

Investigation of the Ionian basin and Calabrian arc lithosphere.

“Potrebbe essere giusto o sbagliato: lo saprai solo se avrai sbagliato.”

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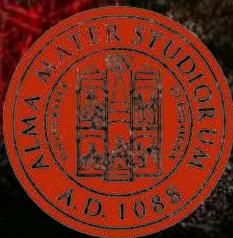
External:

Prof. Christine Thomas (UniMü)

Dr. Rafael Abreu (IPGP).



SPECFEM 3D
Cartesian



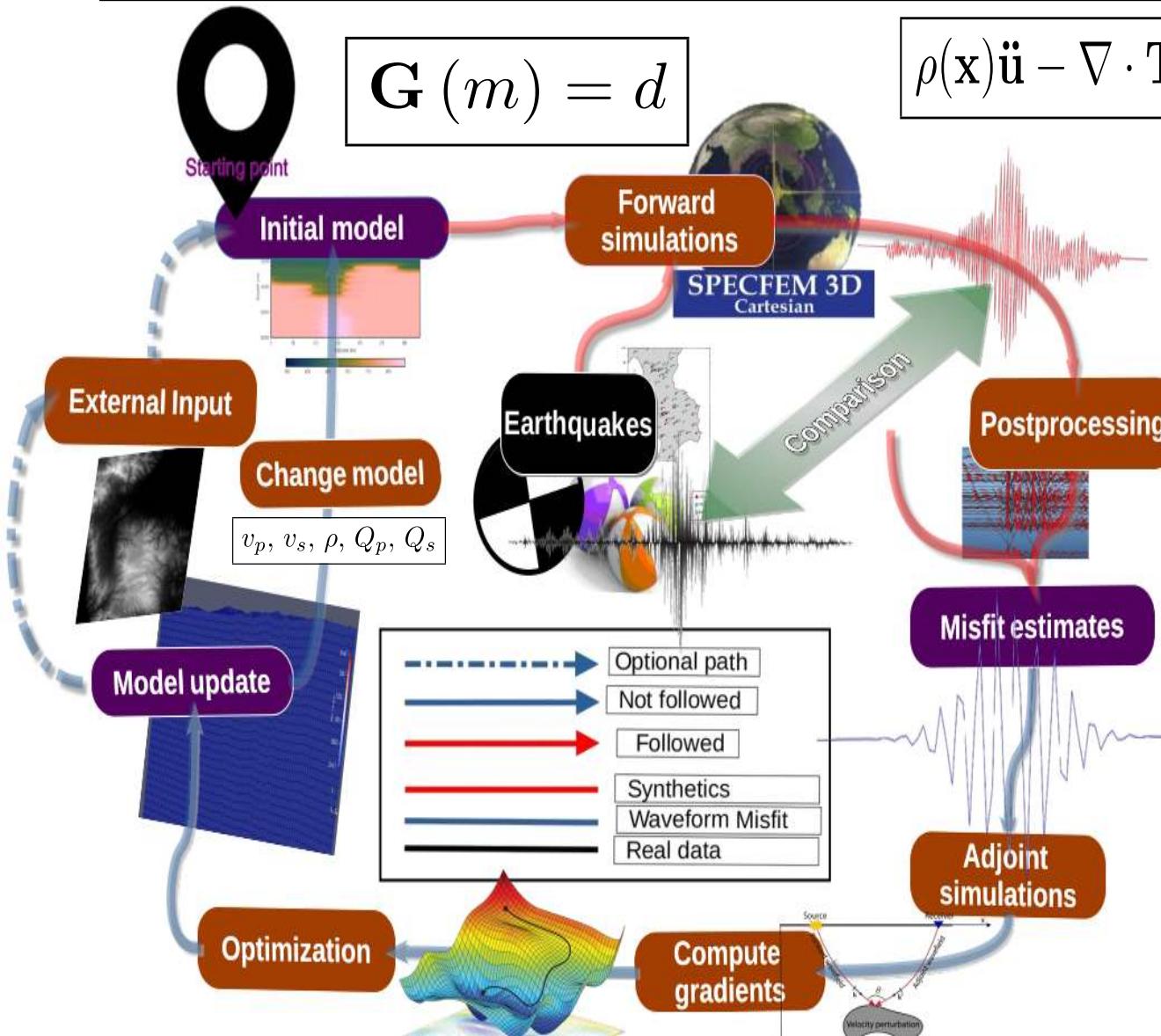
Academic Year 2024/2025

DIPARTIMENTO
DI FISICA



*In collaboration with the Seismology research group at the
University of Münster for a three months
Erasmus traineeship (April - June).
Candidate: Francesco De Rose 256865.*

SEM (Spectral-Element Method) and Full Waveform Inversion (FWI) overview.



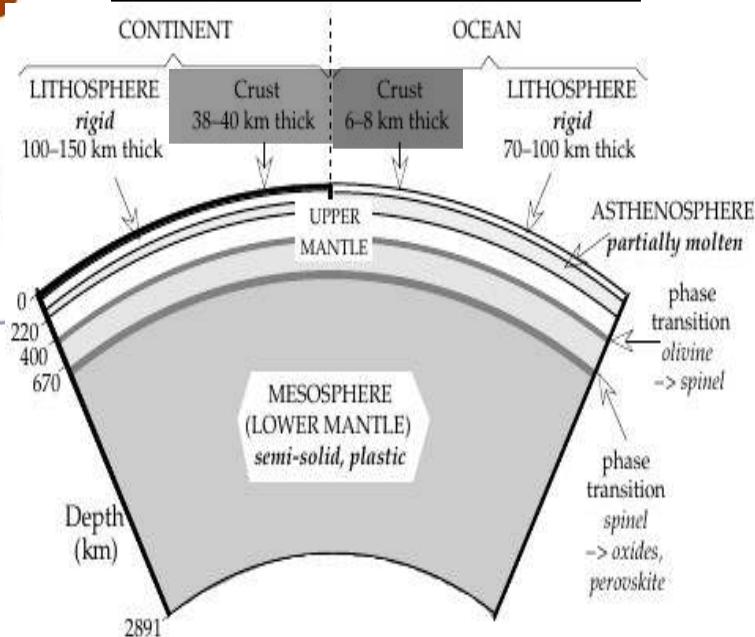
$$\rho(\mathbf{x})\ddot{\mathbf{u}} - \nabla \cdot \mathbf{T}(\mathbf{x}, t) + \mathcal{V} = \mathbf{F}(\mathbf{x}_s, t)$$

$$\hat{\mathbf{n}} \cdot \mathbf{T} = 0$$

$$\mathbf{T}(\mathbf{x}, t) = \mathbf{C}(\mathbf{x}) : \nabla \mathbf{u}(\mathbf{x}, t)$$

$$\mathbf{F}(\mathbf{x}, t) = -\mathcal{S}(t) \mathbf{M} \cdot \nabla \delta(\mathbf{x} - \mathbf{x}_s)$$

$$\mathbf{u}(\mathbf{x}, 0) = \mathbf{0}, \quad \dot{\mathbf{u}}(\mathbf{x}, 0) = \mathbf{0}$$

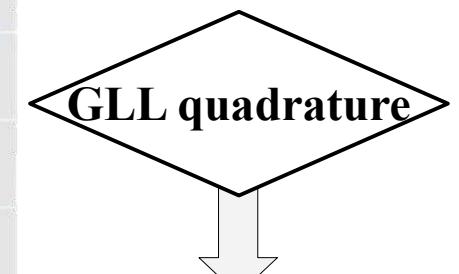
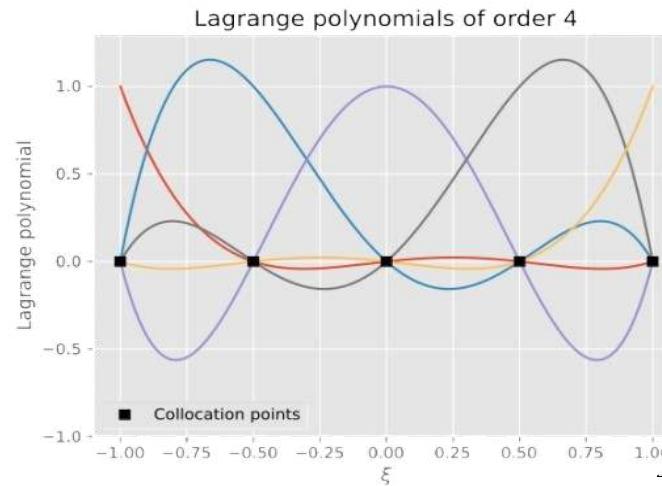
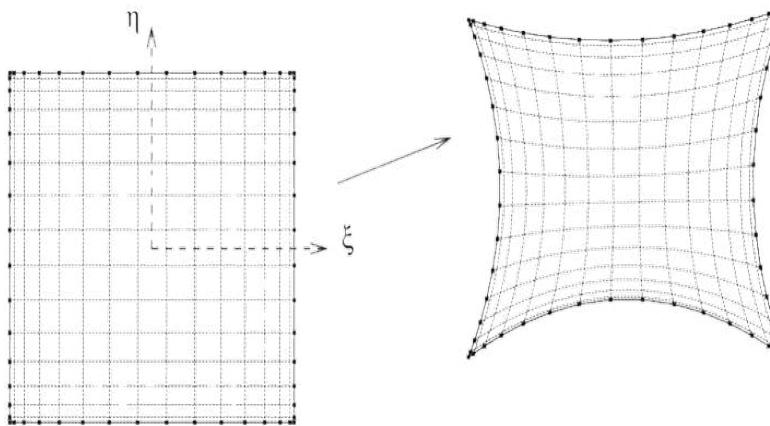


[1] Lowrie 2003

Spatial discretization² and weak formulation of the 3D wave equation.

$$\int_{\Omega} [\mathbf{w} \cdot \mathbf{T} - \mathbf{w} \cdot \mathbf{C} : \nabla \mathbf{u}] d^3x = \int_{\Omega} \mathbf{M} : \nabla \mathbf{w}(\mathbf{x}_s) \mathcal{S}(t) d^3x.$$

Goal: finding the displacement field \mathbf{u} s.t. \forall test function \mathbf{w} this equation is satisfied.



$$u(\mathbf{x}(\xi, \eta, \zeta)) \approx \sum_{i,j,k=0}^{N_i N_j, N_k} u_i(\xi) \ell_j(\eta) \ell_k(\zeta).$$

$$\ddot{\mathbf{u}}(t) = \mathcal{M}^{-1} [\mathcal{F}(t) - \mathcal{K} \mathbf{u}(t)]$$

Meshing strictly with conforming/non conforming hexahedral elements!
Major discontinuities practically feasible though.

Time discretization: Newmark scheme.

Advantages

- Curved finite elements are mapped to unit cube, described by 8 or 27 control points.
- Efficient MPI parallelization.
- ASCII based, serial I/O, ADIOS I/O, HDF5, GPU compatible.

$$\begin{aligned}
 \mathbf{u}_{n+1} &= \mathbf{u}_n + \Delta t \mathbf{v}_n + \frac{\Delta t^2}{2} \mathbf{a}_n \\
 \mathbf{v}_{n+1/2} &= \mathbf{v}_n + \frac{\Delta t}{2} \mathbf{a}_n \\
 \mathbf{a}_{n+1} &= \mathbf{M}^{-1} (\mathbf{F}_{n+1} - \mathbf{C}\mathbf{v}_{n+1/2} - \mathbf{K}\mathbf{u}_{n+1}) \\
 \mathbf{v}_{n+1} &= \mathbf{v}_{n+1/2} + \frac{\Delta t}{2} \mathbf{a}_{n+1}.
 \end{aligned}$$

CMB matrix term.

Disadvantages

- The finer the mesh the stricter the timestep upper value!
- Ill-shaped elements increase the solving time or make it computationally infeasible.
- No stability studies up until now for unstructured heterogeneous medium.

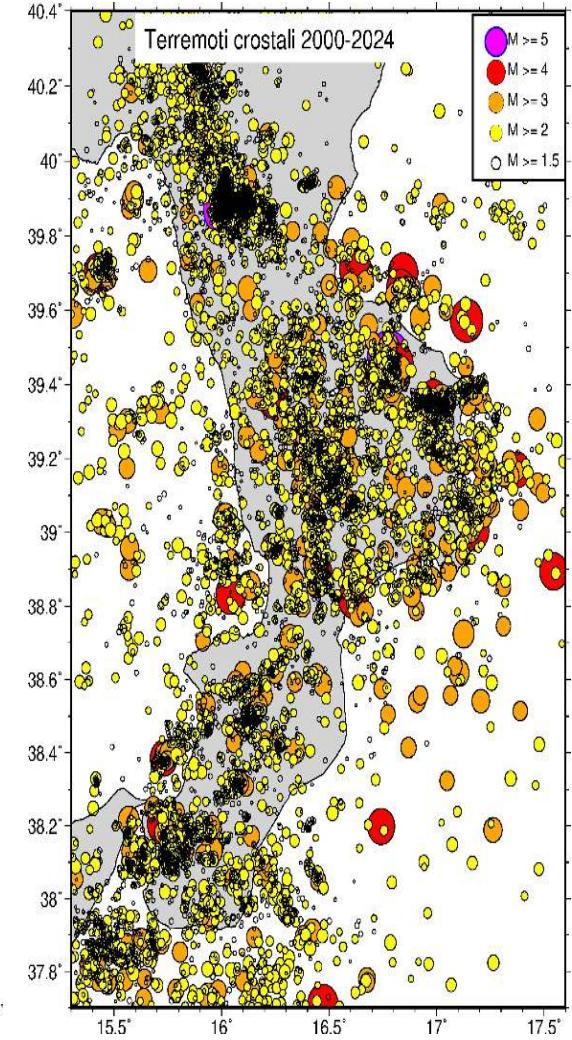
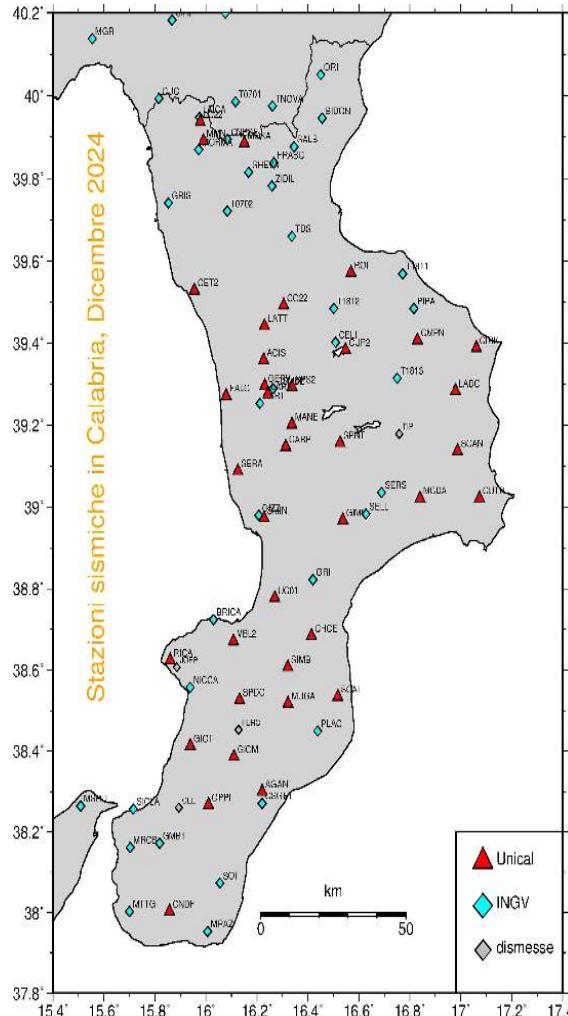
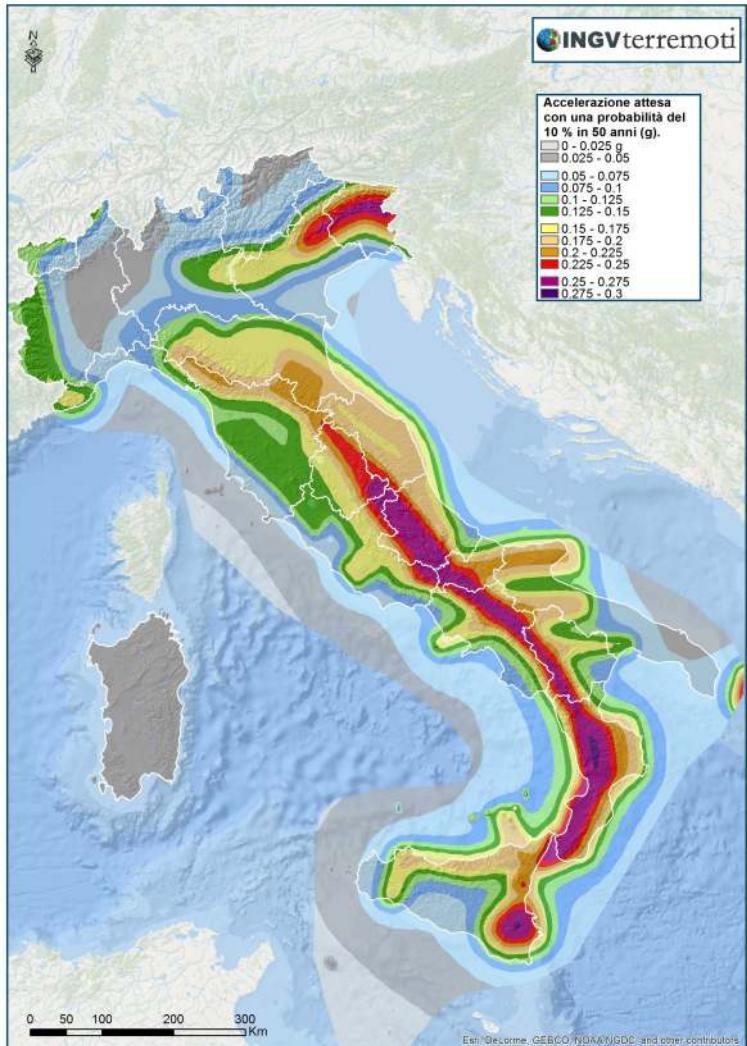
$$h = \frac{N_{GLL} V_{S,min}}{5 f_{max}}.$$

$$C = \Delta t \left(\frac{v_{s,min}}{h_{min}} \right).$$

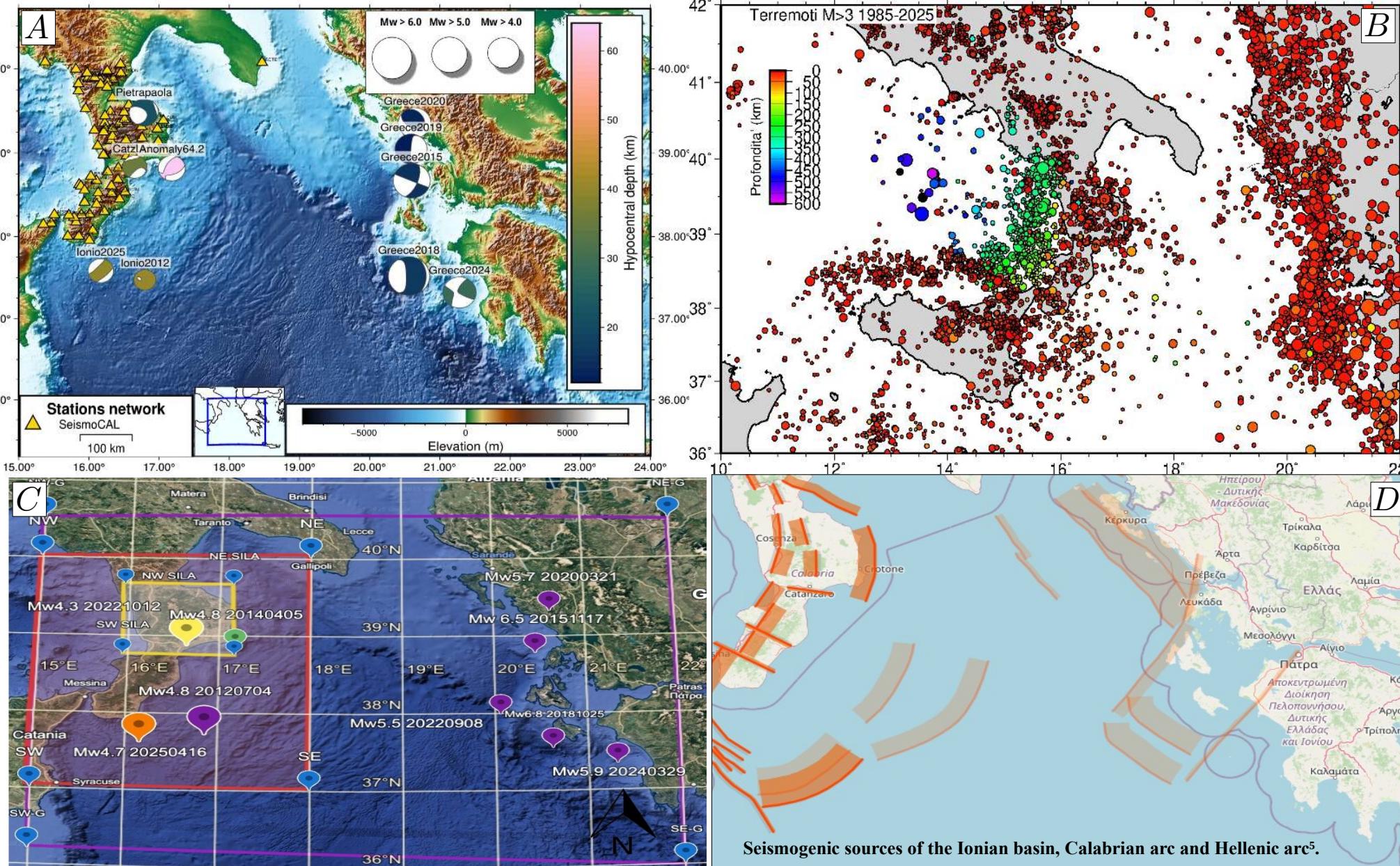
Benchmarks test (regional models)³ C ~ 0.5.

Frequency goal 0.3 Hz.

The importance of the lithospheric structure of the Calabrian arc and Ionian basin⁴.

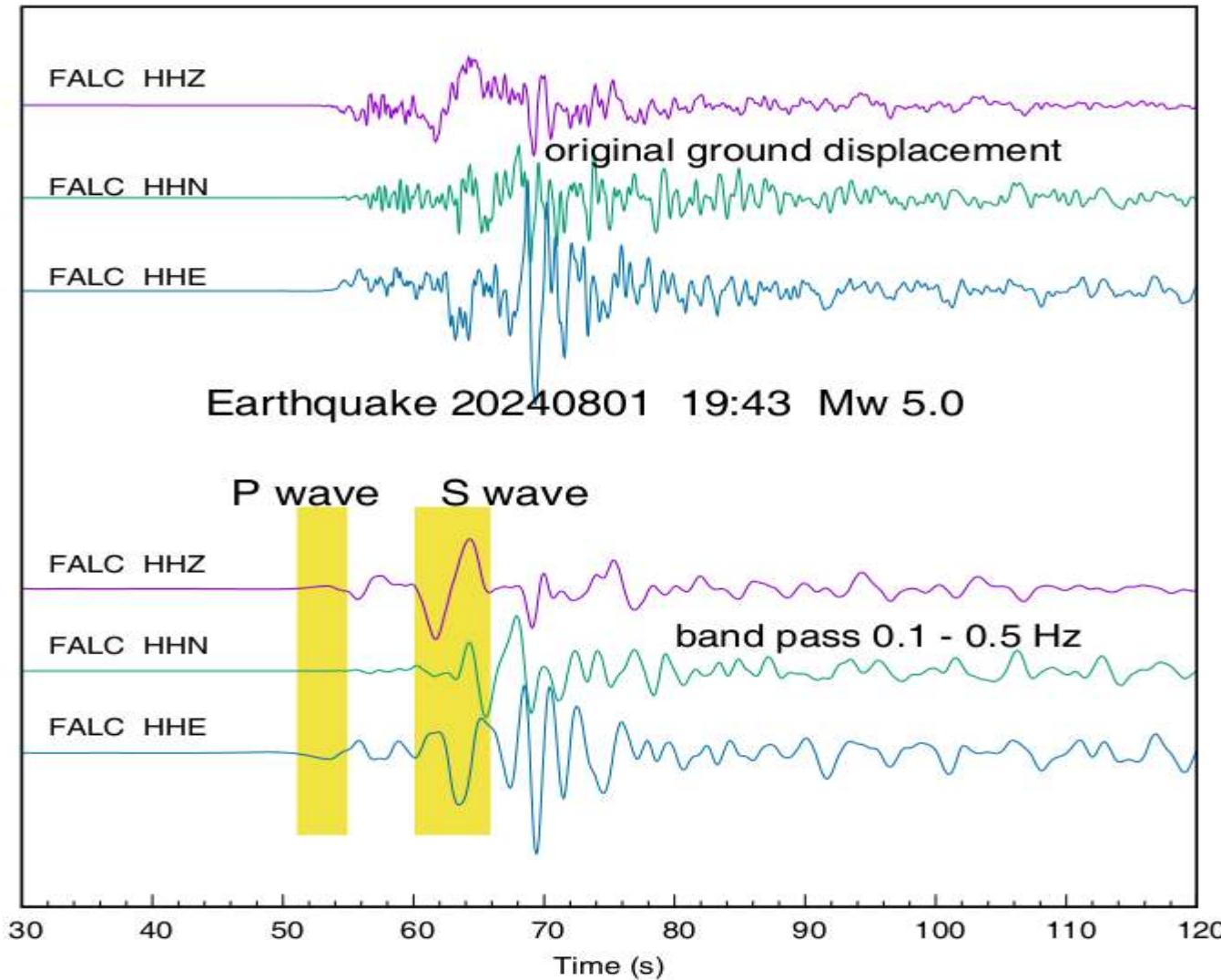


Università della Calabria
Rete Sismica



[5] INGV-DISS source

Observed seismograms contains contribution from the source and from the propagation of seismic waves throughout an highly heterogeneous medium.



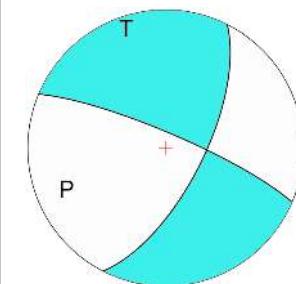
MT components

20240801.194320.0

$$\begin{bmatrix} M_{rr} & M_{rt} & M_{rp} \\ M_{rt} & M_{tt} & M_{tp} \\ M_{rp} & M_{tp} & M_{pp} \end{bmatrix}$$

Values

in	10^{23}	dyne	cm
-1.91	2.170	-3.840	
2.170	4.04	2.410	
-3.840	2.410	-2.13	



EQ20240801 19:43:20 UTC

Mw5.0

$$M_0 = 5.62 \cdot 10^{16} \text{ N}\cdot\text{m}$$

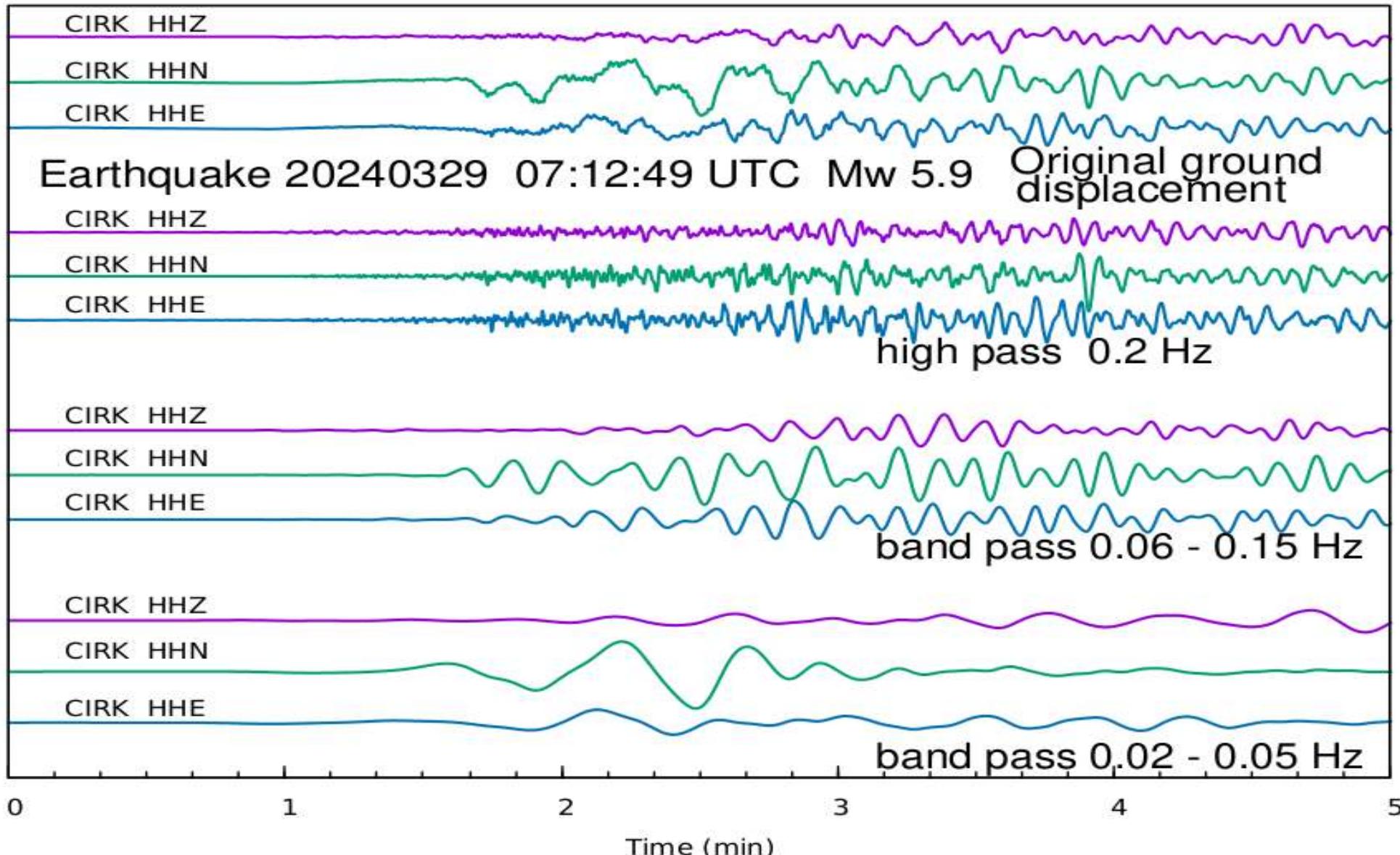
Lat 39.4°N

Long 16.75°E

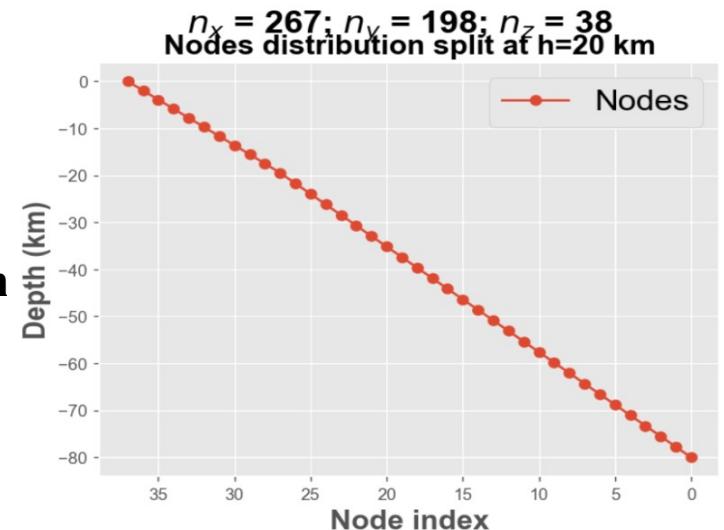
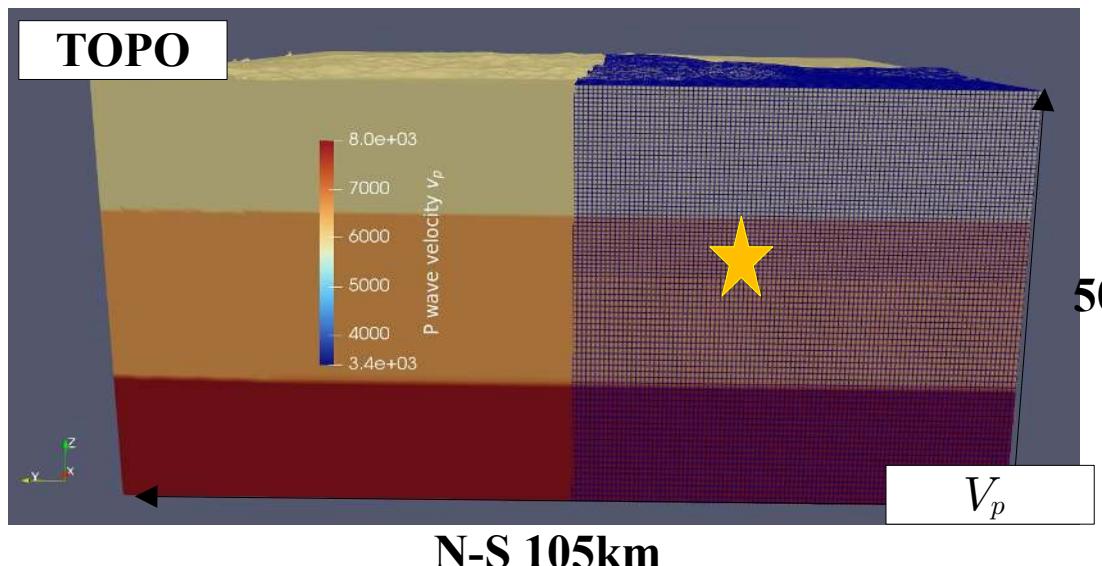
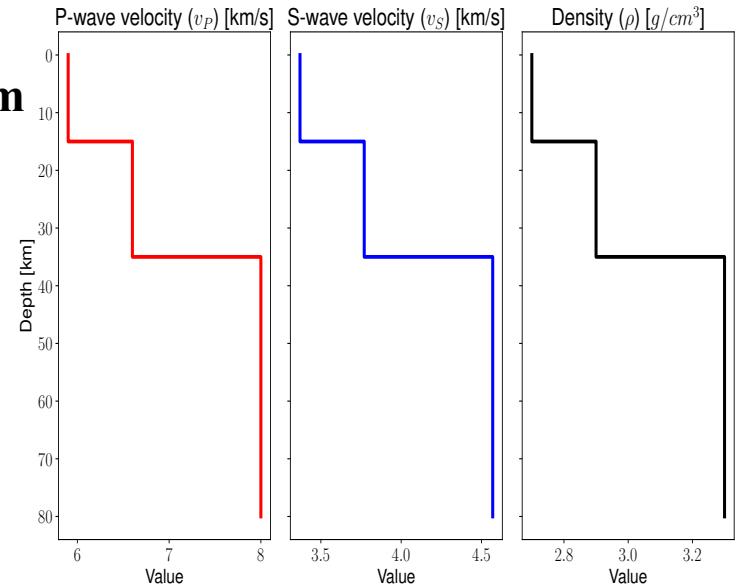
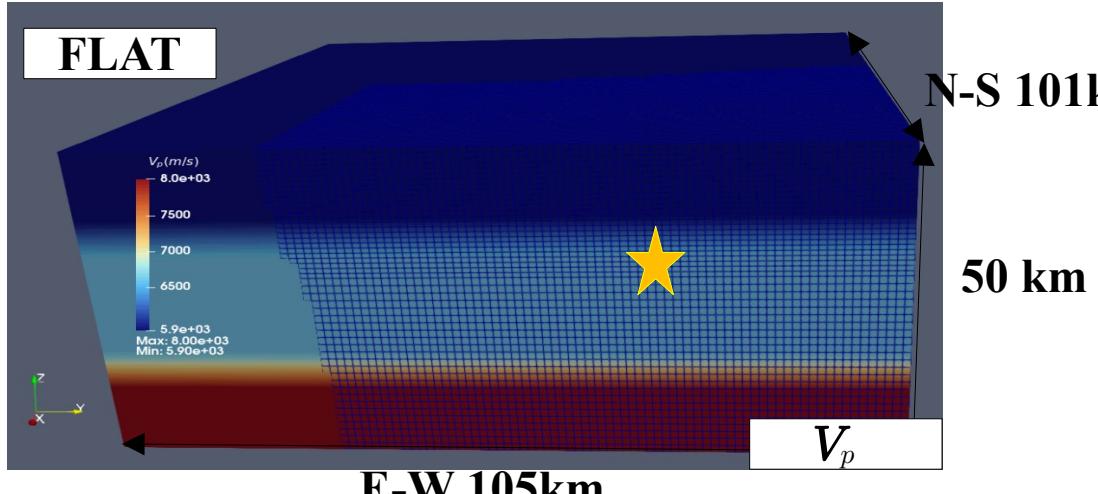
26.3km deep

2kmSE PIPA.

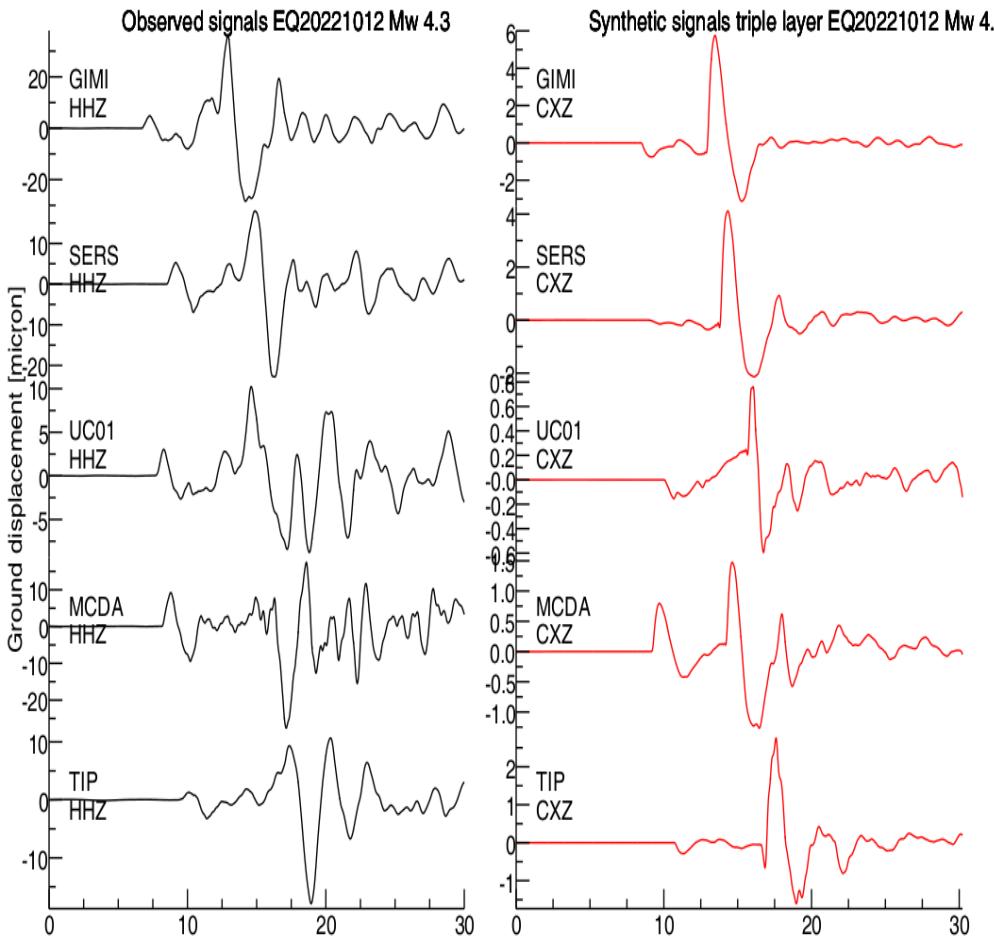
Modeling real seismograms in time domain is challenging.



No topography Sila region test – three homogeneous layers model.



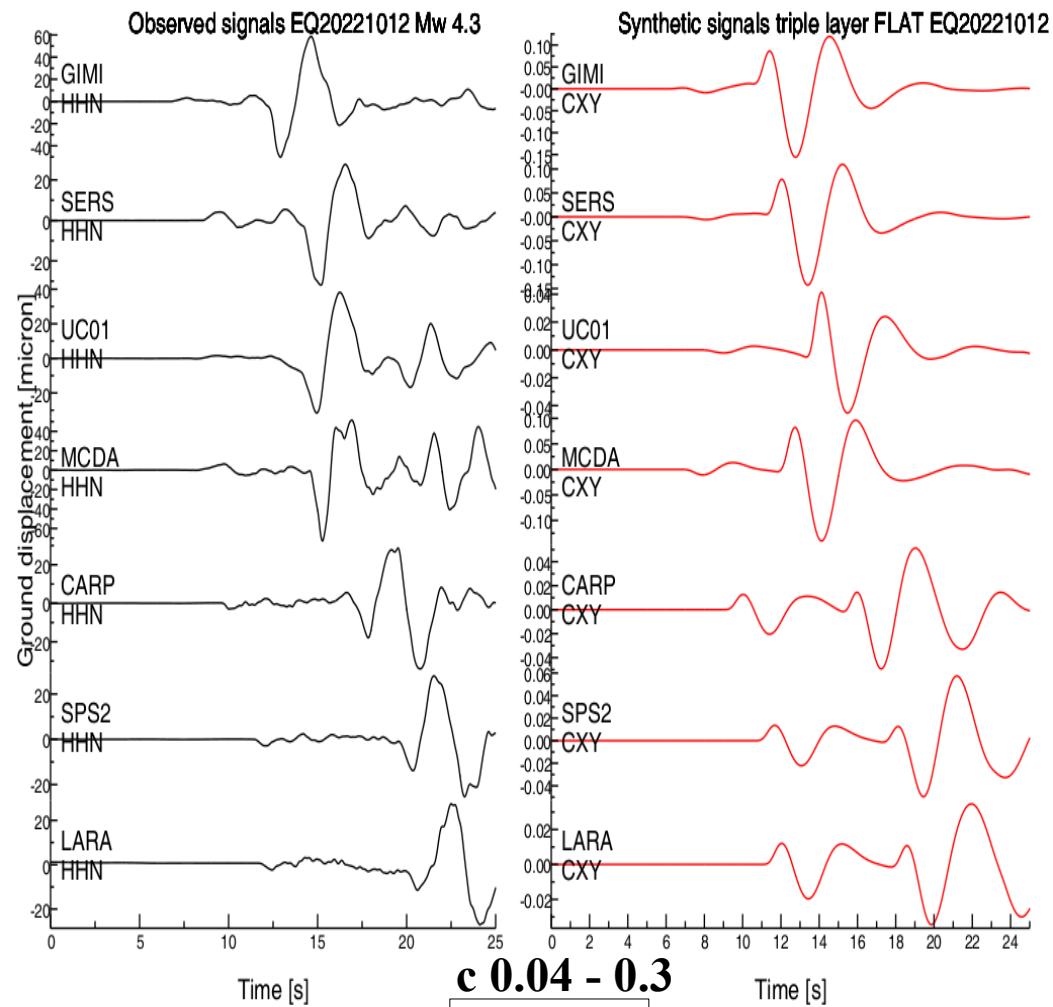
Comparison of real data and synthetics with and without topography.



TOPO

20221012 Mw4.3

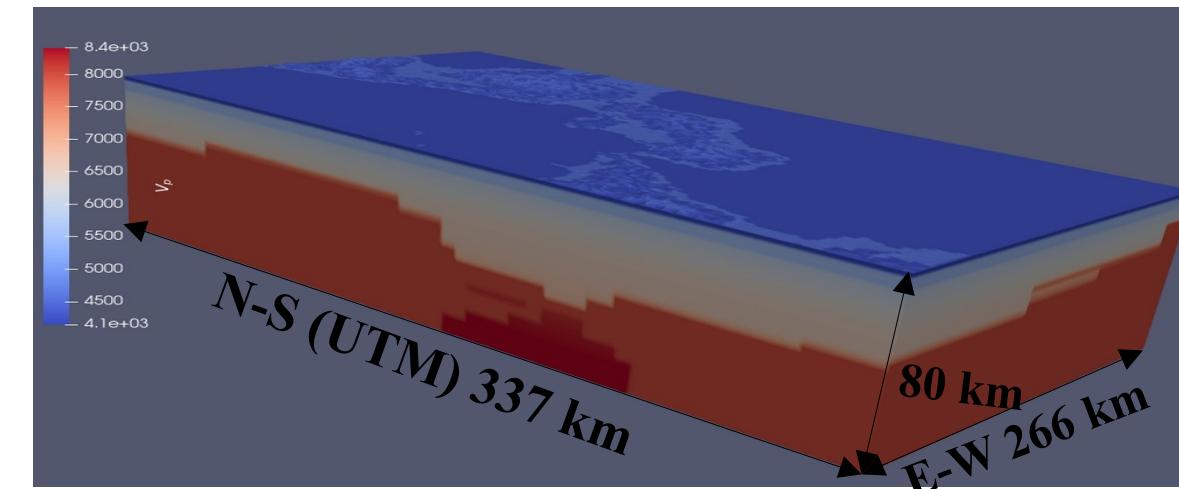
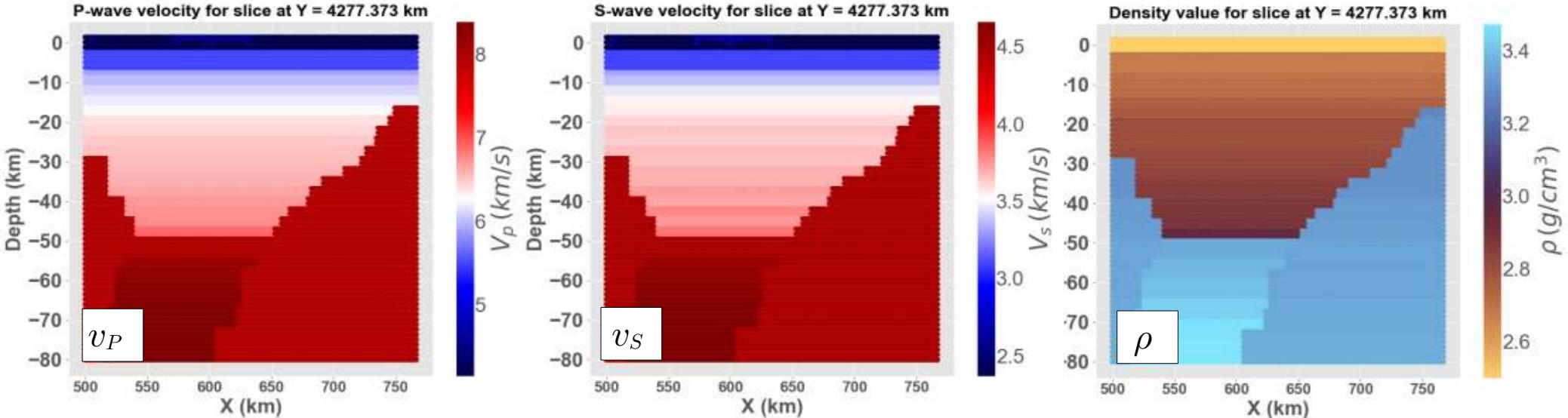
(38.83°N, 16.65°E) D35km 7kmSE CZ.



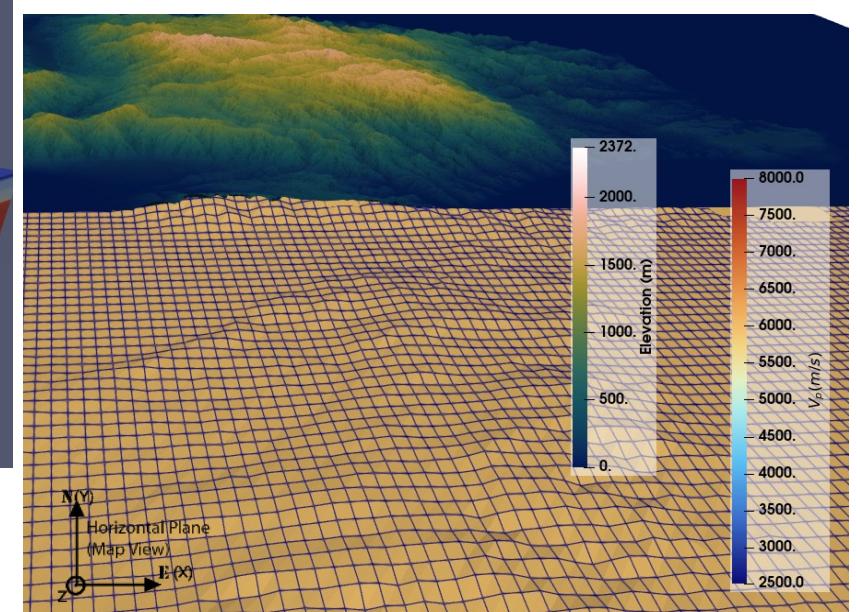
FLAT



Simulations throughout the Calabrian arc - ICLARC 3D model.

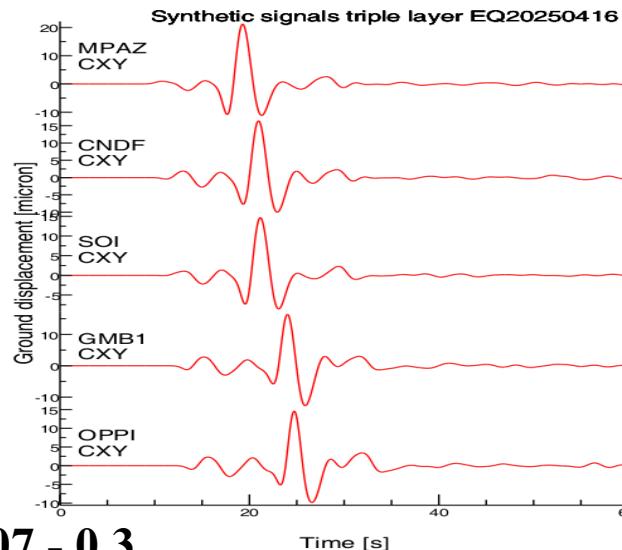
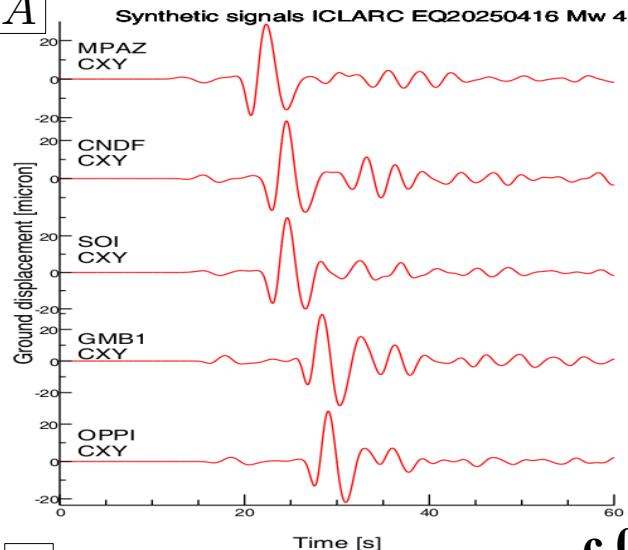


Ionian CaLabrian ARC – ICLARC.
Total number of elements 2.49 million.
Element's spacing $\sim 1.7 \text{ km}$. Nodes split at 20km.

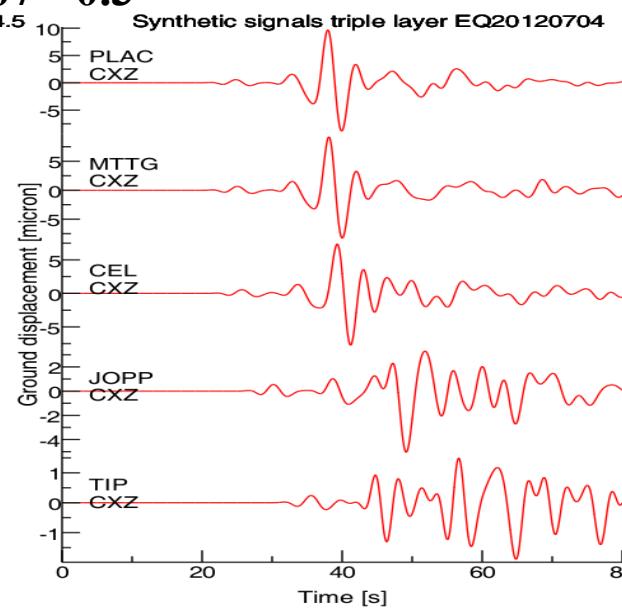
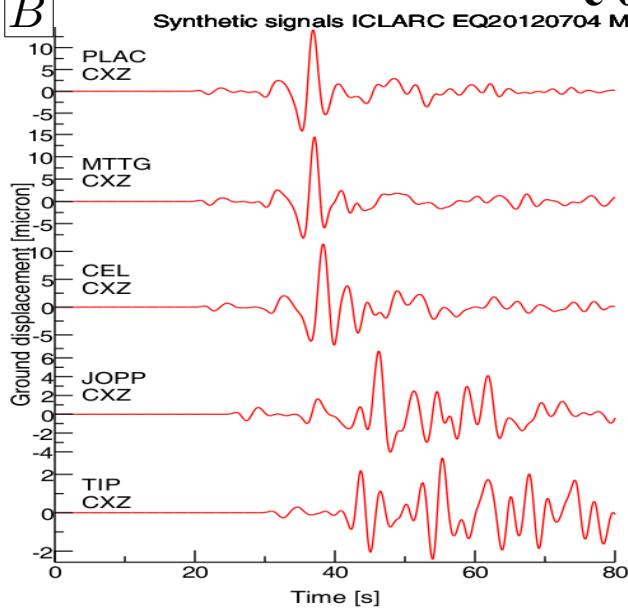


Comparison of synthetics - ICLARC vs 1D model.

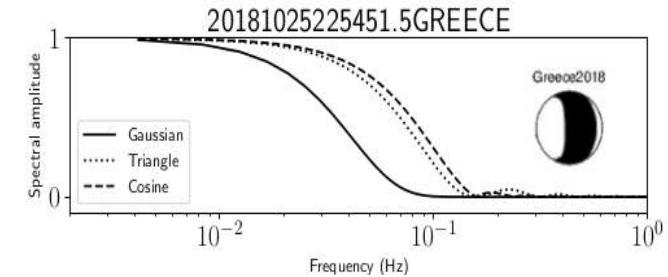
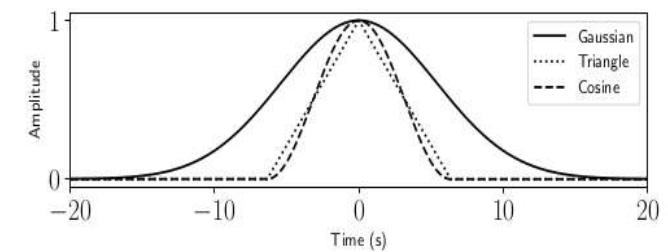
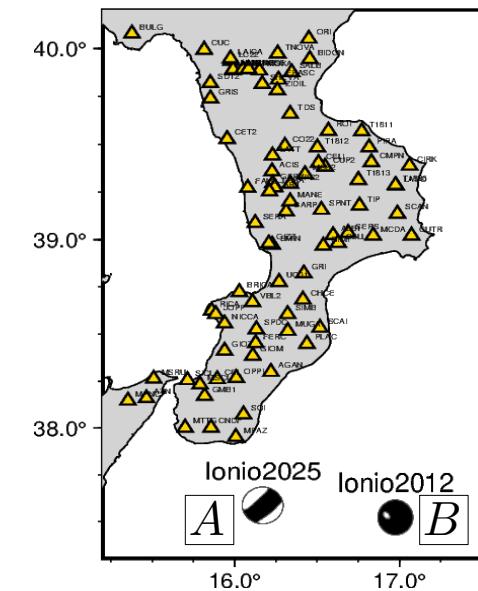
A



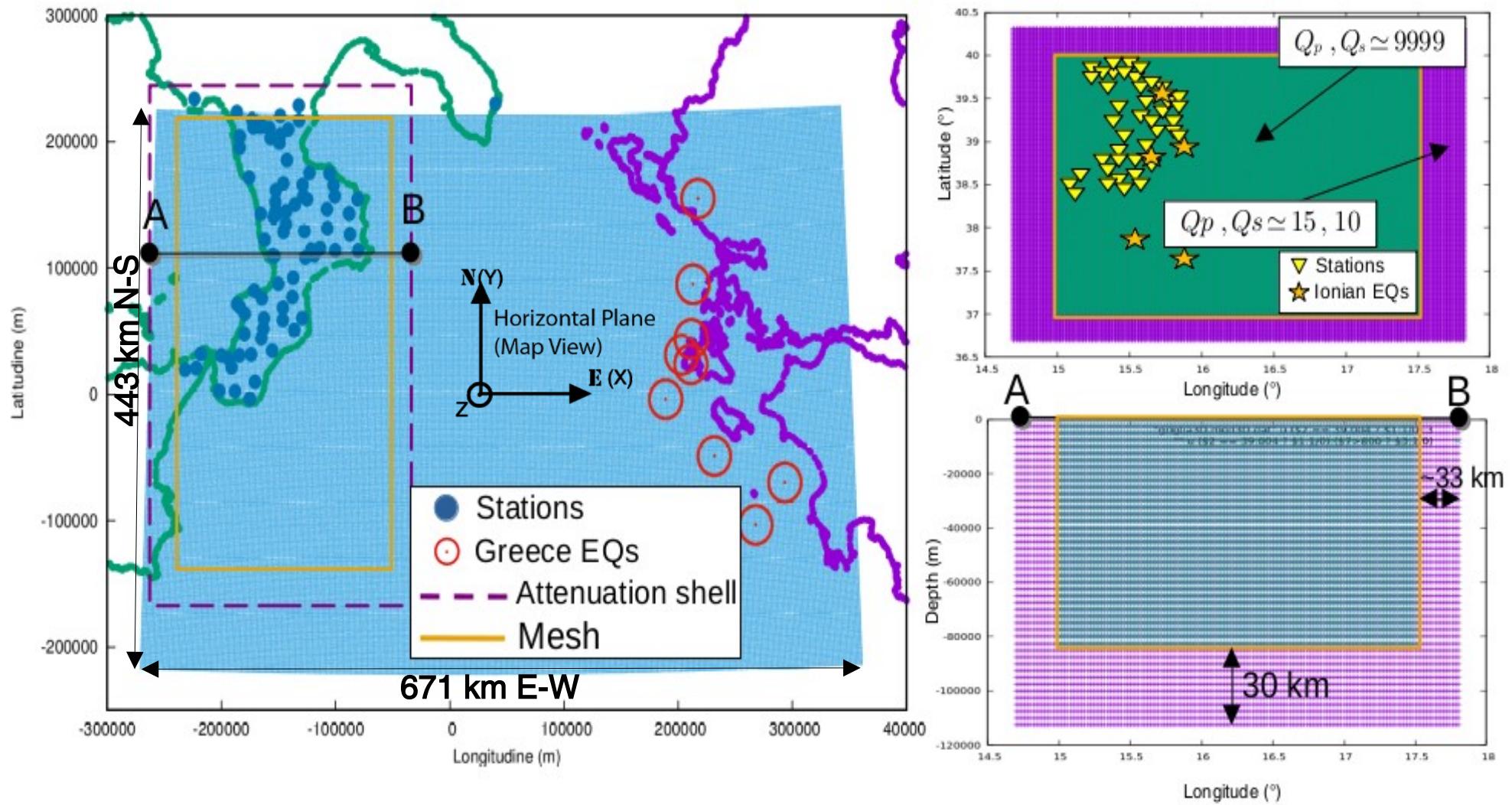
B



Synthetics

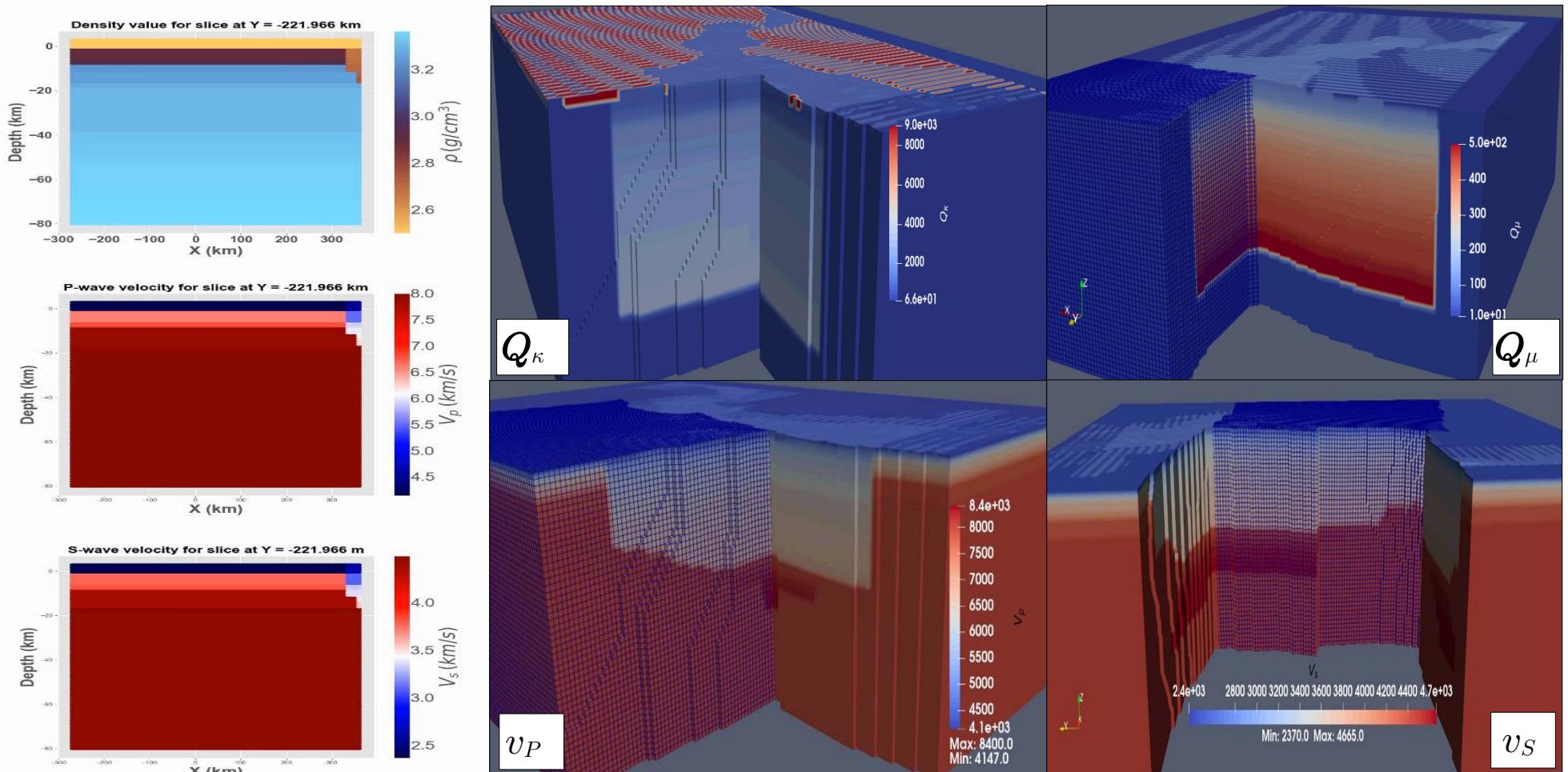


Depth section of the shell attenuation layer - ICLERC & ICLARC.



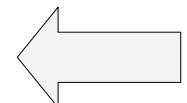
Ionian CaLabrian Ellenic aRC – ICLERC model. From 80 km up to the bottom follows the AK135 model.
 Total number of elements 6.1 million. Elements spacing ~ 2.2 km.

Reflected phases and anelastic attenuating layer for 3D model.



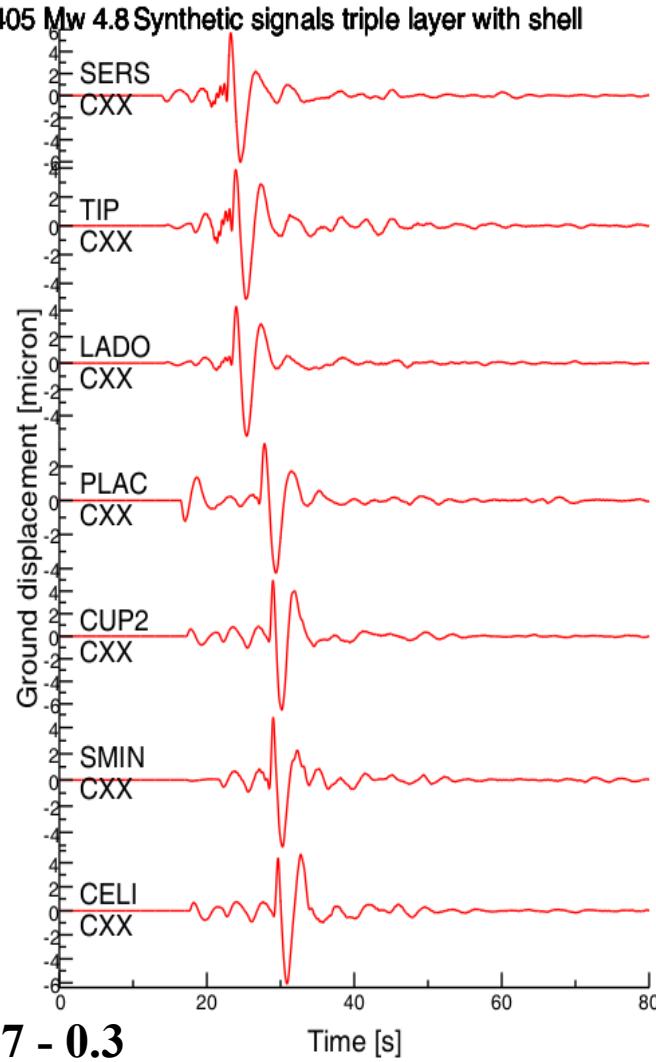
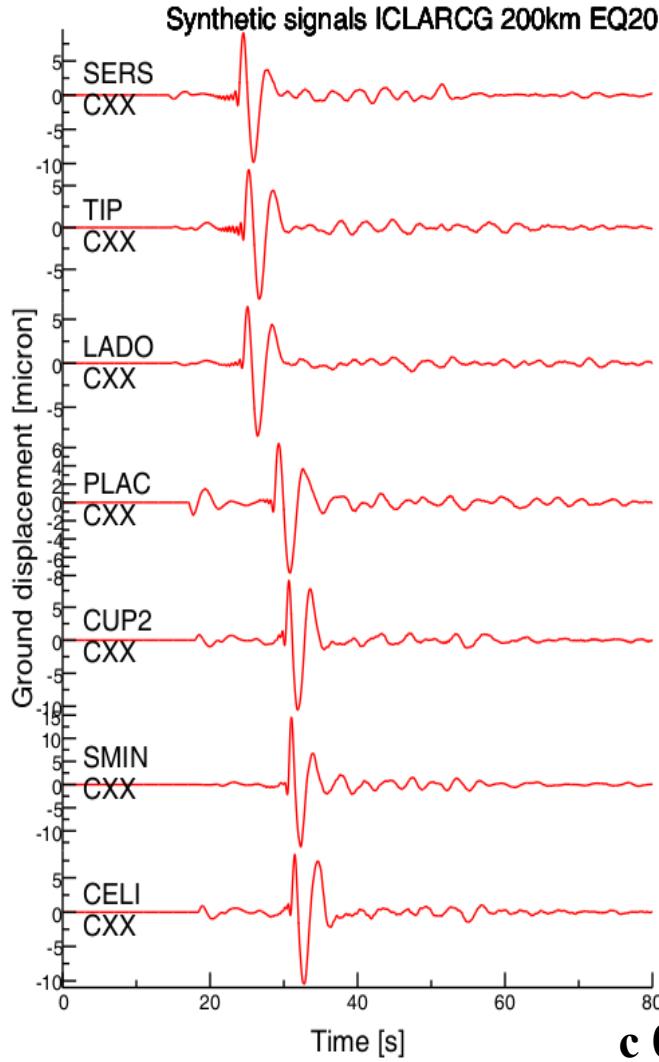
$$Q_P^{-1} = \left(1 - \frac{v_S^2}{v_P^2}\right) Q_\kappa^{-1} + \frac{v_S^2}{v_P^2} Q_\mu^{-1},$$

$$Q_S^{-1} = Q_\mu^{-1}.$$

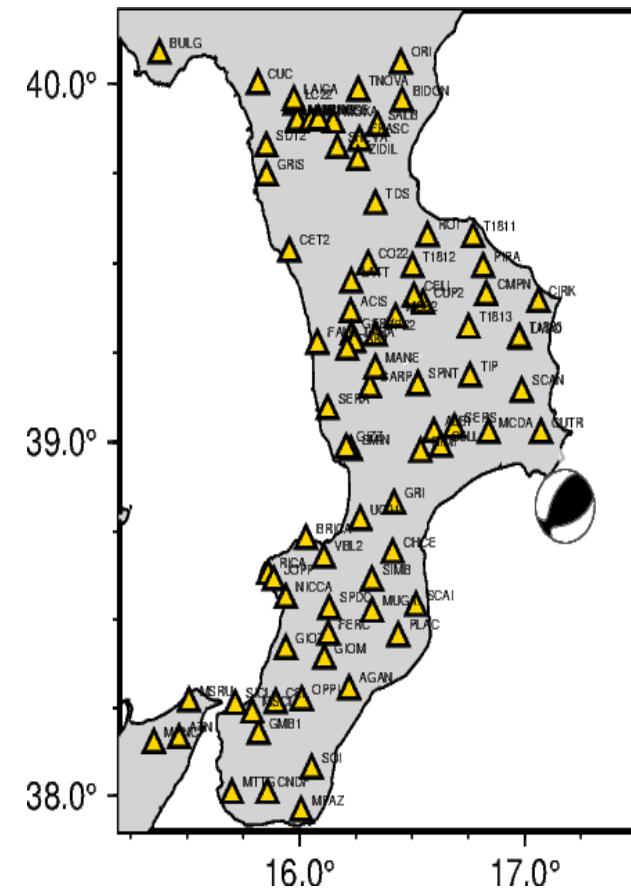


The binary, which interpolates on the GLL nodes, does the conversion internally from Q_P, Q_S (.xyz file) to Q_κ, Q_μ (shown in Figure).

Comparison of synthetics - ICLARC with shell vs 1D model with shell.



EQ20140405 Mw4.8 D64km
(38.82°N, 17.18°E) 35kmS-CUTR.

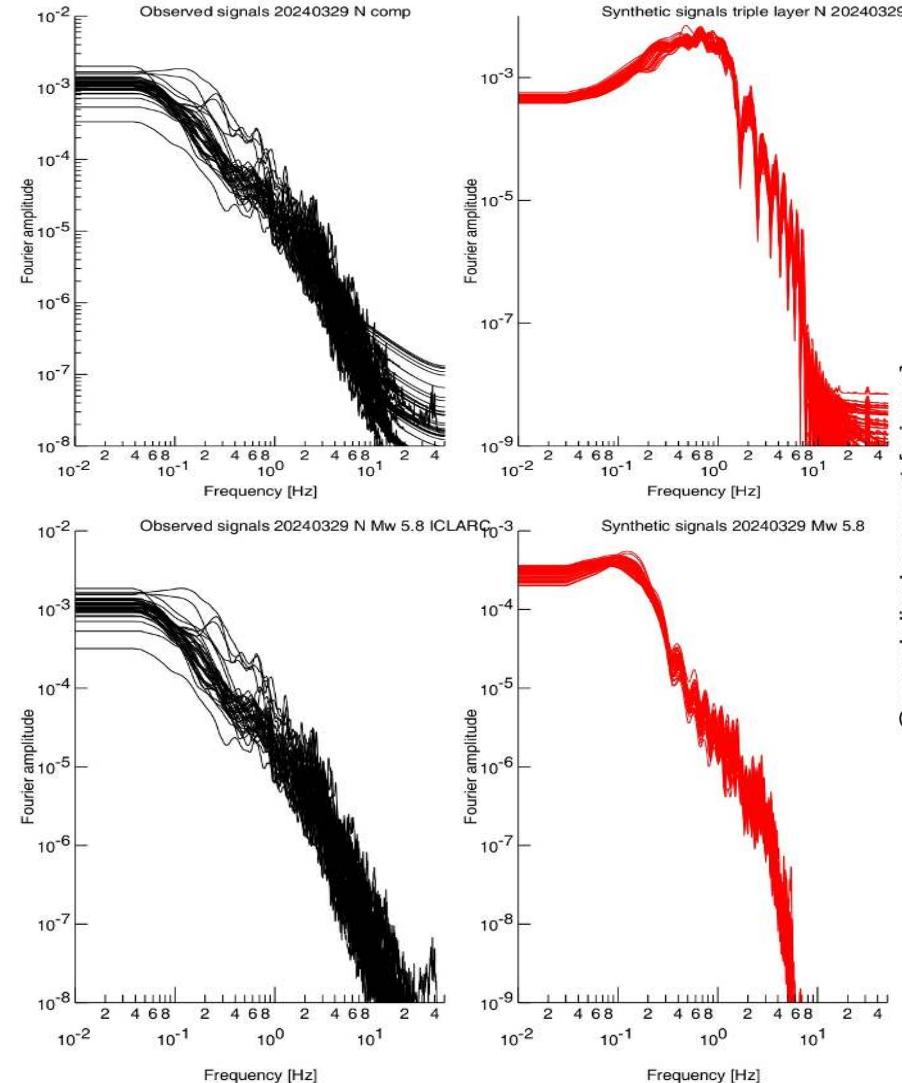


Synthetic spectra and comparison of signals with and without the shell.

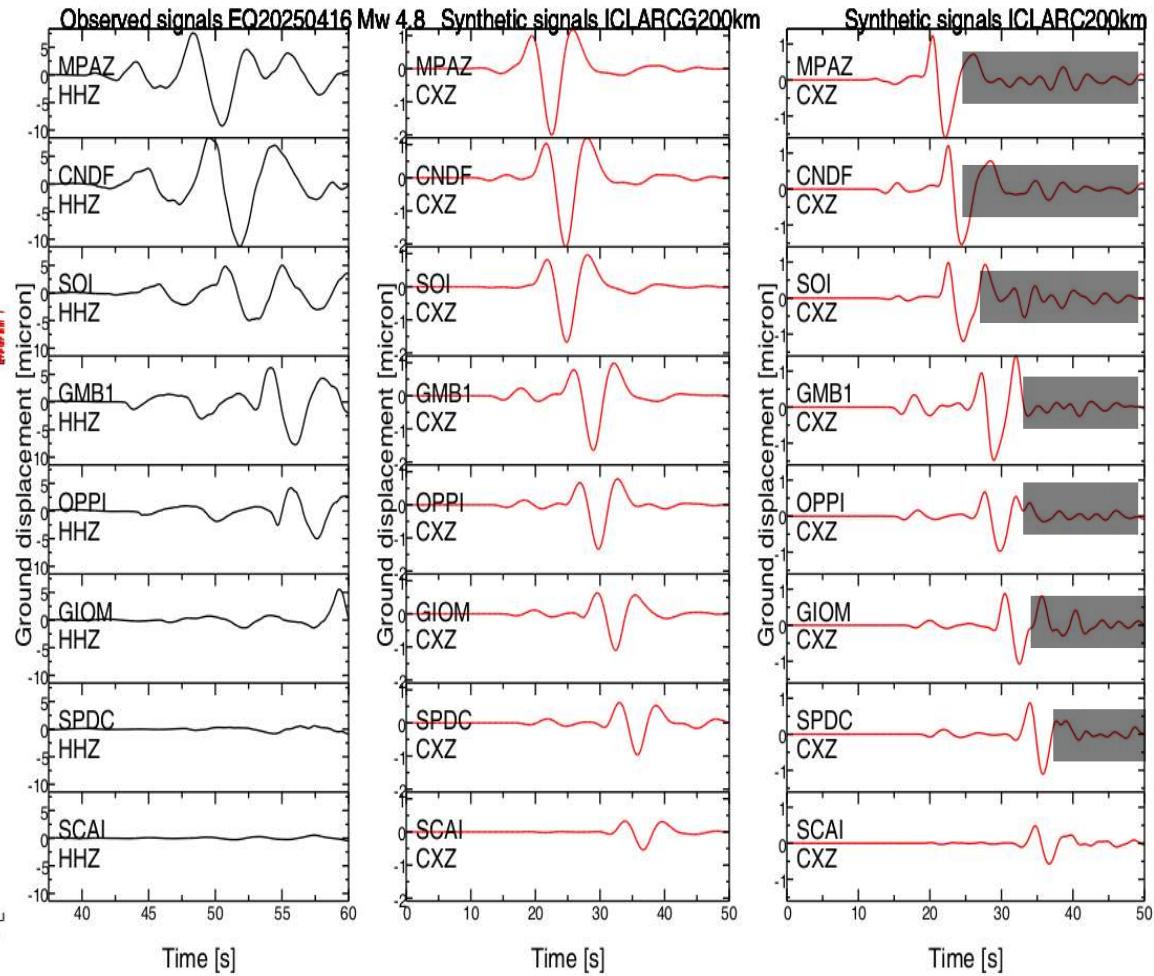
EQ20240329 Mw5.8 (37.32°N,21.31°E) D35km.

Synthetics

Observed



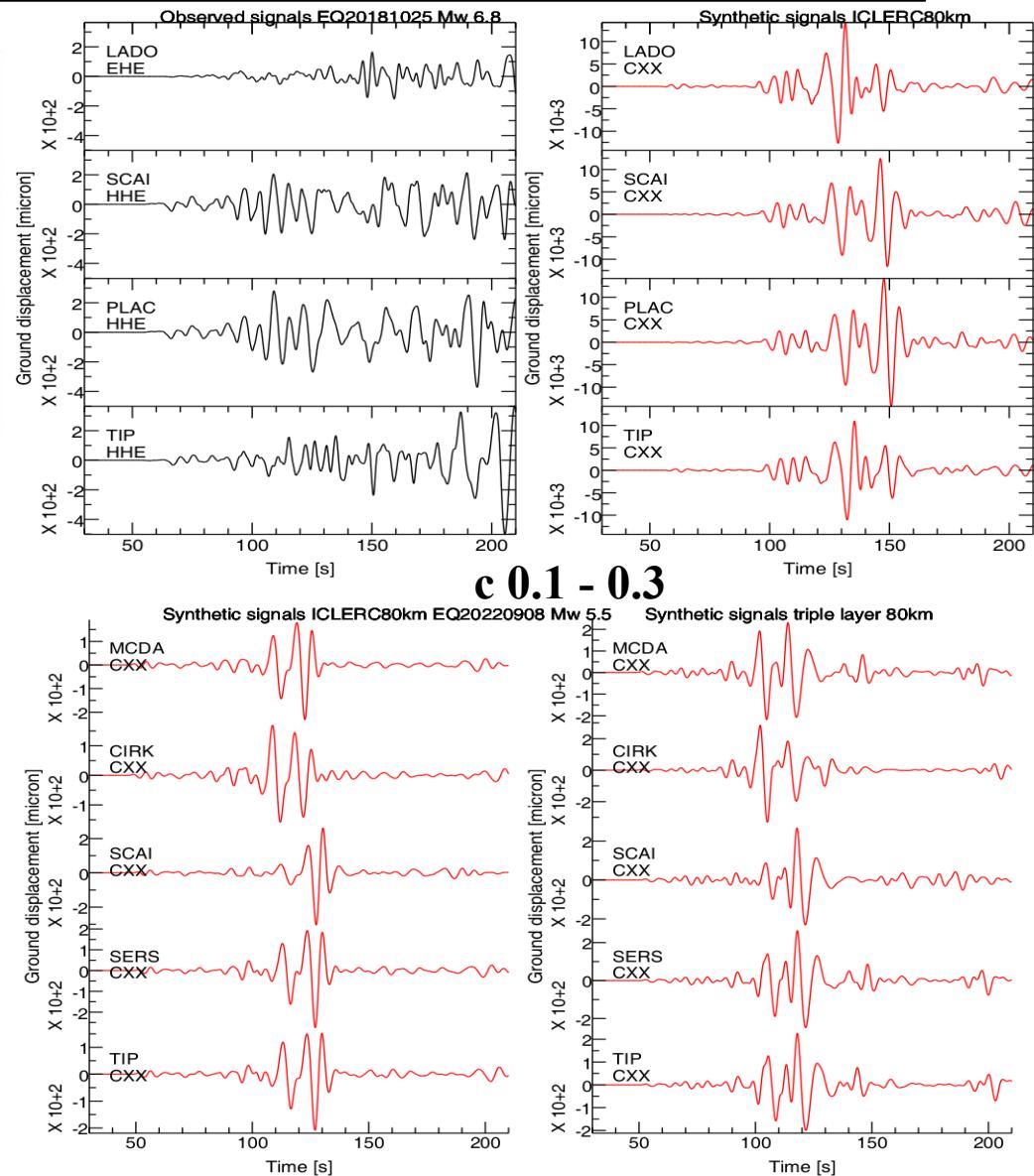
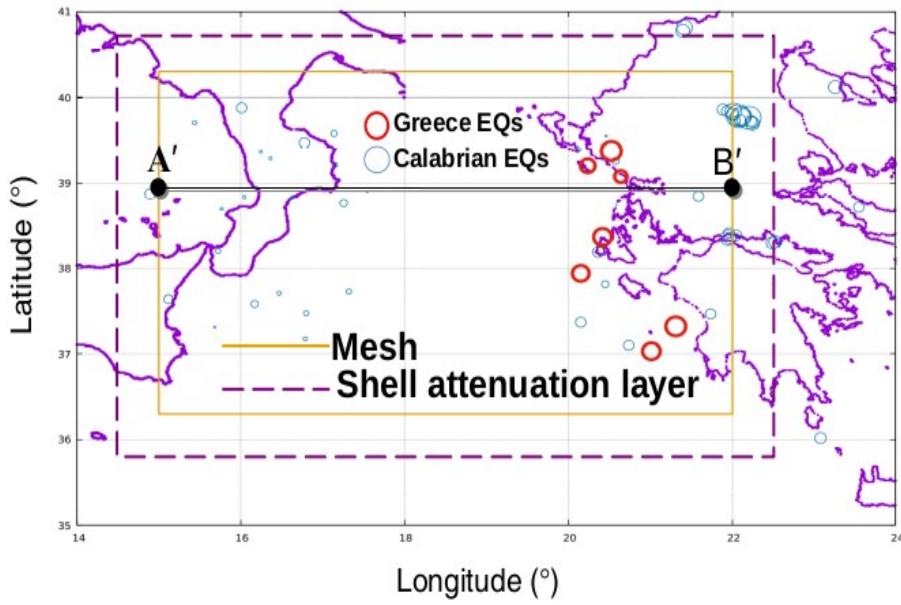
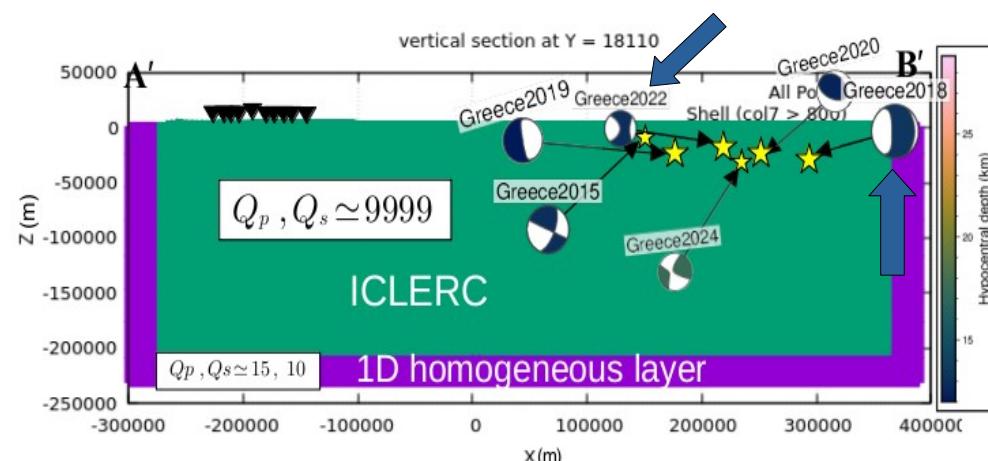
EQ20250416 Mw4.7 (37.58°N,16.17°E) D39km.



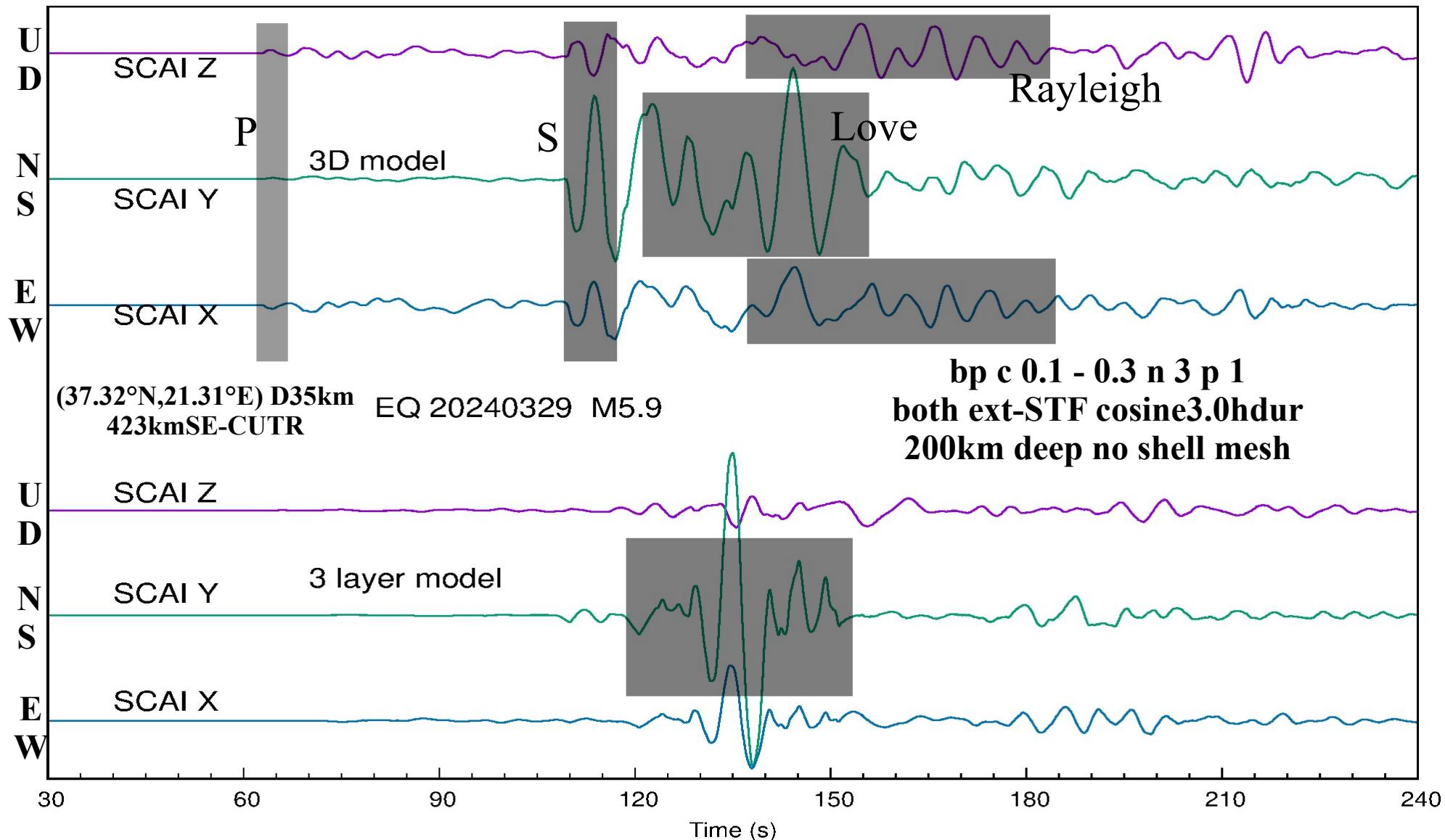
c 0.07 - 0.22

The shell works. However further simulations must be done, perhaps with a Q gradient.

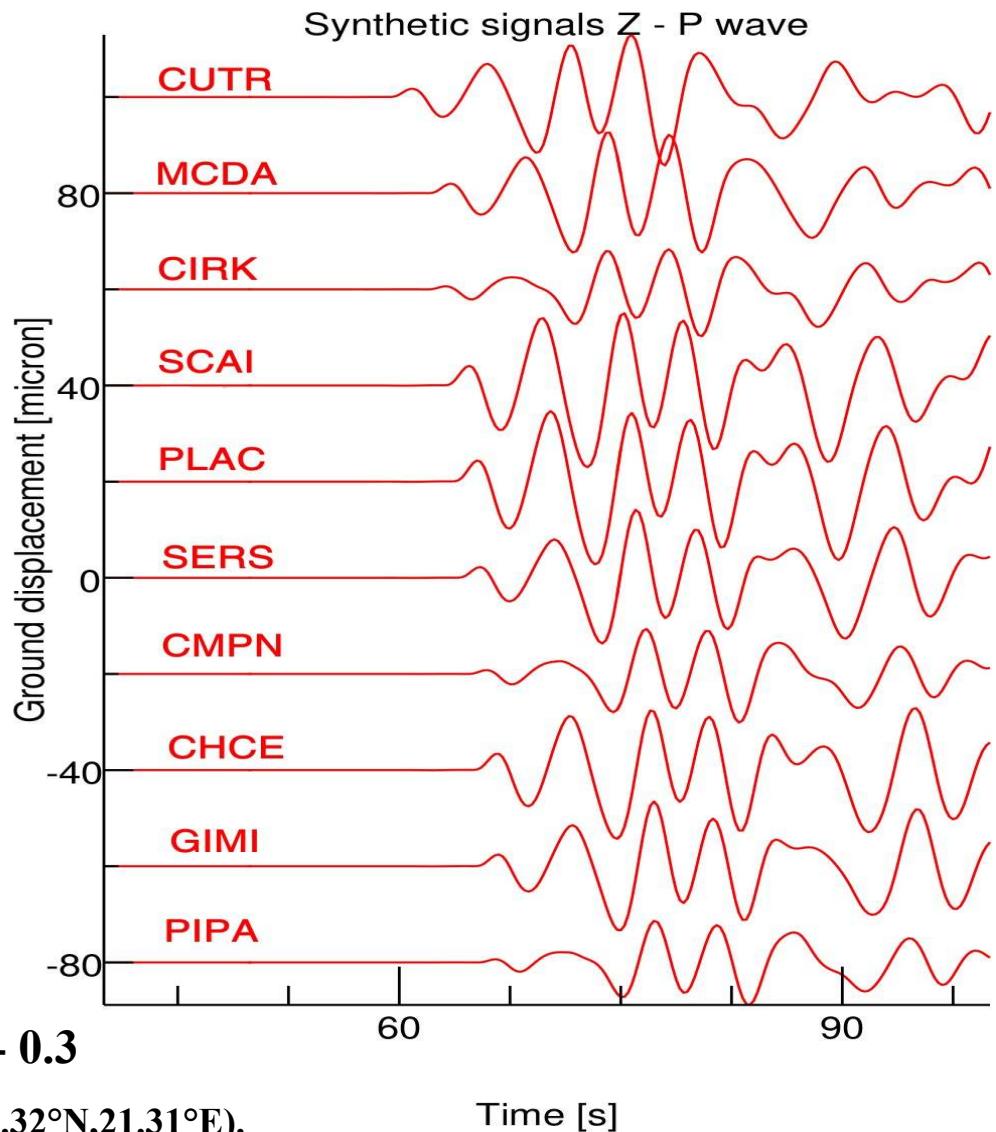
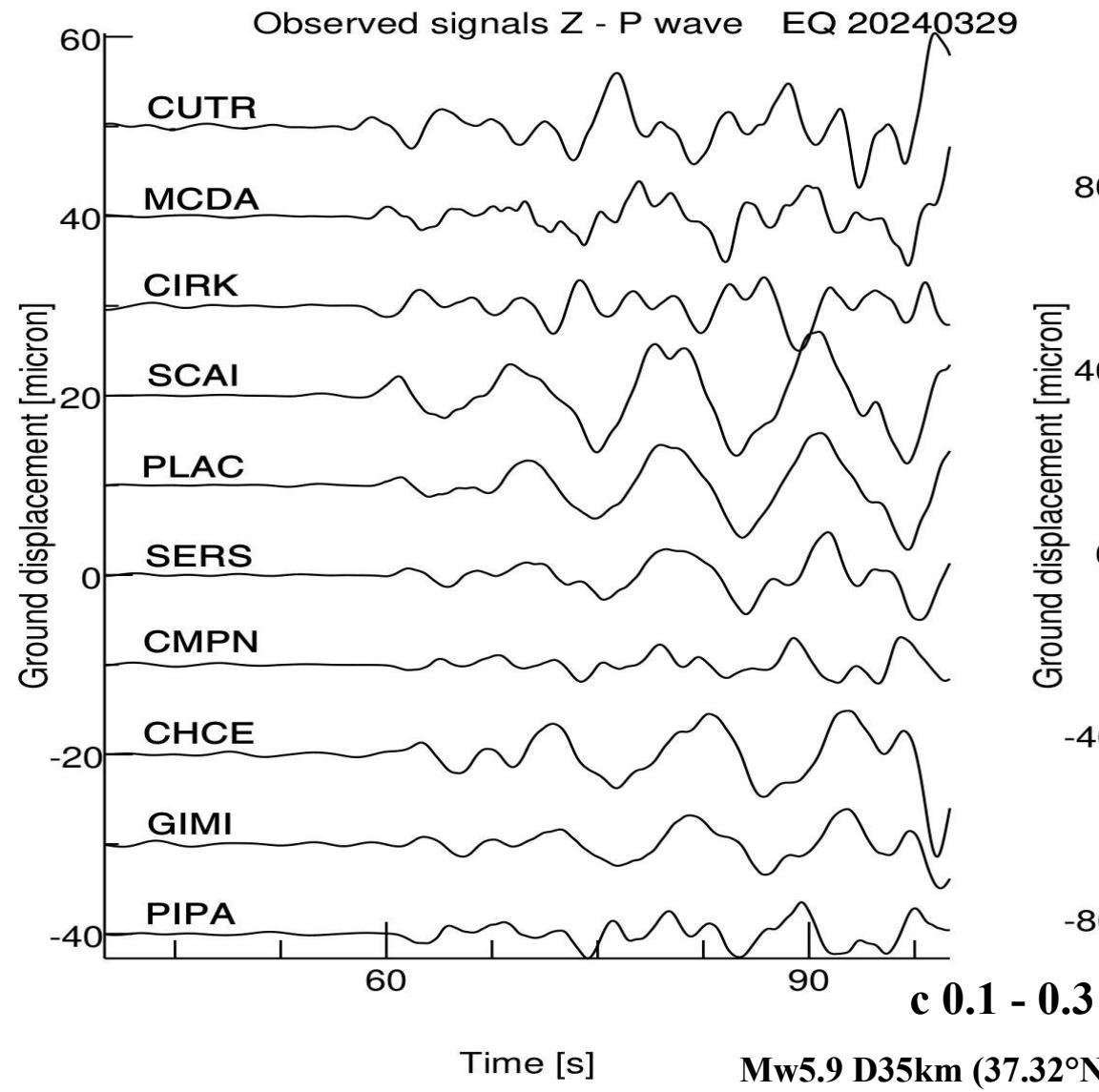
Greece earthquakes - EQ20181025 & EQ20220908 (no shell layer).



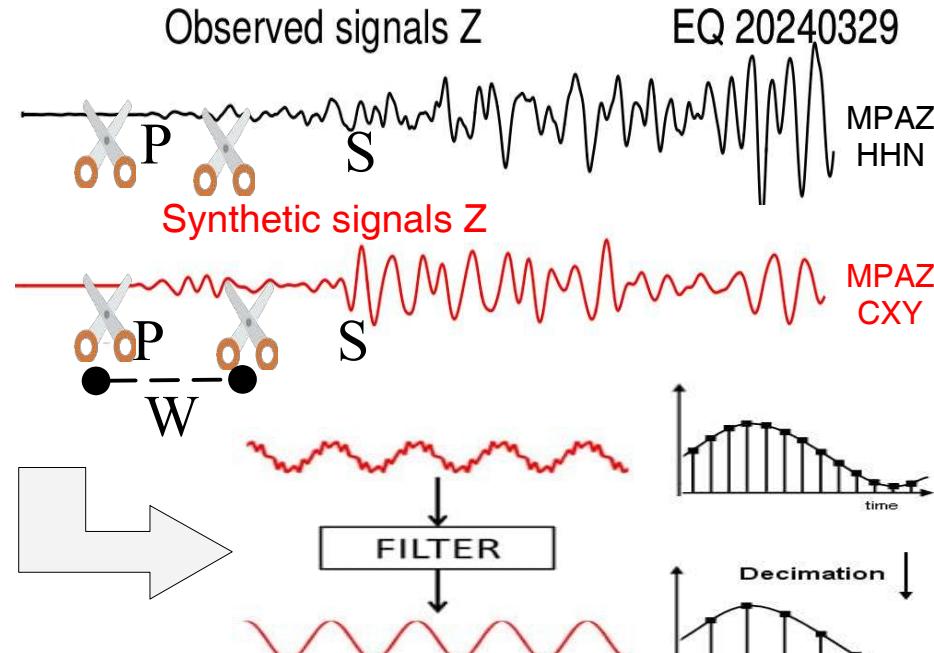
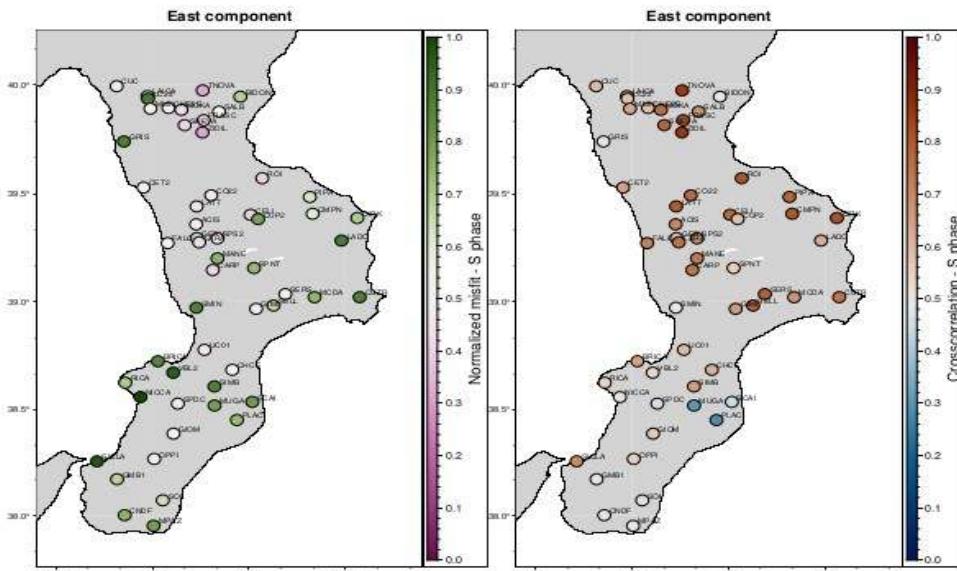
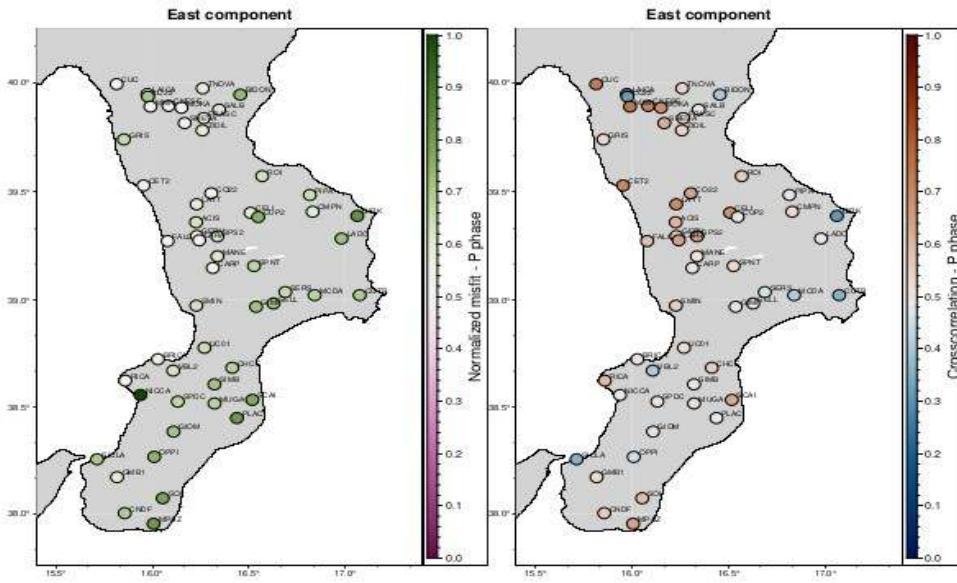
Synthetics from ICLERC 3D model with respect to triple layer model.



Synthetics vs real data - ICLERC (no shell).



Misfit metrics computation - ICLERC EQ20240329 [0.07 - 0.12 / 0.012 - 0.03] Hz [30 / 30] s.



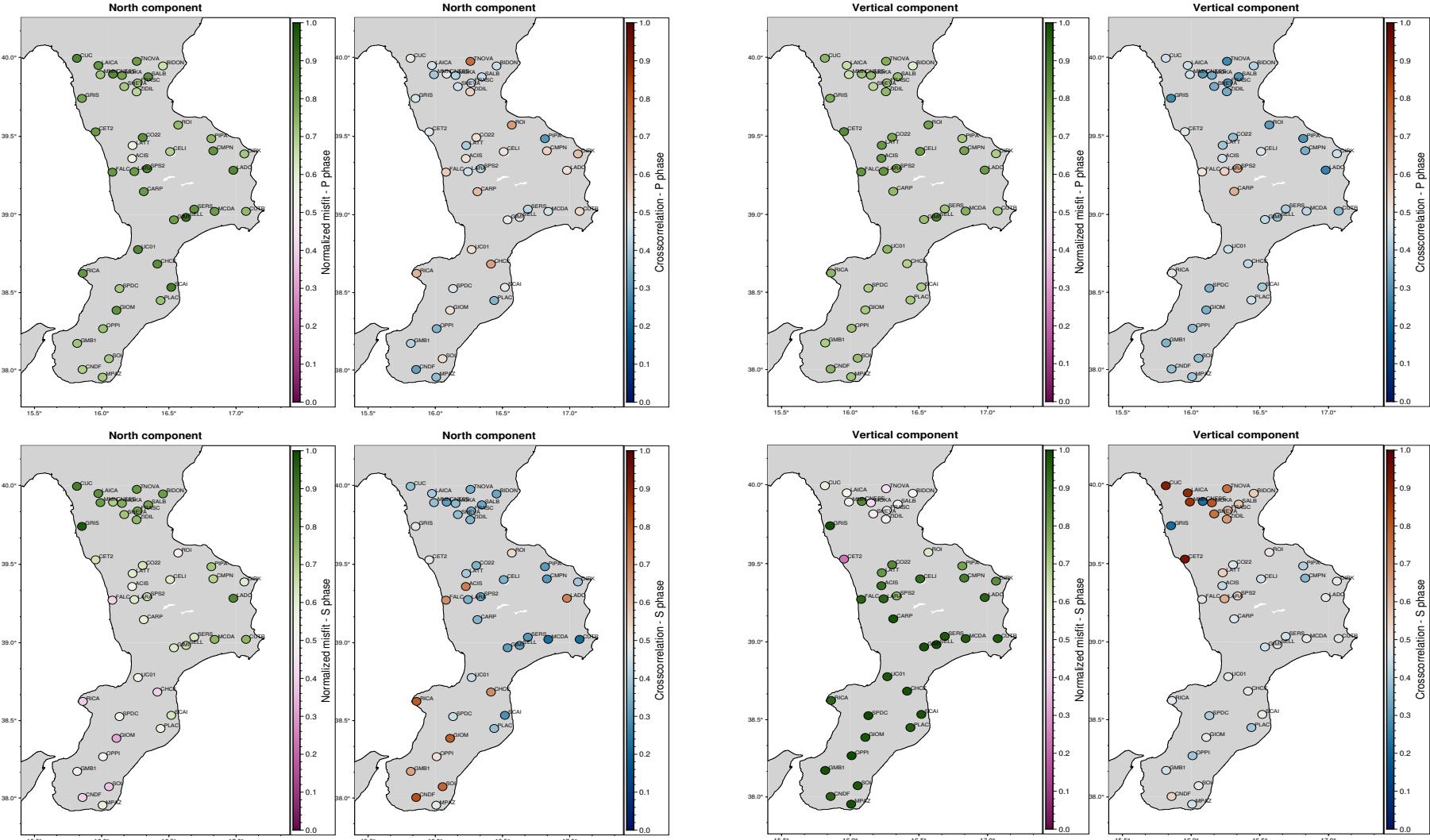
$$\chi(m) = \frac{1}{2} \sum_{i=1}^{N_r} \frac{\sqrt{\int_W |d_i(t) - s_i(t)|^2 dt}}{\sqrt{\int_W d_i(t)^2 dt} + \sqrt{\int_W s_i(t)^2 dt}}$$

Normalized L2 norm

$$NCC(\tau) = \sum_{i=1}^{N_r} \frac{\int_W |d_i(t) s_i(t + \tau)| dt}{\sqrt{\int_W d_i(t)^2 dt} \sqrt{\int_W s_i(t)^2 dt}}$$

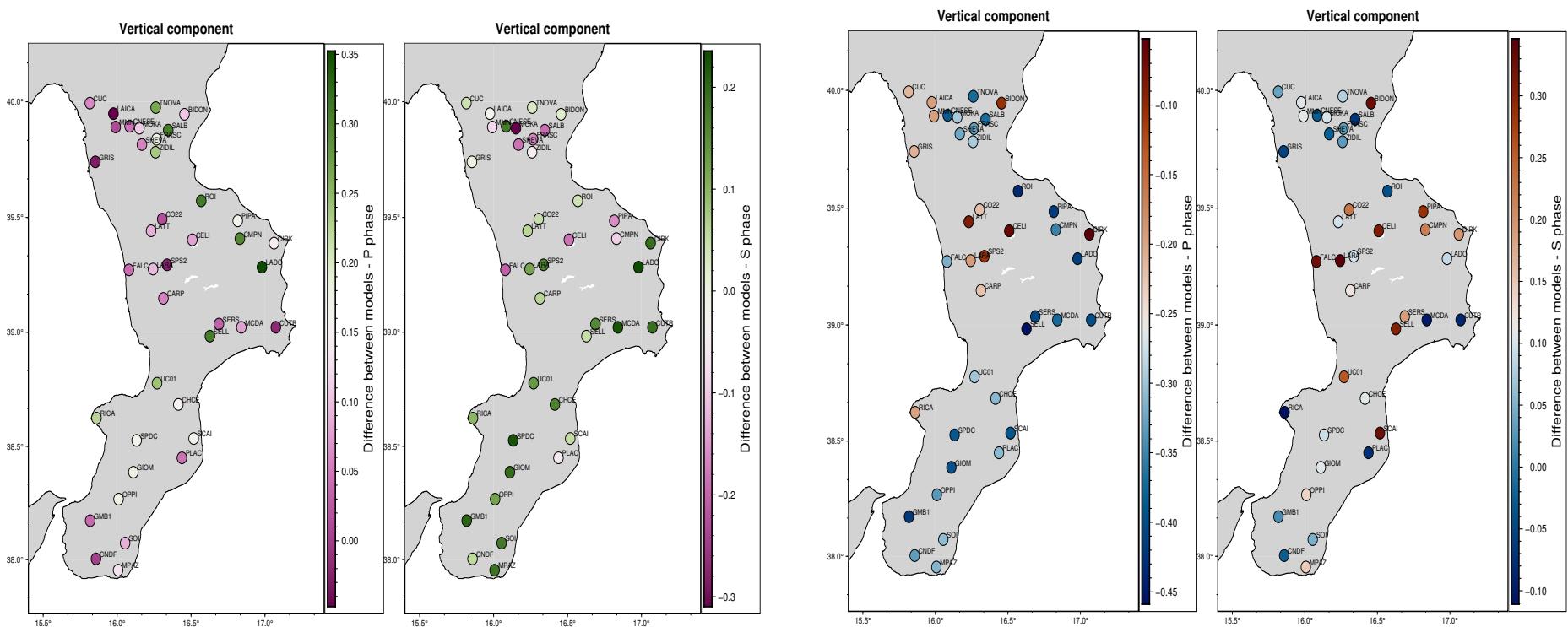
Normalized Cross Correlation

Misfit metrics computation - ICLERC EQ20240329 [0.07 - 0.12 / 0.012 - 0.03] Hz [40 / 40] s.



- Widening the time window for metric computation improves both L2 norm and cross-correlation performance.

Misfit metrics computation – **ICLERC** vs 1D model EQ20240329 [0.07 - 0.22 / 0.02 - 0.10] Hz [30 / 35] s.



- NCC gives better results between the two metrics, by looking at Z component synthetics.
- This was confirmed by looking at polarity and shape for the P wave arrivals.
- Towards Vibo Valentia and the Aspromonte region the L2 differences increase with encouraging results.

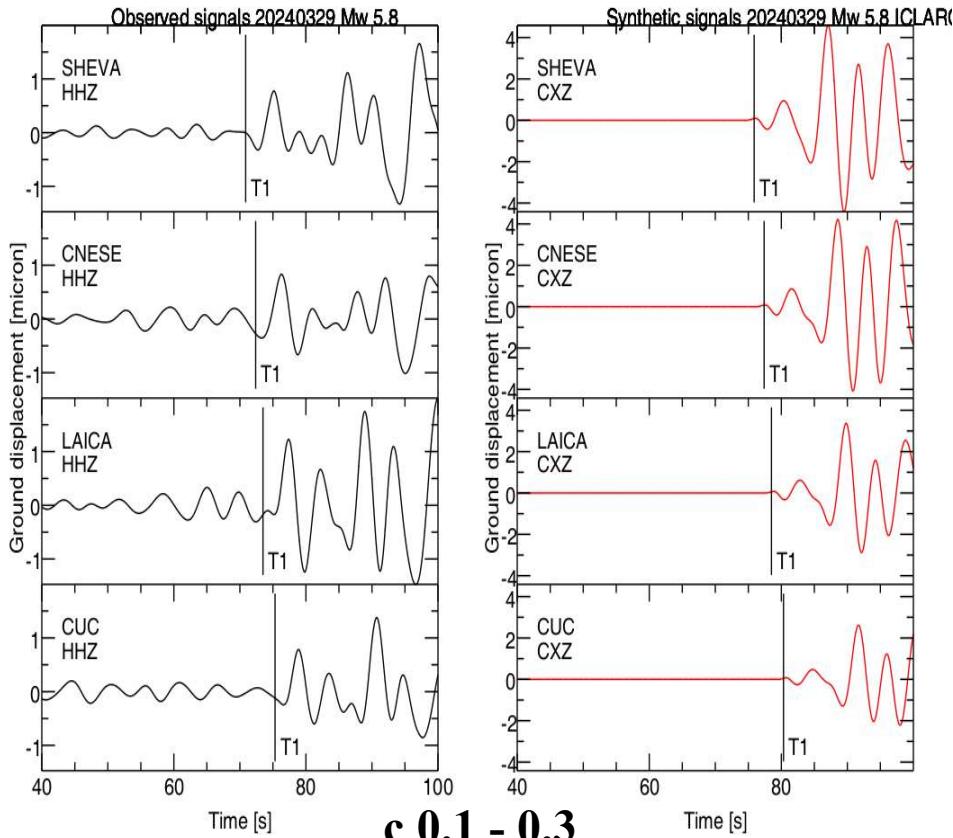
Goal

$$\delta\chi < 0$$

$$\delta NCC > 0.$$

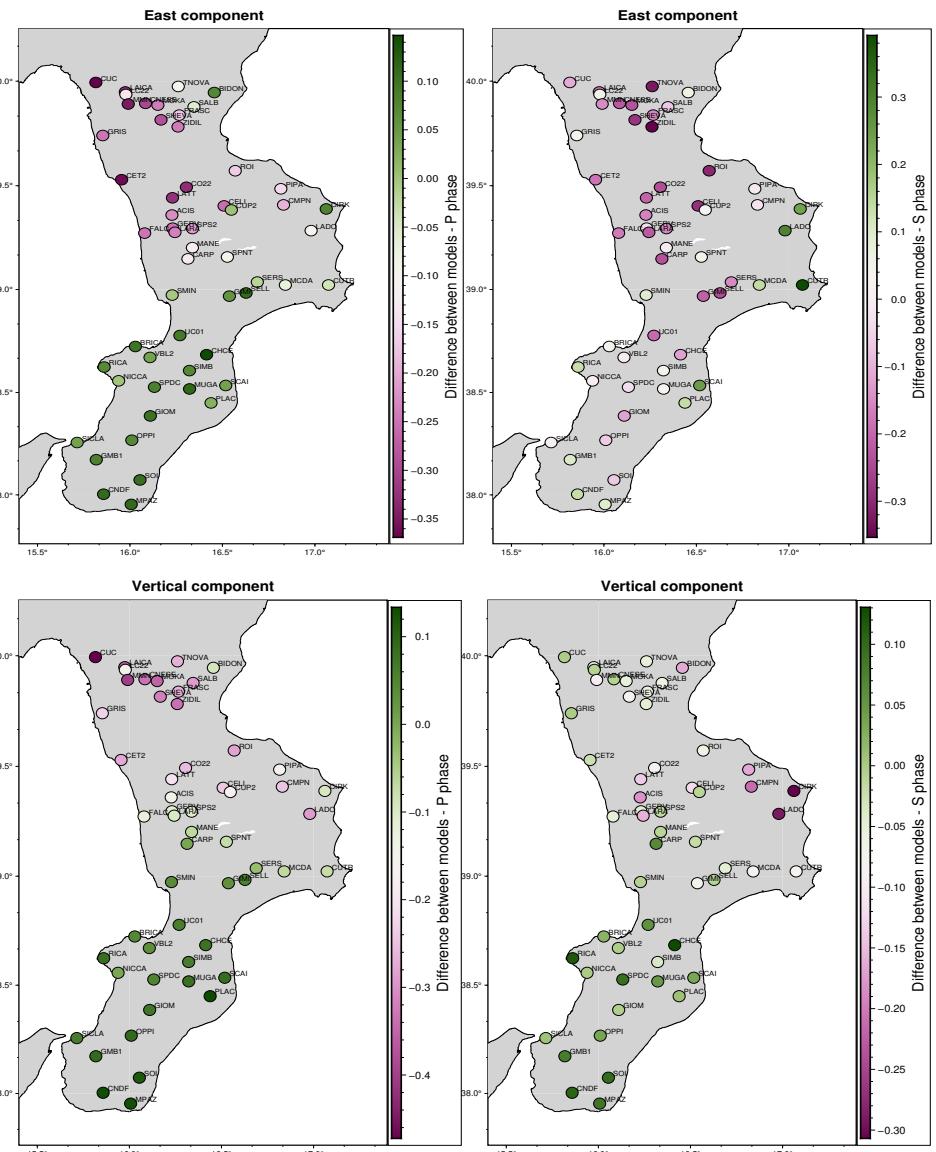
Misfit metrics computation – EQ20240329 [0.07 - 0.22 / 0.02 - 0.10] Hz [30 / 35] s.

Changing the STF type to a Dirac delta one improves significantly the results.



Phase arrival times are consistent between synthetic and observed signals.

However the frequency content of synthetic signals is narrower than observed recordings.

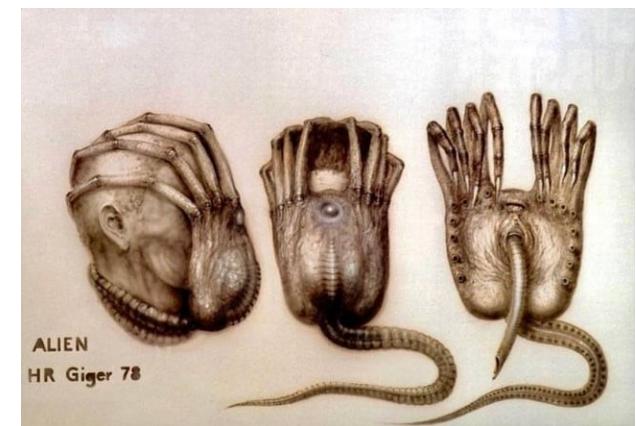


Preliminary conclusions and final remarks.

- Most synthetics show a satisfactory agreement about wave polarity and amplitude compared to real data.
- Undesired phases reflected at mesh boundaries are attenuated but not completely removed by the shell layer.
- Used misfit metrics confirm that preliminary 3D model (ICLARC/ICLERC) is better than the triple layer model.
- Achieving precise 3D model of the Ionian basin requires a lot more observations, possibly even at the Sea bottom, likely feasible in future decades.

Overall, this is still a first small step towards a more refined knowledge of the area.

The uncertainties are many, but that's up to future researchers to deal it with. (<https://github.com/ionianbasin/>)



I acknowledge **Newton cluster**, and the super-computing facility of the University of Calabria (Cosenza), **Mecenate** at Sismolab and the Seismology group for hosting me at the **University of Münster**.
Thank you for your attention.