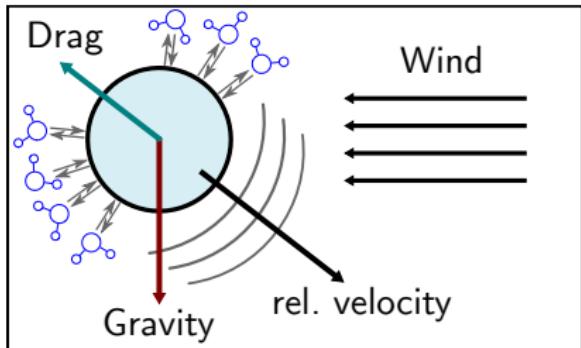


Discrete particle methods for cloud microphysics simulations



Source: World Meteorological Organization (WMO)

Jan Bohrer 21.11.2019

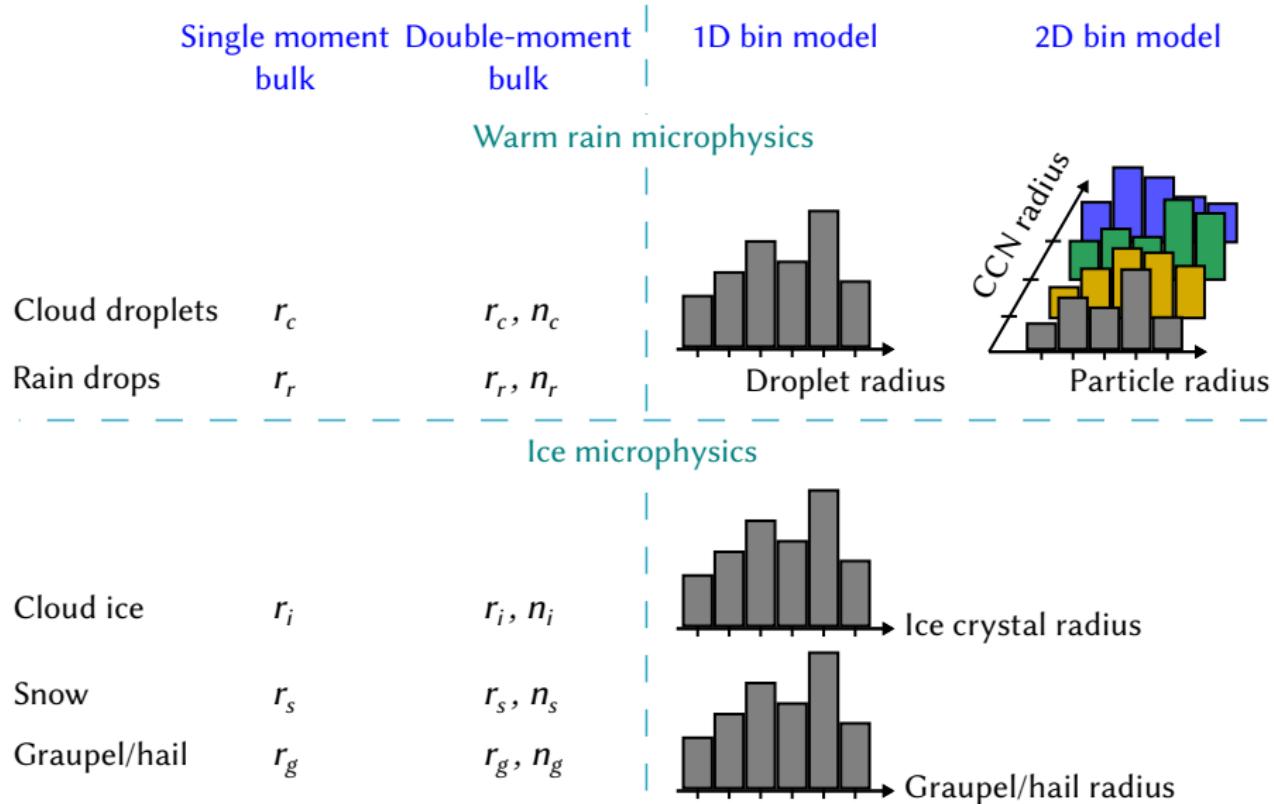
Leibniz Institute for Tropospheric Research Leipzig and
Institute of Physics at the University of Freiburg



Leibniz Institute for
Tropospheric Research



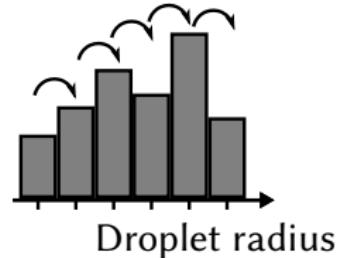
Traditional cloud models



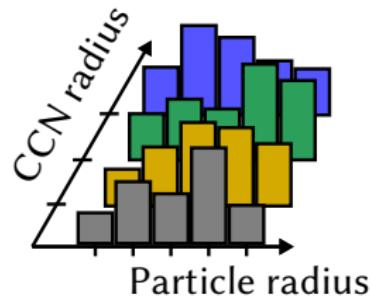
Motivated by [Grabowski *et al.* (2019), Bull. Am. Meteorol. Soc. **100**, 655]

Limitations of bin models

- Numerical diffusion



- "Curse of dimensionality"



Why Lagrangian cloud models?

- No numerical diffusion in radius space
- "Curse of dimensionality" lifted
- Straightforward implementation of multicomponent chemistry
- Detailed description of ice crystals: Shape, phases, growth, ...
- Stochastic treatment of particle collisions and local turbulence

Outline

Atmospheric model

Particle model

Discretization

Collision test in box model

Test case: Drizzling stratocumulus

Conclusions & Outlook

Atmospheric model

Particle model

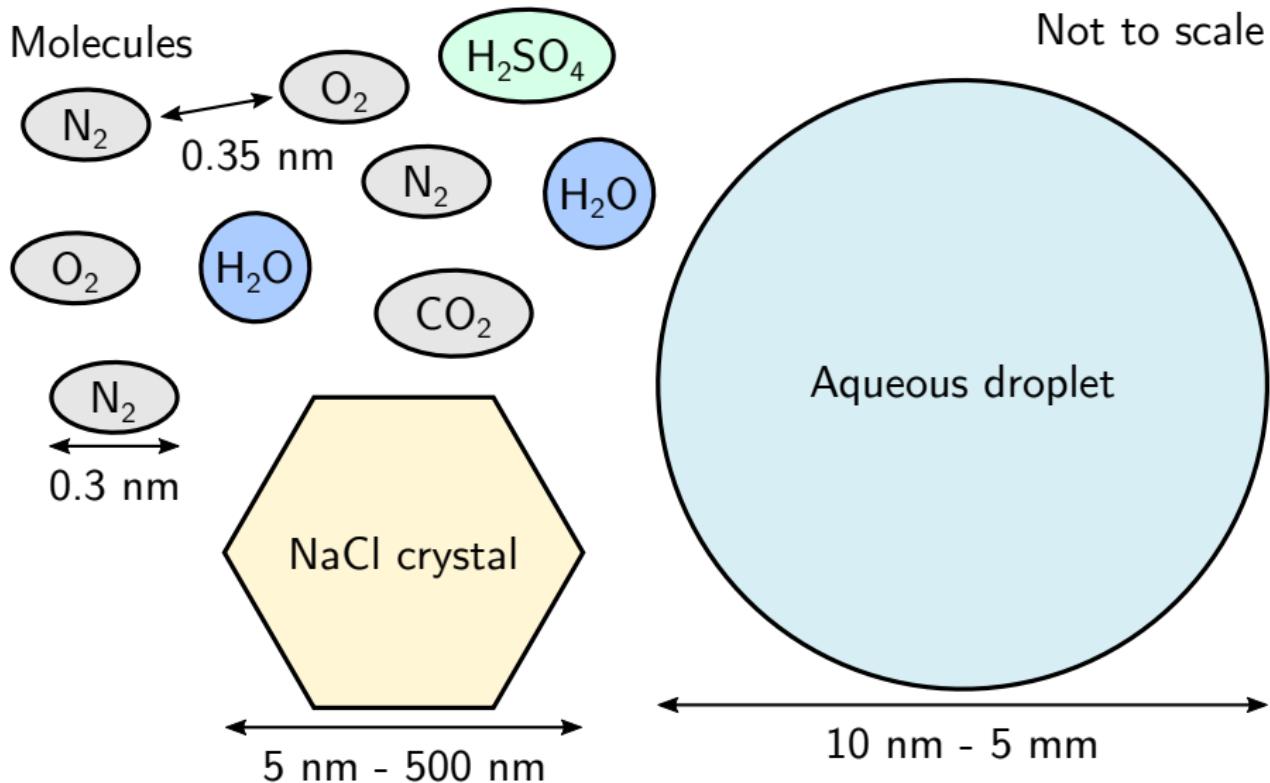
Discretization

Collision test in box model

Test case: Drizzling stratocumulus

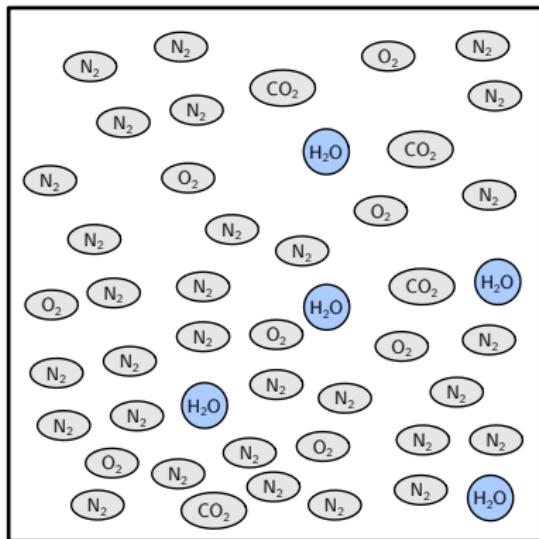
Conclusions & Outlook

Atmospheric composition

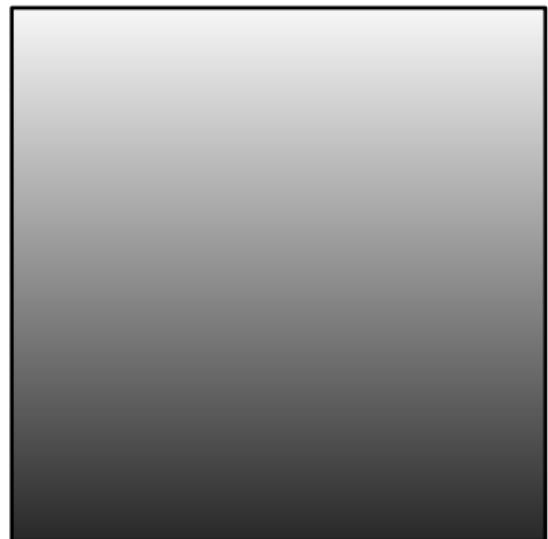


Moist air as continuous fluid

Particle view



Continuous fluid with fields
 $\rho(\mathbf{x}), p(\mathbf{x}), T(\mathbf{x}), \mathbf{u}(\mathbf{x})$



Atmospheric thermodynamics

- Ideal gas equation of state: $p = \tilde{R}_m \rho_m T$

Atmospheric thermodynamics

- Ideal gas equation of state: $p = \tilde{R}_m \rho_m T$
- First law: $dh_e = dq + \frac{dp}{\rho_m} = c_p dT$

Atmospheric thermodynamics

- Ideal gas equation of state: $p = \tilde{R}_m \rho_m T$
- First law: $dh_e = dq + \frac{dp}{\rho_m} = c_p dT$
- Moist potential temperature

$$\Theta_m = T \left(\frac{p^*}{p} \right)^{(\tilde{R}_m / c_p)}$$

Atmospheric thermodynamics

- Ideal gas equation of state: $p = \tilde{R}_m \rho_m T$
- First law: $dh_e = dq + \frac{dp}{\rho_m} = c_p dT$
- Moist potential temperature

$$\Theta_m = T \left(\frac{p^*}{p} \right)^{(\tilde{R}_m / c_p)}$$

- Reversible heat change

$$dq = T ds = c_p \frac{T}{\Theta_m} d\Theta_m$$

Atmospheric thermodynamics

- Ideal gas equation of state: $p = \tilde{R}_m \rho_m T$
- First law: $dh_e = dq + \frac{dp}{\rho_m} = c_p dT$
- Moist potential temperature

$$\Theta_m = T \left(\frac{p^*}{p} \right)^{(\tilde{R}_m / c_p)}$$

- Reversible heat change
- Condensation heat rate (per fluid mass)

$$\dot{q}_{\text{con}} = L_v \frac{\phi_{\text{con}}}{\rho_m}$$

Atmospheric thermodynamics

- Ideal gas equation of state: $p = \tilde{R}_m \rho_m T$

- First law: $dh_e = dq + \frac{dp}{\rho_m} = c_p dT$

- Moist potential temperature

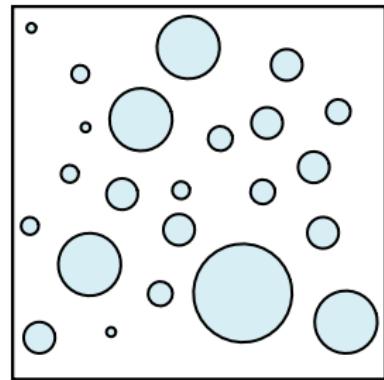
$$\Theta_m = T \left(\frac{p^*}{p} \right)^{(\tilde{R}_m / c_p)}$$

- Reversible heat change

$$dq = T ds = c_p \frac{T}{\Theta_m} d\Theta_m$$

- Condensation heat rate (per fluid mass)

$$\dot{q}_{\text{con}} = L_v \frac{\phi_{\text{con}}}{\rho_m}$$

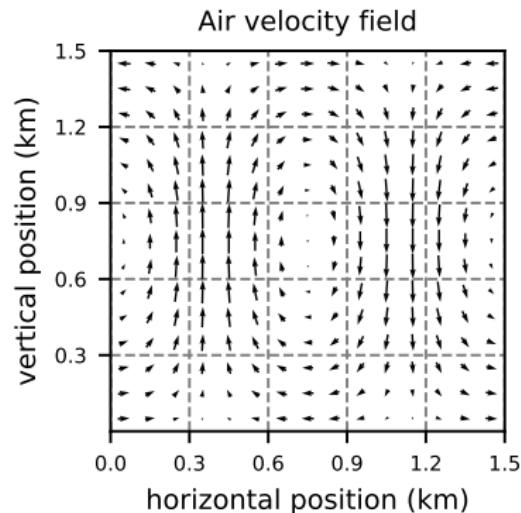
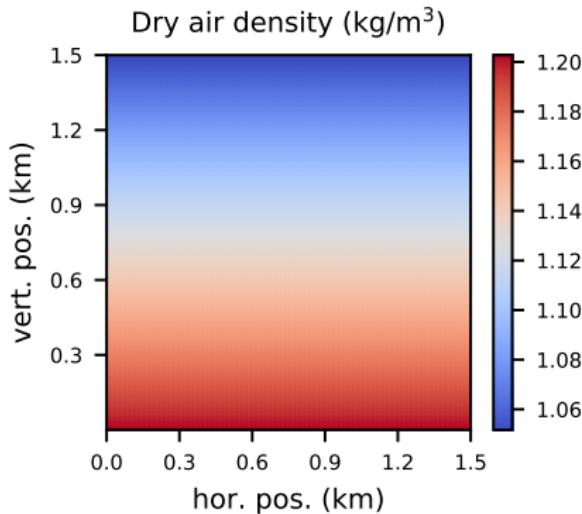


- Condensation mass rate by coarse graining:

$$\Phi = \int_V dV \phi_{\text{con}} = \sum_{x_\alpha \in V} \dot{m}_\alpha$$

Atmospheric kinematic model (2D)

- Test case 1 of the international cloud modeling workshop 2012
[Muhlbauer *et al.* (2013), Bull. Am. Meteorol. Soc. **94**, 45]



- **Stationary** dry air density and dry air mass flux:

$$\rho_{\text{dry}} = \rho_{\text{dry}}(z) \quad \rho_{\text{dry}} \mathbf{u} = \mathbf{j}_{\text{dry}}(x, z) \quad \nabla \cdot \mathbf{j}_{\text{dry}} = 0$$

Atmospheric transport equations

- State variables $\mathbf{G} = (\Theta, r_v)$:

- Dry potential temperature $\Theta = T \left(\frac{p^*}{p_{\text{dry}}} \right)^{(\tilde{R}_{\text{dry}} / c_{\text{dry}})}$
- Water vapor mixing ratio $r_v = \rho_v / \rho_{\text{dry}}$

- Condensation mass density rate ϕ_{con}

Transport equations in the kinematic model

$$\rho_{\text{dry}} \partial_t r_v + \nabla \cdot (\rho_{\text{dry}} \mathbf{u} r_v) = -\phi_{\text{con}}$$

$$\rho_{\text{dry}} \partial_t \Theta + \nabla \cdot (\rho_{\text{dry}} \mathbf{u} \Theta) = \frac{L_v}{c_p (1 + \alpha_v r_v)} \frac{\Theta}{T} \phi_{\text{con}}$$

Atmospheric model

Particle model

Discretization

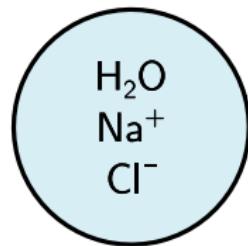
Collision test in box model

Test case: Drizzling stratocumulus

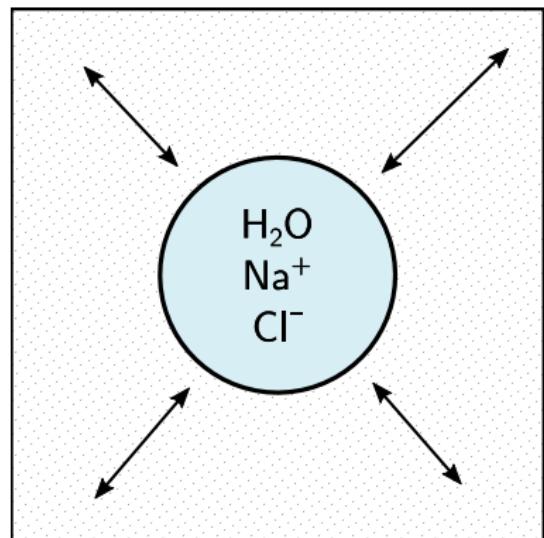
Conclusions & Outlook

Discrete particle model for cloud droplet

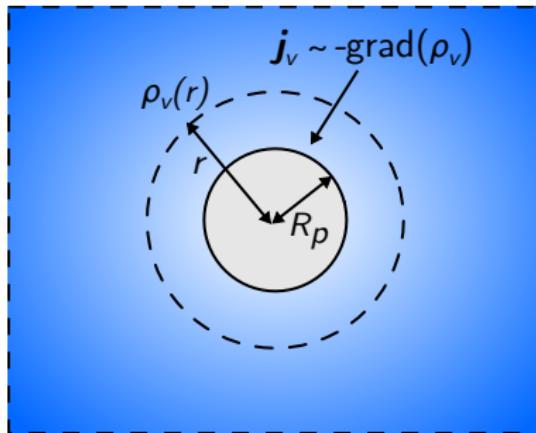
Solution droplet



Fluid-particle interactions?



Mass and heat exchange: Diffusion and conduction

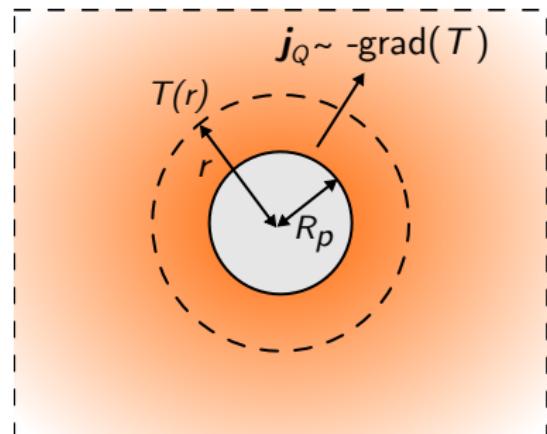
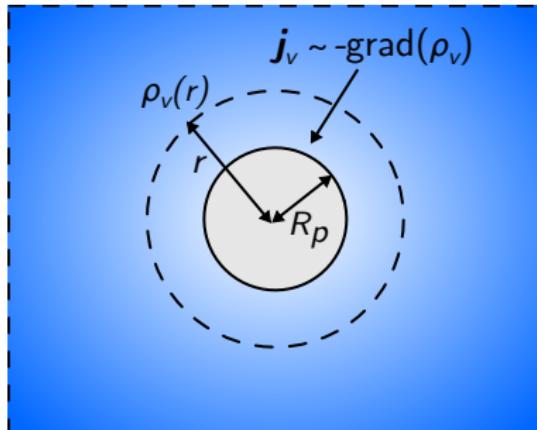


- Diffusion equation for water vapor
(source free region $r > R_p$)

$$\frac{\partial \rho_v}{\partial t} + \nabla \cdot \mathbf{j}_v = 0$$

$$\mathbf{j}_v = -D_v \nabla \rho_v$$

Mass and heat exchange: Diffusion and conduction



- Diffusion equation for water vapor
(source free region $r > R_p$)
- Heat conduction
(source free region $r > R_p$)

$$\frac{\partial \rho_v}{\partial t} + \nabla \cdot \mathbf{j}_v = 0$$

$$\mathbf{j}_v = -D_v \nabla \rho_v$$

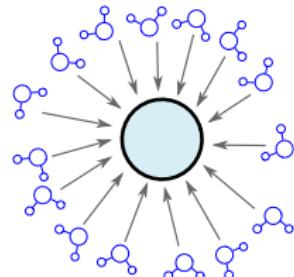
$$\rho c_p \frac{\partial T}{\partial t} + \nabla \cdot \mathbf{j}_Q = 0$$

$$\mathbf{j}_Q = -K \nabla T$$

Droplet growth equation

- Latent heat generation = Thermal conduction flow

$$-L_v \dot{m} = \dot{Q}$$



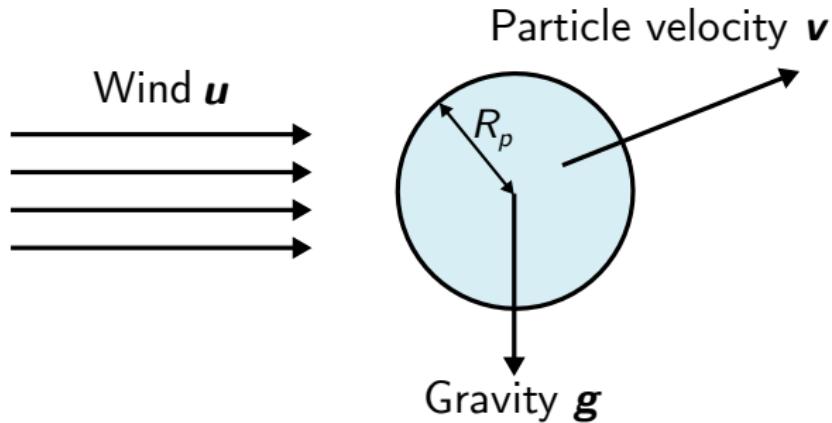
Droplet growth equation

$$\dot{m}_w = 4\pi R_p B(R_p, T, p) (S - S_{\text{eq}})$$

[Fukuta and Walter (1970), J. Atmos. Sci. **27**, 1160]

- Saturation $S = e/e_s(T)$
- Equilibrium saturation $S_{\text{eq}} = f_{\text{Raoult}}(m_s, m_w) f_{\text{Kelvin}}(R_p, T)$

Fluid-particle forces



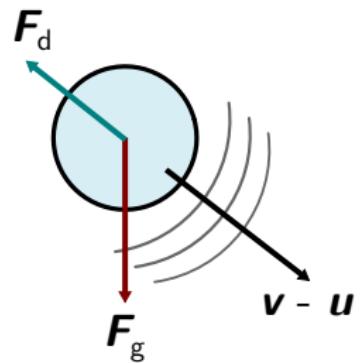
- Relative wind velocity $\tilde{u} = u - v$

- Particle Reynolds number $Re_p = \frac{2\rho_f}{\mu_f} R_p |\tilde{u}|$

Modeled particle forces

- Particle Reynolds number $\text{Re}_p = \frac{2\rho_f}{\mu_f} R_p |\tilde{\mathbf{u}}|$
- Drag force + gravity:

$$\ddot{\mathbf{x}} = \frac{\mathbf{F}_d + \mathbf{F}_g}{m_p} = k_d(R_p, \text{Re}_p) [\mathbf{u}(\mathbf{x}) - \mathbf{v}] + \mathbf{g}$$



Discrete particle model

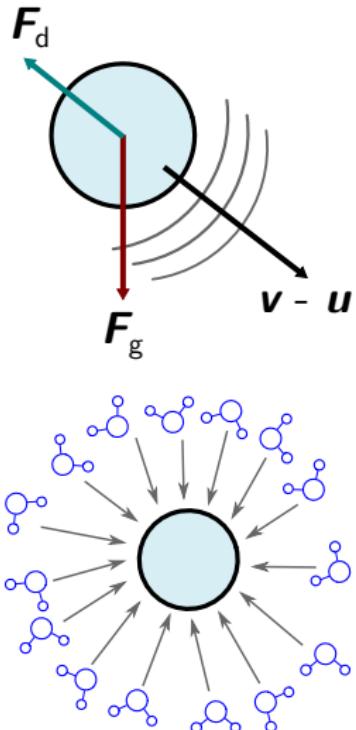
Particle equations of motion

$$\dot{\mathbf{x}} = \mathbf{v}$$

$$\dot{\mathbf{v}} = k_d(R_p, \text{Re}_p) [\mathbf{u} - \mathbf{v}] + \mathbf{g}$$

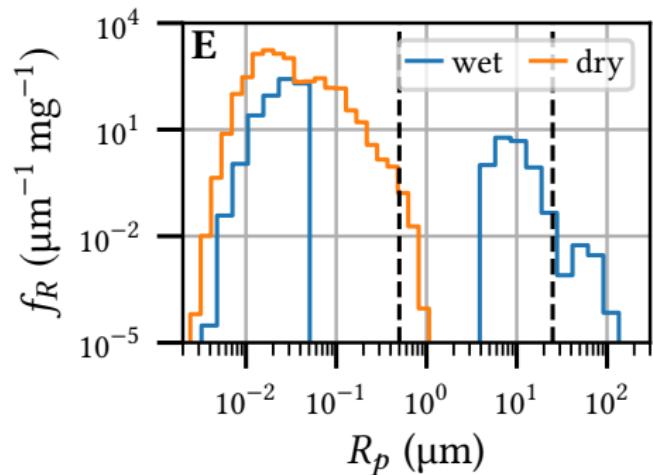
$$\dot{m} = 4\pi R_p B(R_p) [S - S_{\text{eq}}(m)]$$

- **Droplet collisions** coming next



Droplet collision-coalescence

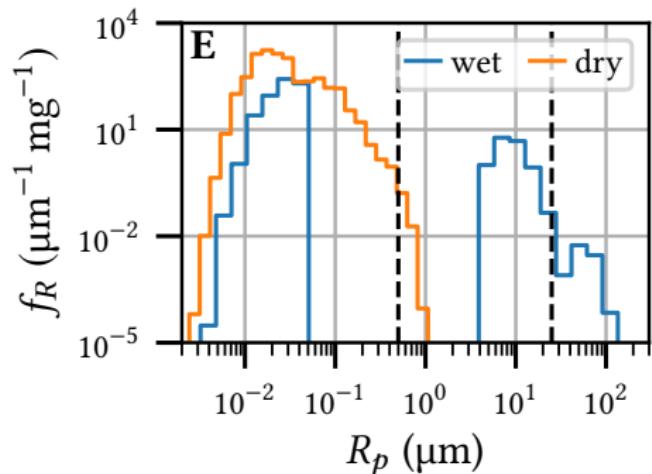
- Droplet size distribution f_R



Droplet collision-coalescence

- Droplet size distribution f_R
- Droplet number concentration

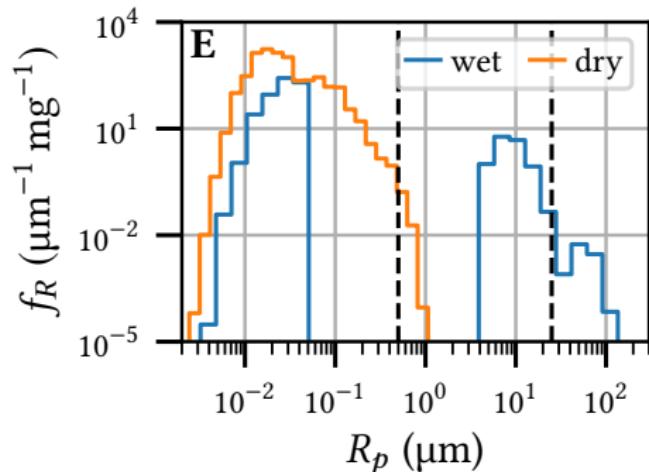
$$\text{DNC}(\mathbf{x}, t) = \int dR f_R(R, \mathbf{x}, t)$$



Droplet collision-coalescence

- Droplet size distribution f_R
- Droplet number concentration

$$\text{DNC}(\mathbf{x}, t) = \int dR f_R(R, \mathbf{x}, t)$$



- Mass distribution $f_m = f_R dR/dm$ with

$$\text{DNC}(\mathbf{x}, t) = \int dm f_m(m, \mathbf{x}, t)$$

Droplet collision-coalescence

- Coalescence equation

$$\begin{aligned}\frac{\partial f_m(m, t)}{\partial t} = & \frac{1}{2} \int_0^m dm' f_m(m', t) f_m(m - m', t) K(m', m - m') \\ & - \int_0^\infty dm' f_m(m, t) f_m(m', t) K(m, m') .\end{aligned}$$

Droplet collision-coalescence

- Coalescence equation

$$\frac{\partial f_m(m, t)}{\partial t} = \frac{1}{2} \int_0^m dm' f_m(m', t) f_m(m - m', t) K(m', m - m') \\ - \int_0^\infty dm' f_m(m, t) f_m(m', t) K(m, m') .$$

- Collection kernel

$$K(m, m') := \lim_{\delta t \rightarrow 0} \frac{P_{\delta t}(m, m') \Delta V}{\delta t}$$

with collection probability $P_{\delta t}(m, m')$

Collision kernels

- Artificial Golovin (sum-of-mass) kernel:

$$K(m, m') = b_G (m + m')$$

- Hydrodynamic collection kernel:

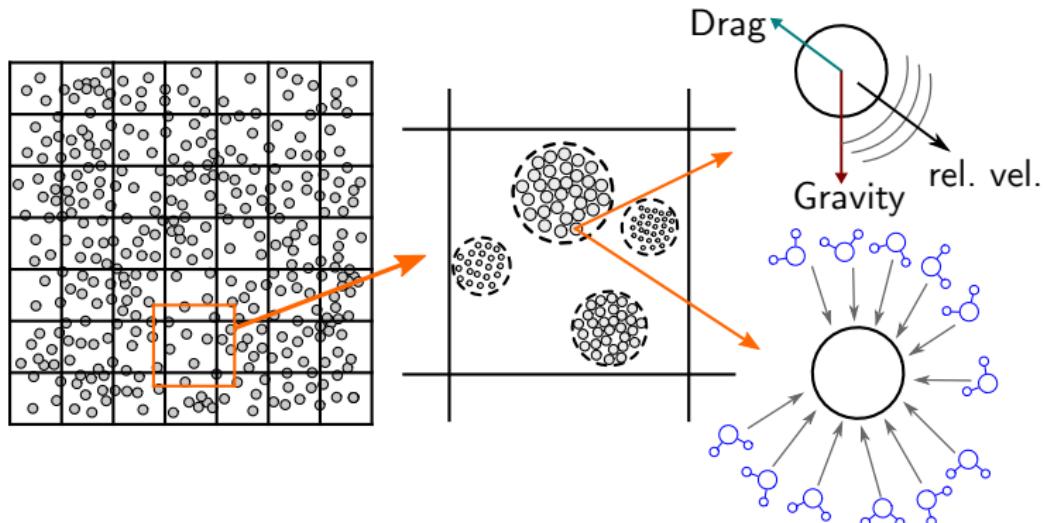
$$K(R_1, R_2) = E_c(R_1, R_2) \pi (R_1 + R_2)^2 |\mathbf{v}_2 - \mathbf{v}_1|$$

Super-droplet approach

- Domain: $1500 \text{ m} \times 1 \text{ m} \times 1500 \text{ m}$
- Cloud condensation nuclei (CCN) number concentration = 100 cm^{-3}
- Total number of particles $\sim 1 \times 10^{14} \Rightarrow$ Problem

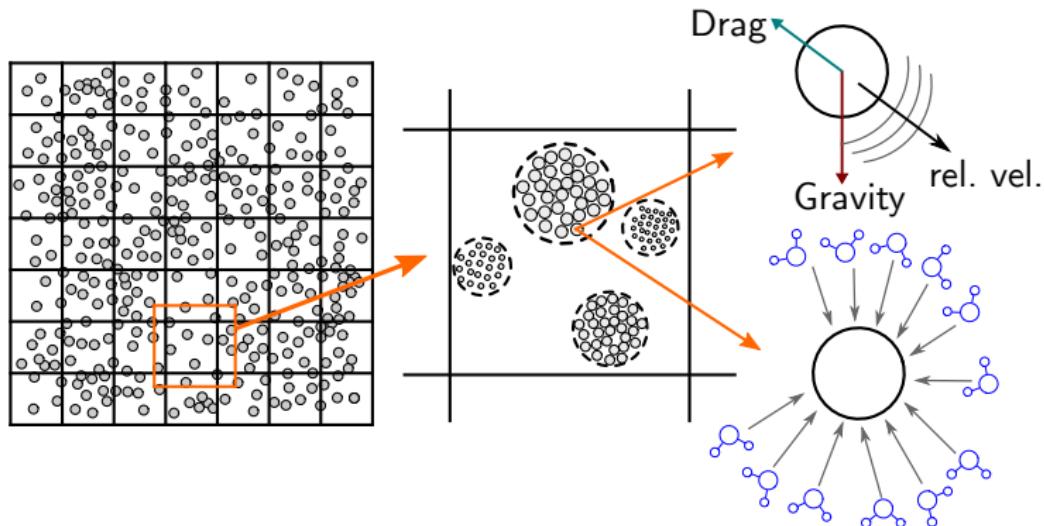
Super-droplet approach

- Domain: $1500 \text{ m} \times 1 \text{ m} \times 1500 \text{ m}$
- Cloud condensation nuclei (CCN) number concentration = 100 cm^{-3}
- Total number of particles $\sim 1 \times 10^{14} \Rightarrow$ Problem
- Super-droplets by [Shima *et al.* (2009), Q. J. R. Meteorol. Soc. **135**, 1307]

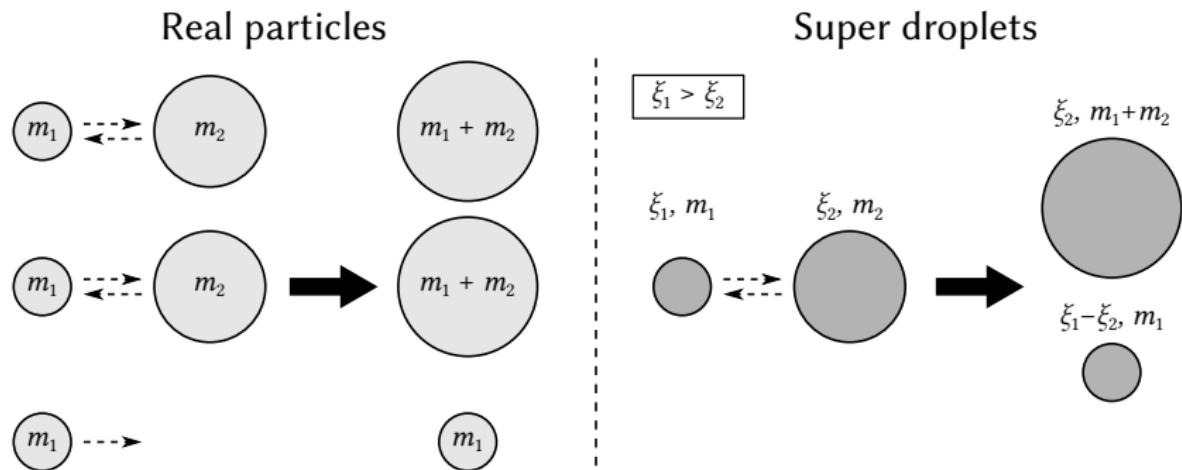


Super-droplet approach

- Super-droplet = Simulation particle (SIP) = Computational particle
- Multiplicity ξ of particles with identical properties
- Equations for (mass m , position \mathbf{x} , velocity \mathbf{v}) as before
- Condensation rate in ΔV : $\phi_{\text{con}} = \sum_{\{i: \mathbf{x}_i \in \Delta V\}} \xi_i \dot{m}_i / \Delta V$



Super-droplet collisions: All-or-nothing approach



- Coalescence of a super-droplets pair during Δt_{col} with

$$P_{\Delta t_{\text{col}}}^S = \xi_1 K(R_1, R_2) \frac{\Delta t_{\text{col}}}{\Delta V}$$

- Total **average** collision rate = $\xi_1 \xi_2 K(R_1, R_2)/\Delta V$

Atmospheric model

Particle model

Discretization

Collision test in box model

Test case: Drizzling stratocumulus

Conclusions & Outlook

Equations of motion

Particles

$$\dot{\mathbf{x}} = \mathbf{v}$$

$$\dot{\mathbf{v}} = k_d(R_p, \mathbf{v}, \mathbf{u}(\mathbf{x})) [\mathbf{u}(\mathbf{x}) - \mathbf{v}] + \mathbf{g}$$

$$\dot{m} = 4\pi R_p B(R_p) [S - S_{\text{eq}}]$$

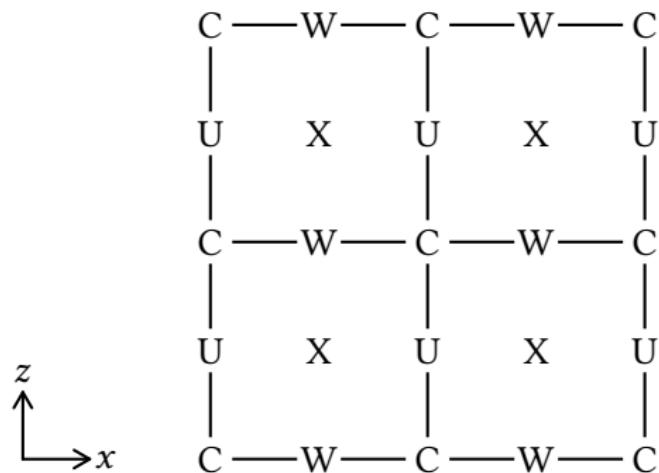
Atmospheric advection

$$\rho_{\text{dry}} \partial_t r_v + \nabla \cdot (\rho_{\text{dry}} \mathbf{u} r_v) = -\phi_{\text{con}}$$

$$\rho_{\text{dry}} \partial_t \Theta + \nabla \cdot (\rho_{\text{dry}} \mathbf{u} \Theta) = \frac{L_v}{c_p (1 + \alpha_v r_v)} \frac{\Theta}{T} \phi_{\text{con}}$$

Spatial grid: 2D rectangular Arakawa C-type grid

- Scalar fields (Θ, r_v) at X
- Flow field $\mathbf{u} = (u, w)^\top$ at (U, W)



Time integration scheme

- State variables $\mathbf{G} = (\Theta, \mathbf{r}_v)$
- Equations of motion (short version):

$$\dot{\mathbf{x}} = \mathbf{v} \quad \dot{\mathbf{v}} = \mathbf{a}(\mathbf{x}, \mathbf{v}, m) \quad \dot{m} = \gamma(\mathbf{x}, m)$$

$$\frac{d\mathbf{G}}{dt} = \mathbf{f}_{\text{adv}}(\mathbf{G}) + \mathbf{f}_{\text{con}}(\mathbf{G}, \phi_{\text{con}})$$

- Time scale separation: **Advection** Δt_{adv} , **collision** Δt_{col} and **condensation** h

Based on [Knoth & Wolke (1998), Appl. Numer. Math. **28**, 327]

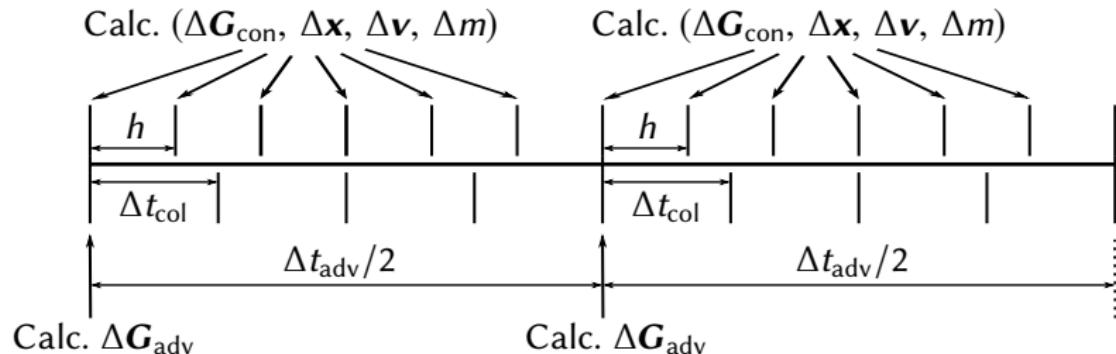
Time integration scheme

- State variables $\mathbf{G} = (\Theta, r_v)$
- Equations of motion (short version):

$$\dot{\mathbf{x}} = \mathbf{v} \quad \dot{\mathbf{v}} = \mathbf{a}(\mathbf{x}, \mathbf{v}, m) \quad \dot{m} = \gamma(\mathbf{x}, m)$$

$$\frac{d\mathbf{G}}{dt} = \mathbf{f}_{\text{adv}}(\mathbf{G}) + \mathbf{f}_{\text{con}}(\mathbf{G}, \phi_{\text{con}})$$

- Time scale separation: **Advection** Δt_{adv} , **collision** Δt_{col} and **condensation** h



Based on [Knoth & Wolke (1998), Appl. Numer. Math. **28**, 327]

Simulation program

- Program written from scratch during master thesis
- Code in pure Python:
 - Numpy data structure
 - Numba for speed (just-in-time compilation)
 - Parallelization possible, but not implemented
- Available on Github:
<https://github.com/JanKBohrer/ProgramMasterthesis>

Atmospheric model

Particle model

Discretization

Collision test in box model

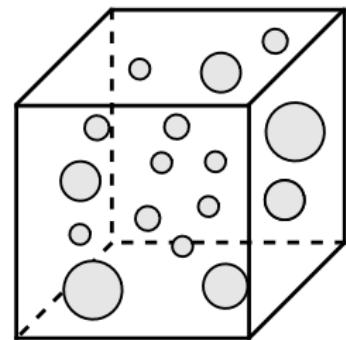
Test case: Drizzling stratocumulus

Conclusions & Outlook

Collision box model

- Particles remain in box ΔV
- Only collisions active
- No condensation/evaporation
- Initial exponential mass distribution

$$f_m(m, t = 0) = \frac{DNC_0}{\bar{m}_0} \exp\left(-\frac{m}{\bar{m}_0}\right)$$



Collision box model

- Particles remain in box ΔV
- Only collisions active
- No condensation/evaporation
- Initial exponential mass distribution

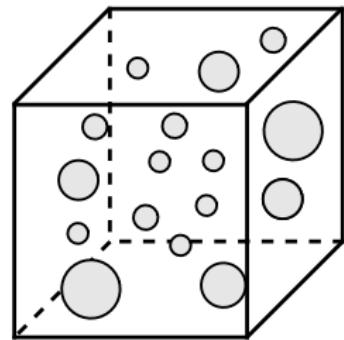
$$f_m(m, t = 0) = \frac{\text{DNC}_0}{\bar{m}_0} \exp\left(-\frac{m}{\bar{m}_0}\right)$$

- Logarithmic radius distribution function

$$g_{\ln(R)} := 3 m^2 f_m$$

- Liquid mass content (total mass per volume)

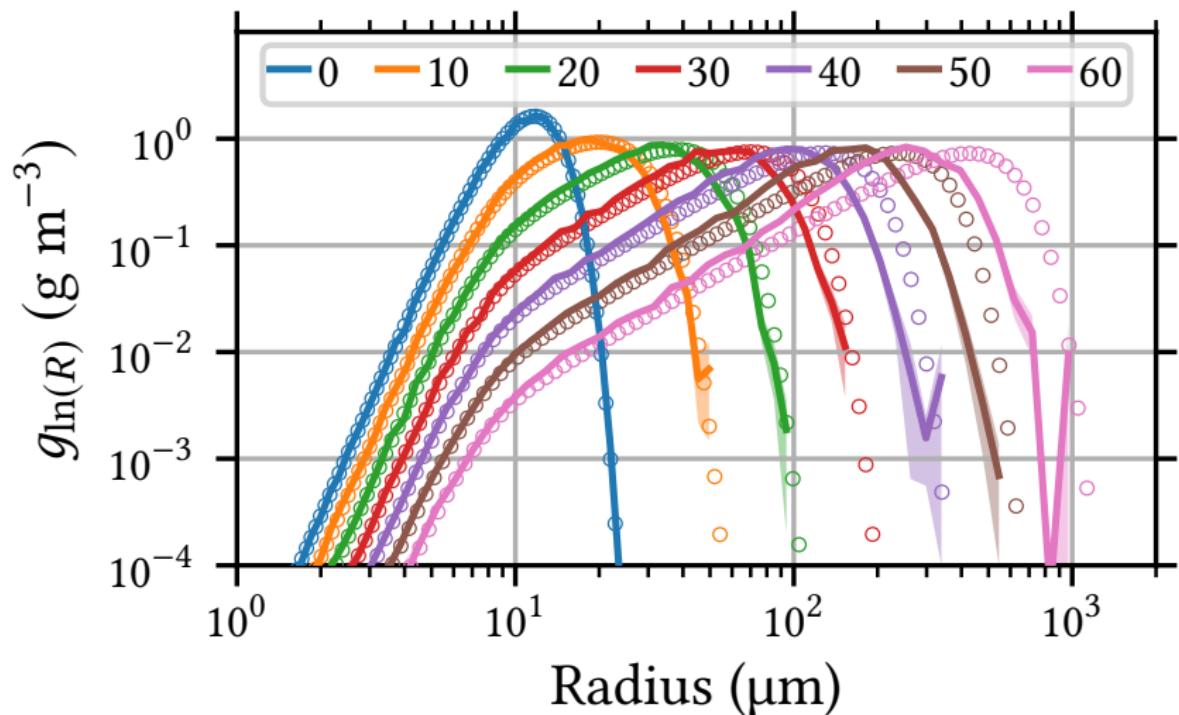
$$\text{LMC} = \int d(\ln R) g_{\ln(R)}(R)$$



Evolution of the size distribution (Golovin collision kernel)

- 50 super-droplets per cell, average over 500 simulation runs

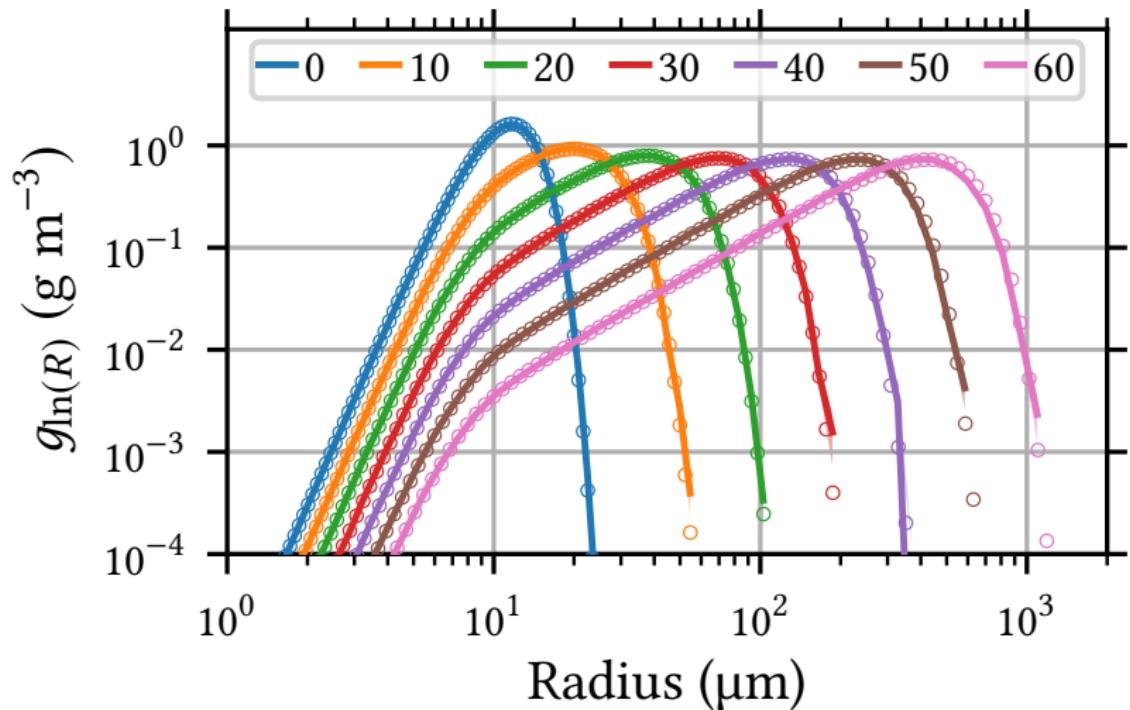
Analytic reference (circles), Legend: Time in minutes



Evolution of the size distribution (Golovin collision kernel)

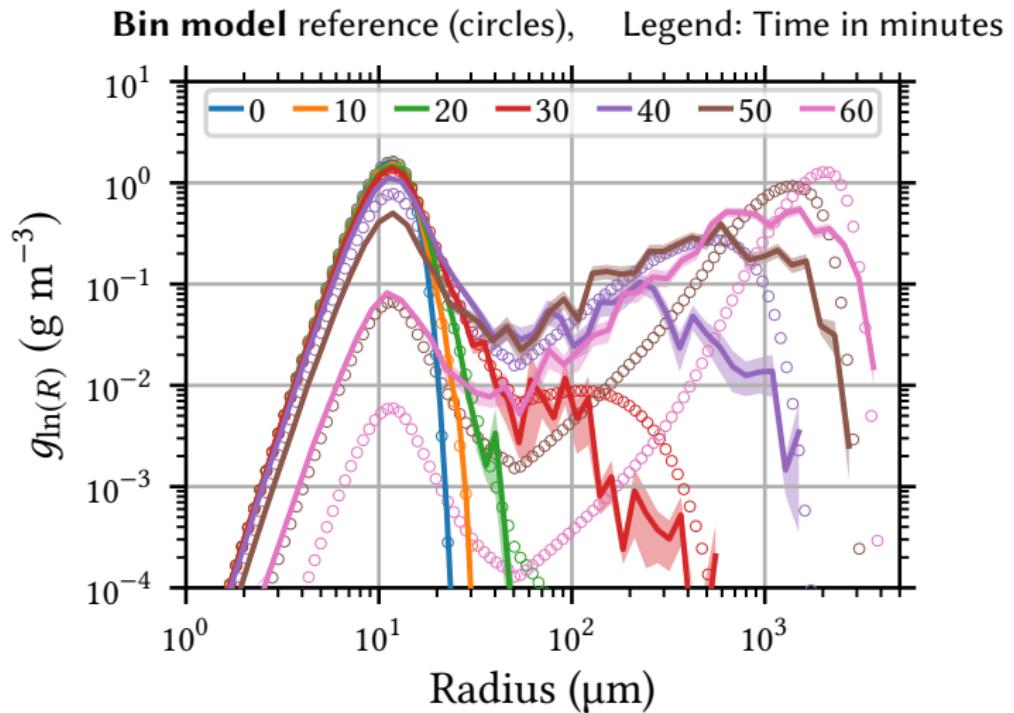
- 1000 super-droplets per cell, average over 500 simulation runs

Analytic reference (circles), Legend: Time in minutes



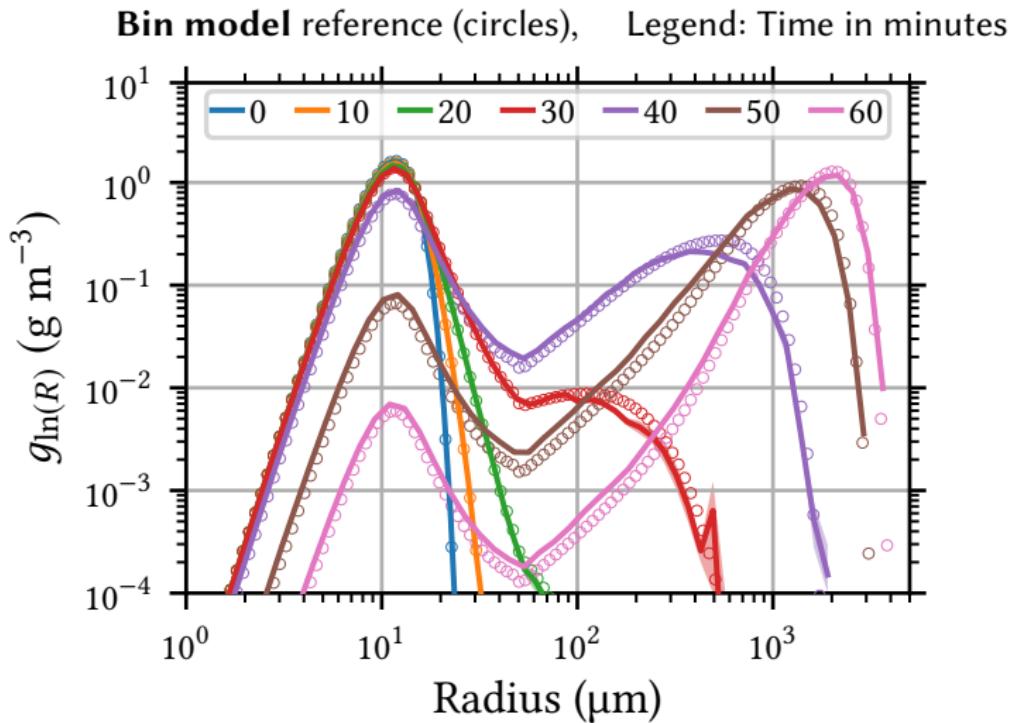
Evolution of the size distribution (Long collision kernel)

- 50 super-droplets per cell, average over 500 simulation runs
- Bin model reference: [Wang, Xue & Grabowski (2007), J. Comp. Phys. **226**, 59]



Evolution of the size distribution (Long collision kernel)

- 15 000 super-droplets per cell, average over 500 simulation runs
- Bin model reference: [Wang, Xue & Grabowski (2007), J. Comp. Phys. **226**, 59]



Atmospheric model

Particle model

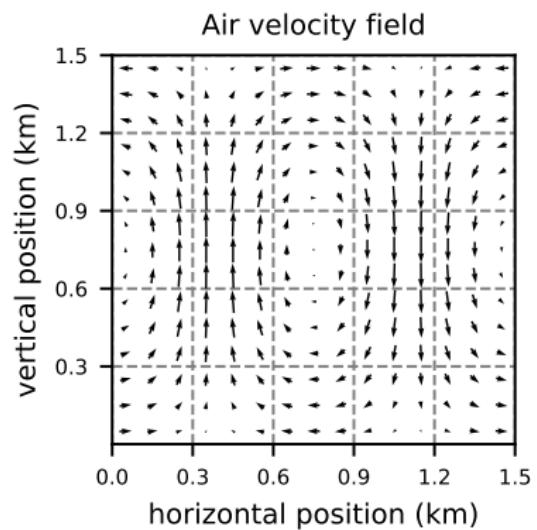
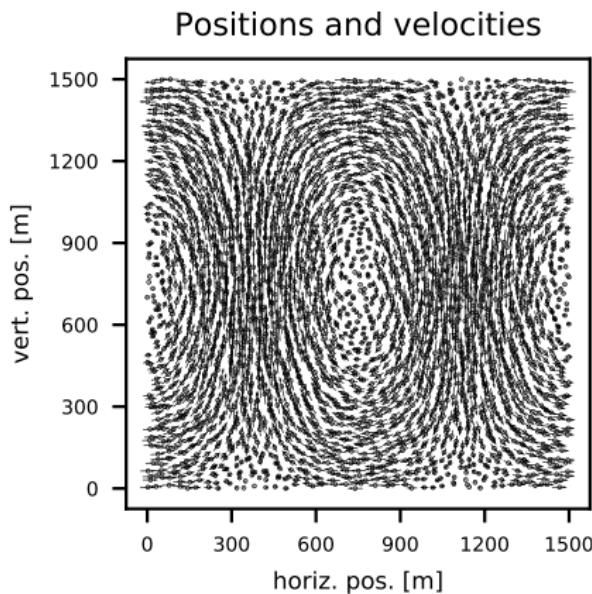
Discretization

Collision test in box model

Test case: Drizzling stratocumulus

Conclusions & Outlook

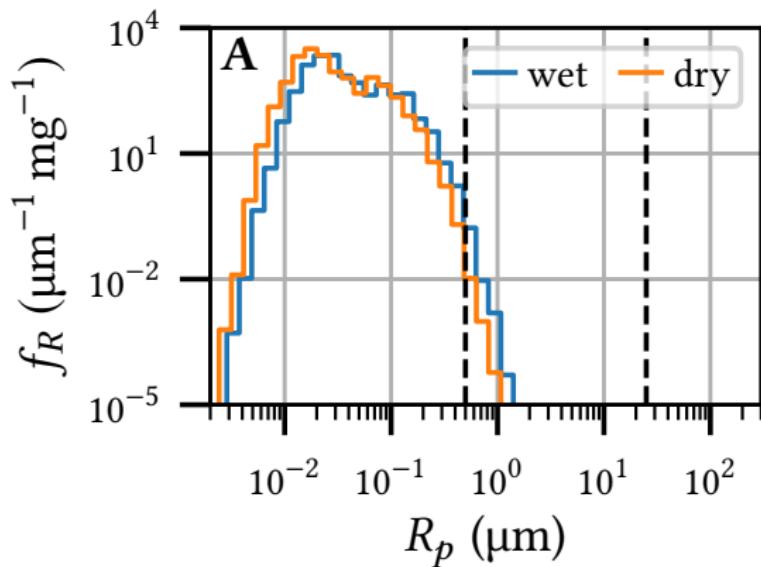
Particle initialization



Particle dry size distribution

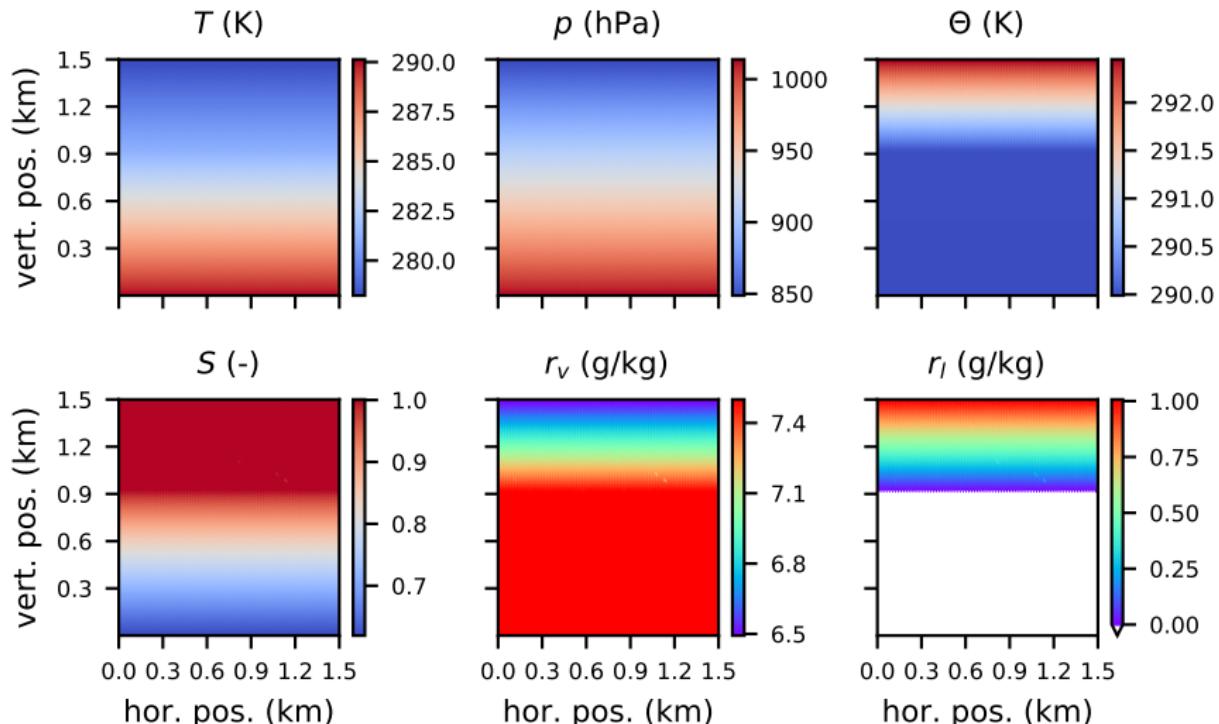
- Bimodal log-normal distribution [Allen *et al.* (2011), *Atmos. Chem. Phys.* **11**, 5237]

$$f_{R,s}(R_s) = \sum_{k=1}^2 \frac{\text{DNC}_k}{\sqrt{2\pi} \ln(\sigma_{R,k}^*) R_s} \exp \left[- \left(\frac{\ln(R_s/\mu_{R,k}^*)}{\sqrt{2} \ln(\sigma_{R,k}^*)} \right)^2 \right]$$



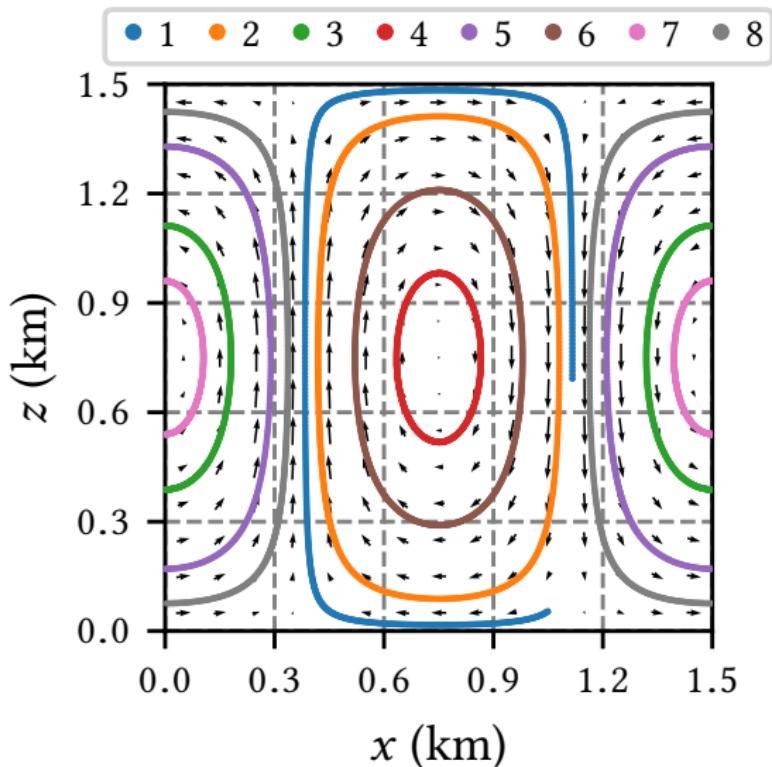
Atmospheric 2D kinematic model: Initialization

- Conditions: $\Theta_{\text{liquid}} = \text{const.} = 289 \text{ K}$, $r_{\text{tot}} = r_v + r_l = \text{const.} = 7.5 \text{ g/kg}$



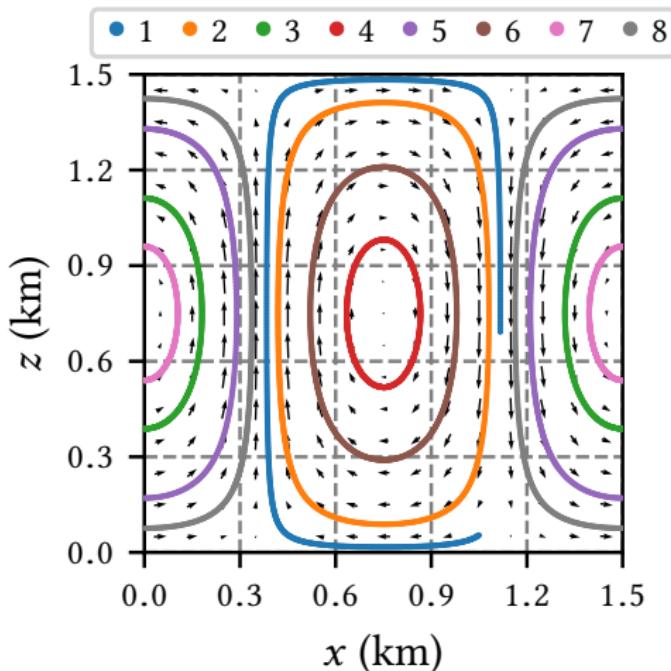
Particle trajectories

- Spin-up phase: No gravity, no collisions

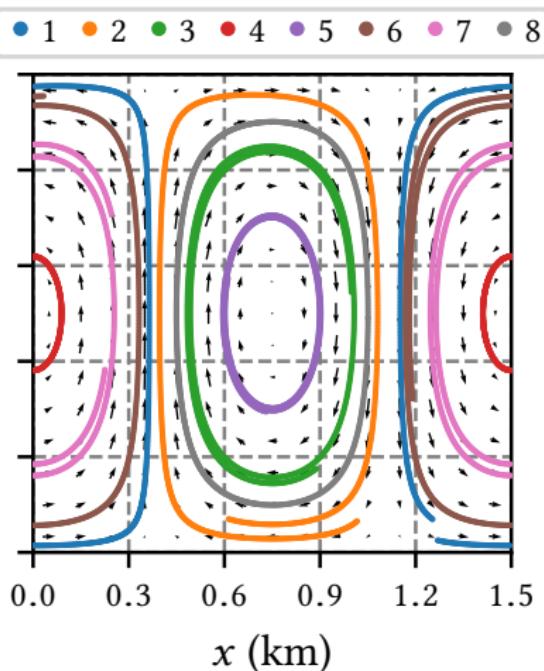


Particle trajectories

no gravity, no collisions

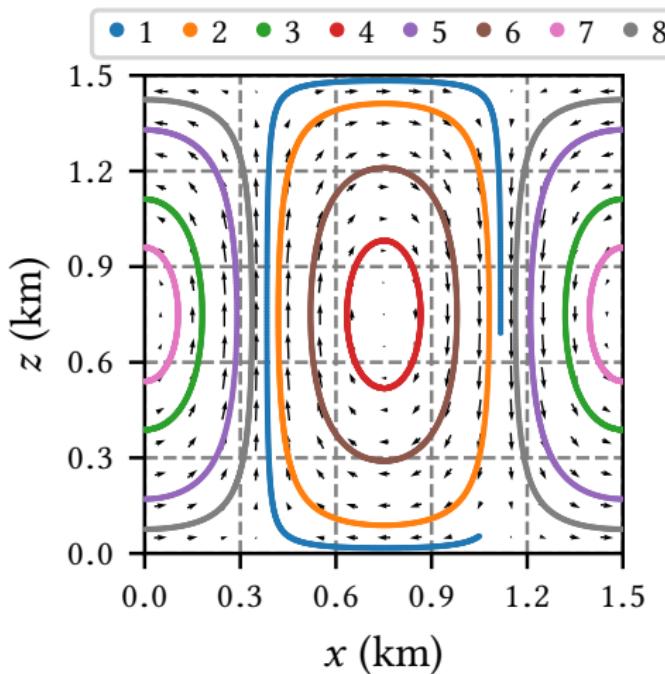


with gravity, no collisions

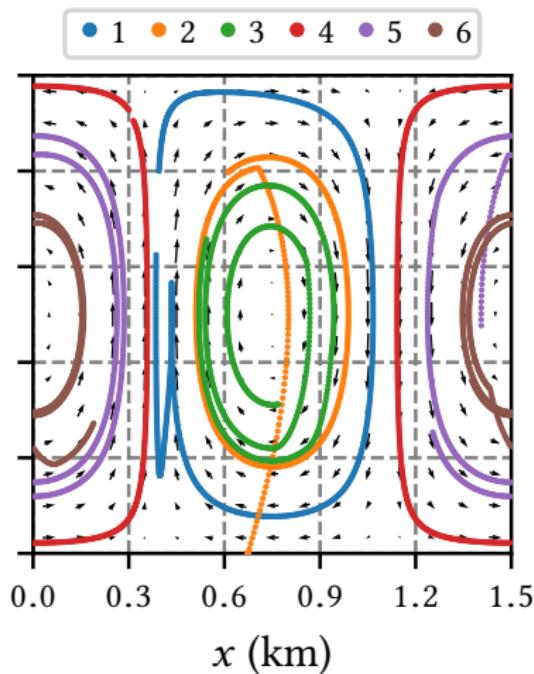


Particle trajectories

no gravity, no collisions



with gravity, with collisions

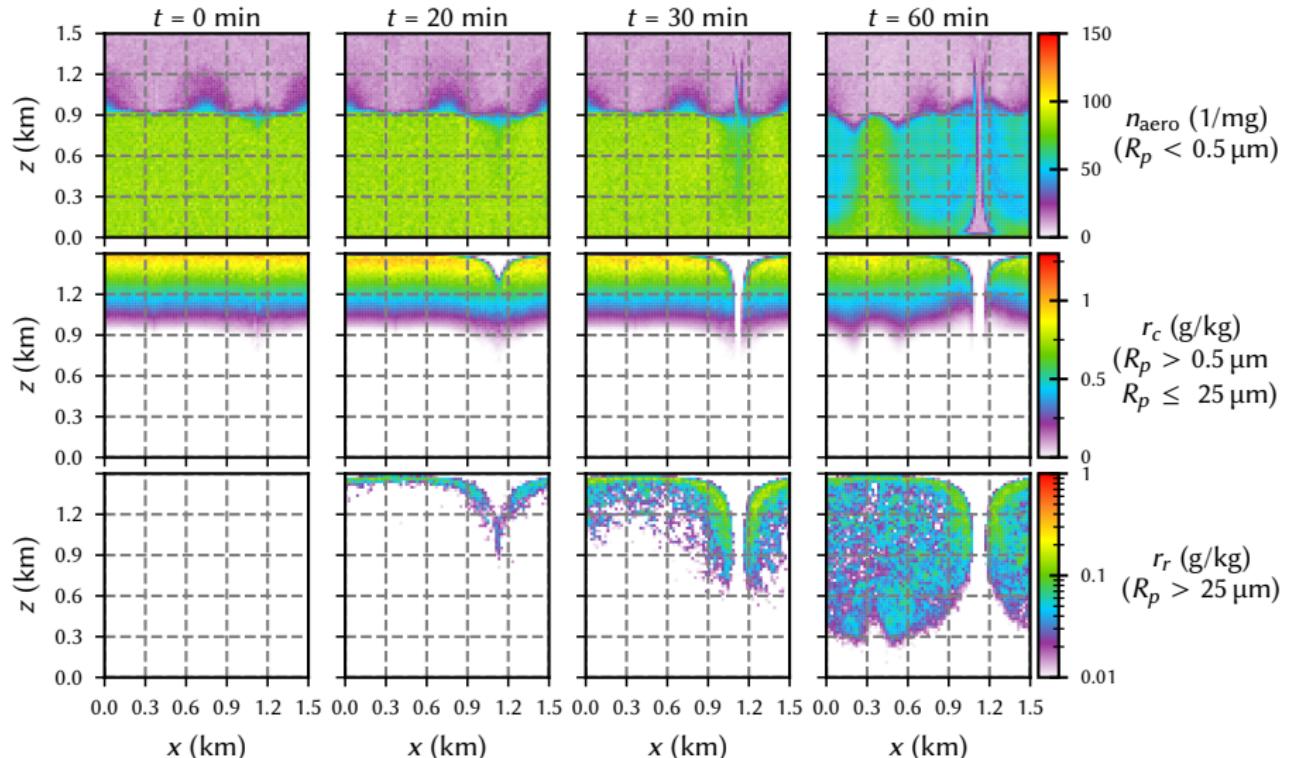


Default parameters

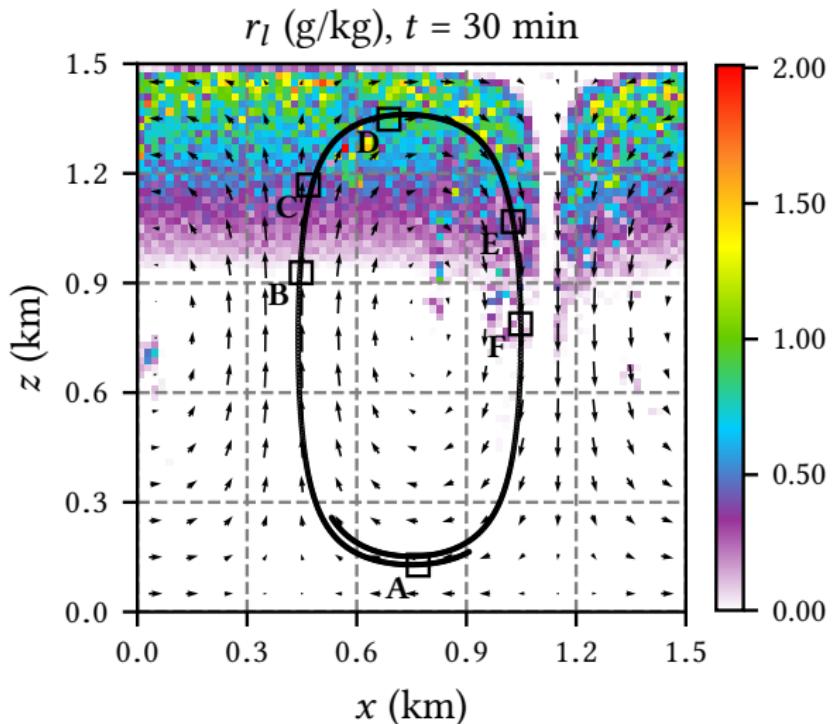
- CCN material: Ammonium sulfate
- $\text{DNC}_1 + \text{DNC}_2 = 60 \text{ cm}^{-3} + 40 \text{ cm}^{-3} = 100 \text{ cm}^{-3}$
- 64 super-droplets per cell
- 75×75 grid cells of $20 \text{ m} \times 20 \text{ m}$
- Time steps: $\Delta t_{\text{adv}} = 1 \text{ s}$, $\Delta t_{\text{col}} = 0.5 \text{ s}$, $h = 0.1 \text{ s}$
- Hydrodynamic Long collision kernel
- 2 hour spin-up phase
- **Results are averages** over 50 independent simulations

Test case: Cloud evolution

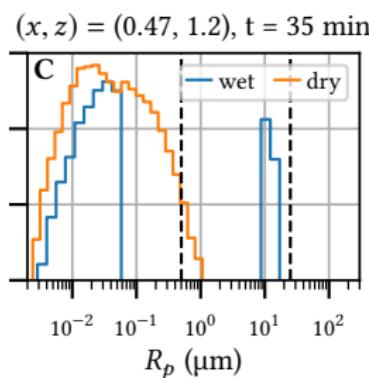
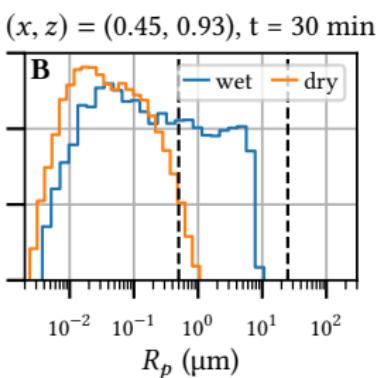
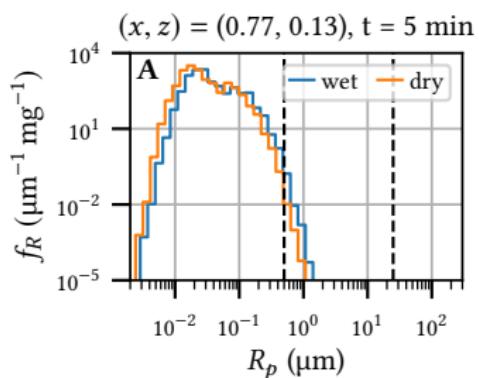
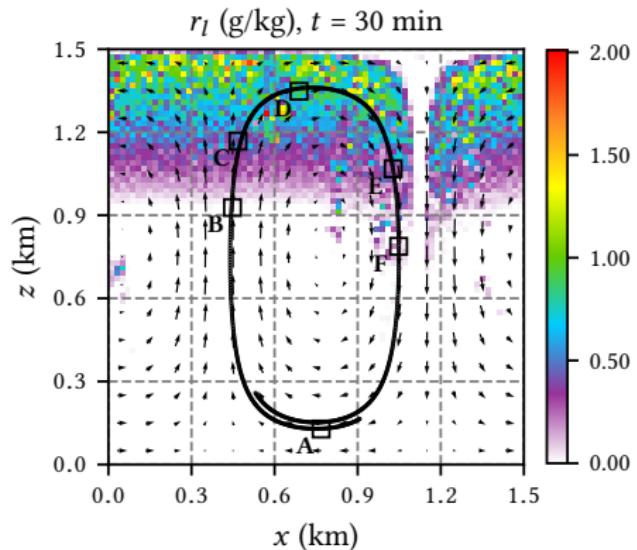
- Ammonium sulfate CCNs, Long kernel, average over 50 simulation runs



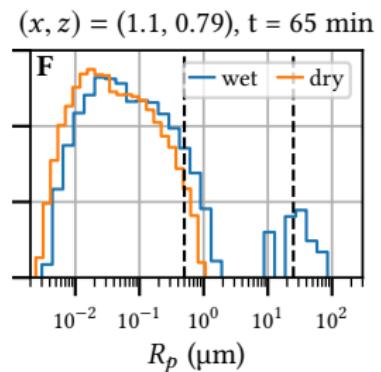
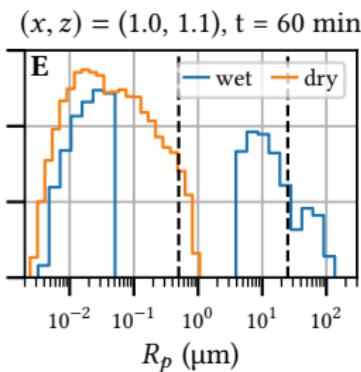
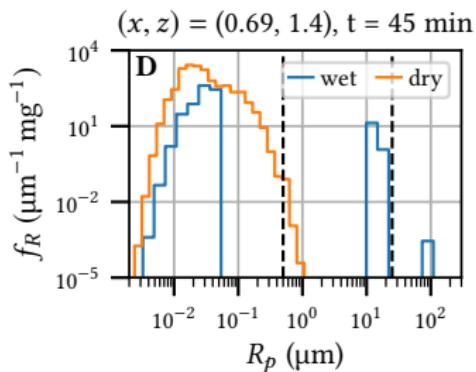
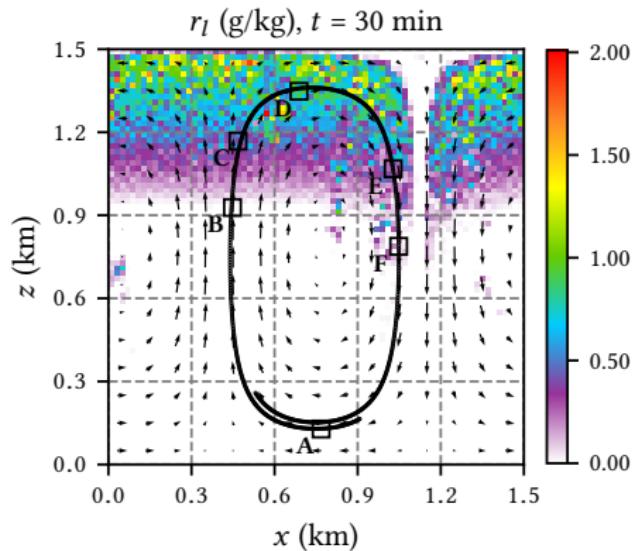
Droplet growth cycle



Droplet growth cycle

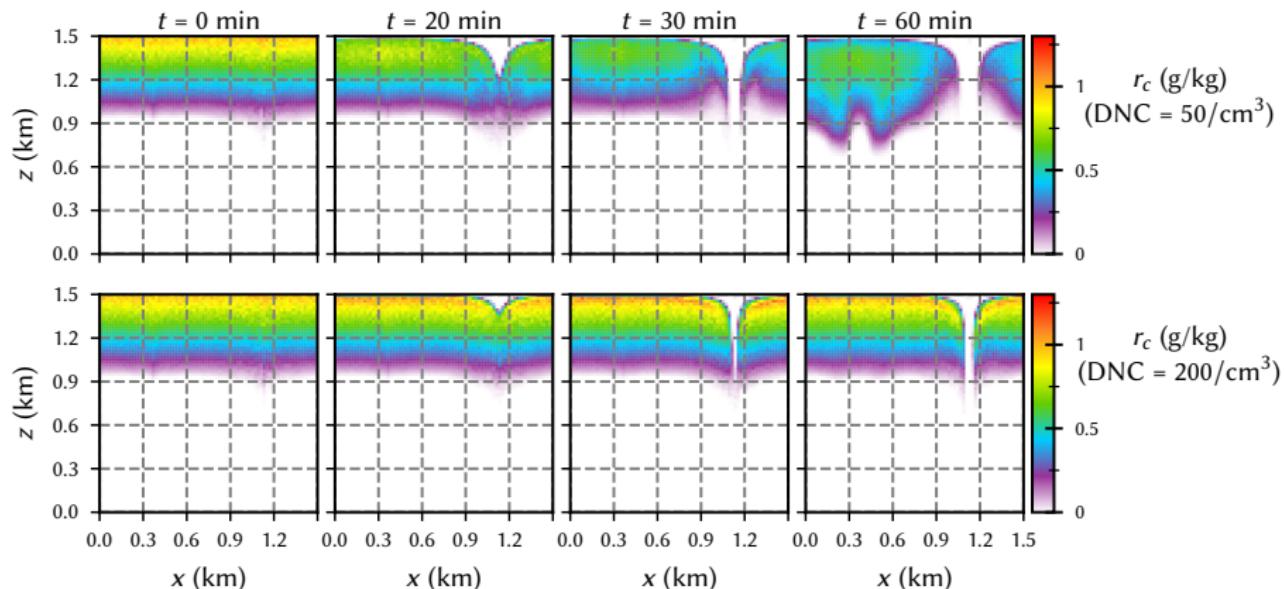


Droplet growth cycle



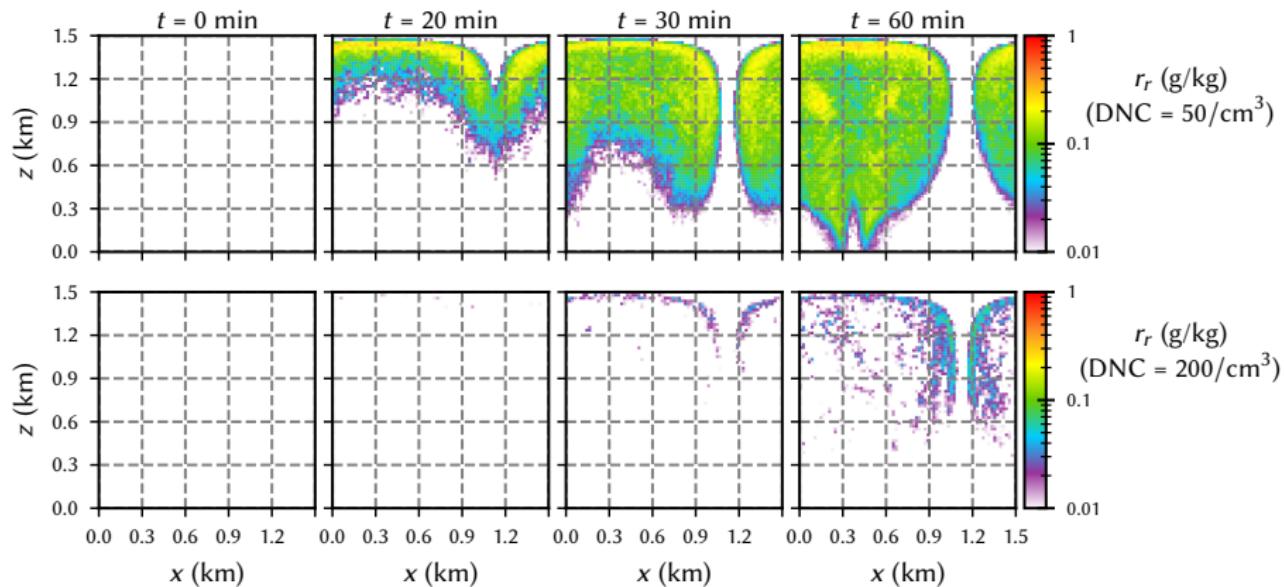
Variation of the CCN number concentration

- Cloud water mixing ratio r_c ($0.5 \mu\text{m} < R_p \leq 25 \mu\text{m}$)



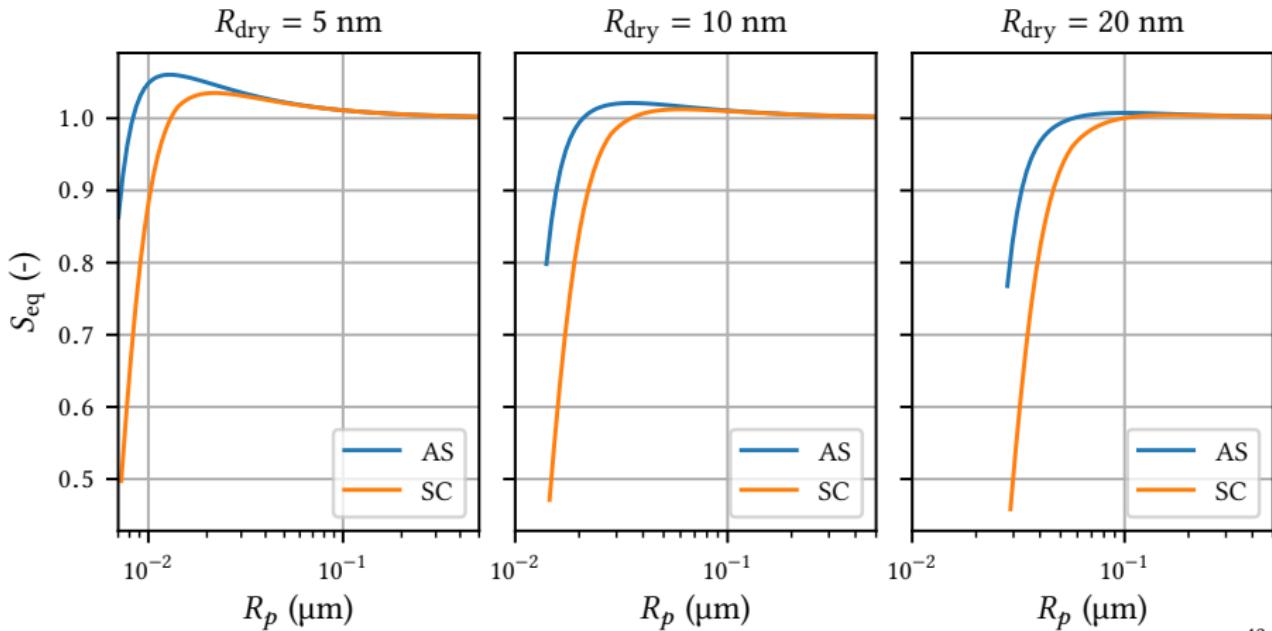
Variation of the CCN number concentration

- Rain water mixing ratio r_r ($R_p > 25 \mu\text{m}$)



Variation of CCN material

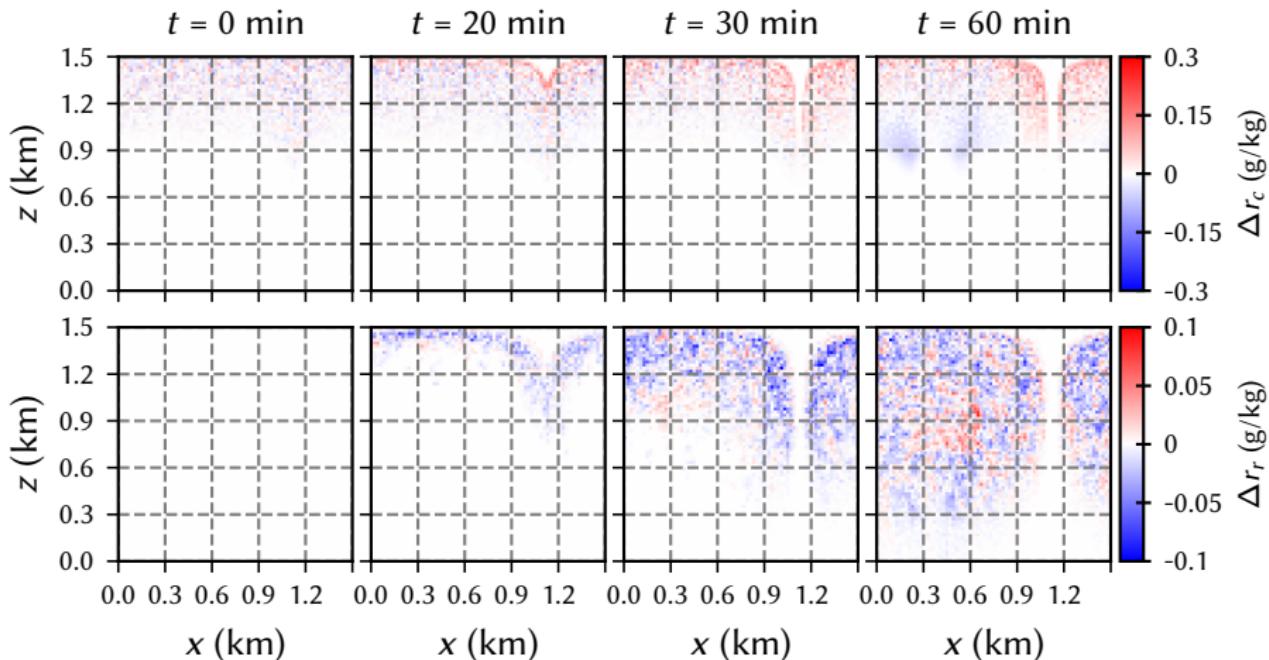
- Equilibrium saturations for
 - **Ammonium sulfate (AS, blue)**
 - **Sodium chloride (SC, orange)**



Variation of CCN material

- Difference for **sodium chloride** (SC) and **ammonium sulfate** (AS):

$$\Delta r = r_{\text{SC}} - r_{\text{AS}}$$



Atmospheric model

Particle model

Discretization

Collision test in box model

Test case: Drizzling stratocumulus

Conclusions & Outlook

Conclusions

- Stable, extendable program
- Agreement of collision box model simulations with most recent results of [Unterstrasser, Hoffmann & Lerch (2017), Geosci. Model Dev. **10**, 1521]
- Results of 2D test case comparable to [Arabas *et al.* (2015), Geosci. Model Dev. **8**, 1677]
- Plausible behavior when varying CCN concentration and material

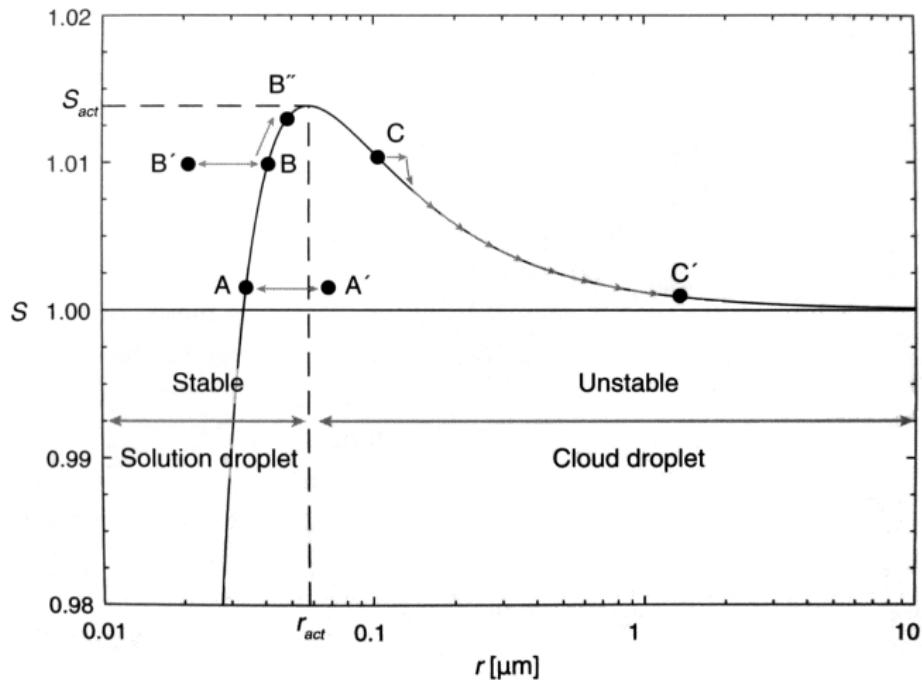
Outlook: Possible extensions

- Parallelization for speed
- 3D model coupled to computational fluid dynamics solver
 - see e.g. [Pressel *et al.* (2015), *J. Adv. Model. Earth Sy.* **7**, 1425]
- Multicomponent chemistry
 - see e.g. [Jaruga & Pawlowska (2018) *Geosci. Model Dev.* **11**, 3623]
- Ice crystal systems
 - see e.g. [Söhlch & Kärcher (2010), *Q. J. R. Meteorol. Soc.* **136**, 2074]
- Local turbulence treatment by stochastic velocity perturbations
 - see e.g. [Grabowski & Abade (2017), *J. Atmos. Sci.* **74**, 1485]

Thank you for your attention!

Any questions?

Equilibrium saturation: Köhler curve



Lohmann, Lüönd, Mahrt:
An introduction to clouds,
 Cambridge Univ. Press 2016

$$S_{eq} = \frac{e_s(T, R_p, m_s)}{e_s(T, R_p = \infty, m_s = 0)} = \frac{m_w}{m_w + m_s i_s M_w / M_s} \exp\left(\frac{a_K}{R_p}\right)$$