# HWRS 505: Vadose Zone Hydrology

Lecture 4

9/5/2024

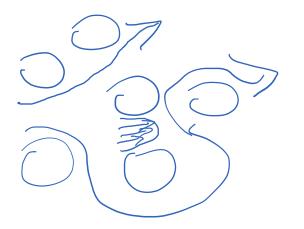
#### Today:

- 1. Wrap up the review of solute transport under saturated flow
- 2. Air-water system in capillary tubes

#### Review of Lecture 3

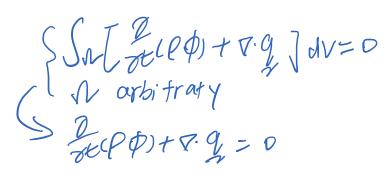
- Derivation of 3D transient groundwater flow.
- Solute transport under saturated flow.
  - Advection ( $v = q/\phi$ )
  - Molecular diffusion
  - Mechanical dispersion

"The dispersion coefficient is a lumped fitting parameter that adequately describes relatively large-scale observations."



$$q_c = \phi (vC - D\nabla C)$$
  $= qC - \phi D\nabla C$ 

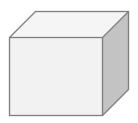




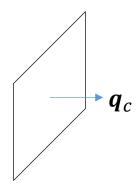
$$\lim_{\Delta x \to 0} \int_{X^* - \Delta t_2}^{X * t_2} f(x) dx = f(x^*) \cdot ox = 0$$

$$\int_{X^* - \Delta t_2}^{X * t_2} f(x) dx = \int_{X^* - \Delta t_2}^{X * t_2} f(x) dx = 0$$

$$\Rightarrow f(x^*)=0$$
 contradicts  $f(x^*)\neq 0$   
 $\Rightarrow f(x)=0$  everywhere



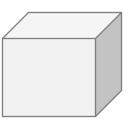
Saturated porous medium



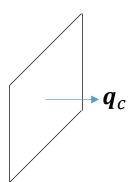
$$\mathbf{q}_{c} = \phi (\mathbf{v}C - \mathbf{D}\nabla C)$$

$$= \mathbf{q}C - \phi \mathbf{D}\nabla C$$

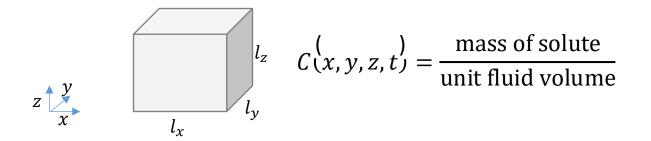
$$\mathbf{D} = \alpha_{T} |_{\mathcal{V}_{1-}} + (c_{L} - \alpha_{T}) \frac{\mathbf{v}\mathbf{v}}{|\mathbf{v}|} + wD_{0}\mathbf{I}$$



Fluid only (e.g., free water)



$$q_c = vC - D_0 \nabla C$$



<u>Mass conservation</u>: Change of mass storage = mass in – mass out.

Rate of mass change:  $\frac{d}{dt} \int_{0}^{t_{x}} \int_{0}^{t_{y}} \int_{0}^{t_{z}} \phi c(x_{y},z_{i}t) dx dy dz = \frac{d}{dt} \int_{0}^{t_{y}} \phi c(x_{y},t) dy$ 

Net fluxes:  $\int_{0}^{C_{x}} \int_{0}^{C_{x}} \int_$ 

$$\Rightarrow 2 \int_{\partial t} \phi c dv = - \int_{\mathcal{N}} \nabla \cdot q dv$$

$$\Rightarrow \sqrt{v + v + v}$$

$$\Rightarrow \frac{1}{2} \phi C + \frac{1}{2} \cdot \frac{1}{2} c = 0$$

$$\begin{cases} \frac{1}{2} (\phi c) + \frac{1}{2} (\phi c) - \frac{1}{2} (\phi c) = 0 \\ \frac{1}{2} (\phi c) + \frac{1}{2} (\phi c) - \frac{1}{2} (\phi c) = 0 \end{cases}$$

$$\begin{cases} \frac{1}{2} (\phi c) + \frac{1}{2} (\phi c) - \frac{1}{2} (\phi c) - \frac{1}{2} (\phi c) = 0 \end{cases}$$

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$$\begin{cases} \frac{1}{2} (\phi c) + \frac{1}{2} (\phi c) + \frac{1}{2} (\phi c) - \frac{1}{2} (\phi c) = 0 \end{cases}$$

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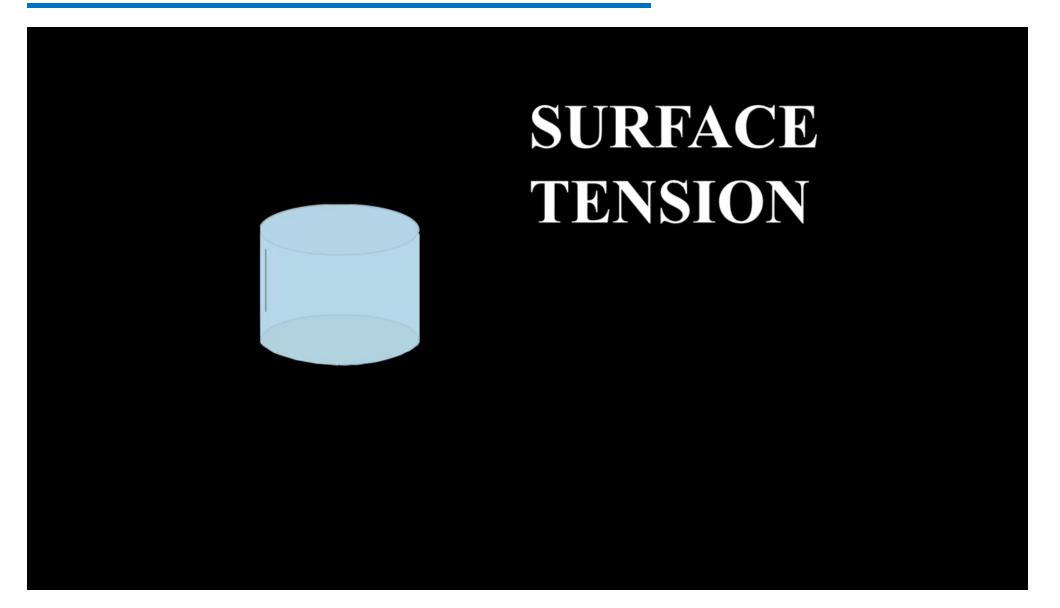
$$\begin{cases} \frac{1}{2} (\phi c) + \frac{1}{2} (\phi c$$

Colute monspore in powers media





- 1. Why does the water try to hold together?
- 2. Why does the water not wet the surface?



Link to the video: https://youtu.be/zMzqiAuOSz0

- > Two and three phase systems: water, oil, air
- Interfacial tension (<u>cohesive</u> forces between fluid molecules)

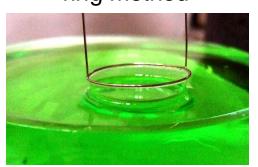


How to measure interfacial tension?

Drop weight method

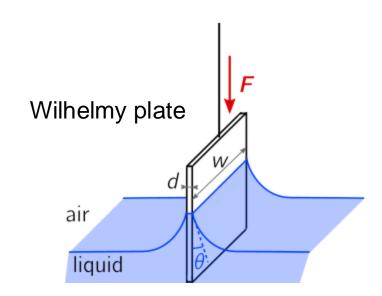


#### ring method

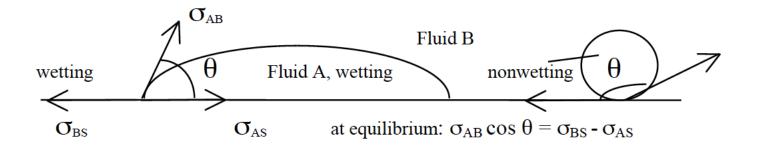


#### Typical values of surface tension:

air-water 0.072 N/m oil-water 0.20 N/m oil-water w/ soap 0.0001 N/m



Wettability (adhesive forces between the fluid and solid surface)



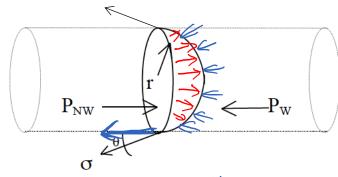
 $\theta < 90^{\rm o}$ : fluid A is wetting with respect to fluid B on the solid S

 $\theta > 90^{\circ}$ : fluid A is nonwetting with respect to fluid B on the solid S

Wettability is a function of the fluid properties, soil properties, and history of contact. For most soils, the relative wettabilities are: water > oil > air

Recommended video for the concepts of *viscosity, cohesive and adhesive forces, surface tension, and capillary action* <a href="https://www.youtube.com/watch?v=P\_jQ1B9UwpU">https://www.youtube.com/watch?v=P\_jQ1B9UwpU</a>

Capillary pressure (difference between the nonwetting and wetting phase pressures)



Force balance at equilibrium:

$$\pi P P m - \pi r^2 P w = 2\pi r \int \cos \theta$$

$$P_{NW}-P_{W}=20000$$

$$P_{C}=20000$$

Young-Laplace Equation

1. More general equation  
for any num interface  
$$Pc = (1 + \frac{1}{7} + \frac{1}{7})$$

$$Pc = (I \cdot \overline{I}, + \overline{I}_2)$$
For the capillary tube:  $Y_1 = Y_2 = \frac{Y_1}{\cos \theta}$ 

2. For a perfectly wetting fluide

$$\theta=0$$
,  $\ell c=\frac{30}{r}$