

Geotechnical objectives

soil Mechanics
4 marks

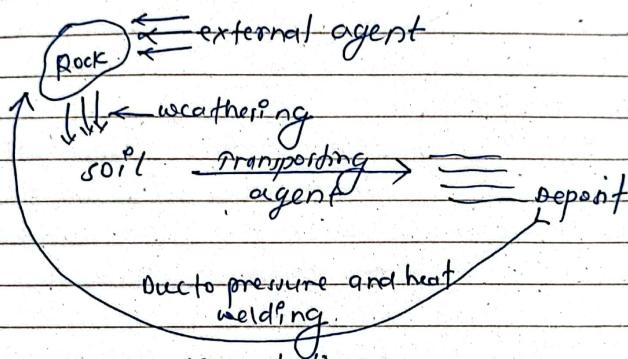
Geology + tunnelling:
4 marks

Chapter: 1

General classification of soil depending on formation, transporting agent and deposit media.

soil: Accumulation of particles formed by disintegration of rocks.

- uncemented or partially cemented: soil.
- fully cemented: rock.



soil formation: weathering

Physical weathering

4 Rock only converts into small particles but minerals remain same as parent rock.

Forms: gravel and sand.

Chemical weathering

4 Rock not only changes into smaller particles but mineral also changes.

Forms: silt and clay.

Physical weathering

causes:

- 4 Abrasion
- 4 wedge action of ice
- 4 spreading of roots of tree
- 4 Heat: regular expansion & contraction

causes:

- 4 chemical reaction
- 4 hydration, oxidation
- 4 carbonation, solution, hydrolysis.

outcrop ← external agent

Residual soil:- soil formed and placed at same place of parent rock.

Colluvial deposit: pulled by gravity and placed at base of slope.

Transporting agent:-

water:

Deposit media:	Name of soil:	Remarks:
River basin	Alluvial soil	Terai
Lake basin	Lacustrine soil	Kathmandu
Sea basin	Marine clay	Finegrained

Air

Desert area

Aeolian soil

Glacier

Glacier deposit
Base of mountain

Glacier deposit

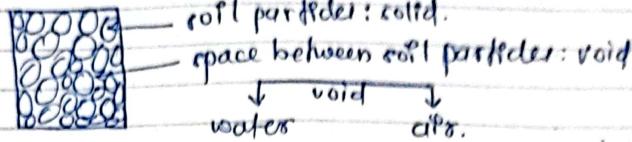
Sea soil: finest soil, sensitivity ↑, shear strength ↓.

Outcrop: Rock exposed to atmosphere.

more than 150% water content
 lacustrine : moist rich : highly compressible, more settlement
 \hookrightarrow pile or pier foundation suitable

Alluvial soft - well foundation because A sandy soil.
 Residual soft: hilly area; shallow foundation.

Chapters 2 Three phases of soil:



In general condition: soil mass consists of solid, water and air. three phase material.

Two extreme condition:

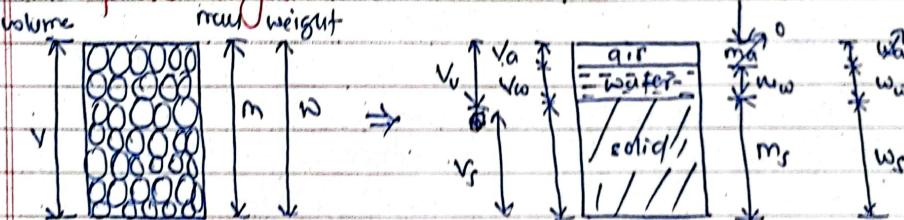
- Two { . Dry condition: Due to regular contact with heat (solid + air)
 phase } . saturated condition: Due to regular contact with water (solid + water)

Theoretically:

One phase:- fully compacted or fully consolidated or zero void.

But practically not possible.

Three phase diagram:



Basic terms:-

i) Void ratio (e):

$$e = \frac{V_v}{V_s}$$

Relation

$$n = \frac{e}{1+e}$$

iii) Porosity (n)

$$n = \frac{V_v}{V}$$

iv) % air voids (N_a)

$$N_a = \frac{V_a}{V} \times 100\%$$

$$N_a = \frac{V_a}{V} \times 100\%$$

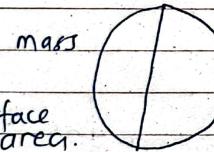
v) Moisture content (ω)

$$\omega = \frac{m_w}{m_s} \times 100\%$$

Note: The more the finer, more is the water content.

When a particle is divided, surface area increases, mass remains constant.

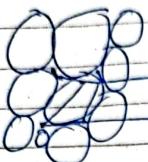
more surface area, more water required to wet.



Specific surface area: $\frac{\text{area}}{\text{mass}}$

$\text{mm}^2/\text{gm.}$

surface attraction (cohesion) is more in finer particles due to negligible gravity and reverse in coarser particles.



void size ($\overline{d^3}$)

No. of voids (less)

coarse particles



void size ($\overline{d^3}$)

No. of voids (more)

finer particles

∴ Overall ~~in~~ finer particles ~~have~~ voids ratio ~~are~~ ~~less~~ in comparison to coarser particles.

(vi) Degree of saturation:

$$S_r = \frac{V_w}{V_v} \times 100\%$$

$$S_r \propto w \therefore S_r \times e = w \times g$$

$S_r = 0$, dry soil.

$S_r = 1$, saturated soil.

$0 < S_r < 1$, general soil.

(vii) Specific gravity of solid (G):

Ratio of mass of solid to mass of water same volume of water at $4^\circ C$.

$$G = \frac{\rho_s}{\rho_w} = \frac{\gamma_s}{\gamma_w} \text{ at } 4^\circ C$$

generally for solid particles ($G=2.7$)

Note:

M, V, W, w and G can be directly measured and other relations are found from the analysis of relationship between them.

for distilled water:

$$\rho_w = 1 \text{ gm/cc}$$

$$= 1 \text{ gm/ml}$$

$$= 1000 \text{ kg/m}^3$$

$$\gamma_w = 9.81 \text{ KN/m}^3$$

(i) General condn.

(ii) Dry condn.

(iii) Saturated condn.

(iv) Submersion condn.

Volumetric Relationship

Soil condⁿ V-r-m Relationship

i) General condⁿ

$$\text{Bulk density } (\gamma) = \frac{(Sr \times e + G) \rho_w}{1+e}$$

ii) Dry condⁿ

$$\text{Dry density } (\gamma_{dry}) = \frac{M_{dry}}{V} = \frac{M_s}{V}$$

$$\therefore \gamma_{dry} = \frac{G \rho_w}{1+e}$$

iii) saturated state condⁿ

$$(\gamma_{sat}) = \frac{M_{sat}}{V}$$

$$\therefore \gamma_{sat} = \left(\frac{e+G}{1+e} \right) \rho_w$$

iv) submerged condⁿ

$$(\gamma_{sub}) \text{ or } \gamma' = \frac{M_{sub}}{V}$$

$$\gamma_{sub} = \gamma_{sat} - \rho_w$$

$$\gamma_{sub} = \left(\frac{G-1}{1+e} \right) \rho_w.$$

V-r-w Relationship:

$$\text{Bulk unit weight } (\gamma) = \frac{w}{V}$$

$$\gamma = \frac{(Sr \times e + G) \rho_w}{1+e}$$

$$\text{Dry unit weight } (\gamma_{dry})$$

$$= \frac{w_{dry}}{V} = \frac{w_s}{V}$$

$$\therefore \gamma_{dry} = \frac{G \rho_w}{1+e}$$

$$\text{saturated unit weight } (\gamma_{sat}) = \frac{w_{sat}}{V}$$

$$\gamma_{sat} = \frac{(e+G) \rho_w}{1+e}$$

$$\text{submerged unit weight } (\gamma_{sub} \text{ m' })$$

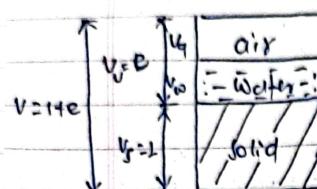
$$\gamma_{sub} = \gamma_{sat} - \rho_w = \frac{w_{sub}}{V}$$

$$\gamma_{sub} = \left(\frac{G-1}{1+e} \right) \rho_w$$

Q. $\gamma_{sat} > \gamma > \gamma_{dry} > \gamma_{sub}$

for same soil when its texture changes.

Three phase diagram:-
i) in terms of void ratio



$$e = \frac{V_v}{V_s}$$

$$V_w = Sr \times e$$

$$\text{if } V_s = 1$$

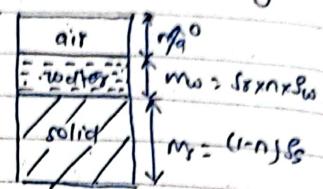
$$\therefore V_v = e \quad V_a = (1-Sr) \times e$$

$$Sr = \frac{V_w}{V_v} \Rightarrow V_w = Sr \times e$$

Q. Volumetric strain = $\frac{\partial e}{V}$

Q. Relation of volumetric strain β :
 $\frac{\partial e}{1+e}$

Three phase diagram:
ii) In terms of porosity.



$$\eta = \frac{V_v}{V}$$

$$V_v = n \quad \therefore V = 1$$

$$V_w = Sr \times \eta$$

$$V_a = (1-Sr) \eta = \frac{1}{1+n} = (1-n)$$

$$\gamma = [Sr \eta + (1-n)G] \rho_w$$

$$\gamma_{dry} = (1-n)G \rho_w = \frac{G \rho_w}{1+n}$$

$$\gamma_{sat} = [(n+1-n)G] \rho_w$$

$$\gamma_{sub} = [(n-1) + (1-n)G] \rho_w$$

Analysis of soil:-

$\rho_s, \omega, G \rightarrow$ measured.

↓

$$S_{dry} = \frac{\rho_s}{1+\omega}$$

↓

$$S_{dry} = \frac{(1-\omega)\rho_w}{1+\omega}, S_{dry} = \frac{G \cdot \rho_w}{1+\omega} \rightarrow S_{dry} = \omega \cdot G.$$

↓

$$\epsilon \rightarrow S_{sat} = \frac{(\epsilon + G) \cdot \rho_w}{1+\epsilon}$$

↓

$$n = \frac{\epsilon}{1+\epsilon} \rightarrow S_{sat} = S_{dry} + \rho_w$$

↓

$$\eta_a = n \cdot \rho_w \rightarrow n$$

↓

$$\rho_{sat}$$

↓

Determination of field density of soils-(e):

(a) core cutter method] Majority used.

(b) sand replacement method

(c) water displacement method.

(Imp) a. How will you measure field density of soil by using

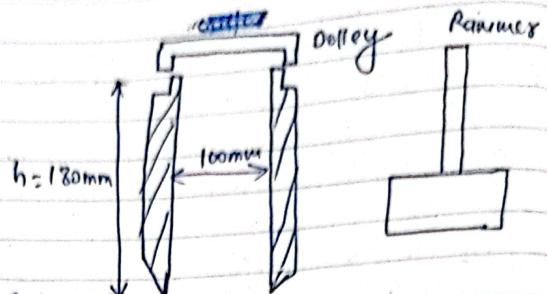
(i) core cutter method [5 marks]

(ii) sand replacement method [5 marks]

a. Core cutter method:

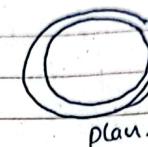
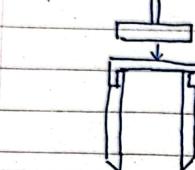
Apparatus:

- (i) core cutter
- (ii) Rammer.



Procedure:

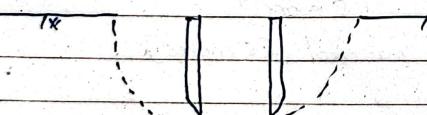
(i) Drive the core cutter full depth into soil in field.



plan.

|| | position of core cutter after driven.

(ii) Take out the driven core cutter by excavation of surrounding soil.



Measure

(i) mass of empty core cutter = m_1 .

(ii) mass of core cutter with full soil = m_2 .

$$\text{(iii) field density of soil } (\delta) = \frac{m}{v}$$

$$= \frac{m_2 - m_1}{v}$$

$$; v = \frac{\pi d^2 h}{4}$$

Limitation:

↳ not used for cohesive soil.

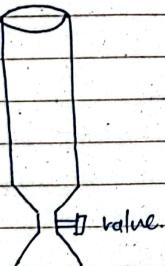
Use:

↳ for cohesive soil only.

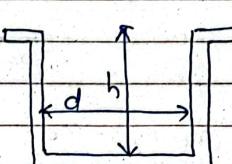
(b) Sand Replacement Method:

Apparatus

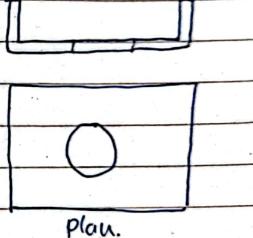
(i) cone apparatus



(ii) metallic pot



(iii) metallic pan



⑥ uniform sand:

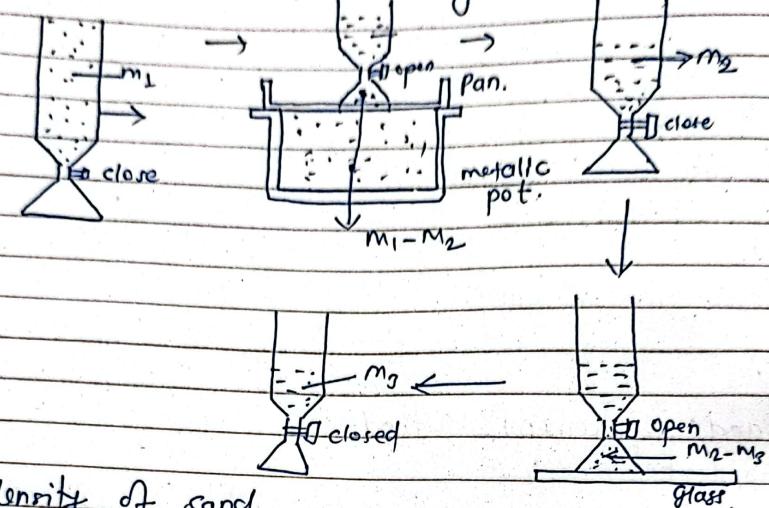
0.45mm - 0.6mm

↑
passing
retained.

Procedure:

a. Calibration of sand:

(i) determination of sand density (δ_{sand}):



density of sand

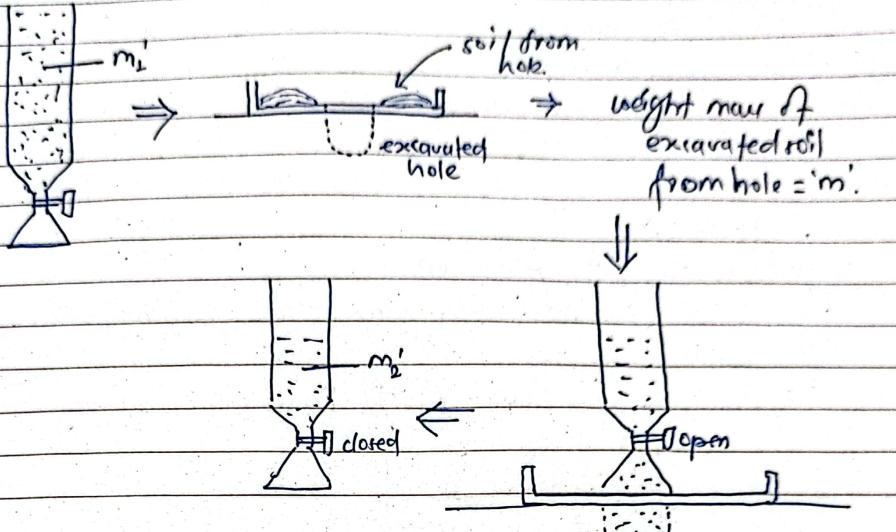
$$\delta_{sand} = \frac{\text{mass of sand to fill pot}}{\text{volume of sand in pot}}$$

$$\delta_{sand} = \frac{(m_1 - m_2) - m_c}{V} ; m_c = m_2 - m_3$$

$$V = \pi d^2 h ; h = \text{height of}$$

m_c = mass of sand to
fill cone section.

(a) Measurement of field density: (8):



$$\text{density of soil } (\rho) = \frac{\text{mass of soil from hole } (m)}{\text{volume of hole } (V_d)}$$

$$; V_d = \frac{\text{mass of sand to fill hole}}{\text{density of sand.}}$$

$$= \frac{(m_1 - m_2) - m_e}{\text{density of sand. } (D_{\text{sand}})}$$

use: for all types of soil.

(c) water displacement method:

use:

↳ for clay.



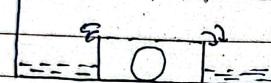
undisturbed soil
piece from field

Procedure:

↳ wet the soil piece (m)

↳ coat the soil with wax. and weight the coated soil (m').

↳ Immmerse the coated soil in water.



↳ measure the displaced water during immersion of coated soil piece into water (V_t).

$$V_t = \text{vol. of } (\text{soil} + \text{wax})$$

$$\text{vol. of soil } (V) = V_t - \text{vol. of wax}$$

$$= V_t - \frac{(m' - m)}{S_{\text{wax}}}$$

$$; S_{\text{wax}} = 0.89 \text{ g/cc.}$$

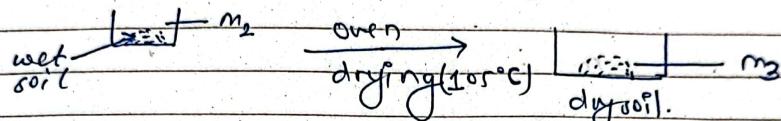
Determination of moisture content: (ω):

- (1) oven dry method → (most accurate method (mostly used))
- (2) pycnometer method
- (3) Radiation method
- (4) calcium carbide method → (fastest method (psc))
within 5 minutes → result.

i. Oven drying method

step:

- (i) wt. empty pan (m_1) :→ (silver pan) $\rightarrow m_1$
- (ii) put little ~~some~~ wet soil in pan & weight (m_2)



- (iii) keep wet soil with pan in oven overnight at temperature 105°C .

- (iv) weight pan with dry soil. (m_3)

$$\text{moisture content } (\omega) = \frac{m_2}{m_3} = \left(\frac{m_2 - m_3}{m_3 - m_1} \right) \times 100\%.$$

$$= \frac{\text{mass of water}}{\text{mass of dry soil.}} \times 100\%.$$

Sp. gravity test:

- (1) Pycnometer method:
- (2) Binnity bottle → more accurate than pycnometer.
- (3) shrinkage limit method. → time consuming.

(ii) Pycnometer method:

Procedure

- 1) mass of empty pycnometer.

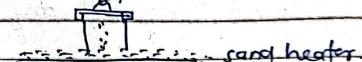


- 2) put some dry soil & weight



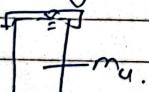
dry soil.

- 3) put water & make expell of entrapped air pocket from soil & weight.



for fast

- 4) clean the pycnometer, put full volume of water & weight



Derivation:

$$m_4 = m_3 - m_2 + V_s \times S_w$$

$$m_4 = m_3 - m_2 + \frac{m_2}{S_s} \times S_w$$

$$m_4 = m_3 - m_2 + \frac{m_2}{G}$$

Q. What are the uses of index properties of soil? (5 marks)

$$\frac{m_s}{G} = \frac{(m_u - m_d) + m_f}{m_u - m_d}$$

$$G = \frac{m_s}{(m_u - m_d) + (m_d - m_f)}$$

$$G = \frac{m_s - m_f}{(m_u - m_d) + (m_d - m_f)}$$

$$; m_s = m_d - m_f$$

Use: mainly for coarse grained soil.

Properties of soil:



- Properties of soil which are not used directly in design of any structure on soil.

Use:

- To classify the soil.
- LL, PL, grain size distribution curve, sensitivity, thixotropy.
- To develop correlation with engineering properties.
e.g. $c_c = 0.009 (LL-LO)$

$$K = 100 D_{10}^2 \text{ Hazen's eqn. i.e. } K \propto D_{10}^2 \text{ sq. of particle size}$$

objective question.

• for research on soil properties.

↓
Aggregate properties
coarse soil ↓ fine soil
Relative density (D_r)

↓ Individual grain properties.

↓
Grain size distribution ↓ shape of soil particles.

↓
consistency limits ↓
LL PL SL

↓ Activity

• Relative density (D_r): (For coarse soil):

• It indicates denseness of soil (sand) in field.

$$D_r = \left(\frac{e_{max} - e}{e_{max} - e_{min}} \right) \times 100 \%$$

$$; e_{max} = \text{max}^{\text{rd}} \text{ void ratio at very loose condition.} = \left(\frac{e_{pw}}{S_{dry(max)}} - 1 \right)$$

$$e_{min} = \text{min}^{\text{rd}} \text{ void ratio at very dense condition.} = \left(\frac{e_{pw}}{S_{dry(min)}} - 1 \right)$$

$$e = \text{natural void ratio of soil in field} \\ = \left(\frac{e_{pw}}{S_{dry}} - 1 \right)$$

↳ Montmorillonite: plastic mineral \rightarrow swelling volume change & water!

Centonite slurry: montmorillonite plastic clay \rightarrow water slurry occurs.
+ stability \rightarrow use often \rightarrow shrinkage & swelling potential high!

D _{r.}	> 85%	65-85%		85-65%	15-35%	< 15%	
state of soil:	Very dense cond ⁿ	Stiff cond ⁿ	—	Medium dense	Loose	Very loose	
consistency	Very good	Good		OK	• Not constructable ***. \rightarrow Recommend for deep compaction.		

consistency limit: (for fine soil)

↳ possibility of high variation of moisture in fine soil:
 \downarrow effect

(A) Volume instability

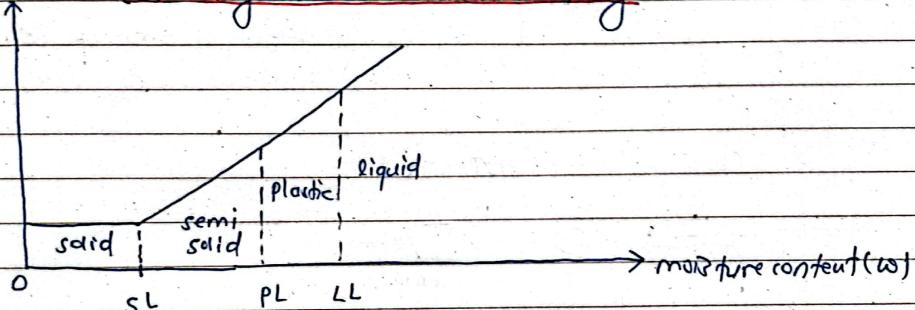
swelling: increase in volume of soil mass with increase of moisture.

shrinkage: decrease in volume of soil mass with decrease of moisture.

Bulking:
Increase in volume of particles.
Especially in sand.

Q. Increase in volume of soil mass due to increase in moisture content R.
 (A) swelling (B) bulking (C) both (D) none.

(B) Change in behaviour: consistency.



Q. Ceramic product \rightarrow clay etc (plasticity high \rightarrow consistency)

Q. The moisture content below which soil mass does not show reduction in volume. R. -

(A) LL (B) PL (C) RL (D) All

Fluidity \leftarrow opposite \rightarrow Rigidity (w \uparrow)

consistency limit:

↳ limit of moisture at which clay shows change in behavior

a. LL
b. PL
c. SI

A Herberg limit.

*: Determination of liquid limit (LL):

↳ Casagrande method:

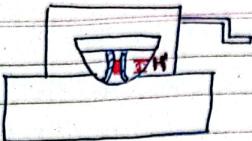
Casagrande's Apparatus:



h = height of drop of Broon2 cup
= 10 mm. (objective question)

cutting tool: ASTM tool.





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$H' = 10\text{mm}$ (if cm^2) portion.

Measurement:

Test No:	No. of blows	Moisture
I	N_1	w_1
II	N_2	w_2
III	N_3	w_3

$N \neq 50$ in any case

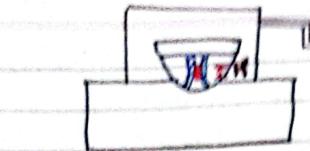
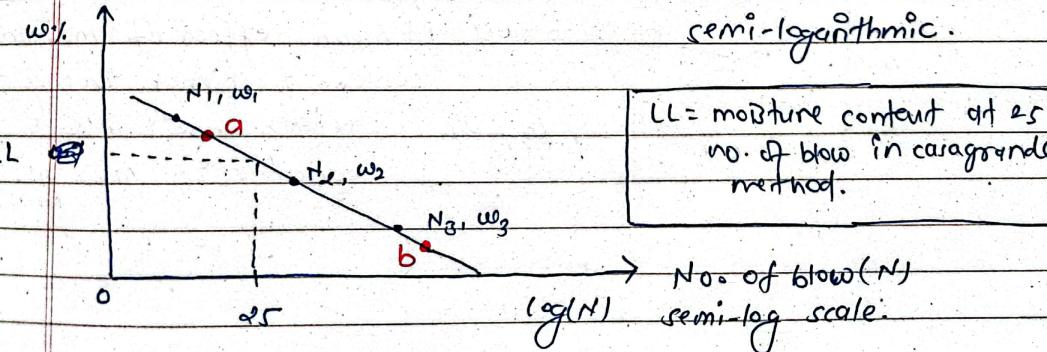
Among 3 values any one N should be greater than or \neq any one N should be less than or \neq .

- a. In casagrande's test for LL, No. of blows should be:

Ans: $N \neq 50$
any one $N < 25$
any. one $N > 25$

} All three are roundabouts.

Result:



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$H' = 10\text{mm}$ (if cm^2) portion.

Measurement:

Test No:	No. of blows	Moisture
I	N_1	w_1
II	N_2	w_2
III	N_3	w_3

$N \neq 50$ in any case

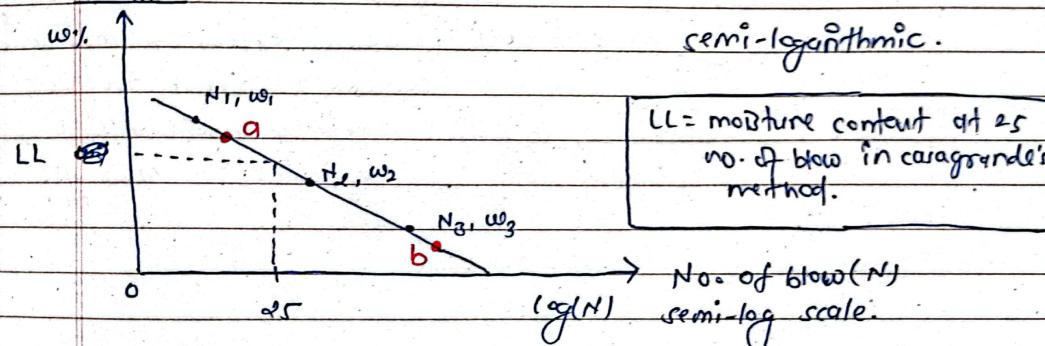
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- a. In casagrande's test for LL, No. of blows should be:

Ans: $N \neq 50$
any one $N < 25$
any. one $N > 25$

} All three are roundabouts.

Result:



for fine ab.

$$\text{slope of ab: flow index. } (f_i) = \frac{w_a - w_b}{\log \left(\frac{N_a}{N_b} \right)}$$

Higher the flow index, it means,

shear strength variation is high

" " " " B 100.

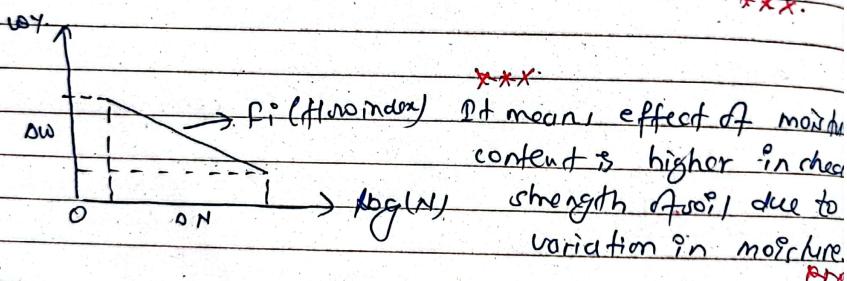
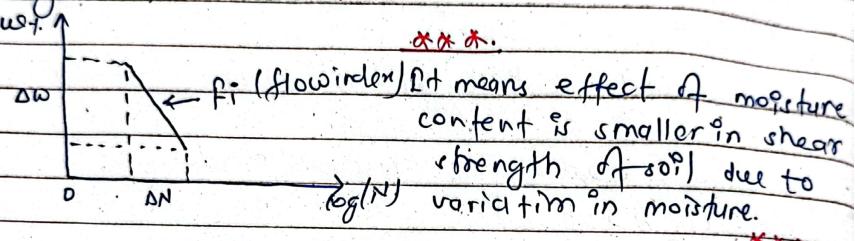
both. (d) None.

lower the flow index, it means,

shear strength variation is low

shear strength variation is high. (c) both. (d) None.

Higher flow index:



(e) Which is not true ...

- (a) Higher flow index, effect of moisture content is smaller in shear strength.
- (b) Lower flow index, effect of moisture content is higher in shear strength.
- (c) Higher flow index, effect of moisture content is higher in shear strength.
- (d) All of above.

a) Determination of plastic limits:-

4. Hit and trial method:-

- Rolling soil on glass below palm changing moisture content.

PL: The moisture at which clay can be rolled to 3mm diameter thread.

some index:

i. Plasticity Index (P_i)

$$P_i = L_L - PL$$

: physical meaning:

4 Range of moisture difference for plastic behavior.

{ for gravel & sand (P_i) = null (0)
for silt = 0 to some limited value
for clay = high. }

ii. liquidity index (L_i)

$$L_i = \frac{w_N - PL}{P_i}$$

; w_N = natural moisture content.

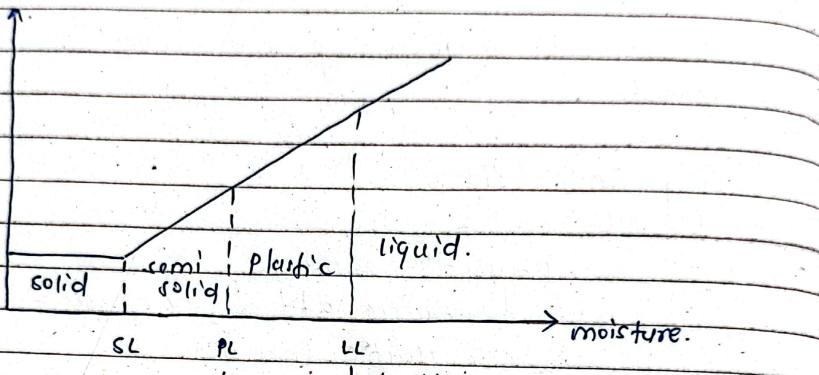
3. consistency index: (C_i)

$$C_i = \frac{LL - w_N}{PL}$$

4. Flow index Toughness index: T_i

$$T_i = \frac{P_i^o}{P_i} \quad (\text{Range: 0 to } 3)$$

P_i^o = flow index obtained from ejection test for liquid limit.



$$Li = \frac{w_N - PL}{PL}$$

$$C_i = \frac{LL - w_N}{PL}$$

$$C_i = \frac{LL - w_N}{PL}$$

*** Note:

$$C_i + Li = 1$$

Always holds true!!!

Sensitivity:

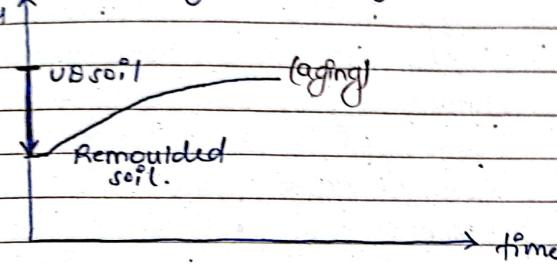
↳ Fine soil especially clay show reduction of strength after remoulding even at same density and moisture, this property is known as sensitivity.
 $S_u = \frac{S_u(\text{soil})}{S_u(\text{remoulded soil})}$

i) Vane shear test: to measure S_u .

for Kmm:
 $S_u = 5$
 $S_u = 4$

Thixotropy:

↳ Remoulded clay starts to regain its lost strength with time/aging



↑ Edge to edge arrangement
 H face to face arrangement:

Remoulding



dispersed structure.

Aging

face to face parallel arrangement

Thixotropy:

0

Parameter		composition	
		Flocculated structure	dispersed structure.
shear strength		Higher	lower.
compressibility		lower	Higher
permeability		Higher	lower
swelling / shrinkage		Higher	-
			Higher.

* Activity :-

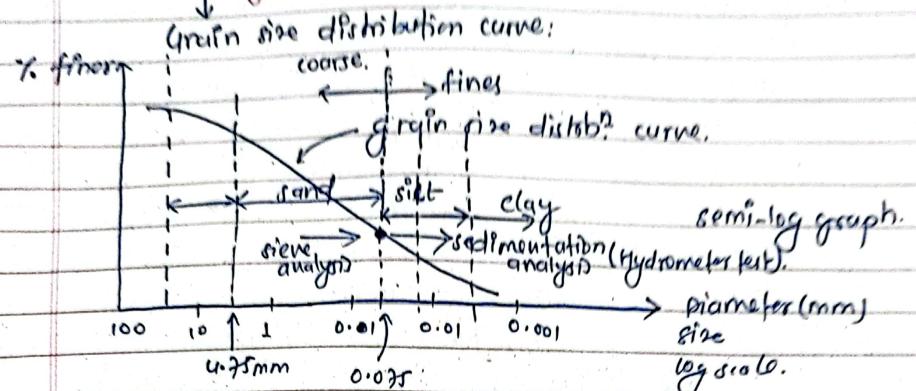
$$A = \rho_i$$

% of clay fraction

; clay fraction: finer
than 2d. size
(0.002mm)

Activity	status of clay	
> 1.25	very active	
0.75 < A ≤ 1.25	moderate active.	
< 0.75	inactive	

Individual grain properties:

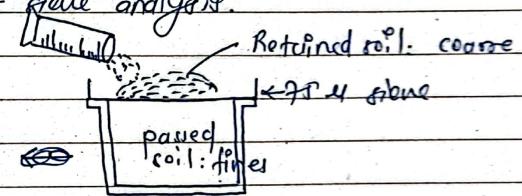


80 - 300 mm : boulders

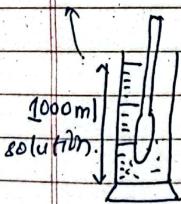
> 300mm : cobbles

if soil mass contains $\geq 5\%$ fines \Rightarrow sieve analysis (dry sieve)

if soil mass contains $\geq 5\%$ fines \Rightarrow fines should be separated \rightarrow wet sieve analysis.



% finer size of particles



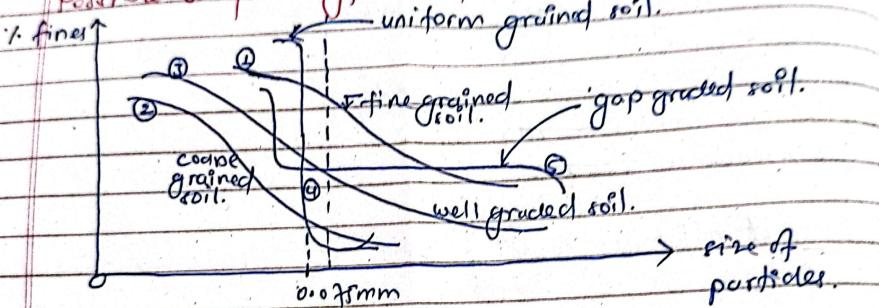
5g dry soil + 40g dispersing agent (sodium hexa metaphosphate) $\times \times \times$

coarse fine set
80 mm
40 mm
20 mm
10 mm
4.9 mm

the steveset:

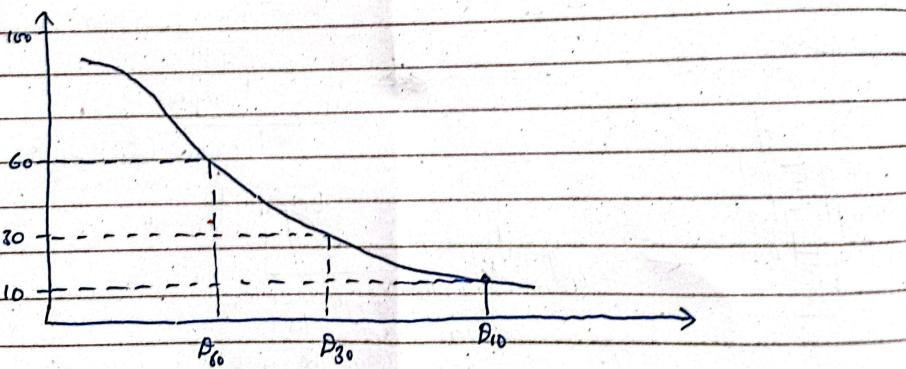
4. 35 mm
0.86 mm
1.18 mm
600 M
425 -
3 db
180°
75°

Possible shape / types of curve:



From grain size distrⁿ curve:

4 Two Coefficients



decrease
more
less

Q. Which water reduces permeability in soil?
→ Adsorbed water.

~~(5)~~ adsorbed water

• Gold area
decreases

Coefficient of uniformity (C_u)

$$C_4 = \frac{D_{60}}{D_{10}}$$

coefficient of curvature (C_c)

$$C_c = \frac{D_{30}^2}{D_{60} \times A_{10}}$$

~~D₁₀~~ = $D_{10} \times A_{10}$ $1 < C_c < 3$ for well graded.

Q. $D_{10} = 0.325$ mm means $\rightarrow 0.825$ mm ~~size~~ ^{10%} particle
 $A_{10} = 10\%$

$D_{60} = \text{size of soil particles}$ coarser soil \rightarrow avg. size $10\text{-}15\text{ mm}$

$D_{10} = \text{size of soil particles}$ fine soil \rightarrow avg. size $1\text{-}2\text{ mm}$

$D_{20} = \text{size of soil particles}$ point of curvature from D_{10} to D_{60}

type of water in miles

readily:

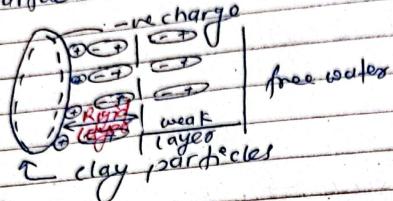
Free water.

- ↳ flow by gravity
- ↳ accumulation of free water : GW

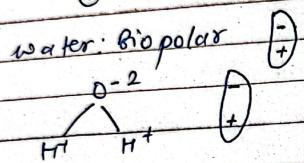
- Adsorbed water
 - ↳ in surface of soil particles
 - ↳ hygroscopic water
- Absorbed water
 - ↳ more water.

Pore water	Hypocropic water	structural water	capillary rise
in void	at interface	H ₂ O component	in void
removed by heat	of soil particles	in mineral of clay	removed by heat
eg: adsorbed water.	removed within oven but taken immediately outside orch from atmosphere.	by heating by heat.	can be removed by breaking structure.
	eg: adsorbed water		

- Ad sorbed water
In surface of soil particles
Hydroscopic water
Diffuse double layer.



cation:



Ad sorbed water:

- In void (void ~~not~~ ~~not~~ air) water.
pore water.

Structure and fabric of soil:-

soil particles arrangement ↓ soil particles arrangement with interparticle forces.

Types of structure:

(A) For fine soil:-

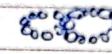
clay:

flaccid structure

disperse structure

silt:

Honeycomb structure



sensitive for dynamic load.

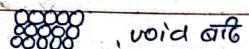
Fine sand:

Honeycomb.

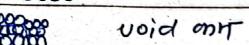
(B) For coarse soil:

sand, gravel

loose structure.



Dense structure.

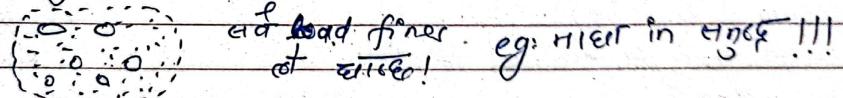


void air

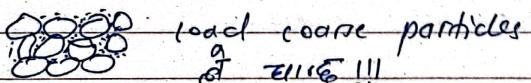
void soil

Mixed soil structure:

(i) clay matrix structure.



(ii) coarse skeleton structure.



Geology (1 mark)

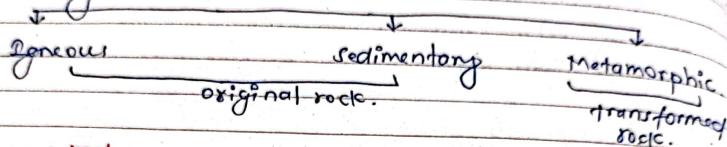
Petrology: study of formation & types & occurrence of rock.

Magma: molten material at crust of earth.

Lava: magma erupted to surface of earth.

Types of Rock:

Depending on formation:



Igneous rock:

↳ formed by cooling and solidification of magma.

- ↳ extrusive rock
- ↳ intrusive rock
- ↳ volcanic rock
- ↳ cooling at large on ground surface.
- ↳ very fast cooling rate (within 1 month)
- ↳ intermediate rate of cooling
- ↳ fine grain size / micrscopic eg: basalt, Rhyolite, Dacite, Trachyte, Grained of soil:

Peridotites

< 1mm	fine
1-5 mm	medium
> 5 mm	coarse.

Depend

↓

concor

↳ Magm

parall

A o

Type

↳ Sill

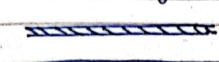
Depending on occurrence:

↳ concordant bodies

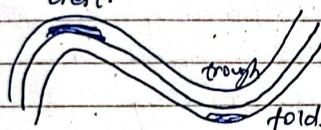
↳ Magma is cooled along or parallel to structural features of old rock layer.

Types:

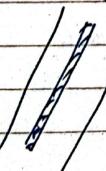
1. Sills: In between two H₂ sedimentary rock layer.



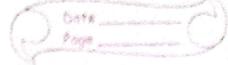
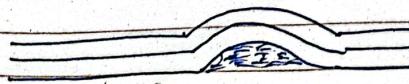
2. Phacoliths: magma cooled & placed at crest & trough of fold.



3. Iopoliths: magma cooled in betw two inclined rock layer (center to earth).



4. Laccoliths: magma cooled in arched up layer.



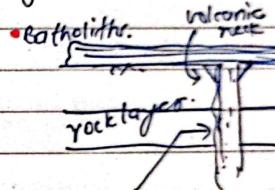
↳ discordant bodies.

↳ does not follow structural feature
↳ magma cooled transverse to structural features.

Types:

1. batholiths:

↳ large body formed by volcanic eruption.
eg: costarica.

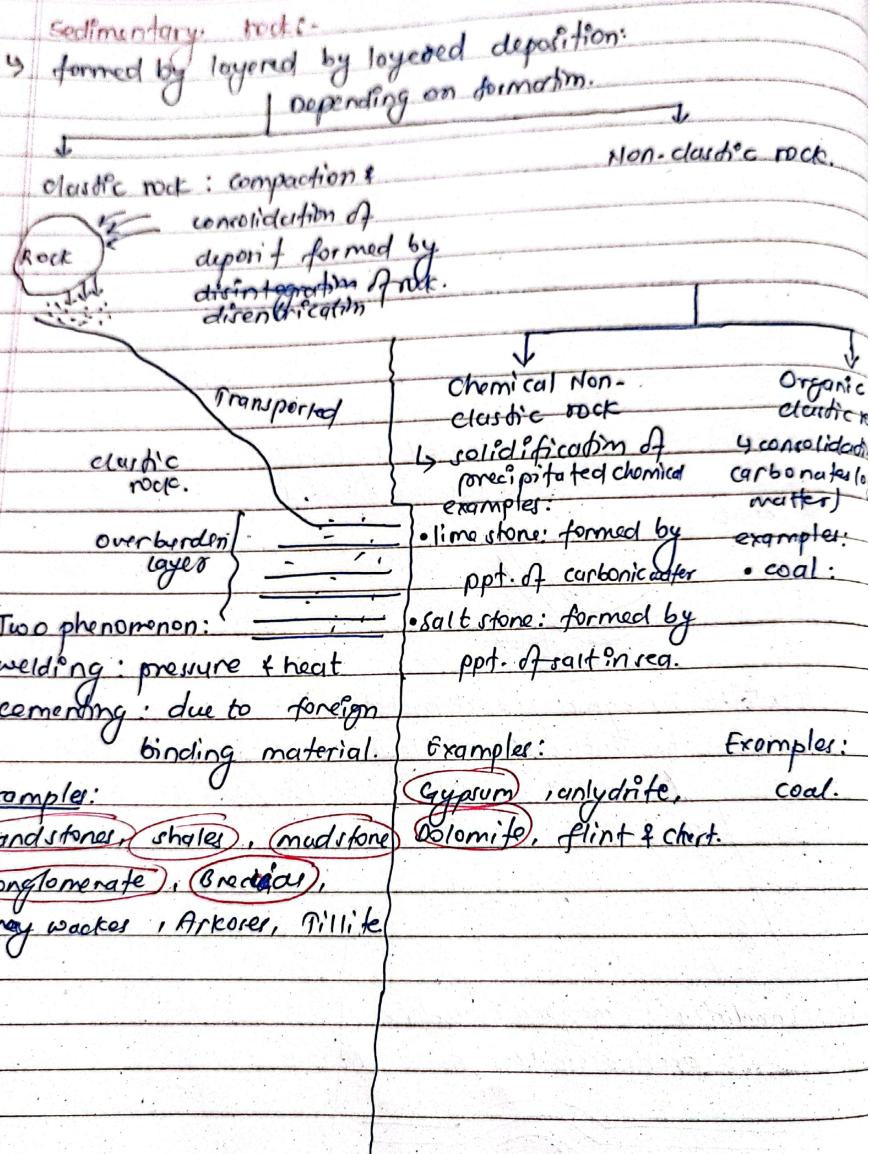


columnar body:
Dikes / dykes

- i) dykes / dykes

ii) volcanic rock.

Batholiths → only extrusive body



Metamorphic rock:

Metamorphism $\xrightarrow{\text{Result}}$ Transformed rock \rightarrow Metamorphic rock.
 \downarrow
 process
Causes:

Heat: Thermal metamorphism
 Pressure: Dynamic " "
 Chemical environment: Chemical "
 Heat + Pressure: Dynamic thermal "

Also called cu "Regional metamorphism" \rightarrow large area HT & SS!

Depending on arrangement of minerals:

\downarrow
 Foliated Rock.

\hookrightarrow Parallel arrangement of minerals

e.g. **Slate**, **Phyllite**, **schist**, **gneiss**

\hookrightarrow Non-foliated Rock.

\hookrightarrow Non-parallel arrangement of minerals.

e.g. **Marble**, **Quartzite**,
 amphibolite, soapstone,
 Hornfels, soapstone.

shales $\xrightarrow{\text{low grade metamorphism}}$ slates $\xrightarrow{\text{dynamic thermal metamorphism}}$ phyllite
 schists. \leftarrow Further metamorphism

shales $\xrightarrow{\text{dynamic thermal metamorphism}}$ phyllite.

Tibetan plate - Indian plate ~~mount~~ during ~~since~~ ~~since~~

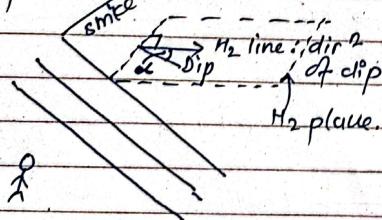
Structural features of Rock:-

- ↓ Primary structures
 - ↳ structural features formed during formation of rock.
 - ↳ mostly in igneous & sedimentary rock.
- ↓ Secondary structures
 - ↳ modified structural features due to change in new to new shape & rearrangement of minerals & grains.
 - ↳ cause: stress & heat.
✓ due to tectonic movement.
 - (i) Tectonic movement
 - (ii) Magmatic eruptions, entrainment
 - (iii) Stress due to huge ^{sediment} deposit.

Structural features:

1. Dip & strike
2. Fold
3. Fault
4. Joint
5. Shear zone
6. Rock cleavage:

Dip and strike:



Strike: Geographical extension of rock layer.

↳ line formed by any assumed H₂ plane & rock layer.

Types of Dip:

(a) True dip (α)

↳ Inclination of rock layer with H₂ line \perp to strike.

(b) Apparent dip (β)

↳ Inclination of rock layer measured with other line beside than H₂ line \perp to strike.

Relation:

$$\tan \beta = \tan \alpha \times \cos \gamma$$

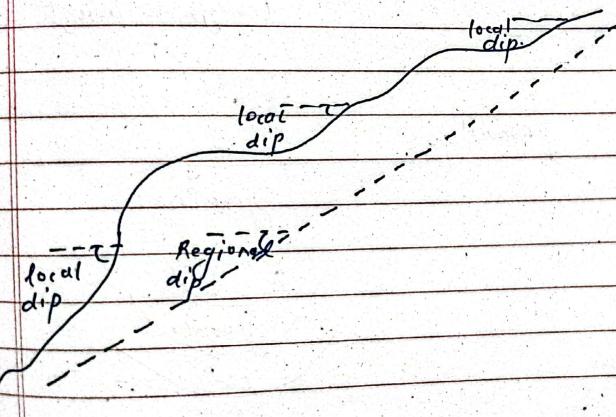
• Primary dip:

Inclination formed during formation of rock. (0° to 20°)

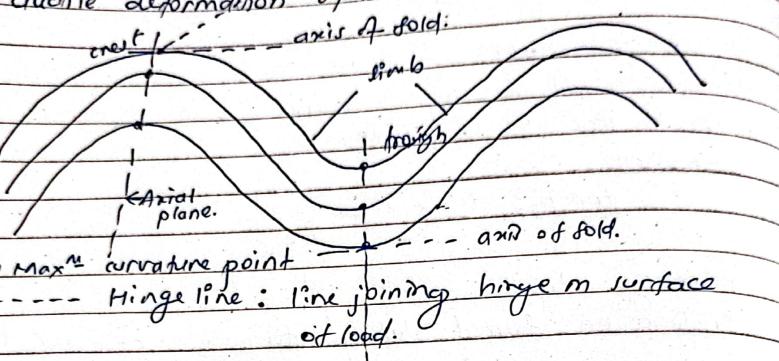
• Secondary dip:

↳ Modified dip due to stress.

↳ upto 90° ($> 40^\circ$ & $\leq 90^\circ$)



Fold:
↳ Undulation, curvature, bend, of rock formed due to ductile deformation of rock.

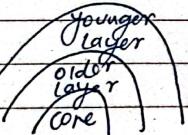


- Hinge: Maxⁿ curvature point
- Hinge line: Line joining hinge in surface of fold.
- Fold on⁹⁰ crest rift - tensile stress
- Fold on⁹⁰ trough rift - compressive stress.

One cosine of fold: distance between crest to crest
" " : " " trough to trough.

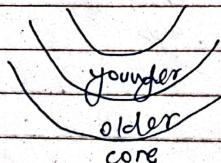
Types of fold:

↓
Anticline



- ↳ convex upward
- ↳ arched up.

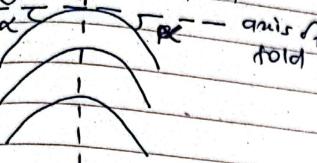
↓
syncline:



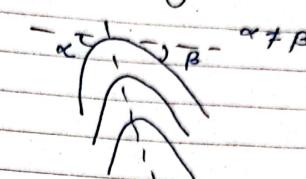
- ↳ concave upward.
- ↳ arched down.

other types:

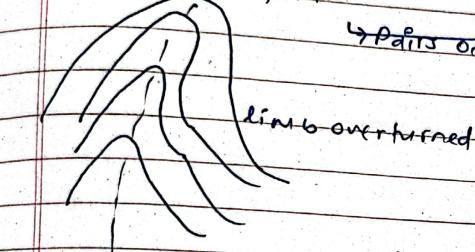
(i) symmetric fold



(ii) Asymmetric fold.



(iii) overturned fold.

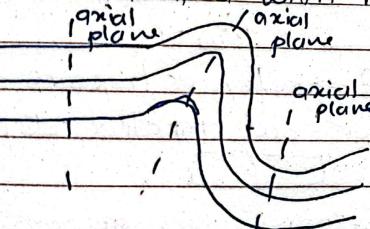


↳ Pairs of fold with non-parallel axial plane.

limbs overturned.

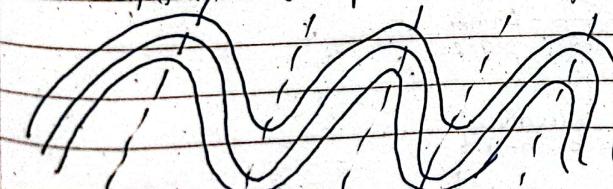
(iv) conjugate fold

↳ Pairs of fold with non-parallel axial plane.

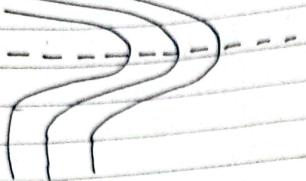


(v) Proclinical fold.

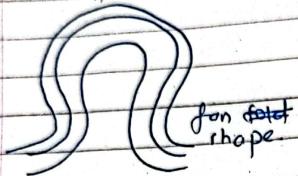
↳ Series of fold with parallel axial plane



(vi) Recumbent fold:
↳ flow with H₂. axial plane.



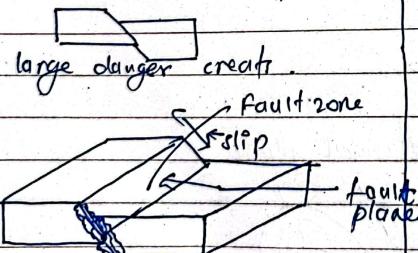
(vii) Fan fold:



a. Fold with H₂. axial plane
① syncline ② recumbent ③ overturned ④ symmetrical

* Fault

↳ fracture / crack / rupture in rock with relative displacement.



Fault zone: clear cut discontinuity which divide rock into blocks.

* Joint:

↳ fracture / crack / rupture in rock without relative displacement.

↳ Relatively less danger

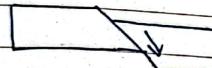
↳ Slip: movement of hanging wall along fault plane.
separation: H₂. distance between two fault plane at same fault rock bed.

Gauge: powder clay like particles formed by rubbing of two blocks.

Fault Breccia: modular fragments (not power).

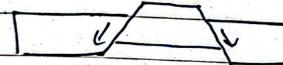
[Types of fault:]

i) Normal fault:



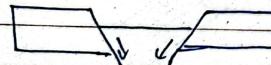
↳ Hanging wall moves downward.

ii) Horst

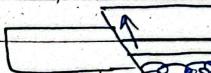


↳ Double hanging wall moving downward.

iii) Graben



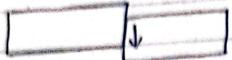
iv) Reverse fault:



↳ If hanging wall moves upward.

(v) strike slip / transcurrent fault:
↳ two blocks move horizontally from each other.

(vi) vertical fault:



↳ if hanging wall moves vertically downward.

Joint:-

(x) Depending on spatial relation

(a) systematic joint:

↳ Regular.

↳ can be mapped.

b

(b) Non systematic joint:

↳ irregular cannot be mapped.

↳ can't be mapped.

x. (ii) Depending on joint geometry:

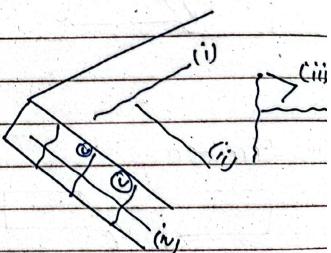
(i) strike joint

(ii) Dip joint

(iii) oblique joint.

(iv) longitudinal joint. or S-joint.

(v) cross joint or Q-joint.



- ↳ depending on origin
(i) tension joint
(ii) compression joint
(iii) shear joint.

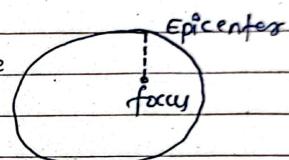
Rock cleavage:

→ special structural features of metamorphic rock.

tendency to split the rock to parallel elements/layer.

Earthquake:-

Focus: A point at crust from where disturbance / cracks occurs/starts.



Epicenter: The point on G.s. vertically above focus on which vibrations reach first.

Seismic wave:

Elastic wave generated due to release of energy stored in rock by elastic deformation

Types of waves:

(a) P-wave:

↳ Primary wave.

↳ pull and push wave.

↳ compression effect.

↳ longitudinal in nature.

*Date _____
Page _____*

longitudinal wave:

- particles vibrate in direction of propagation of wave.

shear wave.

$$v_p = \frac{\sqrt{A + 2\eta}}{s}$$

; $A + 2\eta$: elastic coefficient of medium.

travel fast in hard rock. s = density of medium.

(b) S-wave:

- secondary wave
- shear wave / distortion wave.

Transverse in nature: particles vibrate in per direction with propagation

$$v_s = \frac{\sqrt{\eta}}{s}$$

Body wave:

P & S wave:

travel dipper into depth of earth before reimmerge to ground surface.

velocity increases with depth.

(c) L-wave:

known as long wave.

sluggish wave.

surface wave: it mainly confined to surface of earth.

Types of L-waves:

(i) Rayleigh wave:

(ii) Love wave:

*Date _____
Page _____*

Rayleigh wave:

- complex in nature

- particles vibrate partially in direc of propagation and partially in \perp of propagation.

- zig-zag, wavy

- destroy the surface

Measurement of earthquake:

In two terms:

a) Magnitude.

b) Intensity.

Magnitude: (Richter scale)

log₁₀ A of amplitude recorded at 100 cm seismogram at 100 cm from epicenter.

$$M = \log_{10} A - \log_{10} A_0$$

A = amplitude recorded at seismogram (100cm from epicenter).

A_0 = standard earthquake.

Two types:

(i) M_b = magnitude of body wave.
for small EQ.

(ii) M_s = magnitude of surface wave.
for large EQ.

Type: category	Magnitude (M)
A	≥ 7.7
B	7. to 7.7
C	6 to 7
D	5.8 to 6
E	≤ 5.3

Intensity:

- ↳ Based on quantification of disaster.
- (i) 10 point: Rorrel Parrel scale.
- (ii) 12 point: Mercalli scale.

Types of earthquake:

(a) Shallow eq.
focus: (0 - 60 km) depth

(b) Intermediate eq.
focus: 60 - 300 km

(c) Deep eq.
focus: 300 - 700 km (Rare eq.)

(d) If focus shallow depth.
effect of area on it
ex: disaster air!

Depending on causes of eq:

(a) Tectonic eq:

- ↳ due to movement of tectonic plate.
- ↳ due to mostly occurrence.

(b) Non-tectonic eq:

- ↳ due to huge landslide, huge volcano, atomic explosion.

(Geographical)
geological division of Nepal:-

(i) Terai

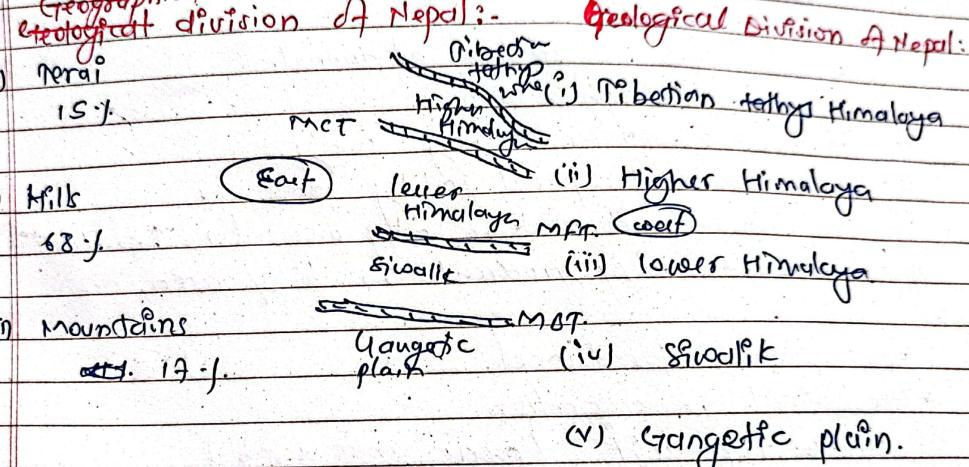
15°S

(ii) Hills

68°S

(iii) Mountains

17°S



Geological division of Nepal:

(i) Tibetan plateau

Higher Himalaya

MCT

Lower Himalaya

Siwalik

Gangotri plain

(ii) Higher Himalaya

MCT

(iii) Lower Himalaya

Siwalik

Gangotri plain

(iv) Siwalik

Gangotri plain

(v) Siwalik