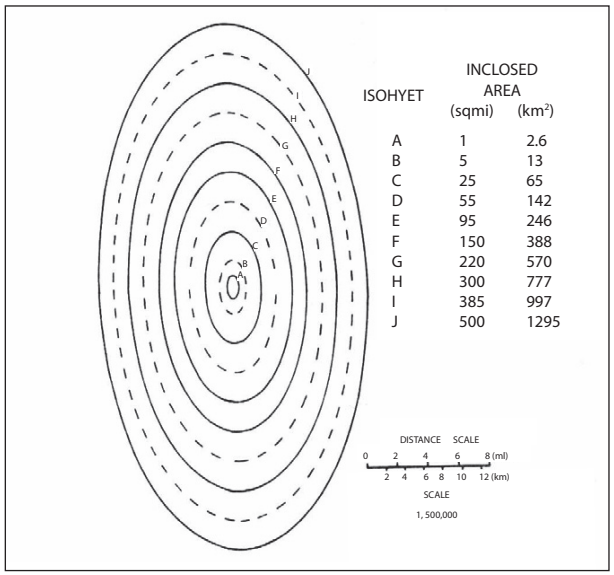

Figure 5.44. Idealized isohyetal pattern for local storm PMP areas up to 1 300 km² (Hansen and others, 1977)

from around central Washington to the Canadian border. The new study results in PMP from 25 to 40 mm lower.

Comparisons were also made with adjoining studies, including HMR No. 49 (Hansen and others, 1977) and HMR No. 55A (Hansen and others, 1988). Some of the different assumptions regarding elevation and durational characteristics have already been discussed.

Relative to HMR No. 49 (Hansen and others, 1977), the differences in 1-hour 2.6-km² PMP are near zero in extreme northern Utah, becoming more positive moving westward to a maximum of around +40 mm along the California–Oregon border area. The primary reason for this discrepancy may come from transposing the Morgan, Utah, storm throughout the southern portions of the Northwest. HMR No. 49 (Hansen and others, 1977) and the present study support a preferred seasonality of storms and do not attempt to apply seasonal curves or nomograms.

No significant PMP differences exist in local storm PMP between the current study and HMR No. 55A (Hansen and others, 1988). No major new storms were found within this general area which would cause any increase in PMP to be made, and no evidence was revealed which might indicate a lowered estimate. Seasonality for HMR No. 55A (Hansen and others, 1988) showed a distinct summer maximum in extreme local storms, a finding in agreement with this.

5.3.7.7 Estimating procedure of local storm PMP for specific basin

(a) The 1-hour 2.6-km² PMP for elevations at or below 1 830 m is determined by locating the basin on Figure 5.45 and determining the basin average 1-hour 2.6-km² PMP local storm index PMP. Linear interpolation is assumed to be applicable.

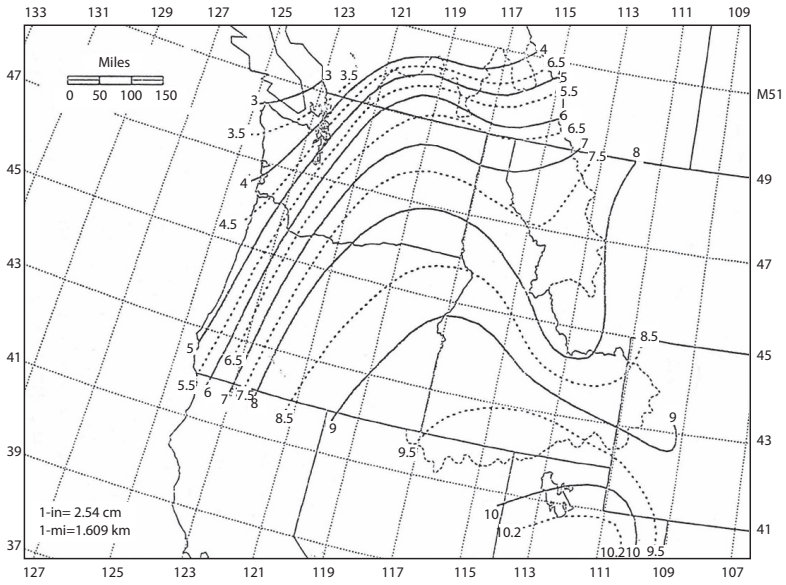

Figure 5.45. One-hour 2.6-km² local storm PMP in inches for elevations to 1 830 m