# HWRS 505: Vadose Zone Hydrology

Lecture 10

9/25/2024

Today: Richards' Equation and steady-state unsaturated flow

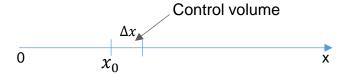
Reading: Chapter 11 (Pinder & Celia, 2006) and Ferre Lecture Notes

#### Comments on Homework #1

#### Problems #1 and #2

Vertical 
$$q_z = -K \frac{\mathrm{d}H}{\mathrm{d}z}$$
  $H = \psi + z$   
Horizontal  $q_x = -K \frac{\mathrm{d}H}{\mathrm{d}x}$   $H = \psi \ (z = 0)$ 

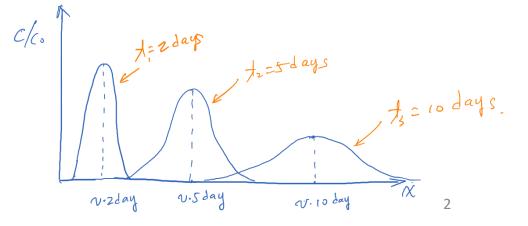
#### Problems #3



Net flux

$$q_x \Big|_{x=x_0} - q_x \Big|_{x=x_0+\Delta x} = -\int_{x=x_0}^{x=x_0+\Delta x} \frac{\mathrm{d}q}{\mathrm{d}x} \,\mathrm{d}x$$

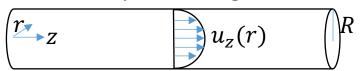
Draw schematics of the solutions of solute concentrations (c(x,t)) at three times, t=2, 5, 10 days.



# Solute transport under saturated flow

Taylor-Aris Dispersion: Dispersion in a capillary tube [G.I. Taylor (1953) and R. Aris (1956)]

Solute transport in "Hagen-Poiseuille" flow



Governing equation for solute transport in the tube:

$$\frac{\partial C}{\partial t} + 2u \left[ 1 - \left( \frac{r}{R} \right)^2 \right] \frac{\partial C}{\partial x} - D_0 \left( \frac{\partial^2 C}{\partial r^2} + \frac{1}{r} \frac{\partial C}{\partial r} + \frac{\partial^2 C}{\partial x^2} \right) = 0$$

At sufficiently long times (derivation via the perturbation method)

$$\frac{\partial \bar{C}}{\partial t} + \bar{u} \frac{\partial \bar{C}}{\partial x} - D_L \frac{\partial^2 \bar{C}}{\partial x^2} = 0, \qquad D_L = D_0 + \frac{R^2 \bar{u}^2}{48D_0}$$

<u>Insights</u>: The average solute concentration spreads out by a dispersion process (radial diffusion and axial advection) and follows Fickian diffusion. The effective diffusivity is not the molecular diffusivity  $D_0$ . Rather, it is  $D_0$  + a quadratic function of the mean velocity.

 $\bar{u}$ : mean velocity in the tube.

 $C_m$ : mean concentration in the tube.

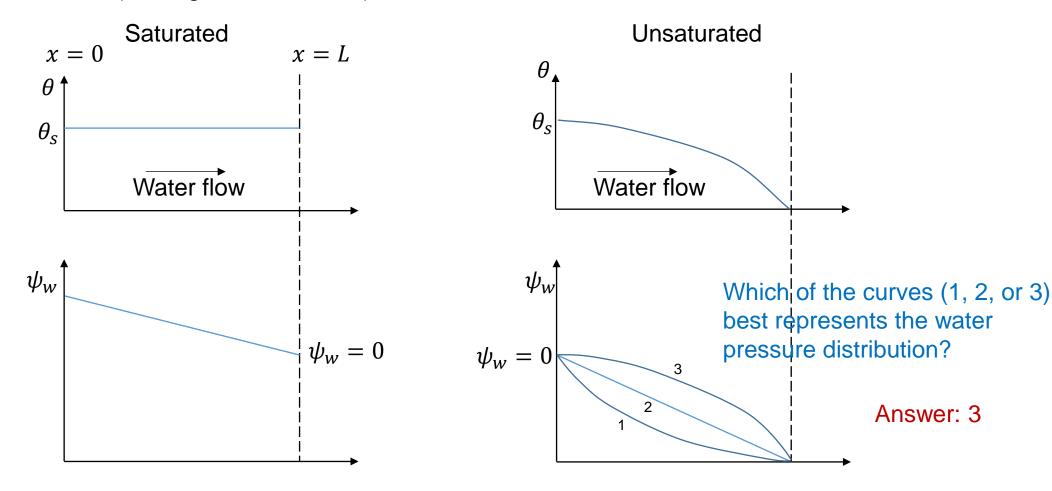
$$C_m = \frac{\int_0^{2\pi} \int_0^R C(r, x) r \, dr \, d\theta}{\int_0^{2\pi} \int_0^R r \, dr \, d\theta} = \frac{2}{R^2} \int_0^R C r \, dr$$

#### Review of Lecture 9

- Three forms of Richards' equation
  - Mixed form
  - Pressure head-based form
    - ✓ Specific moisture capacity
  - Water content-based form
    - ✓ Soil moisture diffusivity (What does the equation have to do with "diffusion"?)
    - ✓ Cannot be used if the domain involves saturated water flow
  - How to include soil and fluid compressibility?
- Richards' assumptions
  - Air pressure remains almost zero everywhere, but air does move.
  - Does "air movement" make the Richards' equation invalid? No, as long as air pressure remains almost zero everywhere.

#### Steady-State Unsaturated Flow

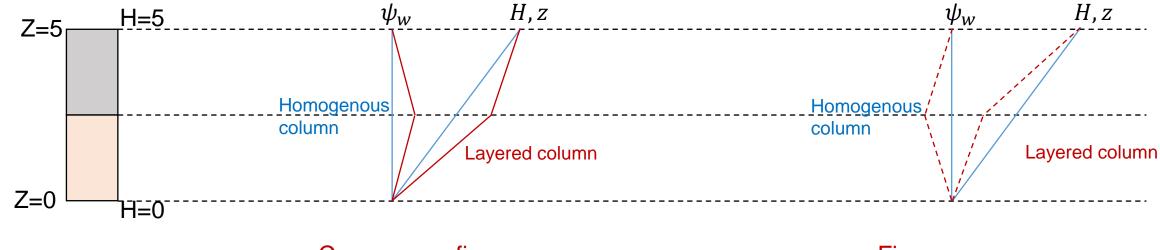
Horizontal flow (Homogeneous column)



Unsaturated flow involves nonlinearities that make their behaviors differ from that of the saturated flow

#### Steady-State Unsaturated Flow

Vertical flow (in layered columns)

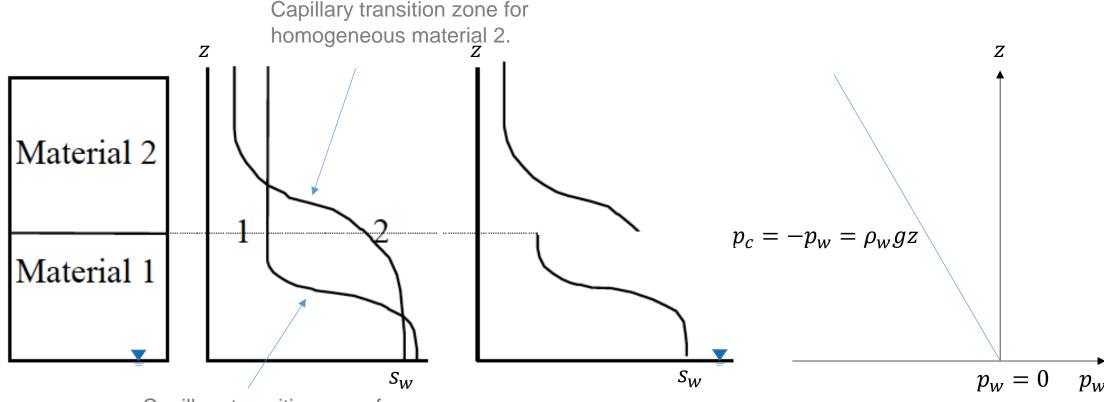


Coarse over fine Fine over coarse

- Hydraulic head and pressure head are both continuous in space.
- Is water saturation continuous?

### Steady-State Unsaturated Flow

#### Hydrostatic unsaturated layered systems



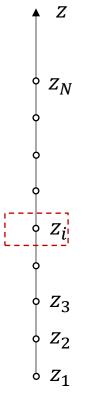
Capillary transition zone for homogeneous material 1.

 $\triangleright$   $S_w$  is discontinuous at the material interface, but the  $p_c$  and  $p_w$  or  $\psi_w$  are continuous.

### Steady-State Unsaturated Flow: Numerical Soln.

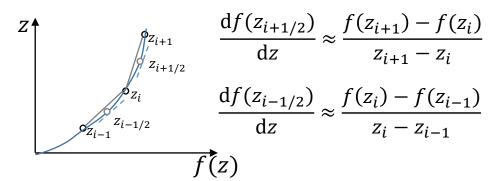
$$\frac{\partial \theta_{w}}{\partial t} - \frac{\partial}{\partial z} \left( K \frac{\partial \psi_{w}}{\partial z} \right) - \frac{\partial K}{\partial z} = 0 \implies \frac{\frac{1}{dz} \left( K \frac{d\psi_{w}}{dz} \right) + \frac{dK}{dz}}{\frac{d}{dz} \left( K \frac{d\psi_{w}}{dz} \right)} + \frac{\frac{1}{dz}}{\frac{dz}{dz}} = 0$$
 (1)

- Equation (1) is a second-order ordinary differential equation in 1D.
- It is nonlinear because  $K = K(\psi_w)$  is a nonlinear function.
- To solve it, we need two boundary conditions and we need to do it in an iterative procedure.



How to solve this 1D nonlinear ordinary differential equation?

- Key idea: Divide the domain into many boxes and convert the differential equation to a system of nonlinear algebraic equations.
- Technique: Use finite difference to approximate derivatives



# Steady-State Unsaturated Flow: Numerical Soln.

$$\frac{\partial \theta_{w}}{\partial t} - \frac{\partial}{\partial z} \left( K \frac{\partial \psi_{w}}{\partial z} \right) - \frac{\partial K}{\partial z} = 0 \implies \frac{\frac{1}{dz} \left( K \frac{d\psi_{w}}{dz} \right) + \frac{dK}{dz}}{\frac{d\zeta}{dz}} = 0 \quad (1)$$

$$\begin{vmatrix} z \\ \frac{d}{dz} \left( K \frac{d\psi_{w}}{dz} \right) \Big|_{z_{i}} \approx \frac{\left( K \frac{d\psi_{w}}{dz} \right)_{i+1/2} - \left( K \frac{d\psi_{w}}{dz} \right)_{i-1/2}}{\Delta z} \approx \frac{K_{i+1/2} \frac{\psi_{w,i+1} - \psi_{w,i}}{\Delta z} - K_{i-1/2} \frac{\psi_{w,i-1} - \psi_{w,i-1}}{\Delta z}}{\Delta z} \\ = \frac{K_{i+1/2} (\psi_{w,i+1} - \psi_{w,i}) - K_{i-1/2} (\psi_{w,i} - \psi_{w,i-1})}{\Delta z^{2}} \\ \frac{dK}{dz} \Big|_{z_{i}} \approx \frac{K_{i+1/2} - K_{i-1/2}}{\Delta z} \\ \Rightarrow \frac{K_{i+1/2} (\psi_{w,i+1} - \psi_{w,i}) - \frac{K_{i-1/2}}{\Delta z}}{\Delta z^{2}} (\psi_{w,i} - \psi_{w,i-1}) + K_{i+1/2} \frac{1}{\Delta z} - K_{i-1/2} \frac{1}{\Delta z} = 0 \\ \frac{z_{3}}{\Delta z^{2}} \Rightarrow \frac{K_{i+1/2}}{\Delta z^{2}} \psi_{w,i+1} - \left( \frac{K_{i+1/2}}{\Delta z^{2}} + \frac{K_{i-1/2}}{\Delta z^{2}} \right) \psi_{w,i} + \frac{K_{i-1/2}}{\Delta z^{2}} \psi_{w,i-1} + K_{i+1/2} \frac{1}{\Delta z} - K_{i-1/2} \frac{1}{\Delta z} = 0$$

This is an algebraic equation with 3 unknowns. We can write such an algebraic equation for each node or box and we can get N algebraic equations for the N unknowns  $(\psi_{w,1}, \psi_{w,2}, ..., \psi_{w,N})$ 

### Steady-State Unsaturated Flow: Numerical Soln.

In matrix form:

 $\Rightarrow F(\psi_w) \cdot \psi_w = R$  How to solve this nonlinear system of algebraic equations? => Employ an iterative methods