



Introduction to Quantitative Geology

Geological advection: Examples of advection in geological processes

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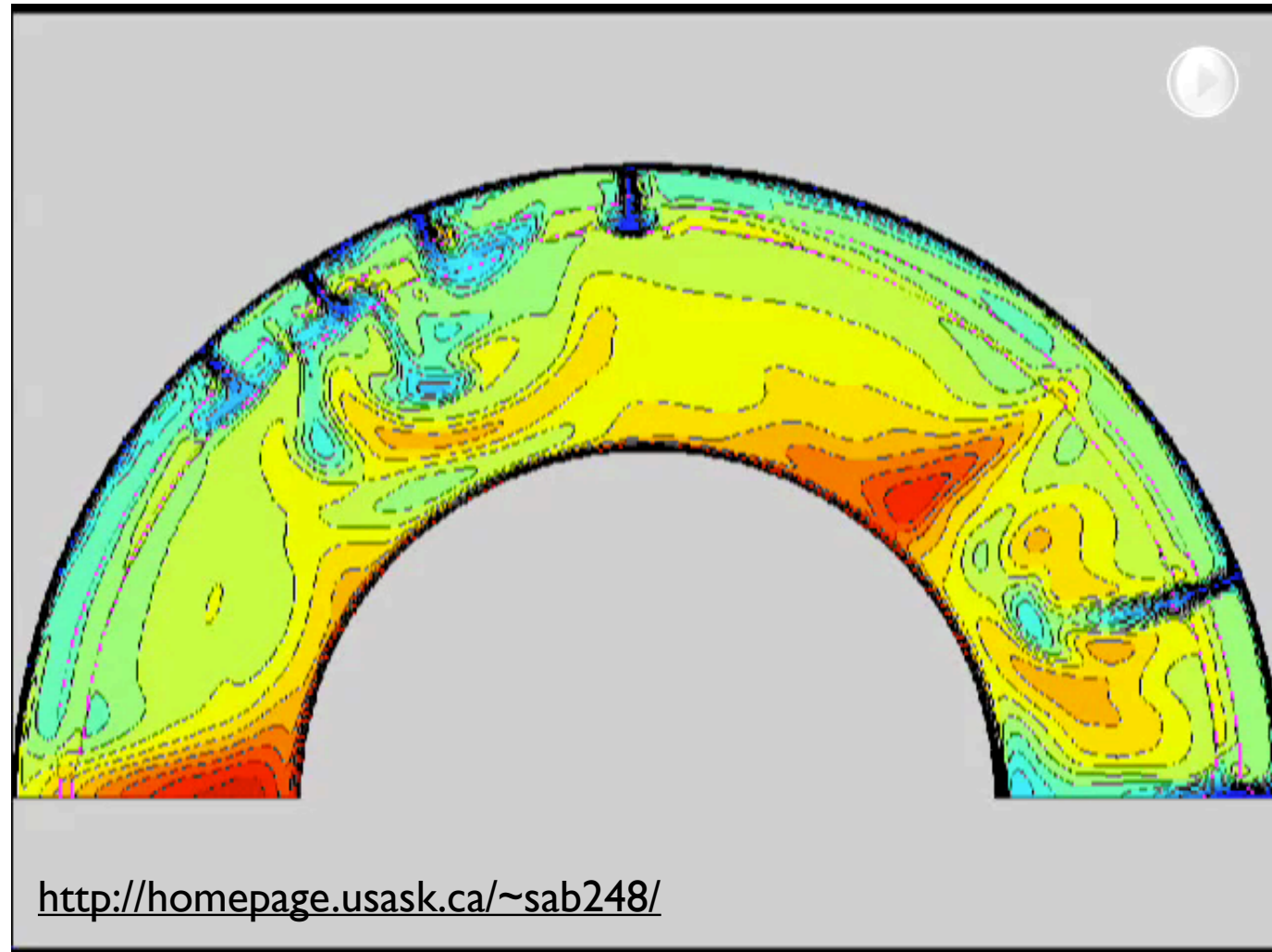


Goals of this lecture

- Introduce the **advection equation**
- Discuss application of the advection equation to **different geological settings**



What is advection?



- Advection involves a lateral translation of some quantity
- For example, the transfer of heat by physical movement of molecules or atoms within a material. A type of convection, mostly applied to heat transfer in solid materials.



Diffusion equation

$$q = -D \frac{\partial C}{\partial x}$$

$$\frac{\partial C}{\partial t} = - \frac{\partial q}{\partial x}$$

- Last week we were introduced to the **diffusion equation**
- Flux (transport of mass or transfer of energy) proportional to a gradient
- Conservation of mass: Any change in flux results in a change in mass/energy



Diffusion equation

Diffusion

$$\frac{\partial C}{\partial t} = D \frac{\partial^2 C}{\partial x^2}$$

$$q = -D \frac{\partial C}{\partial x}$$

$$\frac{\partial C}{\partial t} = - \frac{\partial q}{\partial x}$$

- Substitute the upper equation on the left into the lower to get the classic **diffusion equation**
- q = flux per unit length
 D = diffusivity
 C = concentration
 x = distance
 t = time



Diffusion equation (heat transfer)

Diffusion

$$\frac{\partial T}{\partial t} = \kappa \frac{\partial^2 T}{\partial z^2}$$

- Substitute the upper equation on the left into the lower to get the classic **diffusion equation**
- κ = thermal diffusivity
 T = temperature
 x = distance
 t = time



Diffusion equation (heat transfer)

Diffusion

$$\frac{\partial T}{\partial t} = \kappa \frac{\partial^2 T}{\partial z^2}$$

Advection

$$\frac{\partial T}{\partial t} = v_z \frac{\partial T}{\partial z}$$

- This week we meet the **advection equation**



Diffusion equation (heat transfer)


Diffusion

$$\frac{\partial T}{\partial t} = \kappa \frac{\partial^2 T}{\partial z^2}$$

Advection

$$\frac{\partial T}{\partial t} = v_z \frac{\partial T}{\partial z}$$

- This week we meet the **advection equation**
- Two key differences:
 - Change in mass/energy with time proportional to gradient, rather than curvature (or *change* in gradient)
 - **Advection coefficient** v_z has units of $[L/T]$, rather than $[L^2/T]$



Advection and diffusion equations, a river profile example

River channel profiles

Diffusion

$$\frac{\partial h}{\partial t} = -\kappa \frac{\partial^2 h}{\partial x^2}$$

Advection

$$\frac{\partial h}{\partial t} = c \frac{\partial h}{\partial x}$$

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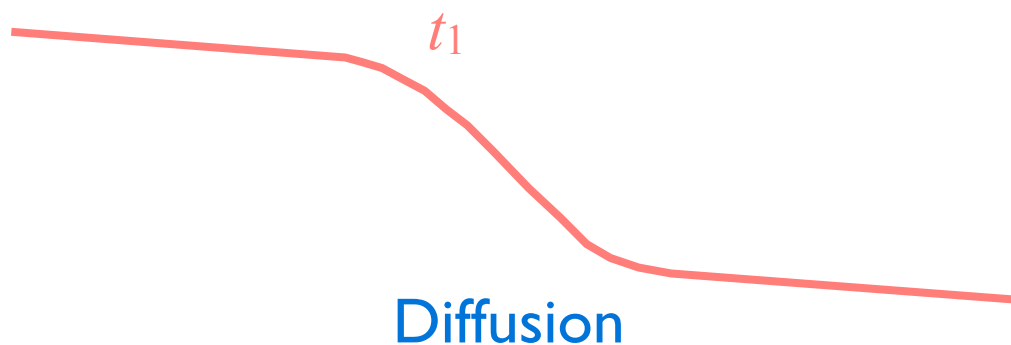



Fig. 1.7, Pelletier, 2008



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River channel profiles

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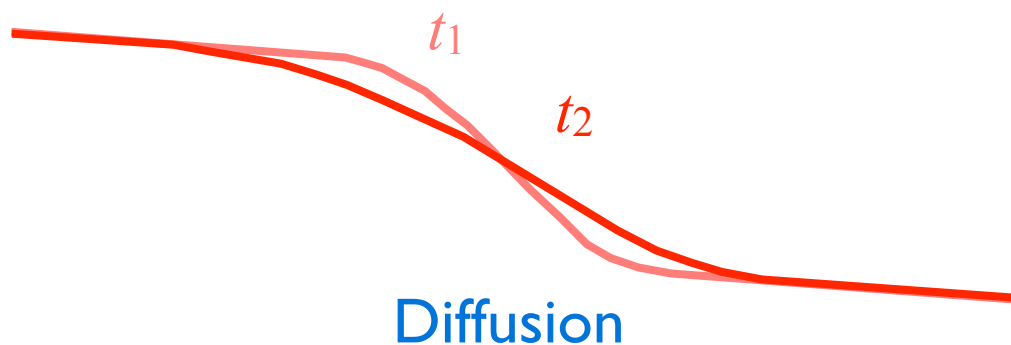



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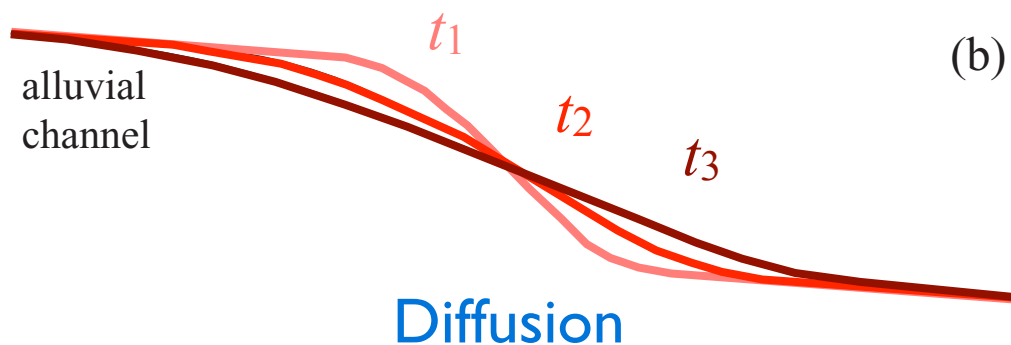



Fig. 1.7, Pelletier, 2008

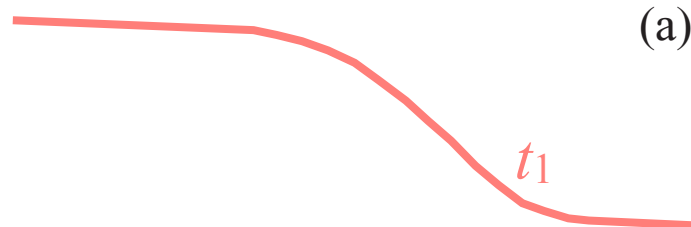
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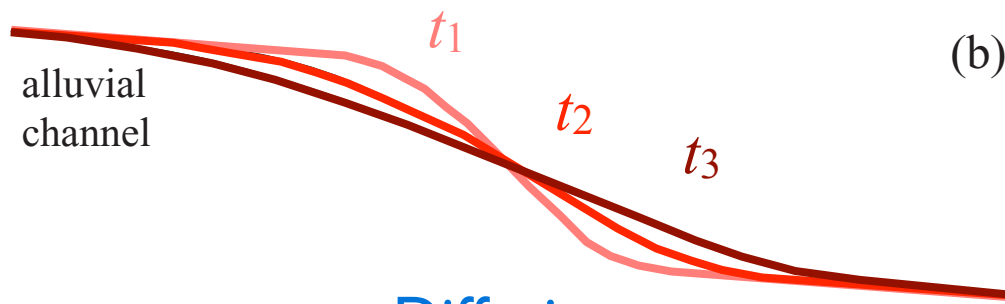
Advection and diffusion equations, a river profile example

River channel profiles

Advection



(a)



Diffusion

alluvial
channel

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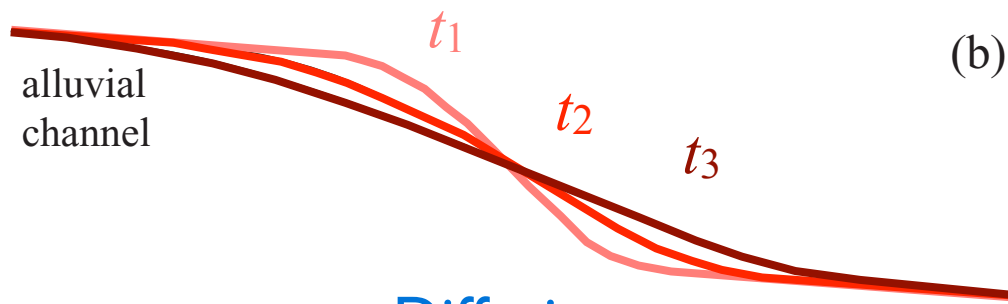
Advection and diffusion equations, a river profile example



River channel profiles

Advection

(a)



Diffusion

(b)

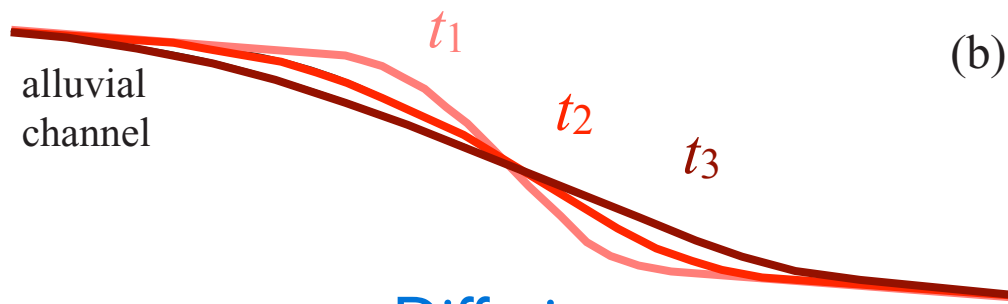


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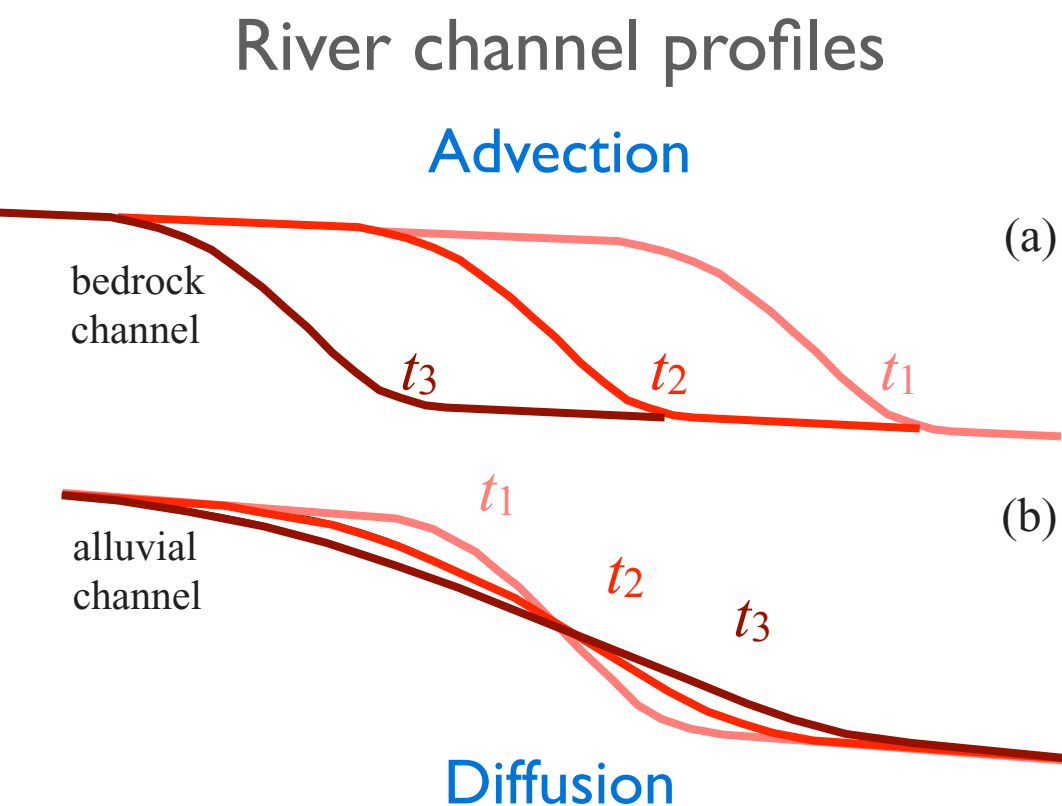


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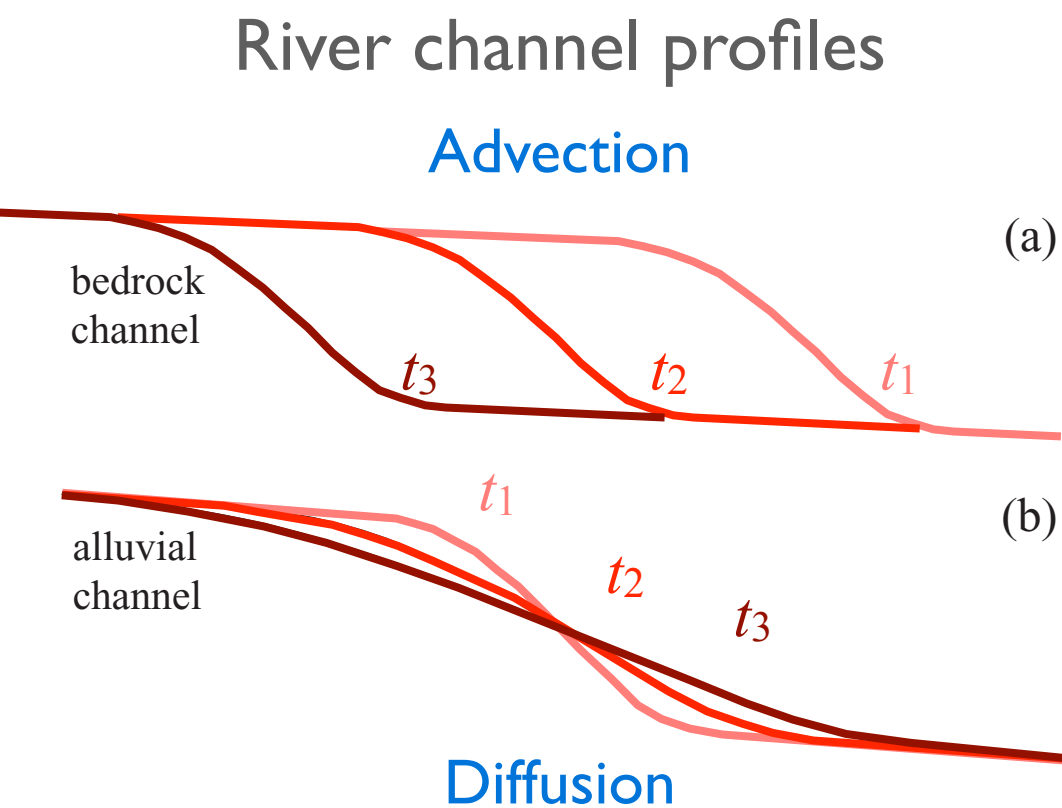


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Advection

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- **Diffusion:** Rate of erosion depends on change in hillslope gradient (curvature)
- **Advection:** Rate of erosion is directly proportional to hillslope gradient
- Also, no conservation of mass (deposition)



Recap

- **What is the main difference between the advection and diffusion equations?**



References

Pelletier, J. D. (2008). *Quantitative modeling of earth surface processes* (Vol. 304). Cambridge University Press.