

Exam 5

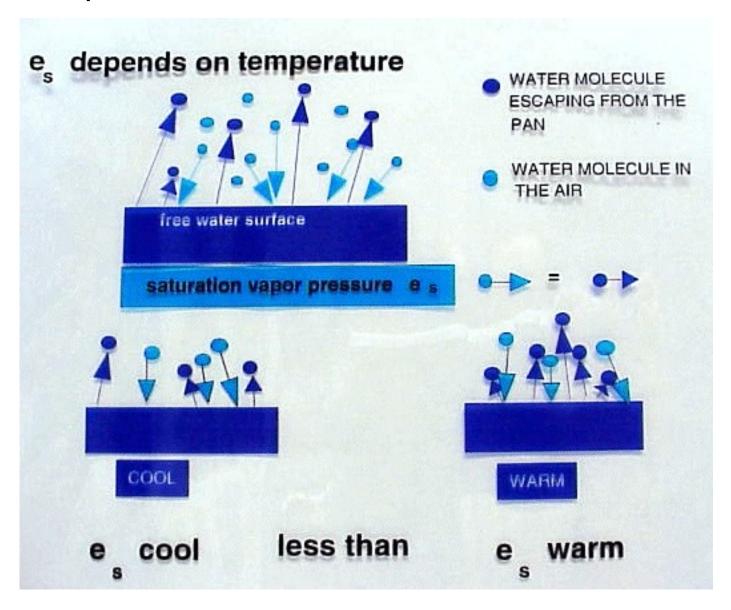
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Online Test on Blackboard
Website: https://uh.edu/blackboard/
Cover:
L6 (Humidity, Condensation, and Clouds-I);
L11 (Air Pressure & Winds-II);
L13 (Atmospheric Circulation-II);
L17 (Hurricanes);
L18 (Light, Color, & Atmospheric Optics)
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- The test is close book and close notes.
- There are 25 multiple choice questions.
- Time limit is 80 mins.
- Once started, the test must be completed in one sitting. Do not leave test before clicking save and submit.
- The test will save and submit automatically when the time expires.
- Please take the test during 1pm-2:30pm on Apr 27.
- You can use a scientific calculator.

Vapor pressure - e

- Air molecules all contribute to pressure p
- Each subset of molecules (e.g., N₂, O₂, H₂O) exerts a partial pressure
- The vapor pressure, e, is the pressure exerted by water vapor molecules in the air
 - similar to atmospheric pressure, but due only to the water vapor molecules
 - 2-30 mb common at surface
 - the larger the vapor pressure is, the more water vapor molecules in the atmosphere

Saturation vapor pressure e_s depends upon temperature higher temperature, higher e_{s_s} more water vapor that the air can hold



Specific Humidity - q

- Ratio of mass of water to total mass of air in a unit volume
- Invariant to change in volume

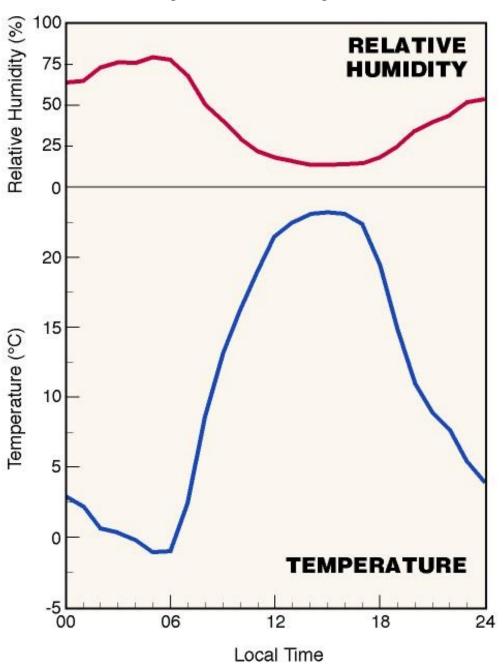


Since q is on the order of 10^{-3} g $_{\rm v}$ /g $_{\rm a}$, we prefer to use $\rm \,g_{\rm v}$ /kg q values normally range from 1 to 20 g $_{\rm v}$ /kg and decreases with increasing height

Change of relative humidity in a day

What time of the day when relative humidity is usually high?

As the air cools during the night, the relative humidity increases. The highest relative humidity occurs in the early morning, during the coolest part of the day.



Dew Point Temperature - T_d

- Temperature to which air must be cooled (at constant pressure and constant water vapor content) to become saturated.
- When $T=T_d$, $e_s(T_d) = e$, $q_s(T_d) = q$, $r_s(T_d) = r$
- T_d is less or equal to T
- Unlike relative humidity which is a measure of how near the air is to being saturated, dew point temperature is a measure of its actual moisture content. The higher the dew point, the more water vapor in the air.
- Dew point depression: T-T_d
- The larger the dew point depression is, the drier the air is, or the air is farther away from saturation

Coriolis Force (CF)

- Apparent force due to the rotation of the earth
- Magnitude depends on latitude and the speed of the air parcel

The higher the latitude, the larger the Coriolis force
Zero at the equator, and maximum at the poles
The faster air moves, the larger the Coriolis force

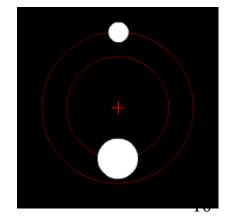
Causes the parcel to deflect

to the right of its intended path in the northern hemisphere to the left of its intended path in the southern hemisphere.

Only influence wind direction, no effect on wind speed!

Centrifugal Force

- Magnitude CENTF = mV²/R
 - m is the mass
 - R the radius of curvature of the curved path
 - V is the speed of the air parcel
- Direction
 - Pointing away from the center of the curve
 - The faster the speed and the tighter the curve of the path traveled (i.e., the smaller R), the larger the centrifugal force.



Frictional Force

Frictional drag of the ground slows wind down.

$$FF = -kV$$

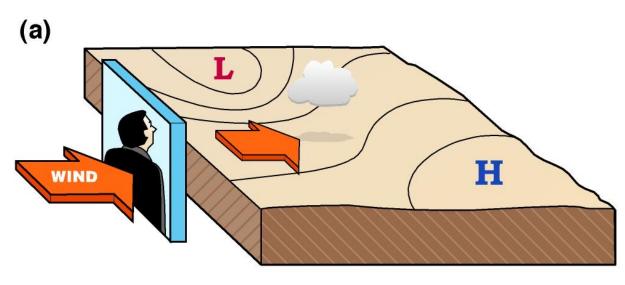
- Magnitude
 - Depends upon the speed of the air parcel (V)
 - Depends upon the roughness of the earth's surface (k)
- Direction
 - Always acts in the direction opposite to the movement of the air parcel (minus sign emphasizes this)
- Important in the friction layer (planetary boundary layer)
 - ~lowest 1000 m of the atmosphere

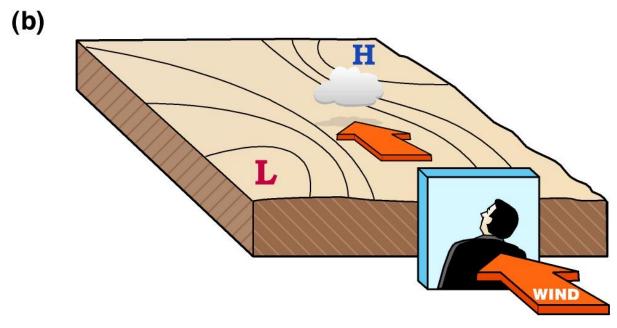
Geostrophic Wind

Geostrophic balance - the most fundamental horizontal force balance when the pressure gradient force is counter balanced by the Coriolis force

$$PGF + CF = 0$$

The resulting wind is called Geostrophic wind. Geostrophic wind is good approximation to winds above earth's surface





Buys Ballot's Law:

Stand with your back towards the wind (upper-level wind), low pressure is on your left and high pressure on your right.

Hadley cell (thermal direct)

Air rises near equator and descends near 30° where a belt of high pressure is found (subtropical high)

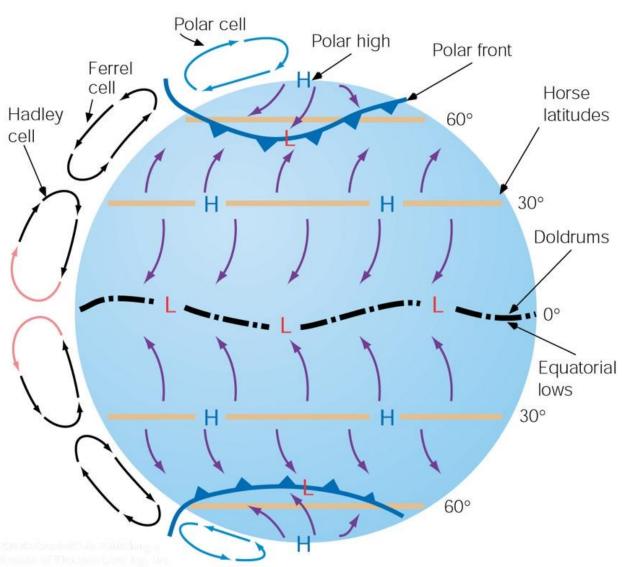
Polar Cell (thermal direct)

Cold air moving southward from the pole rises when it reaches the *polar front* near 60°. The rising air aloft returns to the north where it sinks back to the pole.

Ferrel Cell (thermal indirect)

A reverse circulation between the Hadley and Polar Cell (between 30° and 60°) that carries air northward near the surface and southward aloft

A rotating earth break the single cell into Three Cells



Wind patterns in the three-cell model

equatorial region and 30° latitude belt

Easterly winds dominate

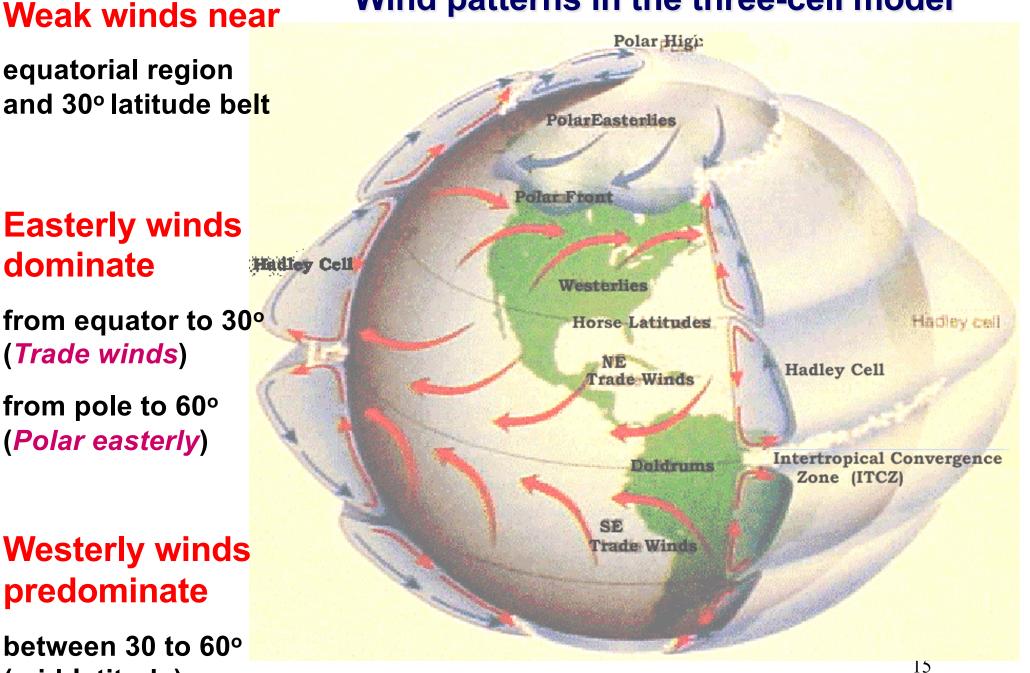
from equator to 30°

(Trade winds)

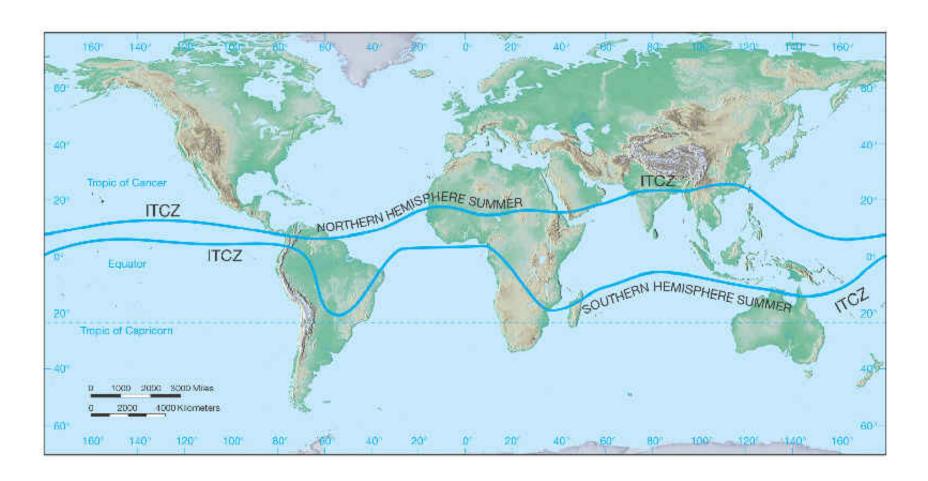
from pole to 60° (Polar easterly)

Westerly winds predominate

between 30 to 60° (mid-latitude)

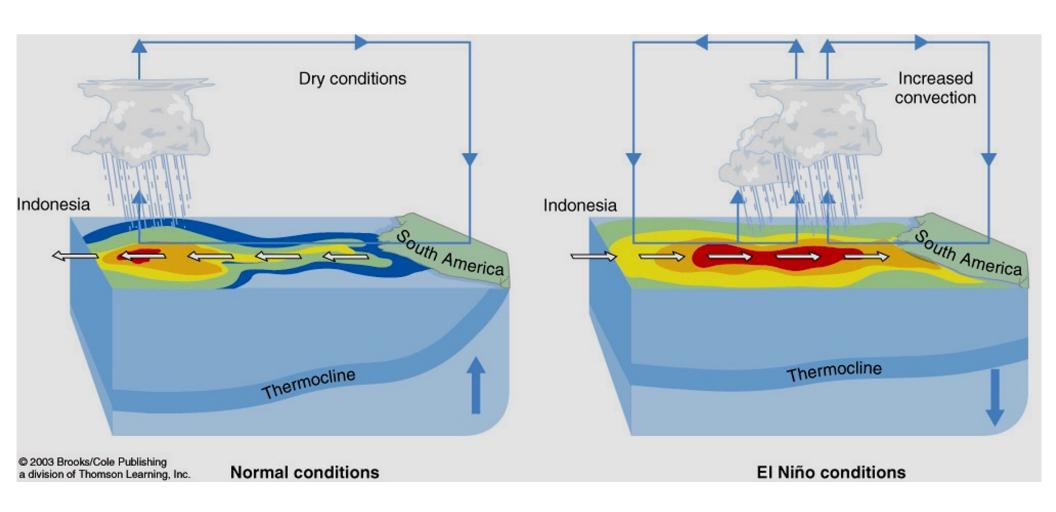


Intertropical Convergence Zone (ITCZ)



- ITCZ moves back and forth across equator following the sun's zenith point.
- Variation in the ITCZ locations affect rainfall in the tropics.
- ITCZ shifts toward south in January and toward north in July.

El Nino

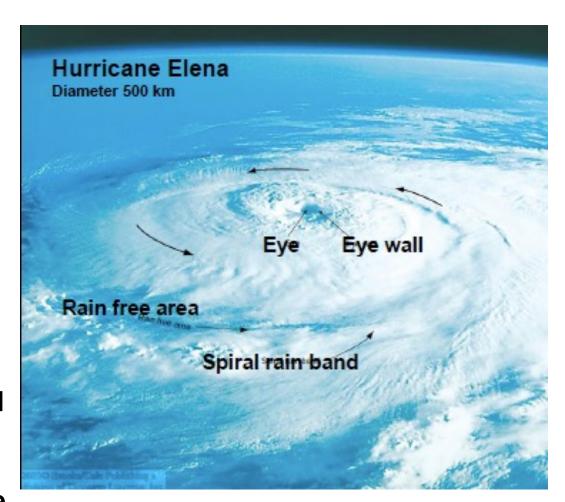


Anatomy of a Hurricane

- Hurricane is an intense storm of tropical origin, with sustained winds exceeding 64 knots (74 mi/hr), which forms over the warm northern Atlantic and eastern North Pacific oceans.
- In the western North Pacific, it is called typhoon, in India a cyclone, and in Australia a tropical cyclone. By international agreement, tropical cyclone is the general term for all hurricane-type storms that originate over tropical waters.

Anatomy of a Hurricane

- Diameter 500 km
- Eye 40 km; Area of broken clouds at the center; Winds are light
- Surface Air Pressure 955 mb
- Surface winds increase in speed as they blow counterclockwise and inward toward this center (NH).
- Adjacent to the eye is the eyewall, a ring of intense thunderstorms that whirl around the storm's center and may extend upward to almost 18 km.
- Within the eyewall, we find the heaviest precipitation and the strongest winds.



Aerosol Scattering

- Refraction depends on wavelength
- 1. Refraction bends short wavelengths (e.g. blue) more than long wavelengths (e.g. red).
- 2. Examples for refraction:
 - --- Dispersion of white light into individual colors through a glass/ice crystal prism
 - --- Refraction of starlight by the atmosphere



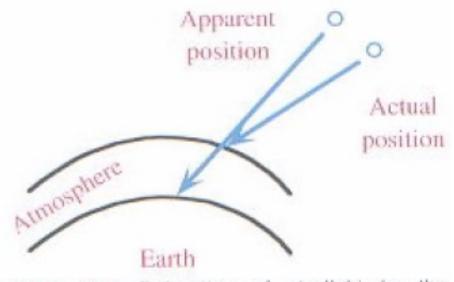


Figure 7.13. Refraction of starlight by the atmosphere makes stars appear to be where they are not.

Scattering

- Rayleigh Scattering
- The selective scattering of blue light by air molecules and very small particles can make distant mountains appear blue.



The blue ridge mountains in Virginia.