

GEO 418 – "Hyperspectral Earth Observation"

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Physics of Optical Remote Sensing

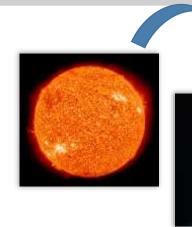
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Image Chain

"Way of a photon

from the sun









through atmosphere

to earth

to **EO** sensor

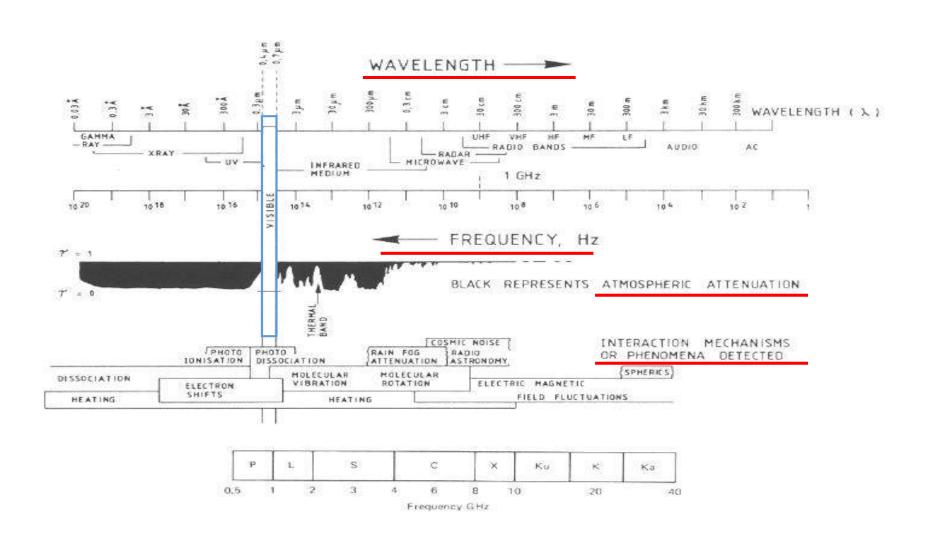
to image"



1st: Radiation Source

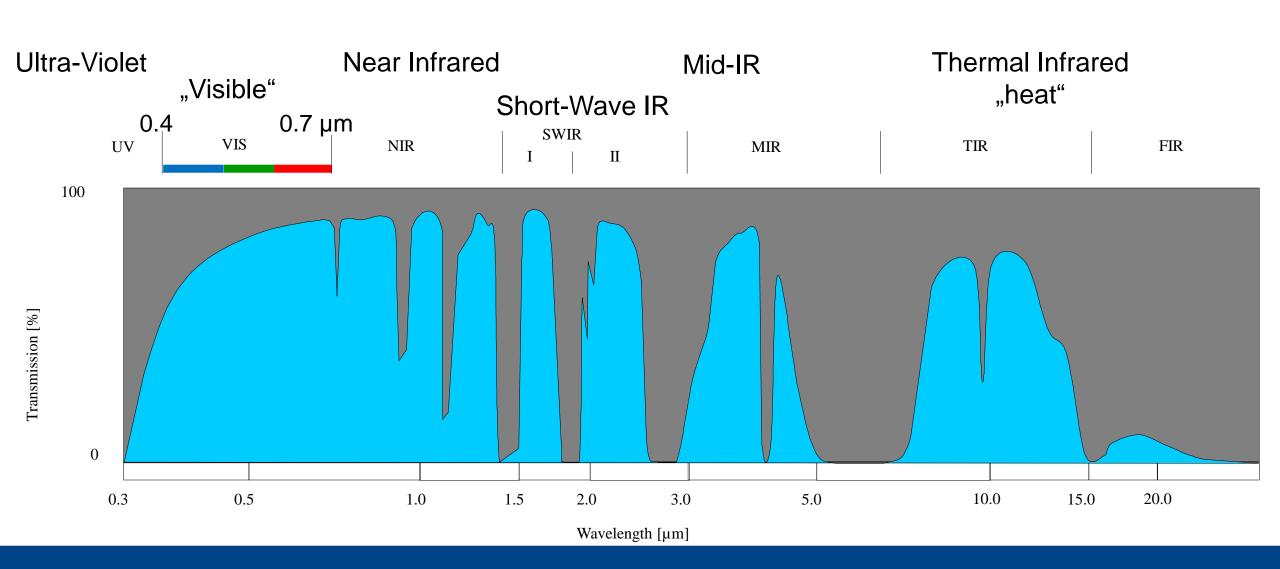


The electro-magnetic spectrum



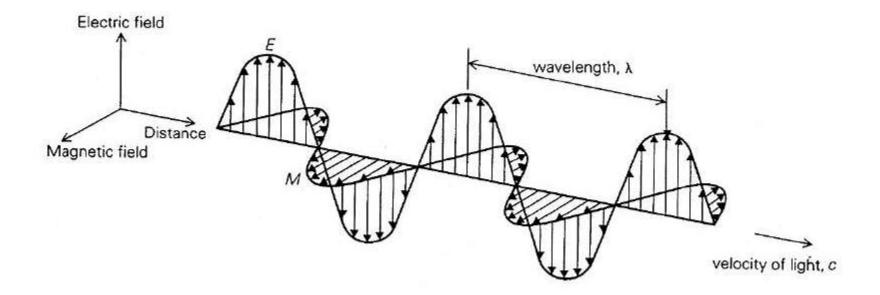


The electro-magnetic spectrum



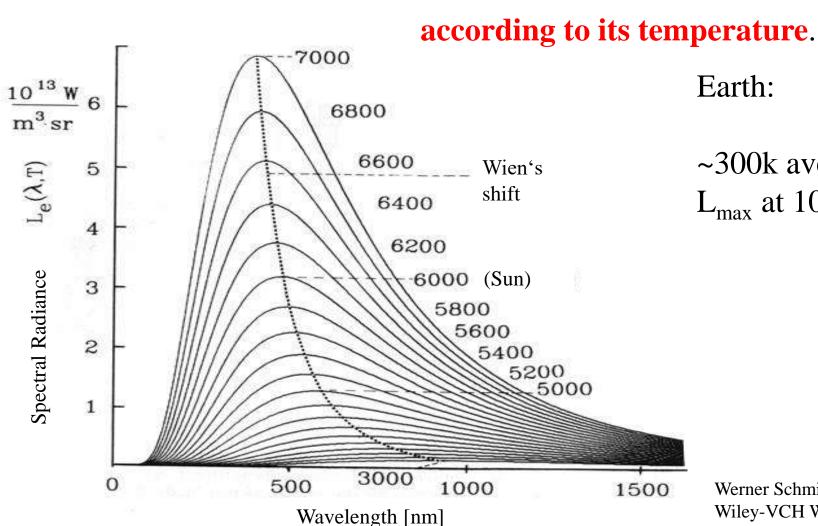


Every object is emitting electro-magnetic radiation





Every object is emitting electro-magnetic radiation



Earth:

~300k average surface temp. L_{max} at 10 µm (=10.000nm)

Werner Schmidt: Optische Spektroskopie – Eine Einführung. Wiley-VCH Weinheim, 2. Auflage 2000

Planck:

Emittance $M = f(\lambda, T)$

Boltzmann:

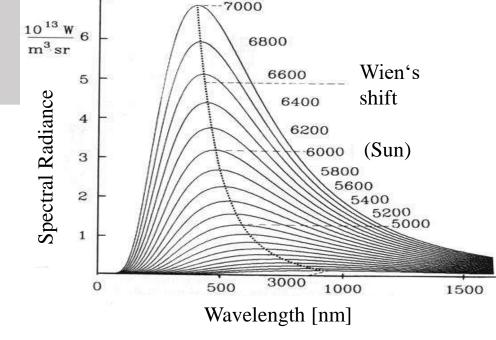
$$M_{blackbody} = \sigma T^{4} [W m^{-2}]$$
 $\sigma = 5.67 * 10^{-8} W m^{-2} K^{-1}$

In general: $M = \epsilon T^n$

n...property of material, ~4

• Wien:

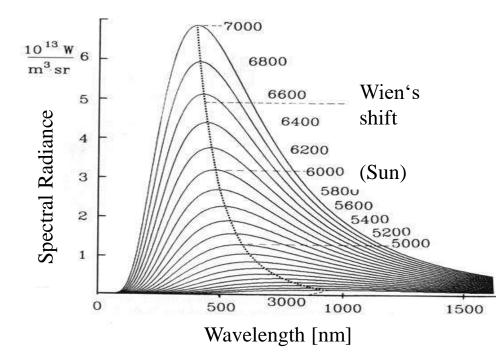
The higher the temperature (T) the shorter the wavelength (λ_{max}) of maximal energy emittance λ_{max} = 2898 / T



Electromagnetic wave (Maxwells Equation)

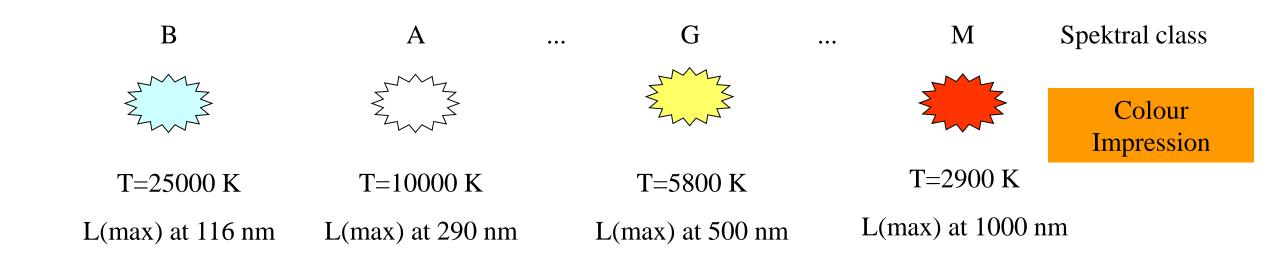
- Energy q = h * v
 h: Planck's Constant: 6.63* 10 -34 [J s]
 => Higher energy per photon with shorter wavelength
- Energy pro Photon: green 0.55 μ m => 3.61 * 10 $^{-19}$ J TIR 12.00 μ m => 1.66 * 10 $^{-20}$ J

=> 22times higher energy!





• Example for Planck- and Wien-Equations: surface temperature of stars



Colour impression: continuum & relative strengths of spectral lines in the VIS



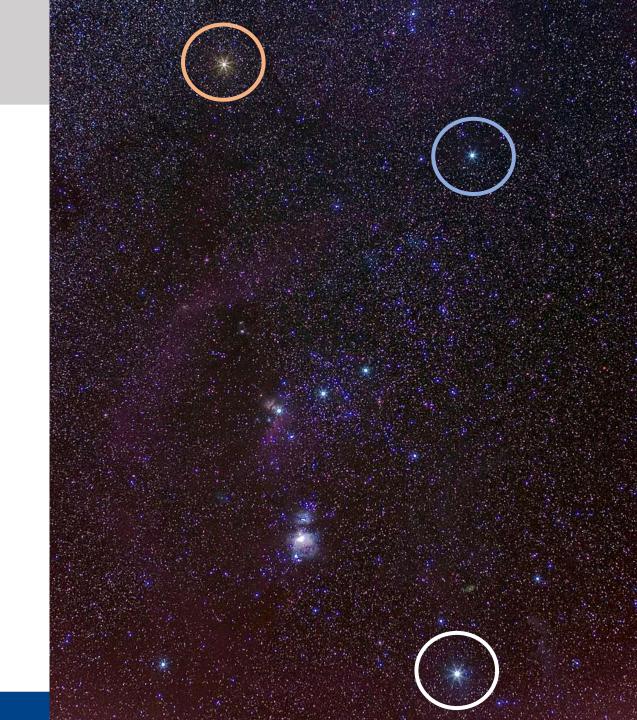
The color of a star is determined by its temperature, according to Wien's law. In the constellation of Orion, one can compare

Betelgeuse (T ≈ 3300 K, upper left) orange

Rigel (T = 12100 K, bottom right) white

Bellatrix (T = 22000 K, upper right) blue

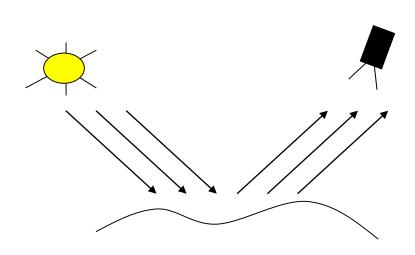
Source (modified): https://en.wikipedia.org/wiki/Wien%27s_displacement_law

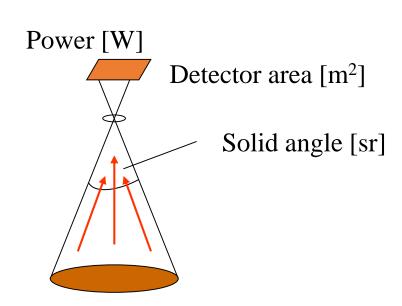


• The sensor detects:

Radiance, At-Sensor Radiance L [W m⁻² sr⁻¹]

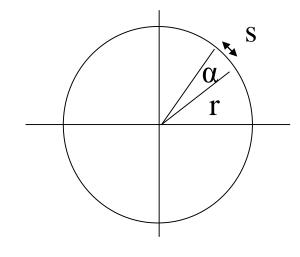
=> Unit after system correction, described as L1B product

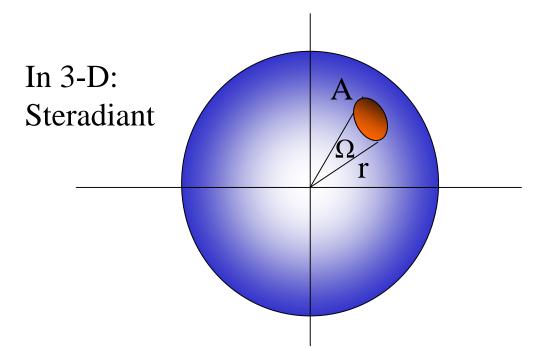




- Solid Angle Ω (Raumwinkel) steradiant = area / radius2
- Sphere = $4\pi r^2 / r^2 = 4\pi [sr]$
- Sky = $2 \pi [sr]$

In 2-D: Radiant







 Radiant Energy Q [J] Strahlungsenergie

i.e. Energy

• Radiant flux, or: photon flux Φ [W] = dQ/dt Strahlungsfluß Theta, d.h. Energie pro Zeiteinheit

i.e. Energy per time interval

Radiant flux density [W m⁻²]
 E: irradiance // M: emittance
 Strahlungsflußdichte, d.h. Energie pro Erfassungsfläche

i.e. Energy per detector area

 Radiant intensity I [W sr⁻¹] Strahlungsintensität

i.e. Energy per solid angle

• Radiance L [W m⁻² sr⁻¹] Strahldichte, d.h. Energie pro Zeit, Fläche, Raumwinkel i.e. Energy per time, detector area, solid angle

When considering <u>spectral</u> measures: additional integration over wavelength interval [μm⁻¹]
 i.e., Spectral Radiance L in [mW cm⁻² sr⁻¹ μm⁻¹] or [W m⁻² sr⁻¹ μm⁻¹]
 ...ATCOR" ...Landsat-8"



Conversion of Radiance L to other radiance units:

```
L = dQ / (dT dA dΩ) = (d2 Φ) / (dA dΩ cos Θ) = (dE) / (dΩ cos Θ) = (dI) / (dA cos Θ)
Radiant Energy Q Radiant Flux Φ Irradiance E / Emittance M Radiant Intensity I
Strahlungsenergie Strahlungsfluß Strahlungsflußdichte Strahlungsintensität I
```

=> Energy, integrated over dT TIME, dA AREA, $d\Omega$ SOLID ANGLE

DLR

Physical basics

For EO applications, measure should be

- Independent of incoming radiation (power and geometry)
- Independent of sensor properties (detector characteristics)
- Independent from atmospheric conditions
 - => material property only!

Thus more suitable measure: reflection p

- No unit, but [%]
- Independent from illumination & sensor
- (Almost) independent from geometry & atmosphere



Note there's a big difference between TOA Top-Of-Atmosphere \mathbf{p} and BOA Bottom-Of-Atmosphere \mathbf{p}

• TOA Reflectance

$$\rho_{TOA} = \frac{\pi L d^2}{E0 \cos \theta_S}$$

d: Earth-Sun distance (astro. units)

E0: Extraterr. solar irradiance

 θ_s : Solar zenith angle

• BOA Reflectance

$$\rho_{BOA} = \frac{\pi \left(d^2 L - L_{path} \right)}{\tau E_g}$$

τ: Ground-to-sensor atmospheric transmittance

L_{path}: Path radiance E_g: Global flux TOA Exosphere Thermosphere Mesosphere Stratosphere Troposhere BOA

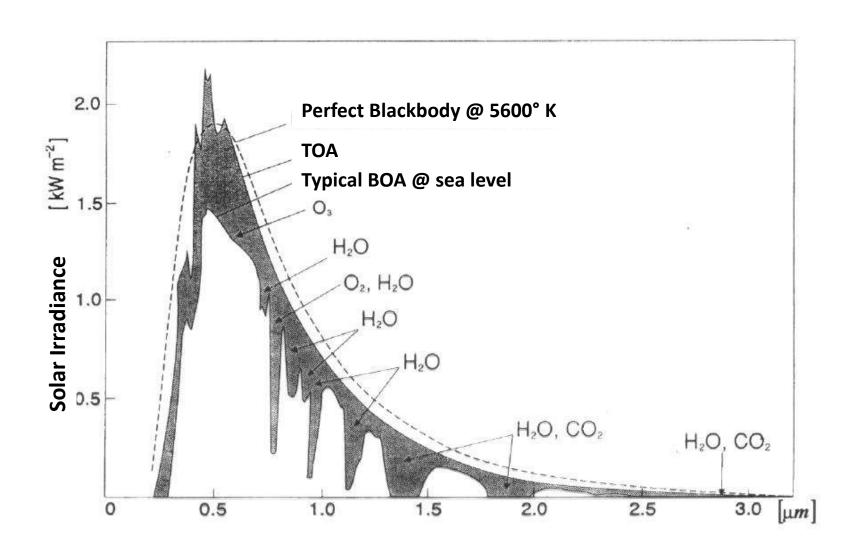
Source: Nasa:GOV ISS030-E-031275



2nd: Atmosphere

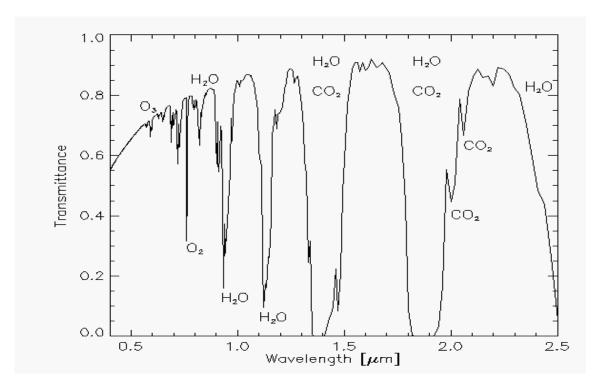


Atmospheric Transmittance

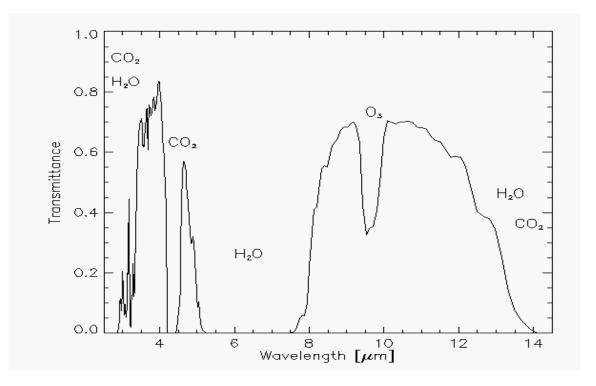




Atmospheric Transmittance



Solar spectral region



Thermal spectral region



Major absorbers in the earth's atmosphere:

```
Water vapor, CO<sub>2</sub>, O<sub>2</sub>, O<sub>3</sub>, CH<sub>4</sub>
```

Scattering: molecular (Rayleigh, mostly N₂, O₂) and aerosol

Physical characteristics of aerosols:

size distribution, refractive index

Optical characteristics:

scattering & extinction coefficient

scattering phase function



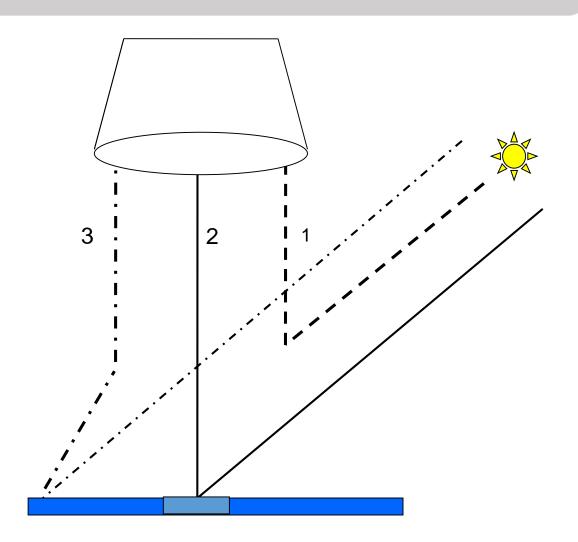
Assuming flat terrain, the signal sensed consists of the following components:

0: sensor calibration & characteristics

1: path radiance (scattered photons)

2: reflected radiance

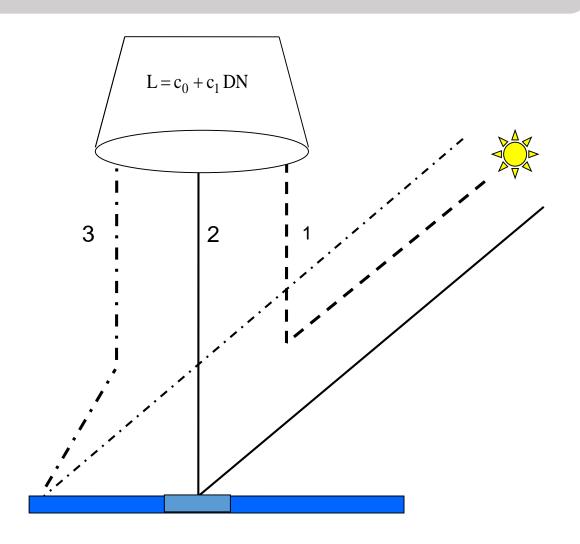
3: adjacency radiance (reflected & scattered)





Sensor calibration:

$$L_{Sensor} = c_0 + c_1 * DN$$





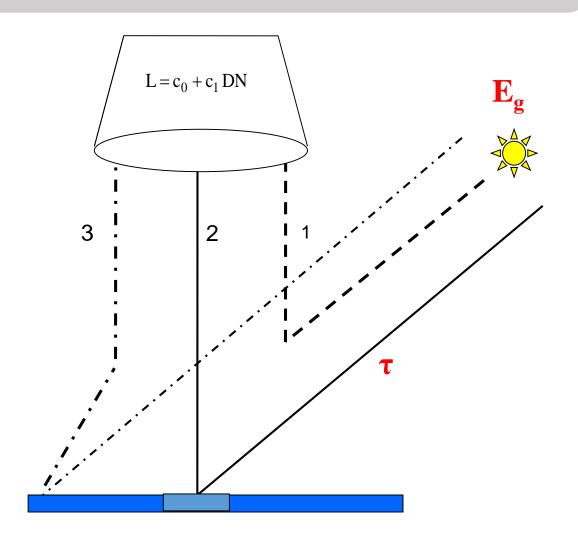
Sensor calibration:

$$L_{Sensor} = c_0 + c_1 * DN$$

At-sensor radiance:

$$L_{Sensor} = L_1 + L_2 + L_3$$

$$L_{Sensor} = L_{path} + L_{reflected} + L_{adjacency}$$



 E_g : global flux τ : total atm. transmittance



Sensor calibration:

$$L_{Sensor} = c_0 + c_1 * DN$$

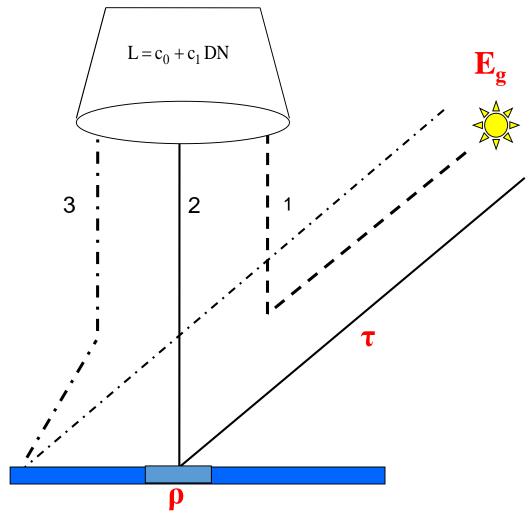
At-sensor radiance:

$$L_{Sensor} = L_1 + L_2 + L_3$$

$$L_{Sensor} = L_{path} + L_{reflected} + L_{adjacency}$$

Target radiance:

$$L_{reflected} = \rho E_g * \tau / \pi$$



 E_g : global flux τ : total atm. transmittance ρ : target reflectance

When neglecting the adjacency effect:

$$L_{Sensor} = L_{path} + L_{reflected} = L_{path} + \rho E_{g} * \tau / \pi = c_{0} + c_{1} * DN$$

When neglecting the adjacency effect:

$$L_{Sensor} = L_{path} + L_{reflected} = L_{path} + \rho E_{g} * \tau / \pi = c_{0} + c_{1} * DN$$

Now solving for the target reflectance:

$$\rho = \frac{\pi((c_o + c_1 * DN) - L_{Path})}{\tau * E_g}$$

This means:

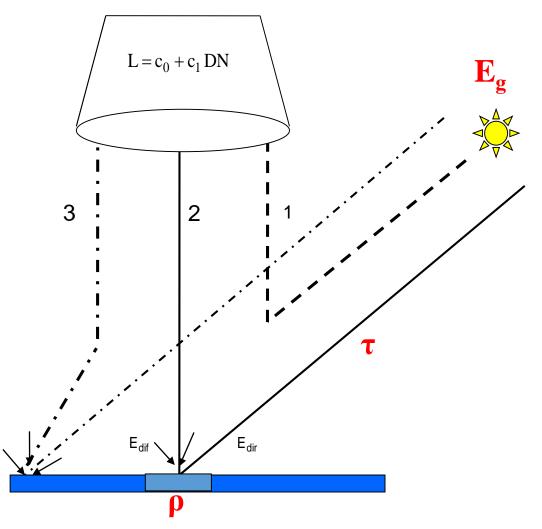
- Estimates of the main atm. parameters (aerosol type, optical thickness and water vapor) required for L_{path} , E_g and τ
- Accurate sensor calibration is mandatory



Adding some details:

$$E_g = E_{direct} + E_{diffuse}$$

$$\tau = \tau_{direct} + \tau_{diffuse}$$



 E_g : global flux τ : total atm. transmittance ρ : target reflectance diffuse + direct



Adding some details:

$$E_g = E_{direct} + E_{diffuse}$$

$$\tau = \tau_{direct} + \tau_{diffuse}$$

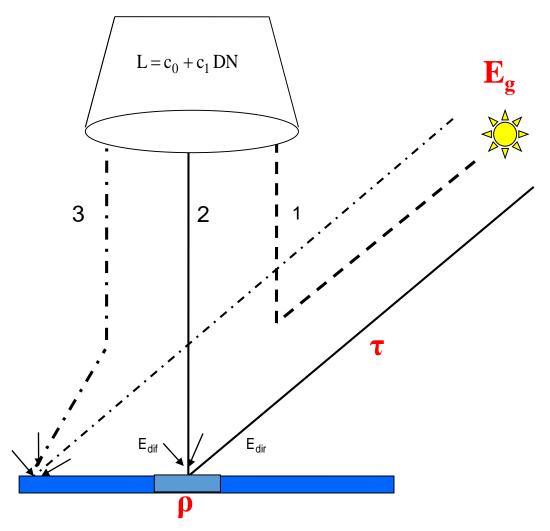
And adding adjacency effects:

$$L_{Sensor} = L_{path} + \rho E_g \tau (1/\pi) * (1/(1-\bar{\rho} s))$$

ρ: background reflection

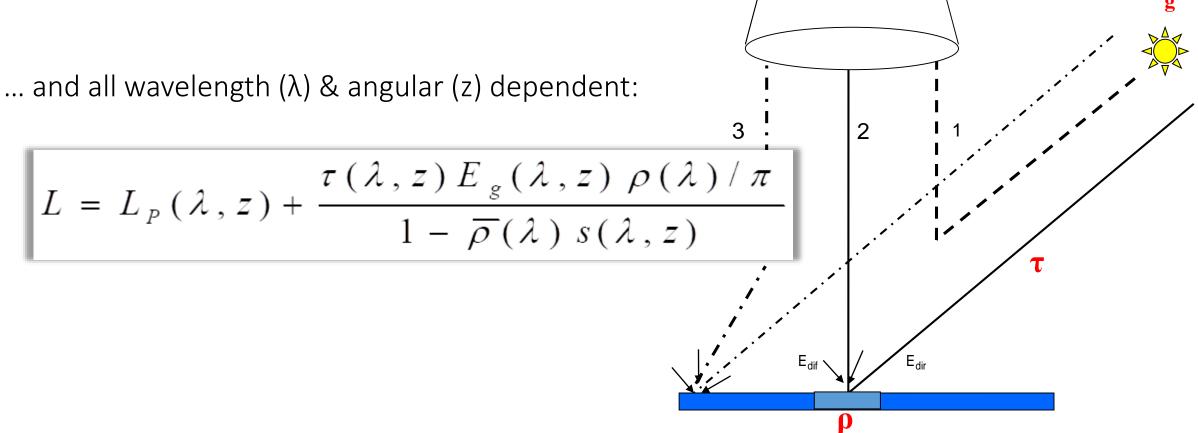
s: spherical albedo

(atm. backscattering to ground)



 E_g : global flux τ : total atm. transmittance ρ : target reflectance diffuse + direct





 E_g : global flux τ : total atm. transmittance ρ: target reflectance

 $L = c_0 + c_1 DN$

Beer's law

$$\tau(\lambda, x) = \exp\left[-\gamma(\lambda) x\right]$$

homogeneous path

 $\gamma(\lambda)$ Extinction coefficient (km⁻¹)

 \mathcal{X} Path length

$$\tau(\lambda, x) = \exp\left[-\int_{0}^{x} \gamma(\lambda, x') dx'\right]$$

inhomogeneous path

Beer's law

$$\tau(\lambda, x) = \exp\left[-\gamma(\lambda) x\right]$$

homogeneous path

$$\gamma(\lambda)$$
 Extinction coefficient (km⁻¹)

$$\xi = \gamma x$$
 Optical thickness

Beer's law

$$\tau(\lambda, x) = \exp\left[-\gamma(\lambda) x\right]$$

homogeneous path

$$\gamma(\lambda)$$
 Extinction coefficient (km⁻¹)

 \mathcal{X} Path length

K

σ

$$\gamma(\lambda) = k(\lambda) + \sigma(\lambda)$$



Beer's law

$$\tau(\lambda, x) = \exp\left[-\gamma(\lambda) x\right]$$

homogeneous path

 $\gamma(\lambda)$ Extinction coefficient (km⁻¹)

 \mathcal{X} Path length

Extinction = Absorption + Scattering

$$\gamma(\lambda) = k(\lambda) + \sigma(\lambda)$$

Separate: M = Molecule (Rayleigh) A = Aerosol

$$\gamma(\lambda) = k_M(\lambda) + k_A(\lambda) + \sigma_M(\lambda) + \sigma_A(\lambda)$$

Beer's law

$$\tau(\lambda, x) = \exp\left[-\gamma(\lambda) x\right]$$

homogeneous path

 $\gamma(\lambda)$ Extinction coefficient (km⁻¹)

 \mathcal{X} Path length

Extinction = Absorption + Scattering

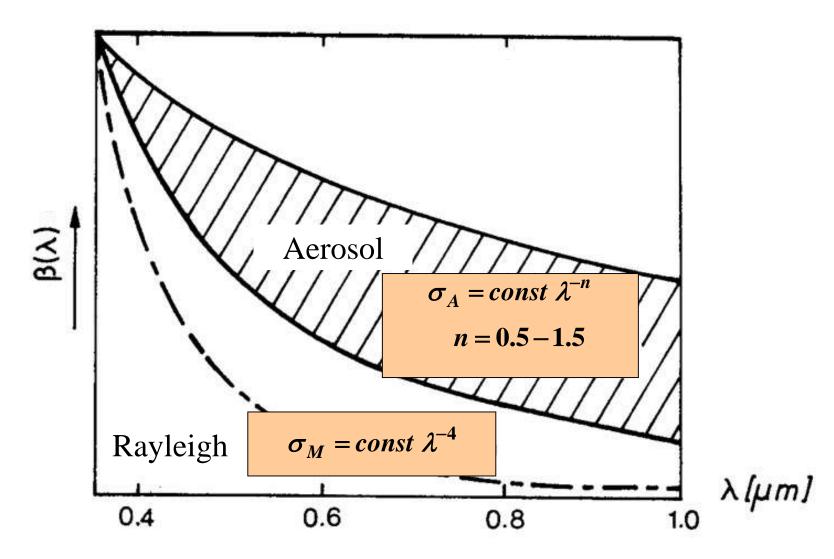
$$\gamma(\lambda) = k(\lambda) + \sigma(\lambda)$$

Single scattering albedo

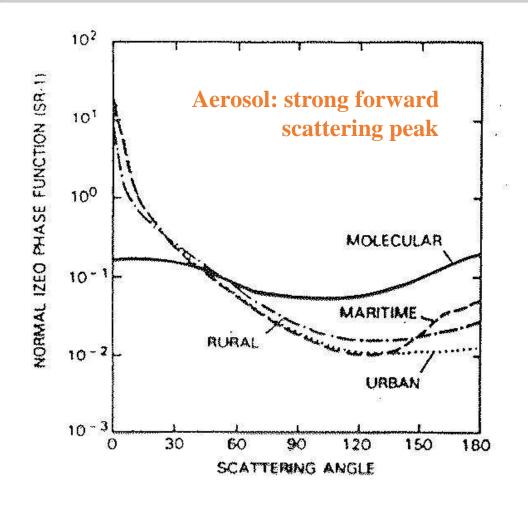
$$\omega(\lambda) = \sigma(\lambda)/\gamma(\lambda)$$



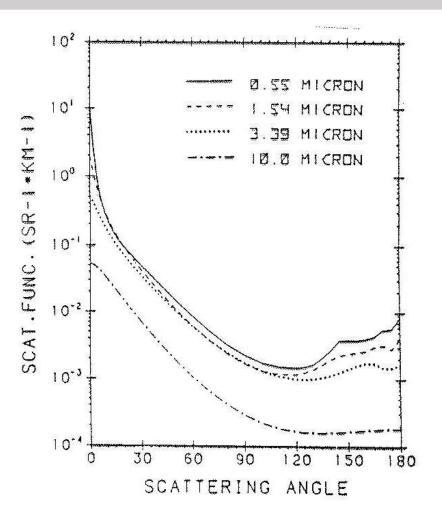
Scattering



Scattering



Rel. humidity = 70%, 550 nm



Maritime Aerosol, Rel. humidity=80%



Scattering phase function

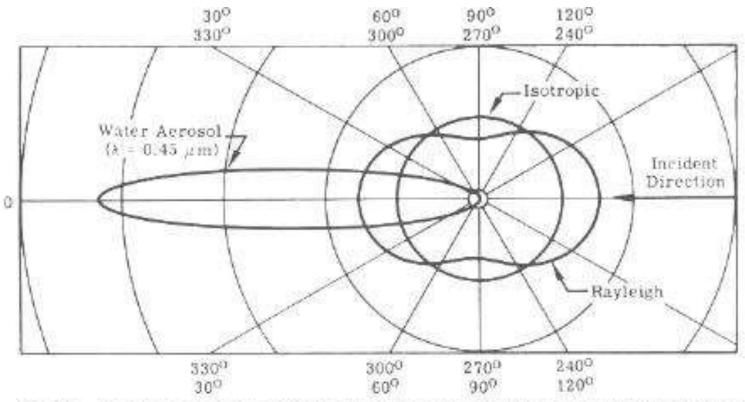


Fig. 8.2 Angular dependence of single-scattering phase functions in any azimuthal plants.
The isotropic and Rayleigh functions have been multiplied by 10. (After LaRocca and Turner, 1975.)



Visibility, meteorological range

Koschmieder Equation:

$$VIS = \frac{1}{\gamma} \ln \left(\frac{1}{\varepsilon} \right) = \frac{3.912}{\gamma}$$

Visibility, meteorological range

Contrast threshold $\varepsilon = 0.02$

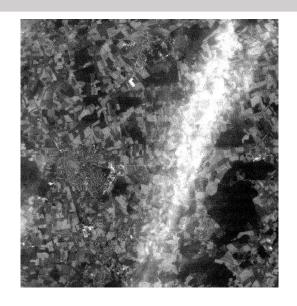
Extinction coefficient (γ at 0.55 μ m)

(depends on complex refractive index, particle size & distribution, concentration)

Wavenumber w (cm⁻¹) and wavelength λ (μ m): $w(cm^{-1}) = \frac{10^4}{\lambda(\mu m)}$



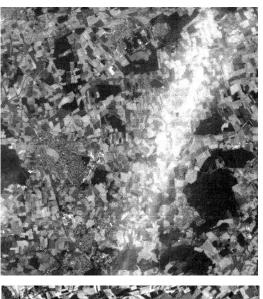
Scattering and Absorption

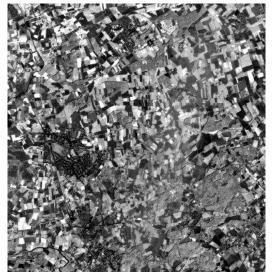


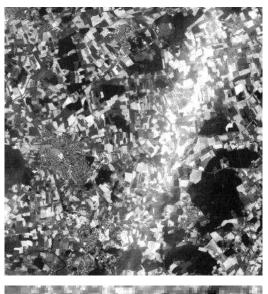
Top: Landsat TM bands 1-3 (480, 560, 660 nm)

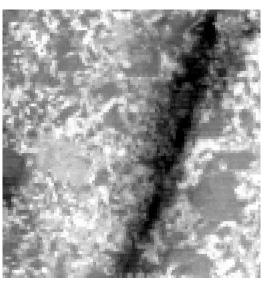
Bottom: TM bands 4, 6 (840 nm, 11.5 μm)

Note: TIR affected by WV, not by dust particles









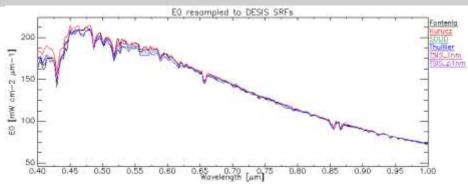


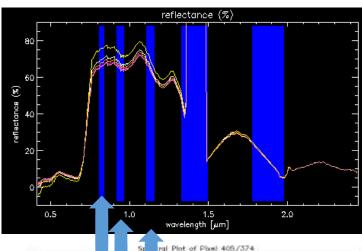
Therefore atm. correction requires

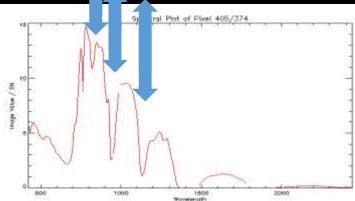
• Solar irradiance (E0)

Aerosol properties

Water vapor column





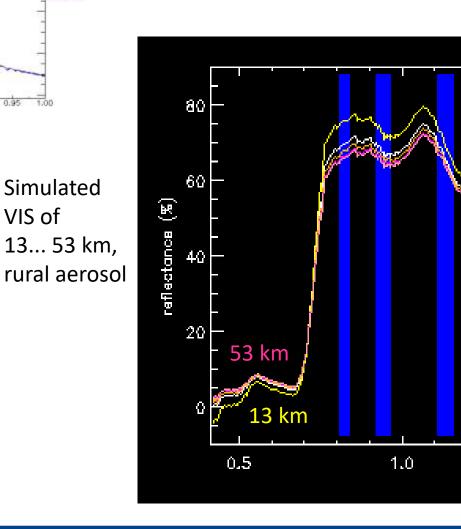


Different E0 models

Simulated

13... 53 km,

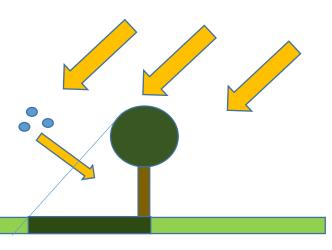
VIS of





• "Blue" shadows: diffuse illumination

Aerosols & molecules Rayleight





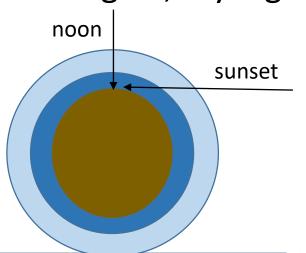
• "Red" sunset & sunrise: path lengths, rayleigh scattering & refraction

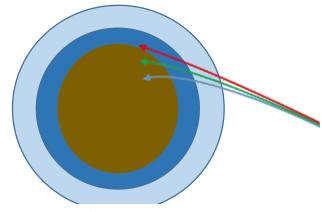
$$\sigma_M = const \lambda^{-4}$$
 Rayleigh

$$\tau(\lambda, x) = \exp[-\gamma(\lambda) x]$$
 Beer's law

 $\gamma(\lambda)$ Extinction coefficient (km⁻¹)

 \mathcal{X} Path length





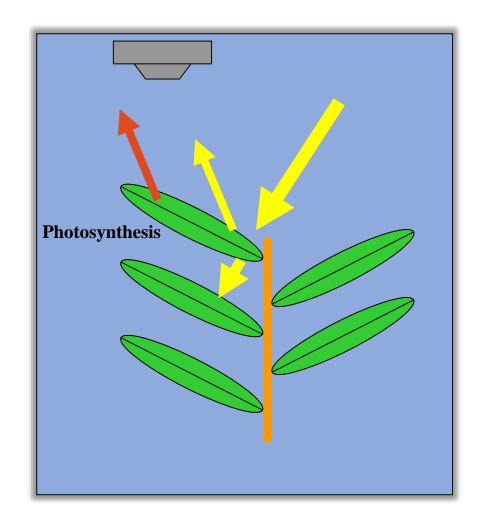


3rd: Interaction with Matter



EM interaction with media

- Absorption (uptake of energy)
 - Electron transfer, rotation, vibration
 - Heating, change in matter
- Emission ("release" of energy)
 - Electron transfer, rotation, vibration
- Reflection
 - Change of direction without energy uptake
- Transmission
 - Transfer without absorption or reflection





EM interaction with media

Energy balance relationship:

$$E_{\text{emitted by sun}} = E_{\text{reflected}} + E_{\text{transmitted}} + E_{\text{absorbed}}$$

E... Incident Energy [W]

$$1 = E_r / E_i + E_t / E_i + E_a / E_i$$

 $1 = R + T + A$

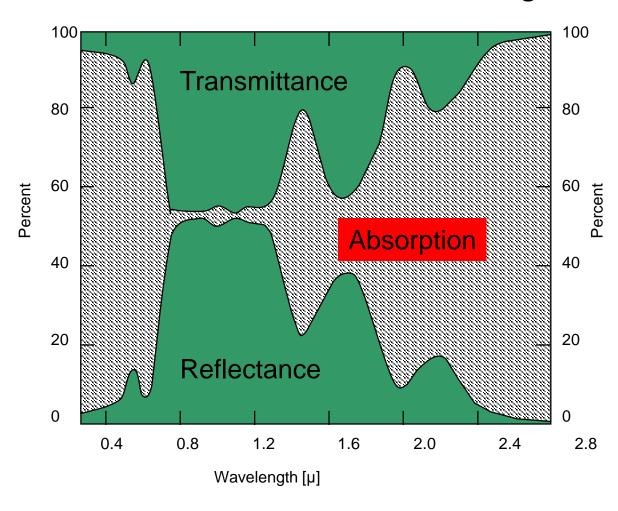
...Reflection-coefficient + Transmission-coefficient + Absorption-coefficient

• Opaque material: T = 0 = 1 = R + A

=> material property, independent of incoming radiant energy!

EM interaction with media

... and now as a function of wavelength:





Emission processes

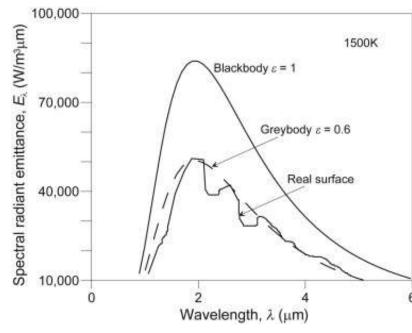
Spectral emissivity:

$$\varepsilon_{\lambda} = M_{\lambda} / M_{\lambda,black}$$

• And thus:

$$M = \varepsilon \sigma T^4$$

- Only for Black Bodies (KIRCHHOFF):
 Blackbody Emission = Absorption = 100%
- Grey Body: ε < 1
- Selective Absorber & Emitter



https://www.sciencedirect.com/topics/engineering/emissivity

=> emissivity depends on temperature and material



Electron transfer:

- Quantized: $\Delta Q = h * f$
- Occur normally < 1 μm
- Rarely detected due to mineral composition => complex overlay of absorption features

Vibration modes - bending & stretching:

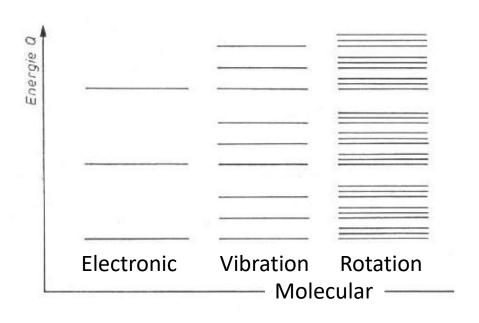
- "Sharper" and frequent (3*N possible for molecules with N atoms)
- Occur normally >1 μm
- Many fundamentals per material, normally $> 3 \mu m$
- Resonance (combinations & harmonics)

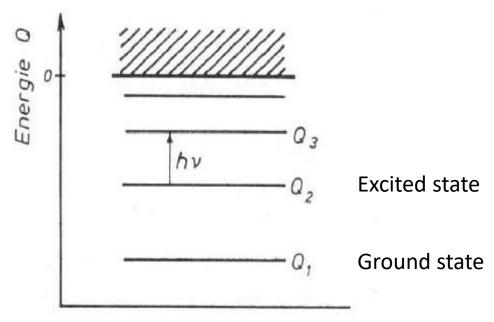
Fundamentals (Grundschwingung)

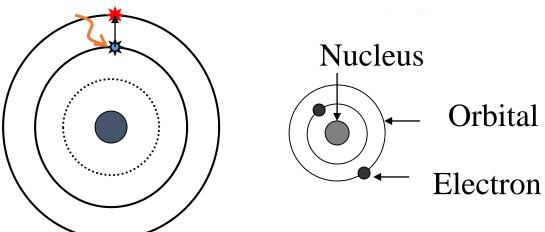
Harmonics (Obertöne): multiples of fundamentals, lower absorption depth 1st harmonic: 1/10 of fundamental absorption, 2nd harmonic: 1/100

Combinations: Sum of fundamentals and/or harmonics

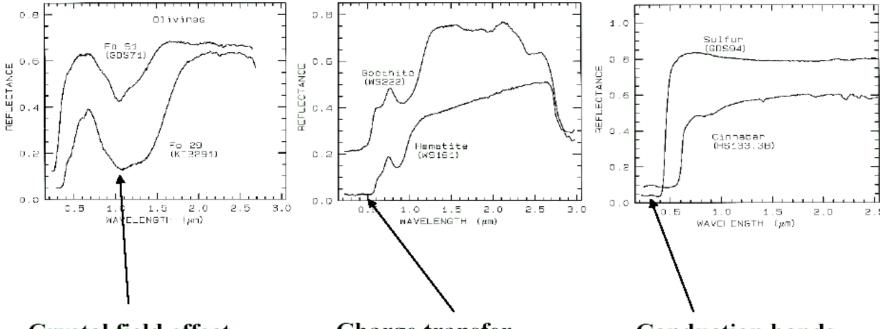












Crystal field effect absorption caused by Fe²⁺. Fe 29 has 53.65% FeO, Fe 91 has 7.93% FeO.

Charge transfer absorption caused by
Fe²⁺. Fe₂O₃ (hematite)
and FeOOH (geothite).

Conduction bands caused by S and HgS.

Example: Liquid Water H₂O

Fundamentals

```
- λ_1: 3.106 μm / 6.4 * 10<sup>-20</sup> J symmetric OH - stretch - λ_2: 6.080 μm / 3.3 * 10<sup>-20</sup> J HOH - bend - λ_3: 2.903 μm / 6.9 * 10<sup>-20</sup> J asymmetric OH - stretch
```

Harmonics

```
(2* f \text{ resp. } \% * \lambda)
- 1.553 \, \mu\text{m} / 1.3 * 10^{-19} \, \text{J}
- 3.040 \, \mu\text{m} / 6.6 * 10^{-20} \, \text{J}
- 1.452 \, \mu\text{m} / 1.4 * 10^{-19} \, \text{J}
```

Combinations

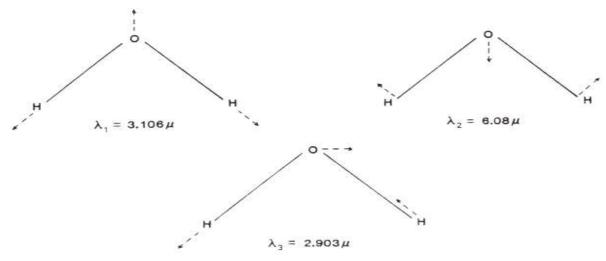


Fig.: ELACHI

Example: Water Vapor

Fundamentals

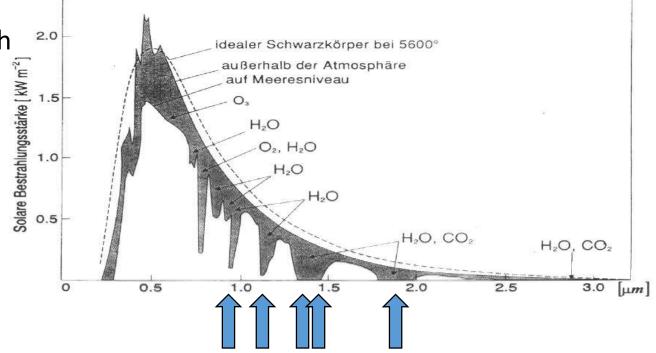
– vi @ 2.74 μm symmetric stretch

– v2 @ 6.27 μm HOH - bending

– v3 @ 2.66 μm asymmetric stretch

Harmonics and Combinations

- $-1.875 \mu m (v2 + v3)$
- $-1.454 \mu m (2*v2 + v3)$
- $-1.380 \mu m (v1 + v3)$
- $-1.135 \mu m (v1 + v2 + v3)$
- $-0.942 \mu m (2*v1 + v3)$



Absorption & transmission processes

Beer-Lambert law (base for transmittance spectroscopy):

• Transmittance $t = I / I_0$

• A = $\log_{10}(I_0/I) = \epsilon c I$

A: absorbance

ε: molecular absorption coefficient for molecule species [mol⁻¹ cm⁻¹]

c: concentration of absorbing molecules [mol]

I: path length through sample [cm]

=> absorbance ~ molecule concentration

c * molecule

• Also valid for reflectances: $A = Log_{10} (1/\rho)$



- Sounds pretty simple, **but**:
 - Some features are related to many materials:
 e.g. OH bond is related to water, cellulose, starch, clays etc.
 - Overlapping features
 - Multiple scattering broadens features
 - No simple relationship between feature depth and component content (saturation effects)
 - In Vivo ≠ In Vitro
 - Noise, spectral bandwidth and position, ...

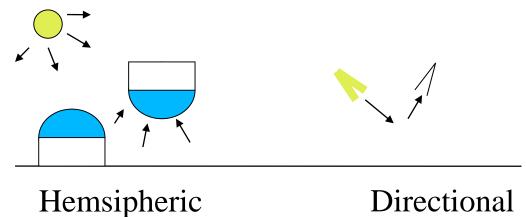


...more on reflectance behavior



Measurement principles

- Bidirectional pointed illumination, observation in one direction
- Directional hemispheric pointed illumination, observation of hemisphere
- Hemispheric directional hemispherical illumination, observation in one direction
- Hemispheric hemispheric hemispherical illumination, observation of hemisphere



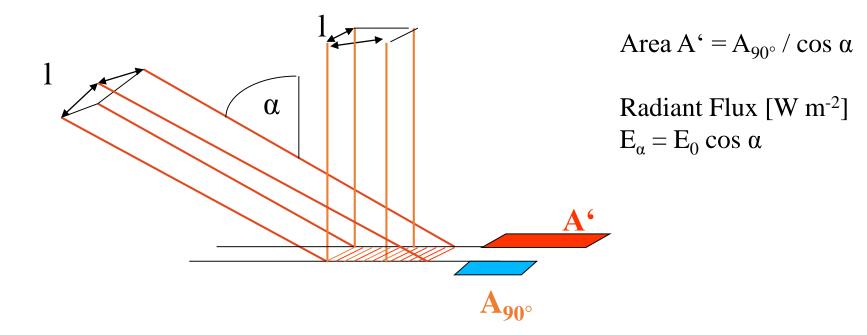


Measurement principles

Incoming/Reflected	Directional	Conical	Hemispherical
Directional	Bidirectional Case 1	Directional-conical Case 2	Directional-hemispherical Case 3
Conical	Conical-directional Case 4	Biconical Case 5	Conical-hemispherical Case 6
	Saso 1		
Hemispherical	Hemispherical-directional Case 7	Hemispherical-conical Case 8	Bihemispherical Case 9

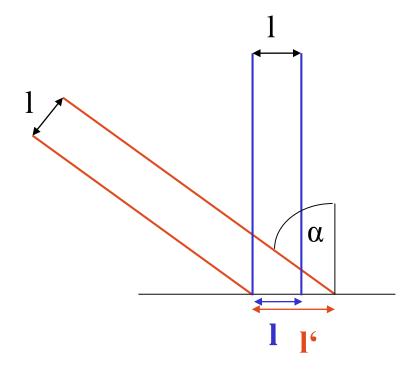
Schaepman-Strub, RSE 2006

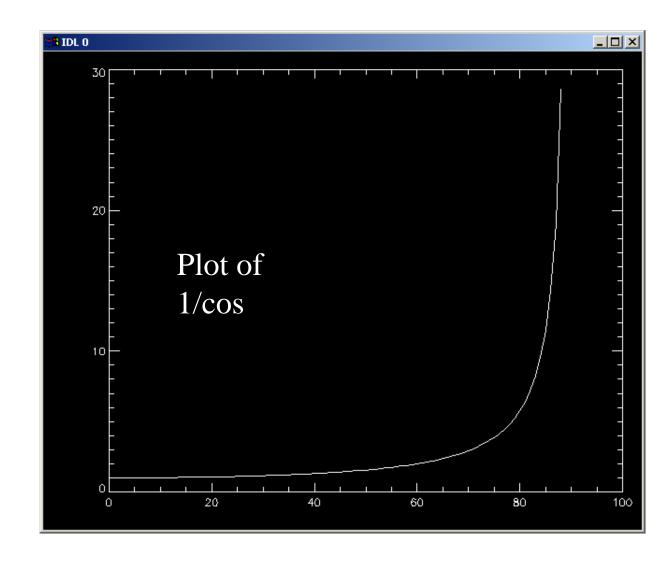
Lambertian Cos.-Law





Lambertian Cos.-Law

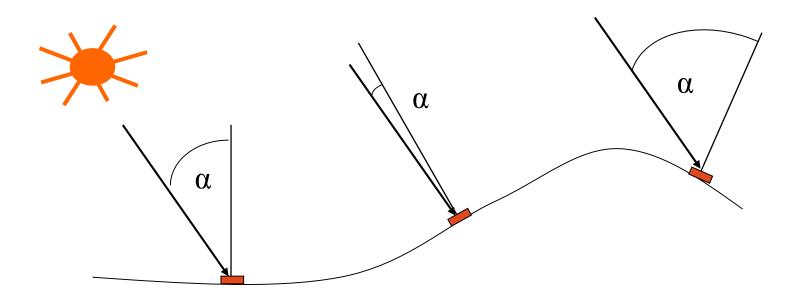




Radiant Flux [W m⁻²] $E_{\alpha} = E_0 \cos \alpha$

$$E_{\alpha} = E_0 \cos \alpha$$

=> Illumination condition cased by terrain & sun angle





Reflection behaviour of surfaces

Specular

angle of incidence = angle of reflection

Diffuse

same brightness in all viewing directions

if $\sigma h < \lambda / (8 * \cos \Theta)$

 Θ ...view angle

 σ h: standard deviation of surface roughness

=> "optical" reflection is normally diffuse

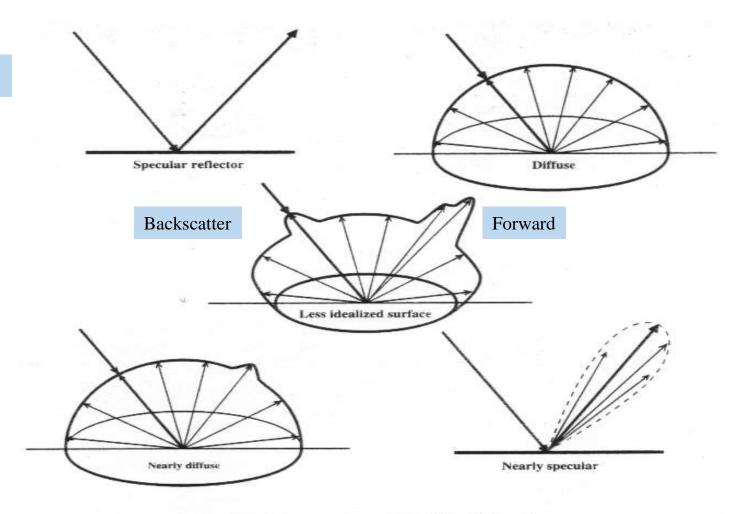
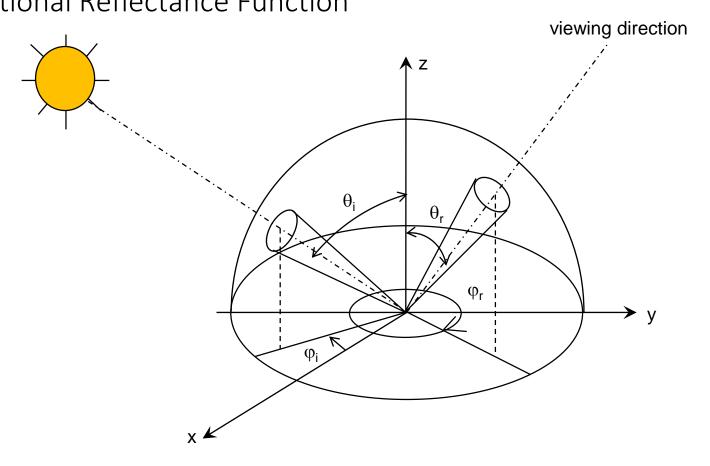


Figure 4.7 Reflectance characteristics of idealized surfaces.



Measurement principles

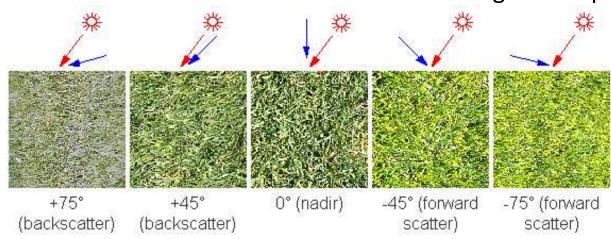
BRDF
Bi-Directional Reflectance Function

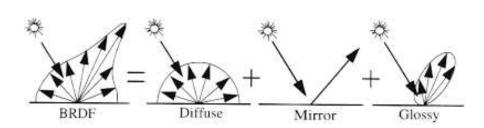




Measurement principles

- BRDF $L_{\lambda} = \rho_{\lambda} E_{\lambda}$ (cos Θ) f_{r} $f_{r} =$ bidirectional distribution function for Lambertian surfaces.: $\rho_{\lambda} = (L \lambda \pi) / (E \lambda \cos \Theta)$
- Shade due to surface roughness
- HotSpot-effect:
 If view direction = illumination direction => no shade visible
 => brighter & specular => no Lambertian





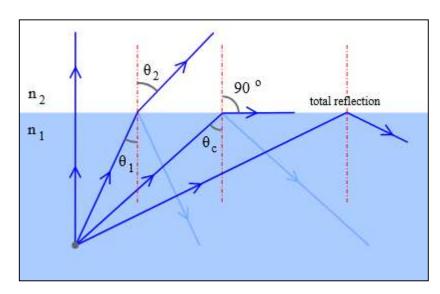
mpi-inf.mpg.de

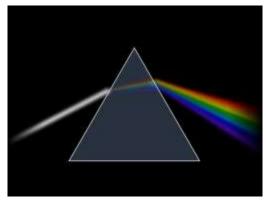
Fig. 2: Bidirectional reflectance effect on a grass lawn, observed under different viewing angles from a FIGOS mounted camera in the solar principal plane. Solar zenith angle is 35°, indicated with red arrows. The view directions are given in blue. The camera is operated in the manual modus keeping aperture, exposure time and focal length constant (k=16, t=1/15, f=135mm).



Other relevant topics:

- Refraction: bending of light at transit from 2 materials with different optical thickness
- Special case: total reflection (from thicker to thinner, e.g. water -> air)
- Prism: bending index changes with wavelength (higher for smaller wavelength)
 - Spectrometer either use prism (e.g., EnMAP) or diffraction grating (e.g., DESIS)





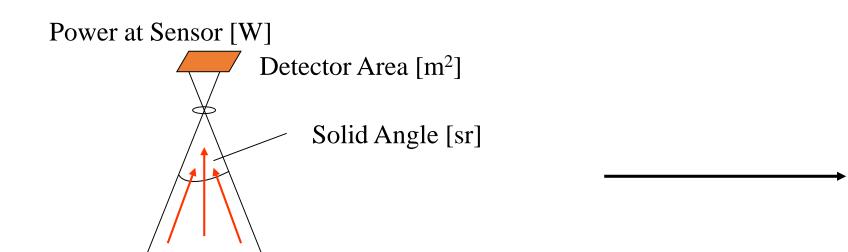


5th: Sensor



Sensor Model

- Input: At-sensor radiance ("Amount of light")
 - Sensor
- Output: DN (Digital Number)



10	23	0	
56	23	23	
22	43	5	
11		17	
			-

DLR

Sensor Model

- Input: At-sensor radiance
 - (Scanner-Mechanics)
 - Optics inclusive Filter & Dispersion
 - Detectors
 - Electronics
 - A/D
 - Platform Motion & Attitude
- Output: DN



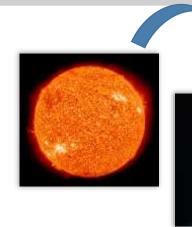
6th: Digital Image Processing



Image Chain

"Way of a photon

from the sun









through atmosphere

to earth

to **EO** sensor

to image"





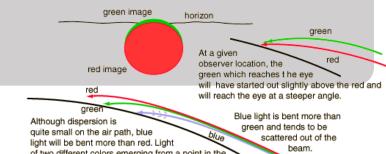
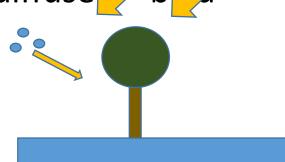


Abbildung Schatten => diffuse b

Aerosols & Molecules Rayleight



FUTURE

• "Red" sunset & sunrise: path lengths, rayleigh scattering & refraction

$$\sigma_M = const \lambda^{-4}$$

$$\tau(\lambda, x) = \exp\left[-\gamma(\lambda) x\right]$$

 $\gamma(\lambda)$ Extinction coefficient (km⁻¹)

X Path length

Beer's law

