Ecosystem modeling

VI. Chemical kinetics and carbonate system modeling

Jang-Geun Choi

Center for Ocean Engineering University of New Hampshire

July, 2025

Modeling chemical reaction (kinetics)

Let us consider an idealized chemical reaction between two elements, A and B, whose compound is AB_2 :

$$A + 2B \longrightarrow AB_2.$$
 (1)

The rate law

$$\frac{dC}{dt} \propto [A]^m [B]^n \tag{2}$$

where C = [A], [B], or $[AB_2]$ (representing the molar concentration of A, B, or AB₂, respectively).

For example, we can set

$$\frac{d[A]}{dt} = -k[A]^m[B]^n \tag{3}$$

where k is the proportional coefficient.



Modeling chemical reaction (kinetics)

$$A + 2B \xrightarrow{k} AB_2$$

$$\frac{d[A]}{dt} = -k[A]^m[B]^n \qquad \text{(4a)}$$

$$\frac{d[B]}{dt} = -2k[A]^m[B]^n \qquad \text{(4b)}$$

$$\frac{d[AB_2]}{dt} = k[A]^m[B]^n \qquad \text{(4c)}$$

where unit of $[AB_2]$ is molar concentration of A (number of element A).

The mass conservation is satisfied by

$$\frac{d}{dt}\left(\left[A\right] + \left[AB_2\right]\right) = 0 \quad \text{(5a)}$$

$$\frac{d}{dt}\left([B] + 2[AB_2]\right) = 0 \quad \text{(5b)}$$

where A_T and B_T are total numbers of element A and B in the system.

Modeling chemical reaction (kinetics)

$$1A + 2B \longrightarrow AB_2$$

$$\frac{d[A]}{dt} = -k[A]^{\mathbf{1}}[B]^2 \qquad (6a)$$

$$\frac{d[B]}{dt} = -2k[A]^{1}[B]^{2}$$
 (6b)

$$\frac{d[AB_2]}{dt} = k[A]^{1}[B]^{2}$$
 (6c)

where unit of $[AB_2]$ is molar concentration of A (number of element A).

The orders, m and n, need to be determined by experimentally, but good a priori is the "molecularity":

$$m=1$$
 and $n=2$,

but is not always true, especially, when the reaction is not elementary (single-step).

Reversible reaction

Let us allow inverse reaction:

$$A + 2B \underset{k_2}{\overset{k_1}{\rightleftharpoons}} AB_2 \tag{7}$$

that can be formulated by

Forward reaction
$$(A + 2B \xrightarrow{k_1} AB_2)$$

$$\frac{d[A]}{dt} = -k_1[A][B]^2 + k_2[AB_2] \qquad (8a)$$

$$\frac{d[B]}{dt} = -2k_1[A][B]^2 + 2k_2[AB_2] \qquad (8b)$$

$$\frac{d[AB_2]}{dt} = k_1[A][B]^2 - k_2[AB_2] \qquad (8c)$$
Inverse reaction $(AB_2 \xrightarrow{k_2} A + 2B)$

Reversible reaction

Once we assume steady-state,

$$\frac{d[A]}{dt} = -k_1[A][B]^2 + k_2[AB_2]$$
 (9a)

$$\frac{d[B]}{dt}^{0} = -2k_1[A][B]^2 + 2k_2[AB_2]$$
 (9b)

$$\frac{d[AB_2]}{dt} = k_1[A][B]^2 - k_2[AB_2]$$
 (9c)

that provide only one equation $k_1[A][B]^2 - k_2[AB_2] = 0$. This can be rewritten as

Equilibrium constant

$$K = \frac{[AB_2]}{[A][B]^2} \tag{10}$$

where $K = k_1/k_2$ referred to as the "equilibrium constant".

$$CO_2 + H_2O \rightleftharpoons HCO_3^- + H^+$$
 (11a)

Not considered in formulation because it it solvent. $HCO_3^{-} = \frac{k_3}{k_4} CO_3^{2-} + H^+$ (11b)

$$\frac{d[\text{CO}_2^{0}]}{dt} = -k_1[\text{CO}_2] + k_2[\text{HCO}_3^{-}][\text{H}^+]$$
 (12a)

$$\frac{d[HCO_3^{1-}]}{dt} = k_1[CO_2] - k_2[HCO_3^{-}][H^+]
- k_3[HCO_3^{-}] + k_4[CO_3^{2-}][H^+]$$
(12b)

$$\frac{d[\text{CO}_3^{2^-}]}{dt} = k_3[\text{HCO}_3^-] - k_4[\text{CO}_3^{2^-}][\text{H}^+]$$
 (12c)

$$\frac{d[\mathbf{H}^{1+}]}{dt} = k_1[\mathbf{CO}_2] - k_2[\mathbf{HCO}_3^-][\mathbf{H}^+] + k_3[\mathbf{HCO}_3^-] - k_4[\mathbf{CO}_3^{2-}][\mathbf{H}^+]$$
(12d)

Two conservative quantities in (12): carbon and electron (charge)

Summing (12a), (12b), and (12c) yields the carbon conservation equation given by

$$\frac{d}{dt} \left([CO_2] + [HCO_3^-] + [CO_3^{2-}] \right) = 0$$

$$\therefore [CO_2] + [HCO_3^-] + [CO_3^{2-}] = C_T$$
(13)

Total (inorganic) carbon (a.k.a. dissolved inorganic carbon)

Weighted summation by the charges¹ yields the electron conservation equation given by

$$\frac{d}{dt} \left([HCO_3^-] + 2[CO_3^{2-}] - [H^+] \right) = 0$$

$$\therefore [HCO_3^-] + 2[CO_3^{2-}] - [H^+] = A_T$$
Carbonate alkalinity
$$\uparrow$$
Total alkalinity

 $^{^{1}0\}times(12a)+1\times(12b)+2\times(12c)-1\times(12d)$

Assuming steady-state in (12a) and (12c) yields

$$0 = -k_1[CO_2] + k_2[HCO_3^-][H^+]$$

$$\therefore K_1 = \frac{[HCO_3^-][H^+]}{[CO_2]}$$
(15)

$$0 = k_3[HCO_3^-] - k_4[CO_3^{2-}][H^+]$$

$$\therefore K_2 = \frac{[CO_3^{2-}][H^+]}{[HCO_3^-]}$$
(16)

where $K_1 = k_1/k_2$ and $K_2 = k_3/k_4$, representing equilibrium constants for each reaction.

Now, we have four equations, (13) to (16), for four variables ($[CO_2]$, $[HCO_3^-]$, $[CO_3^{2-}]$, and $[H^+]$). This is soluble.

Note that this is approximated. In the real ocean, there are more reactions associated with non-carbon:

$$CO_{2} + H_{2}O \rightleftharpoons \stackrel{k_{1}}{\rightleftharpoons} HCO_{3}^{-} + H^{+}$$

$$HCO_{3}^{-} \rightleftharpoons \stackrel{k_{3}}{\rightleftharpoons} CO_{3}^{2-} + H^{+}$$

$$H_{2}O \rightleftharpoons H^{+} + OH^{-}$$

$$B(OH)_{3} + H_{2}O \rightleftharpoons H^{+} + B(OH)_{4}^{-}$$

$$\vdots$$

that plays a role changes the charge conservation equation².

$$[HCO_3^-] + 2[CO_3^{2-}] - [H^+] + [B(OH)_4^-] + [OH^-] \cdots = A_T$$
 (17)

²See Humphreys et al. (2022) for sophisticated carbonate_equilibrium_model.

Coupled carbonate system with ecosystem

An ecosystem model can be coupled with the carbonate equilibrium model by simulating dissolved inorganic carbon C_T and total alkalinity A_T (Fennel et al., 2008).

Note that C_T is subject of photosynthesis and respiration:

$$106 \, \mathrm{CO_2} + 16 \, \mathrm{NO_3}^- + \mathrm{HPO_4}^{2-} + 72 \, \mathrm{H_2O} + 18 \, \mathrm{H}^+$$

$$\xrightarrow{\mathrm{Photosynthesis}} \mathrm{C_{106}H_{175}O_{42}N_{16}P_1} + 150 \, \mathrm{O_2}$$
that can be formulated by
$$\frac{\partial C_T}{\partial t} + \nabla \cdot (\vec{u}C_T) = \nabla \cdot (\vec{K}\nabla C_T) - \mathcal{P} + \mathcal{R} \qquad (19)$$

$$\xrightarrow{\mathrm{Air-sea\ interaction\ (gas\ exchange)}} K_z \frac{\partial C_T}{\partial z} \bigg|_{z=0} = k_C([\mathrm{CO_2}]^* - [\mathrm{CO_2}]) \qquad (20)$$

Coupled carbonate system with ecosystem

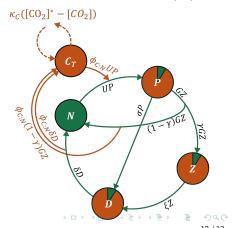
$$\frac{106_{5} \, \text{CO}_{2} + 16 \, \text{NO}_{3}^{-} + \text{HPO}_{4}^{2-} + 72 \, \text{H}_{2}\text{O} + 18 \, \text{H}^{+} \quad \text{Organic matter}}{\stackrel{\text{Photosynthesis}}{\longleftarrow}{\text{Respiration}}} \underbrace{\begin{array}{c} \text{C}_{106} \text{H}_{175} \text{O}_{42} \text{N}_{16} \text{P}_{1} \\ \text{T}_{106} \text{P}_{106} \text$$

In case of NPZD model, ${\mathcal P}$ and ${\mathcal R}$ can be formulated by

$$\mathcal{P} = \phi_{C:N} U P \tag{22}$$

$$\mathcal{R} = \phi_{C:N} \left(\delta D + (1 - \gamma)GZ \right) \tag{23}$$

where $\phi_{C:N} = 106/16$. Total alkalinity A_T can be formulated by same manner or can be prescribed by function of salinity.



References I

- Fennel, Katja et al. (2008). "Denitrification effects on air-sea CO2 flux in the coastal ocean: Simulations for the northwest North Atlantic". In: Geophysical Research Letters 35.24.
- Humphreys, Matthew P et al. (2022). "PyCO2SYS v1. 8:
 Marine carbonate system calculations in Python". In:

 Geoscientific Model Development 15.1, pp. 15–43.