Second Monitoring Committee Meeting

Understanding the information content in diverse observations of forest carbon stocks and fluxes for data assimilation and ecological modelling.

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Project Background

A large amount of data is currently being gathered that is relevant to the carbon balance of forests, with much of this data coming from Eddy covariance flux towers [1]. Attempts are also being made to combine this data with models of forest carbon stocks and fluxes, such as the Data Assimilation Linked Ecosystem Carbon model (DALEC) [6], in a data assimilation (DA) scheme. Currently, however, there are limitations with such schemes as there is a lack of understanding about the additional information provided by different observations. Current DA schemes for ecosystem carbon flux only specify the diagonal elements of the observation error covariance matrix, **R**, which correspond to the individual uncertainties in particular observations. As such, these DA schemes do not specify observation error correlations or covariances, corresponding to the off diagonal elements of **R**. In numerical weather prediction (NWP) it has been shown that the inclusion of observation error correlations can increase the information content from a given data set and reduce analysis error [5]. Better understanding the information content and error correlations of carbon balance observations will form one of the main aims of the project.

The DALEC model

The DALEC model is a simple process-based model describing the carbon balance of an evergreen forest ecosystem. The model is constructed of five carbon pools (foliage (C_f) , fine roots (C_r) , woody stems and coarse roots (C_w) , fresh leaf and fine root litter (C_l) and soil organic matter and coarse woody debris (C_s) linked via fluxes. The gross primary production function (GPP) uses meteorological driving data and the site's leaf area index (a function of C_f) to calculate the total amount of carbon to be allocated at a daily time step. The model equations for the carbon pools at day t+1 are as follows:

$$C_f(t+1) = (1-p_5)C_f(t) + p_3(1-p_2)GPP(C_f(t),\phi),$$

$$C_r(t+1) = (1-p_7)C_r(t) + p_4(1-p_3)(1-p_2)GPP(C_f(t),\phi),$$

$$C_w(t+1) = (1-p_6)C_w(t) + (1-p_4)(1-p_3)(1-p_2)GPP(C_f(t),\phi),$$

$$C_l(t+1) = (1-(p_1+p_8)T(t))C_l(t) + p_5C_f(t) + p_7C_r(t),$$

$$C_s(t+1) = (1-p_9T(t))C_s + p_6C_w(t) + p_1T(t)C_l(t).$$

Where $T(t) = \frac{1}{2}exp(p_{10}T_m(t))$, T_m is daily mean temperature, p_1, \ldots, p_{10} are rate parameters and ϕ represents the meteorological driving data used in the GPP function. The full details of this version of DALEC can be found in [6].

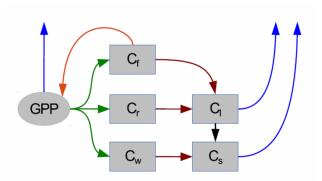


Figure 1: DALEC carbon balance model [2]

Shannon Information Content and Observation Sensitivity

In DA Shannon Information Content (SIC) is a measure of the reduction in entropy given a set of observations. Entropy physically corresponds to the volume in state space taken up by the probability density function (pdf) describing the knowledge of the state [4]. Assuming all pdfs are Gaussian we have,

$$SIC = \frac{1}{2} ln \frac{\left| \mathbf{B} \right|}{\left| \mathbf{A} \right|}$$

where \mathbf{B} is the background error covariance matrix and \mathbf{A} is the analysis error covariance matrix. For a larger reduction in uncertainty in our analysis we have a larger value of SIC. I began by using SIC to understand the information content for different sets of observations at one time when being assimilated with the DALEC model, specifying the state vector for the assimilation as,

$$\underline{x}_b = (C_f, C_r, C_w, C_l, C_s)^T.$$

Where the elements of the state vector have variances, $\sigma^2_{cf,b}, \ldots, \sigma^2_{cs,b}$, respectively. Giving a background error covariance,

$$\mathbf{B} = egin{pmatrix} \sigma_{cf,b}^2 & 0 & 0 & 0 & 0 \ 0 & \sigma_{cr,b}^2 & 0 & 0 & 0 \ 0 & 0 & \sigma_{cw,b}^2 & 0 & 0 \ 0 & 0 & 0 & \sigma_{cl,b}^2 & 0 \ 0 & 0 & 0 & 0 & \sigma_{cs,b}^2 \end{pmatrix}.$$

Here I have assumed a diagonal background error covariance matrix. For A^{-1} we have,

$$\mathbf{A}^{-1} = \mathbf{J}'' = \mathbf{B}^{-1} + \mathbf{H}^T \mathbf{R}^{-1} \mathbf{H}.$$

Where \mathbf{J}'' is the Hessian of \mathbf{J} , the cost function to be minimized in Three-Dimensional Variational Data Assimilation (3D-Var), and \mathbf{H} is the linearized observation operator. One of the main observations made of forest carbon balance at flux tower sites is the net ecosystem exchange (NEE) of CO_2 which can be estimated by DALEC as the difference between GPP and the respiration's of C_l and C_s giving,

$$NEE(t) = (1 - p_2)GPP(C_f(t), \phi) + p_8C_lT(t) + p_9C_sT(t).$$

For a single observation of NEE at one time, t_0 , an analytical expression for the SIC can be derived using,

$$\mathbf{H}_0 = ((1 - p_2)\zeta_0 \quad 0 \quad 0 \quad p_8 T_0 \quad p_9 T_0),$$

where $\zeta_0 = GPP'(C_f(t_0), \phi)$, $T_0 = T(t_0)$ and \mathbf{H}_0 is the linearized observation operator at time t_0 . As we have a single observation at one time our observation error covariance matrix, \mathbf{R} , is just the variance of our observation of NEE, $\sigma_{nee,0}^2$, at time t_0 . So that,

$$\mathbf{R} = \sigma_{nee.0}^2$$

and

$$\mathbf{J}'' \ = \ \mathbf{B}^{-1} + \mathbf{H}^T \mathbf{R}^{-1} \mathbf{H}$$

$$= \ \begin{pmatrix} \sigma_{cf,b}^{-2} + \sigma_{nee,0}^{-2} (1 - p_2)^2 \zeta_0^2 & 0 & 0 & \sigma_{nee,0}^{-2} (1 - p_2) \zeta_0 p_8 T_0 & \sigma_{nee,0}^{-2} (1 - p_2) \zeta_0 p_9 T_0 \\ 0 & \sigma_{cr,b}^{-2} & 0 & 0 & 0 \\ 0 & 0 & \sigma_{cw,b}^{-2} & 0 & 0 \\ \sigma_{nee,0}^{-2} (1 - p_2) \zeta_0 p_8 T_0 & 0 & 0 & \sigma_{cl,b}^{-2} + \sigma_{nee,0}^{-2} p_8^2 T_0^2 & \sigma_{nee,0}^{-2} p_8 p_9 T_0^2 \\ \sigma_{nee,0}^{-2} (1 - p_2) \zeta_0 p_9 T_0 & 0 & 0 & \sigma_{nee,0}^{-2} p_8 p_9 T_0^2 & \sigma_{cs,b}^{-2} + \sigma_{nee,0}^{-2} p_9^2 T_0^2 \end{pmatrix} .$$

We then have,

$$SIC = \frac{1}{2}ln\frac{|\mathbf{B}|}{|\mathbf{A}|} = \frac{1}{2}ln|\mathbf{B}||\mathbf{J''}|.$$

So that,

$$SIC = \frac{1}{2} ln \frac{(p_2 - 1)^2 \zeta_0^2 \sigma_{cf,b}^2 + \sigma_{nee,0}^2 + T_0^2 (p_9^2 \sigma_{cs,b}^2 + p_8^2 \sigma_{cl,b}^2)}{\sigma_{nee,0}^2}.$$

If we assume that the variances and parameters here are fixed we can see that the size of the SIC is dependent on the temperature term, T_0 , and the square of the first derivative of GPP, ζ_0^2 . Generally the value of GPP (and its first derivative) is highest in summer with higher total daily irradiance and higher temperatures. We therefore have that there will be more information content in observations that are taken when temperatures are higher. I have also derived analytical forms for the SIC using different sets of observations at a single time which all have a similar form.

Following this work based at a single time I started looking at the SIC when successive observations are added over a period of time. The model was now built into a Four-Dimensional Variational Data Assimilation (4D-Var) framework where our observation operator, \mathbf{H} , and observation error covariance matrix, \mathbf{R} , are now,

$$\mathbf{H} = \begin{pmatrix} \mathbf{H}_0 \\ \mathbf{H}_1 \mathbf{M}_0 \\ \vdots \\ \mathbf{H}_n \mathbf{M}_{n,0} \end{pmatrix} \quad \text{and} \quad \mathbf{R} = \begin{pmatrix} \mathbf{R}_0 & 0 & 0 & 0 \\ 0 & \mathbf{R}_1 & 0 & 0 \\ 0 & 0 & \ddots & 0 \\ 0 & 0 & 0 & \mathbf{R}_n \end{pmatrix}.$$

Where \mathbf{H}_i is our linearized observation operator at time t_i , $\mathbf{M}_{i,0} = \mathbf{M}_{i-1}\mathbf{M}_{i-2}\cdots\mathbf{M}_0$ is our linearized model evolving the state vector, \underline{x}_b , at time t_0 to time t_i and \mathbf{R}_i is the observation error covariance matrix corresponding to \mathbf{H}_i at time t_i [3]. I first calculated the adjoint model for DALEC analytically as $\mathbf{M}_i = \frac{\delta m_i}{\delta x_i}$. I then wrote a code in Python that calculates \mathbf{H} and \mathbf{R} to find a value of SIC when successive observations of NEE are made each day over a chosen period (To calculate the SIC we do not need the actual observation value). The meteorological driving data used in the model is taken from a ponderosa pine forest in central Oregon for which the DALEC model in [6] is parameterized. Below we have a plot when different starting points and periods are chosen, day 0 represents the start of the year so that day 200 is during the summer,

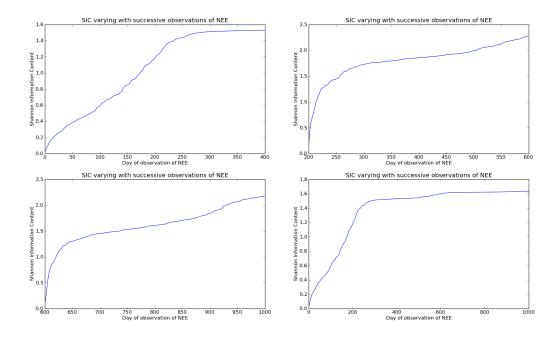


Figure 2: SIC varying as successive observations of NEE are added using driving data from Oregon pine forest.

References

- [1] Dennis Baldocchi. Turner review no. 15. 'breathing' of the terrestrial biosphere: lessons learned from a global network of carbon dioxide flux measurement systems. Australian Journal of Botany, 56(1):1–26, 2008.
- [2] Sylvain Delahaies, Ian Roulstone, and Nancy Nichols. A regularization of the carbon cycle data-fusion problem. In *EGU General Assembly Conference Abstracts*, volume 15, page 4087, 2013.
- [3] John M Lewis, Sivaramakrishnan Lakshmivarahan, and Sudarshan Dhall. *Dynamic data as-similation: a least squares approach*, volume 13. Cambridge University Press, 2006.
- [4] Clive D Rodgers et al. *Inverse methods for atmospheric sounding: Theory and practice*, volume 2. World scientific Singapore, 2000.
- [5] Laura M Stewart, SL Dance, and NK Nichols. Correlated observation errors in data assimilation. *International journal for numerical methods in fluids*, 56(8):1521–1527, 2008.
- [6] Mathew Williams, Paul A Schwarz, Beverly E Law, James Irvine, and Meredith R Kurpius. An improved analysis of forest carbon dynamics using data assimilation. *Global Change Biology*, 11(1):89–105, 2005.