

# Differential Equations in Geophysical Fluid Dynamics

## XIV. Wind-driven circulation: approximated solution to Stommel wind-driven circulation problem

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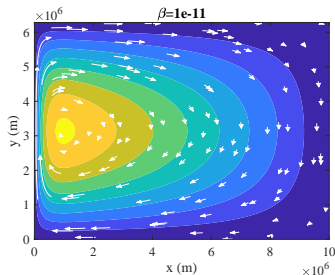
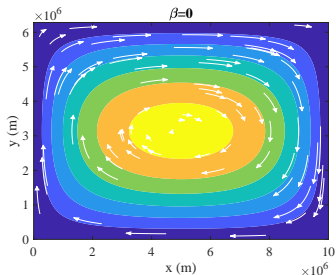
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# Recap

Stommel's wind-driven circulation problem is given by

$$\underbrace{\frac{\gamma}{h} \left( \frac{\partial^2 \psi}{\partial x^2} + \frac{\partial^2 \psi}{\partial y^2} \right)}_{\beta \bar{v}: \text{Planetary } \beta\text{-term}} + \underbrace{\beta \frac{\partial \psi}{\partial x}}_{-(\gamma/h) \nabla \times \vec{u}: \text{Bottom stress curl}} = \underbrace{-\frac{\tau_0 \pi}{\rho_0 h L_y} \sin(\pi y / L_y)}_{\text{Wind stress curl}} \quad (1a)$$

$$\psi|_{x=0} = 0, \quad \psi|_{x=L_x} = 0, \quad \psi|_{y=0} = 0, \quad \psi|_{y=L_y} = 0. \quad (1b)$$

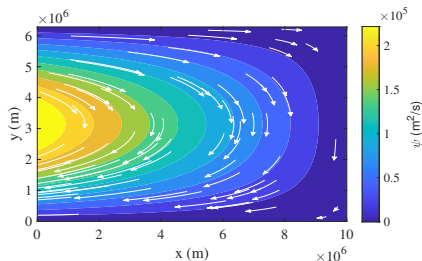


## Scale analysis

Let us nondimensionalize the governing equation using  $x = Lx^*$ ,  $y = Ly^*$ , and  $\psi = \Psi\psi^*$ :

$$\underbrace{\epsilon \left( \frac{\partial^2 \psi^*}{\partial x^{*2}} + \frac{\partial^2 \psi^*}{\partial y^{*2}} \right)}_{O(\epsilon): \text{ small}} + \underbrace{\frac{\partial \psi^*}{\partial x^*}}_{O(1): \text{ big "Sverdrup balance"}} = \nabla \times \vec{\tau}_s^* \quad (2)$$

where  $\epsilon = L_S/L \ll 1$  and  $L_S = \gamma/(h\beta)$ . The order of the forcing term is set to be identical to that of the beta term.



# Scale analysis

What if we set  $x = L_S x^*$  representing narrow western boundary region?

$$\underbrace{\frac{\partial^2 \psi^*}{\partial x^{*2}}}_{O(\epsilon^2): \text{ very small}} + \underbrace{\epsilon^2 \frac{\partial^2 \psi^*}{\partial y^{*2}}}_{O(1): \text{ big}} + \underbrace{\frac{\partial \psi^*}{\partial x^*}}_{O(\epsilon): \text{ small}} = \epsilon \nabla \times \vec{\tau}_s^* \quad (3)$$

So, for the narrow western boundary, the governing equation can be simplified to

$$\frac{\partial^2 \psi}{\partial x^{*2}} + \frac{\partial \psi^*}{\partial x^*} = 0 \quad (4)$$
$$\left( \frac{\gamma}{h} \frac{\partial^2 \psi}{\partial x^2} + \beta \frac{\partial \psi}{\partial x} = 0 \right)$$

## Simplified problem

Consider two components of the stream function: interior component ( $\psi_I$ ) and boundary component ( $\psi_B$ ), so  $\psi = \psi_I + \psi_B$ .

Based on scaling analysis above, simplified governing equations for each component are given by

$$\beta \frac{\partial \psi_I}{\partial x} = -\frac{\tau_0 \pi}{\rho_0 h L_y} \sin(\pi y / L_y) \quad (5)$$

$$\frac{\gamma}{h} \frac{\partial^2 \psi_B}{\partial x^2} + \beta \frac{\partial \psi_B}{\partial x} = 0 \quad (6)$$

and the boundary conditions, also simplified, are given by

$$\psi|_{x=0} = 0 \quad \lim_{x \rightarrow \infty} \psi = \psi_I. \quad (7)$$

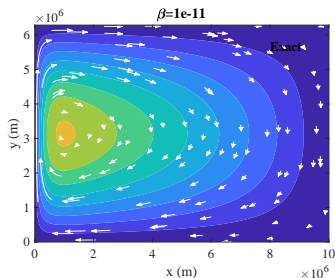
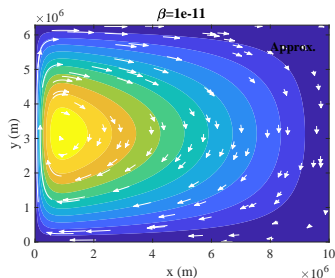
# Simplified problem

Solution to the simplified problem is given by

$$\psi = \psi_I + \psi_B \quad (8a)$$

$$\psi_I = \left( -\frac{\tau_0 \pi}{\rho_0 L_y h \beta} \sin(\pi y / L_y) \right) (x - L_x) \quad (8b)$$

$$\psi_B = -\psi_I e^{x/L_s} \quad (8c)$$



## Summary

For the narrow western boundary region, governing equation is simplified to

$$\frac{\gamma}{h} \frac{\partial^2 \psi_B}{\partial x^2} + \beta \frac{\partial \psi_B}{\partial x} = 0 \quad (9a)$$

$$\psi_B|_{x=0} = -\psi_I \quad (9b)$$

$$\lim_{x \rightarrow \infty} \psi_B = 0 \quad (9c)$$

This implies that the boundary current is forced by interior flow, rather than regional wind stress forcing.

Solution to the problem is given by

$$\psi_B = -\psi_I e^{x/L_S} \quad (10)$$

where  $L_S = \gamma/(h\beta)$  representing scale of width of western boundary current.