

Homework 2

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Exercise 1

① incompressible two-layer fluid $\rho_1 > \rho_2$ Giorgio Checola

$PE = \int g z \rho dV$ $h(y) = \frac{1}{2}H + \gamma y$

a) $PE = \int_0^H \int_{-L}^L \underbrace{g \rho z}_{\text{2D volume}} dy dz \rightarrow \frac{PE}{g} = \int_{-L}^L \rho_1 \int_0^{h(y)} z dz dy + \int_{-L}^L \rho_2 \int_{h(y)}^H z dz dy =$

$$= \int_{-L}^L \rho_1 \left(\frac{H^2}{4} + \gamma^2 y^2 + H\gamma y \right) dy + \int_{-L}^L \rho_2 \left(H^2 - \frac{H^2}{4} - \gamma^2 y^2 - H\gamma y \right) dy =$$

$$= \left[\rho_1 \frac{H^2}{8} y + \rho_1 \frac{\gamma^2 y^3}{6} + \rho_1 \frac{H\gamma y^2}{4} \right]_{-L}^L + \left[\rho_2 \frac{3}{8} H^2 y - \rho_2 \frac{\gamma^2 y^3}{6} - \rho_2 \frac{H\gamma y^2}{4} \right]_{-L}^L =$$

$$= 2\rho_1 \frac{H^2}{8} L + 2\rho_1 \frac{\gamma^2 L^3}{6} + \cancel{\rho_1 \frac{H\gamma L^2}{4}} + 2\rho_2 \frac{3}{8} H^2 L - 2\rho_2 \frac{\gamma^2 L^3}{6} - \cancel{\rho_2 \frac{H\gamma L^2}{4}} =$$

$$= \rho_1 \frac{H^2}{4} L + \rho_1 \frac{\gamma^2 L^3}{3} + \rho_2 \frac{3}{4} H^2 L - \rho_2 \frac{\gamma^2 L^3}{3} = \frac{H^2 L}{4} (\rho_1 + 3\rho_2) + \frac{1}{3} \gamma^2 L^3 (\rho_1 - \rho_2)$$

$$\Rightarrow PE = \rho_1 H^2 L \left(\frac{g}{4} + \frac{3}{4} g \frac{\rho_2}{\rho_1} \right) + \frac{1}{3} \gamma^2 L^3 \rho_1 \left(g - \frac{\rho_2}{\rho_1} g \right) =$$

$$= \rho_1 H^2 L \left(g - \frac{3}{4} g + \frac{3}{4} g \frac{\rho_2}{\rho_1} \right) + \frac{1}{3} \gamma^2 L^3 \rho_1 \left(\frac{\rho_1 - \rho_2}{\rho_1} g \right) =$$

$$= \rho_1 H^2 L \left(g - \frac{3}{4} g \left(g - \frac{\rho_2}{\rho_1} g \right) \right) + \frac{1}{3} \gamma^2 L^3 \rho_1 g' = H^2 L \rho_2 \left(g - \frac{3}{4} g' \right) + \frac{1}{3} \rho_1 g' \gamma^2 L^3$$

$\frac{PE_{\text{final state}}}{g} = \int_{-L}^L \rho_1 \int_0^{\frac{H}{2}} z dz dy + \int_{-L}^L \rho_2 \int_{\frac{H}{2}}^H z dz dy =$

$$= 2L \rho_1 \frac{H^2}{8} + 2L \rho_2 \left(\frac{H^2}{2} - \frac{H^2}{8} \right) =$$

$$= \rho_1 \frac{H^2}{4} L + \frac{3}{4} \rho_2 H^2 L = H^2 L \rho_2 \left(\frac{1}{4} + \frac{3}{4} \frac{\rho_1}{\rho_2} \right)$$

$$\Rightarrow PE_{\text{final state}} = H^2 L \rho_2 \left(\frac{g}{4} + \frac{3}{4} g \frac{\rho_1}{\rho_2} \right) = H^2 L \rho_2 \left(g - \frac{3}{4} g \left(\frac{\rho_1 - \rho_2}{\rho_1} \right) \right) = H^2 L \rho_2 \left(g - \frac{3}{4} g' \right)$$

$\Delta PE = PE_{\text{initial state}} - PE_{\text{final state}} = \frac{1}{3} \rho_1 g' \gamma^2 L^3$ which is independent of the depth H
because it is the amount of PE available to be converted to KE (available to drive motion). It is fundamental only the relation at the interface of the two fluid, not its height

b) $\tan \gamma \approx \gamma = \frac{dz}{dy}$ $\rho_1 = \rho_2$

approximation for small angles definition of tangent

$$\frac{\partial \rho_1}{\partial y} dy + \frac{\partial \rho_1}{\partial z} dz = \frac{\partial \rho_2}{\partial z} dz + \frac{\partial \rho_2}{\partial y} dy$$

$$\rightarrow \frac{\partial \rho_1}{\partial y} - \frac{\partial \rho_2}{\partial y} = \frac{dz}{dy} \left(\frac{\partial \rho_2}{\partial z} - \frac{\partial \rho_1}{\partial z} \right) \Rightarrow \frac{dz}{dy} = \frac{\frac{\partial \rho_1}{\partial y} - \frac{\partial \rho_2}{\partial y}}{\frac{\partial \rho_2}{\partial z} - \frac{\partial \rho_1}{\partial z}} = \frac{\frac{\partial \rho_1}{\partial y} - \frac{\partial \rho_2}{\partial y}}{g(\rho_1 - \rho_2)}$$

hydrostatic balance

$$u_g(\rho_2) - u_g(\rho_1) = -\frac{1}{f} \rho_2 \frac{\partial \rho_2}{\partial y} + \frac{1}{f} \rho_1 \frac{\partial \rho_1}{\partial y} = \frac{1}{f} \left(\frac{\partial \rho_1}{\partial y} \frac{1}{\rho_1} - \frac{\partial \rho_2}{\partial y} \frac{1}{\rho_2} \right)$$

assuming $\rho_2 \approx \rho_1 \rightarrow \approx \frac{1}{f} \frac{1}{\rho_1} \left(\frac{\partial \rho_1}{\partial y} - \frac{\partial \rho_2}{\partial y} \right) = \frac{1}{f} \frac{g(\rho_1 - \rho_2)}{\rho_1} \frac{dz}{dy} = \frac{1}{f} g' \gamma$ Margules relation

lower layer at rest $\Rightarrow u_1 = 0, \frac{\partial p_1}{\partial y} = 0$

from the tangent definition

$$\frac{\partial p}{\partial y} = \left(\frac{dz}{dy} \right)^2 g^2 (p_2 - p_1)^2$$

$$\begin{aligned} \Rightarrow KE &= \int_{h(y)}^H \int_{-L}^L \frac{1}{2} \rho_2 u_z^2 dy dz = \frac{1}{2} \rho_2 \int_{h(y)}^H \int_{-L}^L \frac{1}{f^2 \rho_2^2} \frac{\partial p}{\partial y} dy dz \\ &\quad \uparrow \text{zonal geostrophic velocity} \\ &= \frac{1}{2} \frac{\rho_2}{f^2} g^2 \int_{-L}^L \left(\frac{H}{2} - y \right) dy = \frac{1}{2} \rho_2 \frac{g^2}{f^2} \left(2L \frac{H}{2} - \frac{y^2}{2} \Big|_{-L}^L \right) \\ &= \frac{1}{2} \rho_2 \frac{g^2}{f^2} HL \end{aligned}$$

$$c) \frac{APE}{KE} = \frac{\frac{1}{3} \rho_1 g^2 L^3}{\frac{1}{2} \rho_2 \frac{g^2}{f^2} HL} = \frac{2}{3} \frac{\rho_1}{\rho_2} L^2 \frac{1}{\sqrt{\rho_1 H} \sqrt{g H}} = \frac{2}{3} \frac{\rho_1}{\rho_2} \frac{L^2}{L_R^2} \approx \frac{2}{3} \frac{L^2}{L_R^2}$$

if assuming $\rho_1 \approx \rho_2$

For the atmosphere circulation ($L = 4000 \text{ km}, L_R = 700 \text{ km}$) $\frac{APE}{KE} \approx 21.77 \frac{\rho_1}{\rho_2} \approx 21.77$

$$d) IE = c_v T \cdot m = c_v \int \rho T dV = \frac{c_v}{R} \int p dV \stackrel{\text{hydrostatic balance and flat lower boundary}}{=} \frac{c_v}{R} g \int z \rho dV$$

$$IE + PE = TPE = \frac{c_v}{R} g \int z \rho dV + g \int z \rho dV = \frac{5}{2} \frac{R}{R} g \int z \rho dV + g \int z \rho dV = \frac{7}{2} g \int z \rho dV$$

total potential energy $\rightarrow TPE \text{ is } \frac{7}{2} PE$

Exercise 2

②

a) $\frac{Du}{Dt} + \frac{1}{\rho_0} \frac{\partial p}{\partial x} - f v = 0 \rightarrow \frac{\partial}{\partial y} \frac{Du}{Dt} + \frac{1}{\rho_0} \frac{\partial}{\partial y} \frac{\partial p}{\partial x} - \frac{\partial}{\partial y} (f v) = 0$

$\frac{Dv}{Dt} + \frac{1}{\rho_0} \frac{\partial p}{\partial y} + f u = 0 \rightarrow \frac{\partial}{\partial x} \frac{Dv}{Dt} + \frac{1}{\rho_0} \frac{\partial}{\partial x} \frac{\partial p}{\partial y} + \frac{\partial}{\partial x} (f u) = 0$

$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} = 0 \Rightarrow \frac{\partial}{\partial x} \frac{Dv}{Dt} - \frac{\partial}{\partial y} \frac{Du}{Dt} + u \frac{\partial f}{\partial x} + f \frac{\partial u}{\partial x} + v \frac{\partial f}{\partial y} + f \frac{\partial v}{\partial y} = 0$

$\rightarrow \frac{\partial}{\partial x} \left(\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + w \frac{\partial v}{\partial z} \right) - \frac{\partial}{\partial y} \left(\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} \right) + u \frac{\partial f}{\partial x} + v \frac{\partial f}{\partial y} + f \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) = 0$

$\rightarrow \left(\frac{\partial}{\partial t} \frac{\partial v}{\partial x} + u \frac{\partial}{\partial x} \frac{\partial v}{\partial x} + v \frac{\partial}{\partial y} \frac{\partial v}{\partial x} + w \frac{\partial}{\partial z} \frac{\partial v}{\partial x} \right) - \left(\frac{\partial}{\partial t} \frac{\partial u}{\partial y} + u \frac{\partial}{\partial x} \frac{\partial u}{\partial y} + v \frac{\partial}{\partial y} \frac{\partial u}{\partial y} + w \frac{\partial}{\partial z} \frac{\partial u}{\partial y} \right) + \left(\frac{\partial f}{\partial t} + u \frac{\partial f}{\partial x} + v \frac{\partial f}{\partial y} \right) = 0$

$\Rightarrow \frac{D}{Dt} \left(\frac{\partial v}{\partial x} - \frac{\partial u}{\partial y} \right) + f = 0$

b) $\hat{z} \cdot \nabla \times \vec{u} = \hat{z} \cdot \left(\frac{\partial}{\partial x} \hat{i}, \frac{\partial}{\partial y} \hat{j}, \frac{\partial}{\partial z} \hat{k} \right) \times (u \hat{i}, v \hat{j}, w \hat{k}) =$

$= \hat{z} \cdot \left(\frac{\partial v}{\partial x} \hat{k} - \frac{\partial w}{\partial x} \hat{j} - \frac{\partial u}{\partial y} \hat{k} + \frac{\partial w}{\partial y} \hat{i} + \frac{\partial u}{\partial z} \hat{j} - \frac{\partial v}{\partial z} \hat{i} \right) = \hat{z} \cdot \left(\frac{\partial v}{\partial x} - \frac{\partial u}{\partial y}, \frac{\partial u}{\partial z} - \frac{\partial v}{\partial z}, \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y} \right)$

$= \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y}$ vertical component of the vorticity $\nabla \times \vec{u}$

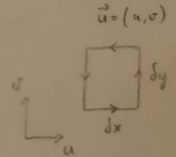
$\frac{\partial v}{\partial x} - \frac{\partial u}{\partial y} + f = \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y} + f$

absolute vorticity relative vorticity planetary vorticity

c) circulation = $u \delta x + \left(v + \frac{\partial v}{\partial x} \delta x \right) \delta y - \left(u + \frac{\partial u}{\partial y} \delta y \right) \delta x - v \delta y$

$= \left(\frac{\partial v}{\partial x} - \frac{\partial u}{\partial y} \right) \delta x \delta y$

area enclosed = $\delta x \delta y$



$\Rightarrow \frac{\text{circulation}}{\text{area enclosed}} = \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y}$

Fluids in solid body rotation have no part in motion relative to any other part of the fluid. In this case we can consider a circle of radius R ,

angular velocity Ω

$\frac{\text{circulation}}{\text{area enclosed}} = \frac{\oint \vec{u} \cdot d\vec{l}}{\pi R^2} = \frac{\int_0^{2\pi} \Omega R d(\lambda R)}{\pi R^2} = \frac{2\pi R^2 \Omega}{\pi R^2} = 2\Omega$

d) $\zeta = \left(\frac{\partial v}{\partial x} - \frac{\partial u}{\partial y} \right) = \frac{\partial}{\partial x} \left(\frac{1}{\rho f} \frac{\partial p}{\partial x} \right) - \frac{\partial}{\partial y} \left(-\frac{1}{\rho f} \frac{\partial p}{\partial y} \right) = \frac{1}{\rho f} \left(\frac{\partial^2 p}{\partial x^2} + \frac{\partial^2 p}{\partial y^2} \right)$

Exercise 3

③ $\langle T_e \rangle = 265 \text{ K}$
 $\langle T_p \rangle = 235 \text{ K}$

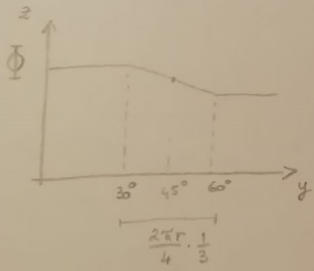
$$\Phi_e - \Phi = R \langle T_e \rangle \ln \frac{1000}{200} = 122406 \frac{\text{m}^2}{\text{s}^2}$$

$$\Phi_p - \Phi = R \langle T_p \rangle \ln \frac{1000}{200} = 108548.54 \frac{\text{m}^2}{\text{s}^2}$$

$$\Delta \Phi = \frac{13857.46}{g} = 1414 \text{ m} = 1.4 \text{ Km}$$

$$u_g = -\frac{1}{f} \frac{\partial \Phi}{\partial y} = -\frac{1}{10^{-4}} \frac{(-13857.46)}{\frac{2\pi r}{4} \cdot \frac{1}{3}} = 41.37 \text{ m/s}$$

mean zonal geostrophic wind on the 200 hPa surface at 45° lat. in the winter hemisphere



Exercise 4

④ $p_1 = 750 \text{ hPa}$
 $p_2 = 500 \text{ hPa}$

$$\frac{\partial T}{\partial x} = -\frac{3 \text{ K}}{100 \text{ Km}} = -0.03 \text{ K/Km}$$

South east $\vec{V} = 20 \text{ m/s}$ at 45°

$$\vec{V} = u_g \hat{i} + v_g \hat{j} = -20 \cos 45^\circ \hat{i} + 20 \sin 45^\circ \hat{j}$$

$$= -14.14 \hat{i} + 14.14 \hat{j}$$

$$u_g(p_2) = u_g(p_1) + \frac{R}{f} \frac{\partial \langle T \rangle}{\partial x} \ln \frac{p_1}{p_2} = 14.14 - 34.9 = -20.76 \text{ m/s}$$

$$v_g(p_2) = v_g(p_1) - \frac{R}{f} \frac{\partial \langle T \rangle}{\partial y} \ln \frac{p_1}{p_2} = -14.14 \text{ m/s}$$

$$|\vec{V}(p_2)| = \sqrt{(-14.14)^2 + (-20.76)^2} = 25.12 \text{ m/s}$$

$$\theta = 180^\circ + \arctan\left(\frac{-20.76}{-14.14}\right) \approx 236^\circ \rightarrow \text{the geostrophic wind at } 500 \text{ hPa is from the northeast}$$
