






Hull_HW_09.md

- Quinn Hull
- Week of 04/05/2021
- River of Dreams


Questions

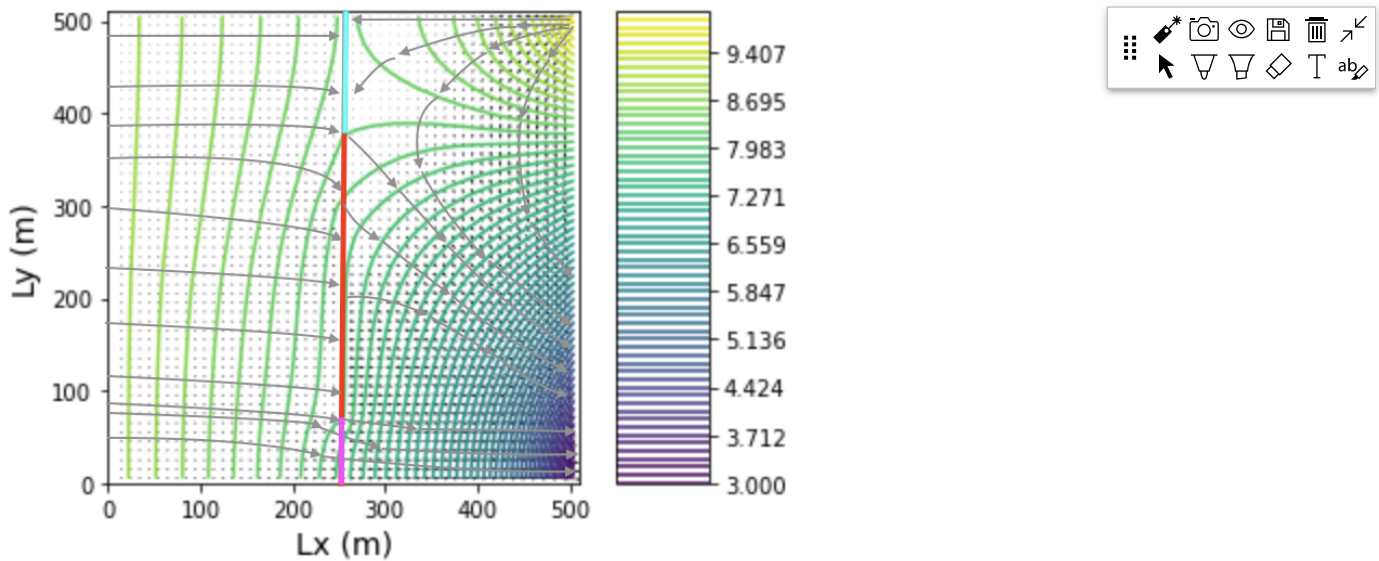
1. I'm a little confused about the usage of 'f' and 'm' in comments, it seems like maybe this refers to the units of length, where f = feet. But this is probably a typo... 
2. Given that the boundaries exist at the left and right sides, I don't really understand how this river originates and terminates at the north and south (so to speak) 
3. I'm not sure about how to interpret the final part of the rivers (3). It seems like there is no flow in the river, and yet the head in the cell is technically above streambed bottom. This could be a conceptual issue in my understanding of the problem. Or it could be that the model doesn't fully capture the dynamics of flow in a stream. Which is it? 

Model Characteristics

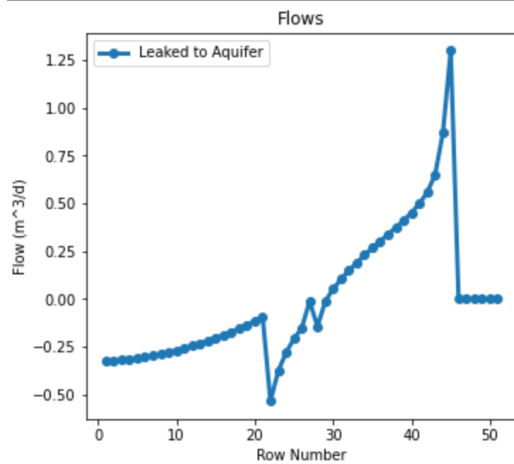
- Single Layer Model, 51x51x1 (each cell 10x10x10 m)
- Stream runs along center of domain (@ Col ~ 25)
 - K changes within three zones along the stream
 - rows 0-20 - low K (1/100)
 - rows 21-26 - medium K (1/10)
 - rows 27-50 - hi K (1)
 - Inflow to first reach = 0 m³/d
 - Streambed slope = 0.0001 m/m
 - Streambed rough = 0.02377
 - Streambed width = 20 m
 - Streambed thickness = 1 m
 - Stream bottom depth = 7 m
- Unconfined Aquifer
- Recharge:
 - Uniform 5e⁻⁵ rows 0-26
 - Uniform 0 rows 27-51
- No ET
- Boundaries:
 - North and South = No flow...
 - Left = Constant Lake Head (=9)
 - Right = Constant Lake Head (=np.linspace(10,3,51))

The Challenges

- a) The code is provided to produce the first set of 'correct' figures. Use these figures to describe the nature (direction/magnitude) of stream/aquifer exchange along the stream. In particular, explain why the leakage changes magnitude or direction where these values change.
- b) Use the head distribution to describe the movement of water across the boundaries and into/out of the stream.
 - For me, it is easier to understand the the nature of stream/aquifer exchange by looking at a more detailed figure of the contour map. 



- Here, the river (column 25) is divided into three parts:
 - Fully gaining - rows 0-12, color sky blue.
 - Partially gaining / losing - rows 12-45, color red.
 - Not flowing - rows 45 - 50, color pink.
- In the fully gaining portion (1), groundwater enters the river from both left and right sides of the domain. (Direction = into River; Magnitude = large on L, large and decreasing on R).
 - As the right boundary decreases in head from 10 to 3, groundwater flow on the right side of the domain starts getting rerouted from the central river back to the right boundary (which experiences a far larger drop in head over its reach). We can show that when the contour line ~8.2 crosses the central river, river begins entering the groundwater system on the right side. This is shown in an accompanying decrease in the rate of increase of flow at that point.
 - It's important to note that the amount of water leaving the domain is buffered by 1) the relatively low conductivity of the sediment material, and 2) the dependable inflow from the left side of the domain, so flow continues to increase along the river.
- In the partially gaining / losing portion (2), flow enters the river from the left side, but leaves the river from the right side (Direction on left = into River; Magnitude on left = large and ~ constant; Direction on right = out of river; magnitude on right = small and increasing)
 - For the first part of this reach, flow is net 'gaining'. We see flow in the river increase from rows 12 to ~30, and we see some important piecewise changes at row = 21, and 26. That change is when the streambed K increases by an order of magnitude to 1/10 and then to 1. With each piecewise change, the river flow rate of change temporarily increases before decreasing - probably a result of the groundwater having less difficulty getting into the river system. But the sharp decrease in head along the right boundary dominates. We can think of this stretch between 12 and ~30 as stretch where more water enters the stream than leaves.
 - Now the second part of the reach, flow is net 'losing'. From row ~30 to row ~45, flow in the river monotonically drops ultimately to 0. An interesting point occurs at ~row 26 (when the streambed K = 1). The river head begins to track at head = 8 until row 45, when flow in the river goes to 0. We can think of this stretch between ~26 and ~45 as when enough water enters the stream to keep the head constant, but more water leaves the stream than enters it, eventually to the point where there ceases to be flow in the river (row 45). I'm not sure how realistic this part of the domain is, @Ty. It seems to me there should be a relationship between the flow in the river and its stage.
 - I think that the period where the head in the cell stays constant at 8 is because for the groundwater system it is essentially 'equally easy' for water to enter and the river as it is for it to stay in the groundwater system.
- In the not flowing part of the river (3), groundwater flow from left to right does not enter into or leave the river. (Direction = none, magnitude = 0). Groundwater is actually going beneath the stream channel. We could think of this as a groundwater system that is no longer intersecting the land surface, and so does not provide water for flow. We can see that point shown in the figure below.



• Other thoughts:

- An important facet that I have not explored is - how does the recharge rate affect these patterns?
 - Recharge probably props up the groundwater system in the upper portion of the reach
- We can actually see the 'net change' in water within the stream at each cell by looking at the magnitude and sign of the difference between the 'flow in' and 'flow out' lines.
- My guess is that the thing that will most impact stream behavior is changing the right boundary condition.

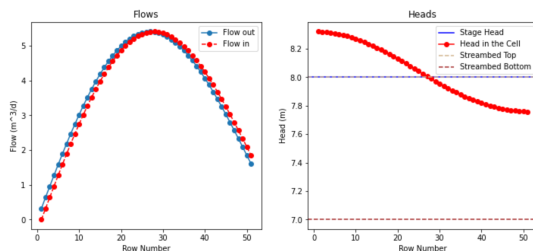


c) Choose two things to explore (e.g. impact of streambed K or inflow into the river or recharge rate). Produce a plot for each to compare to the base plots and use the plots to explain the impact of the hydrologic change.

1. How does turning the streambed conductivity to 1/100 affect streamflow?

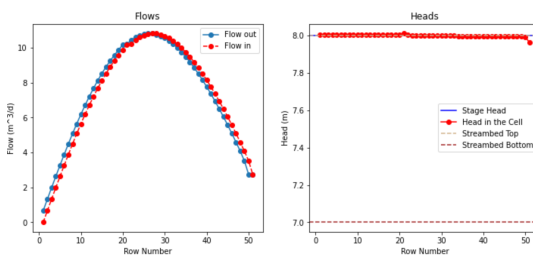


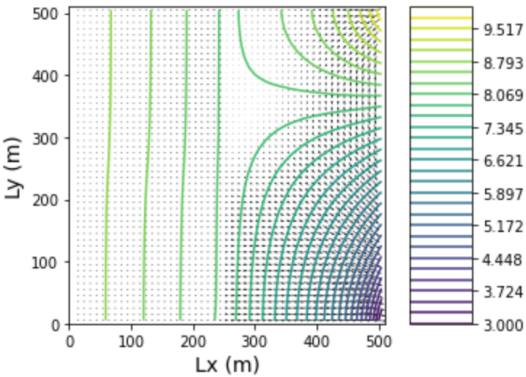
- In general, reducing the K to 1/100 across the whole domain does not greatly change the shape of the curve
- Two important differences:
 - i. The total flow in the river peaks at a smaller amount than the multiple-K condition. This is probably because it is harder for water to get into the river from the groundwater system
 - ii. The river flows to its terminus. This is probably because the low-K condition prevents water flow leaving the river and leaking into the groundwater system.



2. How does turning the streambed conductivity to 1 affect streamflow?

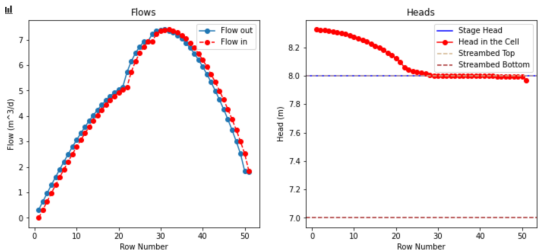
- In general, increasing K to 1 does not change the shape of the curve.
- Two important differences:
 - i. The magnitude of flow peaks at a much higher level (perhaps because it is much easier for water to flow into the river now)
 - ii. The head stays constant at 8 throughout the river.
- Another thought: no matter how we change 'K', the divide at row ~12 seems to be there.





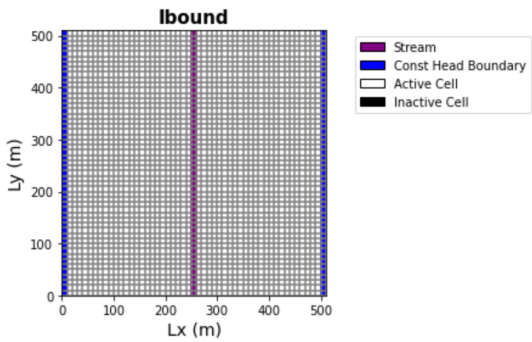
3. How does changing the the recharge zone (increasing its size) affect the flow in the river and the shape of the contour distribution?

- Given the condition of streambed $K = 1$: After testing a recharge zone up to row 36, 46, and 56, we see that it has NO affect on the contour distribution in the domain. The flow in the river is somewhat larger with distance downstream
- Given the initial streambed condition (variable): The size of the recharge zone doesn't fundamentally alter the head distribution in the domain too much, but if we increase the size of the domain enough, it does provide just enough of a boost to keep the stream flowing to its terminus.

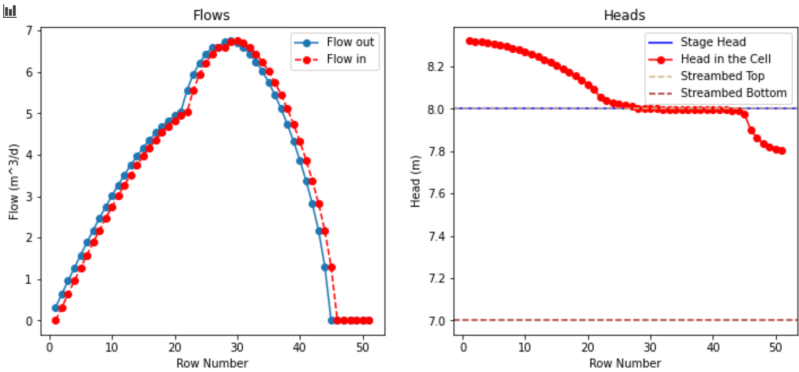


Correct Key Figures

Plan View Map



Plots of Flows and Heads



Head Distribution

