



# Review for Exam 2

*GEOL 1147: Introduction To Meteorology Lab*

## Exam 2

Cover: Labs 3-5

Close-book Exam

You can bring a calculator with you.

Exam counts 22.5% of the total grade.

# Basic laws for radiation

***Stefan-Boltzman law:*** The amount of energy per square meter per second that is emitted by an blackbody is related to the **4<sup>th</sup> power** of its Kelvin temperature

$$E = \sigma T^4$$

where  $E$  is in  $\text{J s}^{-1} \text{m}^{-2}$  or Watts  $\text{m}^{-2}$

$\sigma = 5.67 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4}$  ***Stefan-Boltzman constant***

As  $T$  increases,  $E$  increases by a power of 4. If  $T$  doubles,  $E$  increases by 16 times!

## *Wien's law:*

Wavelength of peak radiation emitted by an object is **inversely** related to temperature

$$\lambda_{\max} = 2897 / T \sim 3000/T$$

( $\lambda_{\max}$  is in  **$\mu\text{m}$**  and T is in **Kelvin**)

Solar radiation :  $\lambda_{\max \text{ sun}} \sim 3000/6000 \text{ K} \sim 0.5 \mu\text{m}$ ,

Earth radiation:  $\lambda_{\max \text{ earth}} \sim 3000/300 \text{ K} \sim 10 \mu\text{m}$ ,

Solar radiation is **shortwave** radiation

Earth radiation is **longwave** radiation

# Greenhouse Gases

Name of the gas	Molecular weight		Percentage
• Water vapor	H <sub>2</sub> O	18.02	< 4.0%
• Carbon dioxide	CO <sub>2</sub>	44.01	0.038%
• Methane	CH <sub>4</sub>	16.04	0.00017%
• Nitrous oxide	N <sub>2</sub> O	44.01	0.00003%
• Ozone	O <sub>3</sub>	48.00	0.000004%

# Albedo

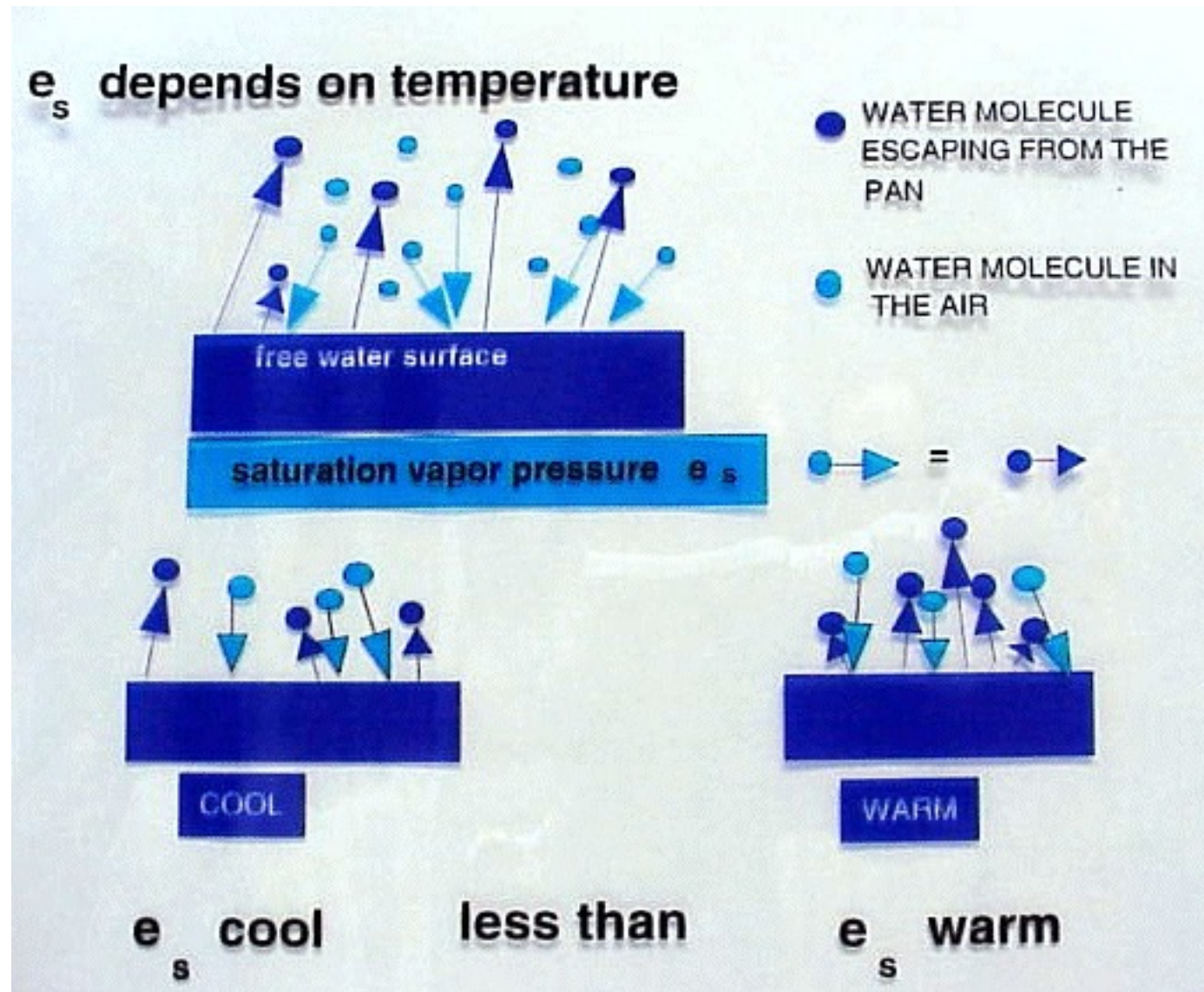
SURFACE	ALBEDO (PERCENT)
Fresh snow	75 to 95
Clouds (thick)	60 to 90
Clouds (thin)	30 to 50
Venus	78
Ice	30 to 40
Sand	15 to 45
Earth and atmosphere	30
Mars	17
Grassy field	10 to 30
Dry, plowed field	5 to 20
Water	10*
Forest	3 to 10
Moon	7
*Daily average.	

# Vapor pressure - e

- Air molecules all contribute to pressure  $p$
- Each subset of molecules (e.g.,  $N_2$ ,  $O_2$ ,  $H_2O$ ) exerts a partial pressure
- The vapor pressure,  $e$ , is the pressure exerted by water vapor molecules in the air
  - similar to atmospheric pressure, but due only to the water vapor molecules
  - 2-30 mb common at surface
  - the larger the vapor pressure is, the more water vapor molecules in the atmosphere



Saturation vapor pressure  $e_s$  depends upon temperature  
higher temperature, higher  $e_s$ ,  
more water vapor that the air can hold





## Mixing Ratio - r

- Ratio of mass of water to mass of dry air in a unit volume
- Invariant to change in volume

$$r = \frac{m_v}{m_d}$$

## Relative Humidity – R.H.

The ratio of the amount of water vapor in the air compared to the amount required for saturation.

R.H. = water vapor content / water vapor capacity

$$\text{R.H.} = \frac{e}{e_s(T)} = \frac{\tilde{n}_v}{\tilde{n}_{vs}(T)} = \frac{q}{q_s(T)} = \frac{r}{r_s(T)}$$

*Higher relative humidity does not necessarily mean more water vapor in the air*

# Dew Point Temperature - $T_d$

- Temperature to which air must be cooled (at constant pressure and constant water vapor content) to become saturated.
- When  $T=T_d$ ,  $e_s(T_d) = e$ ,  $q_s(T_d) = q$ ,  $r_s(T_d) = r$
- $T_d \leq T$
- Unlike relative humidity which is a measure of how near the air is to being saturated, dew point temperature is a measure of its actual moisture content. *The higher the dew point, the more water vapor in the air.*
- **Dew point depression:**  $T - T_d$
- The **larger** the dew point depression is, the **drier** the air is, or the air is farther away from saturation

# Moist adiabatic lapse rate

Decrease of temperature with height for saturated air

$$\Gamma_s \text{ always } < \Gamma_d$$