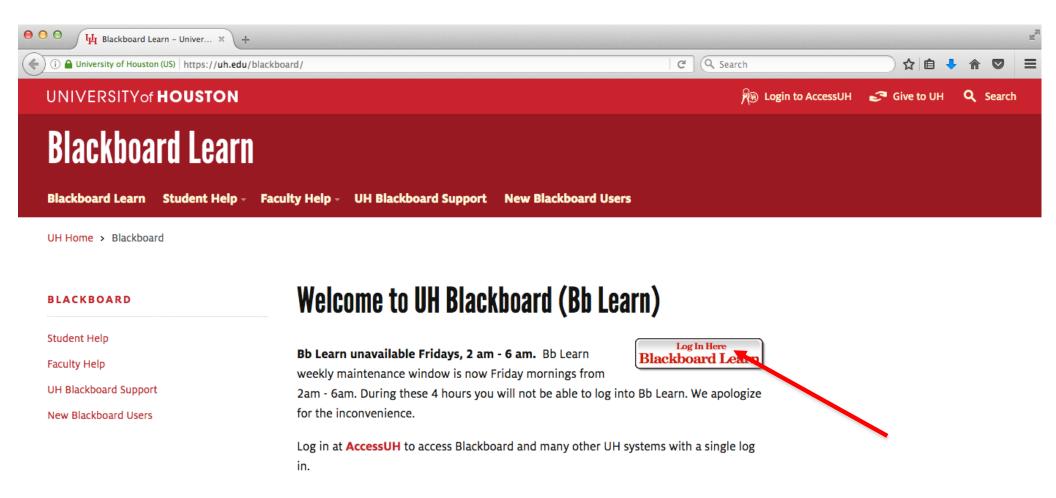
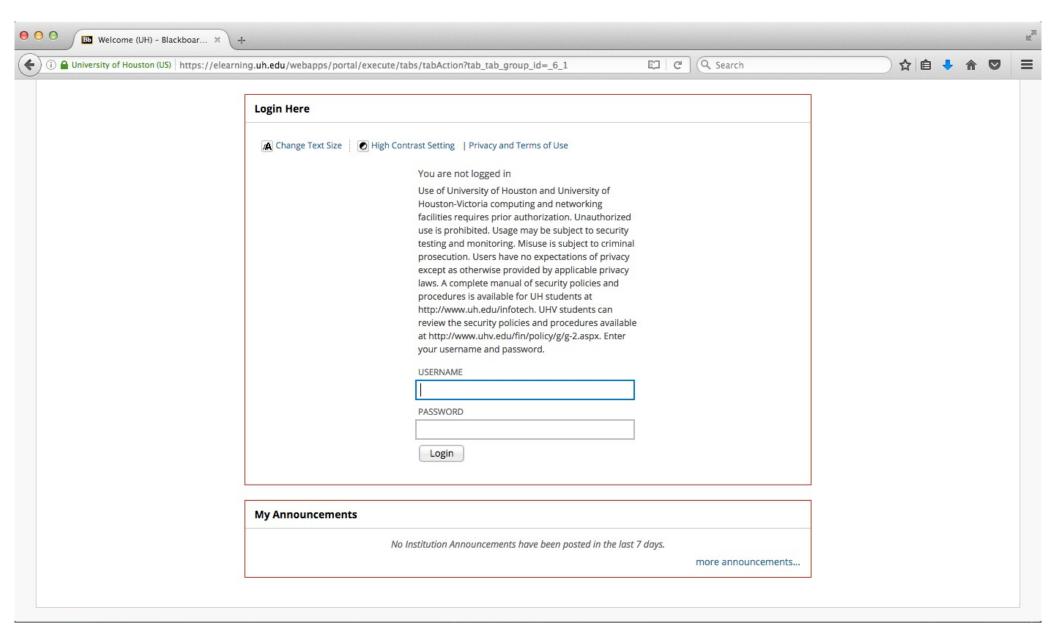


#### Exam 2

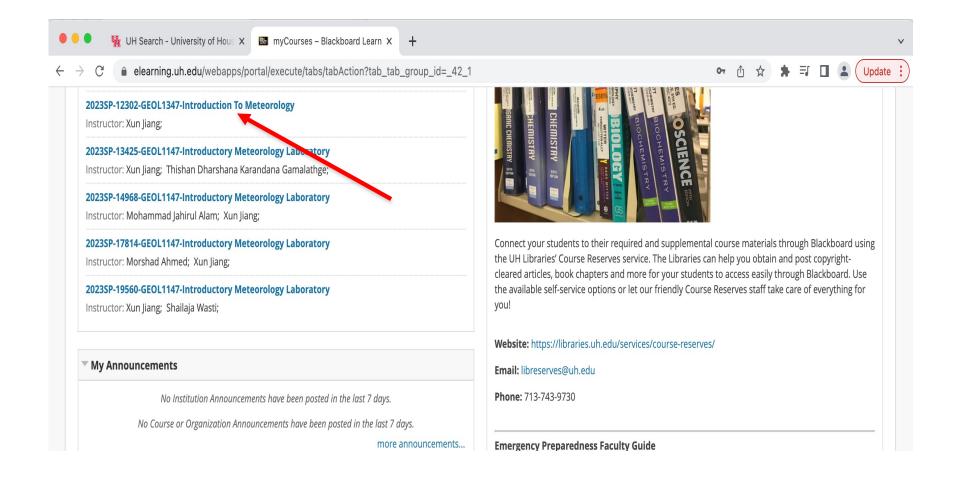
#### Online Test on Blackboard

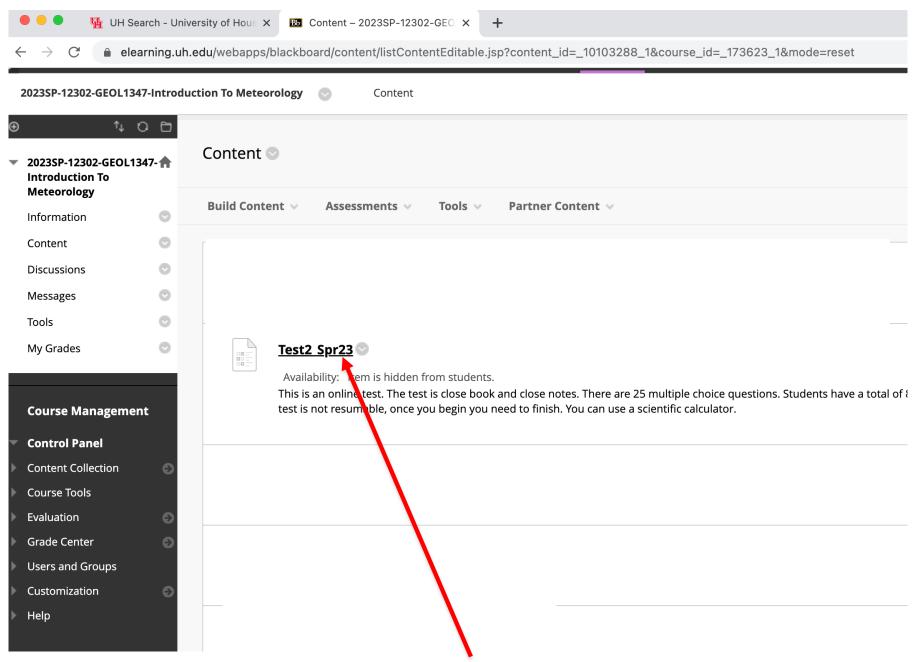
Website: https://uh.edu/blackboard/





Username and password are your Cougarnet username and password.





You should be able to see Test 2 after 1pm on Mar 2. Please finish the test during 1pm-2:30pm.

- The test is close book and close notes.
- There are 25 multiple choice questions.
- Time limit is 80 mins.
- Once started, the test must be completed in one sitting. Do not leave test before clicking save and submit.
- The test will save and submit automatically when the time expires.
- Please take the test during 1pm-2:30pm on Mar 2.
- You can use a scientific calculator.

#### Exam 2

Cover: L5 (Air Temperature); L6 (Humidity, Condensation, and Clouds-I); L7 (Humidity, Condensation, and Clouds-II); L8 (Cloud Development and Precipitation-I)

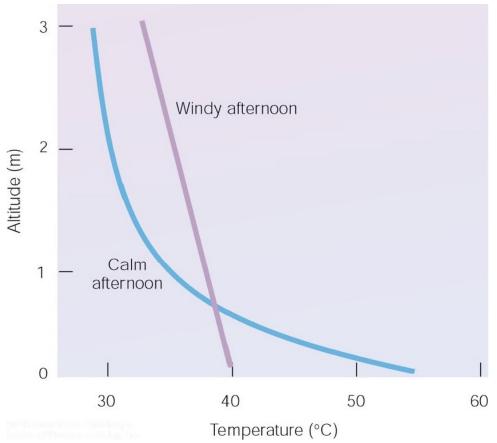
Close-book Exam

Exam counts 25% of the total grade.

### **Daytime Warming**

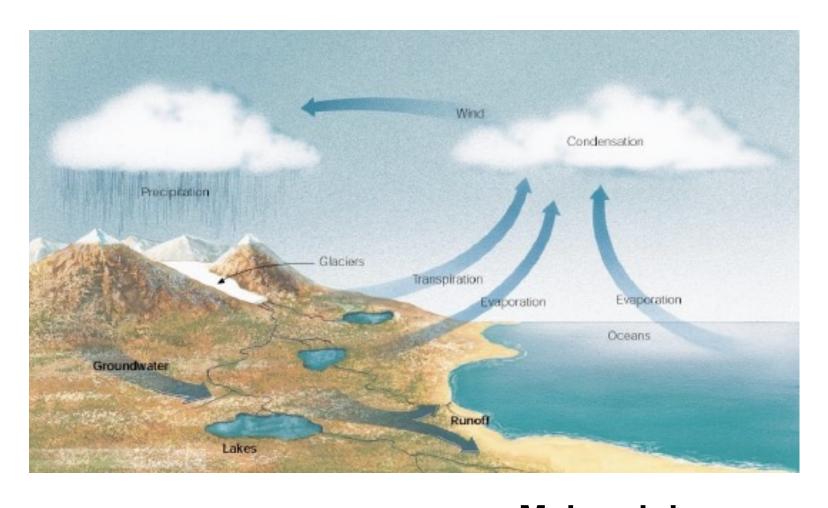
■ On windy days, turbulence eddies are able to mix hot, surface air with cooler air above.

■This form of mechanical stirring, called forced convection, helps the thermals to transfer heat away from the surface more efficiently.



Temperature gradient is smaller in windy day than calm day.

#### Where does the moisture in the atmosphere come from?

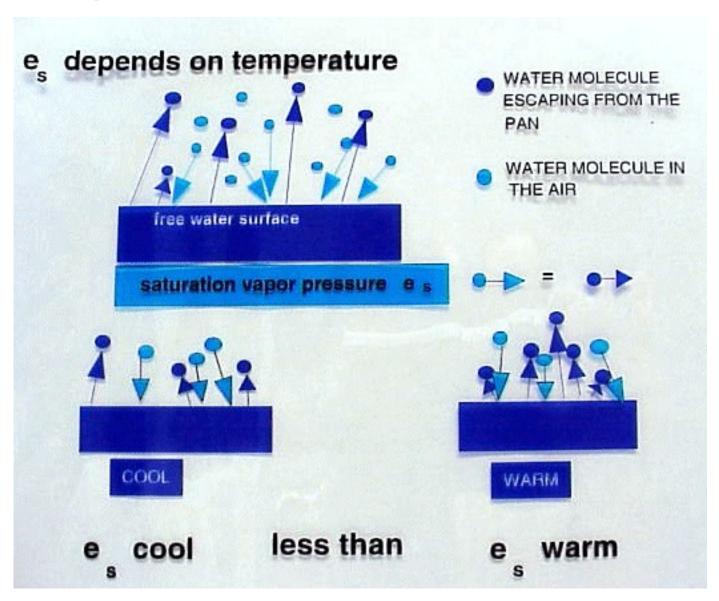


Major Source
Evaporation from ocean

Major sink

Precipitation

# Saturation vapor pressure $e_s$ depends upon temperature higher temperature, higher $e_{s_s}$ more water vapor that the air can hold



- The saturation vapor pressure of water increases with temperature
  - At higher T, faster water molecules in liquid escape more frequently causing equilibrium water vapor concentration to rise
  - We sometimes say "warmer air can hold more water vapor"
- There is also a vapor pressure of water over an ice surface
  - The saturation vapor pressure above solid ice is less than above liquid water,  $e_{s(water)} > e_{s(ice)}$  at all temperatures

# Absolute Humidity - ρ<sub>ν</sub>

- Density of water vapor
  - A measure of the total number (mass) of water vapor molecules in a unit volume of air (1 m<sup>3</sup>)
  - Absolute humidity = mass of water vapor / volume of air

$$\rho_v = m_v / V_{air_s} m_v = n_v M_v$$
,  $M_v = 18$  g/mol

Changes in volume cause changes in absolute humidity

### Specific Humidity - q

- Ratio of mass of water to total mass of air in a unit volume
- Invariant to change in volume

$$q = \frac{m_v}{m}$$

Since q is on the order of  $10^{-3}$  g  $_{\rm v}$ /g  $_{\rm a}$ , we prefer to use  $\rm \,g_{\rm v}$ /kg q values normally range from 1 to 20 g  $_{\rm v}$ /kg and decreases with increasing height

### **Relative Humidity – R.H.**

The ratio of the amount of water vapor in the air compared to the amount required for saturation.

R.H. = water vapor content / water vapor capacity

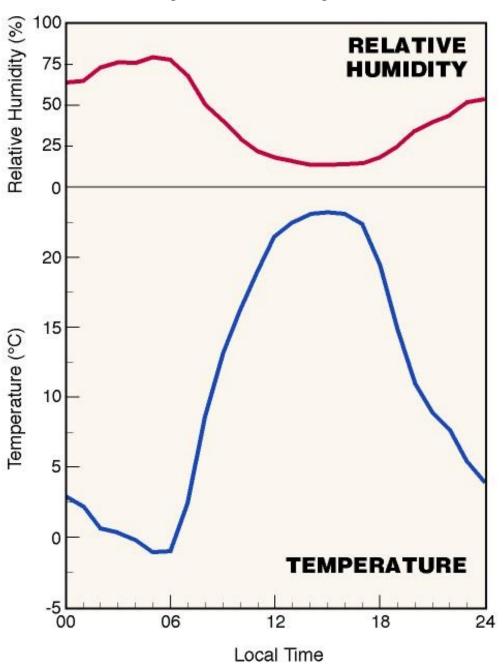
 $R.H = e/e_s$ 

e: Vapor pressure; e<sub>s</sub>: Saturation vapor pressure

#### Change of relative humidity in a day

What time of the day when relative humidity is usually high?

As the air cools during the night, the relative humidity increases. The highest relative humidity occurs in the early morning, during the coolest part of the day.



## **Dew Point Temperature - T<sub>d</sub>**

- Temperature to which air must be cooled (at constant pressure and constant water vapor content) to become saturated.
- When  $T=T_d$ ,  $e_s(T_d) = e$ ,  $q_s(T_d) = q$ ,  $r_s(T_d) = r$
- T<sub>d</sub> is less or equal to T
- Unlike relative humidity which is a measure of how near the air is to being saturated, dew point temperature is a measure of its actual moisture content. The higher the dew point, the more water vapor in the air.
- Dew point depression: T-T<sub>d</sub>
- The larger the dew point depression is, the drier the air is, or the air is farther away from saturation

# Summary of Cloud Types

	Layered	Broken Layer	Separate
High	Cirrostratus	Cirrocumulus	Cirrus
Middle	Altostratus	Altocumulus	
Low	Stratus Nimbostratus	Stratocumulus	
Vertical			Cumulus Cumulo- nimbus

# Dry adiabatic lapse rate – the rate of temperature decrease of a rising *unsaturated* air parcel

$$\Gamma_{\rm d} = -\Delta T/\Delta z = 9.8 \, {\rm ^{\circ}C \, km^{-1}} \approx 10 \, {\rm ^{\circ}C \, km^{-1}}$$

# Stability is determined by comparing parcel's temperature with that of its environment

#### Simply speaking,

$$T_{parcel} > T_{env}$$
 unstable  $T_{parcel} < T_{env}$  stable

$$T_{parcel} = T_{env}$$
 neutral

#### Stability and lapse rate

Absolutely stable  $\Gamma < \Gamma_s$ 

Absolutely unstable  $\Gamma > \Gamma_d$ 

Conditionally unstable  $\Gamma_d > \Gamma > \Gamma_s$ 

 $\Gamma = \Gamma_d$  for unsaturated air

Neutral

 $\Gamma = \Gamma_s$  for saturated air

 $\Gamma$ : Environmental Lapse Rate;  $\Gamma_d$ : Dry Adiabatic Lapse Rate;

 $\Gamma_s$ : Moist Adiabatic Lapse Rate