



# ADVANCED HYDRAULIC SIMULATION MODELS

## 1D hydraulic simulation models of channels and rivers: PART I

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# Content

1. Model hypothesis: basic concepts in channels
2. Steady flow analysis
  1. Energy balance
  2. Uniform flow: Normal Depth
  3. Energy: Critical Depth
  4. Boundary conditions
3. Complete flow equations
  1. Conservation laws
  2. Resolution

# Content

## **1. Model hypothesis: basic concepts in channels**

### 2. Steady flow analysis

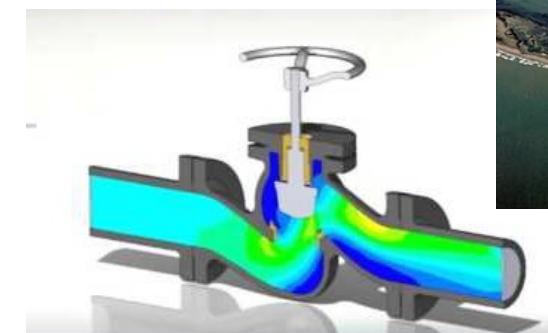
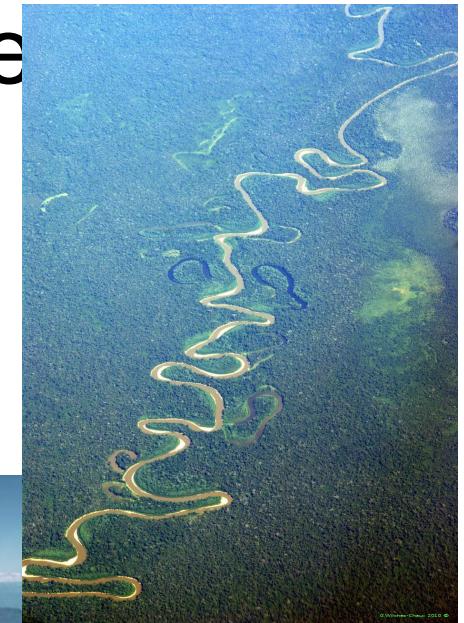
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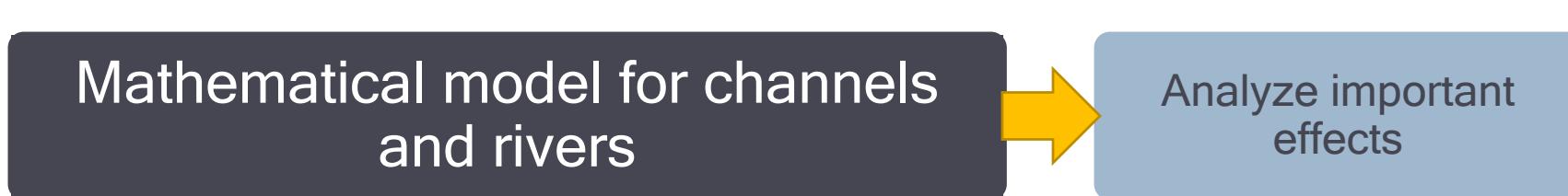
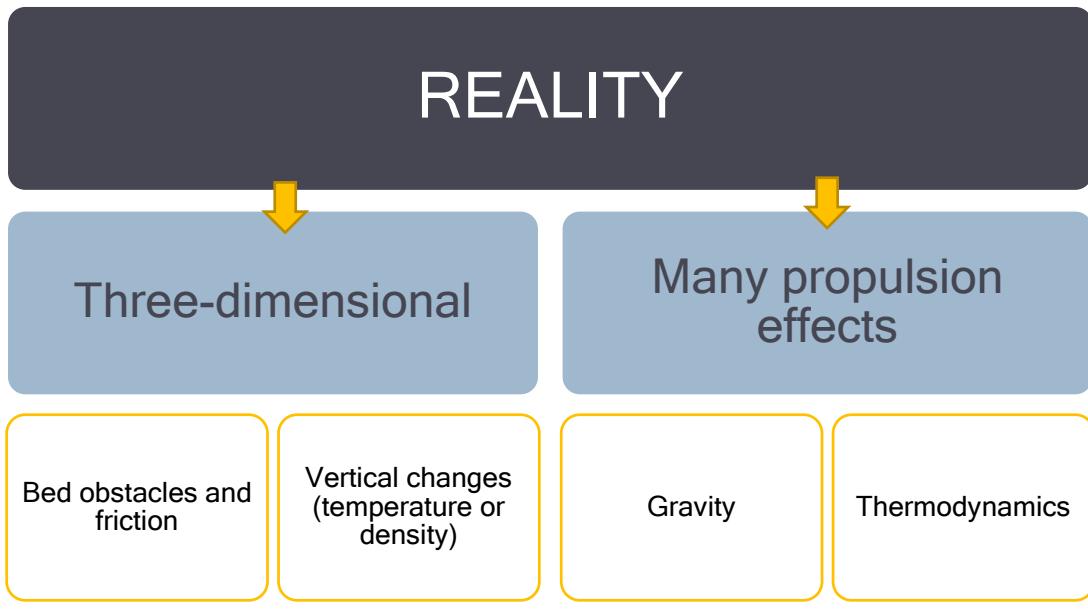
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# What type of flows are we going to simulate

- Pressure flow/free surface flow
- Flow rate, average velocity
- Piezometric level, depth, wetted section
- Stationary/transient flow
- Laminar/turbulent flow
- Ideal/viscous flow



# Use of models for flows



# Characterization of flow in channels and rivers

- One-dimensional movement
- Gravity is the driving force
- Average velocity in section.  $Q=uA$
- Hydrostatic pressure distribution  $p=p(h)$
- Mild slopes:  $\sin(\theta) \approx \tan(\theta)$



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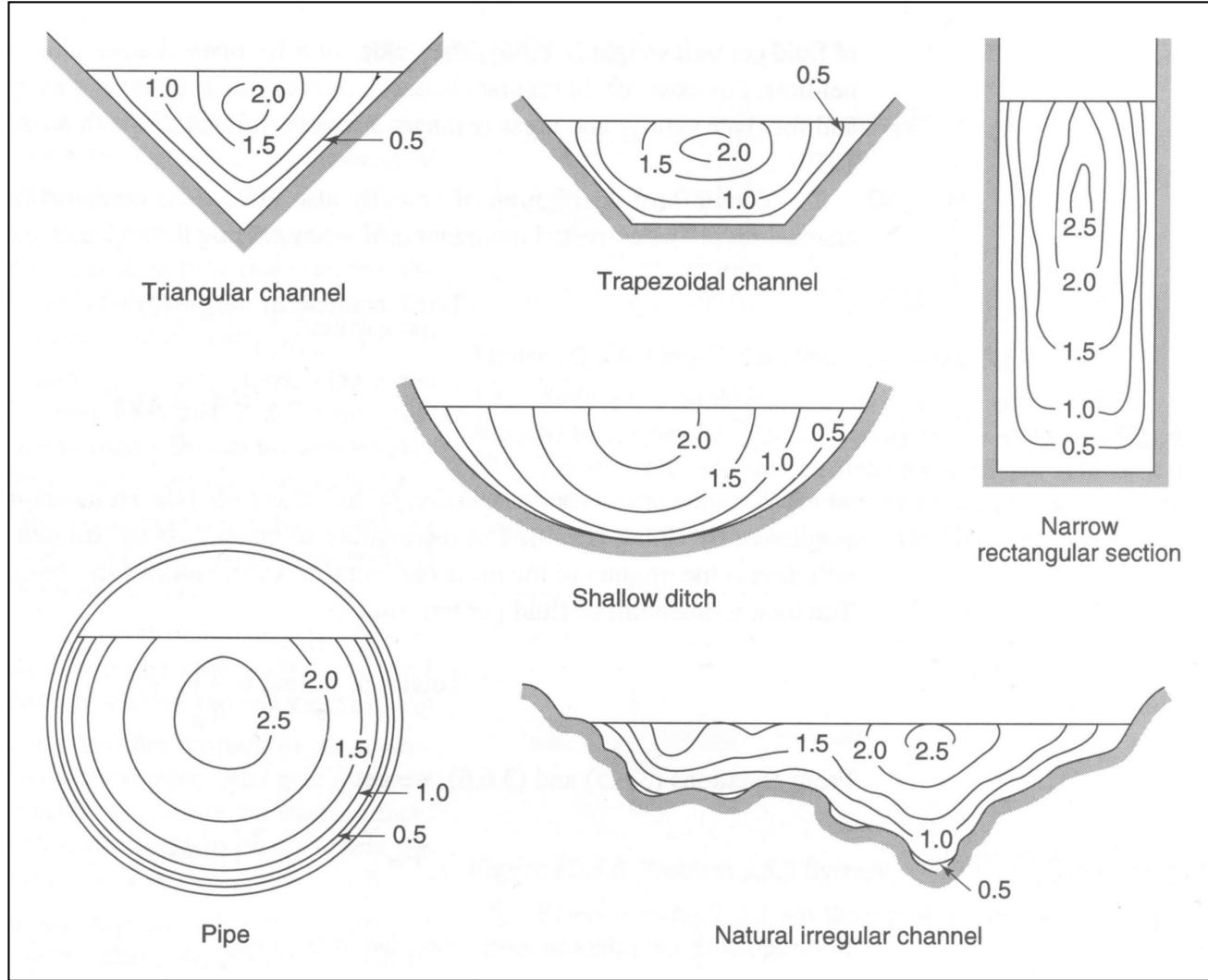


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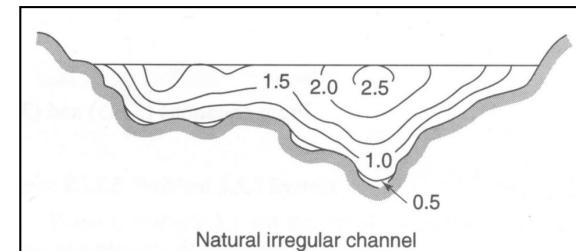
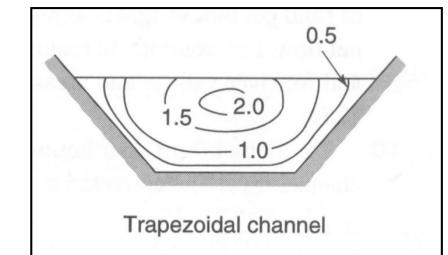


# Velocity distribution across the section



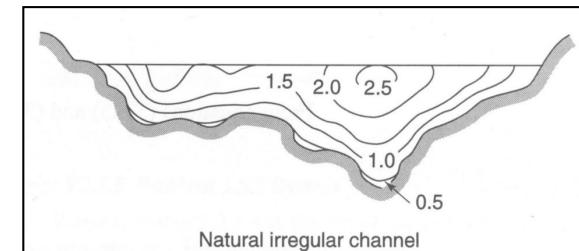
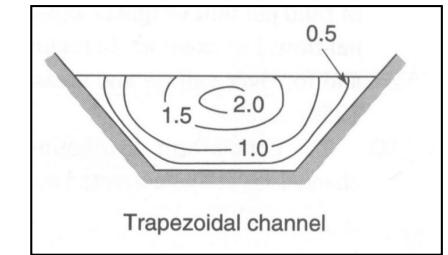
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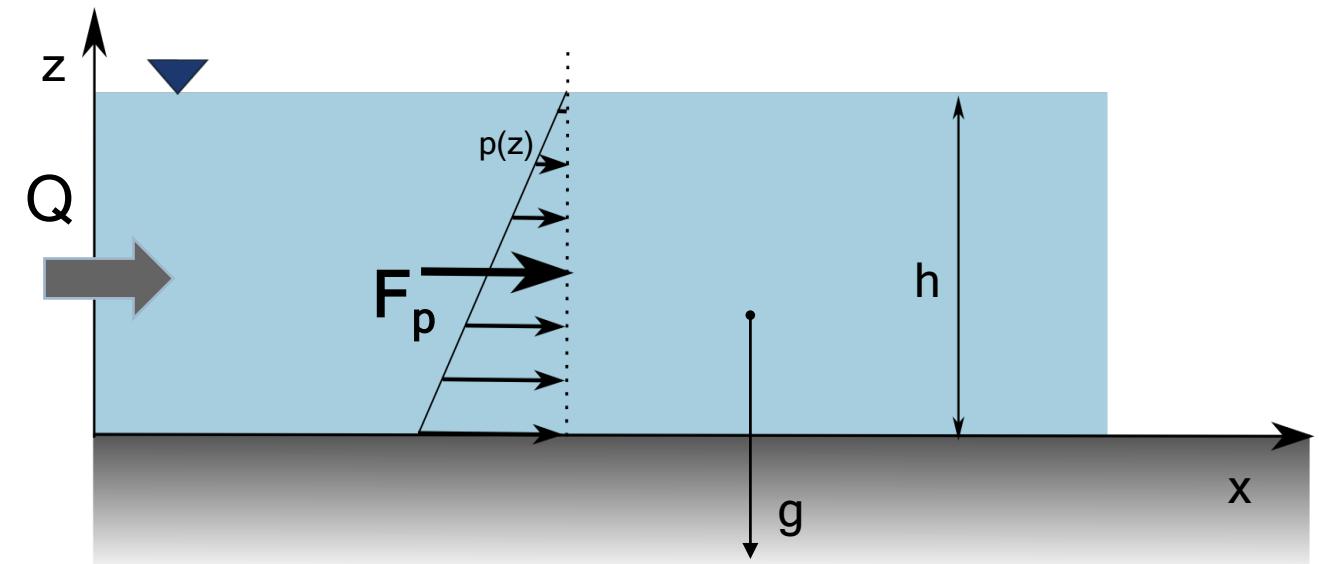
# Hydrostatic pressure

- Typical property of shallow flows: hydrostatic z-pressure

$$p(z) = p(h) + \rho g(h - z) = \rho g(h - z)$$

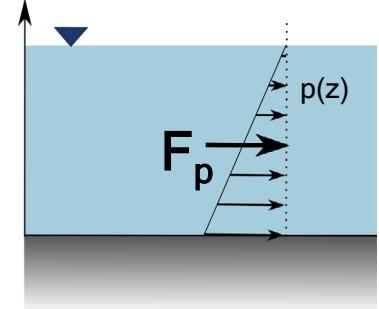
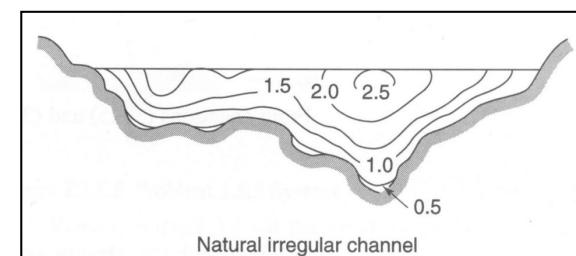
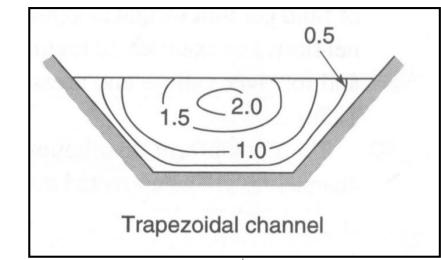
- The pressure at position  $x$ :  $p(x) = p(h(x))$
- By integrating it, we obtain the total pressure force

$$F_p = \frac{1}{2} \rho g h^2$$



# Characterization of flow in channels and rivers

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- Hydrostatic pressure distribution  $p=p(h)$
- Mild slopes:  $\sin(\theta) \approx \tan(\theta)$ ,  $\cos(\theta) \approx 1$



# Basic concepts on channels

Why does the water flow in the rivers?



# Content

1. Model hypothesis: basic concepts in channels

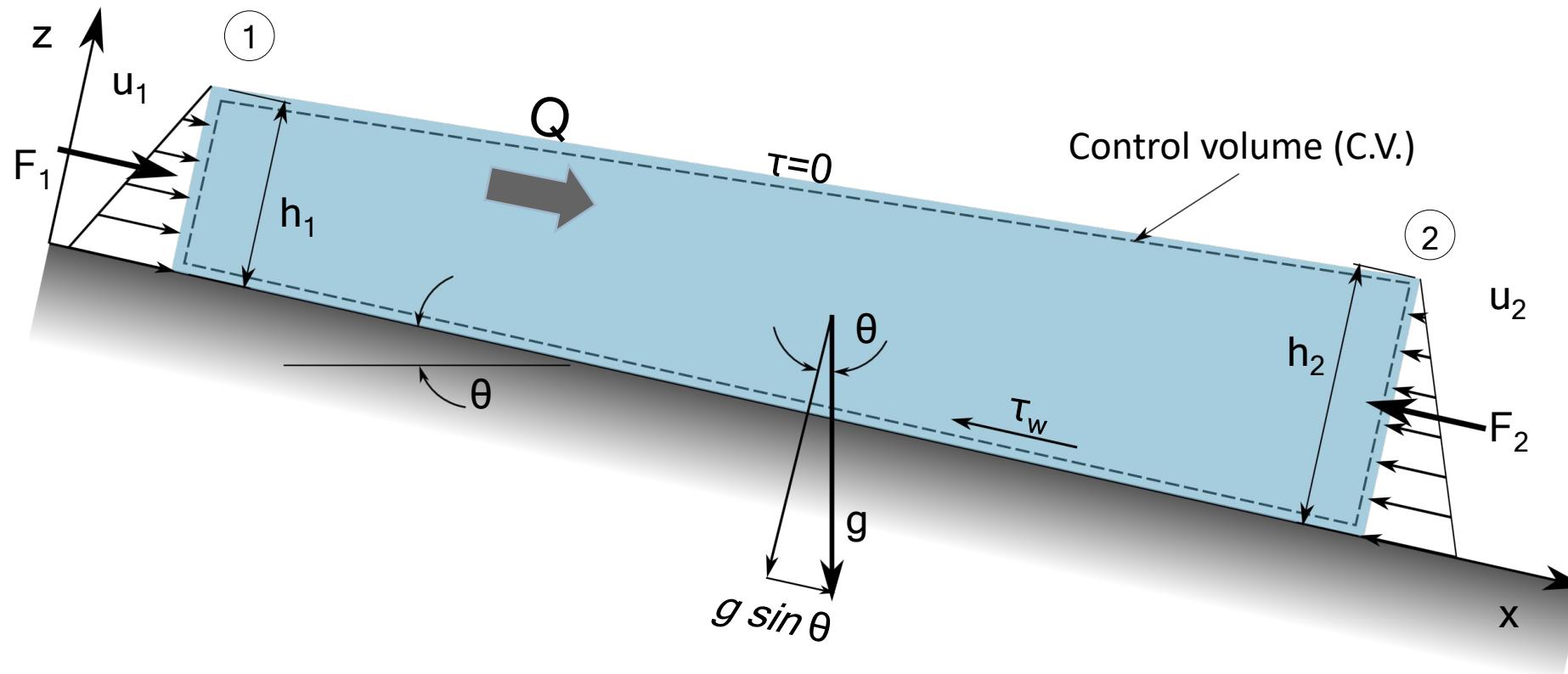
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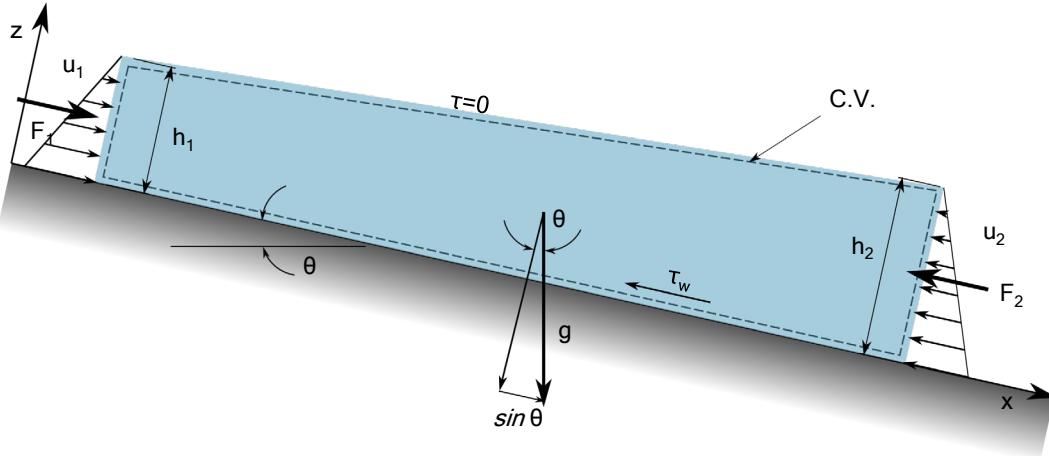
3. Complete flow equations

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# Main forces applied: control volume (C.V.)



# Integral energy balance of a C.V.



$$\int_1 \rho \frac{\vec{u}^2}{2} (\vec{u} \cdot \hat{n}) dS + \int_2 \rho \frac{\vec{u}^2}{2} (\vec{u} \cdot \hat{n}) dS = - \int_1 P^* (\vec{u} \cdot \hat{n}) dS - \int_2 P^* (\vec{u} \cdot \hat{n}) dS - \int_V \phi_{vol} dV$$

$$\left( \frac{u^2}{2g} + h + z_b \right)_1 - \left( \frac{u^2}{2g} + h + z_b \right)_2 = \frac{\int_V \phi_{vol} dV}{\rho g Q}$$

$$Q = cte.$$

Total pressure:

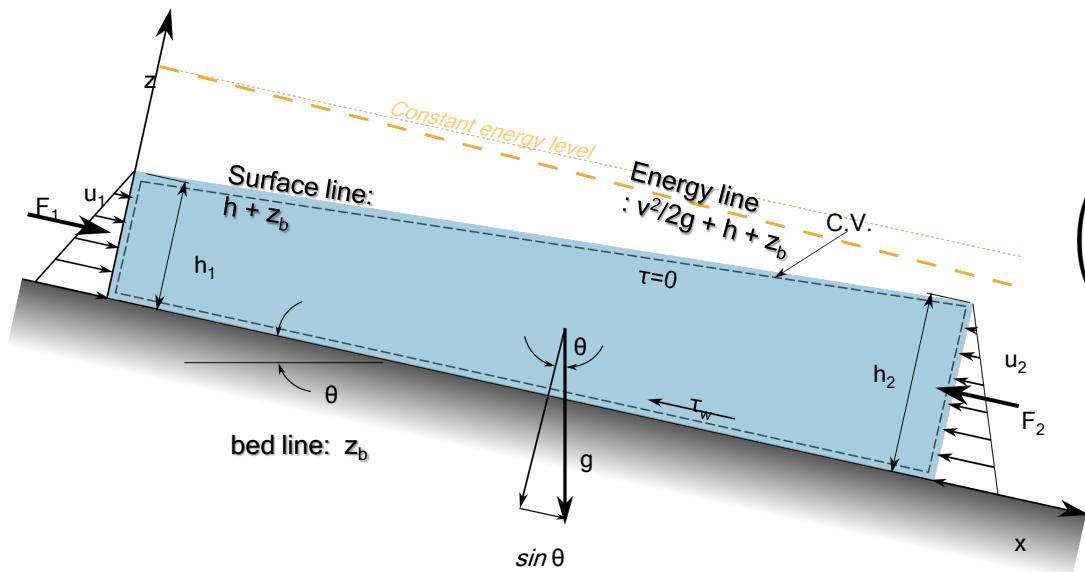
$$P^* = p - \rho g x \sin \alpha + \rho g z \cos \alpha$$

Viscous losses  
In units of length

$$H_1 - H_2 = \Delta H$$

# Formulation of friction losses

$$\frac{\Delta H}{\Delta x} = S_f = \frac{n^2}{R_h^{4/3}} u^2 = \frac{n^2}{R_h^{4/3}} \frac{Q^2}{A^2} \quad n: \text{Manning's Coefficient}$$



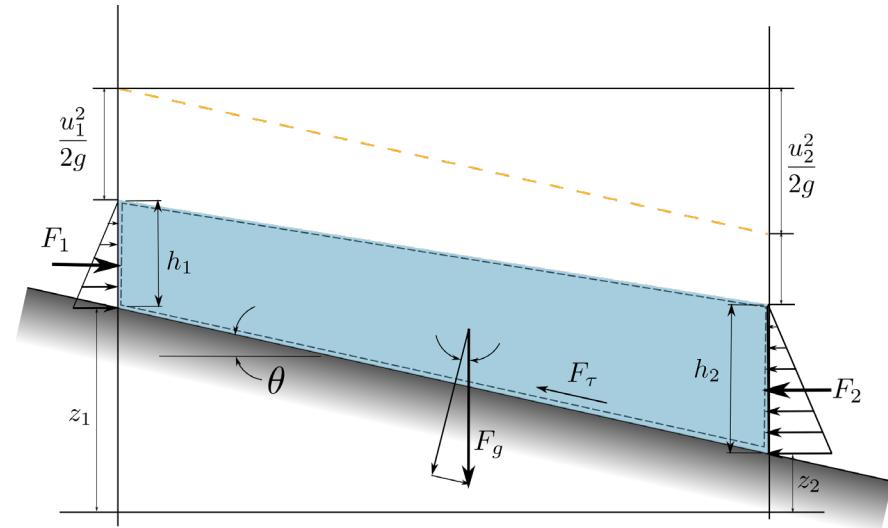
$$\left( \frac{u^2}{2g} + h + z_b \right)_1 - \left( \frac{u^2}{2g} + h + z_b \right)_2 = \Delta H = S_f \Delta x$$

$R_h$ : Hydraulic radius  
 $S_f$ : Friction slope

# Uniform/non-uniform flow

General case: non-uniform flow

$$\left( \frac{u^2}{2g} + h + z_b \right)_1 - \left( \frac{u^2}{2g} + h + z_b \right)_2 = S_f \Delta x$$

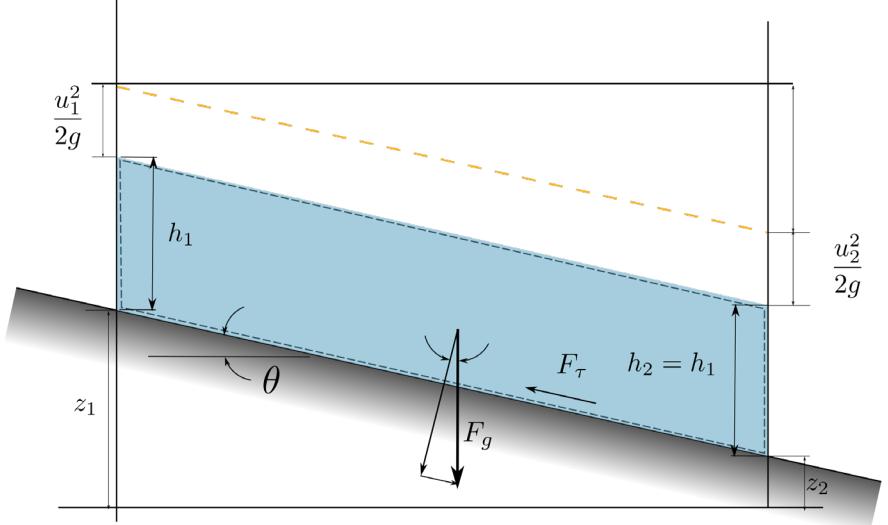


Particular case: uniform flow

$$\left( \cancel{\frac{u^2}{2g}} + \cancel{h} + z_b \right)_1 - \left( \cancel{\frac{u^2}{2g}} + \cancel{h} + z_b \right)_2 = S_f \Delta x$$

$$z_{b1} - z_{b2} = S_f \Delta x$$

$$\frac{z_{b1} - z_{b2}}{\Delta x} = S_f$$



# Uniform flow: normal depth

In the uniform flow

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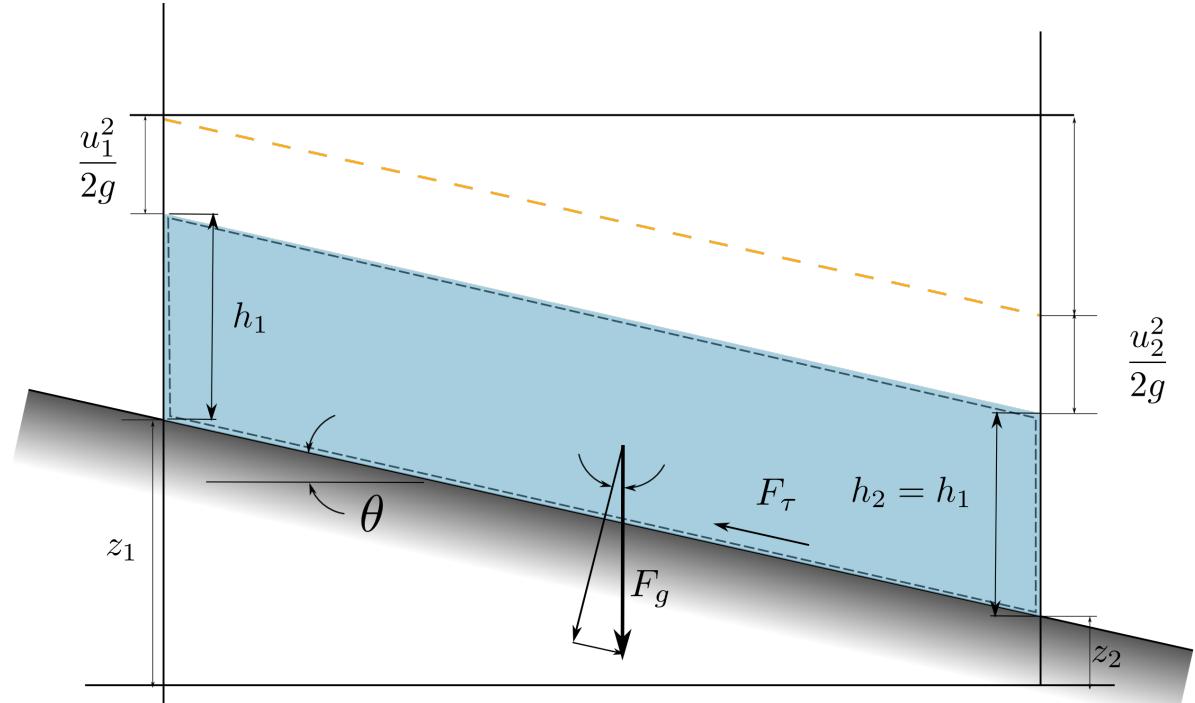
$$z_{b1} - z_{b2} = S_f \Delta x$$

$$\frac{z_{b1} - z_{b2}}{\Delta x} = S_f \quad \rightarrow \quad S_0 = S_f$$

Normal depth,  $h_N$ , ensures equality

$$S_0 = S_f = \frac{n^2 Q |Q|}{R_h^{4/3} A^2}$$

Where  $A=A(h)$ ,  $R_h=R_h(h)$



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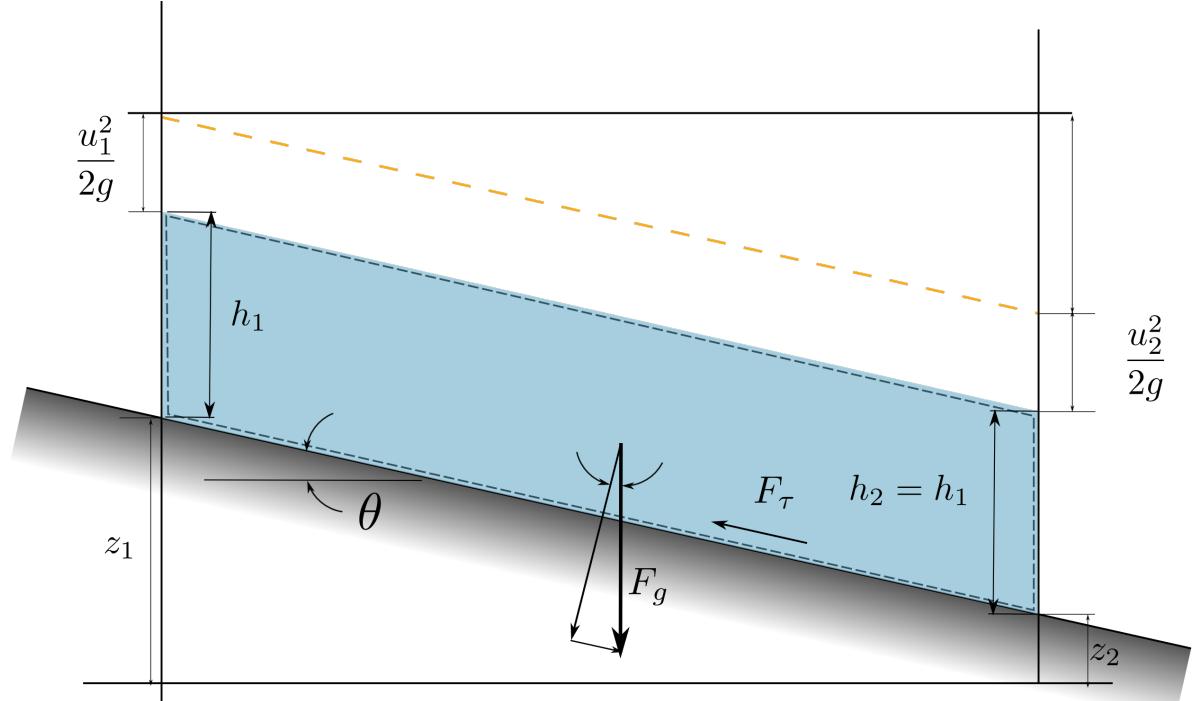
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Note: If  $h \uparrow\uparrow \rightarrow S_f \downarrow\downarrow$

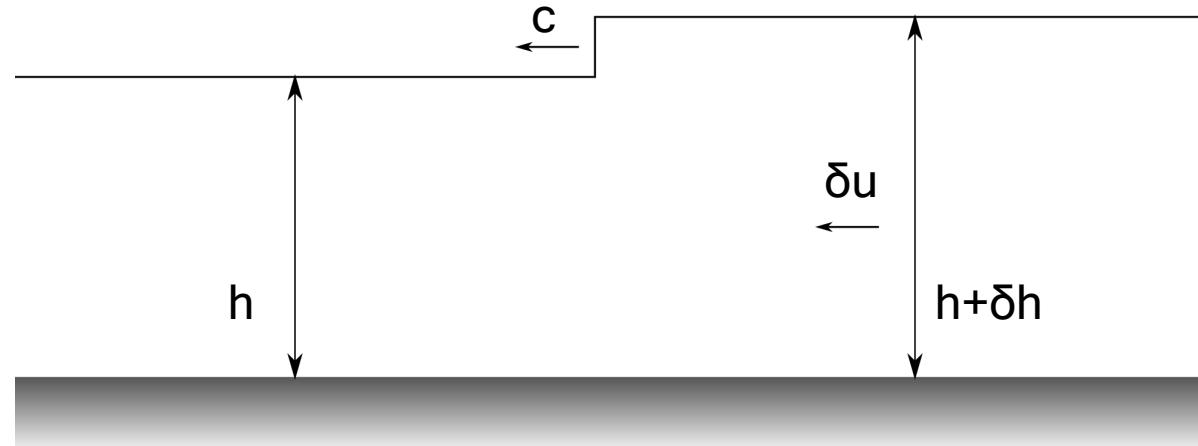
Therefore:

If  $h > h_N \rightarrow S_0 > S_f$   
If  $h < h_N \rightarrow S_0 < S_f$

# Gravity waves

- The velocity of infinitesimal surface waves

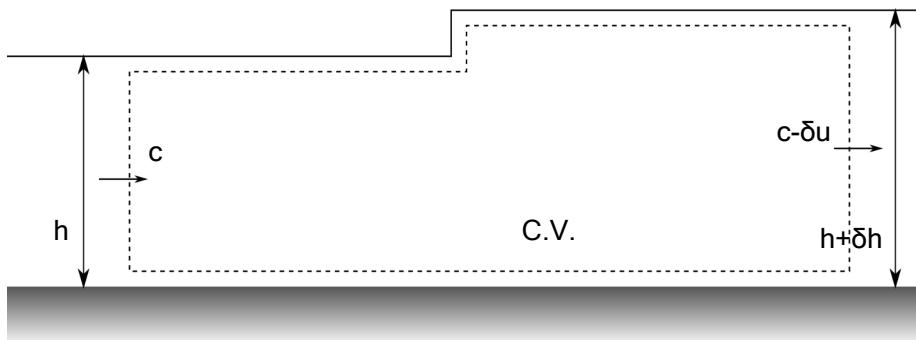
$$\delta h \ll h \quad \Rightarrow \quad c \approx \sqrt{gh}$$



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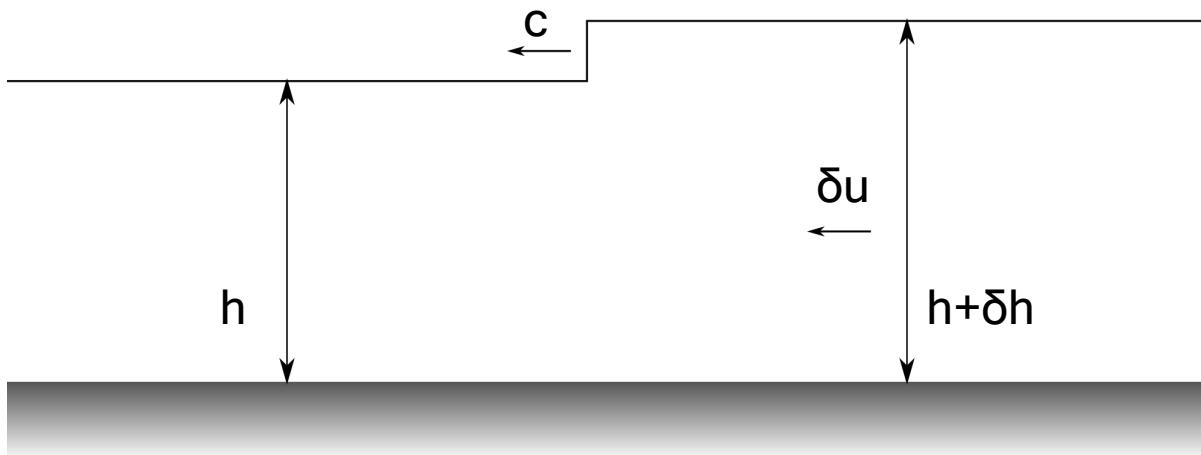
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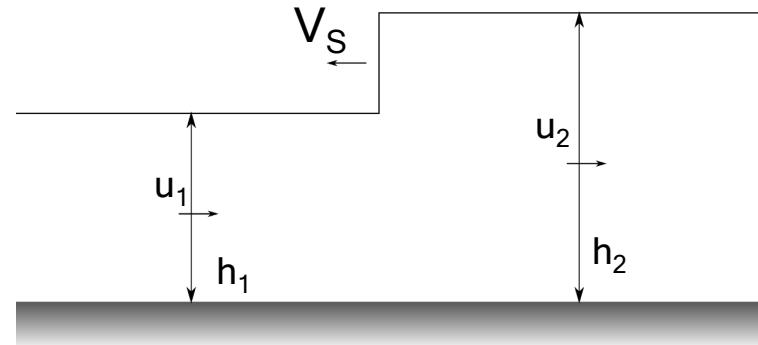
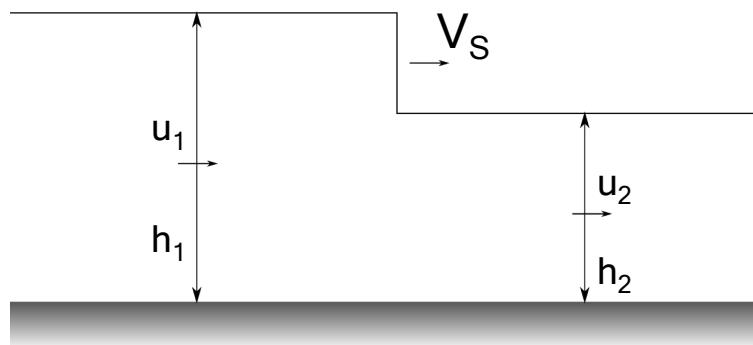
C is **not** the velocity of the flow

# Finite amplitude waves

- The velocity of surface waves

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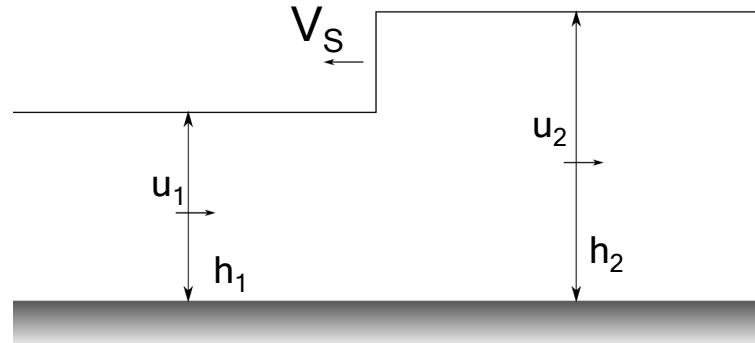
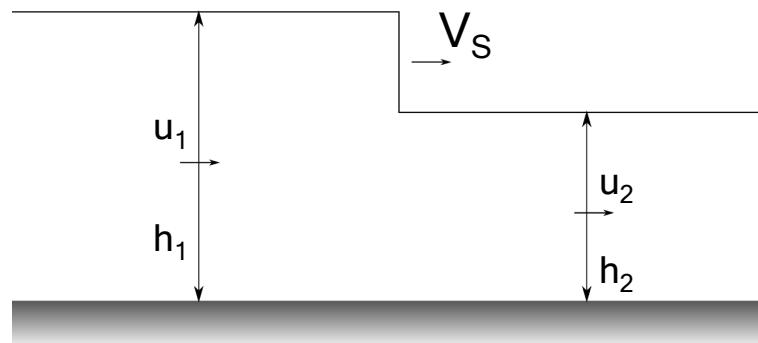


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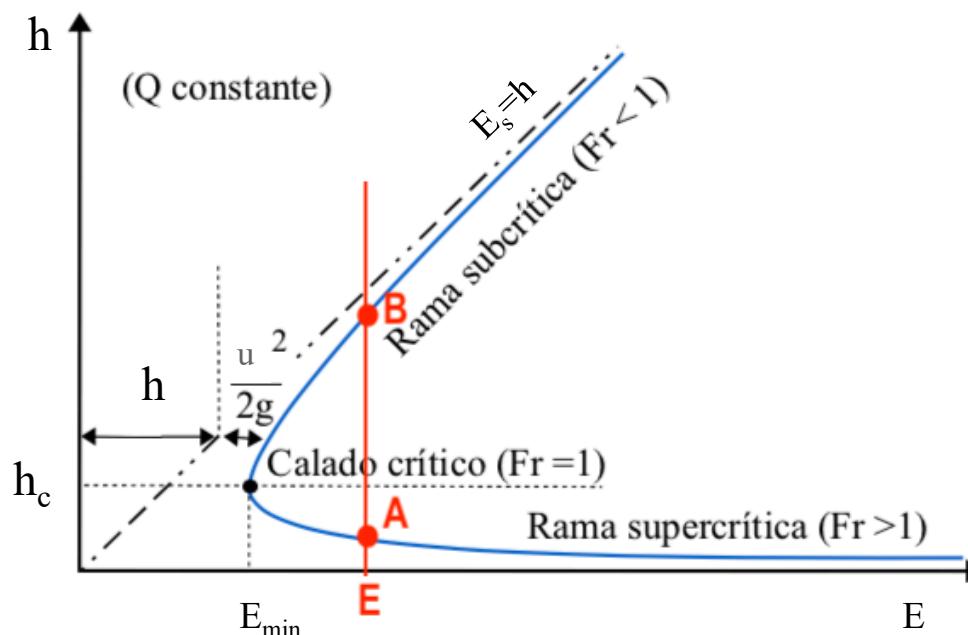
# Specific energy

- Total energy:  $H = E + z_b$
- $E$  is the sum of the kinetic energy and the pressure energy (represented by the depth)
- In a rectangular channel:

$$A=Bh$$

$$Q=qB$$

$$E = \frac{u^2}{2g} + h = \frac{Q^2}{2gB^2h^2} + h = \frac{q^2}{2gh^2} + h$$



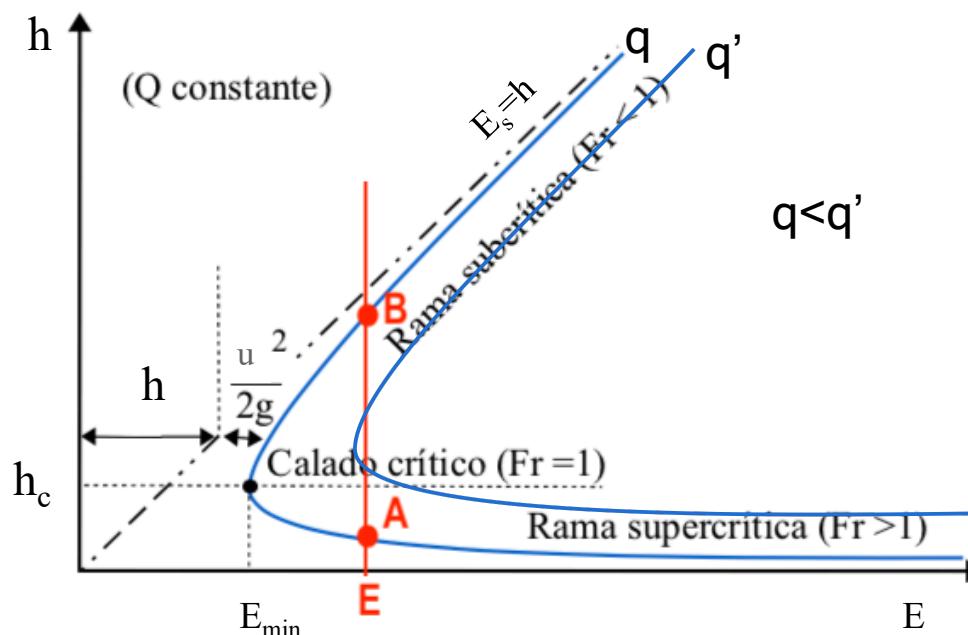
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# Critical depth

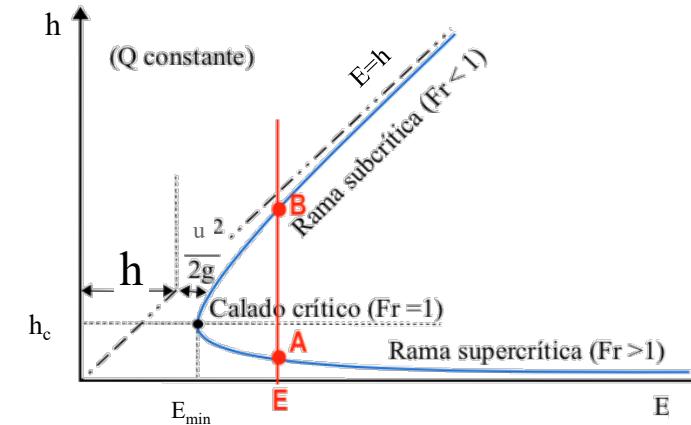
Minimum point of the curve  $E=E(h)$ :  $\frac{dE}{dh} = 0$

$$h_c = \left( \frac{Q^2}{gB^2} \right)^{1/3} \quad E_c = \frac{3}{2} h_c = \frac{3}{2} \left( \frac{Q^2}{gB^2} \right)^{1/3}$$

When  $dE/dh=0$

$$\frac{dE}{dh} = \frac{-Q^2 2h}{2gB^2 h^4} + 1 = 0$$

$$\frac{Q^2 h}{gB^2 h^4} = 1$$



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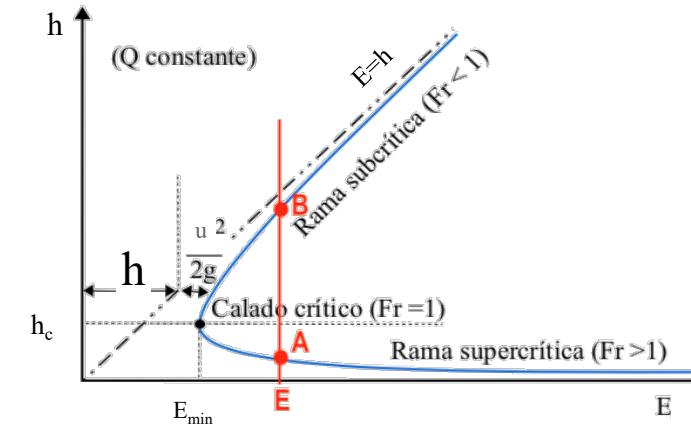
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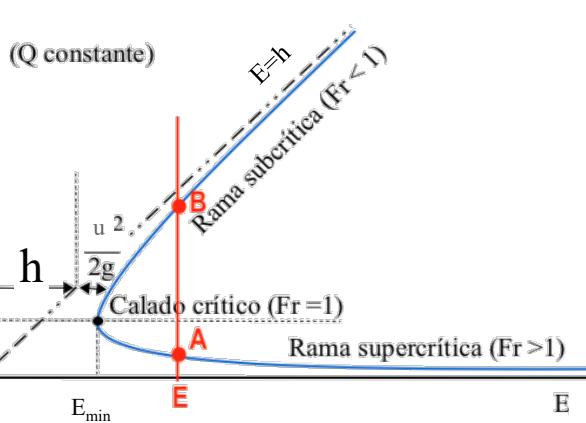
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$$Fr = \frac{u}{\sqrt{gh}}$$

$$\frac{Q^2 h}{gB^2 h^4} = 1 \quad \rightarrow \quad \frac{Q^2 / B^2 h^2}{gh} = 1 \quad \rightarrow \quad \frac{u^2}{gh} = 1 \quad \xrightarrow{\text{downward arrow}}$$

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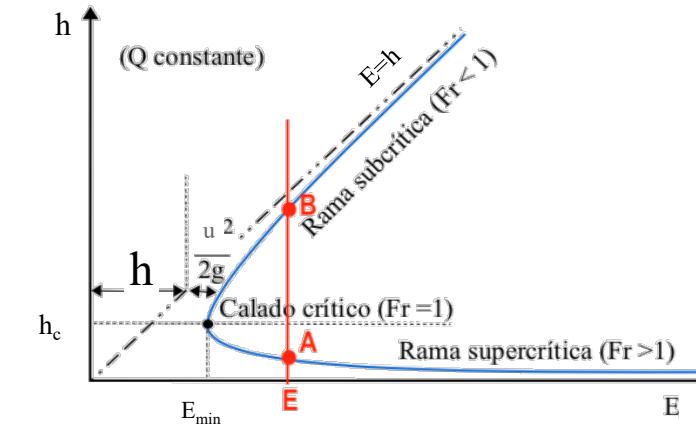
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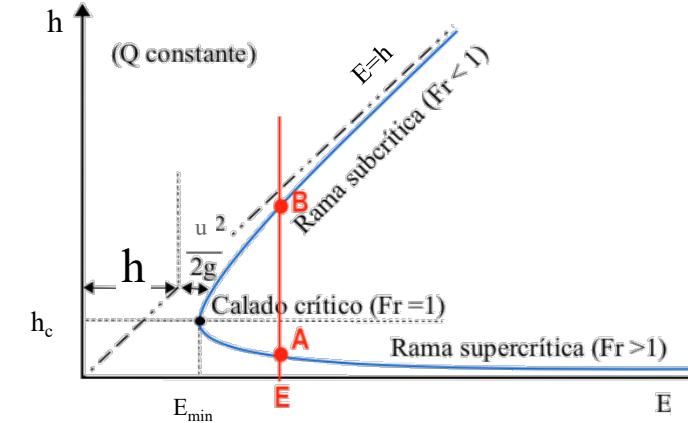
$$\frac{u^2}{gh} = 1 \rightarrow Fr = 1$$

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$Fr > 1$  supercritical

$$\text{Froude number, } Fr = \frac{u}{\sqrt{gh}} = \frac{u}{c} = 1$$

$Fr < 1$  subcritical

- Subcritical flow,  $u < c$  ( $Fr < 1$ ) controlled by downstream conditions
- Supercritical flow,  $u > c$  ( $Fr > 1$ ) controlled by upstream conditions

# Equilibrium profiles

The behaviour of the free surface can be calculated  
Starting from the energy balance between two points

$$\left( \frac{u^2}{2g} + h + z_b \right)_1 - \left( \frac{u^2}{2g} + h + z_b \right)_2 = S_f \Delta x$$

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Knowing :  $Q=uA$

In rectangular channel:  $A=bh$

$$\frac{d}{dx} \left( \frac{Q^2}{2gB^2h^2} + h + z_b \right) = -S_f$$

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And rearranging and replacing definitions

$$\frac{dh}{dx} (1 - Fr^2) + \frac{dz_b}{dx} = -S_f \quad \rightarrow \quad \boxed{\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}}$$

# Equilibrium profiles

Variation of the flow surface:

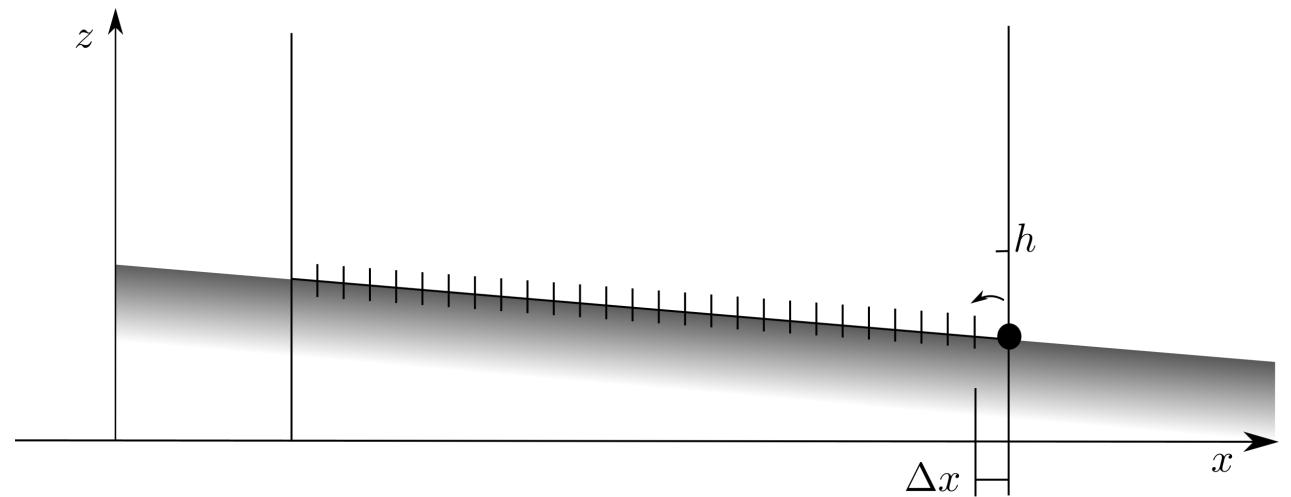
$$\boxed{\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}} \rightarrow \frac{\Delta h}{\Delta x} = \frac{S_0 - S_f}{1 - Fr^2}$$

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$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2} \rightarrow \frac{\Delta h}{\Delta x} = \frac{S_0 - S_f}{1 - Fr^2}$$

Trend analysis of the flow surface slopes



# Equilibrium profiles

Variation of the flow surface:

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

There are two borderline cases:

- $h = h_N \rightarrow S_0 = S_f \rightarrow \frac{dh}{dx} = 0$

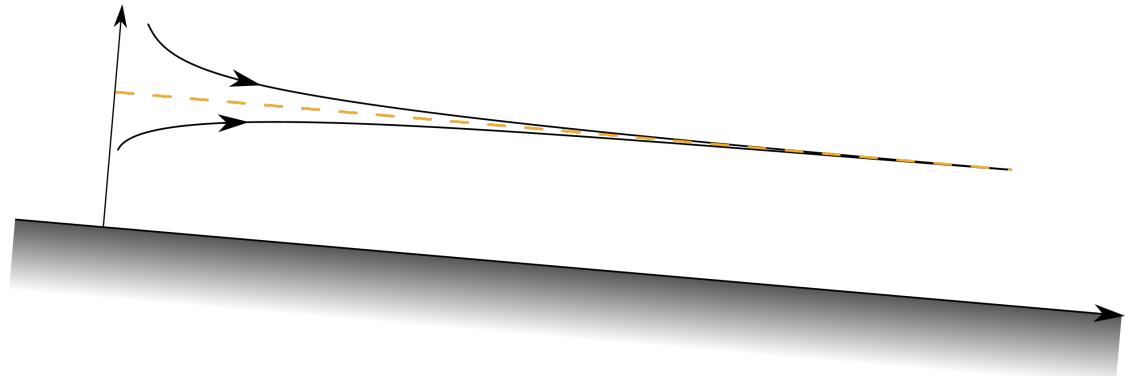
# Equilibrium profiles

Variation of the flow surface:

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

There are two limit cases:

- $h = h_N \rightarrow S_0 = S_f \rightarrow \frac{dh}{dx} = 0$



# Equilibrium profiles

Variation of the flow surface:

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

There are two limit cases:

- $h = h_N \rightarrow S_0 = S_f \rightarrow \frac{dh}{dx} = 0$

$$\frac{dh}{dx} = \infty$$

# Equilibrium profiles

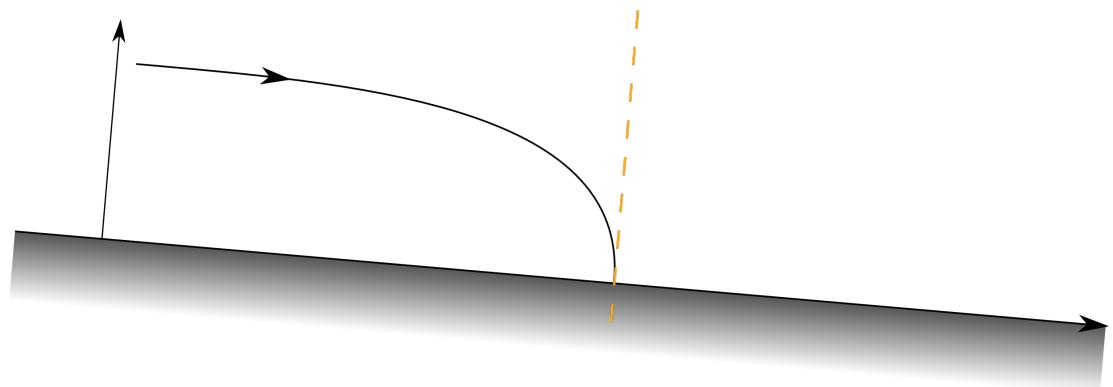
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# Equilibrium profiles

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$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

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Two situations to be analyzed:

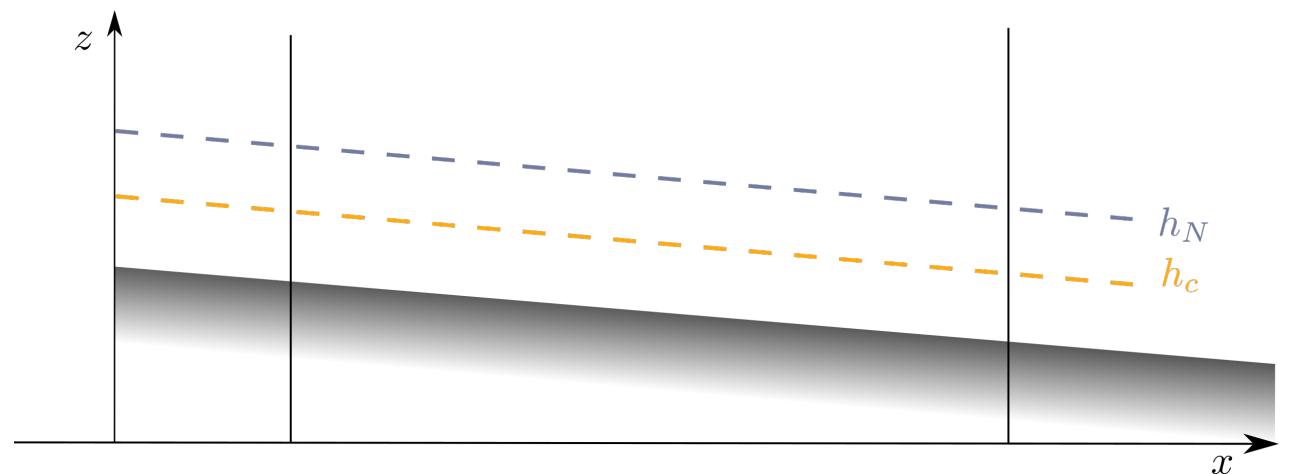
- Mild slopes:  $h_c < h_N$
- Steep slopes:  $h_c > h_N$

$$\frac{dh}{dx} = \infty$$

# Profiles with $h_c < h_N$

Surface variation of the flow on a **mild slope**

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$



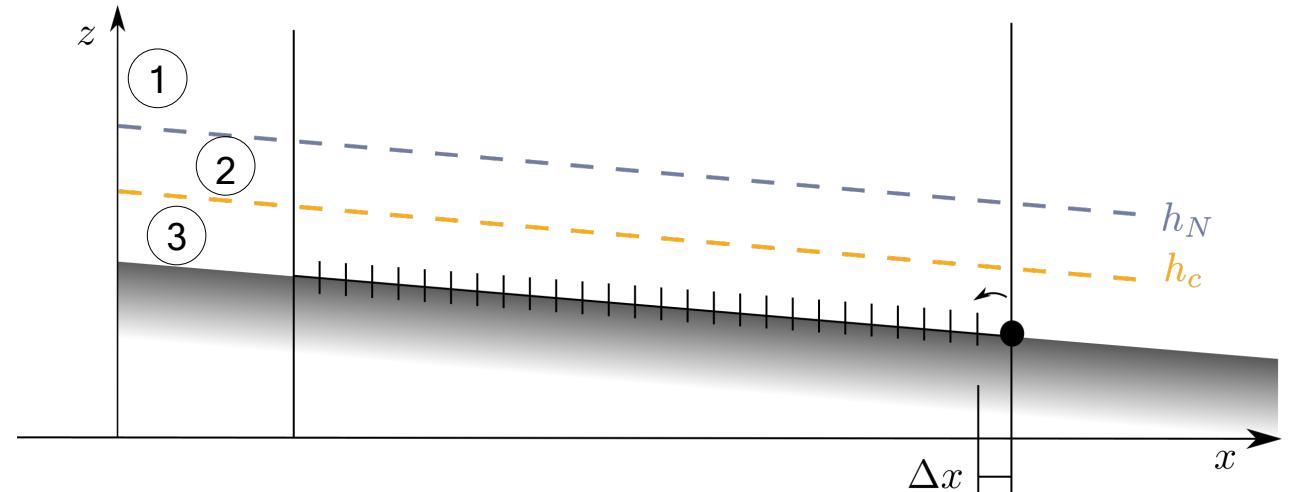
# Profiles with $h_c < h_N$

Surface variation of the flow on a **mild slope**

The equation is solved for each depth region

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

$$\frac{\Delta h}{\Delta x} = \frac{S_0 - S_f}{1 - Fr^2}$$



# Profiles with $h_c < h_N$

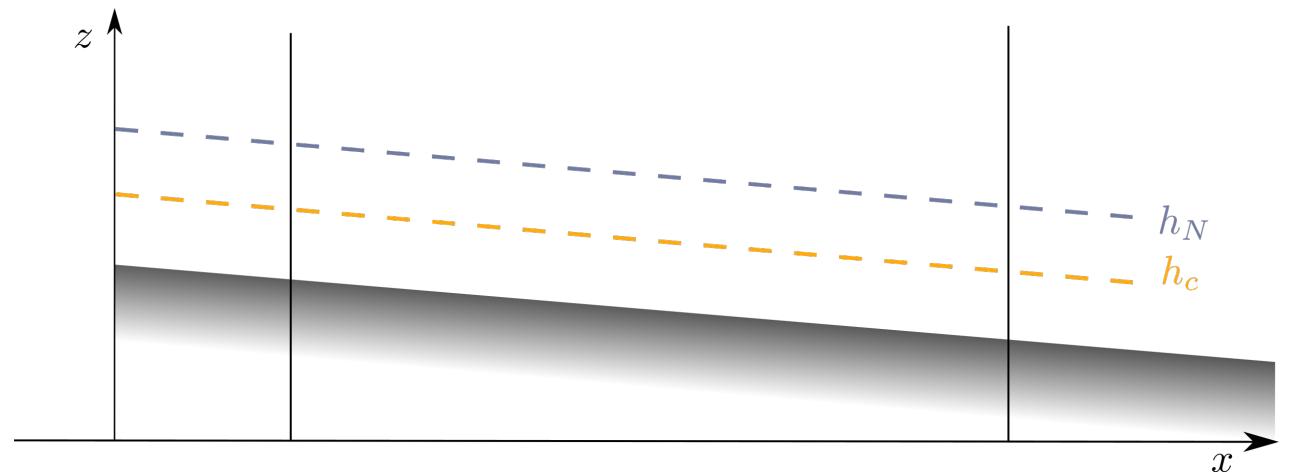
Surface variation of the flow on a **mild slope**

The equation is solved for each depth region

The evolution of the surface profile is analyzed

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

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# Profiles with $h_c < h_N$

Surface variation of the flow on a **mild slope**

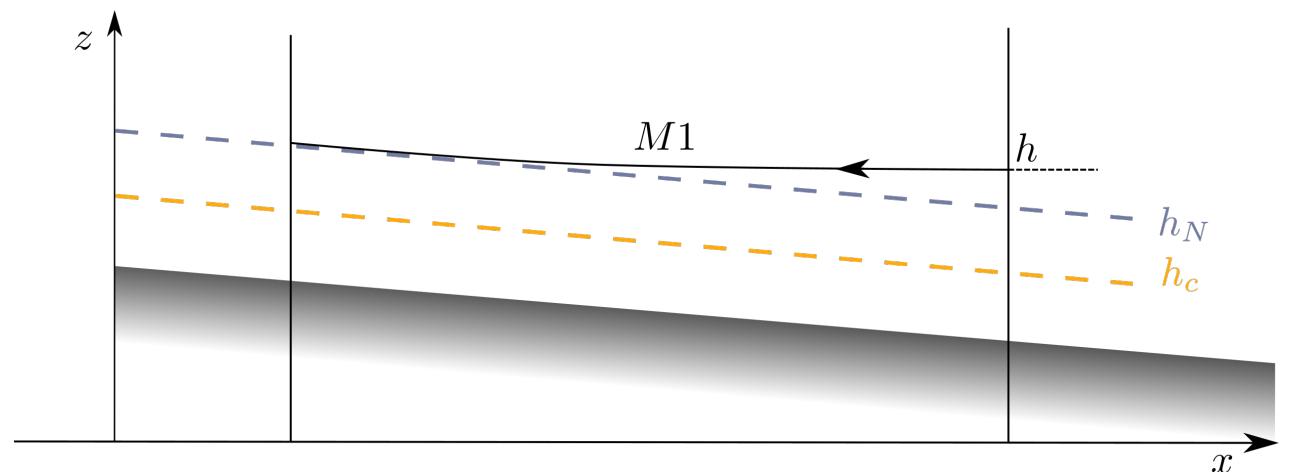
The equation is solved for each depth region

The evolution of the surface profile is analyzed

- **M1:** If  $h > h_N > h_c$ 
  - $S_0 > S_f, Fr < 1 \rightarrow \frac{dh}{dx} > 0$

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

$$\frac{\Delta h}{\Delta x} = \frac{S_0 - S_f}{1 - Fr^2}$$



# Profiles with $h_c < h_N$

Surface variation of the flow on a **mild slope**

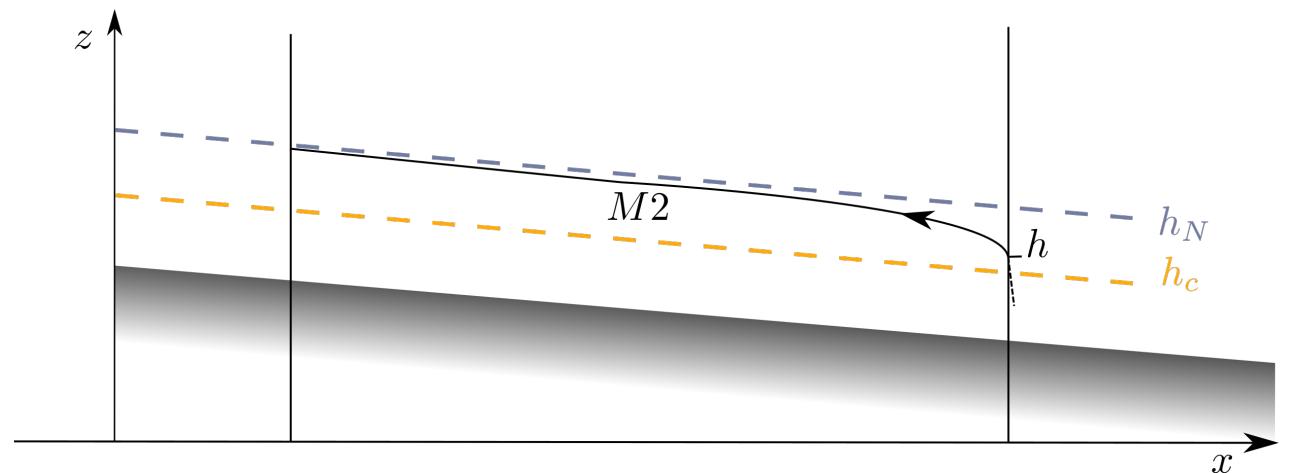
The equation is solved for each depth region

The evolution of the surface profile is analyzed

- **M1:** If  $h > h_N > h_c$ 
  - $S_0 > S_f, Fr < 1 \rightarrow \frac{dh}{dx} > 0$
- **M2:** If  $h_N > h > h_c$ 
  - $S_0 < S_f, Fr < 1 \rightarrow \frac{dh}{dx} < 0$

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

$$\frac{\Delta h}{\Delta x} = \frac{S_0 - S_f}{1 - Fr^2}$$



# Profiles with $h_c < h_N$

Surface variation of the flow on a **mild slope**

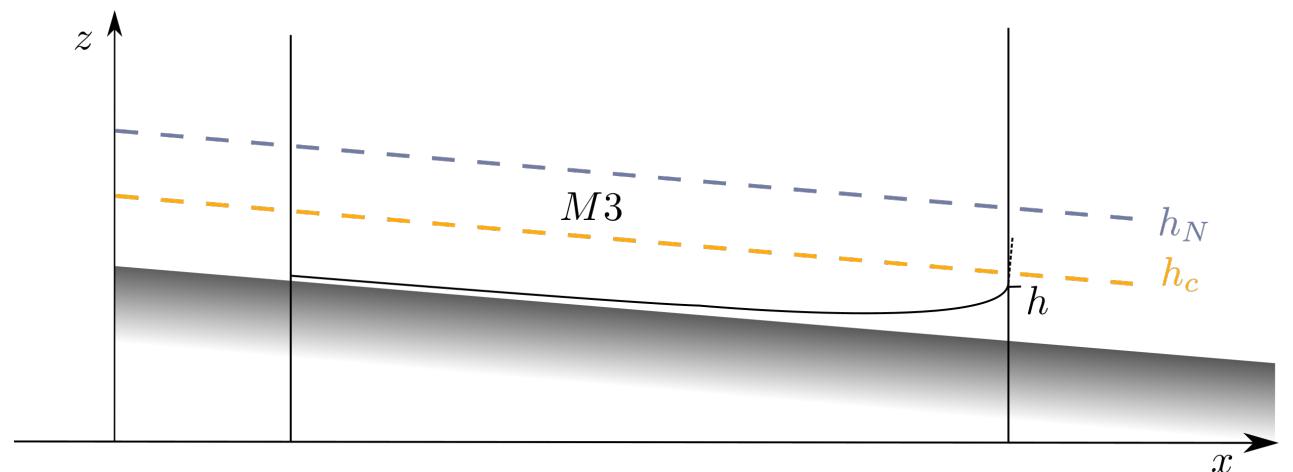
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- **M1:** If  $h > h_N > h_c$ 
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  - $S_0 < S_f$ ,  $Fr < 1 \rightarrow \frac{dh}{dx} < 0$
- **M3:** If  $h_N > h_c > h$ 
  - $S_0 < S_f$ ,  $Fr > 1 \rightarrow \frac{dh}{dx} > 0$

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

$$\frac{\Delta h}{\Delta x} = \frac{S_0 - S_f}{1 - Fr^2}$$



# Profiles with $h_c > h_N$

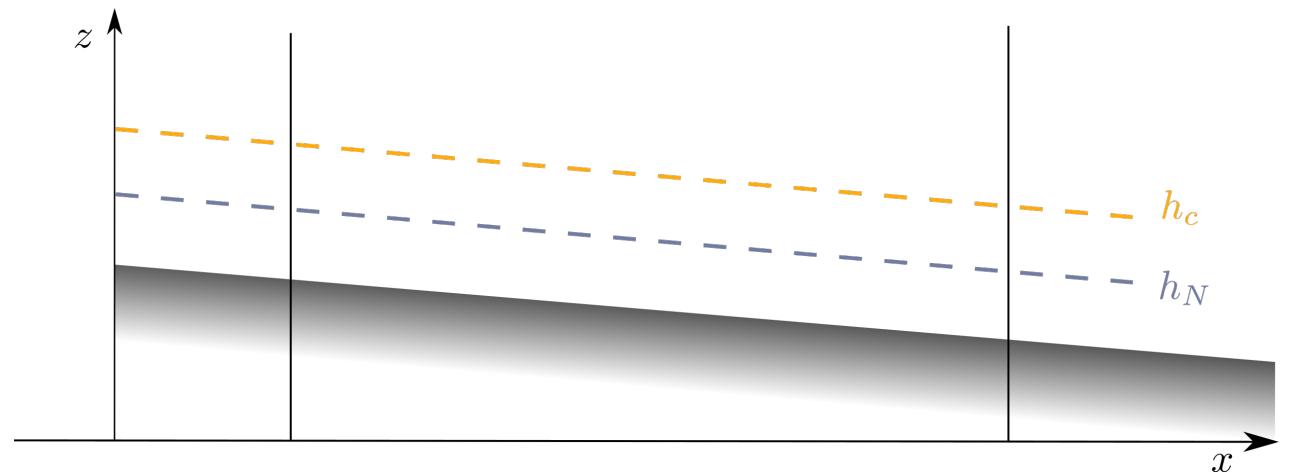
Surface variation of the flow on a **steep slope**

The equation is solved for each depth region

The evolution of the surface profile is analyzed

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

$$\frac{\Delta h}{\Delta x} = \frac{S_0 - S_f}{1 - Fr^2}$$



# Profiles with $h_c > h_N$

Surface variation of the flow on a **steep slope**

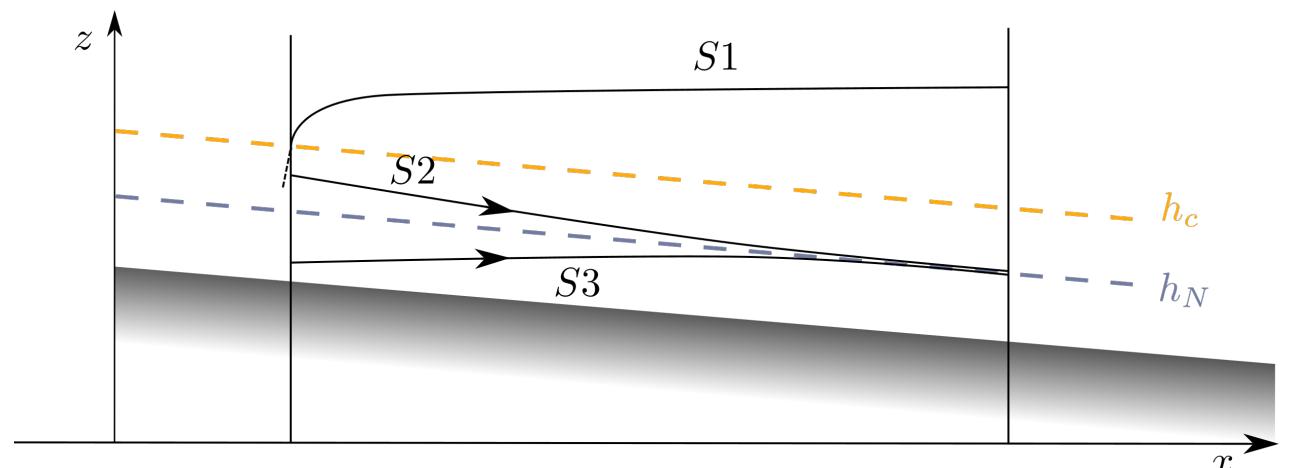
The equation is solved for each depth region

The evolution of the surface profile is analyzed

- **S1:** If  $h > h_c > h_N$ 
  - $S_0 > S_f, Fr < 1 \rightarrow \frac{dh}{dx} > 0$
- **S2:** If  $h_c > h > h_N$ 
  - $S_0 > S_f, Fr > 1 \rightarrow \frac{dh}{dx} < 0$
- **S3:** If  $h_c > h_N > h$ 
  - $S_0 < S_f, Fr > 1 \rightarrow \frac{dh}{dx} > 0$

$$\frac{dh}{dx} = \frac{S_0 - S_f}{1 - Fr^2}$$

$$\frac{\Delta h}{\Delta x} = \frac{S_0 - S_f}{1 - Fr^2}$$



# Hydraulic jump

- Is the physical phenomenon that occurs when the regime changes from supercritical to subcritical.
- The balance profile gives an infinite slope that shows up as a discontinuity in the depth.
- Energy losses can be quantified according to the depths on either side of the jump.



# Steady case: the hydraulic jump

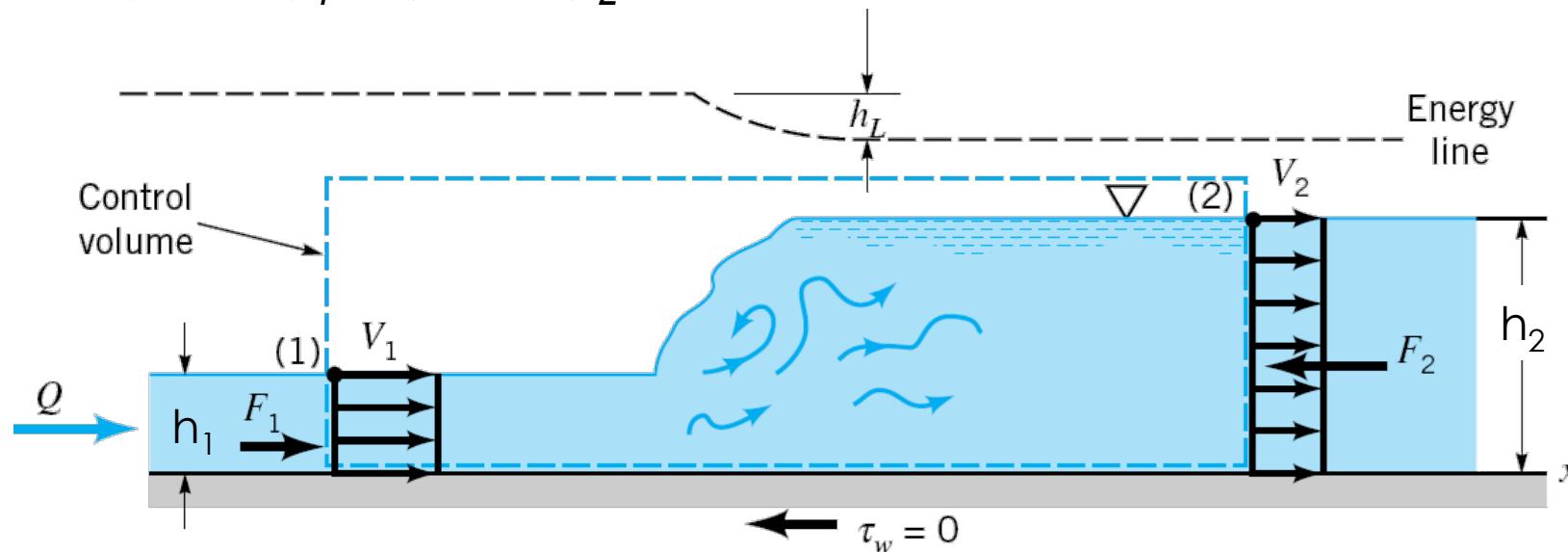
$$Q = V_1 h_1 B = V_2 h_2 B$$

$$Q(V_2 - V_1) = \left( \frac{1}{2} g h_2^2 - \frac{1}{2} g h_1^2 \right) B$$

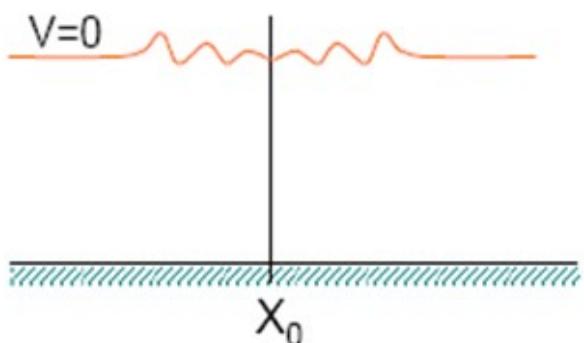
$$\left( \frac{V^2}{2g} + h \right)_1 = \left( \frac{V^2}{2g} + h \right)_2 + \Delta H$$

$$\frac{h_2}{h_1} = \frac{1}{2} \left( \sqrt{1 + 8 Fr_1^2} - 1 \right)$$

$$\Delta H = \frac{(h_2 - h_1)^3}{4h_1 h_2}$$



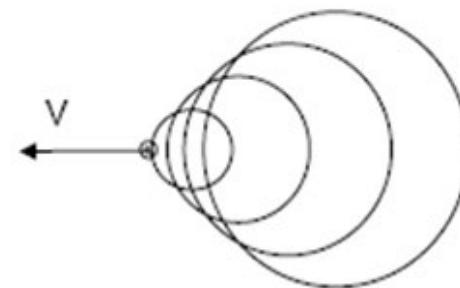
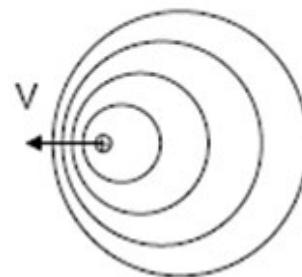
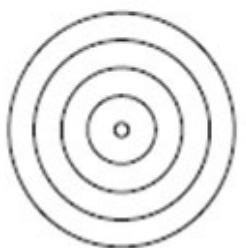
# Subcritical/critical flow



Marque urbaine (Roubaix)

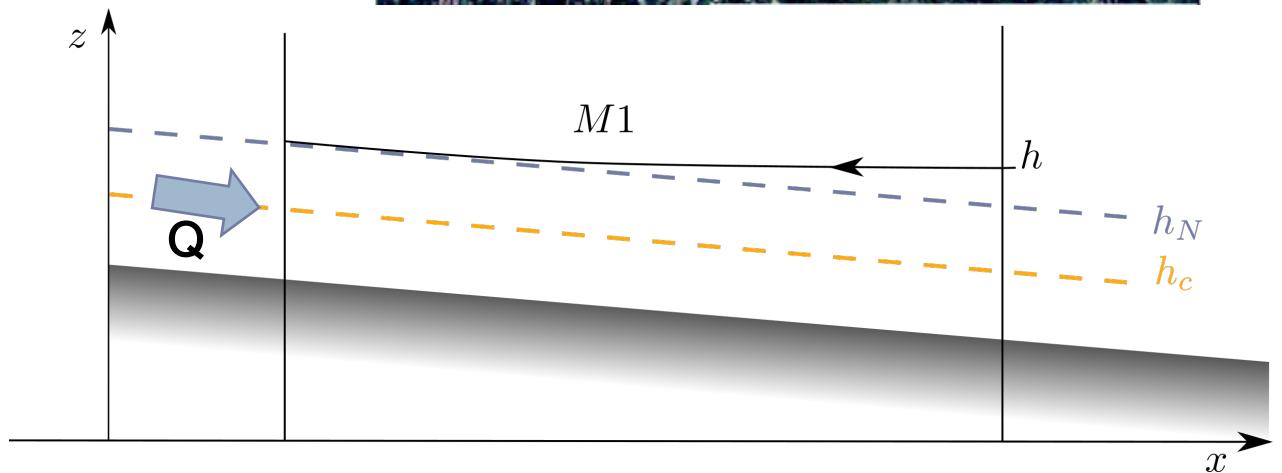


<http://www.eng.vt.edu/fluids/msc/gallery/gall.htm>



# Subcritical flow

- Subcritical flow:
- Decelerated flow
- Low velocity
- Under the influence of the downstream condition
- Backwater curves

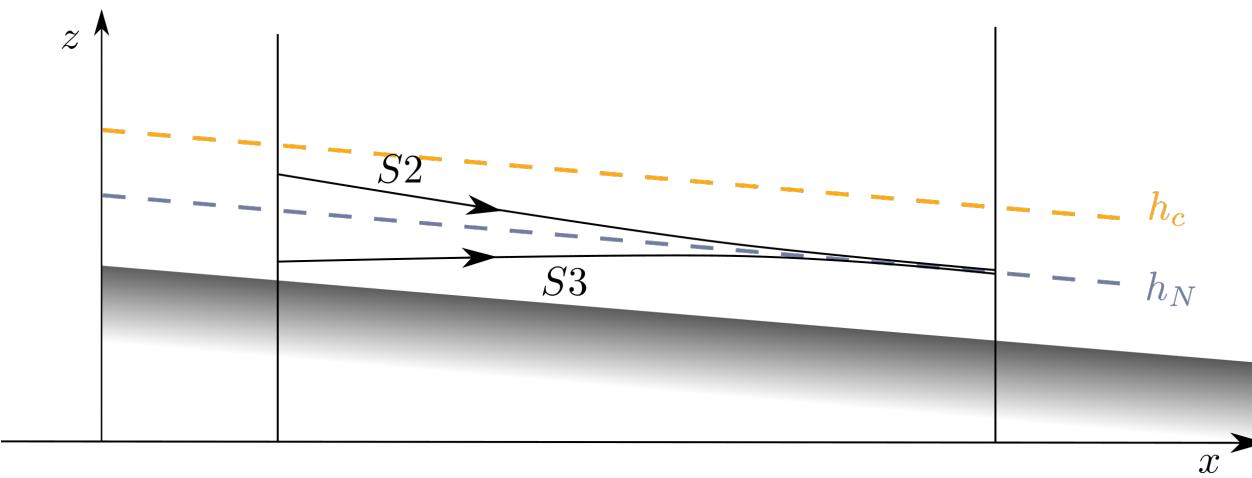


# Supercritical flow

- Supercritical flow:
- Accelerated flow
- High velocity
- Under the influence of the upstream condition



doc.Brett Sanders – <http://www.eng.uci.edu/~bfs/oblique.jpg>



# Variations of $h$ with $E = \text{cte}$

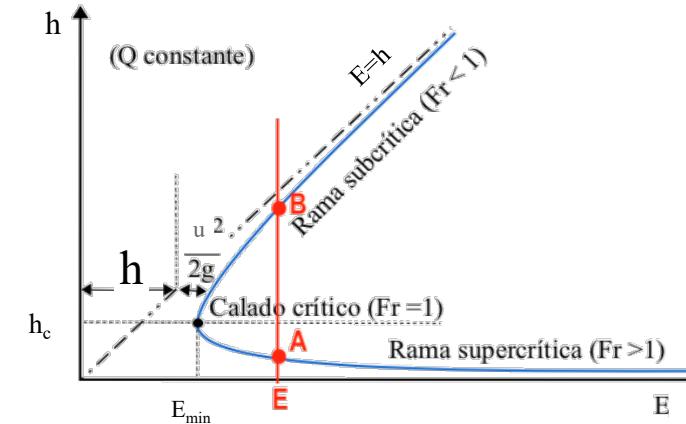
Constant flow rate:

$$Q = qB = \text{cte} \Rightarrow 0 = qdB + Bdq$$

Constant energy:

$$E = \frac{q^2}{2gh^2} + h = \text{cte} \Rightarrow 0 = \frac{q}{gh^2} dq + \left(1 - \frac{q^2}{gh^3}\right) dh$$

$$\left(1 - Fr^2\right) \frac{dh}{dx} = Fr^2 \frac{h}{B} \frac{dB}{dx}$$



**Slow flows ( $Fr < 1$ ):**

- If  $B$  decreases  $h$  also decreases
- If  $B$  increases,  $h$  also increases

**Fast flows ( $Fr > 1$ ):**

- If  $B$  decreases  $h$  increases
- If  $B$  increases  $h$  decreases

# Content

1. Model hypothesis: basic concepts in channels

2. Steady flow analysis

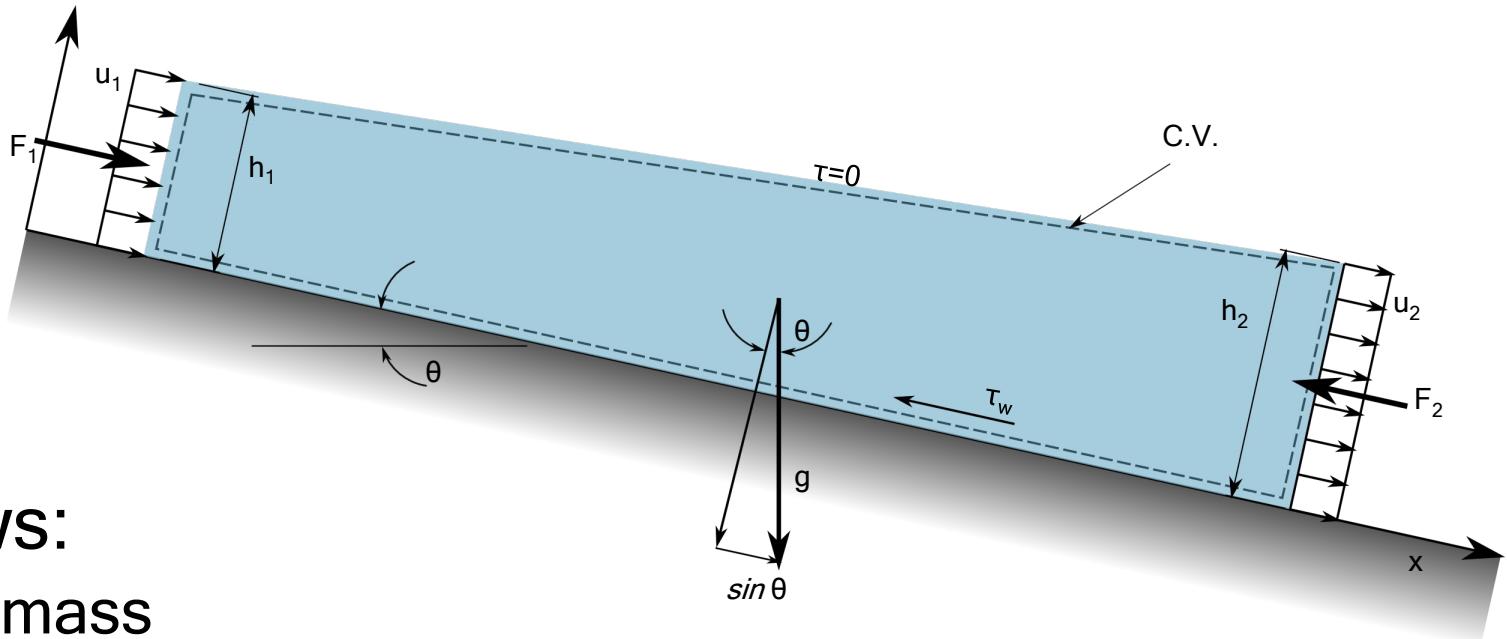
1. Energy balance
2. Uniform flow: Normal Depth
3. Energy: Critical Depth
4. Boundary conditions

3. Complete flow equations

1. Conservation laws
2. Resolution

# Motion equations: conservation laws

- Applied forces:



- Conservation laws:
  - Conservation of mass
  - Conservation of linear momentum
  - Conservation of energy

# Conservation of mass

- The volume of liquid passing through the inlet section per unit of time equals the volume leaving the section,

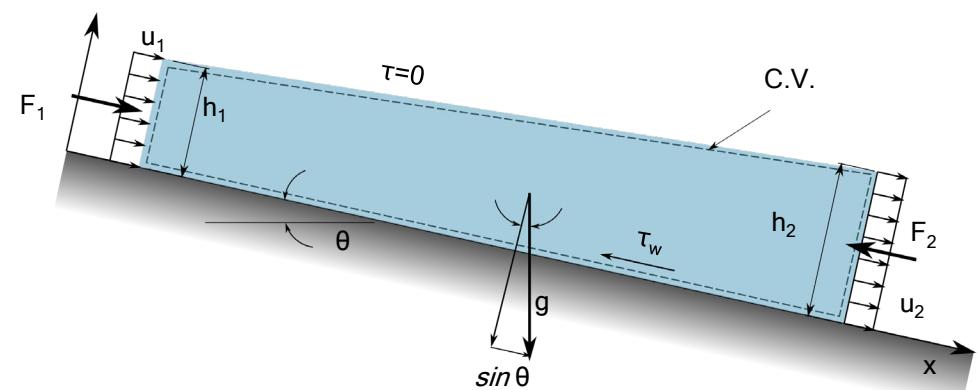
$$Q = \int_{A_e} u dA = \int_{A_s} u dA$$

- The volumetric flow or flow rate through a section is often described as a function of an average velocity over the section

$$Q = u_e A_e = u_s A_s$$

- In general, if the movement is transient

$$\frac{\partial A}{\partial t} + \frac{\partial Q}{\partial x} = 0$$



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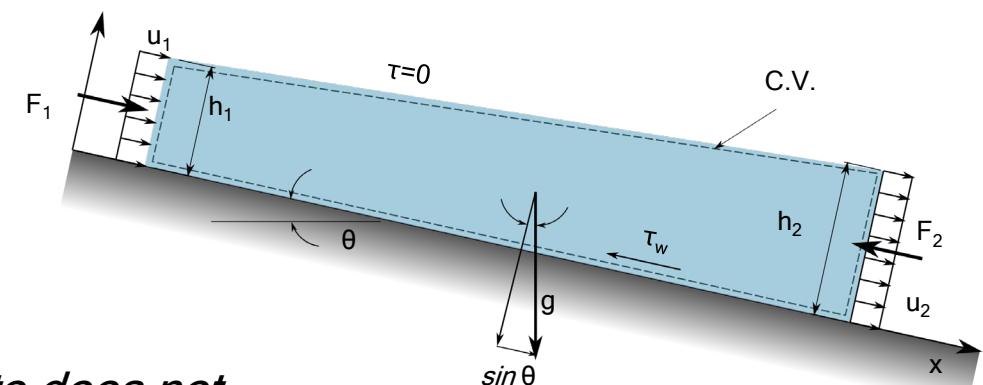
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$$Q = u_e A_e = u_s A_s$$

- In general, if the movement is transient

$$\frac{\partial A}{\partial t} + \frac{\partial Q}{\partial x} = 0$$

*The flow rate does not vary along the channel*



# Conservation of linear momentum

- Newton's second law:

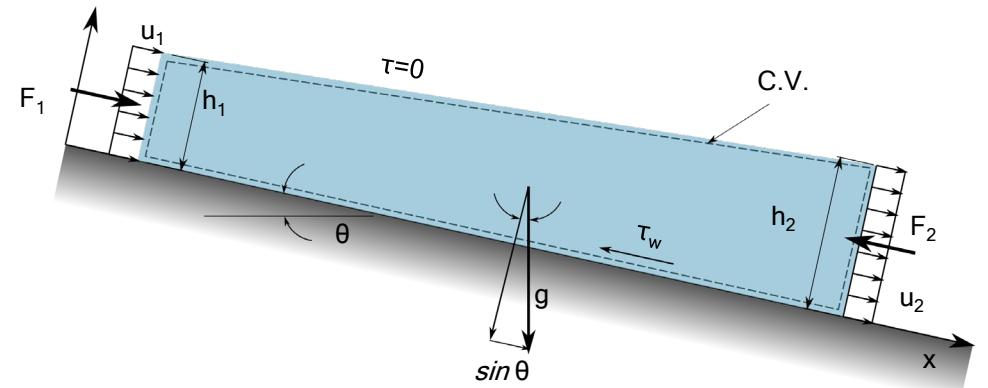
$$\rho \mathbf{a} = \frac{D(\rho \mathbf{V})}{Dt} = \rho \mathbf{F}_V - \nabla p + \mathbf{F}_{vis}$$

- Expressed in the direction of the channel:

$$\rho \left( \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} \right) = -\rho g \underbrace{\frac{\partial z}{\partial x}}_{\text{gravity}} - \underbrace{\frac{\partial p}{\partial x}}_{\text{pressure}} + \underbrace{F_{vis}}_{\text{friction}}$$

- Integrated over the section A:

$$\boxed{\frac{\partial Q}{\partial t} + \frac{\partial}{\partial x} \left( \frac{Q^2}{A} + gI_1 \right) = gA(S_0 - S_f) + gI_2}$$



# Conservation of linear momentum

- Newton's second law states:

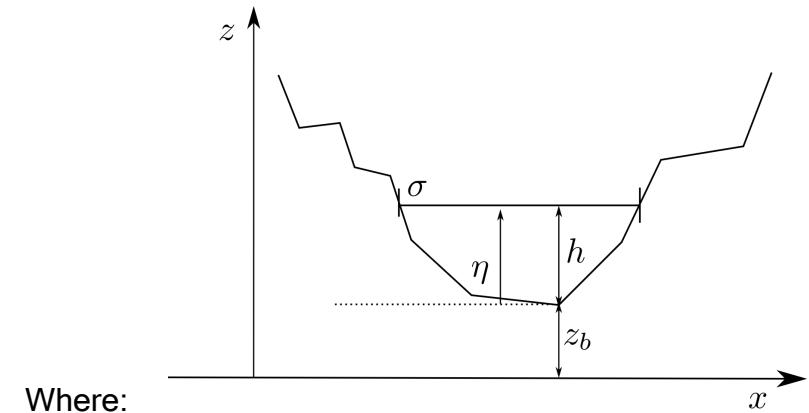
$$\rho \mathbf{a} = \frac{D(\rho \mathbf{V})}{Dt} = \rho \mathbf{F}_V - \nabla p + \mathbf{F}_{vis}$$

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$$\rho \left( \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} \right) = - \underbrace{\rho g \frac{\partial z}{\partial x}}_{\text{gravity}} - \underbrace{\frac{\partial p}{\partial x}}_{\text{pressure}} + \underbrace{F_{vis}}_{\text{friction}}$$

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$$\boxed{\frac{\partial Q}{\partial t} + \frac{\partial}{\partial x} \left( \frac{Q^2}{A} + gI_1 \right) = gA(S_0 - S_f) + gI_2}$$



Where:

$$S_0 = -\frac{\partial z_b}{\partial x} = \operatorname{tg}\theta \quad \text{bed slope}$$

$$S_f = \frac{n^2 Q |Q|}{R_h^{4/3} A^2} \quad \text{Slope of the energy line}$$

$$I_1 = \int_0^h (h - \eta) \cdot \sigma(x, \eta) d\eta$$

$$I_2 = \int_0^h (h(x) - \eta) \frac{\partial \sigma(x, \eta)}{\partial x} d\eta$$

# Conservation of energy

$$\frac{1}{g} \frac{\partial u}{\partial t} + \frac{u}{g} \frac{\partial u}{\partial x} = - \frac{\partial h}{\partial x} + S_0 - S_f$$

$$S_0 = - \frac{\partial z_b}{\partial x} = \operatorname{tg} \theta \quad \text{bed slope}$$

$$S_f = \frac{n^2 Q |Q|}{R_h^{4/3} A^2} \quad \text{Slope of the energy line}$$

# Conservation of energy

In steady state there are no temporary variations

$$\frac{1}{g} \cancel{\frac{\partial u}{\partial t}} + \frac{u}{g} \frac{\partial u}{\partial x} = -\frac{\partial h}{\partial x} + S_0 - S_f$$

Applying the definition of  $S_0$ :  $S_0 = -\frac{\partial z_b}{\partial x}$

And rearranging...

$$\frac{1}{2g} \frac{\partial(u^2)}{\partial x} + \frac{\partial h}{\partial x} + \frac{\partial z_b}{\partial x} = -S_f$$

# Conservation of energy: stationary

The mechanical energy per weight unit varies due to viscous losses

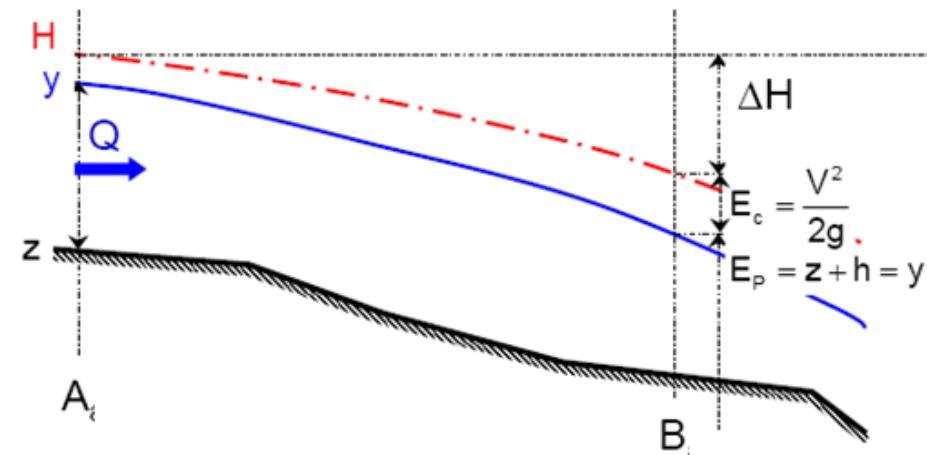
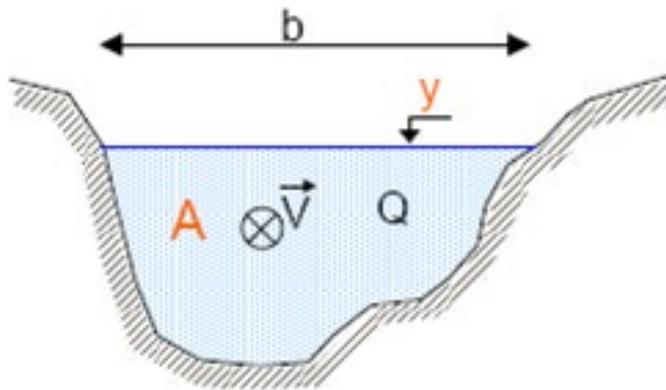
$$\frac{d}{dx} \left( \frac{u^2}{2g} + h + z_b \right) = -S_f$$

Evaluated between points 1 and 2

$$\left( \frac{u^2}{2g} + h + z_b \right)_1 - \left( \frac{u^2}{2g} + h + z_b \right)_2 = h_f = S_f \Delta x$$

# Saint-Venant equations

- Describe the behaviour of the fluid
- Hypothesis for one-dimensional flow:
  - Gentle bed slope
  - Depth and speed only function of  $(x,t)$
  - Uniform cross-sectional velocity



# Simulation model: equations

$$\frac{\partial A}{\partial t} + \frac{\partial Q}{\partial x} = 0$$

$$\frac{\partial Q}{\partial t} + \frac{\partial}{\partial x} \left( \frac{Q^2}{A} + gI_1 \right) = gA(S_0 - S_f) + gI_2$$

- We only need two equations to solve A and Q:
  - We choose mass and linear momentum
- Although in 1D it does not\* matter which one you choose, it is common to choose the momentum one for when the models extend to more dimensions

\*In reality the losses in the energy equation are more complicated to quantify and therefore even in 1D it is sometimes more convenient to choose the equation of linear momentum

$$S_0 = -\frac{\partial z_b}{\partial x} = tg\theta$$

$$S_f = \frac{n^2 Q |Q|}{R_h^{4/3} A^2}$$

$$I_1 = \int_0^h (h - \eta) \cdot \sigma(x, \eta) d\eta$$

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# Simulation model: equations

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$$\frac{\partial Q}{\partial t} + \frac{\partial}{\partial x} \left( \frac{Q^2}{A} + gI_1 \right) = gA(S_0 - S_f) + gI_2$$

- Without analytical solution
- Computer Simulation
- Discretization method of finite volumes

$$S_0 = -\frac{\partial z_b}{\partial x} = tg\theta$$

$$S_f = \frac{n^2 Q |Q|}{R_h^{4/3} A^2}$$

$$I_1 = \int_0^h (h - \eta) \cdot \sigma(x, \eta) d\eta$$

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# ADVANCED HYDRAULIC SIMULATION MODELS

## 1D hydraulic simulation models of channels and rivers: PART I

Pilar Garcia Navarro ([pigar@unizar.es](mailto:pigar@unizar.es))