# Read Me Richards Module PCR-GlobWB2.0

Bright Minds Project

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# **General**

In this document, the modules that I made for the Bright Minds Project and that could be added to the PCRGlobWB2.0 hydrological model are explained. The work was done under supervision of Niko Wanders at the Physical Geography Department of Utrecht University. I will focus on the main structure and provide an overview of the assumptions that were made.

New modules which were added and which will be discussed in this document include:

- Unsaturated Zone: Richards Equation
   Solves the Richards Equation for the transport of water in the unsaturated zone.
- Unsaturated Zone: Heat Function

  Uses advection and diffusion to compute the soil temperature evolution in the unsaturated zone.
- Surface: Energy Balance

Computes the surface energy balance and uses it to determine the skin temperature, which is used as an input for the heat function.

Files which I have adapted include setup\_30min\_seperatefile.ini , landCover.py , landSurface.py , reporting.py , variable\_list.py and meteo.py . Note that all parts that are added are bounded by #%% ADDED BY JOREN: START and #%% ADDED BY JOREN: STOP .

The main file which has been added is richards.py, which includes all relevant functions.

The module is built in Python. Necessary libraries include:

- numpy
- scipy
- multiprocessing
- pcraster

Therefore, in pcrglobwb\_py3.yml (in conda\_env), scipy is added.

# **Updates of Pre-Existing Files**

## setup\_30min\_richards.ini

In this file the input for the model is provided. This part focusses on the changes with respect to the normal *.ini*-file.

Under [landSurfaceOptions] , there are new entries:

- includeRichards = True
  - Boolean: True if you want to include the Richards model (so far only possible if number of soil layers is 2). Model should still be able to run normally if this is not included or set to False.
- layerFactorRichards = 2
  - Integer: number of layers into which the upper and lower soil layer of PCR-GlobWB should be divided. If layerBoundariesRichards is used,

# layerFactorRichards needs to be set to 0.

- layerBoundariesRichards = [0, 0.10, 0.30, 0.40, 1.0]
  - Array [m]; Start with 0 (surface), only one layer between 0.13-0.30m and one between 0.65-1.5m. These ranges are the ranges of thicknesses of the 2 layer unsaturated soil in PCR-GlobWB and make sure that it is easy to switch between the two layer main model and my flexible layer model. **Not used if**

## layerFactorRichards is not 0!

- numberOfCoresRichards = 8
  - Integer: Number of cores you want to run the model on. The Richards model is quite slow and in order to have faster computation times, more cores are recommended. However, 1 core is also possible.
- timestepRichards = 6
  - Integer [hours]: By default, the surface energy balance is computed every 3 hours. The Richards model can be computed in 3,6,12 or 24 hours, in order to increase or decrease the temporal resolution if desired.

Furthermore, under [meteoOptions] new paths for netCDF-files to read are added:

- cloudcoverNC = ../cru\_ts3.21.cld.dat.nc netCDF file containing cloudcover fractions over the basin.
- radiationNC = ../short\_wave\_radiation\_monthly\_climatology\_30arcmin.nc netCDF file containing incoming short wave radiation over the basin.
- vaporNC = ../cru\_ts3.21.vap.dat.nc
   netCDF file containing near surface vapor pressure over the basin.
- sunhoursTable = ../sunhoursfrac.tbl

Table containing the fractions of relative sunshine hours given cloud cover.

Note that the model should still be able to run if these are not provided, as long as includeRichards = False or not mentioned in the file.

## meteo.py

In this file the meteorological input for the model is provided. This part focusses on the changes with respect to the normal meteo.py.

- In the \_\_init\_\_ function, lines are added to read relevant meteo iniItems from the *ini*-file.
- The readExtensiveMeteo() function is included. This reads the relevant extra meteorological input for the correct timestep.

## landCover.py

In this file the landsurface input for the model is provided. This part focusses on the changes with respect to the normal landCover.py .

- At the top, the richards.py file is imported as richards.
- In the \_\_init\_\_ function, lines are added to read relevant land surface iniItems from the ini-file.\
- In the getICsLC -functions, initialiseRichards() is called to initialise the model
- updateSoilStates():
  - The function definition is changes: (self, \*args). This is done to include the meteo object and to include the current time step.
  - If number of soil layers is 2 and the includeRichard = True, the runRichards() is called.

## landSurface.py

This part focusses on the changes with respect to the normal landSurface.py.

• In the \_\_init\_\_ function, lines are added for reporting new variables as netCDF files.

# variable\_list.py

- Extra variables are created for reporting. These new variables are:
  - soilTempUpp: [K] soil temperature of the upper cell in the main model.
  - soilTempLow: [K] soil temperature of the lower cell in the main model.
  - tempDeficit\_6AM: [K] difference between skin and air temperature at 6AM.
  - tempDeficit\_6PM: [K] difference between skin and air temperature at 6PM.
  - $\circ$  netSW: [W/m^2] net short wave radiation.
  - $\circ$   $\mbox{netRad}$  : [W/m^2] net incoming radiation (both short and long wave).
  - longWaveRad: [W/m^2] outgoing long wave radiation.
  - latentHF: [W/m^2] latent heat flux.
  - sensibleHF: [W/m^2] sensible heat flux.
  - groundHF: [W/m^2] ground heat flux.

# reporting.py

• In basic\_post\_processing(), new variables are added for postprocessing.

# The New richards.py File

#### **General**

richards.py is mainly built in *numpy* and not in *pcraster* as the rest of the code. For the most important variables, the reader is referred to the variable list at the end of this document. richards.py consists of multiple seperate functions. The main ones are:

# • initialiseRichards()

Defines all relevant variables which are used in the rest of the model. It starts by converting all relevant variables from *pcraster* to *numpy* and flatten them, while remembering the original order.\

- In the section MODEL DIMENSIONS, the spatial and temporal resolution is defined. In general, units are [m] and [day] if not specified. Note that index [0] means the lower most layer and index [-1] is the top layer.
- In the section *SOIL CONDITIONS*, the relevant soil conditions are defined. It is tried to retrieve most from the PCRGlobWB model, but for variables not defined, conditions are assumed. Van Genuchten parameters are assumed in the function GuelphLoamDrying(). Saturated and residual soil moisture content is taken from the model.
- In the section *VEGETATION TYPES*, vegetation types can be defined. For now, only the albedo is relevant.
- In the section *CONSTANTS*, the relevant constants are defined. The saturated and dry thermal conductivity are now the same for every cell, but this should be updated in a later version.
- In the section *INITIAL VALUES AND BOUNDARY CONDITIONS*, the initial conditions for the soil temperature, richards equation and energy balance are defined.

# • runRichards()

Runs the actual model.

- In the section CONVERSION, the relevant input is converted to numpy.
- INITIALISATION: Data is flattened, empty arrays are defined. Some cells in the model have nan values, since they might fall outside the basin. The model does not work if these are incorporated. Therefore, nanindexUpp is defined to eliminated these. However, the main model still uses these cells and therefore nanindex is used to insert nanvalues in these cells in the POSTPROCESSING section.
- RUN THE MODEL: The actual model is run and it exists of two loops: one for the energy balance (predefined to run in steps of 3 hours) and one for the Richards model in larger timesteps. More details on how the energy balance, Richards model and soil heat function work in their specific sections.
- POSTPROCESSING: The water balance is closed by computing the outgoing percolation. Note that it is first negative, since that is how the model is defined, but later made positive for the PCR-GlobWB model. Futhermore, NaN-values are inserted, the arrays are shaped into the proper dimensions and sizes. In the end, relevant variables are coupled to the landCover object and send to the main model.
- In the section *CONVERSION TO PCRASTER*, the relevant output is converted to *pcraster*.

The smaller functions will be explained in the sections on the Richards model, heat function and energy balance.

#### Wishlist:

- Make the model usable for the 3 layer configuration in PCR-GlobWB.
- Make the multiple layers completely flexible (so that the boundaries of the cells do not have to be included in the *ini*-file).

#### Unsaturated Zone: Richards Model

This model solves the  $\underline{Richards\ equation}$ . It was based on a model on  $\underline{GitHub\ which\ can}$  be found  $\underline{here}$ .

In runRichards(), the main model is the RichardsModel\_dz2(), which is solved with scipy.odeint(). In the case that multiple cores are provided in the *ini*-file, this is done by calling solve() per cell by using functions from the multiprocessing library. If only one core is provided, scipy.odeint() is called directly for all cells at the same time (this is very slow). scipy.odeint() needs high allowed maximum number of steps in order to stabilize the model (now set on mxstep=3500).

RichardsModel\_dz2() uses the matrix potential psi0 at time t[0] (in the function called psi) to compute the fluxes q. Note that the z-axis is positive upwards (all downward fluxes are negative). There convergence is then used to provide the time derivative of psi to compute the following timestep. As boundary conditions, potentials or fluxes at the top (psiTop and qTop) or at the bottom (psiBot and qBot). The model now uses qTop as infiltration and has free drainage at the bottom. The model uses the Van Genuchten parameters to compute the hydraulic conductivity (K from KFun1()) and d\$\theta\$/d\$\psi\$ (C from CFun()). Hydraulic conductivity values, and therefore fluxes are averaged between two adjacent cells to compute the correct flux between them. Furthermore, no flow is allowed if the soil temperature is below 0 degrees Kelvin, so then the hydraulic conductivity is set to 0 for that layer to assume freezing.

In the end the new matrix potential psi is returned to runRichards() for time t[-1]. With thetaFun(), this can be converted to water content.

fluxModel2() is basically the same model as RichardsModel\_dz2(), but only computes and returns the fluxes at one moment. Note that we therefore only have the fluxes at the beginning and end of each run of RichardsModel\_dz2(). Therefore, we do not exactly know how much water has left the soil, and this can only be assessed at the end by assuming the water balance to be closed per definition and computing how much has left the model.

# Main Assumptions:

- Water Balance is closed by definition. Therefore, all water that is not in any other variable, is put in percolation.
- Matrix potential psi can be larger than 0. However, if this happens and the is returned to a water content, the rest of the water is converted to either interflow (bottom layer) or saturation excess (top layer) after RichardsModel\_dz2 is run. Due to this the timestep has some influence on the amount of interflow and saturation excess from the model.
- No flow if soil is frozen.
- There is no interaction between the unsaturated zones of neighbouring cells.

#### Wishlist:

- If the surface is flooded, psiTop could be changed automatically to increase infiltration (instead of only having rain).
- Model is still quite slow and could be made faster in a future version.

## Unsaturated Zone: Heat Function

The main function for this model is heatFun3(). It solves for the soil temperature and also uses scipy.odeint() to compute the new state. However, since it stabilises much faster than RichardsModel\_dz2, there is no need for multiprocessing and the maximum allowed number of steps can be low (now set to 1000, but can be lower).

Based on the fraction of water in the soil the thermal conductivity (ThermalConduct()) and heat capacity (heatcapacity1()) is computed for the soil layers. Note that it is assumed that the soil part consists of dry sand when computing the heat capacity. Furthermore, if it freezes the heat capacity and thermal conductivity of ice is used for the water fraction, which is fixed as of now.

#### Main Assumptions:

- Top layer has the skin temperature. It does not change its soil temperature during the run of the heat function.
- Bottom layer soil temperature is fixed in the calculation. However, at the end, it receives the same dTdt as the layer just above it, in order to change through time.
- Advective fluxes are computed from the state at the end of the previous run of RichardsModel\_dz2. However, the real fluxes deviate probably a small amount from this.

## Wishlist:

- Soil temperature should be influenced by the energy released from freezing or melting.
- The soil temperature in the bottom layer should change by itself, instead of following the layer above it.
- Small correction could be done to the moisture content if it freezes, since ice has a different density. This could have some influence on the conductivity.

# Surface: Energy Balance

By using the Newton-Raphson method, the skin temperature is in such a way that the suface energy balance is closed. For this, scipy.optimize.newton() is used and the functions radiation\_f2() (computes how much energy is unacounted for in the energy balance) and the radiation\_deriv2() (computes the derivative with respect to the skin temperatures). This is very fast, so no multiprocessing is needed (now mxiter=150).

The actual fluxes are computed in the function ComputeFluxes2(). The latent heat flux is taken from the actual evapotranspiration computed by the original PCR GlobWB model itself (bare soil evaporation and soil transpiration). It is not changed during the solution of the energy balance. The incoming short wave radiation is taken from input netCDF files. Longwave radiation (longWave\_out2(): based on cloudfraction and relative humidity), the sensible heatflux (computeH()) and the ground heatflux are computed (GFlux()).

Fluxes are positive away from the surface. Only the incoming short wave radiation is positive towards the surface.

## **Main Assumptions**

- There is no melting or snow.
- Since latent heat is taken from the actual evapotranspiration in the PCR GlobWB model, there is no dependence on the other variables of the energy balance.
- GFlux() uses the full depth of the first layer to compare the skin temperature against (based on equations in the literature and fitted to the Cabauw data). It further assumes that the heatflux at that depth is neglible over the times step (dG/dz=(G(top)-0)/dz)
- The aerodynamic resistance ra in computeH() taken as constant.
- A fixed daily cycle is assumed: Tair\_list assumes deviations from the mean air temperature during the day and SW\_fractions devides the incoming solar radiation over the cells.
- Incoming shortwave radiation is not affected by cloud cover.

#### Wishlist:

- Time step size of the energy balance could be given in the ini-file.
- ComputeFluxes2() is now called double. It is very fast, so it does not take time but this could be made more efficient in a future model.
- Energy used for melting should be included in the energy balance.
- Snow layer albedo should be coupled in the energy balance. Now, always the albedo of grass is used.
- The energy balance could be decoupled from the Richards equation, so it can also be used in the normal model.
- Make the aerodynamic resistance ra a function of soil cover and wind.
- A larger variety and a higher resolution meteorological input could be used in the future.

## Variable list

List of the most important variables.

Variable Name	Unit	Description
initialiseRichards.py()		
self.basin_shape		Tuple: the dimension of the basin in PCR GlobWB2.0.
self.nCells		Integer:The number of cells in the original model.
dz	[m]	Array: containing the thickness of the layers in the model.
self.layerfractions	[m]	Array: to shift between the two layer model and the arbitrary layer Richards model.
self.layerFactorRichards		Integer: Number of layers in which the bottom layer is divided (the top layer as well if activated in <i>ini</i> -file).
n		Tuple: the number of layers ( $len(dz)$ ) in the vertical and the number of cells ( $self.nCells$ ).
dt	[hours]	Integer: number of hours between runs of the

		RichardsModel_dz2()
t	[days]	Array: times for which the RichardsModel_dz2() is evaluated in scipy.odeint().
timeunit	[s]	Integer: number of seconds per timestep of the RichardsModel_dz2().
p		Dictionary containing soil characteristics (Van Genuchten Parameters).
thetaS		Array: saturated moisture content.
thetaR		Array: residual moisture content.
theta_soil		Array: fraction of soil: 1-thetaS.
alpha		Array: surface albedo.
dT	[K]	Integer: temperature step; needed to compute the derivative in radiation_deriv2().
self.soiltemperature	[K]	Array: soil temperature.
qBot	[m/day]	Array: fixed bottom drainage. For free drainage, set it to None
psiBot	[m]	Array: fixed bottom matrix potential. For free drainage, set it to None
psiTop	[m]	Array: fixed upper matrix potential. For fixed infiltration, set it to None
self.storTotRich	[m]	Array: water content (without residual water!) for all layers in the Richards Model.
dt_short	[hours]	Integer: number of hours between runs of the energy balance
t_short	[days]	Array: times for which the energy balance is evaluated in scipy.optimize.newton().
timeunit_short	[s]	Integer: number of seconds per timestep of the energy balance.
SW_fractions	[]	Array: fractions of incoming short wave radiation. Assumed daily cycle.
Tair_list	[K]	Array: deviations of air temperature from the mean. Assumed daily cycle.
runRichards.py()		
cloudCover		Array: contains the cloudcover as a fraction.
vaporPressure	[kPa]	Array: actual vapor pressure.
S0	[W/m^2]	Array: incoming short wave radiation.
Tair	[K]	Array: air temperature.

		Array: indices where the storage array is NaN in the upper layer.
nanindexLow		Array: indices where the storage array is NaN in the lower layer.
nanindex		Array: indices where to insert nanvalues later (is different that nanindexUpp).
n1		Tuple: same as n but corrected for NaN values.
ra	[s/m]	Integer: aerodynamic resistance.
RH		Array: relative humidity.
plantevap	[m]	Array: total evaportranspiration per layer.
ЕТрт	[m]	Array: total evapotranspiration per cell.
theta0		Array: initial water content per day.
excess	[m]	Array: saturation excess and interflow together (from oversaturated layers).
oversaturated		Array: indices of oversaturated layers.
qТор	[m/day]	Array: infiltrating water.
Ts	[K]	Array: skin temperature.
Rn	[J/m^2]	Array: net radiation.
Н	[J/m^2]	Array: sensible heat.
G	[J/m^2]	Array: ground heat.
Sn	[J/m^2]	Array: short wave radiation.
Ln	[J/m^2]	Array: long wave radiation.
LE	[J/m^2]	Array: latent heat.
S0_3H	[J/m^2]	Array: as so but with daily cycle.
Tair_3H	[K]	Array: as Tair but with daily cycle.
ETpm_3H	[m]	Array: as ETpm but with daily cycle.
Т0	[K]	Array: intial soil temperature.
theta_run		Array: soil moisture for during the computation.
psi0	[m]	Array: matrix potential for during the computation.
theta_top_new	[m]	Array: soil moisture content for during the computation.
q0	[m/day]	Array: flux from psi0.
q	[m/day]	Array: flux from psi.

_mp		Multiprocessing
psi	[m]	Array: matrix potential for during the computation.
tempDeficit_6AM	[K]	Array: Difference between the skin temperature and the air temperature at 6AM.
tempDeficit_6PM	[K]	Array: Difference between the skin temperature and the air temperature at 6PM.
percTot	[m/day]	Array: total percolation per cell
percUpp	[m/day]	Array: percolation for upper cell (in PCR GlobWB)
percLow	[m/day]	Array: percolation for lower cell (in PCR GlobWB)
storTot_new	[m]	Array: the new self.storTotRich translated to the 2 cells of PCR GlobWB

# Contact

If anything is unclear and questions remain, please contact one of the following by email:

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Furthermore, if desired, multiple older versions of the code can be send as well (e.g. for data from the Raam network or Cabauw in the Netherlands, which are not coupled to the PCR-GlobWB model).