

# Importance of Dispersion for Shoaling Waves

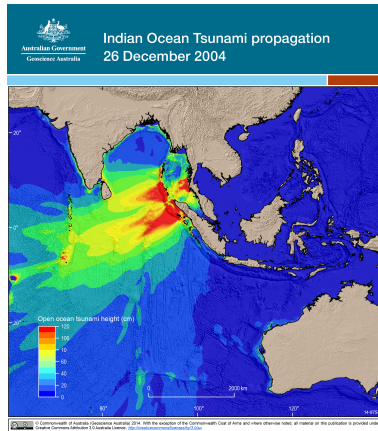
Jordan Pitt, Stephen Roberts and Christopher Zoppou  
Australian National University

January 18, 2018

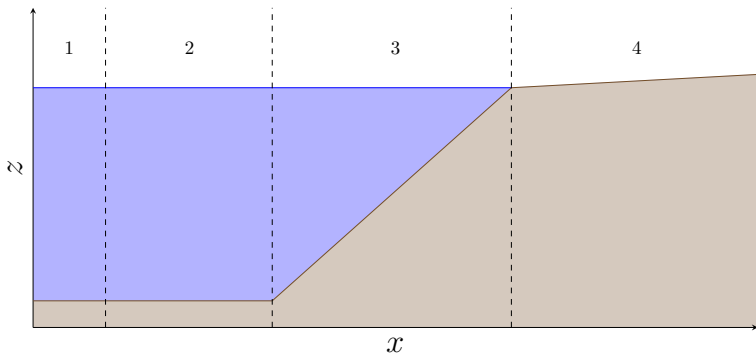
# Introduction

- ▶ Motivation : Tsunamis
- ▶ Model : Shallow Water Wave and Serre equations
- ▶ Experiment : Comparison of numerical solutions

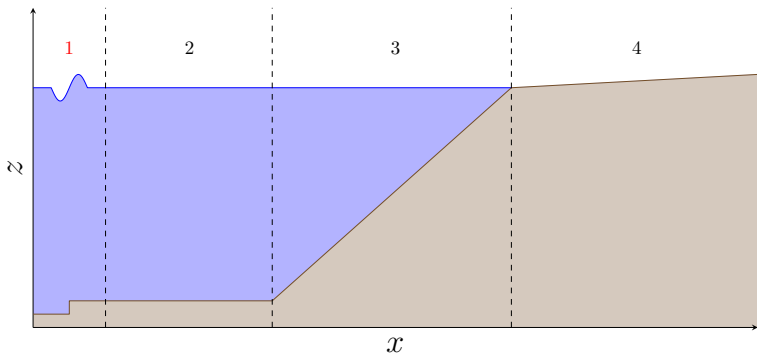
# Indian ocean tsunami



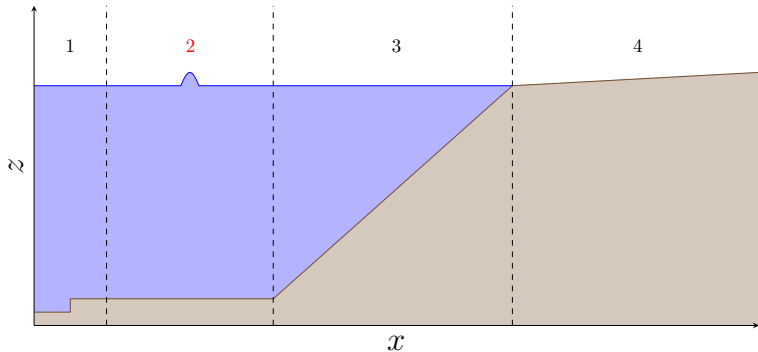
# Tsunami diagram



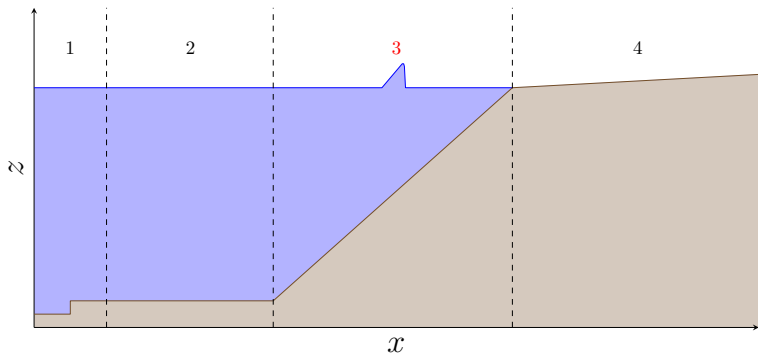
# 1 : Generation



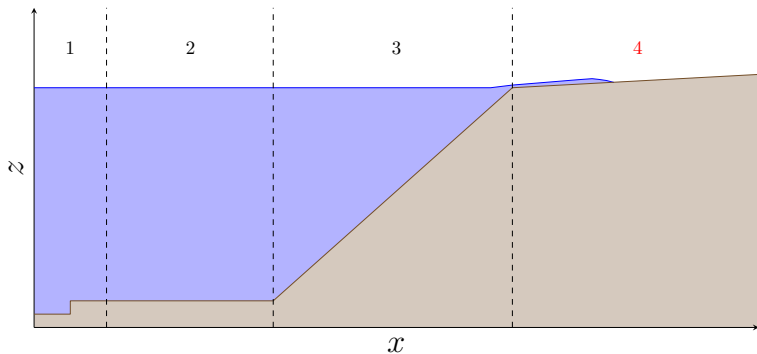
## 2 : Propagation far from coast



### 3 : Propagation near coast

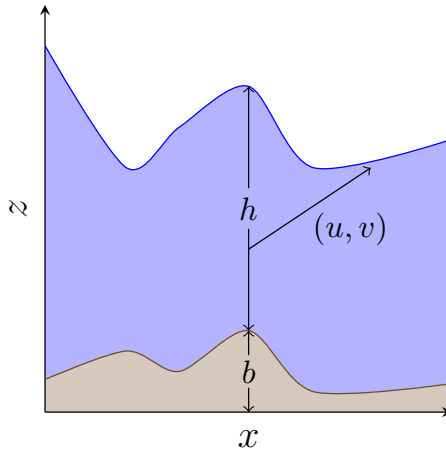


## 4 : Inundation





# Depth averaged equations



# Shallow water wave equations

- ▶ Wavelengths ( $\lambda$ )  $\gg$  Water depth ( $H$ ) ( $\lambda \geq 20H$ )
- ▶ Horizontal velocity constant over  $z$
- ▶ Vertical velocity is 0
- ▶ Pressure is hydrostatic  $p(z) = \rho g(h + b - z)$

# Shallow water wave equations

Conservation of Mass

$$\frac{\partial h}{\partial t} + \frac{\partial(uh)}{\partial x} = 0$$

Conservation of Momentum

$$\frac{\partial(uh)}{\partial t} + \frac{\partial}{\partial x} \left( u^2 h + \frac{gh^2}{2} \right) + gh \frac{\partial b}{\partial x} = 0$$

## Serre equations

- ▶ No restrictions
- ▶ Horizontal velocity is constant over  $z$
- ▶ Vertical velocity is linear in  $z$

$$v'(z) = u \frac{\partial b}{\partial x} - (z - b) \frac{\partial u}{\partial x}$$

- ▶ Pressure

$$p(z) = \rho g(h + b - z) + \rho(h + b - z)\Psi + \frac{\rho}{2}(h + b - z)(h - b + z)\Phi$$

$$\Psi = \frac{\partial b}{\partial x} \left( \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} \right) + u^2 \frac{\partial b}{\partial x}, \quad \Phi = \frac{\partial u}{\partial x} \frac{\partial u}{\partial x} - u \frac{\partial^2 u}{\partial x^2} - \frac{\partial^2 u}{\partial x \partial t}$$

# Serre equations

Conservation of Mass

$$\frac{\partial h}{\partial t} + \frac{\partial(uh)}{\partial x} = 0$$

Conservation of Momentum

$$\begin{aligned} \frac{\partial(uh)}{\partial t} + \frac{\partial}{\partial x} \left( u^2 h + \frac{gh^2}{2} \right) + gh \frac{\partial b}{\partial x} \\ + \frac{\partial}{\partial x} \left( \frac{h^2}{2} \psi + \frac{h^3}{3} \phi \right) + \frac{\partial b}{\partial x} \left( h \psi + \frac{h^2}{2} \phi \right) = 0 \end{aligned}$$

$$\psi = \frac{\partial b}{\partial x} \left( \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} \right) + u^2 \frac{\partial b}{\partial x}, \quad \phi = \frac{\partial u}{\partial x} \frac{\partial u}{\partial x} - u \frac{\partial^2 u}{\partial x^2} - \frac{\partial^2 u}{\partial x \partial t}$$

# Differences

Differences:

- ▶ Dispersion
- ▶ Higher order terms

Are they important for tsunamis?

# Aim

- ▶ Compare Shallow Water Wave and Serre equations
- ▶ Highlight different behaviours
- ▶ Highlight possible impacts these differences could make on current simulations

## Numerical Solvers

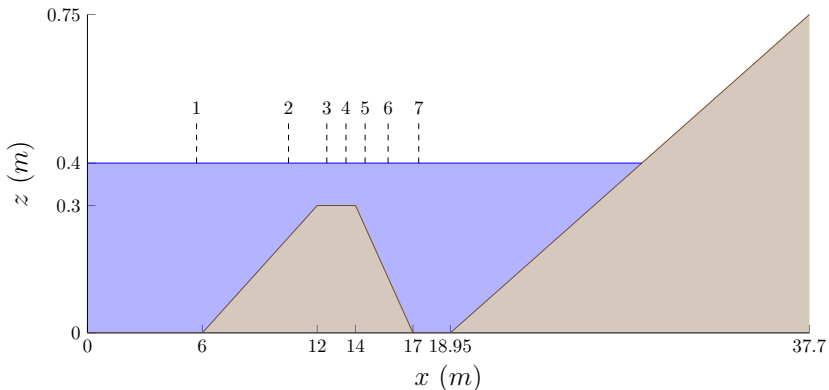
- ▶ Shallow water wave equations: ANUGA, second-order finite volume method
- ▶ Serre equations: second-order finite volume method (same technique as ANUGA) and a second-order finite difference method

## Experiments

- ▶ Experimental results of Beji and Battjes (1994)
- ▶ Artificial example replicating common phenomena



# Periodic waves over a submerged bar: initial conditions



## Wave gauge 1: boundary condition

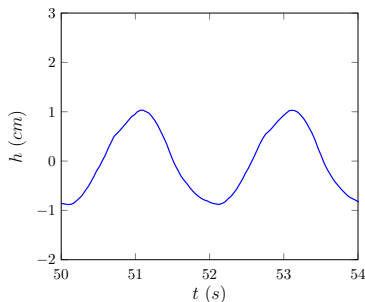


Figure: Low frequency  $\lambda = 3.69m$   
and  $H = 0.4m$

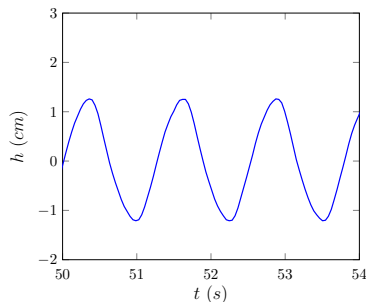


Figure: High frequency  $\lambda = 2.05m$   
and  $H = 0.4m$

## Wave gauge 2: experimental result

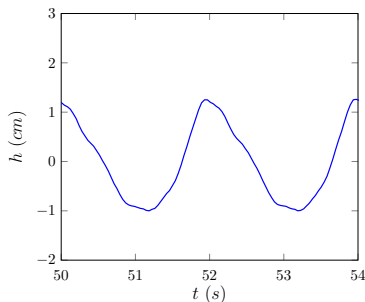


Figure: Low frequency

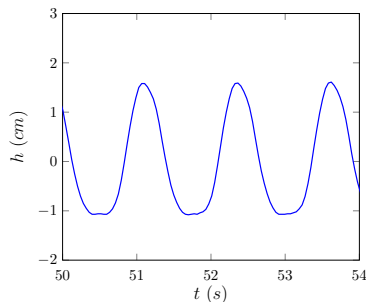
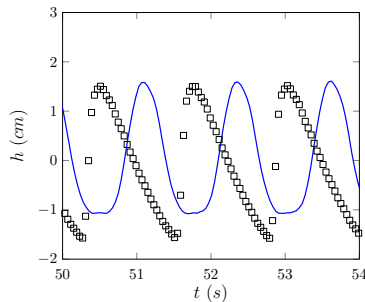
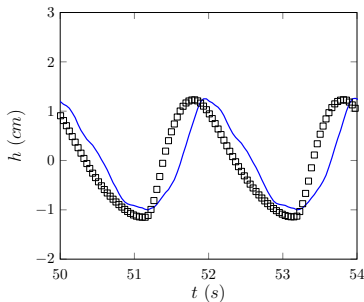


Figure: High frequency

## Wave gauge 2: shallow water wave equation



## Wave gauge 2: all results

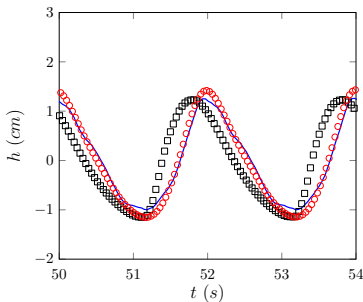


Figure: Low frequency

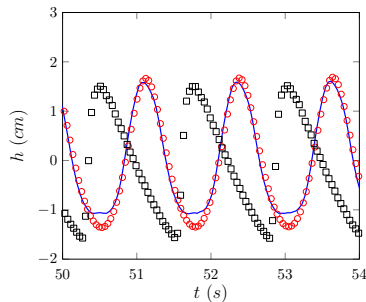


Figure: High frequency

## Wave gauge 3: experimental result

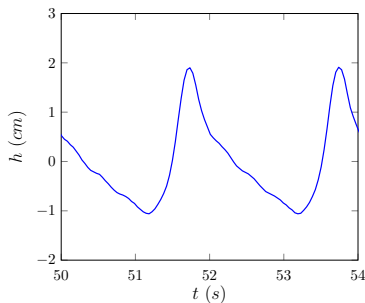


Figure: Low frequency

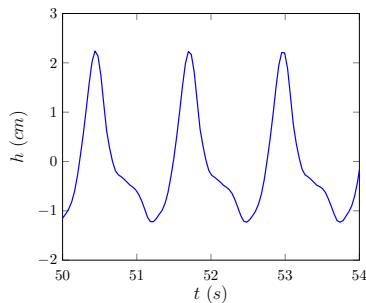
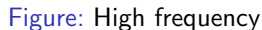


Figure: High frequency

### Wave gauge 3: shallow water wave equation



## Wave gauge 3: all results

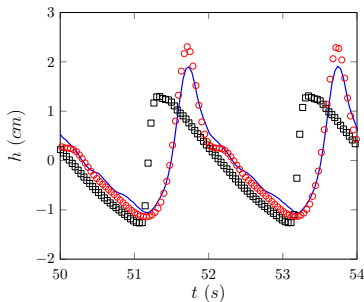


Figure: Low frequency

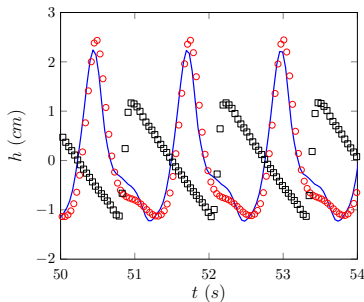


Figure: High frequency



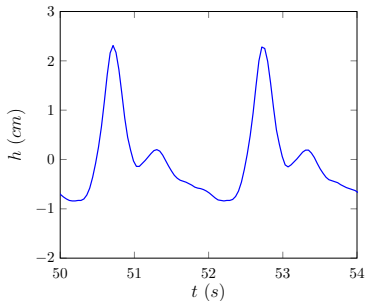


Figure: Low frequency

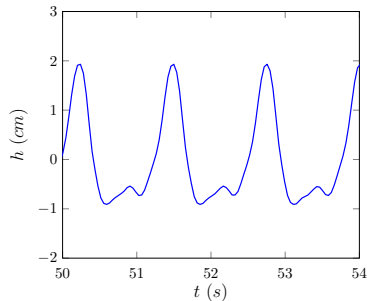


Figure: High frequency

# Wave gauge 4: shallow water wave equation

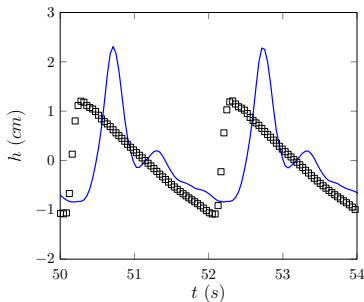


Figure: Low frequency

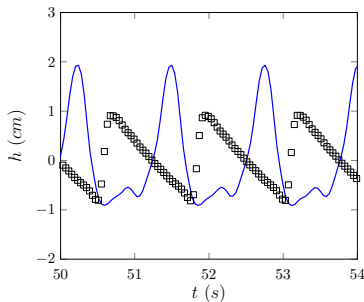


Figure: High frequency

# Wave gauge 4: all results

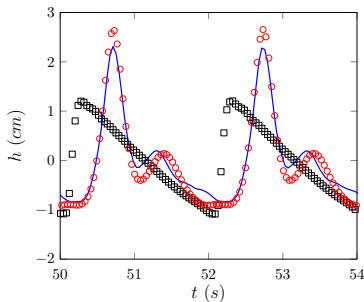


Figure: Low Frequency

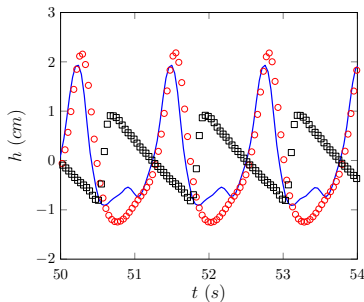
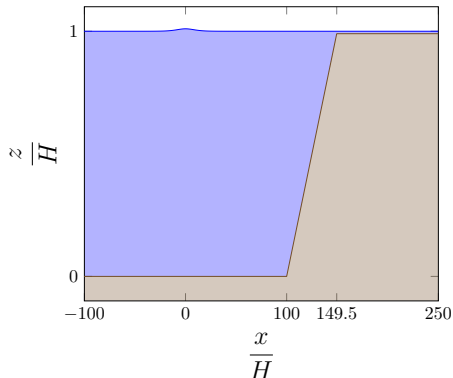


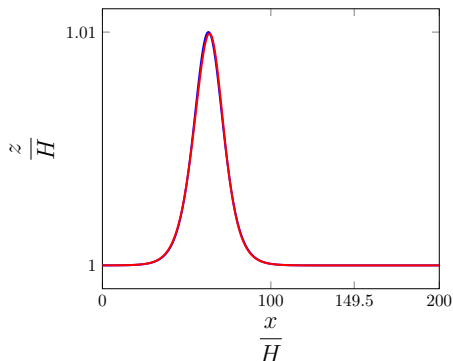
Figure: High Frequency

# Solitary wave over a constant slope: initial conditions



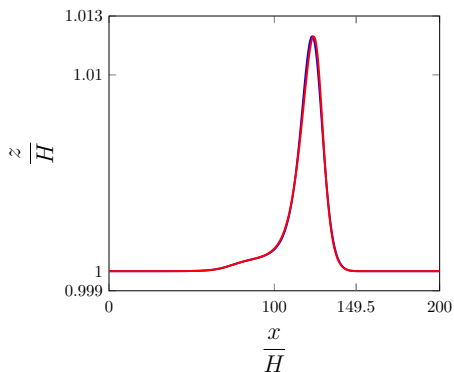
## Solitary wave over a constant slope

Before slope



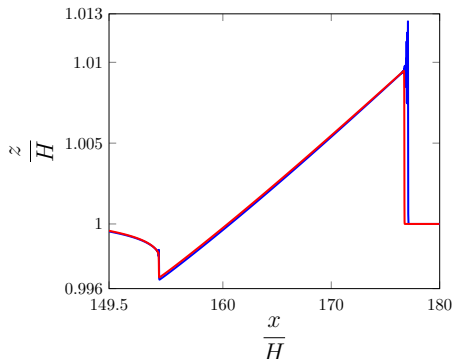
## Solitary wave over a constant slope

## Shoaling



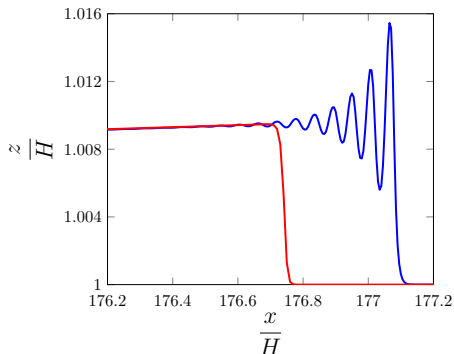
# Solitary wave over a constant slope

## Bore formation



## Solitary wave over a constant slope

## Front of bore





## Conclusion

- ▶ Dispersion plays an important role when wavelengths are not long compared to water depths
- ▶ Dispersion is not important for shoaling of long wavelength waves
- ▶ Dispersion is an important effect for waves after shoaling has occurred.
- ▶ For shoaled waves our current models may underestimate wave amplitude and predict later arrival times.