Tropopause Evolution in a Rapidly Intensifying Tropical Cyclone: A Static

Stability Budget Analysis

- Patrick Duran* and John Molinari
- 4 University at Albany, State University of New York, Albany, NY

^{5 *}Corresponding author address: Department of Atmospheric and Environmental Sciences, Univer-

⁶ sity at Albany, State University of New York, 1400 Washington Avenue, Albany, NY.

E-mail: pduran2008@gmail.com

ABSTRACT

8 Enter the text of your abstract here.

9 1. Introduction

There will be a whole bunch of papers cited here...

11 2. Model Setup

- Put description of Fig. 1 in this section.
- Don't forget to mention 1-2-1 smoother.

14 3. Budget Computation

Add details of budget computation here...

16 4. Results

a. Static stability evolution

The average N² over the first day of the simulation (Fig. 2a) indicates the presence of a static 18 stability maximum about 400 m above the cold-point tropopause. This lower-stratospheric stable 19 layer had begun to erode during the initial spin-up period, with the maximum destabilitzation occurring at the innermost radii. This decrease in static stability continued into the second day 21 of the simulation (Fig. 2b) as the storm intensified to hurricane strength (Fig. 1). Destabilization was particularly pronounced over the developing eye, where the time-mean cold-point tropopause height increased by up to 400 m compared to the previous day. Over the developing eyewall 24 and outer rainband regions, meanwhile, the tropopause height remained nearly constant. During 25 the third day of the simulation (Fig. 2c), static stability over the eye continued to decrease, and the cold-point tropopause height rose to 18.3 km at the storm center. The tropopause sloped 27 sharply downward over the innermost radii, reaching the 16.4-km level near the 50-km radius. This local minimum in tropopause height corresponded to the eyewall region, where upper-tropospheric static stability increased during this time period. Outside of the eyewall region, static stability
began to increase in the layer immediately overlying the cold-point tropopause. This stable layer
sloped upward with radius, which corresponded to an upward-sloping tropopause radially outside
of the eywall region. Over the next 24 hours (Fig. 2d), as the storm's maximum 10-m wind speed
leveled off near 80 m s⁻¹ (Fig. 1), the upper-tropospheric static stability within the eyewall region
continued to strengthen, along with the static stability within the lower stratosphere radially outside
of the eyewall. As the stable layer strengthened, its altitude rose slightly, which corresponded to
a slight increase in tropopause height outside of the eyewall during this period. Within the upper
troposphere radially outside of the eyewall, meanwhile, static stability decreased such that it was
nearly neutral in a thin layer between the 120- and 150-km radii. The eye region likewise continued
to destabilize, and the cold-point tropopause height increased to a level above 18.5 km. This static
stability evolution closely follows that observed in Hurricane Patricia (2015; Duran and Molinari
2018).

b. Static stability budget analysis

- The left column of Fig. 3 depicts 24-hour changes in N² over each of the four days of the simulation. These represent bulk changes computed by subtracting the instantaneous N² at the initial time from the instantaneous N² at the final time. The middle column of Fig. 3 represents the change in N² computed using Eq. XXX and the method described in Section 3. The residual between these two computations (Fig. 3, right column) is much smaller than the change in N², meaning that the budget performs well within the analysis domain.
- To determine which of the budget terms are most important, a time series of the contribution of
 each of the budget terms in Eq. XXX to the tropopause-layer static stability tendency is plotted
 in Fig. 4. For this figure, each of the budget terms is computed using the method described in

Section 3, except with 1-hour averaging intervals instead of 24-hour intervals. The absolute values of these tendencies are then averaged over the radius-height domain depicted in Fig. 3 and plotted as a time series¹. Advection (Fig. 4, red line) plays an important role in the mean tropopause-55 layer static stability tendency at all times, and vertical turbulence (Fig. 4, blue line) and radiation (Fig. 4, dark green line) both become important after 48 hours. Although the contribution from 57 horizontal turbulence (Fig. 4, purple line) becomes more important after 72 hours, it is confined to a very small region immediately surrounding the eyewall tangential velocity maximum (not shown), and is negligible throughout the rest of the tropopause layer. The remaining two processes microphysics and dissipative heating (Fig. 4, orange and light green lines, respectively) - lie atop 61 one another near zero. These time series indicate that, at all times, three budget terms dominate the tropopause-layer static stability tendency: advection, vertical turbulence, and radiation. Variations in the magnitude and spatial structure of these terms drive the static stability changes depicted in Fig. 2; subsequent sections will focus on these variations and what causes them.

- 66 (i) 0-24 hours
- 67 (ii) 24-48 hours
- 68 (iii) 48-72 hours
- 69 (iv) 72-96 hours
- 70 Acknowledgments. Start acknowledgments here.

 $^{^{1}}$ It will be seeen in subsequent figures that each of the terms contributes both positively and negatively to the N^{2} tendency within the analysis domain. Thus, taking an average over the domain tends to wash out the positive and negative contributions. To circumvent this problem, the absolute value of each of the terms is averaged, yielding a time series of the mean magnitude of each budget term.

71 References

- Duran, P., and J. Molinari, 2018: Dramatic inner-core tropopause variability during the rapid
- intensification of Hurricane Patricia (2015). Mon. Wea. Rev., XXX (X), XXX–XXX.

LIST OF FIGURES

75	Fig. 1.	The maximum 10-m wind speed (top panel; m s ⁻²) and minimum sea-level pressure (bottom
76		panel; hPa) in the simulated storm (blue lines) and from Hurricane Patricia's best track (red
77		stars)
78	Fig. 2.	Twenty-four-hour averages of squared Brunt-Väisälä frequency (10 ⁻⁴ s ⁻²) over the first four
79		days of the simulation. Orange lines represent the cold-point tropopause computed from the
80		mean temperature field over the same time periods
81	Fig. 3.	Left panels: Twenty-four-hour changes in squared Brunt-Väisälä frequency (10 ⁻⁴ s ⁻²) over
82		(a) 0-24 hours, (b) 24-48 hours, (c) 48-72 hours, (d) 72-96 hours. Middle Panels: The N^2
83		change over the same time periods computed using Eq. XXX. Right Panels: The budget
84		residual over the same time periods, computed by subtracting the budget change (middle
85		column) from the model change (left column)
86	Fig. 4.	Time series of the contribution of each of the budget terms to the time tendency of the
87		squared Brunt-Väisälä frequency (N^2 ; 10^{-4} s ⁻²). For each budget term, the absolute value
88		of the N ² tendency is averaged both temporally over 1-hour periods (using output every
89		minute), and spatially within the radius-height domain depicted in Fig. 3
90	Fig. 5.	
91	Fig. 6.	
92	Fig. 7.	
93	Fig. 8.	

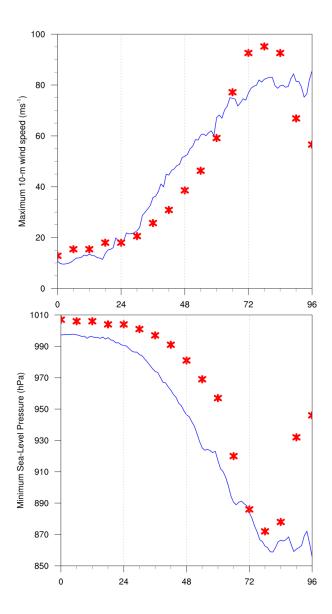


FIG. 1. The maximum 10-m wind speed (top panel; m s⁻²) and minimum sea-level pressure (bottom panel;

hPa) in the simulated storm (blue lines) and from Hurricane Patricia's best track (red stars).

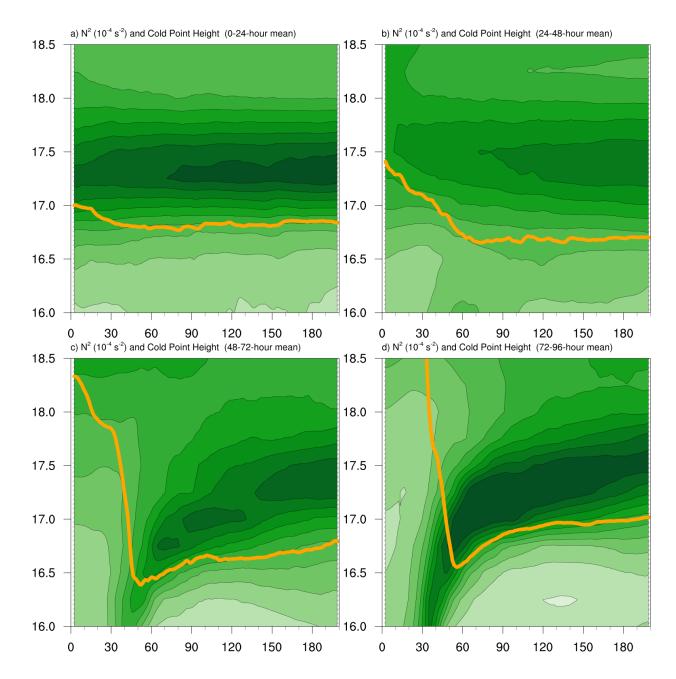


FIG. 2. Twenty-four-hour averages of squared Brunt-Väisälä frequency (10⁻⁴ s⁻²) over the first four days of the simulation. Orange lines represent the cold-point tropopause computed from the mean temperature field over the same time periods.

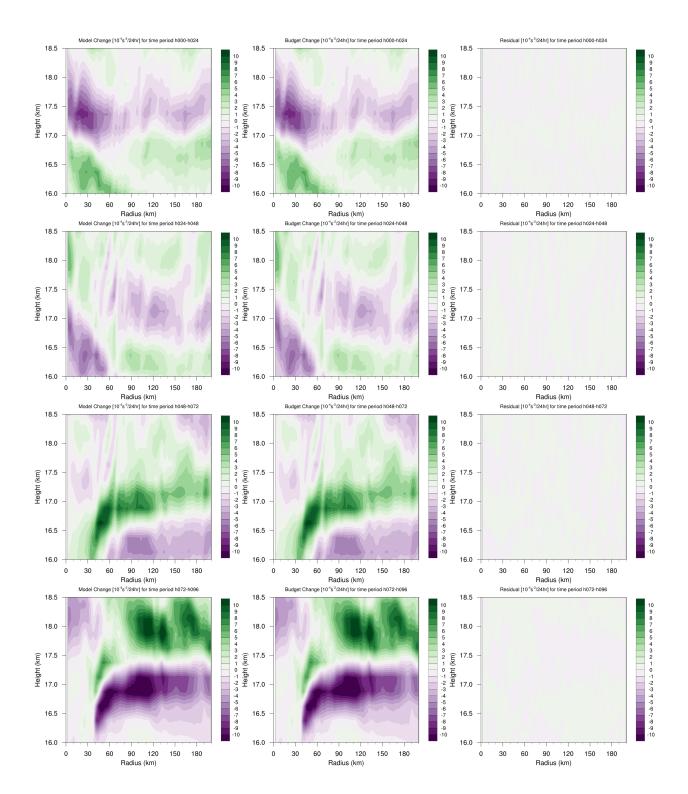


FIG. 3. Left panels: Twenty-four-hour changes in squared Brunt-Väisälä frequency (10⁻⁴ s⁻²) over (a) 0-24 hours, (b) 24-48 hours, (c) 48-72 hours, (d) 72-96 hours. Middle Panels: The N² change over the same time periods computed using Eq. XXX. Right Panels: The budget residual over the same time periods, computed by subtracting the budget change (middle column) from the model change (left column).

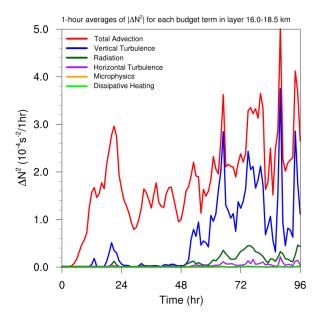


FIG. 4. Time series of the contribution of each of the budget terms to the time tendency of the squared Brunt-Väisälä frequency (N^2 ; 10^{-4} s⁻²). For each budget term, the absolute value of the N^2 tendency is averaged both temporally over 1-hour periods (using output every minute), and spatially within the radius-height domain depicted in Fig. 3.

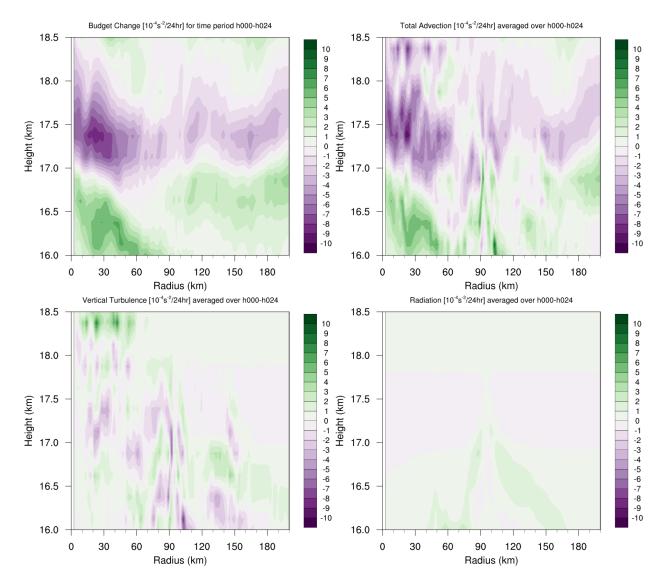


FIG. 5.

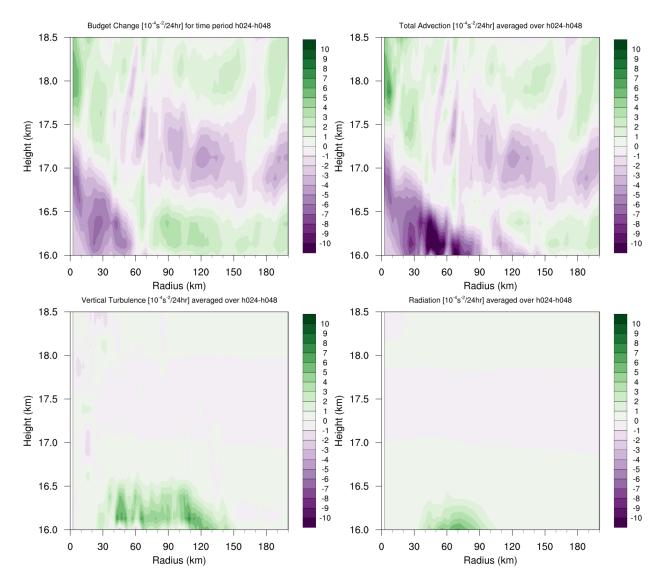


FIG. 6.

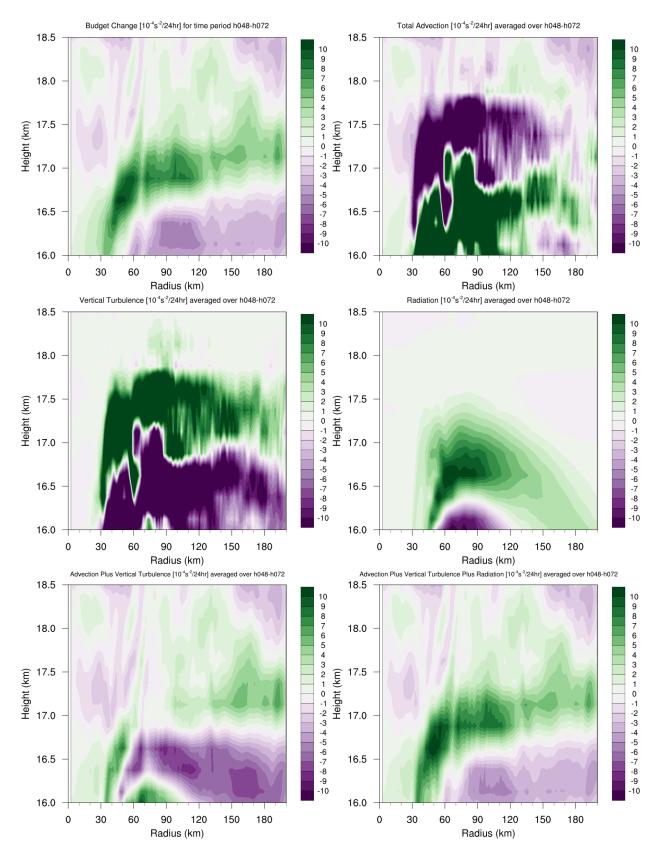


FIG. 7.

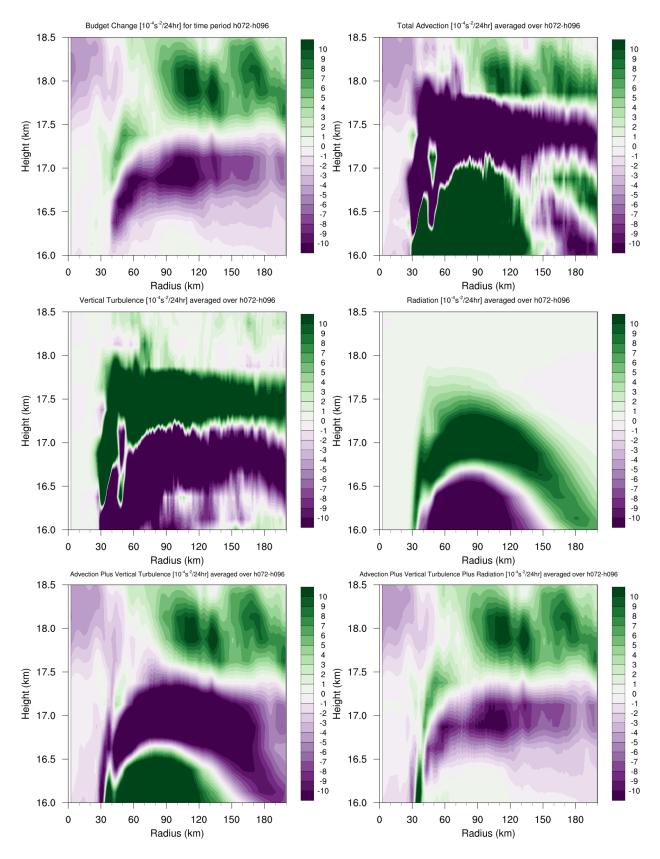


FIG. 8.