Deforestation Reduces Downwind Precipitation and

2 Soybean Yields in Brazil

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ABSTRACT

Rapid agricultural expansion has driven unprecedented forest loss worldwide. However, deforestation reduces evaporation locally, potentially triggering precipitation declines in downwind croplands far beyond agricultural frontiers, with consequent impacts on regional crop yields. Using a novel integration of Lagrangian atmospheric moisture tracking and statistical crop modeling, we quantified impacts of evaporation declines on precipitation and soybean yields across Brazilian soybean-producing states during 1982–2015. Our analysis reveals that tree evaporation (i.e., transpiration and canopy interception) contributes roughly one-third of precipitation during the soybean growing season, but recent deforestation has decreased seasonal precipitation by 171 to 326 mm (9-16%). This impact was most pronounced in tropical states near the arc of deforestation in the southern Amazonia, yet the largest reductions in soybean yields (288 kton, or 8% loss) occurred in the southern Brazilian state of Rio Grande do Sul. Cumulatively, deforestation-driven precipitation declines resulted in a total soybean production loss of ~900 kton (3%) in Brazil. These findings demonstrate a self-reinforcing feedback between upwind deforestation and downwind agricultural productivity—revealing a spatial trade-off where agricultural frontier expansion into forests undermines crop yields particularly in established croplands, which in turn may drive demand for additional forest clearing to compensate for these losses. Under continued deforestation and climate change, the feedback loop between forest clearing and agricultural losses is likely to intensify, threatening Brazil's soybean production into the future.

Main Text

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Global food security faces unprecedented challenges, requiring a substantial increase in agricultural production by 2050 to meet rising population needs and biofuel consumption^{1–3}. This challenge is intensified by increasing climate extremes, such as droughts and heatwaves, during crop growing seasons^{4,5}. In response, many countries, particularly in South America, are expanding their agricultural frontiers into natural vegetation regions. As the world's leading soybean producer (~34% of global production), Brazil has converted over 20 million hectares of forests and savannas into soybean croplands since 2000⁶. However, a growing body of research suggests that this forest cover loss can have profound impacts on regional precipitation patterns, reducing precipitation volumes^{7,8}, increasing drought frequency^{9,10}, and delaying rainy season onset^{11–13}. Despite these documented precipitation effects, their consequent impacts on agricultural productivity remain largely unquantified at regional scales.

The Amazon rainforest acts as a critical moisture pump, evaporating and recycling 23 vast amounts of water that supports rainfall across key agricultural regions in South America¹⁴. In particular, Amazonian moisture transport through the South American 25 low-level jet provides 16–25% of annual precipitation over the La Plata river basin¹⁵, 26 where Brazil's southernmost croplands are located. However, as the agricultural frontier expands at the expense of forests, reductions in tree evaporation (i.e., the sum of tree transpiration and canopy interception) have the potential to alter precipitation patterns, far beyond the deforestation regions to the south 16-18. Agricultural systems 30 critically depend on reliable growing-season rainfall 19-21, particularly as documented in southern Brazil²², yet quantifying the extent to which changes in tree evaporation propagate through the atmosphere to affect downwind agricultural productivity remains a key research challenge.

To address this gap, we developed an innovative framework combining atmo-

spheric moisture tracking²³ and statistical crop modeling to quantify the cascading impacts of deforestation-driven tree evaporation declines on regional precipitation and agricultural productivity. This framework, demonstrated here using Brazilian soybean production²⁴, provides a quantitative foundation for understanding forest—precipitation interactions and agricultural implications worldwide. Our findings reveal that agricultural expansion into forests has the potential to create a self-reinforcing feedback in Brazil, where deforestation can reduce regional precipitation and crop yields, potentially enhancing food demand and exacerbating the need for further agricultural expansion²⁵. This mechanism likely operates in other major deforestation hotspots for agricultural expansion, such as the Congo Basin and Southeast Asian rainforests, suggesting that maintaining forest cover is crucial for sustaining global agricultural productivity under escalating climate change.

Tree cover and evaporation declines

During the period of 1982–2015, South America experienced unprecedented levels of 49 tree cover loss, with the arc of deforestation (ARC, hatches in Figure 1a)—located 50 between southern Amazonia and the northern Cerrado—emerging as one of Earth's most intensively deforested regions²⁶. Using remotely sensed forest cover observa-52 tions²⁷, we found that tree cover in this region declined by \sim 30% during the study 53 period, while deforestation progressively expanded southward into Paraguay and northern Argentina (Figure 1a). Agricultural expansion, primarily driven by global commodity demands, was the primary driver of tree cover loss in South America^{28,29}. Consequently, deforestation led to widespread reductions in tree evaporation during the soybean growing season (August–March), with the most significant reductions exceeding 300 mm—observed in the ARC again (Figure 1b). These declines also persisted across all soybean growth stages (Supplementary Figure S1).

To assess the implications of precipitation changes on soybean yield variability

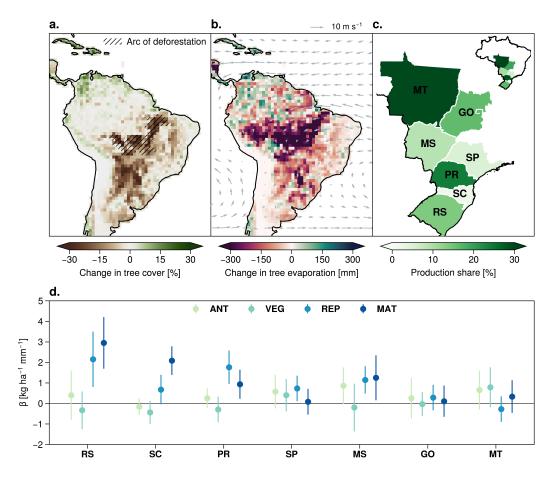


Figure 1. Deforestation, tree evaporation declines, and soybean yield sensitivity to precipitation. a. Change in tree cover fraction (%), calculated as the linear trend multiplied by the study period length (34 years, 1982–2015). The arc of deforestation (ARC) in southern Amazonia is highlighted with black hatches. b. Change in tree evaporation (mm) during the soybean growth period (August–March), with arrows indicating prevailing wind direction and magnitude. Growth stage-specific evaporation changes are presented in Supplementary Figure S1. c. Share (%) of national soybean production across major soybean-producing states (PR: Paraná, SC: Santa Catarina, RS: Rio Grande do Sul, SP: São Paulo, MS: Mato Grosso do Sul, MT: Mato Grosso, GO: Goiás), calculated by multiplying soybean yield (kg ha⁻¹) by harvested area (ha) in 2000 at 0.5-degree resolution and aggregated to state and country levels. **d.** Sensitivity of soybean yields to precipitation (β , kg ha⁻¹ mm⁻¹) during distinct growth stages (ANT: antecedent—August and September, VEG: vegetative—October and November, REP: reproductive—December and January, MAT: maturation—February and March). Points represent mean values with 95% confidence intervals.

across Brazilian soybean states (Figure 1c), we developed a regional-scale statistical crop modelling approach, using the least absolute selection and shrinkage operator (LASSO) regression (see Methods). These models demonstrated strong performance against observations (Supplementary Figure S2 and Supplementary Figure S3). Our results show that, soybean yields in southern Brazilian states—Rio Grande do Sul (RS), Santa Catarina (SC), and Paraná (PR)—exhibited strong sensitivity to precipitation during reproductive and maturation stages, with sensitivity coefficients exceeding 2.0 kg ha⁻¹ mm⁻¹ during these critical periods (Figure 1d) when atmospheric water demand peaks and reproductive tissues are most vulnerable to drought stress^{30,31}. In contrast, tropical soybean yields, particularly in Mato Grosso (MT) and Goiás (GO), showed lower sensitivity to precipitation (Figure 1d) and positive sensitivity to temperature (Supplementary Figure S4), reflecting their energy-limited growth conditions^{5,32}.

₇₅ Impacts of tree evaporation changes on precipitation

To identify moisture source regions where precipitation originated from tree evaporation, or in other words, where tree evaporation (E_{tree}) contributed to precipitation 77 (hereafter referred to as $E2P_{tree}$), in Brazil's major soybean-producing states, we 78 employed a Lagrangian-based moisture tracking model integrated with satellite ob-79 servations (see Methods). Results show, tree evaporation over South America was 80 a major contributor to precipitation across states, where moisture source regions 81 extended from the Amazon through central Brazil into the La Plata basin, with the highest contributions (>100 mm) concentrated in southern Amazonia (Figure 2a-d). The contribution varied seasonally: 403 mm (37% of the accumulated precipitation of 1090 mm), increasing to 711 mm (33%) in the vegetative stage, peaking at 949 mm (32%) during reproduction, before declining to 690 mm (27%) in maturation (pie charts in Figure 2a–d). The importance of tree evaporation for precipitation becomes

even more crucial at the state level (Figure 2e). MT, bordering Amazonia, exhibited
the strongest dependence, receiving 47–265 mm (39–56%) precipitation from tree
evaporation across growth stages, but this percent decreased to 64–94 mm (26–37%)
in the southernmost state of RS. Moreover, the ARC region alone contributed 17–
26% to MT's precipitation (92–110 mm, outlined bars in Figure 2e), highlighting the
importance of the deforestation hotspot for agricultural water supply in Brazil.

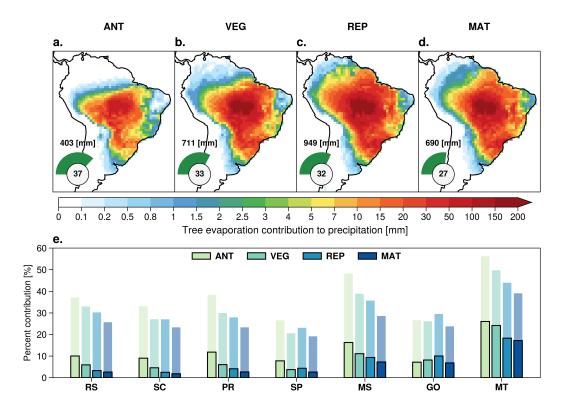


Figure 2. Tree evaporation contribution to precipitation. a–d. Spatial pattern of climatological tree evaporation contributions to precipitation ($E2P_{tree}$, mm) during soybean growth stages: a) antecedent (ANT), b) vegetative (VEG), c) reproductive (REP), and d) maturation (MAT), aggregated across all seven states (state-specific patterns in Supplementary Figure S5). Pie charts show the percentage contribution (%) relative to total precipitation (P), with absolute values (mm) indicated above. e. Growth stage-specific tree evaporation contributions (%) across states. Shaded bars represent the percentage of precipitation that originated from tree evaporation across continental South America (relative to total precipitation), while outlined bars show only the portion from the arc of deforestation (ARC). For example, if total precipitation is 500 mm, with 100 mm from South American tree evaporation (including 40 mm from the ARC), the shaded and outlined bars would show 20% and 8%, respectively. Absolute $E2P_{tree}$ values (mm) are provided in Supplementary Figure S6.

We next developed a statistical decomposition method to quantify the impacts of

tree evaporation changes under historical deforestation on precipitation (see Methods). Our results showed that tree evaporation declines substantially reduced precipitation over soybean regions, with the most pronounced impacts—exceeding 70 mm declines—observed within the ARC (Figure 3a-d). These reductions followed similar seasonal variability, with -171 mm (16% of accumulated precipitation) during the 99 antecedent period, -229 mm (11%) in the vegetative stage, -326 mm (11%) during 100 reproduction, and -226 mm (9%) during maturation. The magnitude of these im-101 pacts also varied significantly across states (shaded bars in Figure 3e), with MT 102 experiencing the largest reductions up to 107 mm (\sim 18%) during reproduction, and 103 the effects declining southward (20–41 mm, 7–16%). Importantly, tree evaporation 104 changes within the ARC accounted for more than half of the precipitation decline in 105 tropical states (outlined bars in Figure 3e). The magnitude and consistency of these 106 deforestation-induced precipitation declines suggest that their impacts on agricultural 107 productivity are likely to be widespread. 108

Estimated reductions in soybean yields under deforestation

We next examined how widespread tree evaporation declines affected soybean 110 yields in Brazil, by combining the yield sensitivity to precipitation (Figure 1d) with 111 absolute precipitation reductions driven by tree evaporation changes (Supplementary 112 Figure S9). In tropical states, although yield sensitivity to precipitation was relatively 113 lower, substantial precipitation reductions led to a considerable yield loss (Figure 4a). 114 For instance, in Mato Grosso do Sul (MS), tree evaporation declines resulted in a total 115 yield loss of 119 kg ha⁻¹ during the growing season. In contrast, southern states, 116 despite experiencing smaller precipitation reductions, suffered greater losses due to their higher yield sensitivity. For example, RS lost 125 kg ha⁻¹ in total, with the most 118 severe reductions occurring during the critical reproductive (46 kg ha⁻¹ losses) and 119 maturation phases (63 kg ha^{-1} losses).

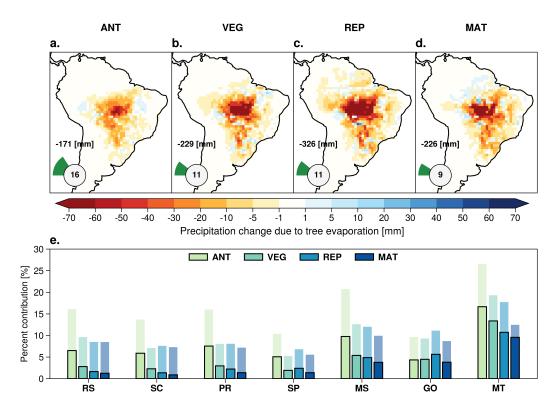


Figure 3. Precipitation reductions caused by tree evaporation declines. a–d. Spatial patterns of precipitation changes (mm) caused by tree evaporation during soybean growth stages: a) antecedent (ANT), b) vegetative (VEG), c) reproductive (REP), and d) maturation (MAT), aggregated across all seven Brazilian states (state-specific patterns in Supplementary Figure S8). Pie charts show these reductions as percentages of total precipitation (*P*), with absolute values (mm) indicated above. e. Growth stage-specific precipitation reductions (%) across states during 1982–2015. Shaded bars represent the percentage reduction in precipitation (relative to total precipitation) caused by tree evaporation across continental South America, while outlined bars show only the portion from the arc of deforestation (ARC). For example, in a region receiving 500 mm total precipitation, if tree evaporation decline causes a 60 mm reduction (including 40 mm from the ARC), the shaded and outlined bars would show -12% and -8%, respectively. Absolute reduction values are provided in Supplementary Figure S9.

We further scaled these yield impacts to estimate total production losses using soybean harvested areas circa 2000³³ as the reference baseline. The yield reductions translated to substantial production losses totaling 923 kton across all studied states (Figure 4b). RS experienced the largest absolute decline (288 kton, 8% of regional production), followed by MT (243 kton, 3%), PR (194 kton, 3%), and MS (124 kton, 5%). Based on average soybean prices during 2000–2015 (~400 per ton), these losses amounted to approximately 369 million dollars, demonstrating that deforestation creates hidden economic costs extending far beyond the deforestation frontier. Given substantial expansion of soybean cultivation after 2000⁶, these estimates are probably conservative.

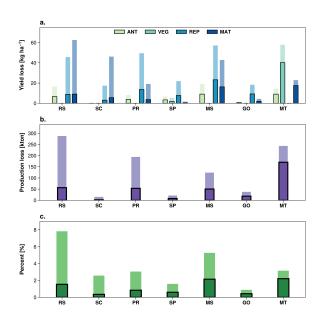


Figure 4. Soybean productivity losses due to tree evaporation-driven precipitation declines. a. Estimated soybean yield reductions (kg ha⁻¹) during growth stages, calculated by multiplying precipitation sensitivity (Figure 1d) by precipitation decreases caused by tree evaporation (Supplementary Figure S9). **b.** Absolute soybean production declines (kton), calculated as the product of total yield declines across stages and harvest area circa 2000. **c.** Percent of production declines relative to total production across states. Shaded bars represent impacts from tree evaporation declines across continental South America, while outlined bars show the portion specifically from the arc of deforestation (ARC).

Self-reinforcing feedback of agricultural expansion

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Our analysis reveals that forest clearing for agricultural expansion undermines soybean productivity at larger spatial scales through atmospheric moisture recycling and transport. Across Brazil, deforestation-induced evaporation declines substantially reduce precipitation (9–16%) in agricultural regions, resulting in a total soybean production loss of 923 kton (approximately 3% of national output), with impacts extending far beyond deforestation frontiers.

These yield reductions exhibited considerable spatial variability across Brazil. While direct precipitation reductions were most apparent in the immediate regions that are closest to the deforestation front (Figure 3e), the largest deforestation-induced reductions in soybean yield occurred in drier regions (Figure 4), where strong moisture limitations have increased the sensitivity of crops to water availability^{30,31}. This heightened sensitivity produced some of the most severe implications in southern Brazil's agricultural heartland, where even moderate precipitation reductions led to significant yield losses. Such vulnerability is particularly pronounced because these southern regions operate closer to critical water thresholds for optimal soybean development, particularly during reproductive stages (Figure 3e) when water stress most severely impacts pod formation and filling.

The study specifically focuses on how deforestation-driven evaporation declines affect precipitation and agricultural productivity. While deforestation also influences atmospheric (thermo)dynamics particularly in the small scale^{17,34}, previous research at comparable regional scales demonstrates that evaporation changes serve as the dominant driver of precipitation patterns in these contexts^{7,16,18}, supporting our methodological emphasis on evaporation impacts on precipitation and soybean in Brazil.

Consequently, these findings demonstrate that expansion into forests is directly

limiting the capacity for soybean production across broad spatial scales. While cur-157 rent deforestation-induced yield reduction (\sim 3%) may be temporarily offset by newly 158 cleared agricultural land, this balance is inherently unsustainable and increasingly 159 precarious²⁵. As deforestation continues alongside climate change^{35,36}, these hydro-160 logical impacts will likely intensify non-linearly³⁷, progressively constraining regional 161 agricultural productivity. Declining soybean productivity may further drive additional 162 demand for cropland expansion, triggering a self-reinforcing feedback loop that has 163 the potential to accelerate over the coming decades. This positive feedback operates 164 as agricultural forest clearing continues, causing precipitation reductions that lower 165 soybean yields, which in turn enhances the demand for further deforestation. 166

In conclusion, our findings reveal that forest cover loss significantly reduces downwind precipitation and soybean productivity far beyond deforestation regions. Therefore, maintaining forest cover is not merely an environmental concern but also an economic and political imperative for ensuring long-term agricultural outputs and food security³⁸. The self-reinforcing feedback loop between forest clearing and agricultural losses demonstrated here calls for integrated forest-agricultural management policies that recognize the connection between forest conservation and sustained agricultural productivity. Future agricultural planning must carefully consider the trade-offs between short-term expansion gains and the long-term productivity losses that result from disrupting critical hydrological cycles^{17,25}.

Methods

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178 Soybean states and growth stages

Our study focused on Brazil's seven major soybean-producing states, which account for approximately 90% of national production (Figure 1c). The southern states—Rio Grande do Sul (RS), Santa Catarina (SC), and Paraná (PR)—have a subtropical climate where soybean yields are highly sensitive to precipitation variability. In contrast,

the central and tropical states—São Paulo (SP), Mato Grosso do Sul (MS), Goiás (GO) and Mato Grosso (MT)—experience a seasonal wet-dry climate (Supplementary Figure S11), where agriculture is generally energy-limited during the soybean growing season. MT, Brazil's leading soybean producer, accounts for one-third of national output and intersects with the arc of deforestation (Figure 1a), a region experiencing substantial tree cover loss between 1982 and 2015. PR and RS contribute 13% and 11% of national production, respectively (Figure 1c).

We analyzed soybean yield response to precipitation across four critical growth 190 stages: the antecedent period (ANT, August-September) for establishing baseline soil 191 moisture conditions; the vegetative stage (VEG, October-November) characterized by 192 emergence and leaf development; the reproductive phase (REP, December-January) 193 encompassing flowering and pod formation; and the maturation stage (MAT, February-194 March) focusing on pod filling and yield determination. While growing season lengths 195 vary across regions, this standardized framework enabled systematic cross-regional 196 comparison while capturing essential phenological stages. For temporal consistency, 197 we referenced soybean growth years by their harvest year (e.g., 2001 refers to October 2000 to March 2001). Cropland extent was based on a gridded dataset circa 2000³³. The arc of deforestation in the south of Amazonia (ARC) is identified in this study when tree cover²⁷ loss exceeds 10% over 0.5-degree grids during the study period.

202 Quantifying yield response to precipitation

Here, we developed the following state-specific linear regression models to predict soybean yields:

$$Y = \sum_{s=1}^{s=4} (\beta_{P_s} \times P_s) + \beta_{T_g} \times T_g + \beta_{year} \times year + \beta_{year^2} \times year^2$$
 (1)

where Y represents soybean yield (kg ha⁻¹), P_s is precipitation (mm) during four growth stages: ANT (s=1), VEG (s=2), REP (s=3), and MAT (s=4), and T_g is

the average growing-season (daily mean) air temperature (°C). The time term (year and year²) accounts for technological advancements in agronomic practices over time, such as improved cultivars, fertilization, and management techniques. We used growth-stage-aggregated rather than monthly precipitation to avoid overfitting while capturing the sensitivity of soybean yields to water availability. We also tested more complex formulations (including quadratic temperature terms and precipitation-temperature interactions) but adopted this simpler model structure as the alternatives showed minimal R^2 improvements while introducing potential overfitting issues. This model also allows us to more directly quantify the effects of precipitation changes on soybean yields—a critical consideration for assessing deforestation impacts on agricultural productivity through atmospheric moisture recycling.

We implemented the Least Absolute Shrinkage and Selection Operator (LASSO) regression to optimize model performance through variable selection and regularization. Model uncertainty was quantified using 1000 bootstrap iterations, with each iteration randomly allocating 2/3 of the responses-predictor pairs for calibration, and the remaining 1/3 for validation. Model performance was assessed through determination of coefficients (R^2) between observed and predicted yields for both calibration and validation datasets. The iterations also provided distributions of regression coefficients (β) and R^2 , allowing calculation of their confidence intervals (CIs). Our model exhibits strong performance, achieving the high determination coefficients in both train and test data (mean $R^2 > 0.80$ and 0.60, Supplementary Figure S2), and accurately predicting historical soybean yields ($R^2 > 0.70$ across states, Supplementary Figure S3).

Soybean yield statistics were obtained from Brazil's national agricultural database Instituto Brasileiro de Geografia e Estatística²⁴. For climate variables, we used precipitation data from the Multi-Source Weighted-Ensemble Precipitation (MSWEP) at 0.1° spatial resolution³⁹, and 2-m air temperature data from the European Center for Medium-Range Weather Forecasts ReAnalysis (ERA) 5 at 0.25° spatial resolution⁴⁰.

We calculated cropland area-weighted precipitation totals for each state across four growth stages, while temperature was averaged over the entire soybean growth period (August–March) to capture thermal conditions throughout crop development.

238 Atmospheric moisture tracking framework

In this study, we employed the Lagrangian particle dispersion model FLEXPART 239 version 10.423 driven with reanalysis at 1° resolution to track air parcels through the 240 atmosphere. The model was initialized with approximately 3 million air parcels (each 24 with mass 1.697x10¹² kg) distributed homogeneously across the globe. Each air 242 parcel carries properties including specific humidity, potential temperature and density, 243 with FLEXPART tracking their three-dimensional positions and properties over time. The outputs provided information on parcel positions (longitude, latitude, height), 245 properties (density, specific humidity), and boundary layer characteristics, enabling 246 the following construction of source-sink relationship of precipitation⁴¹. While the 247 simulations incorporated both 6-hourly reanalysis and 3-hourly forecasts to improve 248 trajectory accuracy through better representation of turbulence, our source-sink 249 relationship analysis of precipitation used only 6-hourly reanalysis outputs. 250

Next, we constructed the source-sink relationship of precipitation based on the 251 trajectories tracked, through three steps: (1) diagnosis, (2) attribution, and (3) bias 252 correction⁴². The diagnosis step evaluates all air parcels at consecutive 6-hourly time 253 steps, quantifying precipitation and evaporation to establish detection metrics. During 254 the step of attribution, we selected air parcels over target regions (that is, Brazilian 255 soybean states under study, Figure 1c) and then tracked their backward trajectories. 256 This identifies the evaporation and precipitation events through specific detection 257 criteria: evaporation occurs when the change of specific humidity is higher to 0 g kg $^{-1}$ 258 in a 6-hourly time interval, and precipitation events are identified when relative humidity 259 exceeds 80%^{42,43}. In the third step, the framework conducts bias correction using 260

observational datasets: MSWEP for precipitation and GLEAM over land⁴⁴ and OAFlux over ocean⁴⁵ for surface evaporation. Furthermore, using daily tree evaporation (that is, the sum of tree transpiration and interception loss) fraction (relative to total evaporation), we identified moisture source regions where tree evaporation contributes to precipitation, or precipitation originates from tree evaporation^{46,47}.

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We quantified tree evaporation contributions to precipitation ($E2P_{tree}$) across soy-266 bean growth stages at both state and continental scales. At the state level, we 267 calculated the relative contribution of each source region (Figure 2e) by comparing 268 absolute $E2P_{tree}$ values (Supplementary Figure S6) to total precipitation received by 269 individual states (Supplementary Figure S11). For continental-scale patterns, we 270 aggregated moisture sources across all seven study states while preserving the 271 spatial identity of source regions (Supplementary Figure S5). We then evaluated the 272 relative importance of these contributions by comparing them to the total precipitation 273 received by all states combined, conducting separate analyses for each growth stage 274 (Figure 2a-d). 275

276 Uncovering tree evaporation impacts on precipitation

Since precipitation is inherently a fraction of evaporation, we express $E2P_{tree}$ as the product of tree evaporation (E_{tree}) and a parameter (λ) that represents the transport and conversion efficiency of evaporated moisture to precipitation:

$$E2P_{tree} = \lambda \times E \tag{2}$$

To analyze long-term changes in $E2P_{tree}$, we used a linearization approach that assumes independence between E_{tree} and λ at climatological timescales (>30 years). This assumption enables us to express temporal changes in $E2P_{tree}$ by differentiating

Equation 2 with respect to time:

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$$\frac{\partial E2P_{tree}}{\partial t} = \frac{\partial E_{tree}}{\partial t} \times \lambda + \frac{\partial \lambda}{\partial t} \times E_{tree}$$
(3)

Taking time averages, Equation 3 can be rewritten as follows:

$$S_{E2P_{tree}} = S_{E_{tree}} \times \bar{\lambda} + S_{\lambda} \times \bar{E_{tree}} + \delta$$
 (4)

where S represents the slope of the linear regression (trend) for $E2P_{tree}$, E_{tree} and λ , respectively, with bars denoting climatological averages and δ being the residual of the approximation. The term $S_{E_{tree}} \times \bar{\lambda}$ specifically quantifies the impact of E_{tree} changes on precipitation.

Note that, the transformation from Equation 3 to 4 involves two simplifications suited to our climatological analysis: we approximate the temporal derivatives using linear regression slopes (e.g., $\frac{\partial E_{tree}}{\partial t}$ as $S_{E_{tree}}$), and decompose time-averaged products into products of averages (denoted by bars). However, the robustness of these approximations is empirically validated in Supplementary Figure S10, which shows that residual terms account for less than 1% of the total $E2P_{tree}$ changes (Supplementary Figure S7). This also confirm that nonlinear interactions between E_{tree} and λ changes play a minor role at the climatological scale.

Similarly, we quantified the impact of tree evaporation declines on precipitation at both state and continental scales. For individual states (Figure 3e), we calculated precipitation reductions (i.e., $S_{E_{tree}} \times \bar{\lambda}$, Supplementary Figure S9) relative to seasonal totals (Supplementary Figure S11). At the continental scale (Figure 3a–d), we aggregated these changes across all study states (Supplementary Figure S8), with analyses conducted separately for each soybean growth stage.

Quantifying tree evaporation impacts on soybean yields

To quantify how tree evaporation changes affect soybean productivity, we combined 307 crop-climate sensitivity analysis with precipitation changes driven by tree evaporation declines. We estimated soybean yield sensitivity to precipitation (β , kg ha⁻¹ mm⁻¹) 309 across growth stages using LASSO regression (see above and Figure 1d), and 310 then quantified how historical tree evaporation declines altered precipitation patterns 311 (Supplementary Figure S9). For each growth stage, yield impacts were calculated by 312 combining precipitation reductions with corresponding yield sensitivities. For example, 313 with a yield sensitivity of 2 kg ha^{-1} mm⁻¹ to precipitation, a 10 mm reduction in 314 precipitation due to decreased tree evaporation would result in a yield loss of 20 kg 315 ha⁻¹. We included only positive sensitivities exceeding 0.10 kg ha⁻¹ mm⁻¹, as weak or negative responses typically indicate energy-limited conditions where increased 317 precipitation reduces solar radiation^{5,48}. Total state-level production losses were 318 estimated by multiplying yield reductions with harvested areas circa 2000³³. 319

Data availability

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Soybean yield statistics were obtained from https://www.ibge.gov.br. The harvest area is available from https://www.dante-project.org/datasets/mirca2K. MEaSUREs Vegetation Continuous Fields is acquired from https://lpdaac.usgs.gov/products/vcf5kyrv001. GLEAM data are available through https://www.gleam.eu. OAFlux data can be retrieved from https://oaflux.whoi.edu/data-access. ERA data were accessed from https://cds.climate.copernicus.eu/cdsapp#!/dataset. MSWEP data are available through http://www.gloh2o.org.
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329 Code availability

The FLEXPART model can be downloaded via https://www.flexpart.eu. The version of the moisture tracking framework used for analysis of FLEXPART data is preserved at https://doi.org/10.5281/zenodo.5788506. Python scripts for the analysis are available upon request from Dr. Hao Li.

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Author contributions

H.L. conceived the study, designed the experiments, and conducted the analysis.
H.L., and C.L., contributed to data analysis and interpretation. All authors contributed to writing the original draft and participated in several rounds of revisions.

346 Competing interests

The authors declare no competing interests.

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