

Information content for observations of forest carbon stocks and fluxes when assimilated with the DALEC carbon balance model

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1 Introduction

A large amount of data is currently being gathered that is relevant to the carbon balance of forests, with much of this data coming from Eddy covariance flux towers [1, 2, 10, 13]. Attempts are also being made to combine this data with models of forest carbon stocks and fluxes, such as the Data Assimilation Linked Ecosystem Carbon model (DALEC), in a data assimilation scheme [7, 14]. Currently, however, there are limitations with such schemes as there is a lack of understanding about the additional information provided by different observations. Better understanding of the information content of carbon balance observations will help inform measurement campaigns of when and which observations to take in order to gain the most possible information about the system. In this report we will look at different information content measures which have been used in meteorological data assimilation [5, 9, 11] and apply these to carbon balance observations assimilated with DALEC. Although we use DALEC in this report the results will be similar for other carbon balance models which use similar driving data and equations. We begin by introducing the DALEC model which will initially be used to look at the information content in different observations.

2 The DALEC Model

The DALEC model is a simple process-based model describing the carbon balance of an evergreen forest ecosystem [14]. The model is constructed of five carbon pools (foliage (C_f), fine roots (C_r), woody stems and coarse roots (C_w), fresh leaf and fine root litter (C_l) and soil organic matter and coarse woody debris (C_s)) linked via fluxes. The gross primary production function (GPP) uses meteorological driving data and the site's leaf area index (a function of C_f) to calculate the total amount of carbon to be allocated at a daily time step.

The model equations for the carbon pools at day $t + 1$ are as follows:

$$C_f(t + 1) = (1 - p_5)C_f(t) + p_3(1 - p_2)GPP(C_f(t), \phi), \quad (1)$$

$$C_r(t + 1) = (1 - p_7)C_r(t) + p_4(1 - p_3)(1 - p_2)GPP(C_f(t), \phi), \quad (2)$$

$$C_w(t + 1) = (1 - p_6)C_w(t) + (1 - p_4)(1 - p_3)(1 - p_2)GPP(C_f(t), \phi), \quad (3)$$

$$C_l(t + 1) = (1 - (p_1 + p_8)T(t))C_l(t) + p_5C_f(t) + p_7C_r(t), \quad (4)$$

$$C_s(t + 1) = (1 - p_9T(t))C_s + p_6C_w(t) + p_1T(t)C_l(t), \quad (5)$$

where $T(t) = \frac{1}{2}exp(p_{10}T_m(t))$, T_m is daily mean temperature, p_1, \dots, p_{10} are rate parameters and ϕ represents the meteorological driving data used in the GPP function. The full details

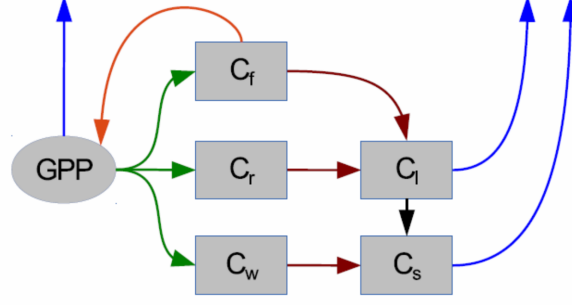


Figure 1: Representation of the carbon fluxes in the DALEC carbon balance model. Green arrows represent C allocation, dark red and black arrows represent litterfall and decomposition fluxes, blue arrows represent respiration fluxes and the light red arrow represents the feedback of foliar carbon to the *GPP* function. [3]

of this version of DALEC can be found in [14], it is parameterized for data from a young pine stand in Ponderosa, Oregon, we include values of the model parameters and the equations used to calculate *GPP* in the appendix. We now see how DALEC can be implemented in a four-dimensional variational data assimilation (4D-Var) framework.

2.1 DALEC in a variational assimilation framework

In 4D-Var we aim to maximise the probability of our initial state \mathbf{x}_0 (for DALEC our state \mathbf{x}_0 corresponds to the initial values of the five carbon pools, $\mathbf{x}_0 = (C_f(t_0), C_r(t_0), C_w(t_0), C_l(t_0), C_s(t_0))^T$) given a set of observations \mathbf{y} , $P(\mathbf{x}_0|\mathbf{y})$, over some time window, N . We do this by minimising a cost function $J(\mathbf{x})$ derived from Baye's Theorem [8]. The cost function is given as,

$$J(\mathbf{x}_0) = \frac{1}{2}(\mathbf{x}_0 - \mathbf{x}_b)^T \mathbf{B}^{-1}(\mathbf{x}_0 - \mathbf{x}_b) + \frac{1}{2} \sum_{i=0}^N (\mathbf{y}_i - h_i(\mathbf{x}_i))^T \mathbf{R}_i^{-1}(\mathbf{y}_i - h_i(\mathbf{x}_i)), \quad (6)$$

where \mathbf{x}_b is our background and acts as our initial guess to our state \mathbf{x}_0 , \mathbf{B} is the background error covariance matrix and quantifies our knowledge of the error in our background, h_i is our observation operator at time t_i and maps our state vector evolved by our nonlinear model (\mathbf{x}_i) to the observations at this time \mathbf{y}_i and \mathbf{R}_i is the observation error covariance matrix and represents our knowledge of the uncertainty in the observations. The state that minimises the cost function is called the analysis and is denoted as \mathbf{x}_a , this state is found using a minimisation routine that takes the cost function, our initial guess (\mathbf{x}_b) and also the gradient of the cost function defined as,

$$\nabla J(\mathbf{x}_0) = \mathbf{B}^{-1}(\mathbf{x}_0 - \mathbf{x}_b) - \sum_{i=0}^N \mathbf{M}_{i,0}^T \mathbf{H}_i^T \mathbf{R}_i^{-1}(\mathbf{y}_i - h_i(\mathbf{x}_i)), \quad (7)$$

where $\mathbf{H}_i = \frac{\delta h_i(\mathbf{x}_i)}{\delta \mathbf{x}_i}$ is our linearized observation operator, $\mathbf{M}_{i,0} = \mathbf{M}_{i-1} \mathbf{M}_{i-2} \cdots \mathbf{M}_0$ is our tangent linear model with $\mathbf{M}_i = \frac{\delta \mathbf{m}_i(\mathbf{x}_i)}{\delta \mathbf{x}_i}$. We can calculate the linearized model for DALEC from equations

1 to 5 as,

$$\mathbf{M}_i = \begin{pmatrix} (1-p_5) + p_3(1-p_2)\zeta_i & 0 & 0 & 0 & 0 \\ p_4(1-p_3)(1-p_2)\zeta_i & (1-p_7) & 0 & 0 & 0 \\ (1-p_4)(1-p_3)(1-p_2)\zeta_i & 0 & (1-p_6) & 0 & 0 \\ p_5 & p_7 & 0 & (1-(p_1+p_8)T_i) & 0 \\ 0 & 0 & p_6 & p_1T_i & (1-p_9T_i) \end{pmatrix}, \quad (8)$$

where $\zeta_i = GPP'(C_f(t_i), \phi)$ and $T_i = T(t_i)$.

Once we have performed the minimisation of the cost function and determined \mathbf{x}_a we can also calculate the analysis error covariance matrix, \mathbf{A} , to quantify the uncertainty in our new estimate of the state. We can define the analysis error covariance matrix as,

$$\mathbf{A} = (\mathbf{J}'')^{-1} = (\mathbf{B}^{-1} + \hat{\mathbf{H}}^T \hat{\mathbf{R}}^{-1} \hat{\mathbf{H}})^{-1}, \quad (9)$$

where $\hat{\mathbf{H}}$ is the matrix of linearized observation operators evolved by the tangent linear model and $\hat{\mathbf{R}}$ is the block diagonal matrix of observation error covariance matrices,

$$\hat{\mathbf{H}} = \begin{pmatrix} \mathbf{H}_0 \\ \mathbf{H}_1 \mathbf{M}_0 \\ \vdots \\ \mathbf{H}_N \mathbf{M}_{N,0} \end{pmatrix} \quad \text{and} \quad \hat{\mathbf{R}} = \begin{pmatrix} \mathbf{R}_0 & 0 & 0 & 0 \\ 0 & \mathbf{R}_1 & 0 & 0 \\ 0 & 0 & \ddots & 0 \\ 0 & 0 & 0 & \mathbf{R}_N \end{pmatrix}. \quad (10)$$

It is important to note that we do not need any actual observations to calculate \mathbf{A} , \mathbf{J}'' or \mathbf{B} only the values of the standard deviations for the observations and the stand deviations for our background values.

3 Information Content Measures

Information content measures are already being used to quantify the different levels of information provided by observations in the development of satellite instruments [4, 12] and in operational data assimilation schemes [5, 11]. In these fields two of the more widely used measures are Shannon Information Content (also known as entropy reduction) and the degrees of freedom for signal. We will apply both methods for observations assimilated with DALEC.

3.1 Shannon Information Content

In DA Shannon Information Content (*SIC*) is a measure of the reduction in entropy given a set of observations. Entropy physically corresponds to the volume in state space taken up by the probability density function (*pdf*) describing the knowledge of the state. When a measurement is made the volume of this *pdf* decreases, the information content of the measurement is a measure of the factor by which it decreases [9]. If $P_b(x)$ is our knowledge of the state before an observation and $P_o(x|y)$ is our knowledge after an observation then we have entropies,

$$S[P_b(x)] = - \int P_b(x) \log_2[P_b(x)] dx \quad \text{and} \quad S[P_o(x|y)] = - \int P_o(x|y) \log_2[P_o(x|y)] dx.$$

The entropy reduction, or *SIC*, due to the observation is then,

$$SIC = S[P_b(x)] - S[P_o(x|y)]. \quad (11)$$

If we assume all *pdfs* are Gaussian and use the natural logarithm as opposed to \log_2 (for algebraic convenience) [9] the entropy of a multivariate Gaussian distribution for a vector \mathbf{x} with n elements (before and after observations) can be derived as,

$$S[P_b(\mathbf{x})] = n \ln(2\pi e)^{\frac{1}{2}} + \frac{1}{2} \ln |\mathbf{B}| \quad (12)$$

and

$$S[P_o(\mathbf{x}|\mathbf{y})] = n \ln(2\pi e)^{\frac{1}{2}} + \frac{1}{2} \ln |\mathbf{A}| \quad (13)$$

where \mathbf{B} is the background error covariance matrix and \mathbf{A} is the analysis error covariance matrix. Combining equations 11, 12 and 13 we can write the *SIC* as,

$$SIC = \frac{1}{2} \ln \frac{|\mathbf{B}|}{|\mathbf{A}|}. \quad (14)$$

From equation 9 we can see the *SIC* can also be written as,

$$SIC = \frac{1}{2} \ln |\mathbf{B}| |\mathbf{J}''|. \quad (15)$$

3.2 Degrees of Freedom for Signal

The degrees of freedom for signal (*DFS*) indicates the number of elements of the state that have been measured by the observations. If we consider a state vector \mathbf{x} with n elements (or n degrees of freedom) then the maximum value the *DFS* could obtain would be n , in this case all elements of the state would have been measured. Conversely if $DFS = 0$ then no elements of the state would have been measured by our observations [6].

We have symmetric positive definite background and analysis error covariance matrices \mathbf{B} and \mathbf{A} , the eigenvalues of each matrix gives a representation for the uncertainty in the direction of the associated eigenvector, thus by comparing the eigenvalues of both matrices we can determine the reduction in uncertainty given a set of observations [12].

In order to do this we take $\mathbf{B}^{-\frac{1}{2}}$ such that $\mathbf{B}^{-1} = \mathbf{B}^{-\frac{1}{2}} \mathbf{B}^{-\frac{1}{2}}$. We now take \mathbf{Q} to be the orthogonal matrix composed of the eigenvectors of $\mathbf{B}^{-\frac{1}{2}} \mathbf{A} \mathbf{B}^{-\frac{1}{2}}$ we have,

$$\mathbf{Q}^T \left(\mathbf{B}^{-\frac{1}{2}} \mathbf{A} \mathbf{B}^{-\frac{1}{2}} \right) \mathbf{Q} = \mathbf{\Lambda}, \quad (16)$$

$$\mathbf{Q}^T \left(\mathbf{B}^{-\frac{1}{2}} \mathbf{B} \mathbf{B}^{-\frac{1}{2}} \right) \mathbf{Q} = \mathbf{I}_{n \times n} \quad (17)$$

where $\mathbf{\Lambda}$ is a diagonal matrix. Each diagonal element of our transformed \mathbf{B} is equal to one and corresponds to one degree of freedom. The diagonal elements of $\mathbf{\Lambda}$ correspond to the matrixs eigenvalues and can be interpreted as the relative reduction in variance for each of the n degrees of freedom [11]. We can then define the *DFS* as,

$$\begin{aligned} DFS &= \text{trace}(\mathbf{I}_{n \times n} - \mathbf{\Lambda}) \\ &= n - \text{trace}(\mathbf{\Lambda}) \\ &= n - \text{trace}(\mathbf{B}^{-\frac{1}{2}} \mathbf{A} \mathbf{B}^{-\frac{1}{2}}) \\ &= n - \text{trace}(\mathbf{B}^{-1} \mathbf{A}). \end{aligned} \quad (18)$$

4 Shannon Information Content for DALEC

We begin by using *SIC* to understand the information content for different observations when being assimilated with the DALEC model. For these experiments the model is set up as in section 2.1 where the elements of the state vector have standard deviations, $\sigma_{cf,b}, \dots, \sigma_{cs,b}$, respectively. These standard deviations represent the uncertainty in our initial background estimate and are taken as a percentage of the initial carbon pools (these values can be found in the appendix). The background error covariance matrix is then taken as the diagonal matrix of the variances of the carbon pools,

$$\mathbf{B} = \begin{pmatrix} \sigma_{cf,b}^2 & 0 & 0 & 0 & 0 \\ 0 & \sigma_{cr,b}^2 & 0 & 0 & 0 \\ 0 & 0 & \sigma_{cw,b}^2 & 0 & 0 \\ 0 & 0 & 0 & \sigma_{cl,b}^2 & 0 \\ 0 & 0 & 0 & 0 & \sigma_{cs,b}^2 \end{pmatrix}. \quad (19)$$

Our initial experiments look at the *SIC* in observations taken at single time, the 4D-Var data assimilation then becomes three-dimensional variational assimilation (3D-Var) as we are not summing over a time window and have no time component.

4.1 *SIC* for a single observation at one time

If we first consider one observation of C_f (the first element of our state vector \mathbf{x}) at time t_0 , we can derive an analytical expression for the *SIC* using,

$$\mathbf{H}_0 = \frac{\delta C_f(t_0)}{\delta \mathbf{x}_0} = \begin{pmatrix} 1 & 0 & 0 & 0 & 0 \end{pmatrix}, \quad (20)$$

where \mathbf{H}_0 is our linearized observation operator at time t_0 . As we have a single observation at one time our observation error covariance matrix, \mathbf{R}_0 , is just the variance of our observation of C_f at time t_0 ($\sigma_{cf,o}^2$). Therefore,

$$\mathbf{R}_0 = \sigma_{cf,o}^2. \quad (21)$$

We then have from equation 9,

$$\begin{aligned} \mathbf{J}'' &= \mathbf{B}^{-1} + \hat{\mathbf{H}}^T \hat{\mathbf{R}}^{-1} \hat{\mathbf{H}} \\ &= \mathbf{B}^{-1} + \mathbf{H}_0^T \mathbf{R}_0^{-1} \mathbf{H}_0 \\ &= \begin{pmatrix} \sigma_{cf,b}^{-2} + \sigma_{cf,o}^{-2} & 0 & 0 & 0 & 0 \\ 0 & \sigma_{cr,b}^{-2} & 0 & 0 & 0 \\ 0 & 0 & \sigma_{cw,b}^{-2} & 0 & 0 \\ 0 & 0 & 0 & \sigma_{cl,b}^{-2} & 0 \\ 0 & 0 & 0 & 0 & \sigma_{cs,b}^{-2} \end{pmatrix}. \end{aligned} \quad (22)$$

We can now derive the *SIC* using equation 15 as,

$$SIC = \frac{1}{2} \ln |\mathbf{B}| |\mathbf{J}''| = \frac{1}{2} \ln \frac{(\sigma_{cf,o}^2 + \sigma_{cf,b}^2)}{\sigma_{cf,o}^2} = \frac{1}{2} \ln \left(1 + \frac{\sigma_{cf,b}^2}{\sigma_{cf,o}^2} \right). \quad (23)$$

We see the *SIC* for an observation of a single observation of C_f is dependant on the ratio between the observation and background variances, the *SIC* will have the same form for all other direct observations of carbon pools contained in the state. As our background standard deviations are set

to being the same percentage for each carbon pool (SEE APPENDIX), the carbon pool observation which will give us the highest SIC is the pool that we can measure most accurately as this will maximise the ratio $\frac{\sigma_{c,b}^2}{\sigma_{c,o}^2}$ by minimising $\sigma_{c,o}^2$.

One of the main carbon balance observations made at forest flux tower sites is the net ecosystem exchange (NEE) of CO_2 , which can be estimated by DALEC as the difference between GPP and the respiration of C_l and C_s , giving,

$$NEE(t) = -(1 - p_2)GPP(C_f(t), \phi) + p_8C_lT(t) + p_9C_sT(t). \quad (24)$$

For a single observation of NEE at one time, t_0 , we can again derive an analytical expression for the SIC using,

$$\mathbf{H}_0 = \frac{\delta NEE(t_0)}{\delta \mathbf{x}_0} = \begin{pmatrix} -(1 - p_2)\zeta_0 & 0 & 0 & p_8T_0 & p_9T_0 \end{pmatrix}, \quad (25)$$

where $\zeta_0 = GPP'(C_f(t_0), \phi)$, $T_0 = T(t_0)$ and \mathbf{H}_0 is the linearized observation operator at time t_0 . Again our observation error covariance matrix, \mathbf{R}_0 , is just the variance of our observation of NEE , $\sigma_{nee,o}^2$, at time t_0 . Therefore,

$$\mathbf{R}_0 = \sigma_{nee,o}^2 \quad (26)$$

and again,

$$\begin{aligned} \mathbf{J}'' &= \mathbf{B}^{-1} + \mathbf{H}_0^T \mathbf{R}_0^{-1} \mathbf{H}_0 \\ &= \begin{pmatrix} \sigma_{cf,b}^{-2} + \sigma_{nee,o}^{-2}(1 - p_2)^2\zeta_0^2 & 0 & 0 & \sigma_{nee,o}^{-2}(1 - p_2)\zeta_0p_8T_0 & \sigma_{nee,o}^{-2}(1 - p_2)\zeta_0p_9T_0 \\ 0 & \sigma_{cr,b}^{-2} & 0 & 0 & 0 \\ 0 & 0 & \sigma_{cw,b}^{-2} & 0 & 0 \\ \sigma_{nee,o}^{-2}(1 - p_2)\zeta_0p_8T_0 & 0 & 0 & \sigma_{cl,b}^{-2} + \sigma_{nee,o}^{-2}p_8^2T_0^2 & \sigma_{nee,o}^{-2}p_8p_9T_0^2 \\ \sigma_{nee,o}^{-2}(1 - p_2)\zeta_0p_9T_0 & 0 & 0 & \sigma_{nee,o}^{-2}p_8p_9T_0^2 & \sigma_{cs,b}^{-2} + \sigma_{nee,o}^{-2}p_9^2T_0^2 \end{pmatrix}. \end{aligned} \quad (27)$$

We then have,

$$SIC = \frac{1}{2} \ln |\mathbf{B}| |\mathbf{J}''| = \frac{1}{2} \ln \frac{(p_2 - 1)^2 \zeta_0^2 \sigma_{cf,b}^2 + \sigma_{nee,o}^2 + T_0^2 (p_9^2 \sigma_{cs,b}^2 + p_8^2 \sigma_{cl,b}^2)}{\sigma_{nee,o}^2}. \quad (28)$$

If we assume that the variances and parameters here are fixed we can see that the size of the SIC is dependent on the temperature term, T_0 , and the square of the first derivative of GPP , ζ_0^2 . Generally, the value of GPP (and its first derivative) is highest in summer with higher total daily irradiance and higher temperatures. We therefore have that there will be more information content in observations that are taken when temperatures are higher. Physically this makes sense as more NEE takes place when temperatures are higher (to a point) so measurements are of greater magnitude and give us more information about carbon fluxes. **We have also conducted further experiments with different sets of observations at one time and over a time window (should I include these?).**

By plotting the SIC for a single observation of NEE varying with three years of meteorological driving data and the temperature term ($T(t_i)$) for the same period of the same data we can see that both are closely linked in figure 2. However the relationship is not linear as the magnitude of GPP 's first derivative is also dependent on daily irradiance and the value of the foliar carbon pool (C_f). This shows that observations of NEE made in the summer are much more valuable than those made in the winter assuming warmer temperatures, higher daily irradiance and a higher amount of foliar carbon in the summer. In the next section we consider a series of observations over some time window. We will see that it takes up to 30 observations of NEE in winter to gain the same level of information as 1 observation of NEE in the winter.

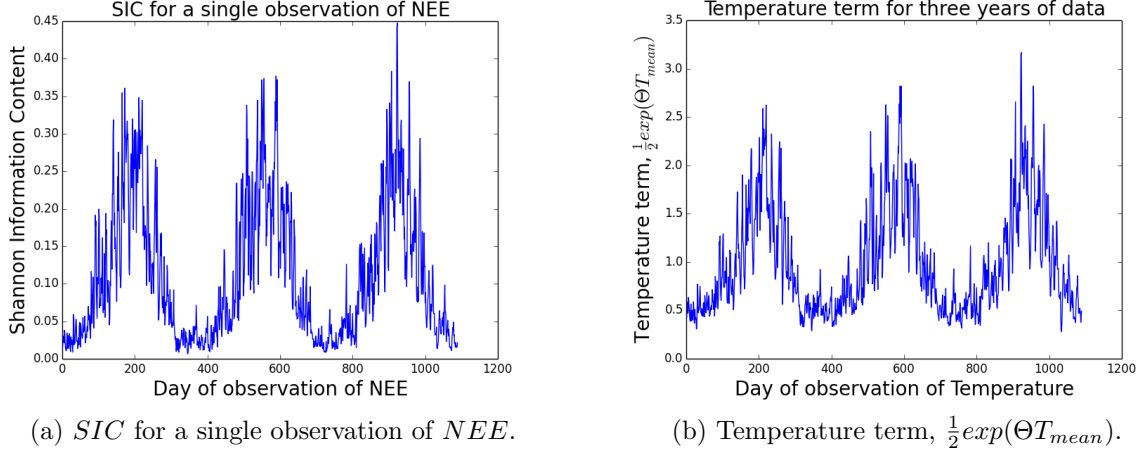


Figure 2: SIC and temperature varying over three years using driving data from Oregon pine forest.

4.2 SIC for successive observations over a time window

Following the results for SIC based at a single time, we now consider the SIC when successive observations are added over a period of time. We begin by considering two successive observations of C_f in time. From section 2.1 and 4.1 we have our linearized observation operator and observation error covariance matrix at time t_i ,

$$\mathbf{H}_i = \frac{\delta C_f(t_i)}{\delta \mathbf{x}_i} = \begin{pmatrix} 1 & 0 & 0 & 0 & 0 \end{pmatrix} \quad \text{and} \quad \mathbf{R}_i = \sigma_{cf,o}^2. \quad (29)$$

Then for two successive observations of C_f we have,

$$\hat{\mathbf{H}} = \begin{pmatrix} \mathbf{H}_0 \\ \mathbf{H}_1 \mathbf{M}_0 \end{pmatrix} = \begin{pmatrix} 1 & 0 & 0 & 0 & 0 \\ (1 - p_5) + p_3(1 - p_2)GPP'(C_f(t_0), \phi) & 0 & 0 & 0 & 0 \end{pmatrix} \quad (30)$$

and

$$\hat{\mathbf{R}} = \begin{pmatrix} \mathbf{R}_0 & 0 \\ 0 & \mathbf{R}_1 \end{pmatrix} = \begin{pmatrix} \sigma_{cf,o}^2 & 0 \\ 0 & \sigma_{cf,o}^2 \end{pmatrix}. \quad (31)$$

We then have,

$$SIC = \frac{1}{2} \ln |\mathbf{B}| |\mathbf{J}''| = \frac{1}{2} \ln \left(1 + \frac{\sigma_{cf,b}^2}{\sigma_{cf,o}^2} + \frac{\sigma_{cf,b}^2 \eta_0^2}{\sigma_{cf,o}^2} \right), \quad (32)$$

where $\eta_i = (1 - p_5) + p_3(1 - p_2)GPP'(C_f(t_i), \phi)$. We see this is similar to equation 23 but with an extra term, η_0^2 , which is the square of the first element of our linearized model \mathbf{M}_0 . We can continue adding more observations at successive times and we start to see a pattern. For three observations at successive times we have,

$$SIC = \frac{1}{2} \ln \left(1 + \frac{\sigma_{cf,b}^2}{\sigma_{cf,o}^2} + \frac{\sigma_{cf,b}^2 \eta_0^2}{\sigma_{cf,o}^2} + \frac{\sigma_{cf,b}^2 \eta_0^2 \eta_1^2}{\sigma_{cf,o}^2} \right), \quad (33)$$

for four,

$$SIC = \frac{1}{2} \ln \left(1 + \frac{\sigma_{cf,b}^2}{\sigma_{cf,o}^2} + \frac{\sigma_{cf,b}^2 \eta_0^2}{\sigma_{cf,o}^2} + \frac{\sigma_{cf,b}^2 \eta_0^2 \eta_1^2}{\sigma_{cf,o}^2} + \frac{\sigma_{cf,b}^2 \eta_0^2 \eta_1^2 \eta_2^2}{\sigma_{cf,o}^2} \right). \quad (34)$$

Using a simple proof by induction we find that for n observations we have,

$$SIC \text{ for } n \text{ observations of } C_f = \frac{1}{2} \ln \left(1 + \frac{\sigma_{cf,b}^2}{\sigma_{cf,o}^2} \left(1 + \sum_{k=0}^{n-2} \prod_{i=0}^k \eta_i^2 \right) \right). \quad (35)$$

We have plotted the SIC for increasing numbers of observations of C_f , using three years of meteorological driving data from the Oregon pine forest, as seen in figure 3.

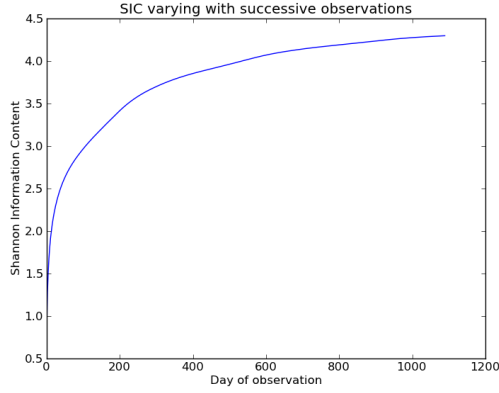


Figure 3: SIC varying as successive observations of C_f are added using driving data from Oregon pine forest.

We can see after around 250 observations of C_f have been made our SIC starts to level off as we are getting a smaller and smaller reduction in entropy for the extra observations of C_f being made. We this effect when making successive observations of any of the carbon pools, eventually we cannot reduce the entropy of the system anymore by just adding more of the same observation.

As before with a single observation at one time we can repeat this with successive observations of NEE instead of C_f this is plotted in figure 4. Here we see a similar levelling off in SIC as figure 3, we can also see the seasonal cycle of information content in NEE as in figure 2. Once the SIC has levelled off no information is added during the winter months and the SIC is only increased with observations made during the summer.

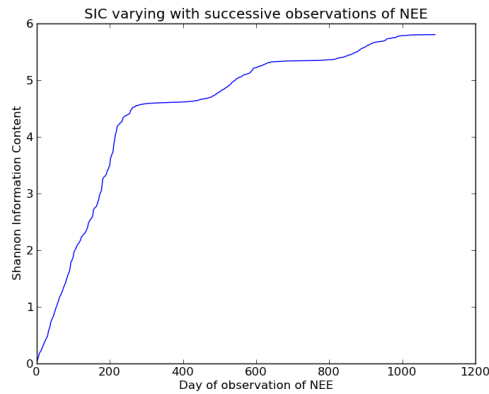


Figure 4: SIC varying as successive observations of NEE are added using driving data from Oregon pine forest.

We can also repeat this with successive observations of both NEE and C_f this can be seen in figure 5. We attain a higher value of SIC more quickly than before but ultimately reach a similar value of SIC as in figure 4 and reduce the entropy of our system very little when adding extra observations of C_f and NEE .

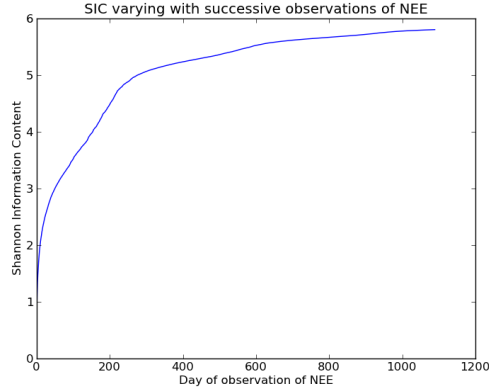


Figure 5: SIC varying as successive observations of NEE and C_f are added using driving data from Oregon pine forest.

In order to see the difference in information content between NEE observations made in the summer and those made in the winter in figure 6 we have plotted the increasing SIC as 50 observations are made from December 21st onwards and also the the constant line of SIC for one observation of NEE made in summer at a mean daily temperature of 26°C on July 12th.

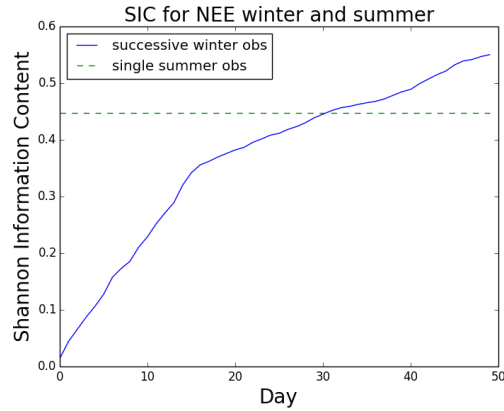


Figure 6: SIC varying as successive observations of NEE are added during winter using driving data from Oregon pine forest. The dashed line shows the SIC for a single observation of NEE made in summer.

Figure 6 shows us that for the driver data used in this experiment we require 31 days of NEE observations to gain the same reduction in entropy for our analysis as 1 observation made in the summer.

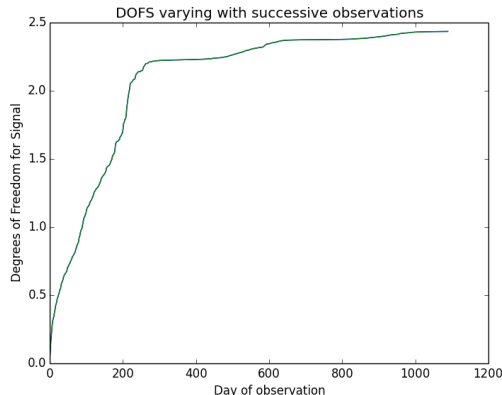


Figure 7: *DOFS* varying as successive observations of *NEE* are added using driving data from Oregon pine forest.

5 Degrees of freedom for signal for DALEC

6 Conclusion

We have seen that after continually adding the same observation the information we get from the observation starts to get smaller and eventually we can reduce the entropy of our system no more with more of the same observations. It would be better to take a variety of observations if possible. *NEE* observations made in summer are much more valuable than those made in winter. Although we have used DALEC for these experiments *NEE* is a product of *GPP* and the respirations and the *GPP* function in DALEC is essentially the same as such functions in other carbon models so we would expect to see very similar results. etc etc

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