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Sensitivity and linearity analysis of ozone in East Asia: The effects of domestic emission and intercontinental transport

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In this study, ozone (O₃) sensitivity and linearity over East Asia (EA) and seven urban areas are examined with an integrated air quality modeling system under two categories of scenarios: (1) The effects of domestic emission are estimated under local emission reduction scenarios, as anthropogenic NO_x and volatile organic compounds (VOC) emissions are reduced by 20%, 50%, and 100%, respectively and independently; and (2) the influence of intercontinental transport is evaluated under Task Force on Hemispheric Transport of Air Pollution (TF HTAP) emission reduction scenarios, as anthropogenic NO_x emission is reduced by 20% in Europe (EU), North America (NA), and South Asia (SA), respectively. Simulations are conducted for January and July 2001 to examine seasonal variation. Through the domestic O₃ sensitivity investigation, we find O₃ sensitivity varies dynamically depending on both time and location: North EA is VOC limited in January and NO_x limited in July, except for the urban areas Beijing, Shanghai, Tokyo, and Seoul, which are VOC limited in both months; south EA is NO_x limited in both January and July, except for the urban areas Taipei, which is VOC-limited in both months, and Pearl River Delta, which is VOC limited in January. Surface O₃ change is found to be affected more by NO_x than by VOC over EA in both January and July. We also find different O₃ linearity characteristics among urban areas in EA: O₃ at Beijing, Tokyo, and Seoul shows a strong negative linear response to NO_x emission in January; O₃ at Shanghai, Pearl River Delta, and Taipei shows a strong positive response to VOC emission in both January and July. Through the long-range transport investigation, monthly O₃ changes over EA resulting from different source regions indicate the largest source contribution comes from NA (0.23 ppb), followed by SA (0.11 ppb) and EU (0.10 ppb). All of the three regions show higher impacts in January than in July.

Implications: This study examine O₃ sensitivities and linear response of NO_x and VOC emission over EA and seven urban areas based on regional air quality modeling system MM5/CMAQ. We also quantify the intercontinental transport effect from EU, SA, and NA over EA. The result provide a theoretical basis for emission control strategy design in EA, and also reveal the O₃ special nonlinearity features for further related studies that are applicable to other continents. The HTAP multimodel experiments need to examine the potential impacts on ground-level O₃ of changes in meteorology and transport patterns expected as a result of the regional scale.

Introduction

One of the most significant criteria pollutants affecting human health and ecosystems, tropospheric ozone (O₃) is formed through a sequence of chemical reactions involving volatile organic compounds (VOC), methane (CH₄), and carbon monoxide (CO) in the presence of nitrogen oxides (NO_x) and sunlight (National Research Council, 1991). Excessive regional anthropogenic emissions of VOC and NO_x may amplify or restrain the photochemical production of O₃ formation, depending on the VOC/NO_x ratio (Liu et al., 2004; So et al., 2003; Wang et al., 2005). Therefore, it is essential to identify VOC-limited and NO_x-limited regimes to provide scientific basis to support development of effective O₃ control strategies for regional air quality improvement. While the O₃–NO_x–VOC relationship has been extensively studied in the

United States (Jacob et al., 1995; Sillman, 1995; Sillman, 1999; Sillman and He, 2002), much less attention has been paid to East Asia (EA). In a 2000 study conducted by Luo et al. (2000) with local anthropogenic emission of VOC and NO_x reduced by 35%, it was found that in wintertime rural areas in south China are NO_x limited, while rural areas in north China are more likely to be VOC limited. Model simulations from Wang et al. (2005) indicate that urban areas in the Pearl River Delta (PRD) region are VOC limited in March, while nonurban areas in PRD are NO_x limited. These reported O₃ sensitivity features, however, are limited to specific seasons or local areas, which may not be applicable, for instance, for summertime's higher temperatures and more intensive photochemical activities, or appropriate for describing the O₃ sensitivity over whole EA domain. In addition, none of these studies provide a comprehensive investigation of the seasonal variation of ozone

over EA. To achieve better air quality, it is important to investigate the characteristics of the O₃–VOC–NO_x photochemical relationships (Hakami et al., 2004) and its seasonality on a regional scale (Cohan et al., 2006) over EA. Additionally, although the dependence of O₃ production on domestic NO_x and VOC over the United States is strongly nonlinear based on a number of previous studies (Sillman, 1999; Jacob et al., 1995), very few studies have been conducted over EA. Xu et al. (2008) reported the long-term linear trend of maximum O₃ concentration from a regional background observational station in China relevant to NO_x emission, but this study was limited to the rural site and did not consider VOC emission. Xing et al. (2011) recently conducted simulations over eastern China to characterize the NO_x and VOC emission impacts on O₃, reported the nonlinear response of O₃ regarding to local and regional emissions, but only focused on O₃ isopleths for three cities. Yamaji et al. (2012) conducted a modeling study of O₃ response in Japan by controlling anthropogenic emission in China from –100 to +100%, and found a strong linear response in spring and summer over the range between –30 and +100% emission changes, but the finding is based on subdomain averaged values and may not be applicable for urban areas.

While local VOC and NO_x emissions play an important role in O₃ formation, intercontinental transport may also affect the background O₃ significantly in downwind regions (Bernsten et al., 1999; Fiore et al., 2002; Heald et al., 2006; Wu et al., 2009). In order to evaluate the long-range transport effects of air pollutants and source–receptor (SR) relationships for O₃ in the Northern Hemisphere, the Task Force on Hemispheric Transport of Air Pollution (TF HTAP) was established under the United Nations Economic Commission for Europe (UNECE) Convention on Long-Range Transboundary Air Pollution (LRTAP). Applying 21 global chemical transport models (CTMs), Fiore et al. (2009) reported the monthly ensemble mean O₃ changes from 21 global model outputs for all SR scenarios. Although these global CTMs provide an essential broad outline to estimate the intercontinental SR relationships, it is reported that there are significant uncertainties regarding the representation of both chemical mechanisms (Emmerson and Evans 2009) and physical parameterizations due to the global-scale spatial resolution of 2° × 2.5° or coarser. Moreover, since rapid industrialization over EA has introduced large amounts of anthropogenic emissions of air pollutants and precursors (Streets et al., 2003; Streets et al., 2007) within the recent decade, most of the studies focus on EA as a major source region for air pollutants (Jaffe et al., 1999; Jacob et al., 1999). Much less attention has been paid concerning the impacts from other continents on EA, which could also suffer as a receptor region from long-range transport of air pollutants from Europe (EU), North America (NA), and South America (SA) (Fiore et al., 2009; Holloway et al., 2008). Fu et al. (2011) reported SA could contribute 8–18 ppbv on O₃ over southeastern EA in the biomass burning period; Holloway et al. (2008) found that EU and NA emissions contribute up to 9 ppbv of O₃ over western Mongolia in summer, and 5–7 ppbv over Japan in the spring and late fall, based on results from global Model of Ozone and Related Tracers (MOZART). These studies imply the significance of intercontinental transport effects over EA, but still lack detailed information

about SR relationships regarding EA in different seasons. Likewise, they may retain uncertainties from the global model.

We have undertaken this study to conduct a regional modeling assessment and long range transport analysis by applying the fifth-generation Mesoscale Model version/Community Multiscale Air Quality (MM5/CMAQ) modeling system over EA for two reasons: (1) to link O₃ sensitivities and linearity with local VOC and NO_x emissions in EA for both winter and summer, and (2) to quantify the impacts from EU, NA, and SA on surface O₃ concentration in EA at a finer-scale resolution under TF HTAP scenarios, and examine its relative importance compared with domestic emissions. We also select seven urban areas of most air pollution concern within the EA domain because of their condensed population and developed economic and industries. This is the first attempt to model O₃ in different seasons under both domestic and intercontinental transport emission removal scenarios to characterize the O₃ sensitivity features and linearity of O₃ responses to local NO_x and VOC emission change, and also to quantify the long-range transport effects spread over EA, as well as the impacts in urban areas.

Methods

Model configuration

CMAQ is a regional air quality model developed by the U.S. Environmental Protection Agency (EPA) (Byun and Schere, 2006; Byun and Ching, 1999). It has been applied in numerous governmental and research studies for making regulatory decisions and conducting scientific atmospheric research in the United States. In our study, CMAQv4.6 is driven by meteorological fields from MM5 version 3.6. Model configurations are listed in Table 1. The National Centers for Environmental Prediction (NCEP) Final Analyses Data (ds083.2) with a spatial resolution of 1.0 × 1.0° and a 6-hourly temporal resolution was used to drive the MM5 model, along with 6-hourly observational data including NCEP Automated Data Processing (ADP), surface observations data set (ds464.0), and upper air observations (ds353.4). The one-way nested approach with four-dimensional data assimilation (FDDA) in MM5 was performed from a mother domain with a 108 km × 108 km horizontal resolution (ranging from 174 to 46° E, 3 to 65° N) over Asia nest-down to 36 km × 36 km (72.5–147° E, 15–49.5° N). Meteorology/Chemistry Interface Processor (MCIP) 2.3 (Byun and Schere 2006) was used to process MM5 output, and the outcome was used as the input for CMAQ.

The emission inventory in this study for CMAQ was developed from the Transport and Chemical Evolution over the Pacific (TRACE-P) project, which has been successfully applied and validated in other Asia-related studies (Streets et al., 2003, 2007; Fu et al., 2009). Eleven major chemical species were involved in the TRACE-P emission inventory, including gas-phase species such as methane (CH₄), carbon dioxide (CO₂), sulfur dioxide (SO₂), NO_x, carbon monoxide (CO), ammonia (NH₃), non-methane volatile organic compounds (NMVOC), particulate matter with aerodynamic diameter less than or equal to 2.5 μm

Table 1. Model configurations for MM5 and CMAQ

MM5 configuration	
Vertical resolution	34 σ_p levels
Horizontal resolution	36 km
Projection	Lamber conformal
FDPA	Analysis nudging
Cumulus	Kain–Fritsch (Kain and Fritsch, 1993)
Microphysics	Mixed phase
Radiation	RRTM
PBL	Pleim–Xiu
LSM	Pleim–Xiu LSM
LULC	USGS 25-category
CMAQ configuration	
Horizontal resolution	36 km
Vertical resolution	14 sigma-pressure levels (with the top pressure of 100 mb)
Projection	Lambert conformal conic
Depth of first later	36 m
Advection	Piecewise parabolic scheme
Vertical diffusion	K-theory
Gas-phase chemistry	CB-IV with EBI solver
Dry deposition	Wesely (Wesely, 1989)
Wet deposition	Henry's law
Aqueous chemistry	Walcek (Walcek and Aleksic, 1998)
Advection	Global mass-conserving scheme
PBL scheme	Blackadar (Zhang and Anthes, 1982)
Explicit moisture scheme	Mixed phase (Reisner et al., 1998)
Atmospheric radiation scheme	Cloud (Dudhia, 1993)
Cumulus scheme	Grell
Multilayer soil temperature model	Five-layer soil model

(PM_{2.5}), PM₁₀, black carbon aerosol (BC), and organic carbon aerosol (OC) (Streets et al., 2003; Woo et al., 2002). NMVOC emissions were further speciated into 19 subcategories based on chemical reactivity and functional groups. The main anthropogenic emission sources were categorized according to five types of fossil fuels, which included power plants, industry, residential, transportation, and agricultural emissions. Biomass burning was evaluated independently in three major categories: the burning of forest, savanna/grassland, and crop residues. Biogenic emissions, such as isoprene and terpene emissions, were prepared by a geographic information system (GIS) based on Global Emissions Inventory Activity (GEIA; www.geiacenter.org) and land-use data. Due to the compatibility between TRACE-P chemical species and the Carbon Bond CB05 chemical species in CMAQ, mapping tables were developed to convert TRACE-P emissions into CB05 emissions, described in Fu et al. (2009).

Simulation scenarios

Initial and boundary (IC/BC) conditions of CMAQ were derived from GEOS-Chem, one of the 21 CTMs used in TF

HTAP and also used in U.S. EPA regulatory modeling. GEOS-Chem incorporated various emission inventories for different sectors, such as using the Emission Database for Global Atmospheric Research (EDGAR) for anthropogenic emission, Model of Emissions of Gases and Aerosols from Nature (MEGAN) for biogenic emission, and Global Fire Emissions Database (GFED) for biomass burning emissions. Simulation results from GEOS-Chem were downscaled to transfer the inter-continental transport information of SR relationships from global scale to the regional scale in CMAQ. The downscaling methodology of IC/BC from GEOS-Chem to CMAQ has been described in Lam and Fu (2009, 2010).

GEOS-Chem output with 2001 original emissions was downscaled to produce IC/BC for the base scenario, which is denoted as SR1 hereafter. For local NO_x and VOC reduction scenarios, SR1 IC/BC is the unchanged base case. O₃ sensitivity analyses were conducted under 20%, 50%, and 100% local anthropogenic emission reduction of NO_x and VOC, respectively. For long-range transport scenarios, IC/BC was downscaled from GEOS-Chem outputs from three HTAP emission control scenarios: (1) 20% anthropogenic NO_x emission reduction from Europe (SR3EU), (2) 20% reduction from North America (SR3NA), and (3) 20% reduction from South Asia (SR3SA).

January and July 2001 were selected as simulation periods for all of the sensitivity scenarios to characterize the O₃ seasonal variation over EA, and to examine the impacts from different seasonal monsoons, that is, the cold and dry winter monsoon and the warm and wet summer monsoon (Lin et al., 2008). Unlike surface ozone in the United States where summer ozone is of the major concern, winter ozone over EA is also of interest because it may be close to summer peak values. It has been reported by Li et al. (2011) that wintertime ozone (0.052 mg/m³) in Shanghai is close to that in summer (0.054 mg/m³), and Zhang et al. (2011) reported surface ozone at 10 observation sites in Pearl River Delta all peak in October. So it is necessary to examine the seasonality of ozone in EA.

Since it has been reported that European emissions exerted the most significant impacts on northern Asia (Holloway et al., 2008), three megacities—Beijing (BJ), Tokyo, and Seoul—in the north part of EA are selected, along with four urban areas in the south part of EA: Chengdu (CD), Pearl River Delta (PRD), Taiwan, and Shanghai (SH). These are the most developed urban areas in EA since they sustain concentrated populations and substantive anthropogenic emissions. The CMAQ model simulation is performed from December 23, 2000, to January 31, 2001, in winter, and from June 22, 2001, to July 31, 2001, in summer. The first few days of the simulation periods (December 23 to 31, 2000, and June 22 to 30, 2001) serve as spinning off days to protect against inappropriate initializations. Simulation results from January 1–31 and July 1–31, 2001, are used for analysis. The full simulation domain is shown in Figure 1, including most of the East Asian regions: China, Japan, North Korea, South Korea, Taiwan, and parts of Laos, Myanmar, and Vietnam.

MM5/CMAQ model evaluation

It is essential to evaluate models with observations to demonstrate credibility of model results. GEOS-Chem model

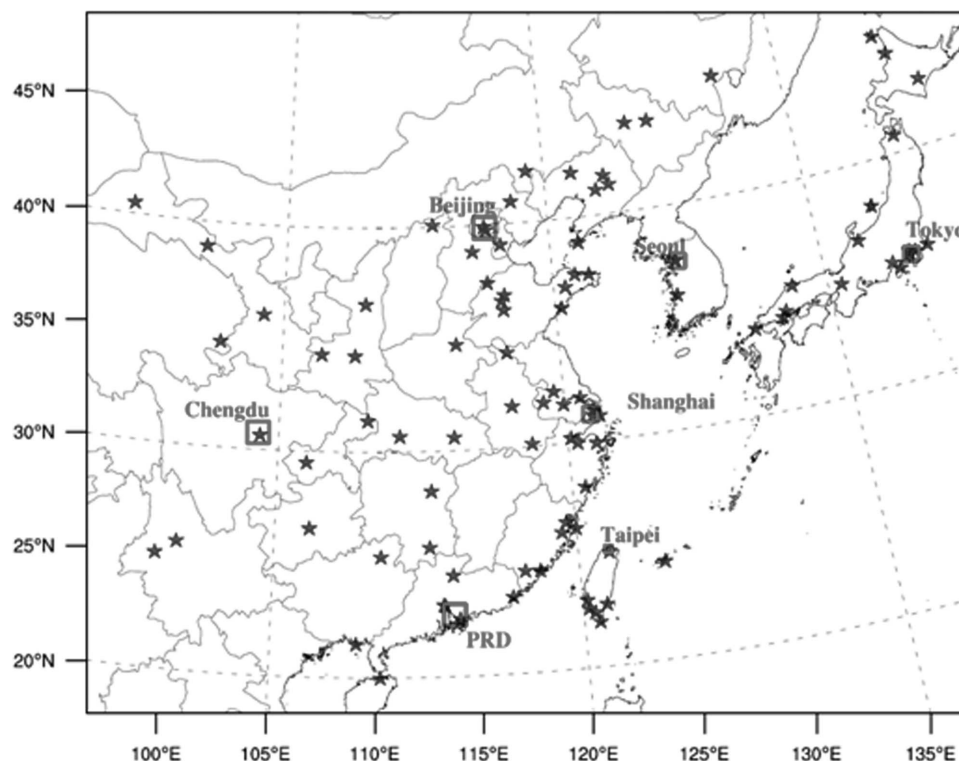


Figure 1. Model domain for the application of CMAQ in East Asia.

performance under TF HTAP scenarios has been evaluated in previous studies (Fiore et al., 2009). Thus, in this study, we evaluate the model with surface observations from National Climate Data Center (NCDC) for the MM5 meteorological field, and the Acid Deposition Monitoring Network (EANET) for CMAQ simulated surface O_3 in East Asia. Locations of the observational sites are shown in Figure 1.

We compare the evaluation results with the established statistical benchmarks (BM) for MM5 evaluation used by Kim et al (Kim et al., 2010) to demonstrate the credibility of MM5 performance in this study, as shown in Table 2.

According to very limited surface observational data over EA, we perform CMAQ model evaluations with hourly data at Taipei for O_3 and NO_x , as shown in Figure 2. CMAQ generally performs well but underestimates NO_x and overestimate O_3 in both January and July. Large negative biases of NO_x may be mainly due to underestimated emissions, as indicated in Streets et al.

(2003). In order to better determine the reason for overestimation of O_3 , we performed statistical analysis by separating the daytime O_3 and nighttime O_3 . Daytime O_3 has a much smaller bias (9.3 in January, 0.19 in July) than nighttime O_3 (16.02 in January, 9.38 in July), indicating the model overestimation of O_3 may be mainly due to the underestimation of NO_x , which usually has a strong titration effect during the night time, especially in the most populated yet polluted urban areas in EA.

Results and Discussion

Seasonal variation of surface O_3 over EA

It has been recognized that surface O_3 distribution is closely related to meteorological conditions, and that wind patterns and temperature are the most important factors (Dawson et al., 2007). EA is greatly affected by monsoons in both winter and summer.

Table 2. Model evaluation for MM5

	Wind speed (m/s)			Wind direction (degree)			Temperature (K)		
	BM	JAN	JUL	BM	JAN	JUL	BM	JAN	JUL
RMSE	≤ 2	2.18	2.17						
Bias	≤ 0.5	0.19	0.06	≤ 10	6.68	0.4	≤ 0.5	0.98	0.68
Gross error				≤ 0	52.39	49.35	≤ 2	2.56	2.18
IOA	≥ 0.6	0.60	0.62				≥ 0.8	0.96	0.83

Note: RMSE: root mean square error. IOA: index of agreement (Willmont, 1981).

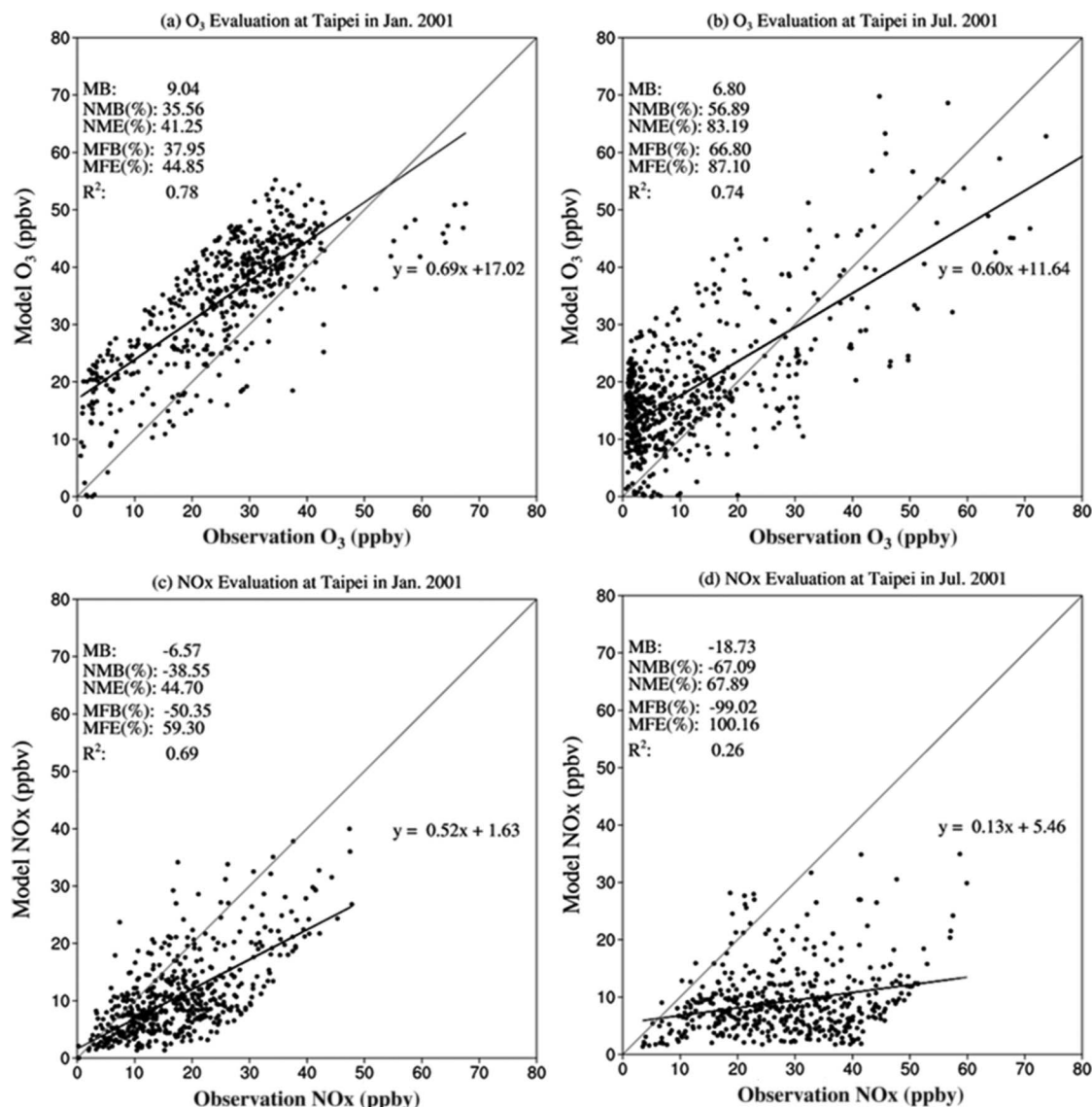


Figure 2. CMAQ model simulations evaluate against surface observations at Taipei for O₃ (first row) and NO_x (second row) in January (left column) and July (right column), 2001.

In fact, this is one of the major controlling mechanisms affecting the spatial distribution of EA surface O₃. Temperature also varies greatly between winter and summertime for EA. Thus, the combined effects from the temperature gradient and wind patterns lead to different spatial patterns of surface O₃ in January and July. As shown in Figure 3, (a) and (b) illustrate the spatial distribution of monthly average wind and temperature from MM5; (c) and (d) show the monthly average O₃ from CMAQ. In January, temperature decreases gradually from around 300 K in the southern region of EA to 250 K in the northern region. During the same period, a Siberian high-pressure system drives air from northern China southeastward. Pollution contained in this air mass is the result of extensive utilization of coal for heating within this region. The temperature gradient and southeastward wind lead to lower O₃ due to weak photochemical production in the north part of EA, and a higher O₃ in the south part due to downstream pollution from the upper latitudes, especially along the east coast

of China, where surface O₃ gradually increases from north to south. In July, temperature is almost within the same scale (around 300 K) for most areas of EA, resulting in similar photochemical production rates for O₃ in the north and south part. In summer, an intensified North Pacific high-pressure system expands westward and leads to the southeast summer monsoon in EA. Along with the northeastward SA summer monsoon coming from the Indian Ocean, the prevailing monsoon brings a clean marine air mass, and leads to the gradually increasing pattern of O₃ from south to north over EA. In addition, titration resulting from high emission of NO_x and low emission of VOC is another important factor in determining ozone concentration in the northern part of EA especially in winter. As demonstrated in Figure 3c, the most populated areas including BJ and SH actually have the lowest ozone concentration, indicating a VOC-limited sensitivity case in these areas. We continue this discussion in the next subsection.

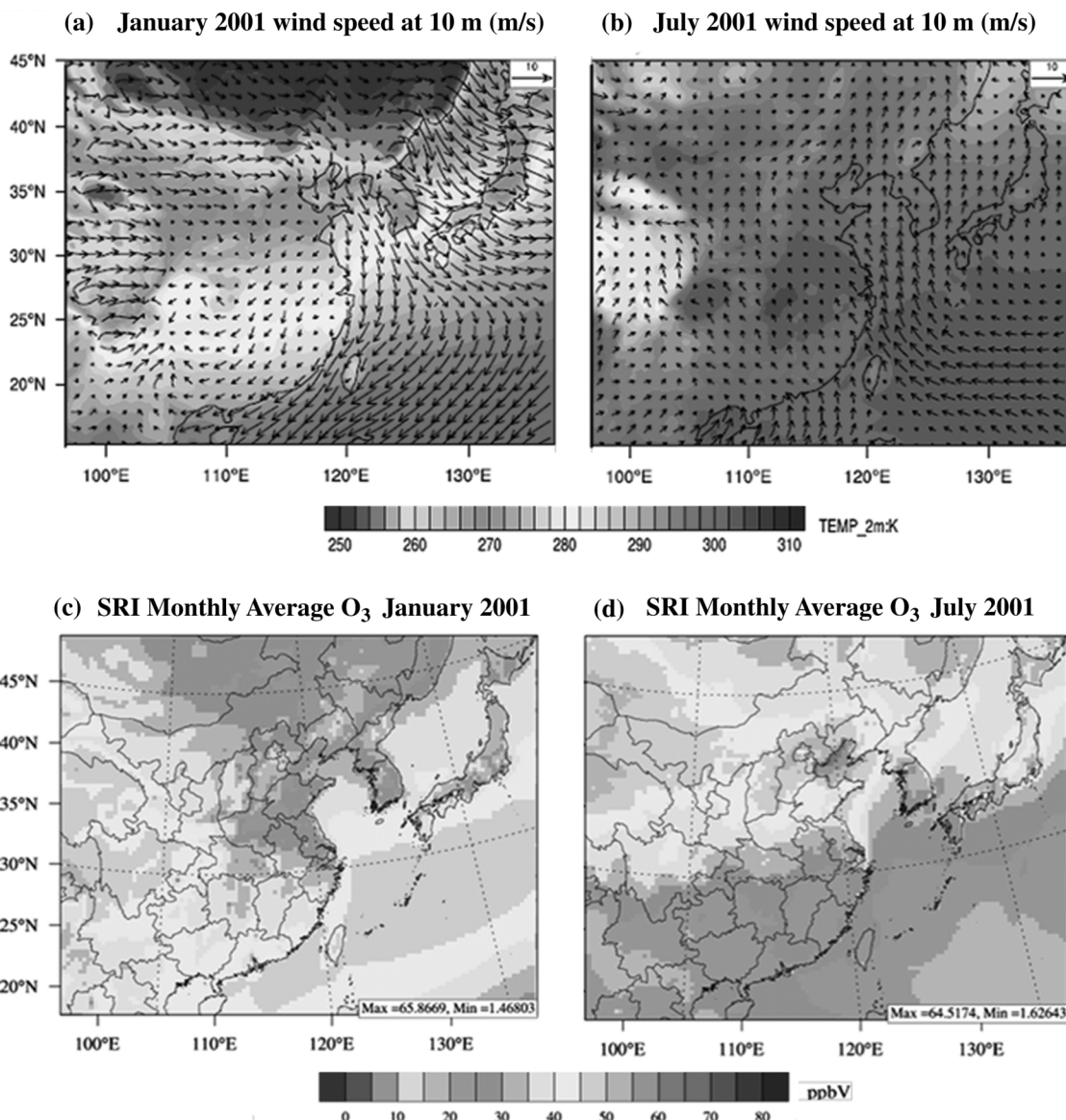


Figure 3. (a) Monthly mean wind speed, wind directions, and temperature in January and (b) July 2001. (c) Monthly mean hourly O_3 spatial distribution in January and (d) July 2001.

Analysis of O_3 sensitivity and linearity

In addition to meteorological conditions, chemical production is another dominant mechanism determining surface O_3 concentration, which is mainly affected by precursors including NO_x and VOC. Figure 4 illustrates the O_3 changes over EA as a result of emission reduction of NO_x or VOC under different control scenarios, which shows the contributions of NO_x and VOC in O_3 production over different areas of EA.

As shown in Figure 4, in January, urban areas including Japan, South Korea, northeastern China, Yangtze River Delta, and PRD have O_3 concentration increased by up to 28.88 ppbv under the 100% NO_x emission reduction scenario, and decreased by up to 12.52 ppbv under the 100% VOC emission reduction scenario, indicating a VOC-limited feature over these areas. This is consistent with our expectation, since

most of the developed industries and heavy transportations are distributed within the areas just mentioned, and have an extensive amount of NO_x emission. On the other hand, vegetation coverage is relatively lower in those areas compared to the south part of EA, which provides limited VOC emission especially in the wintertime. The other areas of EA are NO_x limited due to a large amount of VOC emission from vegetation and less NO_x emission from industry activities and vehicles. In July, as shown in Figures 4g to 4l, most of the areas that are VOC limited in January change to being NO_x limited, due to the significantly increased amount of VOC emission from vegetation in the summertime in these mid-latitude regions. However, some large metropolitan areas, including BJ, SH, TP, Tokyo, and Seoul, remain VOC limited due to the extensive amount of NO_x emission from vehicles and industries in the summer time.

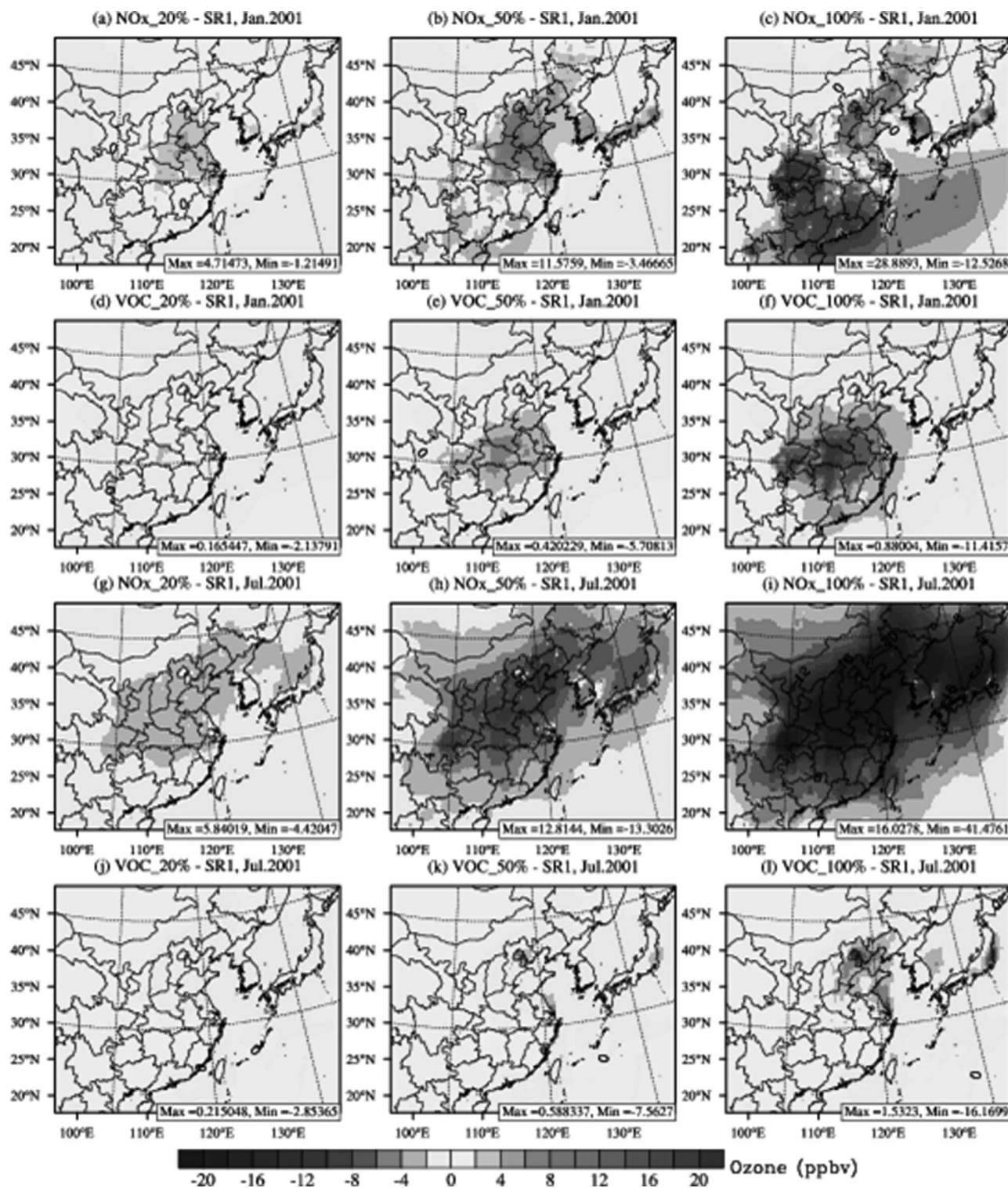


Figure 4. Spatial distribution of monthly average O₃ responses to local NO_x emission reduction at rate of (a) 20%, (b) 50%, and (c) 100% in January, and (g) 20%, (h) 50%, and (i) 100% in July 2001; and the O₃ responses to VOC emission reduction at rate of (d) 20%, (e) 50%, and (f) 100% in January, and (j) 20%, (k) 50%, and (l) 100% in July 2001.

In January, NO_x reduction scenarios show a more beneficial effect on the surface concentration of O₃ in the south part of EA, while VOC reductions are more beneficial in the north part. In July, NO_x reduction is beneficial for almost the whole EA. In

order to reveal the features of O₃ sensitivity and linear (or non-linear) response in the major urban areas as well as whole EA domain, we further analyze the O₃ concentration changes for these areas as demonstrated in Figure 5.

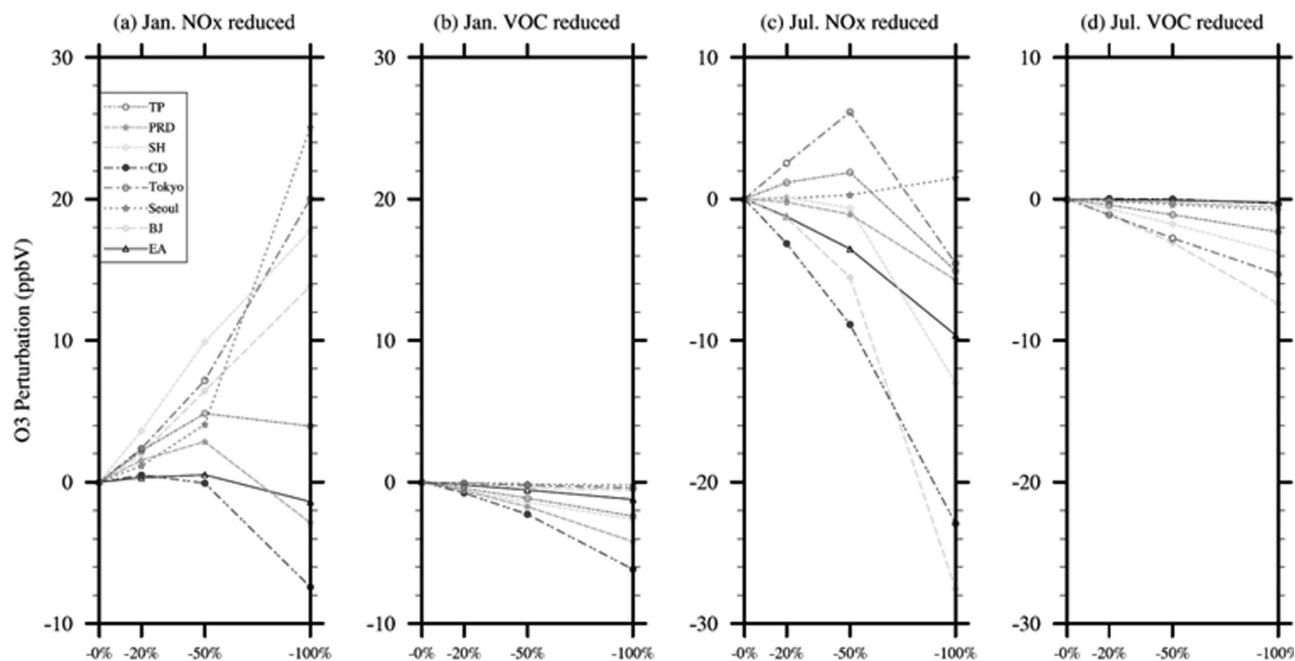


Figure 5. O₃ response (sensitivity cases: base) for EA and seven urban areas, resulted from local emission reduction for (a) NO_x and (b) VOC in January, and for (c) NO_x and (d) VOC in July 2001.

As shown in Figure 5a, in January, O₃ concentrations in most of the urban areas increase as the anthropogenic NO_x emission is reduced 20% at the beginning, implying typical VOC-limited conditions there. However, these urban areas exhibit different results in terms of linearity of O₃ production depending on NO_x emission reduction ratios, which could be divided into two groups. The first group includes Seoul, SH, BJ, and Tokyo, where surface O₃ concentrations keep increasing as the anthropogenic NO_x emission reduction ratio increases from 20% to 100%. This monotonic increment indicates that these areas have a significant amount of NO_x emission and are seriously VOC limited. Jacob et al. (1995) claimed that the nonlinearity of O₃ production response for NO_x emission changes in the United States. In this study, however, we find that in EA, for city areas including BJ, SH, and Tokyo, O₃ response is almost linear with respect to NO_x reduction. For Seoul, the O₃ concentration increment is even faster when the reduction rate change moves from 50% to 100%. The linearity in these areas indicates the existence of significantly high NO_x concentration, which indicates that the production of HNO₃ is the dominant sink for odd hydrogen, and which results in VOC-limited conditions of O₃ production for these urban areas. Under the NO_x reduction scenarios, O₃ depletion through titration in nighttime is slowed down in these areas with the removal of anthropogenic NO_x emission. The second group includes TP, CD, and PRD, which are also VOC limited in the beginning where O₃ concentrations increase with the NO_x emission reduction increases from 20% to 50%, but gradually change to being NO_x limited when the reduction rate of NO_x continues to increase from 50% to 100%, implying a nonlinear response for O₃ production for these urban areas.

In July, both VOC and NO_x reductions lead to an O₃ decrease in most of the areas except for TP, Tokyo, and Seoul, where there

is a slight increase in O₃ under 20% and 50% NO_x reduction scenarios. TP is located in the south part of EA where temperature and humidity do not differ significantly between January and July, so it is still VOC limited under 20% and 50% NO_x reduction scenarios, but when the removal rates continually increase to 100%, TP switches to being NO_x limited, which is similar for Tokyo. Seoul shows monotonic increase of O₃ under all NO_x reduction scenarios, indicating a highly unbalanced ratio between NO_x and VOC emission.

In contrast to the obvious negative correlation between O₃ production and NO_x reduction, VOC emission reduction shows a strong linear response for all of the urban areas in both January and July, which is consistent with implications from NO_x reduction scenarios indicating the typical VOC-limited conditions in city areas. However, under the same percentage of emission removal rate at 100%, VOC reductions show less significant impact on the change of O₃ than that from NO_x. For example, 100% reduction of NO_x could result in up to 27.5 ppbv reduction of O₃ in July for BJ, while 100% reduction of VOC only leads to 7.2 ppbv O₃ reduction there. Consequently, anthropogenic NO_x contributes more than VOC to the surface O₃ over these areas, but emission control strategy design should be more careful since it is generally impossible to have an actual reduction rate of 100%, and the real reduction may fall into VOC-limited conditions and thus increase the surface O₃ concentration in EA, especially for large metropolitan areas.

Effects of long-range transport

In order to better understand the O₃ change due to NO_x reduction, three sensitivity cases are conducted to study the effects of long range transport from EU, SA, and NA under the

HTAP SR3NA, SR3SA, and SR3EU scenarios (Fiore et al., 2009).

Figure 6 illustrates the spatial distribution of monthly average surface O₃ response over EA as a receptor region to 20% reduction of anthropogenic NO_x emission from EU, SA, and NA as source regions respectively. All contributions discussed in this section refer to the resulting O₃ change from 20% NO_x reduction scenarios in this section. As shown in Figure 6a, long-range transport from EU widely impacts air quality over the whole

EA in January, contributing up to 0.34 ppbv of the surface O₃ in the northwest part of EA, and 0.1–0.2 ppbv of the surface O₃ over the northern China plain, Yangtze River Delta (YRD), and the East Ocean along the trace of winter time southeastward monsoon, reflecting its upwind proximity. While in July as shown in Figure 6b, the transport effect from EU is constrained by the summer monsoon from the South China Sea within the northwest part of EA but with a higher contribution of up to 0.60 ppbv for the surface O₃, indicating that the dominant distribution

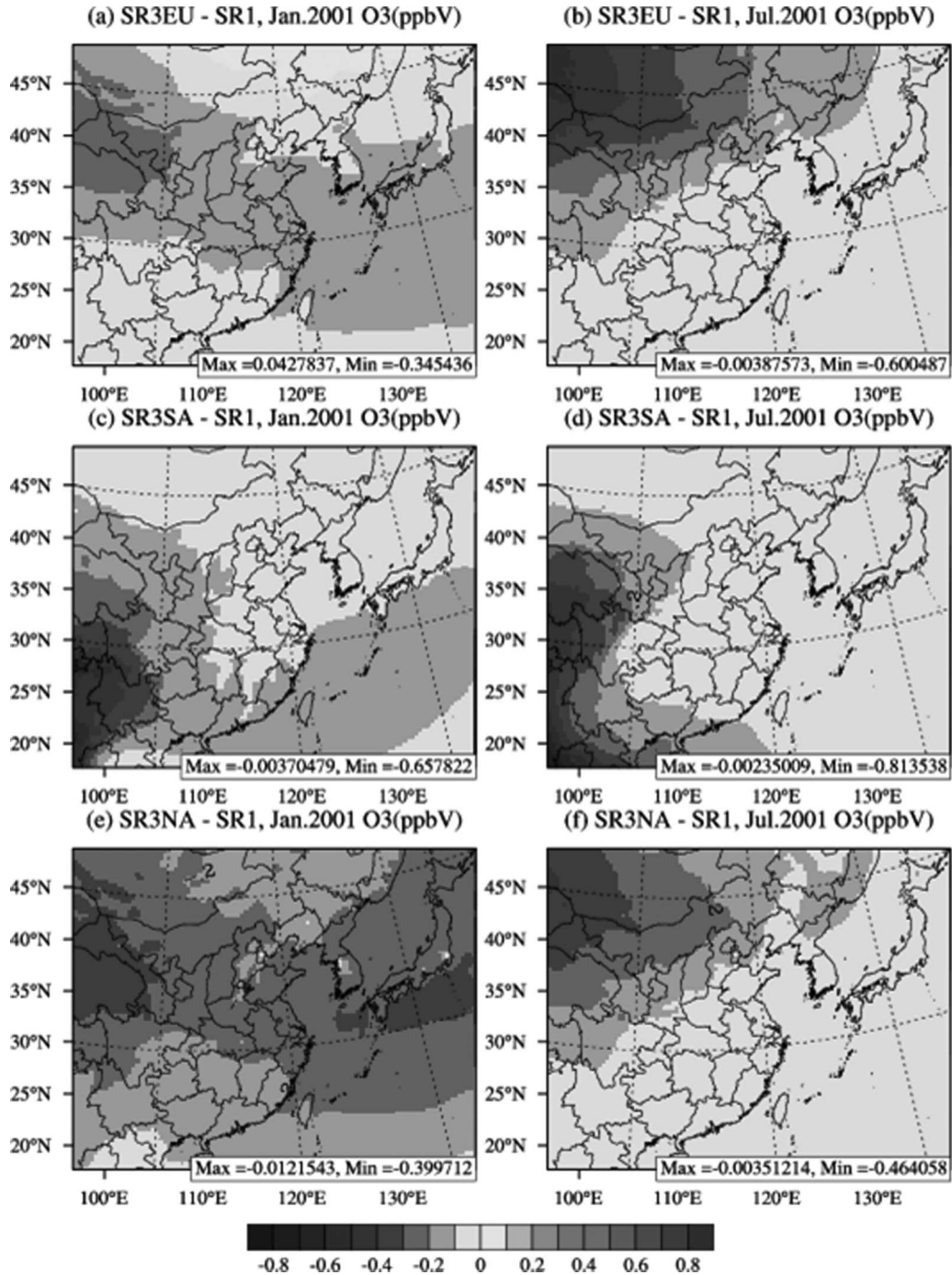


Figure 6. Monthly average O₃ responses over EA as receptor region from 20% reduction of anthropogenic NO_x emissions of (a) EU, (c) SA, and (e) NA in January, and (b) EU, (d) SA, and (f) NA in July 2001.

of the transport from EU is mainly determined by the monsoon. Contribution of transport from SA is up to 0.65 ppbv in January and 0.81 ppbv in July, which is constrained in the southwest part of EA by the Himalayas in both months, although the winter monsoon helps to extend the long-range transport along the coast of China. NA has the most significant impact over EA in January, contributing to 0.1–0.3 ppbv for the whole domain. However, the impact of NA is also constrained in the northwest part of EA in the summer time.

In order to examine the relative importance of intercontinental influence from each source region, long-range transport scenarios are also compared with each other and also with the local NO_x reduction scenarios to demonstrate the contributions of anthropogenic NO_x emissions from local and long-range transport over EA.

Figures 7a, 7b, and 7c illustrate the monthly average and maximum daily O₃ responses for EA and the urban areas under SR3EU, SR3SA, and SR3NA scenarios, respectively. The 20% NO_x emission reduction from EU leads to a monthly average surface O₃ decrease of 0.11 ppbv in January and 0.10 ppbv in July over EA, with the maximum daily O₃ changes ranging from 0.23 ppbv in January to 0.28 ppbv in July. This long-range transport effect varies a lot for the seven urban areas. The most affected city is BJ, where resulting monthly average O₃ response is 0.07 ppbv in January and 0.14 ppbv in July. BJ also has the largest maximum daily O₃ response, which is 0.31 ppbv in January and 0.49 ppbv in July, reflecting a downwind location of BJ for both the summer monsoon and the EU emission transport pathway. The transport effect of SA is within the same scale as that of EU, which is 0.13 ppbv in January and 0.11 ppbv in July, but mainly affects the southwest part of EA and is not much affected by the monsoon, resulting in similar O₃ changes in January and July. CD is the most affected urban area by SA in both January and July, where resulting O₃ changes are

0.18 and 0.12 ppbv, respectively, and maximum daily changes reach up to 0.5 and 0.65 ppbv, respectively. Among the three long-range transport scenarios, NA contributes the most as a source region to surface O₃ concentration over EA, and the urban areas in January as the monthly mean O₃ response reached up to 0.23 ppbv and 0.5 ppbv for maximum daily response. In July, transport effects over EA from the NA falls into same scale as other source regions. Thus, impacts from intercontinental transport from NA show the largest seasonal variations, especially for the coastal areas including SH, PRD, and TP, which are located in the upwind direction in the summer monsoon. For example, the 20% NO_x emission reduction of NA leads to 0.38 ppbv daily maximum O₃ decrease for TP in January, while the value is less than 0.01 ppbv in July. O₃ responses caused by NA at the urban areas are generally within the same scale as EA, indicating that the long-range transport effect of NA is widely spread over the whole EA. Intercontinental transport effects are generally stronger in January than that in July, indicating that the wintertime monsoon dominates the imported intercontinental transport of O₃ and its precursors for EA.

Figure 7d shows the monthly average responses under the 20% of local NO_x emission reduction scenario (denoted as SR3local). For the entire EA domain, O₃ response to local emission change is 0.32 ppbv in January, which is even smaller than the impact from NA as 0.35 ppbv for daily maximum O₃ response. The impacts from EU and SA are 0.23 and 0.31 ppbv, respectively. Fiore et al. (2009) also reported the annual surface O₃ decrease from a 21-CTM ensemble over EA; values are 0.11, 0.11, and 0.10 ppbv for SR3NA, SR3EU, and SR3SA respectively, and 0.60 ppbv for domestic 20% NO_x emission reduction, which is consistent with our CMAQ result, but indicates a smaller significant factor for intercontinental influences based on the global model results. This difference implies uncertainty within the scalability of global model regarding quantifying

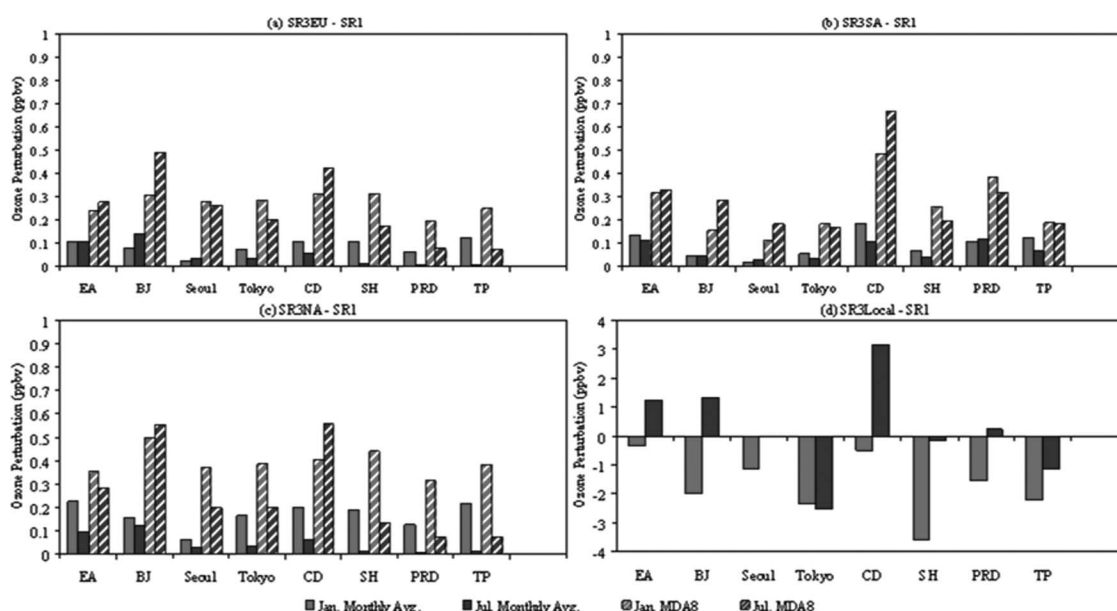


Figure 7. Monthly average and maximum daily O₃ responses under intercontinental transport scenario from (a) EU, (b) SA, and (c) NA, and (d) O₃ perturbation under local 20% NO_x emission reduction over EA.

the contributions from source regions, but shows, nevertheless, the comparable impacts as local emission.

For those selected cities in EA, local NO_x emission reduction generally has an impact of more than 2 ppbv, while the contributions from outside source regions are usually less than 0.2 ppbv for monthly average and less than 0.7 ppbv for maximum daily, indicating that local emission is the dominant contributor of air pollutants in city areas, but that long-range transport likely contributes greatly to the local pollution as well. For example, SH has less than 0.1 ppbv monthly average O₃ response under local 20% NO_x emission reduction in July, while the transport impacts under 20% NO_x emission reduction from EU, SA, and NA could reach around 0.2 ppbv for maximum daily O₃ response in SH in July. However, small values for monthly average O₃ response under local NO_x emission reduction do not necessarily indicate that local emissions are unimportant, as SH also demonstrated a large daily fluctuation for O₃ response in both January and July, a fluctuation that is not shown in this paper due to its minimal relevance to this study. This indicates the necessity to conduct more comprehensive designed sensitivity simulations over EA to investigate the detailed temporal and spatial variations of local O₃ sensitivities.

This comparison illustrates that although the long-range transport effect is generally smaller in urban areas than the local emission effect, its contribution is still important especially in winter, and this influence could be as high as the impact caused by local emission in certain urban areas, indicating the necessity to take into consideration of long-range transport impact for strategy design since the transport effect could be comparable to the local emission.

Conclusions

In this study, the regional modeling system MM5/CMAQ is applied with 20%, 50%, and 100% emission reduction scenarios of NO_x and VOC, respectively and independently, in both January and July 2001 to understand the O₃ formation over EA. Both monthly average and maximum daily O₃ responses are examined for EA and seven urban areas in order to understand the sensitivity and linearity of O₃ production dependence on NO_x and VOC, and to provide a theoretical basis for the development of the integrated emission control strategies.

Seasonal variation between January and July suggests that spatial distribution and transport of surface O₃ is strongly affected by monsoon and temperature over EA, resulting in lower O₃ concentration in the north part of EA in January due to less ultraviolet (UV) radiation and southeastward monsoon, and higher O₃ concentration in the same area in July due to increasing temperature and northwestward monsoon.

Investigation of O₃ responses to local NO_x emission changes demonstrates that the north part of EA behaves differently from the south part in terms of both seasonal variation and O₃ sensitivity. The northeast part of EA, including north China, South Korea, and Japan, is mainly VOC limited in January and NO_x limited in July, although the metropolitan areas including BJ, Seoul, and Tokyo are always VOC limited due to significant amounts of NO_x emission from power plants and vehicles. The

southeast part is generally NO_x limited in both January and July, except for urban areas including PRD and TP, due to large NO_x emissions with the seasonal monsoon. The west part of EA is NO_x limited in both January and July due to less developed industry and transportation.

Linearity of O₃ response is examined for seven urban areas regarding both NO_x and VOC emission changes independently. We find strong negative linearity of O₃ response to NO_x reduction in BJ, SH, Seoul, and Tokyo in January, and strong positive linearity of O₃ response to VOC emission reduction in all the seven urban areas in both January and July. Moreover, O₃ responses under 20% of local NO_x emission reduction scenarios ranges from 0.7 ppbv in January to −1.2 ppbv in July, which is on a larger scale than that under 20% of VOC emission reduction scenarios, ranging from −0.2 ppbv in January to −0.02 ppbv in July, indicating surface O₃ is more affected by NO_x emission than by VOC in EA. However, this assessment should be taken carefully while making emission control strategies, since applicable NO_x reduction rates may fall into the VOC-limited condition, which causes surface O₃ increases with decreasing anthropogenic NO_x emission.

Intercontinental transport effects from EU, SA, and NA over EA and seven urban areas are also examined under HTAP SR3EU, SR3SA, and SR3NA scenarios, and compared with the local emission reduction scenario to distinguish the contribution of domestic emission from that of long-range transport on surface O₃. As a receptor region, EA benefits from NO_x emission reduction in EU, SA, and NA since the surface O₃ concentration decreases in both January and July. The transport effect is found to be higher in January for all source regions than that in July, and the daily maximum O₃ change from long-range transport could be as high as the monthly O₃ change from local NO_x emission reduction at the same rate. Contributions from EU and SA are mainly trapped by the dominant monsoon, but NA is found to affect EA, spreading over the entire domain. EU transport effect could reach as far as Japan in January, but is mainly constrained in northwest part of EA in July, while SA transport effect is trapped by the Himalayas in the southwest part of EA. The overall intercontinental transport effect is smaller than the effect from domestic NO_x emission reduction by the same rate, which is consistent among the regional model CMAQ and the TF HTAP 21 global CTM ensemble, but could still cause important O₃ changes comparable to local emission in January, and this impact may even reach the same scale as local emission in urban areas. Taking into consideration that global CTMs may underestimate their contributions from source regions due to uncertainties from coarse resolution, the intercontinental transport influences may be larger than that reported by HTAP.

This study illustrates the effectiveness of the regional modeling system MM5/CMAQ to analyze the O₃ sensitivity and linearity over EA, and also examines the impacts of North Hemispheric transport of O₃. Pending is further investigation of the detailed chemical schemes and physical transport processes regarding O₃ and PM_{2.5} as a response to different spatial resolutions at finer scale (12 km), and revealing the seasonal variations through a whole year.

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