

1 **Tropopause Evolution in a Rapidly Intensifying Tropical Cyclone: A Static**
2 **Stability Budget Analysis**

3 Patrick Duran* and John Molinari

4 *University at Albany, State University of New York, Albany, NY*

5 **Corresponding author address:* Department of Atmospheric and Environmental Sciences, Univer-
6 sity at Albany, State University of New York, 1400 Washington Avenue, Albany, NY.

7 E-mail: pduran2008@gmail.com

ABSTRACT

⁸ We have some cool results!

9 **1. Introduction**

10 Perhaps introduce upper-tropospheric static stability and its relationship to the diurnal cycle
11 before going into Patricia? Include references to Dunion, Navarro, and O'Neill here.

12 After undergoing a remarkably rapid intensification (RI), Hurricane Patricia (2015) set a new
13 record as the strongest tropical cyclone (TC) ever observed in the Western Hemisphere (Kim-
14 berlain et al. 2016; Rogers et al. 2017). High-altitude dropsonde observations taken during
15 the Tropical Cyclone Intensity (TCI) Experiment captured this RI in unprecedented detail (Doyle
16 et al. 2017). These observations revealed remarkable changes in the structure of the cold-point
17 tropopause and upper-level static stability as the storm intensified (Duran and Molinari 2018). At
18 tropical storm intensity, shortly before RI commenced, a strong inversion layer existed just above
19 Patricia's cold-point tropopause, which was located near 17.2 km. During the first half of the RI
20 period, this inversion layer weakened throughout Patricia's inner core, with the weakening most
21 pronounced over the developing eye. By the time the storm reached its maximum intensity, the
22 inversion layer over the eye had disappeared almost completely, which was accompanied by an
23 increase in the tropopause height to a level at or above the highest-available dropsonde data point
24 (18.3 km) at two locations. Meanwhile over the eyewall region, the static stability re-strengthened
25 and the tropopause was limited to a level at or below 17.5 km. The mechanisms that led to these
26 changes in upper-level static stability and tropopause height are the subject of the current paper.

27 More recently, Dunion et al. (2014) documented a periodic oscillation of infrared brightness
28 temperature in hurricanes, which they call the "TC diurnal pulse." There will be a whole bunch of
29 papers cited here...

30 At some point (probably in the Discussion) mention the possible importance of static stability
31 asymmetries, in the context of the Dunion diurnal pulse

32 2. Model Setup

33 The numerical simulations were performed using version 19.4 of Cloud Model 1 (CM1) de-
34 scribed in Bryan and Rotunno (2009). The equations of motion were integrated on a 3000-km-
35 wide, 30-km-deep axisymmetric grid with 1-km horizontal and 250-m vertical grid spacing. The
36 computations were performed on an f -plane at 15°N latitude, over a sea surface with constant
37 temperature of 30.5°C, which matches that observed near Hurricane Patricia (2015; Kimberlain
38 et al. 2016). Horizontal turbulence was parameterized using the Smagorinsky scheme described in
39 Bryan and Rotunno (2009, pg. 1773), with a prescribed mixing length that varied linearly from 100
40 m at a surface pressure of 1015 hPa to 1000 m at a surface pressure of 900 hPa. This formulation
41 allows for realistically-large horizontal mixing lengths near the hurricane’s inner core, consistent
42 with the results of Bryan (2012), while not over-representing horizontal turbulence in convection at
43 outer radii. Vertical turbulence was parameterized using the formulation of Markowski and Bryan
44 (2016, their Eq. 6), using an asymptotic vertical mixing length of 100 m. A Rayleigh damping
45 layer was applied outside of the 2900-km radius and above the 25-km level to prevent spurious
46 gravity wave reflection at the model boundaries. Microphysical processes were parameterized
47 using the Thompson et al. (2004) microphysics scheme and radiative heating tendencies were
48 computed every two minutes using the Rapid Radiative Transfer Model for GCMs? (RRTMG)
49 longwave and shortwave schemes (Iacono et al. 2008). A horizontally-homogeneous temperature
50 and humidity field was initialized with a mean sounding computed using all dropsondes deployed
51 during the TCI flight conducted within and around Tropical Storm Patricia on 21 October, 2015
52 (see Doyle et al. 2017 for details.) Above 19 km, where few TCI observations were available,
53 the temperature profile was taken from the Climate Forecast System Reanalysis (CFSR) grid point
54 nearest Patricia’s storm center, valid at 18 UTC 21 October, 2015. Since relative humidity mea-

surements were unreliable at temperatures below -40°C (Bell et al. 2016), relative humidity was set equal to 50% above 11.5 km (the level above which temperature dropped below -40°C). The vortex described in Rotunno and Emanuel (1987, their Eq. 37) was used to initialize the wind field, setting all parameters equal to the values used therein.

3. Budget Computation

The static stability can be expressed as the squared Brunt Väisälä frequency:

$$N_m^2 = \frac{g}{T} \left(\frac{\partial T}{\partial z} + \Gamma_m \right) \left(1 + \frac{T}{R_d/R_v + q_s} \frac{\partial q_s}{\partial T} \right) - \frac{g}{1 + q_t} \frac{\partial q_t}{\partial z}, \quad (1)$$

where g is gravitational acceleration, T is temperature, R_d and R_v are the gas constants of dry air and water vapor, respectively, q_s is the saturation mixing ratio, q_t is the total condensate mixing ratio, and Γ_m is the moist-adiabatic lapse rate:

$$\Gamma_m = g(1 + q_t) \left(\frac{1 + L_v q_s / R_d T}{c_{pm} + L_v \partial q_s / \partial T} \right) \quad (2)$$

In the tropopause layer, q_s , $\partial q_s / \partial T$, and $\partial q_t / \partial z$ approach zero. In this limiting case, Eq. 1 reduces to:

$$N^2 = \frac{g}{\theta_v} \frac{\partial \theta_v}{\partial z}, \quad (3)$$

where θ_v is the virtual potential temperature. To compute N^2 , CM1 uses Eq. 1 in saturated environments and Eq. 3 in sub-saturated environments, but for simplicity only Eq. 3 will be used for the budget computations herein¹.

Taking the time derivative of Eq. 3 yields the static stability tendency:

$$\frac{\partial N^2}{\partial t} = \frac{g}{\theta} \frac{\partial}{\partial z} \frac{\partial \theta}{\partial t} - \frac{g}{\theta^2} \frac{\partial \theta}{\partial z} \frac{\partial \theta}{\partial t}, \quad (4)$$

where the potential temperature tendency, $\partial \theta / \partial z$, can be written:

$$\frac{\partial \theta}{\partial t} = HADV + VADV + HTURB + VTURB + MP + RAD + DISS \quad (5)$$

¹The validity of this approximation will be substantiated later in this section.

Each term on the right-hand side of Eq. 5 represents a potential temperature budget variable, each of which is output directly by the model every minute. HADV and VADV are the radial and vertical advective tendencies, HTURB and VTURB are the radial and vertical tendencies from the turbulence parameterization, MP is the tendency from the microphysics scheme, RAD is the tendency from the radiation scheme, and DISS is the tendency due to turbulent dissipation. This equation neglects Rayleigh damping, since this term is zero everywhere below 25 km, and the analysis domain does not extend to that level. Each term in Eq. 5 is substituted for $\partial\theta/\partial t$ in Eq. 4, yielding the contribution of each budget term to the static stability tendency. These terms are summed, yielding an instantaneous "budget change" in N^2 every minute. The budget changes are then averaged over 24-hour periods and compared to the total model change in N^2 over that same time period using a residual, i.e.:

$$\Delta N_{budget}^2 = \sum_{t=t_0}^{t_0+\delta t} \frac{\partial N^2}{\partial t} |_{t, NEEDBAROVERSUMMATION} \quad (6)$$

$$\Delta N_{model}^2 = N_{t_0+\delta t}^2 - N_{t_0}^2, \quad (7)$$

$$Residual = \Delta N_{model}^2 - \Delta N_{budget}^2 \quad (8)$$

where t_0 is an initial time and δt is 24 hours. In Eq. 6, Eq. 1 in saturated environments and Eq. 3 in sub-saturated environments; t_0 is an initial time and δt is 24 hours.

CONTINUATION OF THIS SECTION IS IN MY RESEARCH NOTEBOOK

4. Results

Put description of Fig. 1 in this section.

Don't forget to mention 1-2-1 smoother.

90 *a. Static stability evolution*

91 The average N^2 over the first day of the simulation (Fig. 2a) indicates the presence of a static
92 stability maximum about 400 m above the cold-point tropopause. This lower-stratospheric stable
93 layer had begun to erode during the initial spin-up period, with the maximum destabilization
94 occurring at the innermost radii. This decrease in static stability continued into the second day
95 of the simulation (Fig. 2b) as the storm intensified to hurricane strength (Fig. 1). Destabilization
96 was particularly pronounced over the developing eye, where the time-mean cold-point tropopause
97 height increased by up to 400 m compared to the previous day. Over the developing eyewall
98 and outer rainband regions, meanwhile, the tropopause height remained nearly constant. During
99 the third day of the simulation (Fig. 2c), static stability over the eye continued to decrease, and
100 the cold-point tropopause height rose to 18.3 km at the storm center. The tropopause sloped
101 sharply downward over the innermost radii, reaching the 16.4-km level near the 50-km radius. This
102 local minimum in tropopause height corresponded to the eyewall region, where upper-tropospheric
103 static stability increased during this time period. Outside of the eyewall region, static stability
104 began to increase in the layer immediately overlying the cold-point tropopause. This stable layer
105 sloped upward with radius, which corresponded to an upward-sloping tropopause radially outside
106 of the eyewall region. Over the next 24 hours (Fig. 2d), as the storm's maximum 10-m wind
107 speed leveled off near 80 m s^{-1} (Fig. 1), the upper-tropospheric static stability within the eyewall
108 region continued to strengthen, as did the static stability just above the cold-point tropopause
109 radially outside of the eyewall. As the stable layer strengthened, its altitude rose slightly, which
110 corresponded to a slight increase in tropopause height outside of the eyewall during this period.
111 Within the upper troposphere radially outside of the eyewall, meanwhile, static stability decreased
112 such that it was nearly neutral in a thin layer between the 120- and 150-km radii. The eye region

likewise continued to destabilize, and the cold-point tropopause height increased to a level above 18.5 km. This static stability evolution closely follows that observed in Hurricane Patricia (2015; Duran and Molinari 2018).

b. Static stability budget analysis

The left column of Fig. 3 depicts 24-hour changes in N^2 over each of the four days of the simulation. These represent bulk changes computed by subtracting the instantaneous N^2 at the initial time from the instantaneous N^2 at the final time. The middle column of Fig. 3 represents the change in N^2 computed using Eq. XXX and the method described in Section 3. The residual between these two computations (Fig. 3, right column) is much smaller than the change in N^2 , meaning that the budget performs well within the analysis domain.

To determine which of the budget terms are most important, a time series of the contribution of each of the budget terms in Eq. XXX to the tropopause-layer static stability tendency is plotted in Fig. 4. For this figure, each of the budget terms is computed using the method described in Section 3, except with 1-hour averaging intervals instead of 24-hour intervals. The absolute values of these tendencies are then averaged over the radius-height domain depicted in Fig. 3 and plotted as a time series². Advection (Fig. 4, red line) plays an important role in the mean tropopause-layer static stability tendency at all times, and vertical turbulence (Fig. 4, blue line) and radiation (Fig. 4, dark green line) both become important after 48 hours. Although the contribution from horizontal turbulence (Fig. 4, purple line) becomes more important after 72 hours, it is confined to a very small region immediately surrounding the eyewall tangential velocity maximum (not shown), and is negligible throughout the rest of the tropopause layer. The remaining two processes

²It will be seen in subsequent figures that each of the terms contributes both positively and negatively to the N^2 tendency within the analysis domain. Thus, taking an average over the domain tends to wash out the positive and negative contributions. To circumvent this problem, the absolute value of each of the terms is averaged, yielding a time series of the mean magnitude of each budget term.

134 - microphysics and dissipative heating (Fig. 4, orange and light green lines, respectively) - lie atop
135 one another near zero. These time series indicate that, at all times, three budget terms dominate the
136 tropopause-layer static stability tendency: advection, vertical turbulence, and radiation. Variations
137 in the magnitude and spatial structure of these terms drive the static stability changes depicted in
138 Fig. 2; subsequent sections will focus on these variations and what causes them.

139 *(i) 0-24 hours* The first 24 hours of the simulation was characterized by a weakening of the
140 lower-stratospheric static stability maximum above 17 km (Fig. ??a, purple shading) and an in-
141 crease in static stability below (green shading). Although these tendencies extended out to the
142 200-km radius, they were particularly pronounced at innermost radii. A comparison of the contri-
143 butions of advection (Fig. ??b), vertical turbulence (Fig. ??c), and radiation (Fig. ??d) reveals that
144 advection is primarily responsible for the change in static stability during this period. ...Explain
145 this in the context of radial and vertical velocities...

146 *(ii) 24-48 hours* During the second day of the simulation, the lower-stratospheric stable layer
147 continued to weaken (Fig. 6a). This weakening trend in the 16.75-17.75-km layer extended from
148 the 50 km radius outward to past 200 km, and was primarily driven by advection (Fig. 6b). Below
149 this layer, static stability began to increase slightly. This stabilization had contributions from both
150 vertical turbulence (Fig. 6c) and radiation (Fig. 6d) in the 16-16.5-km layer. ...Explain this in
151 context of mean vertical mixing coefficient and mean radiative heating tendency... Meanwhile,
152 radially inward of 60 km, static stability below 17.5 km continued to weaken, primarily due to
153 advective processes.

154 *(iii) 48-72 hours* The third day of the simulation marked a dramatic change in the structure of the
155 tropopause-layer static stability tendencies. During this time, static stability increased markedly
156 in an upward-sloping region within the 30-60-km radial band (Fig. 7a), and also increased within

157 the 16.75-17.5-km layer out to at least the 200-km radius. As this layer stabilized, the layer
158 immediately below it destabilized in a broad region extending from 60-200 km. Examination
159 of the contribution from total advection (Fig. 7b) reveals that advection no longer dominates the
160 static stability tendencies. Instead, a combination of vertical turbulence (Fig. 7c) and radiation
161 (Fig. 7d) overcomes the destabilizing influence of advection to create the layer of increasing static
162 stability. Meanwhile, the destabilizing influence of vertical turbulence in a broad region below
163 17 km combines with a small region of destabilization due to radiation in the 50-120-km radial
164 band combine to destabilize the layer below 16.5 km in the 50-200-km radial band. Comparing
165 the sum of advection and vertical turbulence (Fig. 7e) to the sum of advection, vertical turbulence,
166 and radiation (Fig. 7f) reveals that radiation plays a fundamental role in the re-strengthening of the
167 lower-stratospheric stable layer during this time.

168 *(iv) 72-96 hours*

169 **5. Discussion**

170 Dunion et al. speculate that the diurna pulse only occurs in mature storms. Maybe the develop-
171 ment of the near-tropopause stable layer could partially explain the reason for this.

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173 support of Cloud Model 1. We also thank Drs. Jeffrey Kepert, Robert Fovell, and Erika Navarro
174 for fruitful conversations related to this work. ADD GRANT NUMBER

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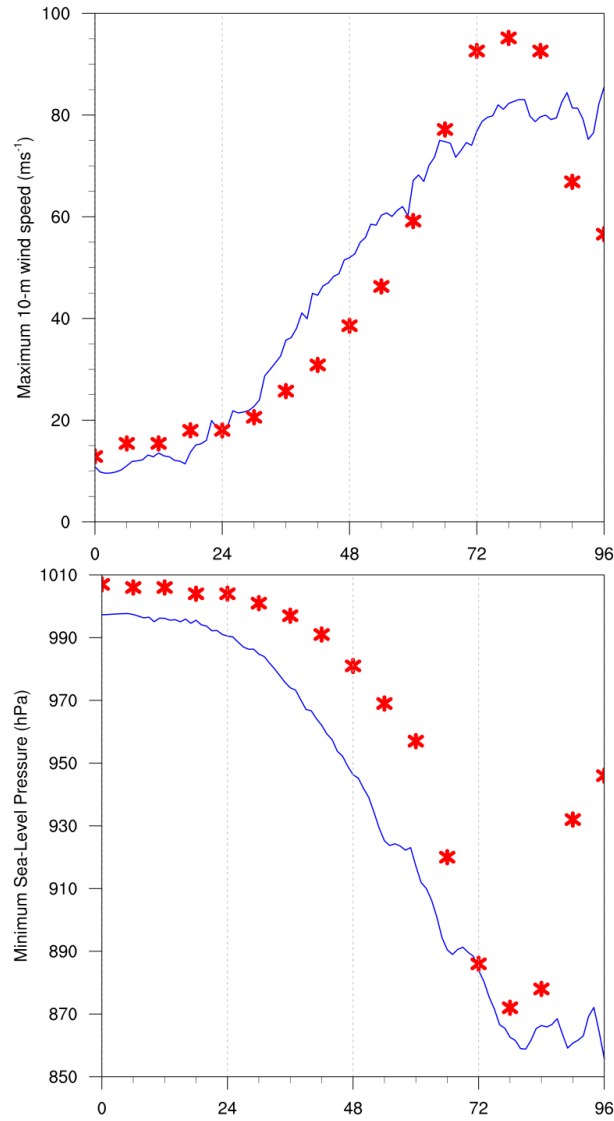


FIG. 1. The maximum 10-m wind speed (top panel; m s^{-2}) and minimum sea-level pressure (bottom panel; hPa) in the simulated storm (blue lines) and from Hurricane Patricia's best track (red stars).

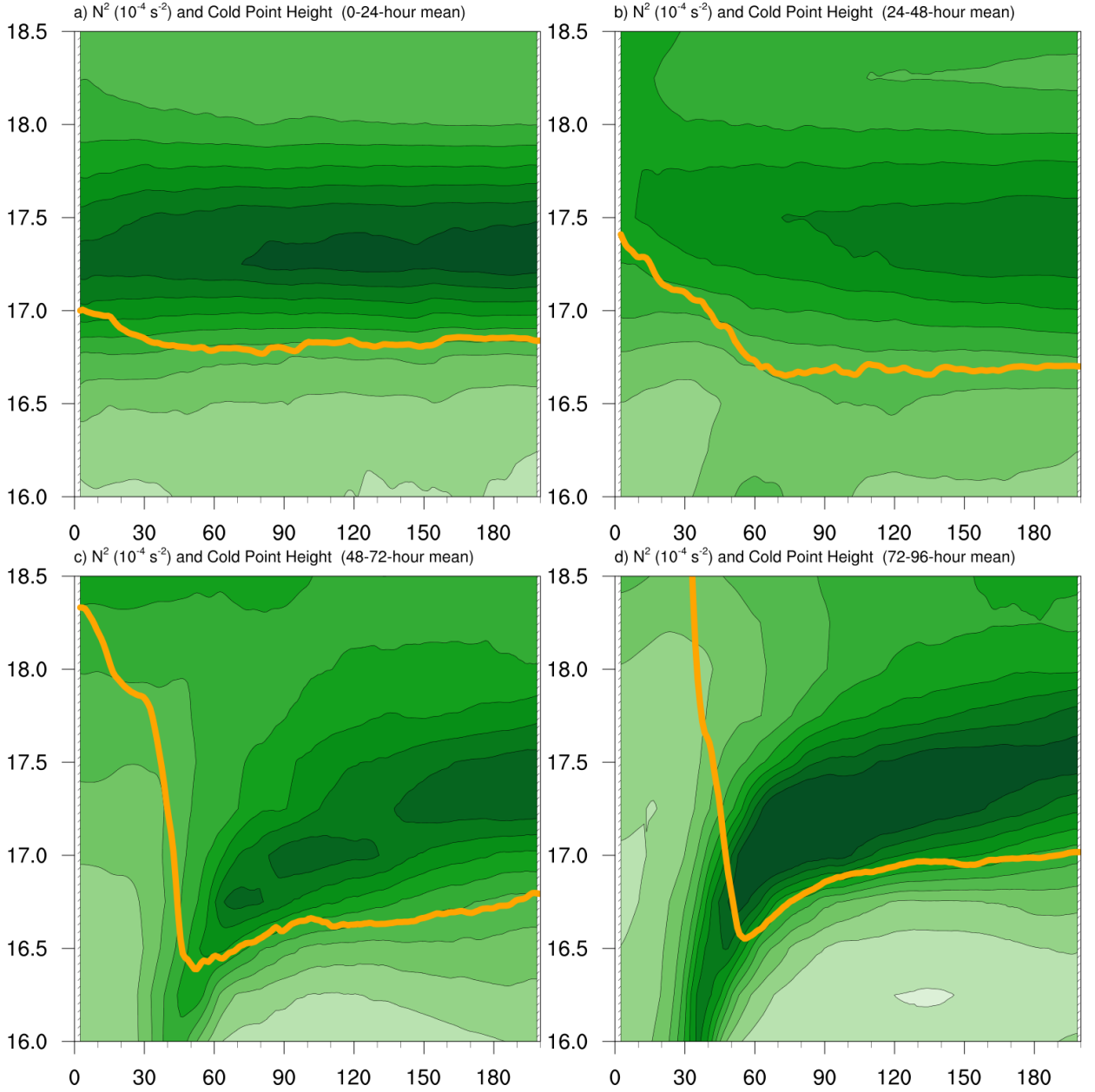


FIG. 2. Twenty-four-hour averages of squared Brunt-Väisälä frequency (10^{-4} s^{-2}) over the first four days of the simulation. Orange lines represent the cold-point tropopause computed from the mean temperature field over the same time periods.

figures/fig03_R-Z_mod+bud+res.png

244 FIG. 3. Left panels: Twenty-four-hour changes in squared Brunt-Väisälä frequency (10^{-4} s^{-2}) over (a) 0-24
245 hours, (b) 24-48 hours, (c) 48-72 hours, (d) 72-96 hours. Middle Panels: The N^2 change over the same time
246 periods computed using Eq. XXX. Right Panels: The budget residual over the same time periods, computed by
247 subtracting the budget change (middle column) from the model change (left column).



figures/fig04_AVG_budterms.png

248 FIG. 4. Time series of the contribution of each of the budget terms to the time tendency of the squared
249 Brunt-Väisälä frequency (N^2 ; 10^{-4} s^{-2}). For each budget term, the absolute value of the N^2 tendency is averaged
250 both temporally over 1-hour periods (using output every minute), and spatially within the radius-height domain
251 depicted in Fig. 3.

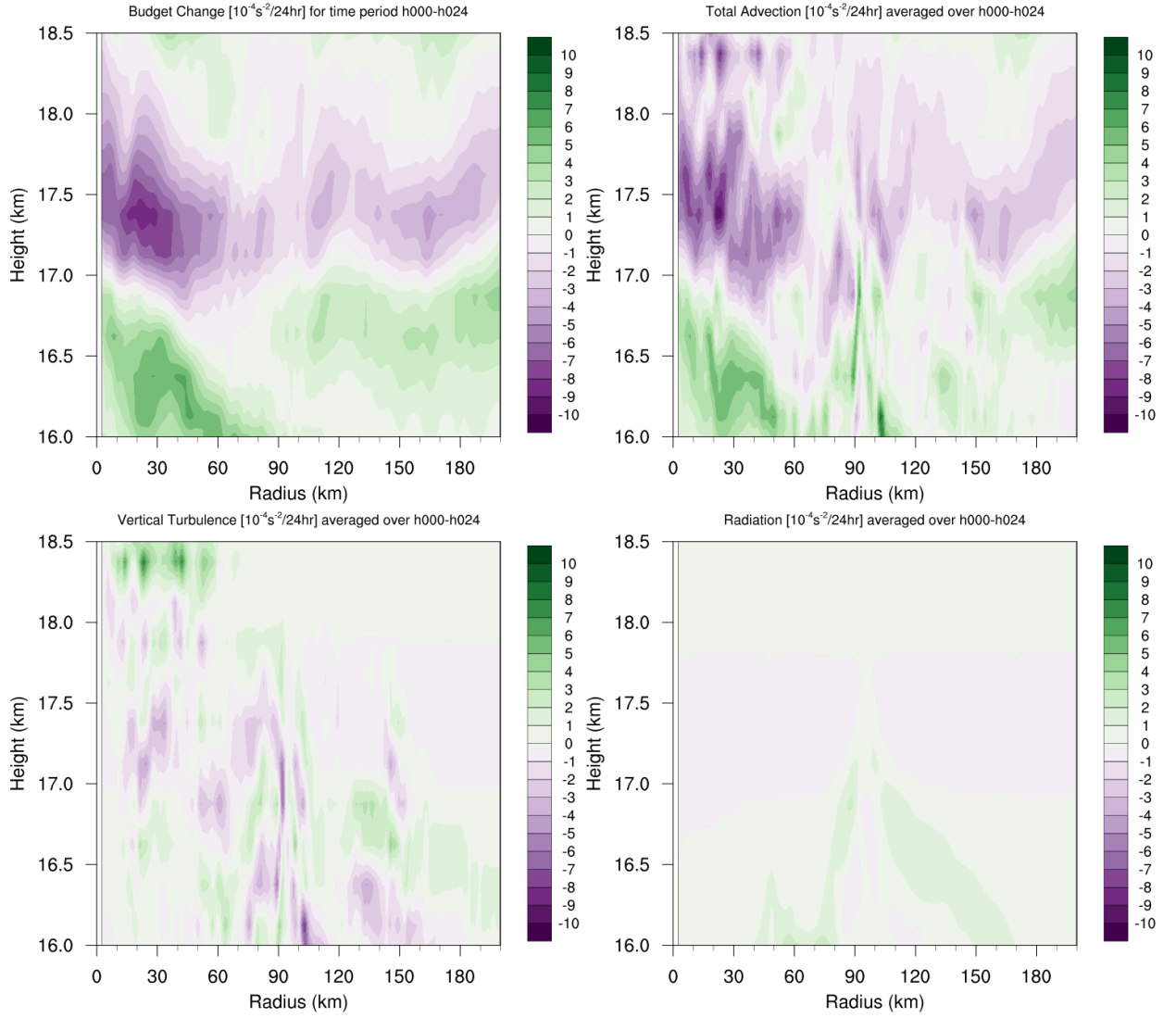


FIG. 5. (a) Total change in N^2 over the 0-24-hour period ($10^{-4} \text{ s}^{-2} (24 \text{ hr})^{-1}$) and the contributions to that change from (b) the sum of horizontal and vertical advection, (c) vertical turbulence, and (d) the sum of longwave and shortwave radiation.

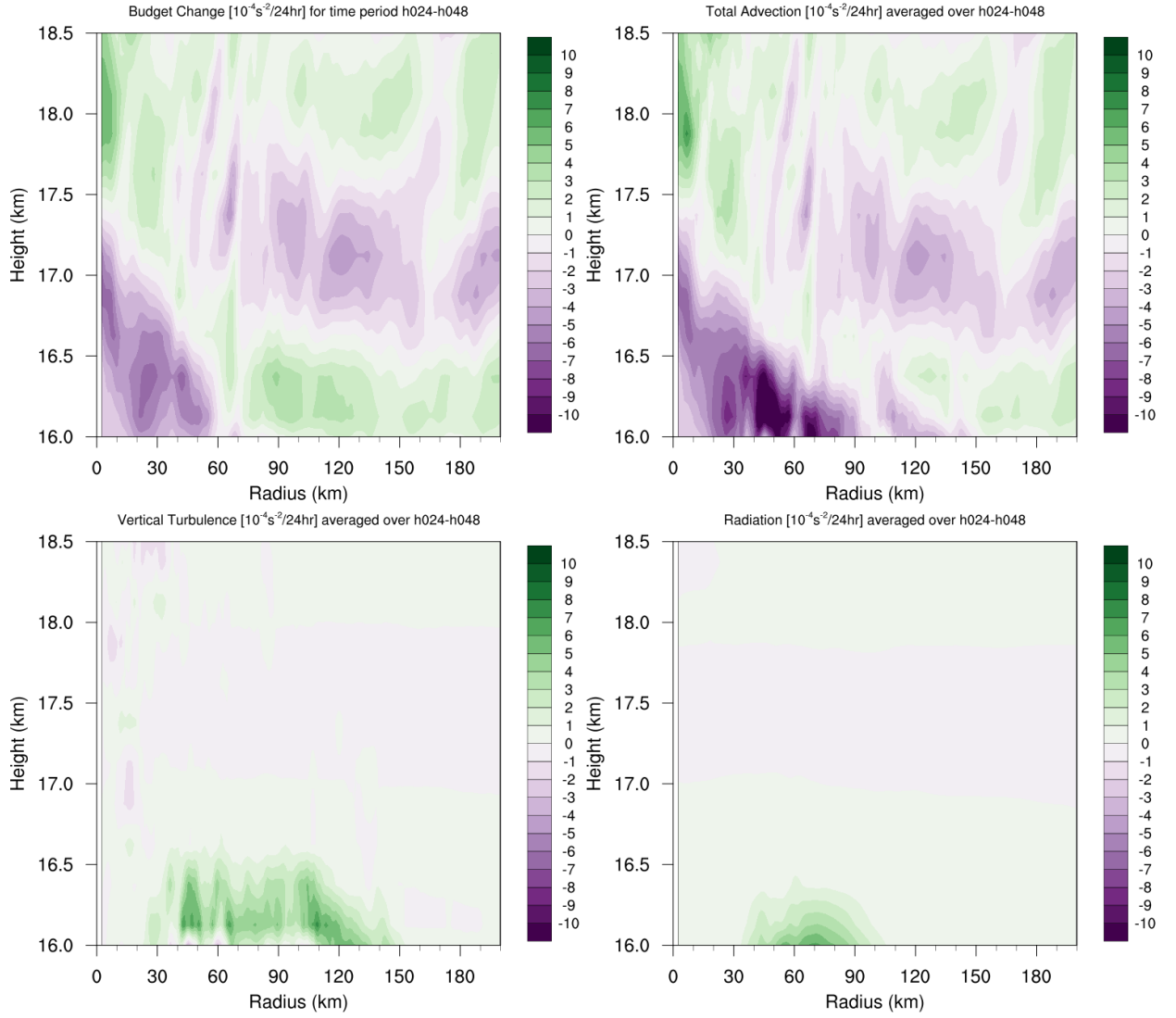


FIG. 6. As in Fig. 5, but for the 24-48-hour period.

figures/fig07_h048-h072-budgetterms.png

255 FIG. 7. (a) Total change in N^2 over the 48-72-hour period ($10^{-4} \text{ s}^{-2} (24 \text{ hr})^{-1}$) and the contributions to that
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257 and shortwave radiation, (e) the sum of horizontal advection, vertical advection, and vertical turbulence, and (f)
258 the sum of horizontal advection, vertical advection, vertical turbulence, and longwave and shortwave radiation.

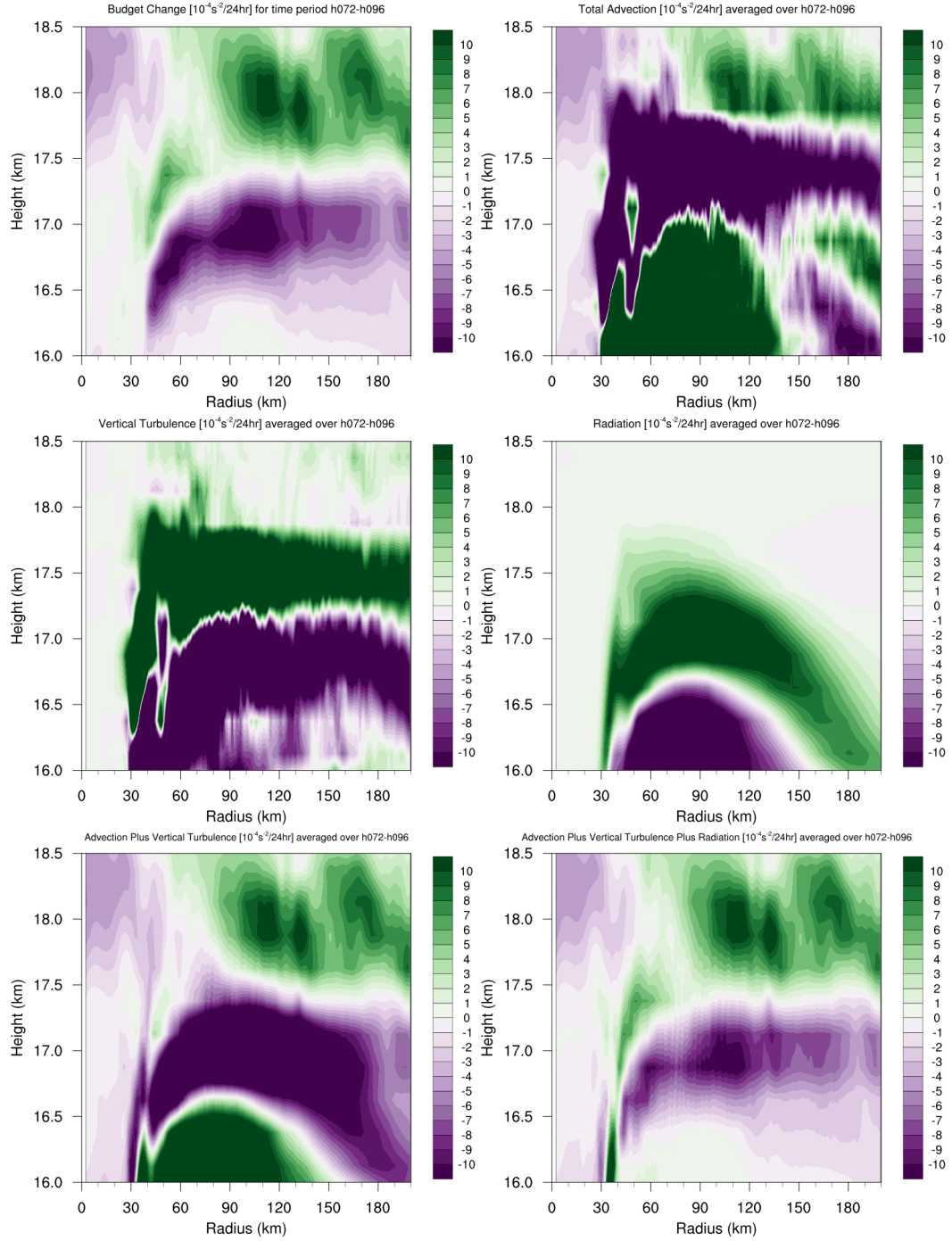


FIG. 8. As in Fig. 7, but for the 72-96-hour period.

figures/fig09_u.png

259 FIG. 9. Radial velocity (m s^{-1} ; filled contours), potential temperature (K; thick black contours), and cold point
260 tropopause height (orange line) averaged over (a) 0-24 hours, (b) 24-48 hours, (c) 48-72 hours, and (d) 72-96
261 hours.

figures/fig10_w.png

262 FIG. 10. Vertical velocity (cm s^{-1} ; filled contours), potential temperature (K; thick black contours), and cold
263 point tropopause height (orange line) averaged over (a) 0-24 hours, (b) 24-48 hours, (c) 48-72 hours, and (d)
264 72-96 hours.

figures/fig11_qtot.png

265 FIG. 11. Total condensate mixing ratio (g kg^{-1}) and cold point tropopause height (orange line) averaged over
266 (a) 0-24 hours, (b) 24-48 hours, (c) 48-72 hours, and (d) 72-96 hours.