.
- [12] Donald D. Gray and Aldo Giorgini, *The Validity of the Boussinesq Approximation for Liquids and Gases*, Int. J. Heat Mass Transfer **19**, 545–551.
- [13] H. Hanche-Olsen, *Buckingham’s Pi-theorem*, <http://www.math.ntnu.no/~hanche/notes/buckingham/buckingham-a4.pdf>.
- [14] L. Gary Leal, *Advanced Transport Phenomena*, Cambridge University Press, 2007.
- [15] W. V. R. Malkus, *Boussinesq equations, Notes on the 1964 Summer Study Program in Geophysical Fluid Dynamics at the Woods Nole Oceanographic Institution*, U.S. National Technical Information Service **1** (1964), 1–12.
- [16] ———, *A scaling and expansion of equations of motion to yield the Boussinesq equations, Notes on the 1969 Summer Study Program in Geophysical Fluid Dynamics at the Woods Hole Oceanographic Institution*, U.S. National Technical Information Service **1** (1969), 23–28.
- [17] R. H. M. Miesen, L. P. J. Kamp, and F. W. Sluijter, *On the application of the Boussinesq approximation for nonlinear gravity waves in the atmosphere*, Physica Scripta **38** (2006), no. 6, 857.

- [18] J. M. Mihaljan, *A rigorous exposition of the Boussinesq approximations applicable to a thin layer of fluid*, Astrophys. Journal **136** (1962), 1126–1133.
- [19] A. Oberbeck, *Über die Wärmeleitung der Flüssigkeiten bei Berücksichtigung der Strömungen infolge von Temperatur Differenzen*, Ann. Phys. Chem. **7** (1879), 271–292.
- [20] A. Pellew and R. V. Southwell, *On maintained convective motion in a fluid heated from below*, Proc. Roy. Soc. A. **176** (1940), no. 966, 312–343.
- [21] E. J. Plate, *Aerodynamic Characteristics of Atmospheric Boundary Layers*, U.S. National Technical Information Service, 1971.
- [22] Lord Rayleigh, *On convection currents in a horizontal layer of fluid, when the higher temperature is on the under side*, Phil. Mag., Ser. **6** (1916), no. 32, 529–546.
- [23] Gerald Schubert, Donald L. Turcotte, and Peter Olson, *Mantle Convection in the Earth and Planets*, Cambridge University Press.
- [24] A.A. Sonin, *The Physical Basis of Dimensional Analysis*, MIT Department of Mechanical Engineering, 2001.
- [25] E. A. Spiegel and G. Veronis, *On the Boussinesq approximation for a compressible fluid*, Astrophys. Journal **131** (1960), 442–447.
- [26] Radyadour Kh. Zeytounian, *Joseph Boussinesq and his approximation: a contemporary view*, Comptes Rendus Mécanique **331** (2003), no. 8, 575–586.
- [27] V.A. Zorich, *Mathematical Analysis of Problems in the Natural Sciences*, MC-CME, Moscow, Russian Federation, 2008.