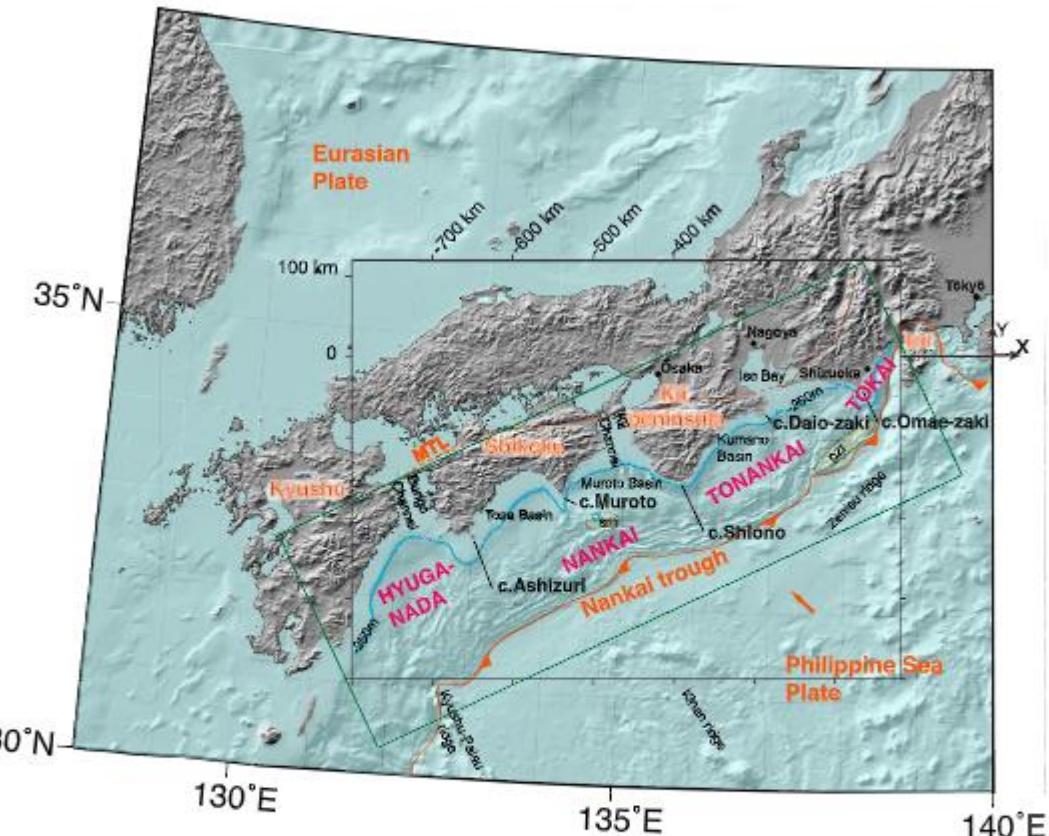


# 일본 난카이 지역에서 발생 가능한 대지진에 대한 합성 파형 시뮬레이션 및 한반도 지역 지진동에 미치는 영향에 관한 고찰

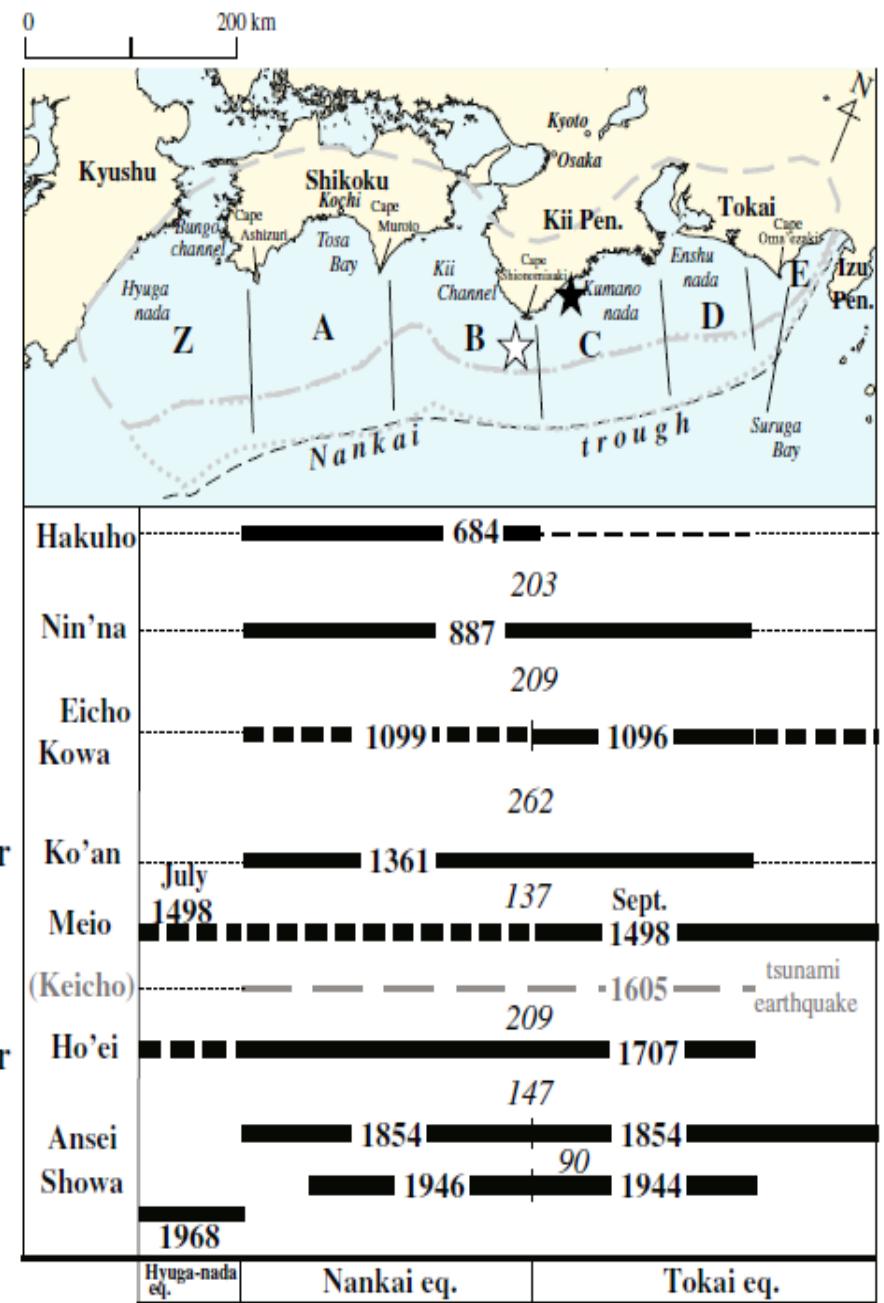
서울대학교 물리천문학부  
이재석

# Nankai Trough Earthquake



**Figure 1.** Topographic map of southwest Japan. The  $-250$  m isodepth in the shelf bank is underlined with thick blue-gray contour. Abbreviations ‘sm,’ ‘pzs’ and ‘c.’ stand for subducted seamount, subducted paleo-Zenisu ridge and cape, respectively. Black rectangle shows the computation area shown in the following, and the XY coordinate system used, whose origin is located at  $(35^{\circ}\text{N}, 140^{\circ}\text{E})$ . Rotated green rectangle shows the area covered by free surface elements.

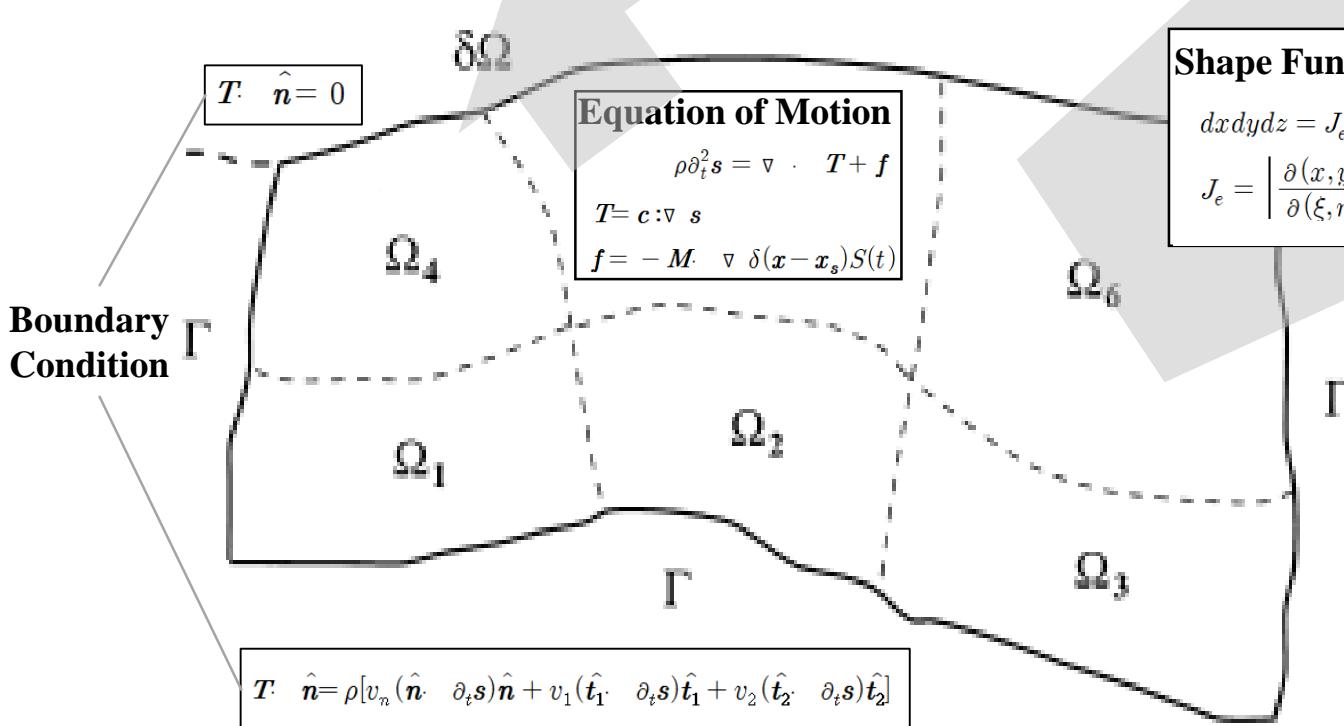
(Hok et al., 2011) / (Hyodo et al., 2014)



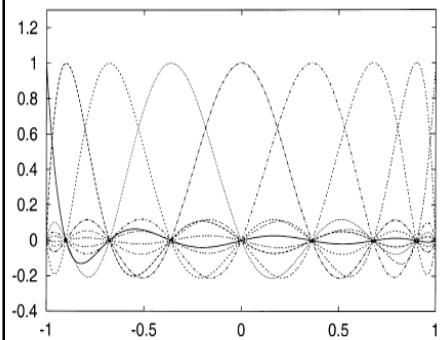
# Spectral Element Method (SEM)

## Weak formulation

$$\int_{\Omega} \rho \mathbf{w} \cdot \partial_t^2 \mathbf{s} d^3x = - \int_{\Omega} \nabla \cdot \mathbf{w} T d^3x + \mathbf{M} \nabla \cdot \mathbf{w}(\mathbf{x}_s) S(t) \\ + \int_{\Gamma} \rho [v_n (\hat{\mathbf{n}} \cdot \partial_t \mathbf{s}) \hat{\mathbf{n}} + v_1 (\hat{\mathbf{t}}_1 \cdot \partial_t \mathbf{s}) \hat{\mathbf{t}}_1 + v_2 (\hat{\mathbf{t}}_2 \cdot \partial_t \mathbf{s}) \hat{\mathbf{t}}_2] \cdot \mathbf{w} d^2x$$



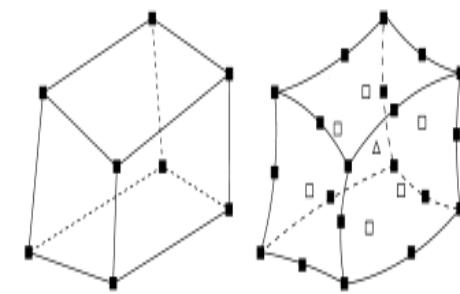
## Lagrange Polynomial Approximation



$$l_a^{nl}(\xi) = \frac{(\xi - \xi_0) \dots (\xi - \xi_{\alpha-1})(\xi - \xi_{\alpha+1}) \dots (\xi - \xi_{nl})}{(\xi_\alpha - \xi_0) \dots (\xi_\alpha - \xi_{\alpha-1})(\xi_\alpha - \xi_{\alpha+1}) \dots (\xi_\alpha - \xi_{nl})}$$

$$l_a^{nl}(\xi_\beta) = \delta_{\alpha\beta}$$

$$\mathbf{x}(\xi, \eta) = \sum_{a=1}^{n_a} N_a(\xi, \eta) \mathbf{x}_a$$



## Gauss-Lobatto-Legendre Point

$$(1 - \xi^2) P'_{nl}(\xi) = 0$$

$$f(\mathbf{x}(\xi, \eta, \zeta)) \approx \sum_{\alpha, \beta, \gamma=0}^{n_l} f^{\alpha\beta\gamma} l_\alpha(\xi) l_\beta(\eta) l_\gamma(\zeta)$$

$$f^{\alpha\beta\gamma} = f(\mathbf{x}(\xi_\alpha, \eta_\beta, \zeta_\gamma))$$

## Shape Function

$$dxdydz = J_e d\xi d\eta d\zeta$$

$$J_e = \left| \frac{\partial(x, y, z)}{\partial(\xi, \eta, \zeta)} \right|$$

## GLL Integration Rule

$$\int_{\Omega_e} f(\mathbf{x}) d^3x = \int_{-1}^1 \int_{-1}^1 \int_{-1}^1 f(\mathbf{x}(\xi, \eta, \zeta)) J_e(\xi, \eta, \zeta) d\xi d\eta d\zeta \\ \approx \sum_{\alpha, \beta, \gamma=0}^{n_l} \omega_\alpha \omega_\beta \omega_\gamma f^{\alpha\beta\gamma} J_e(\xi_\alpha, \eta_\beta, \zeta_\gamma)$$

(Komatitsch & Tromp, 1999)

# Earthquake Event

- 2016\_04\_01 Nankai Earthquake

-Mw 5.9

-Lat:  $33.38^{\circ}$  Lon:  $136.39^{\circ}$  Depth: 18.8km

-necis.kma.go.kr 기상청 지진관측소 자료

-32개 광대역 속도관측소 & 시추공 광대역 속도관측소

-Point Source Model (Global CMT Catalog)

-East Asia Model (Kim et al., 2016) / 1D-PREM (Dziewonski & Anderson, 1981 )

- 1946\_12\_20 Nankai Earthquake

-Mw 8.4

-Lat:  $33.03^{\circ}$  Lon:  $135.62^{\circ}$  Depth: 30.0km

-Finite Fault Model (Tanioka & Satake, 2001. | Baba et al., 2002)

-Rupture Velocity: Vs (3.496km/s) |  $0.75 \times$  Vs (2.622km/s)

-Half-Duration: 0.7s

Preliminary Determination of Epicenter								body-wave magnitude	surface-wave magnitude
year	day	min	sec	latitude	longitude	depth	mb	Ms	PDE event name
PDE 2001	9	9	23	59	17.78	34.0745	-118.3792	6.4	4.2
event name: 9703873									
time shift: 0.0000									
half duration: 0.0000									
latorUTM: 34.0745									
longorUTM: -118.3792									
depth: 5.4000									
Mrr:	-0.002000e+23								
Mtt:	-0.064000e+23								
Mpp:	0.066000e+23								
Mrt:	-0.090000e+23								
Mrp:	-0.002000e+23								
Mtp:	0.188000e+23								

Harvard CMT solution

$\mathbf{M} = \begin{bmatrix} M_{rr} & M_{r\theta} & M_{r\phi} \\ M_{r\theta} & M_{\theta\theta} & M_{\theta\phi} \\ M_{r\phi} & M_{\theta\phi} & M_{\phi\phi} \end{bmatrix}$

$$\mathbf{M}_0 = \frac{1}{\sqrt{2}} (\mathbf{M} : \mathbf{M})^{1/2} \approx 2.18 \times 10^{22} \text{ dyne cm}$$
$$M_w = \frac{2}{3} (\log_{10} M_0 - 16.1) \approx 4.19$$

(SPECFEM3D-Cartesian Manual)

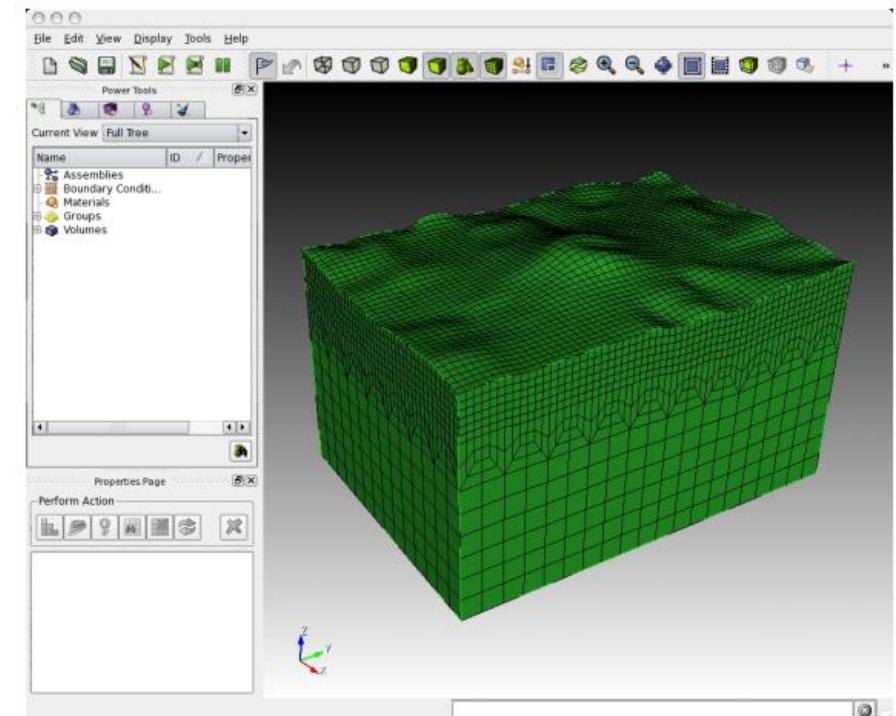
# Simulation

- Mesh Generation – Trelis

- 한반도와 일본을 포함한 1300km\*900km\*170km mesh
- Topography Data from SRTM(<http://srtm.csi.cgiar.org>)
- element size 4km → computed maximum frequency = (element size)/(wave velocity)

- Computation – SPECFEM3D Cartesian

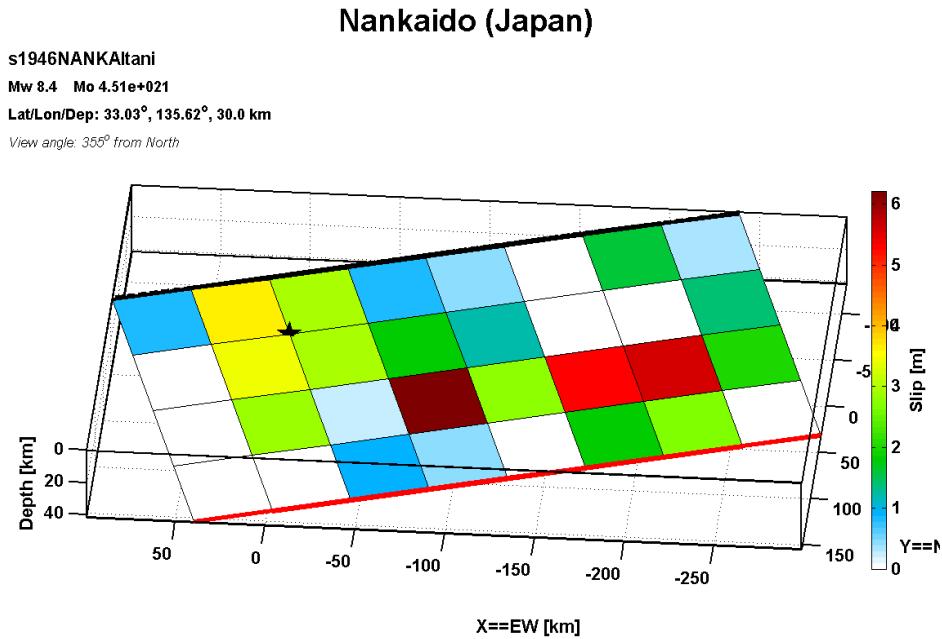
- dT = 0.04s
- Nstep = 20000
- cpu 200: ~10hrs



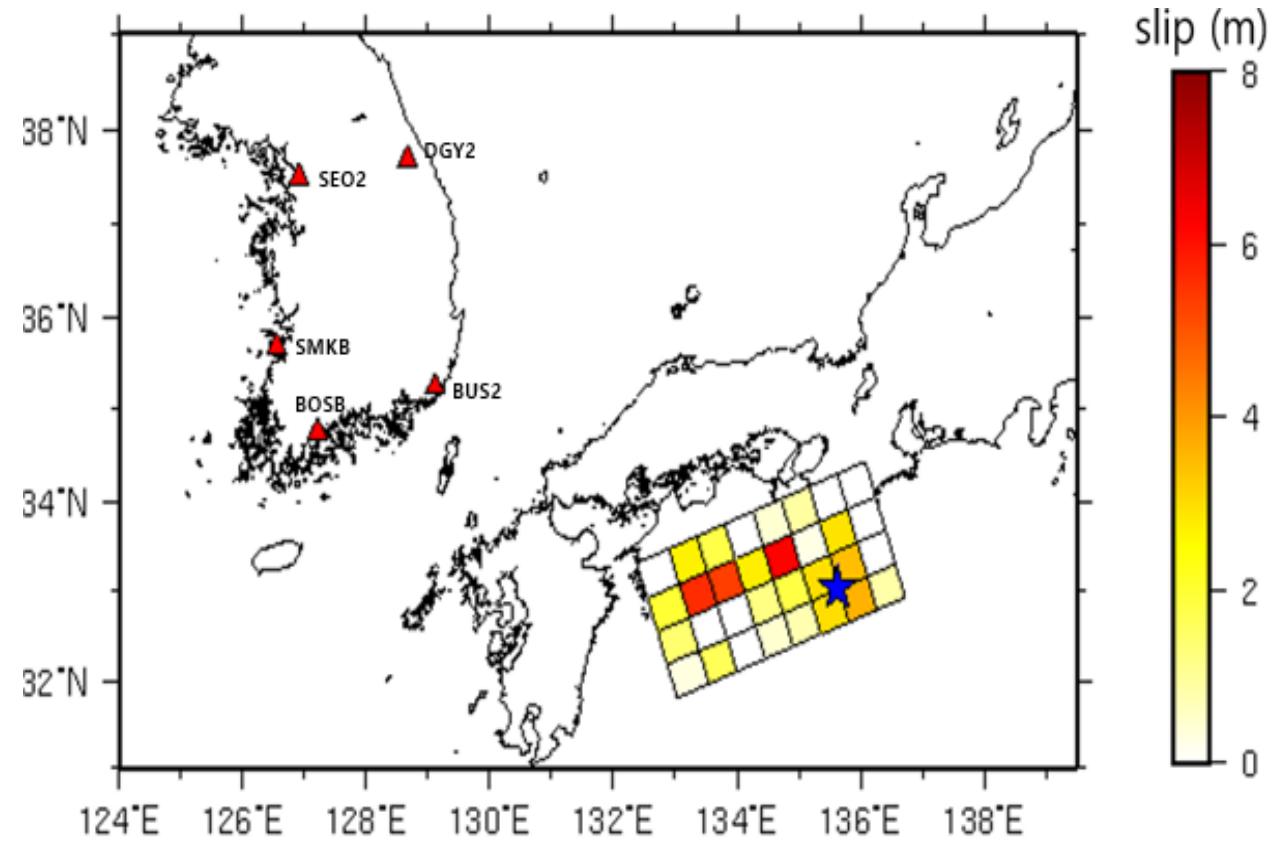
(SPECFEM3D-Cartesian Manual)

# 1946 Nankai Earthquake Finite Fault Model (Tanioka & Satake, 2001)

- Tsunami waveform inversion
- Strike:  $250.00^\circ$  / Dip:  $33.38^\circ$  / Rake:  $120.00^\circ$
- 29 subfaults -  $45\text{km} \times 45\text{km}$
- Slip distribution:  $0 \sim 6.22\text{m}$



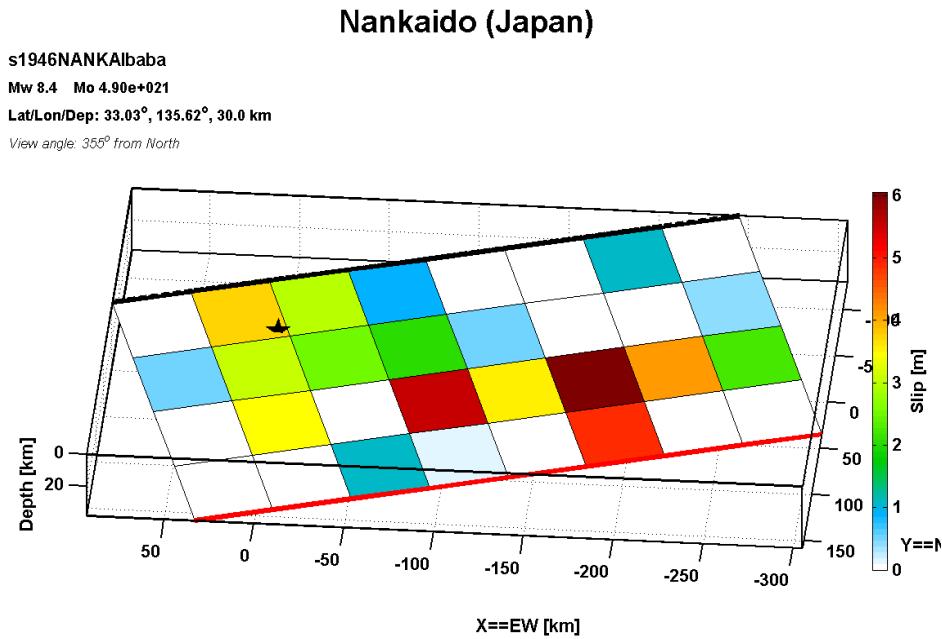
(Tanioka & Satake, 2001)



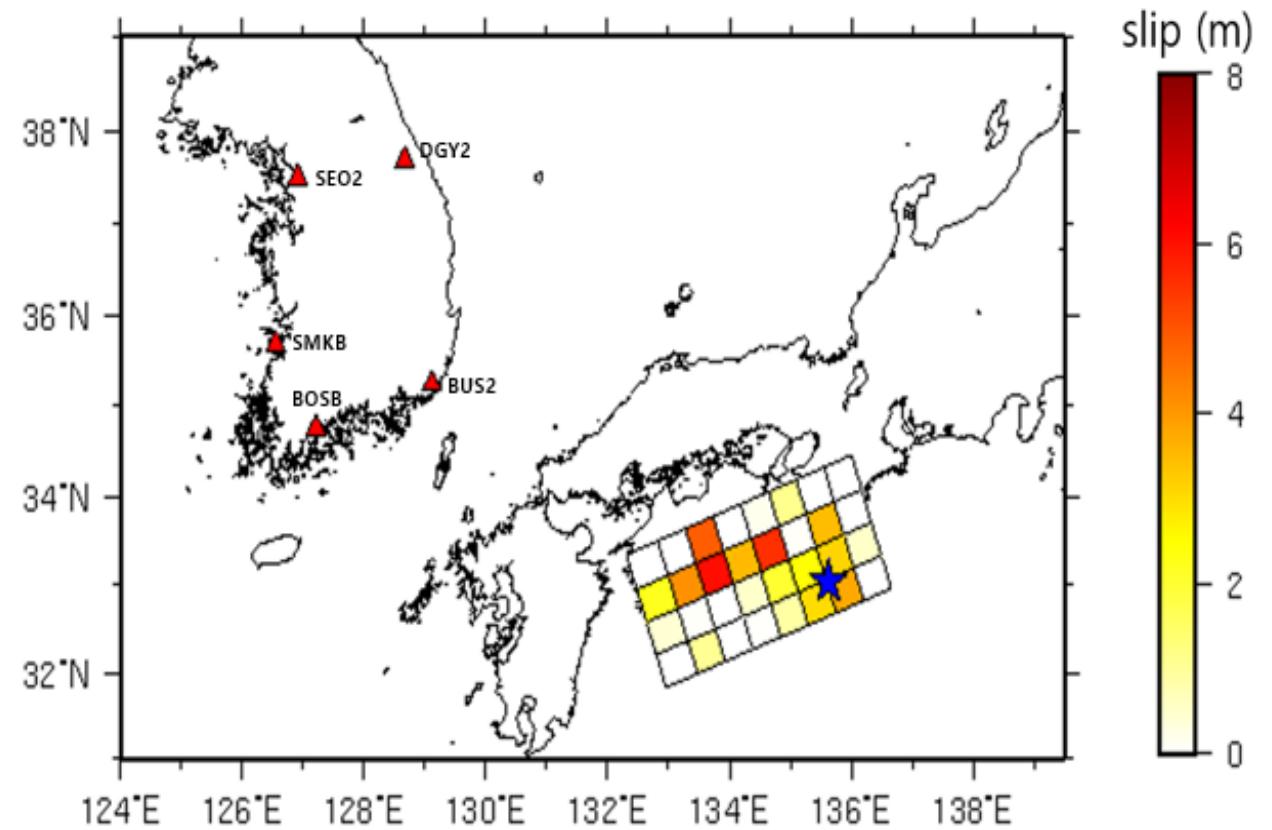
# 1946 Nankai Earthquake Finite Fault Model -(Baba et al., 2002)

- Modified Philippine Sea Plate Model

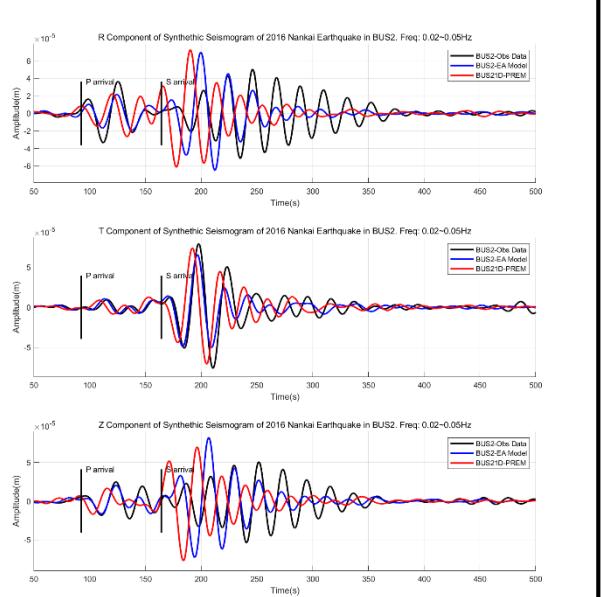
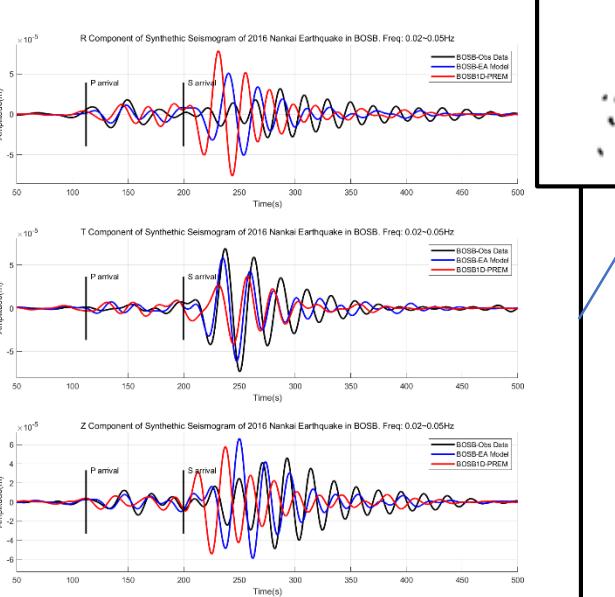
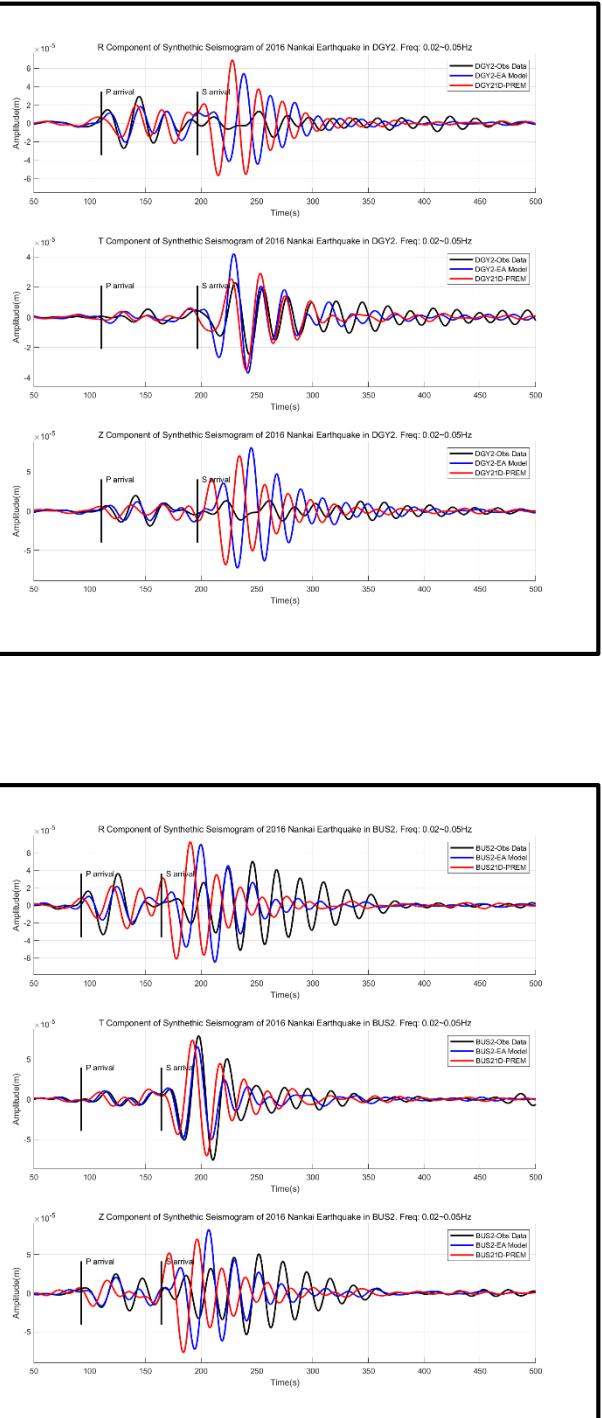
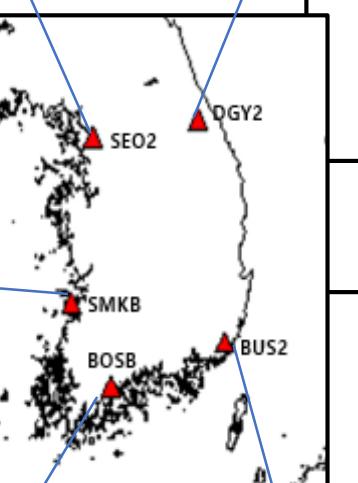
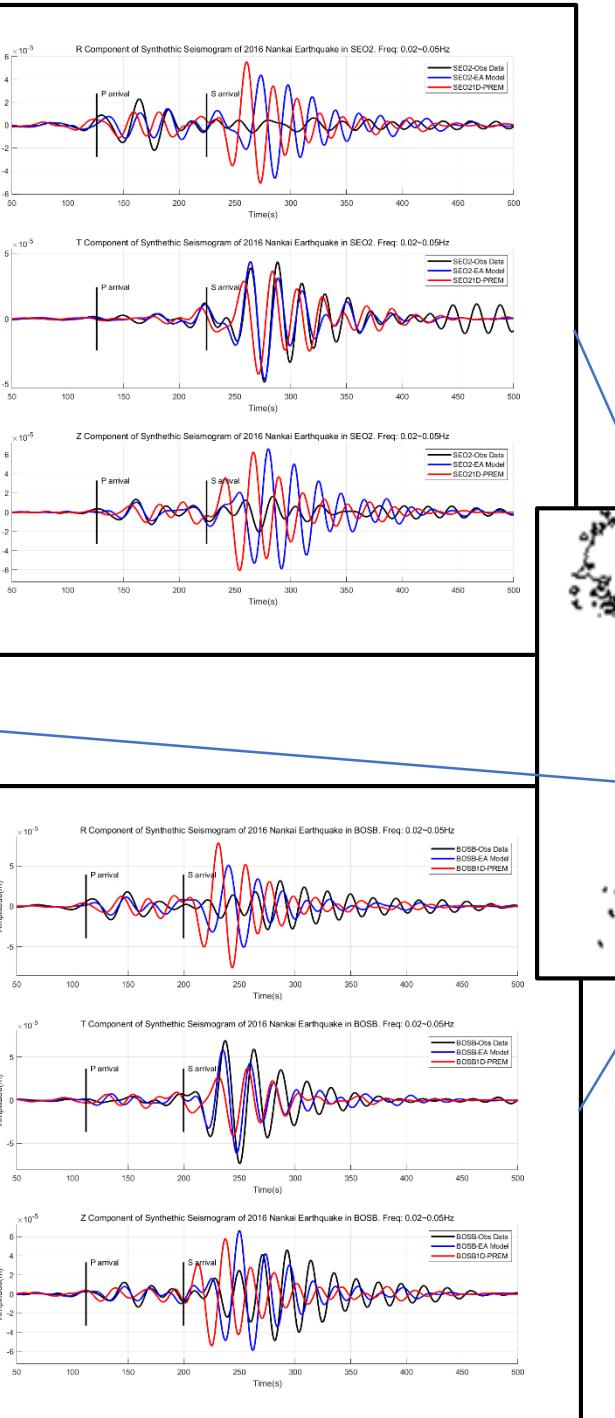
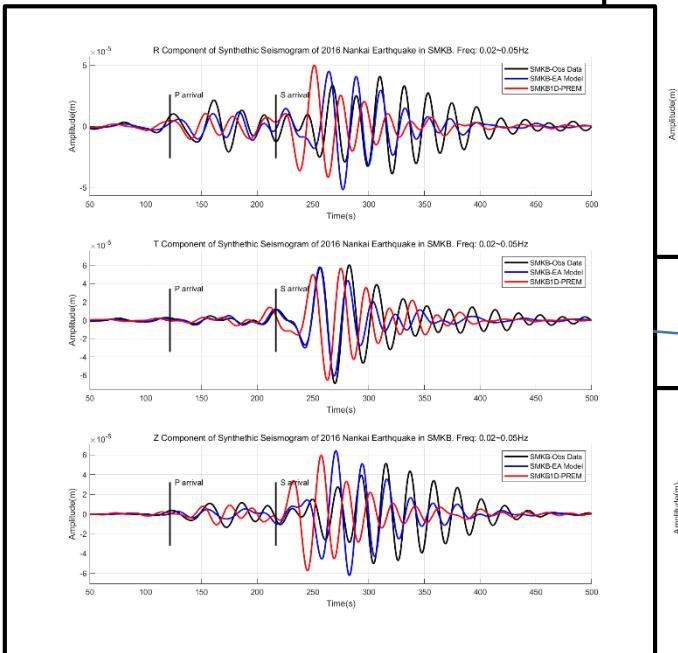
- combined marine seismic survey results and seismicity data
- Dip:  $13.00^\circ \rightarrow 12.00^\circ$
- Seismic Moment:  $4.51E+21 \rightarrow 4.9E+21$
- Slip distribution variance:  $132.22\text{m}^2 \rightarrow 112.95\text{m}^2$



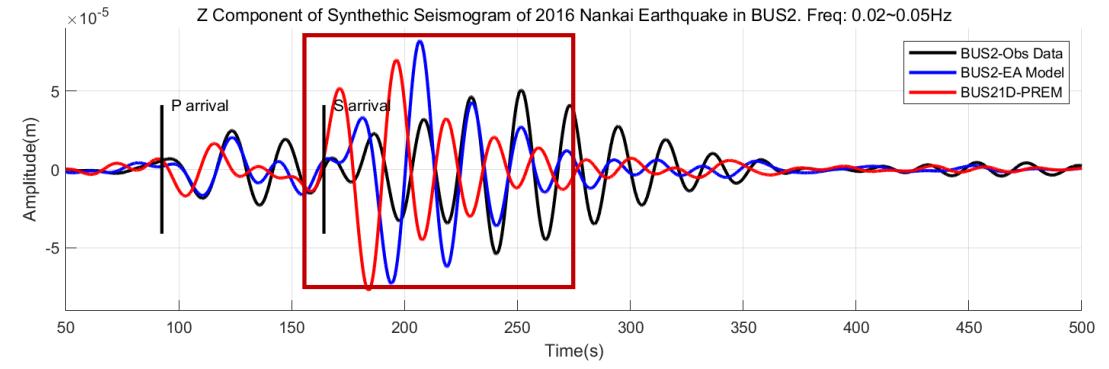
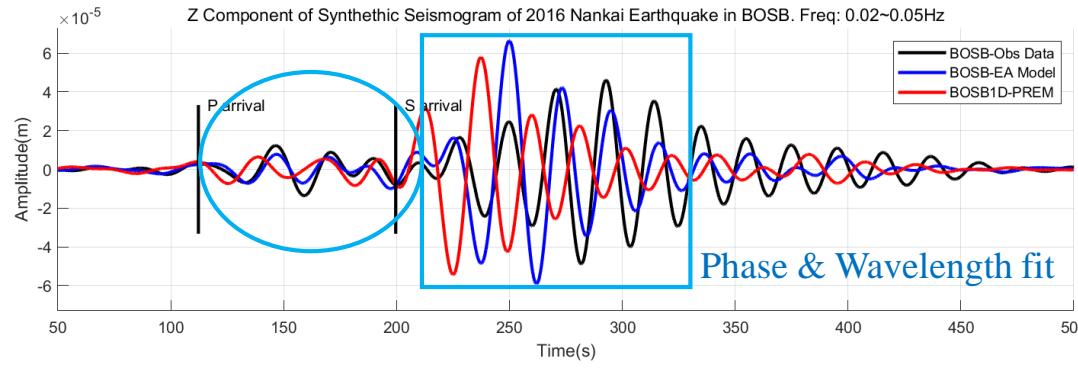
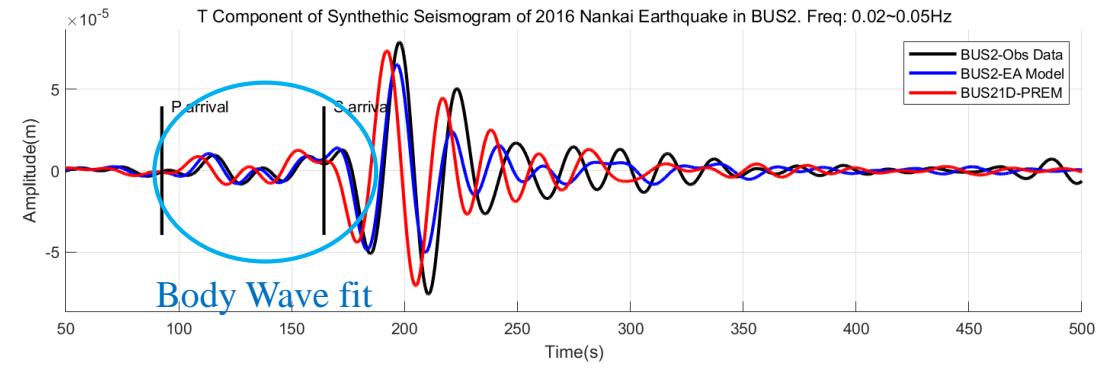
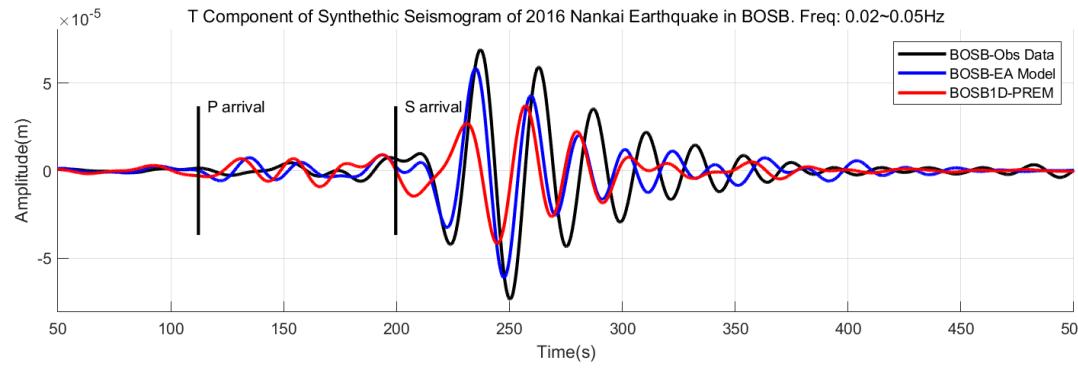
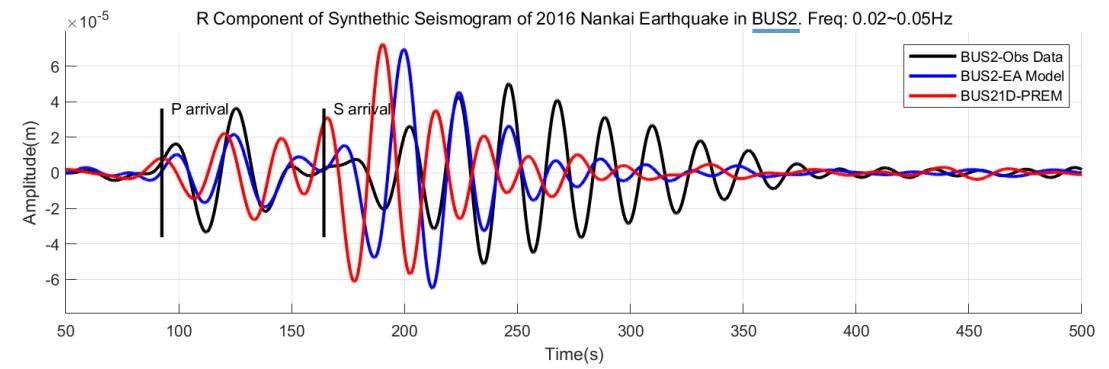
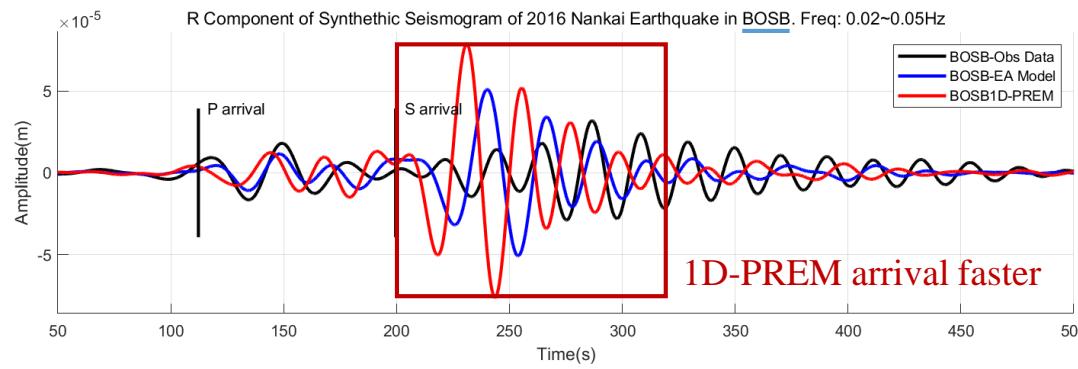
(Baba et al., 2002)

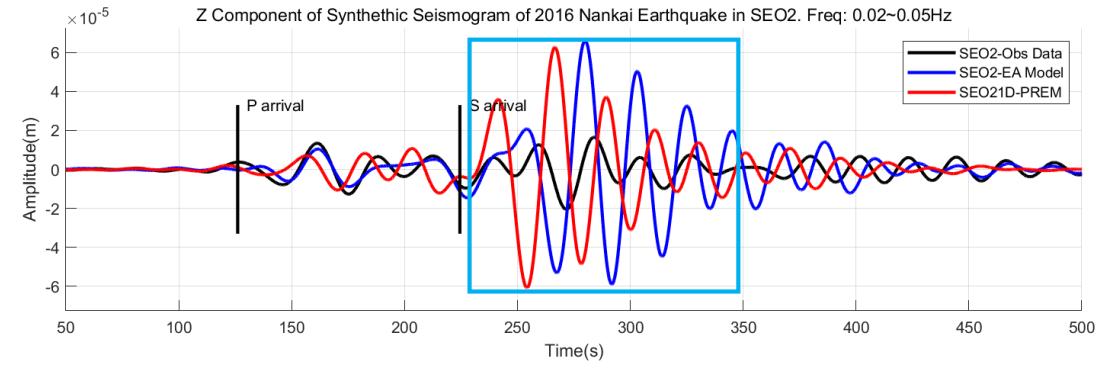
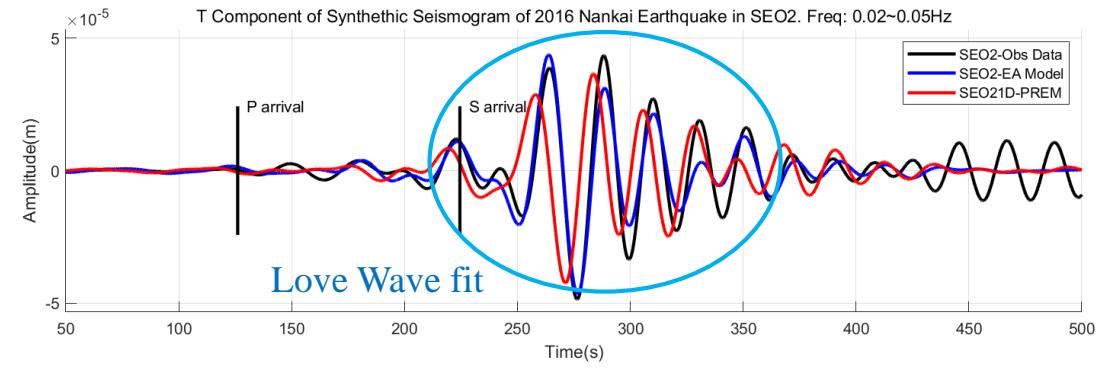
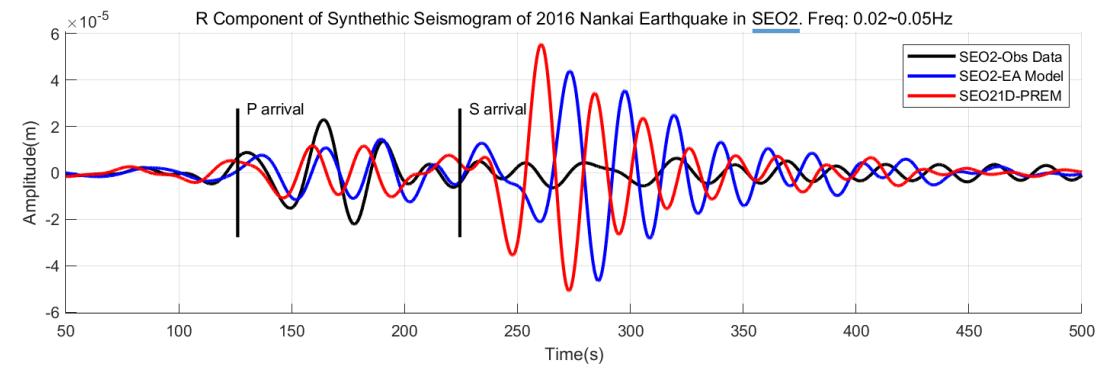
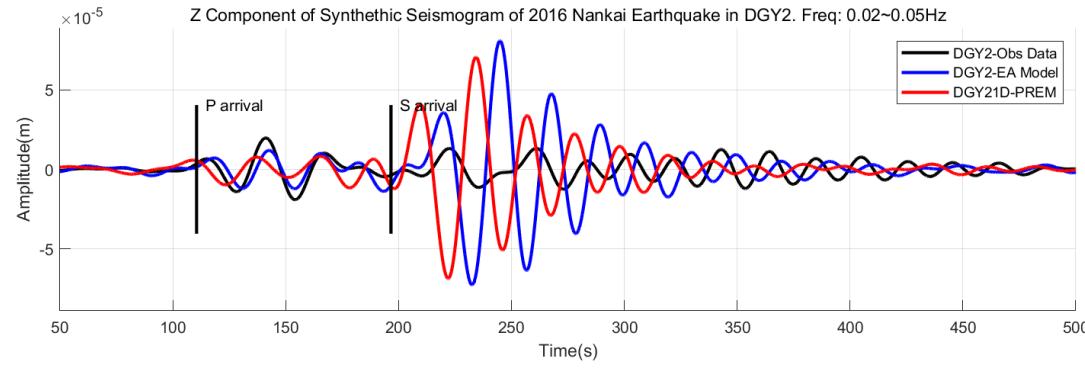
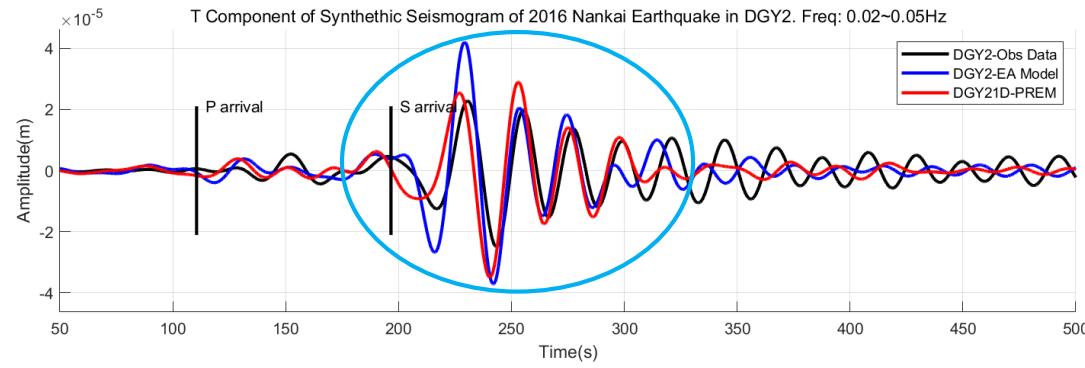
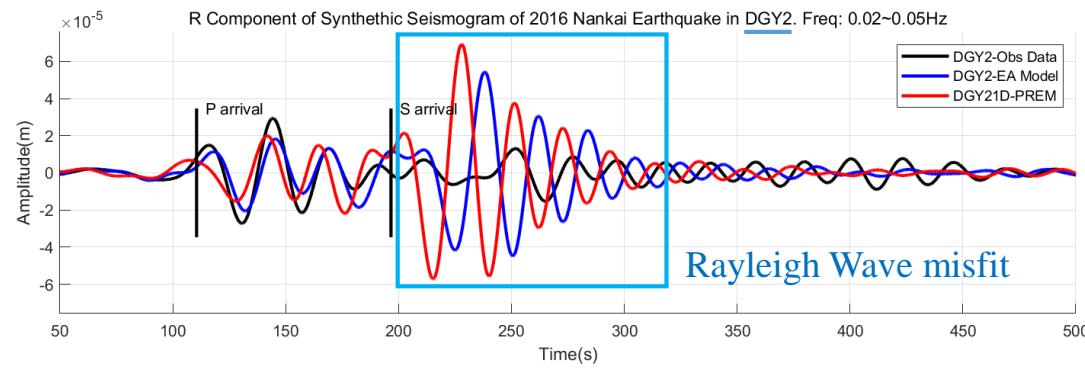


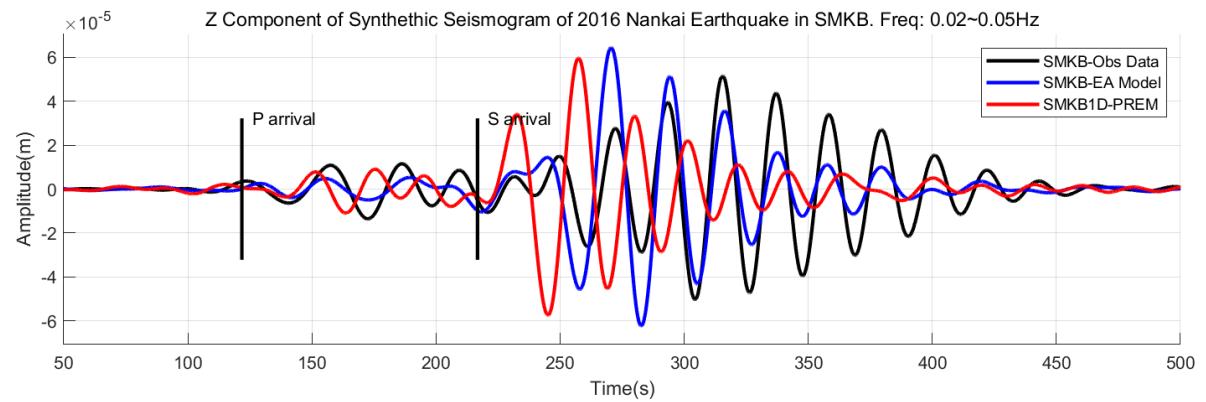
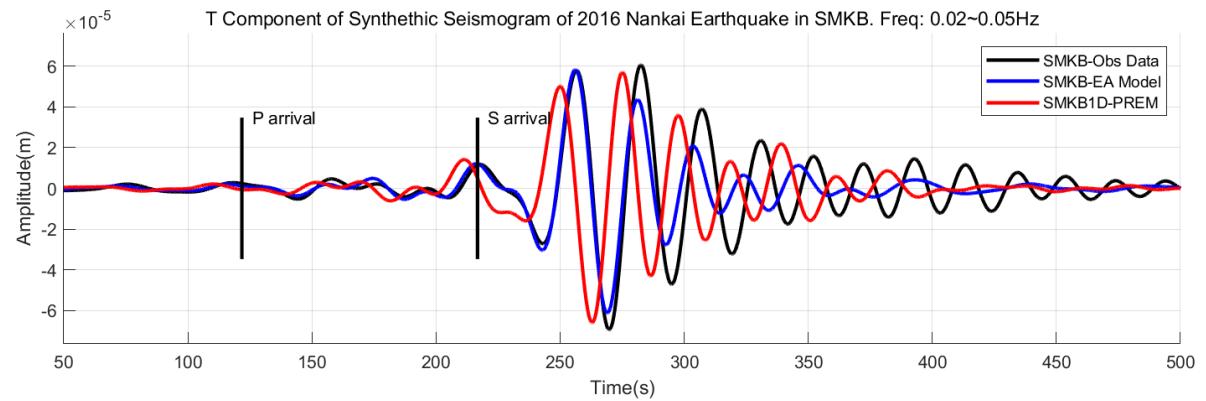
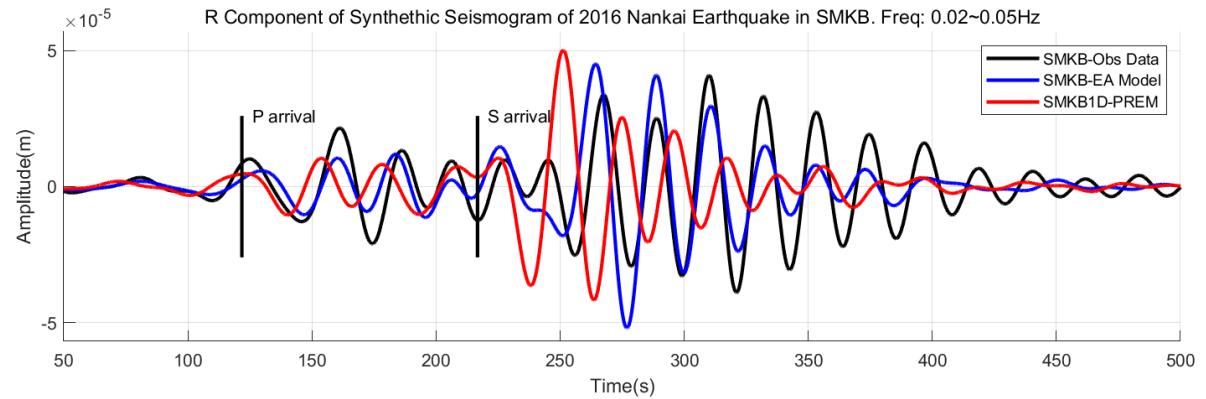
# Synthetic Seismogram Of 2016 Nankai Earthquake (Freq: 0.02~0.05Hz)



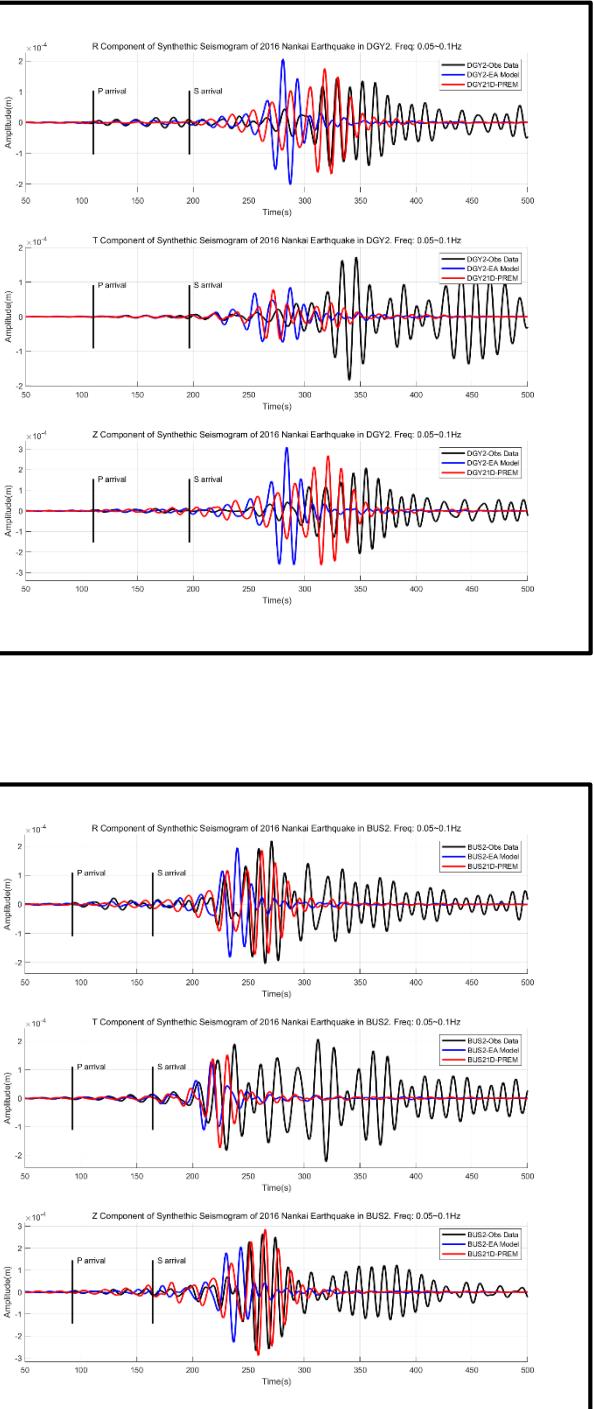
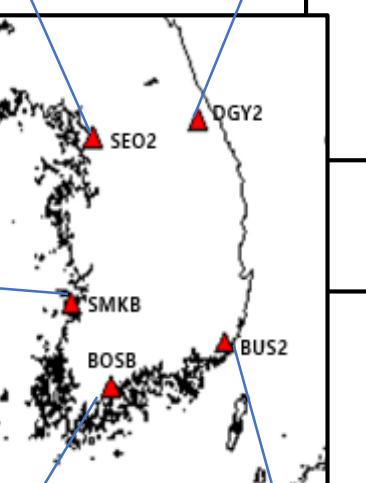
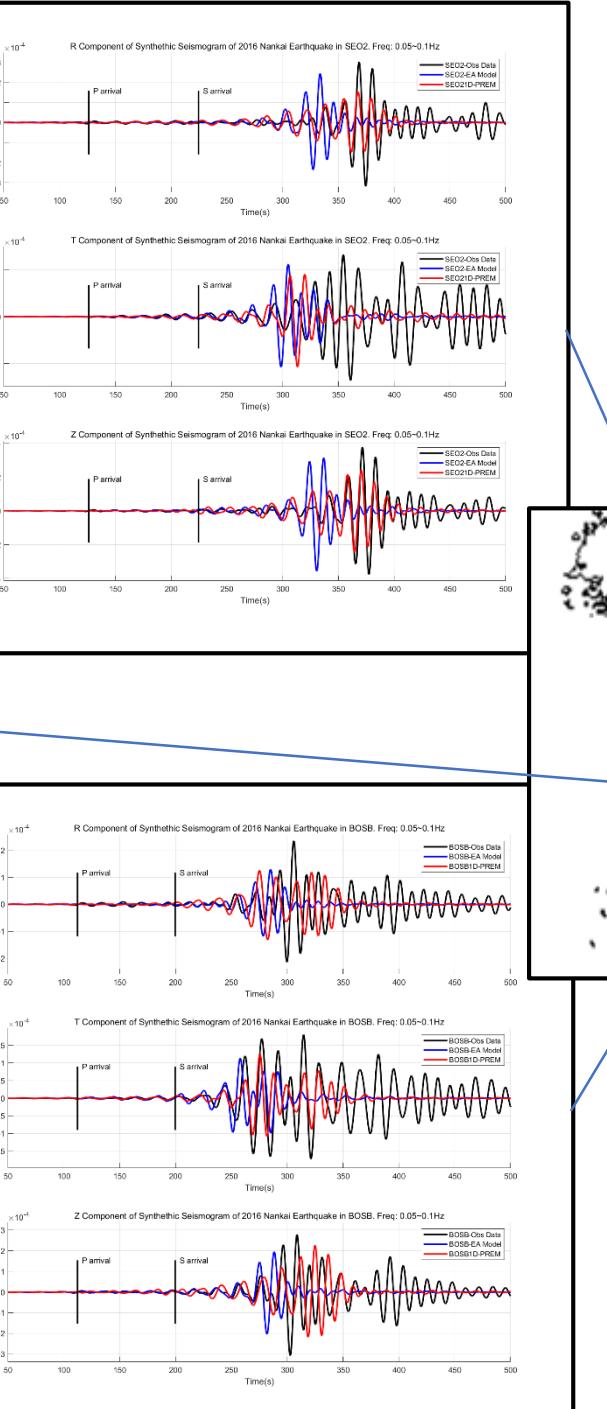
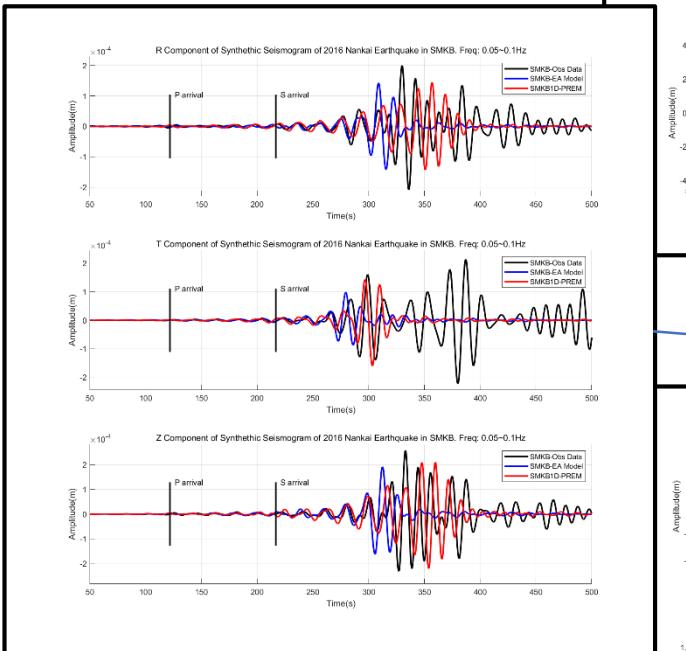
## Stations closer to event fit better

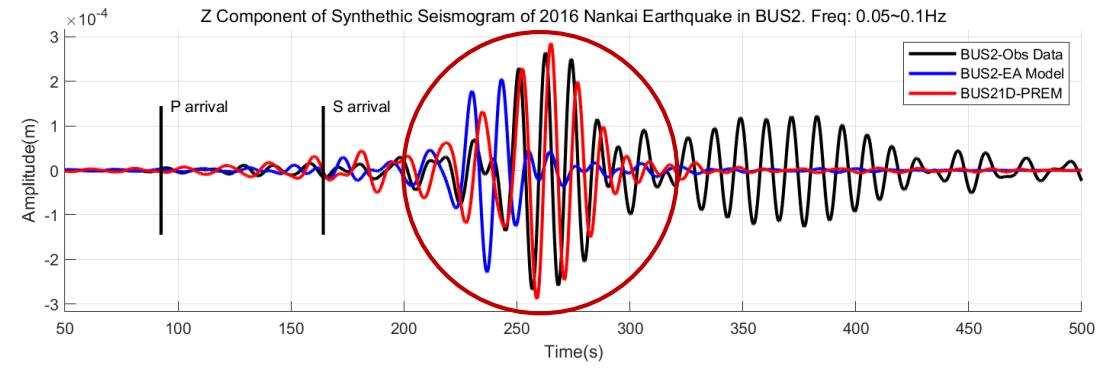
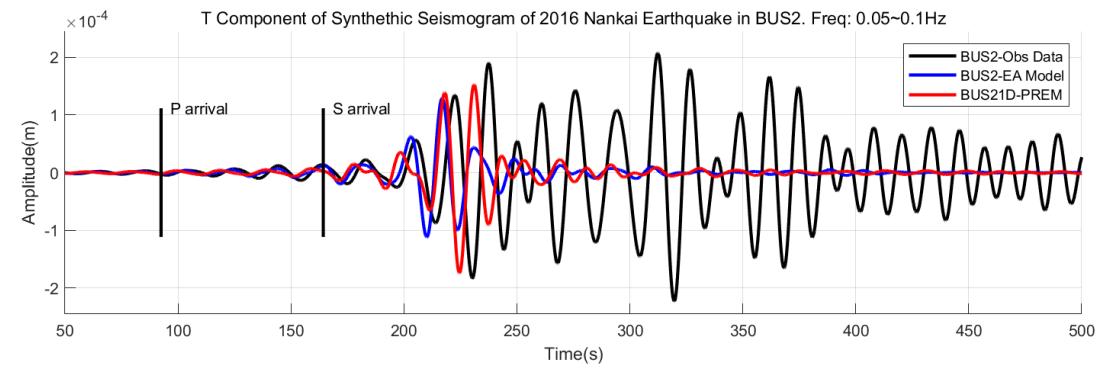
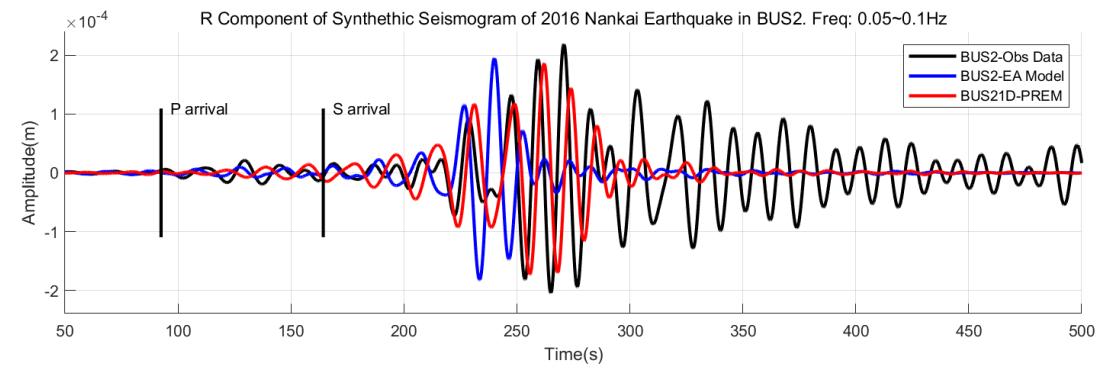
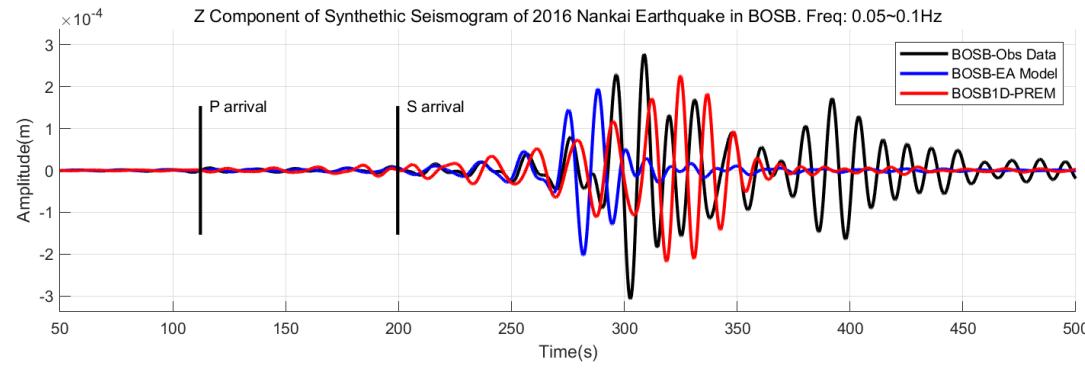
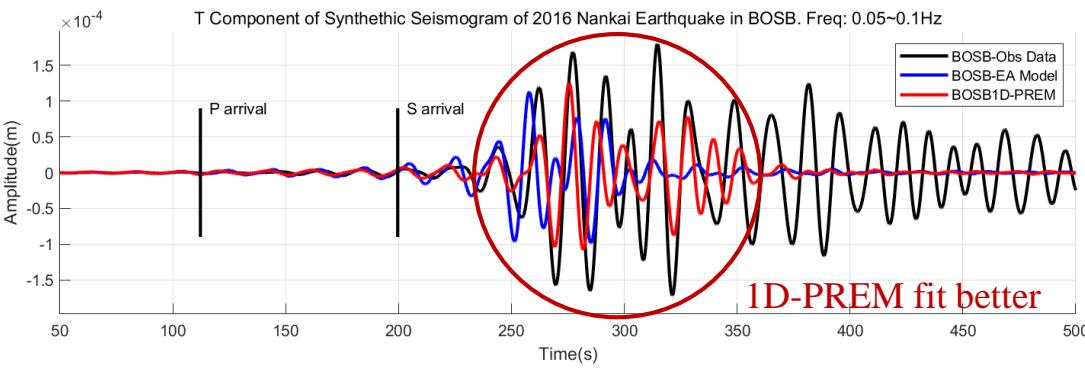
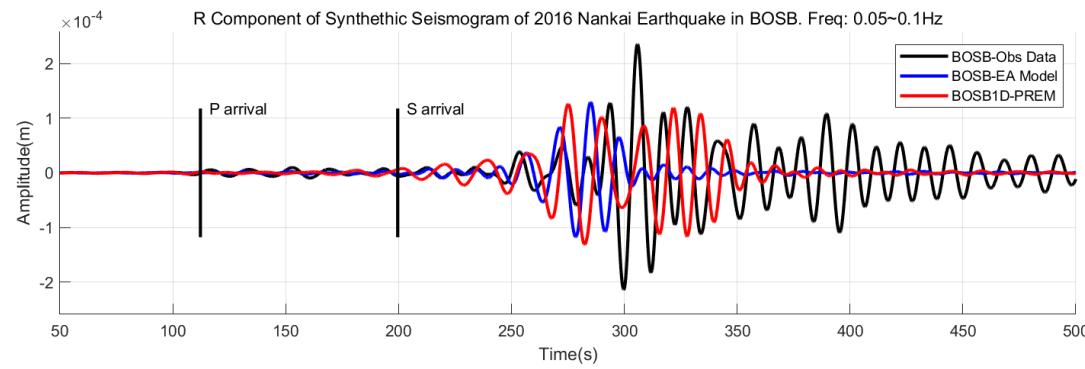


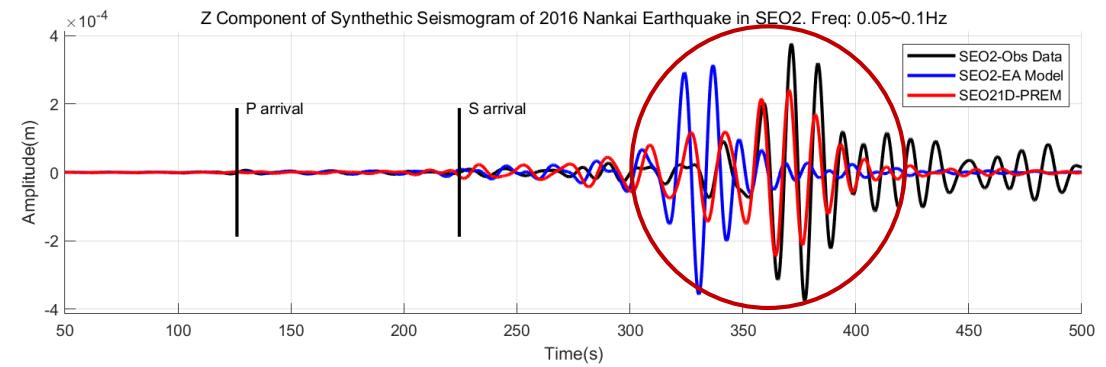
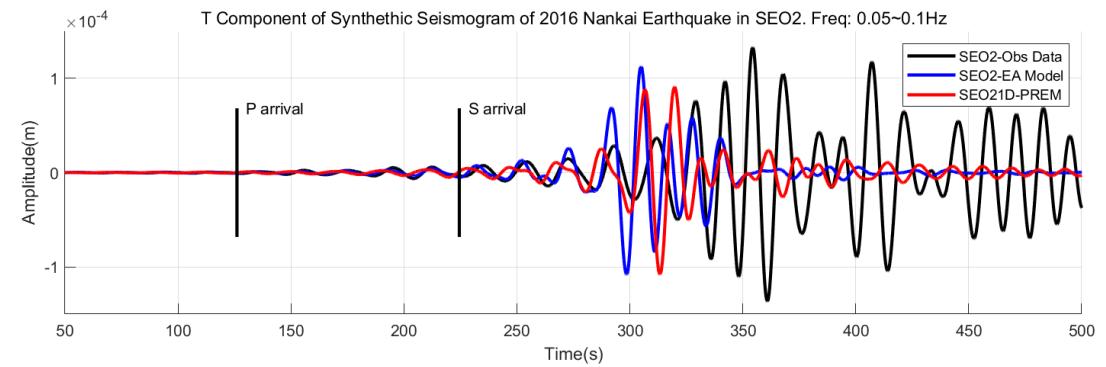
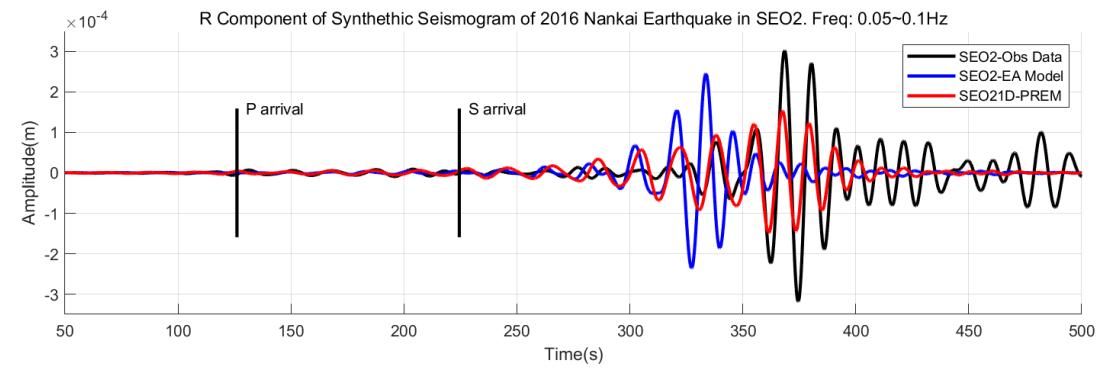
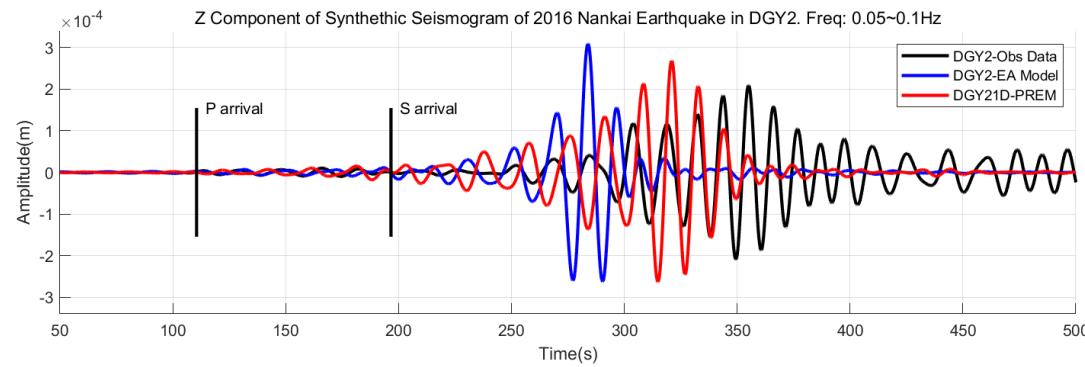
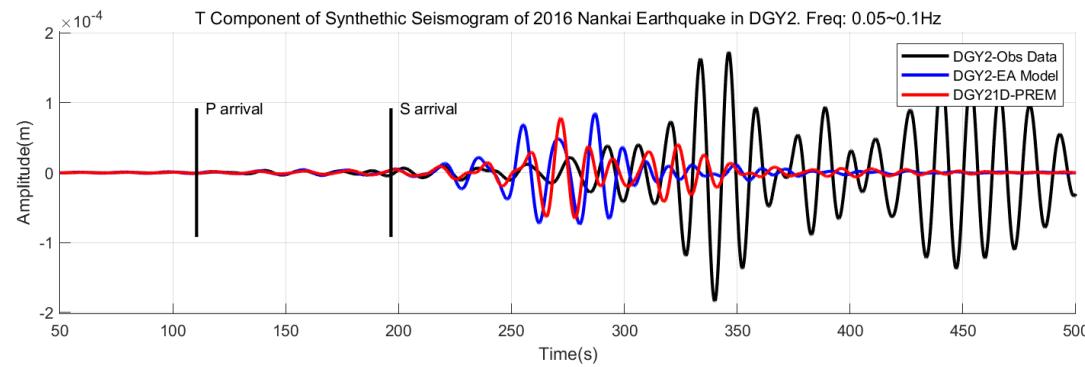
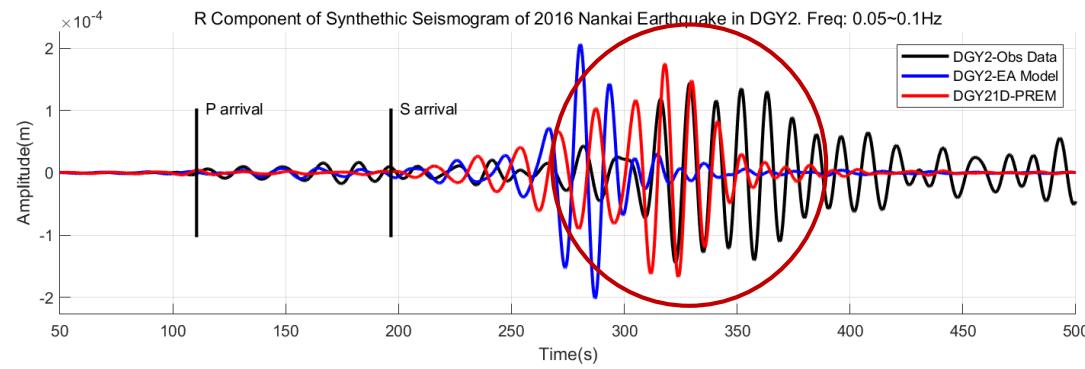


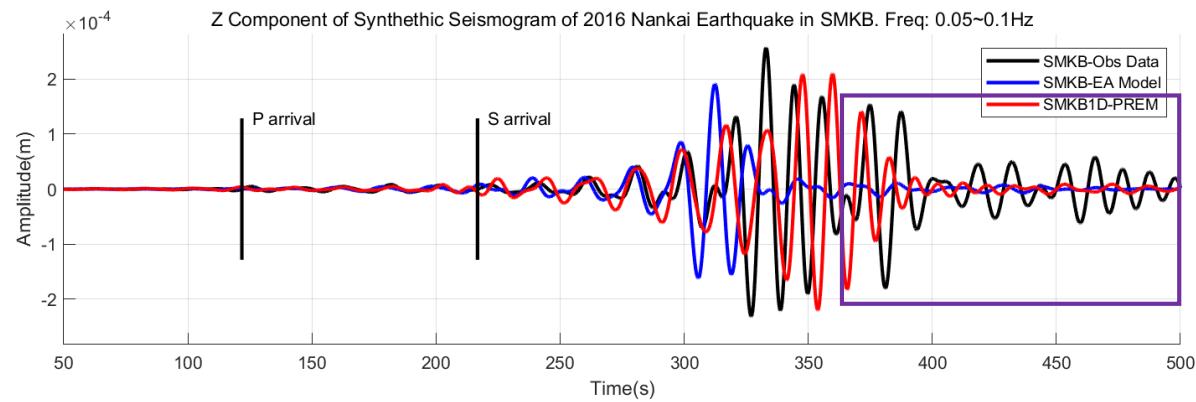
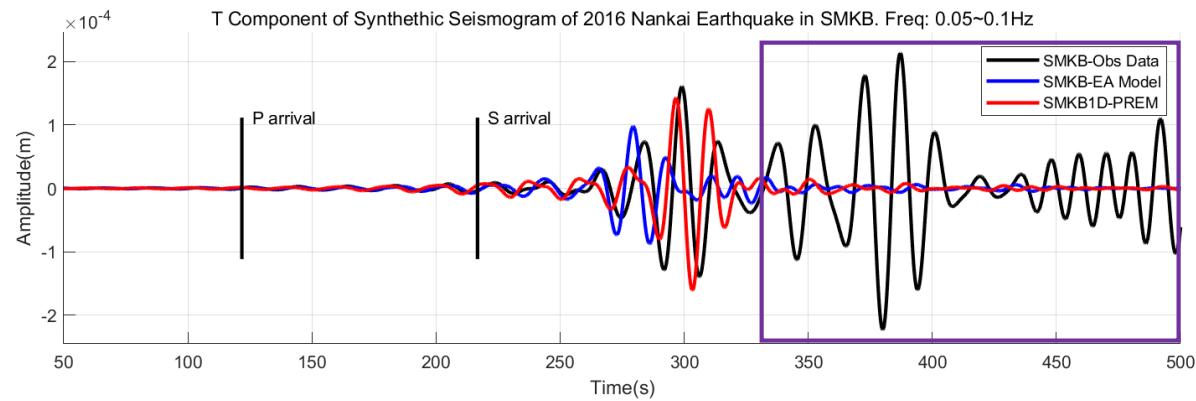
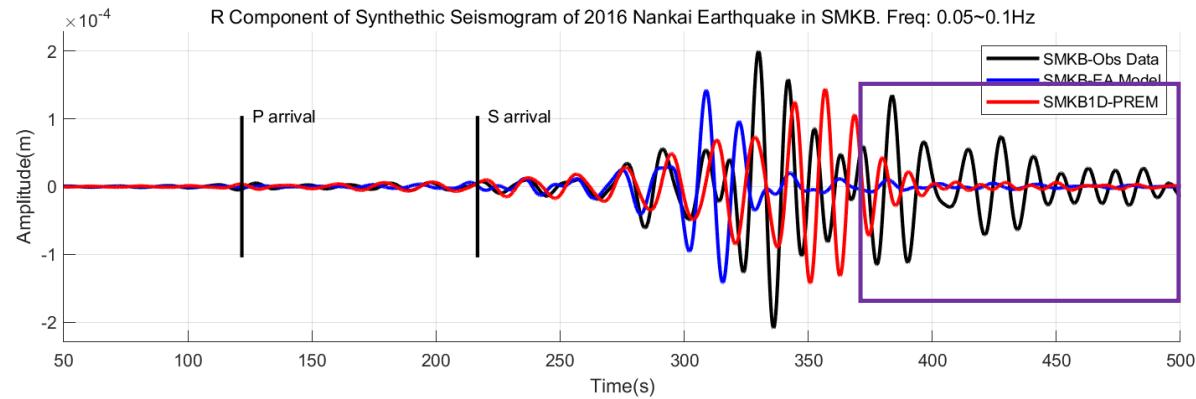


# Synthetic Seismogram Of 2016 Nankai Earthquake (Freq: 0.05~0.1Hz)



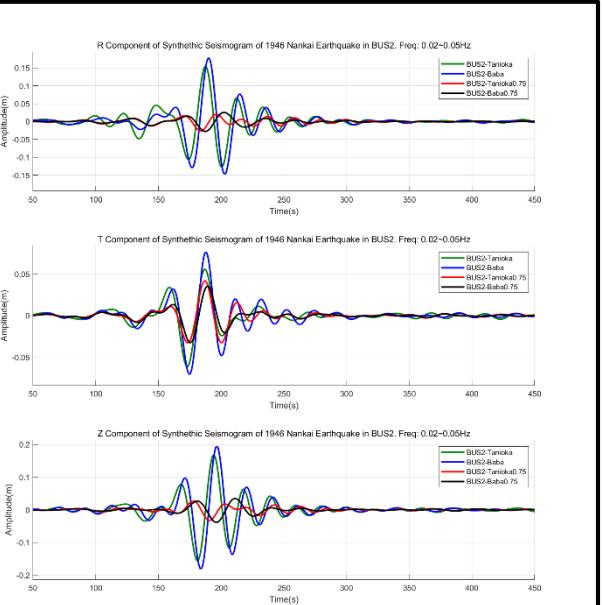
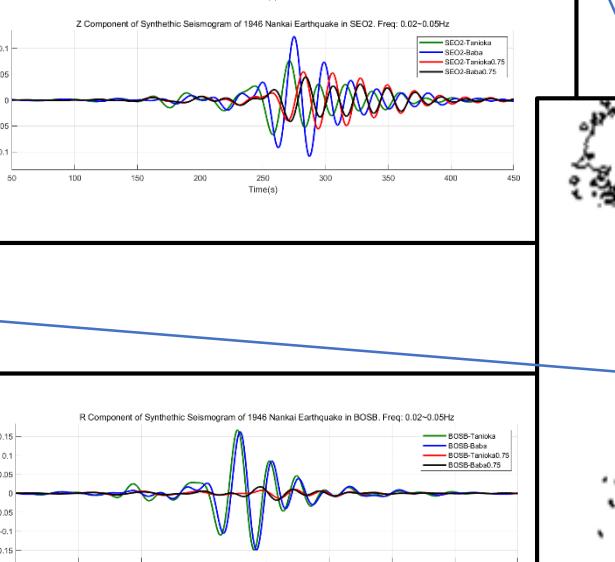
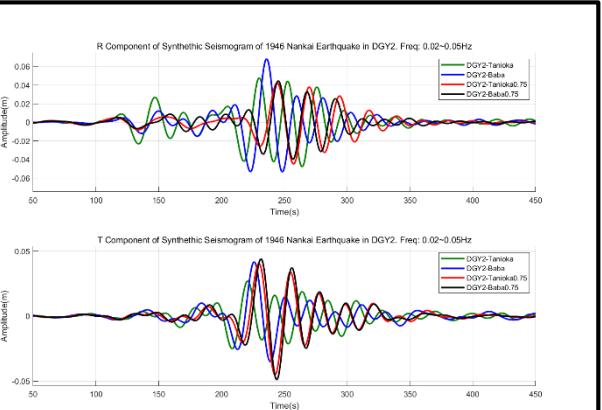
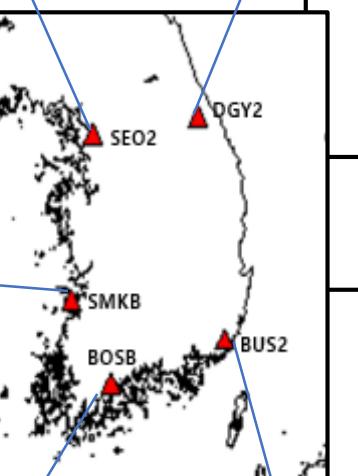
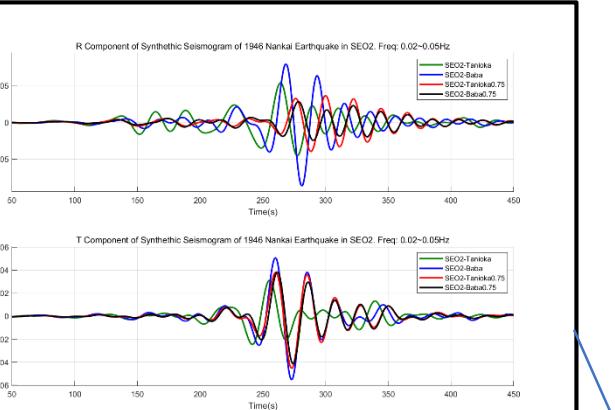
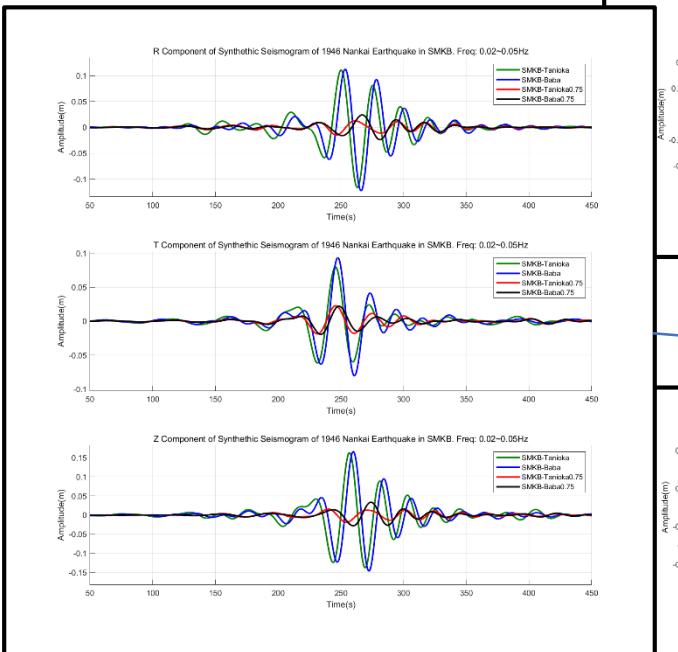


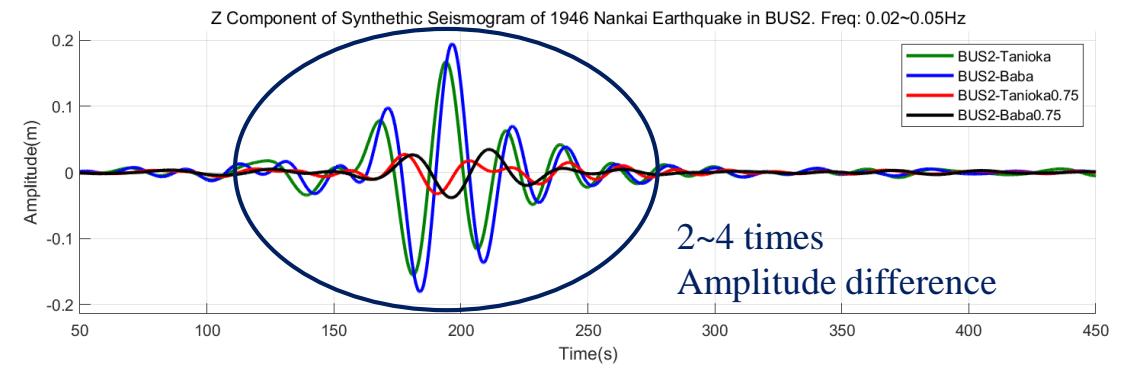
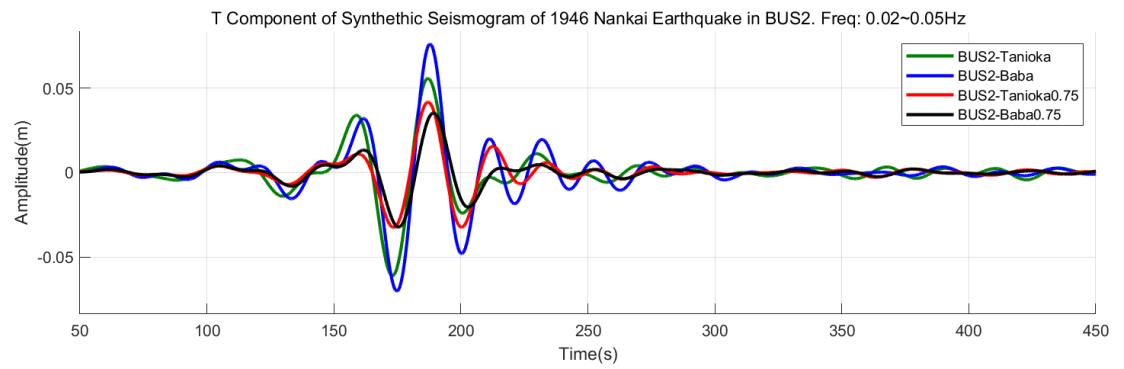
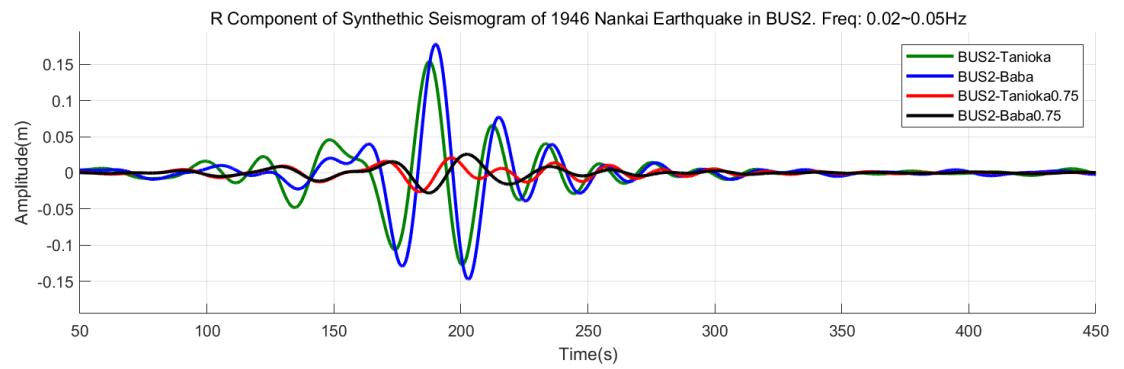
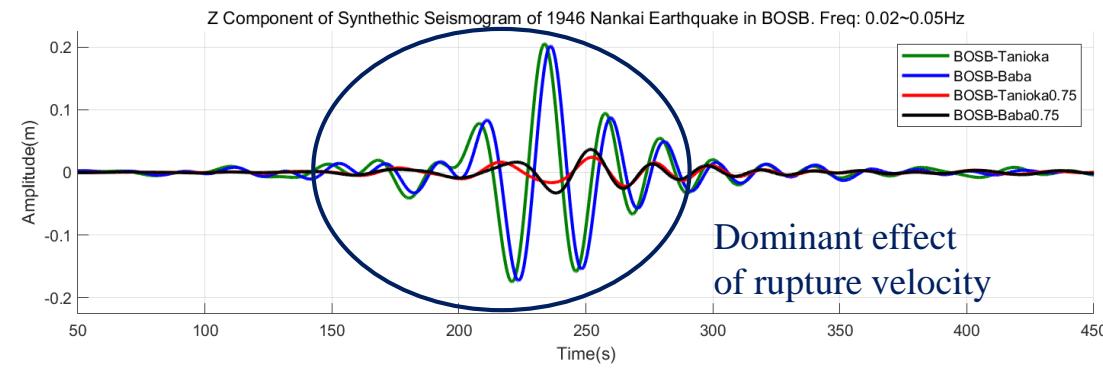
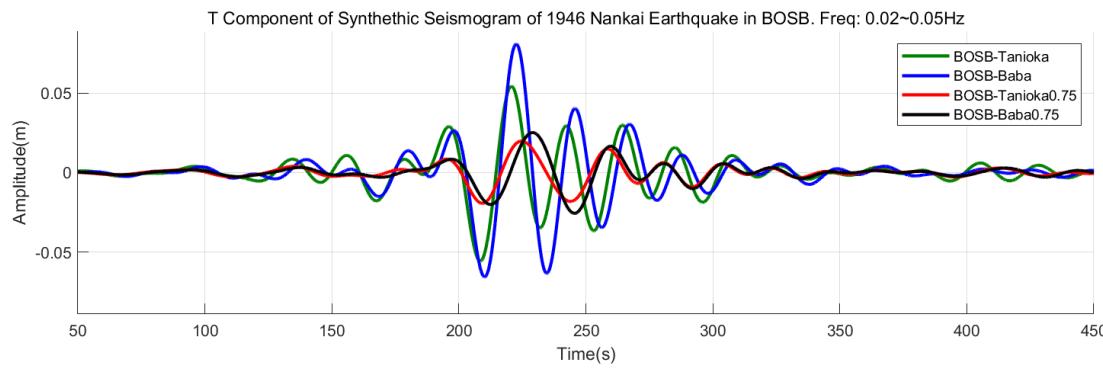
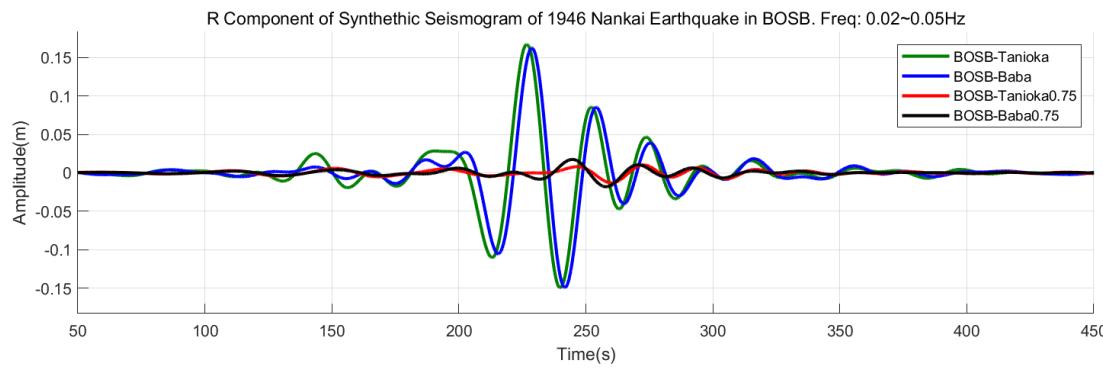


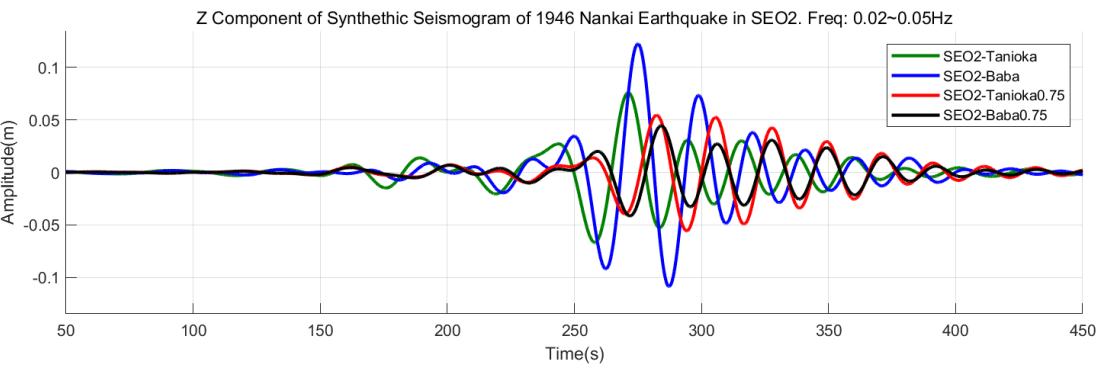
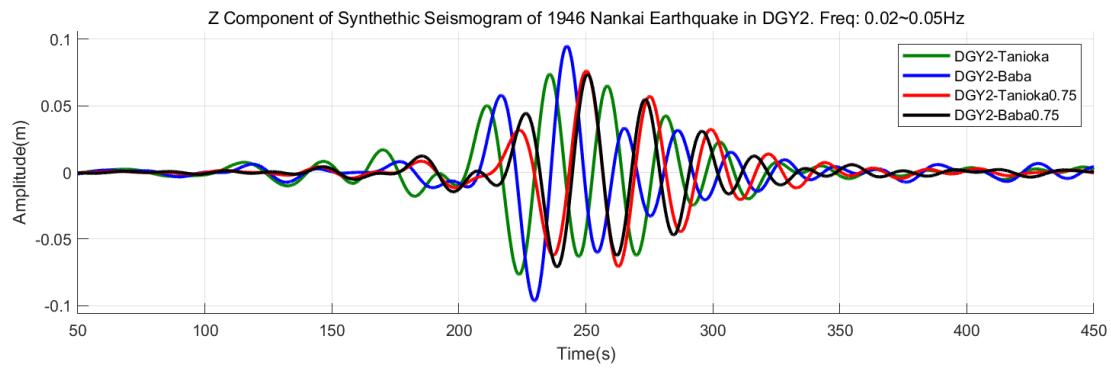
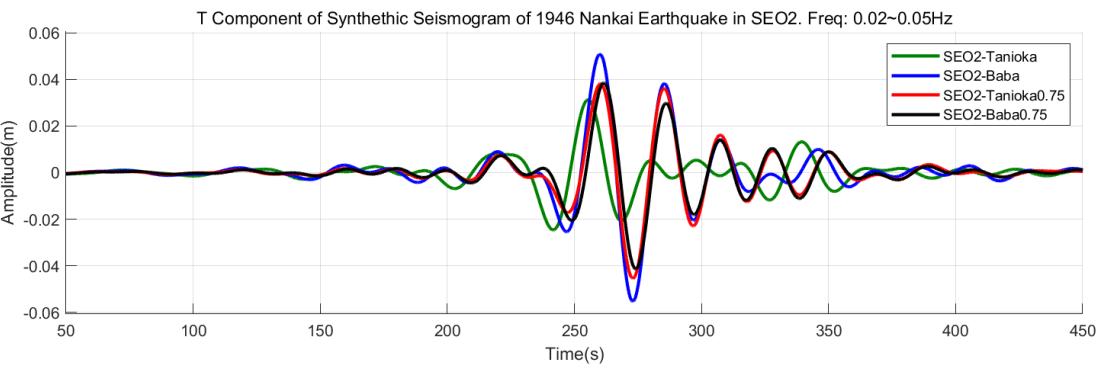
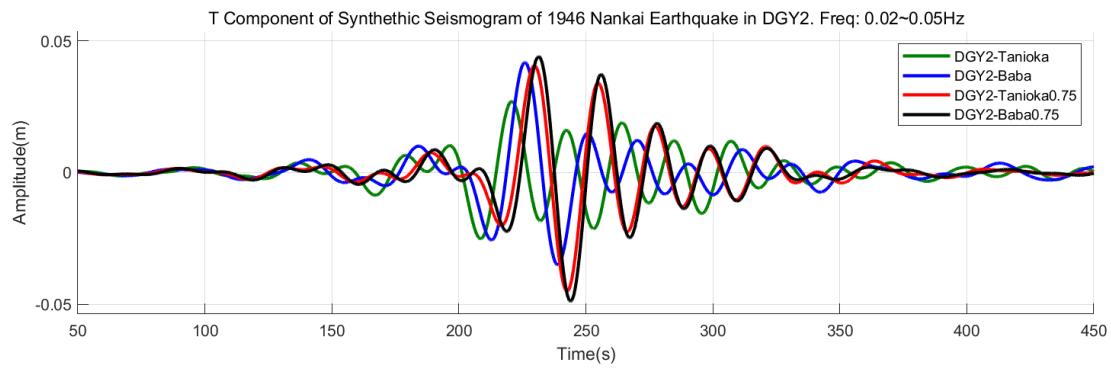
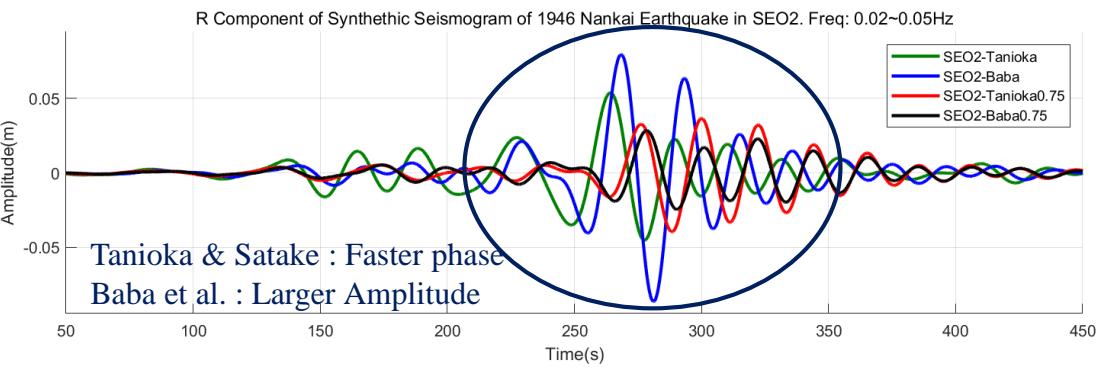
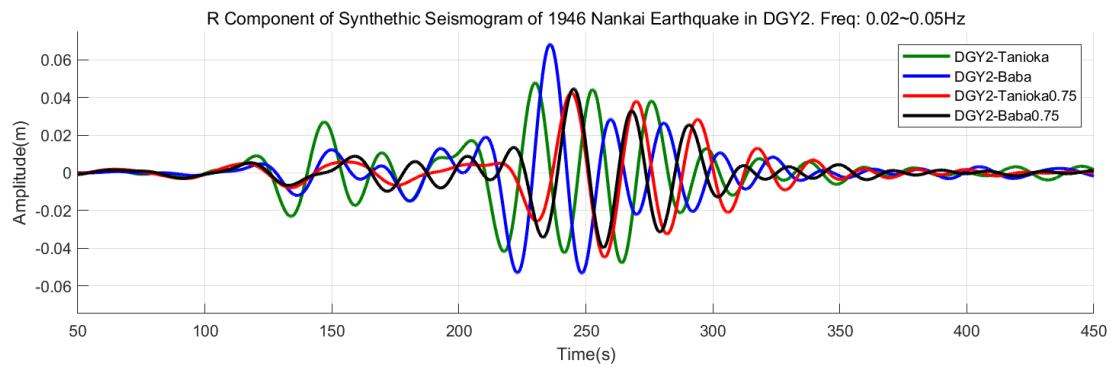


Surface wave dispersion  
→ crust effect?

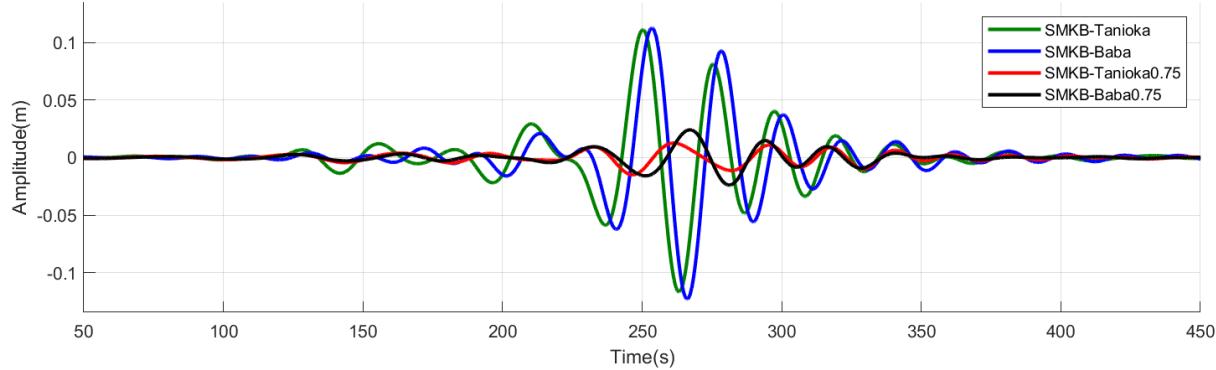
# Synthetic Seismogram Of 1946 Nankai Earthquake (Freq: 0.02~0.05Hz)



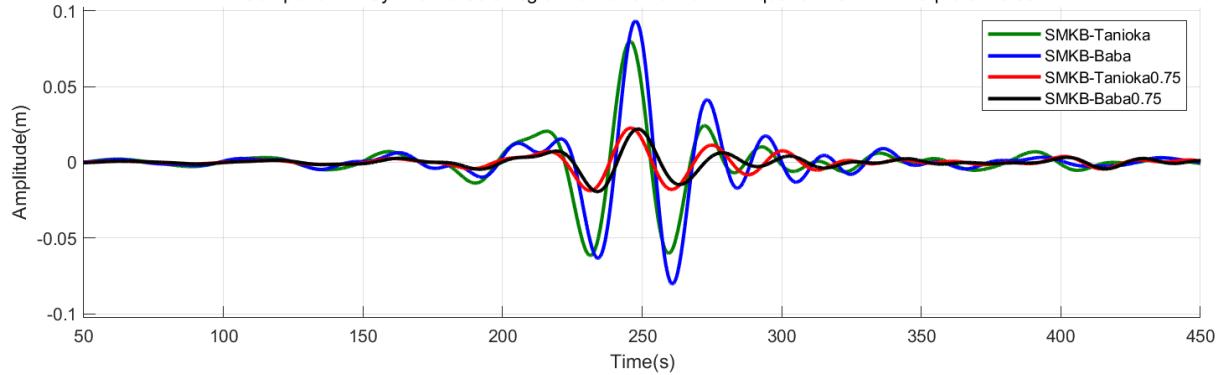




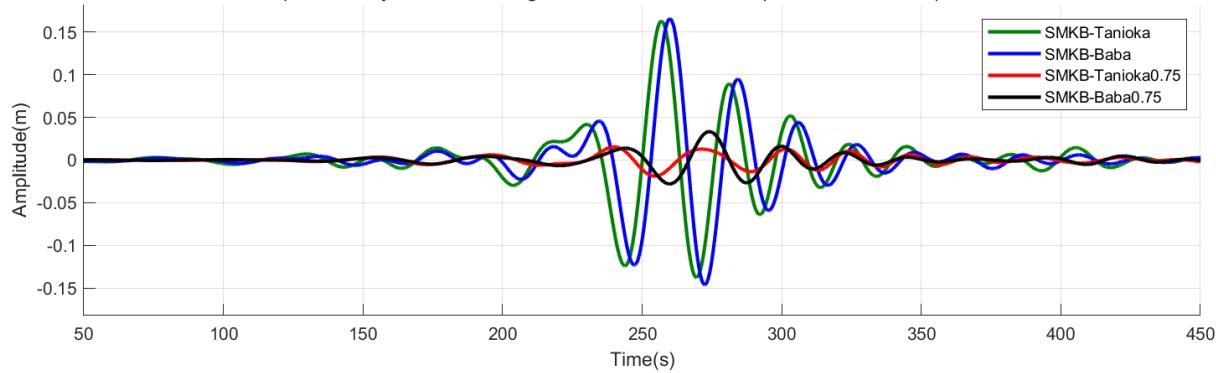
R Component of Synthetic Seismogram of 1946 Nankai Earthquake in SMKB. Freq: 0.02~0.05Hz



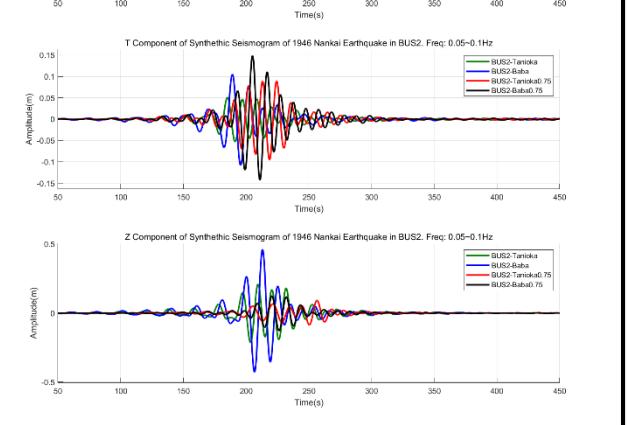
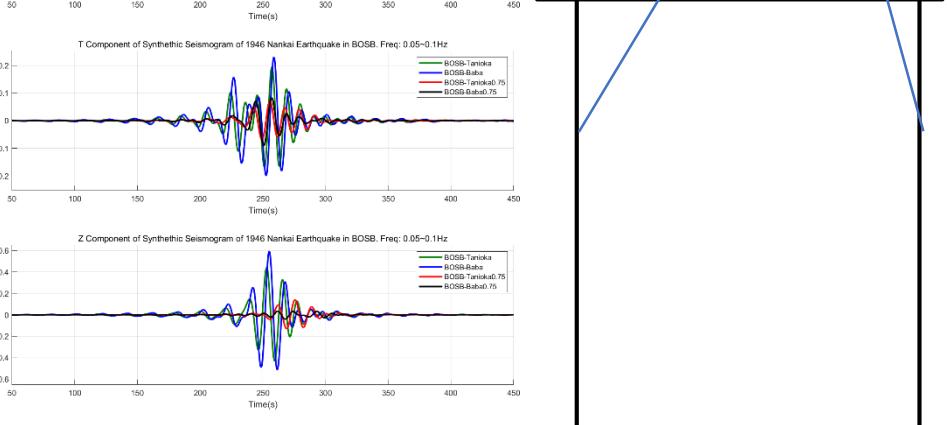
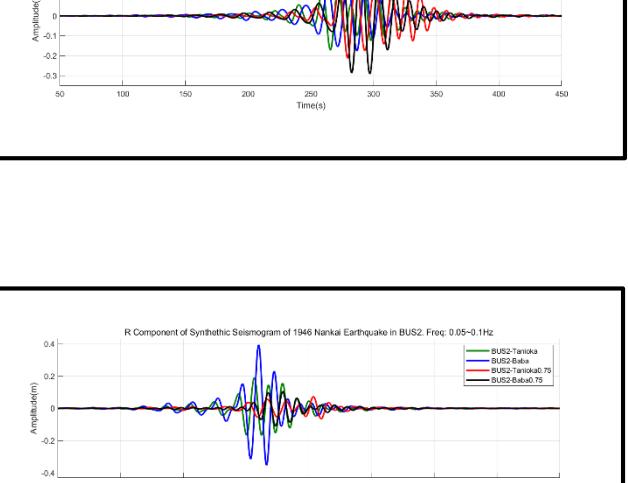
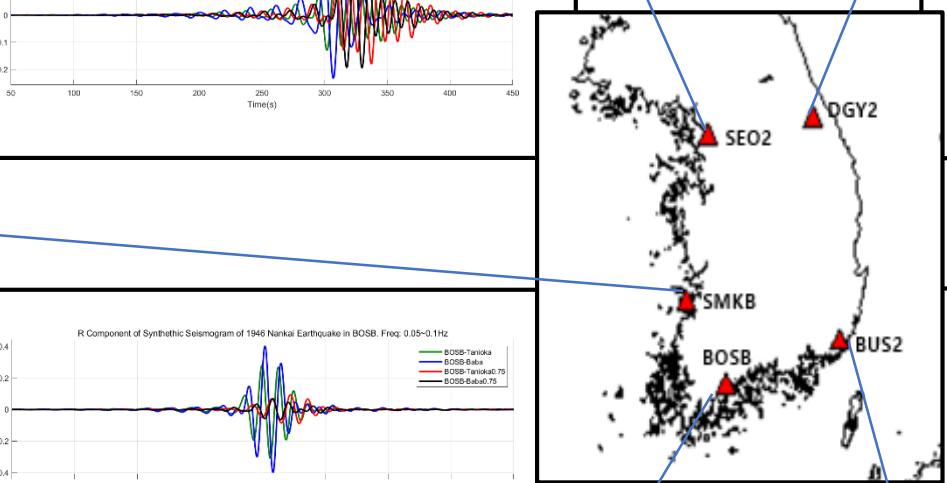
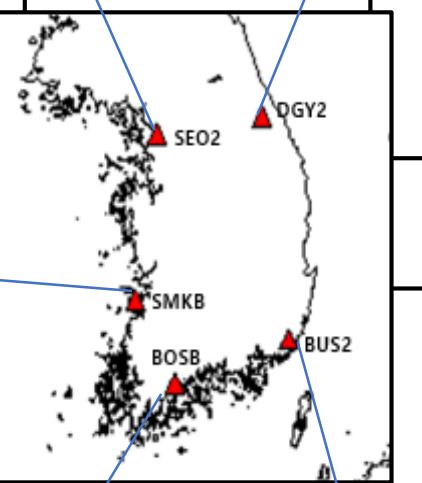
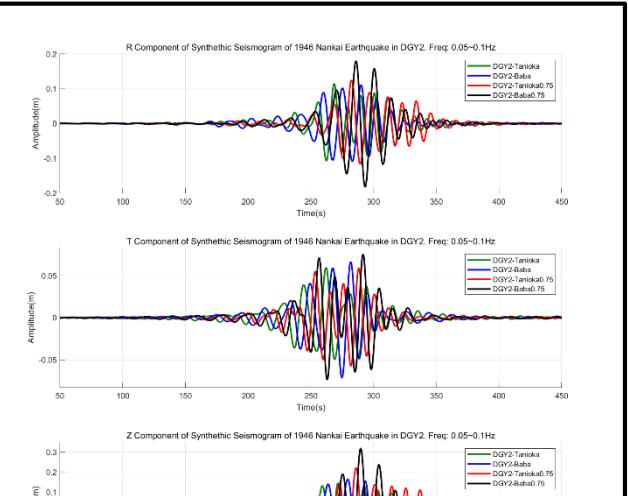
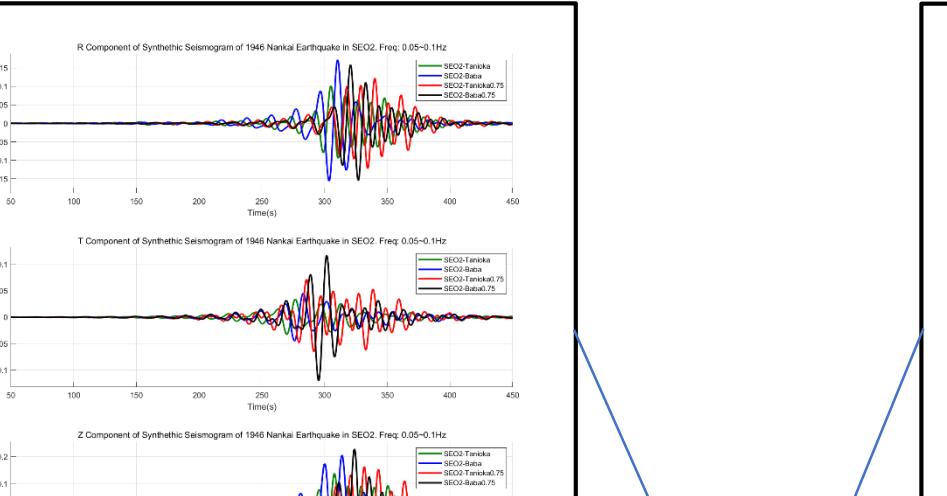
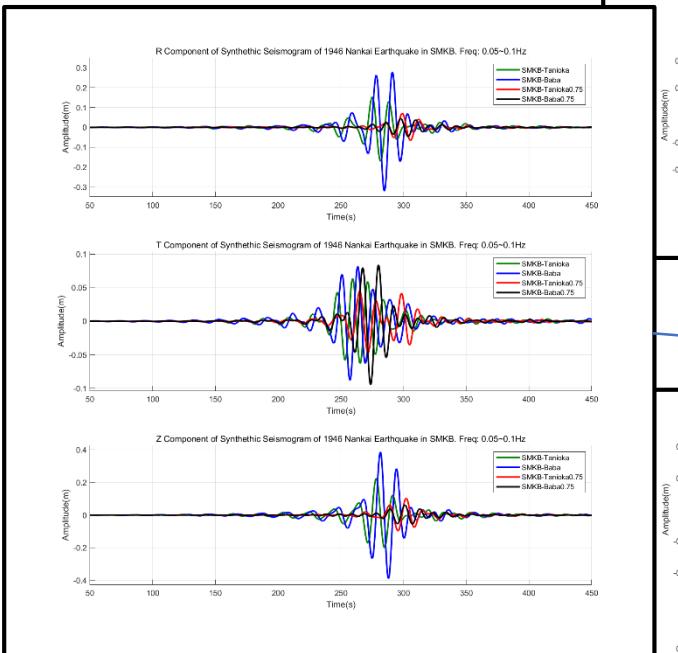
T Component of Synthetic Seismogram of 1946 Nankai Earthquake in SMKB. Freq: 0.02~0.05Hz



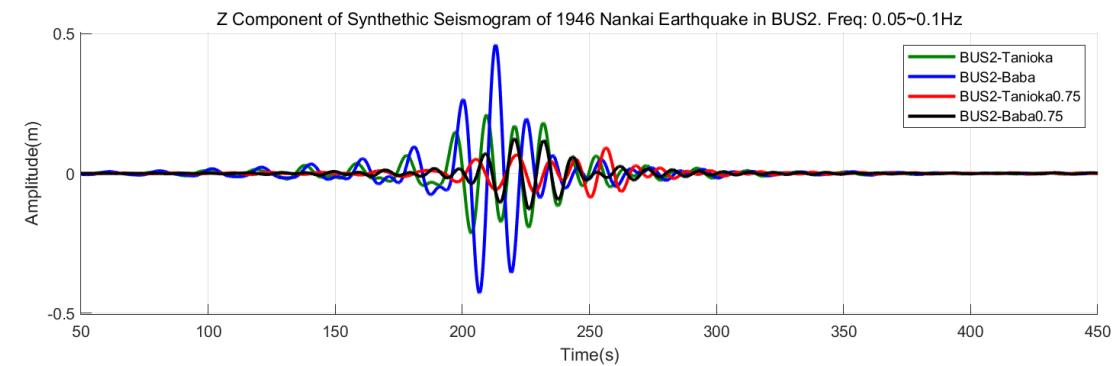
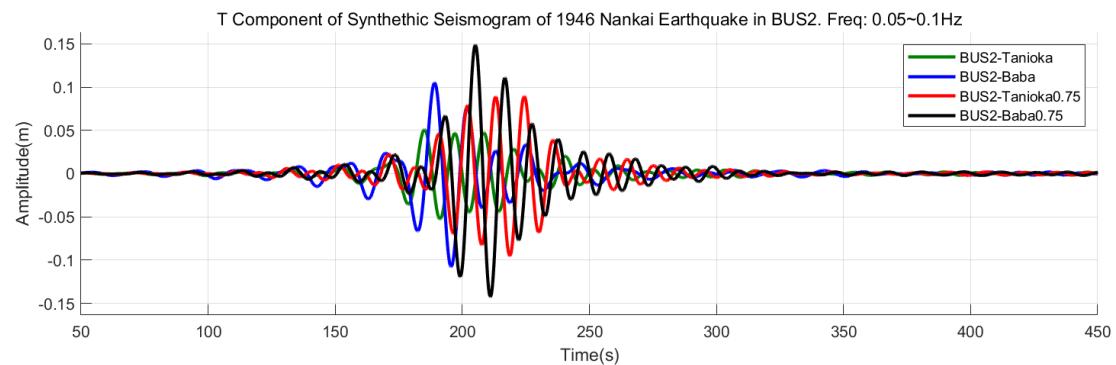
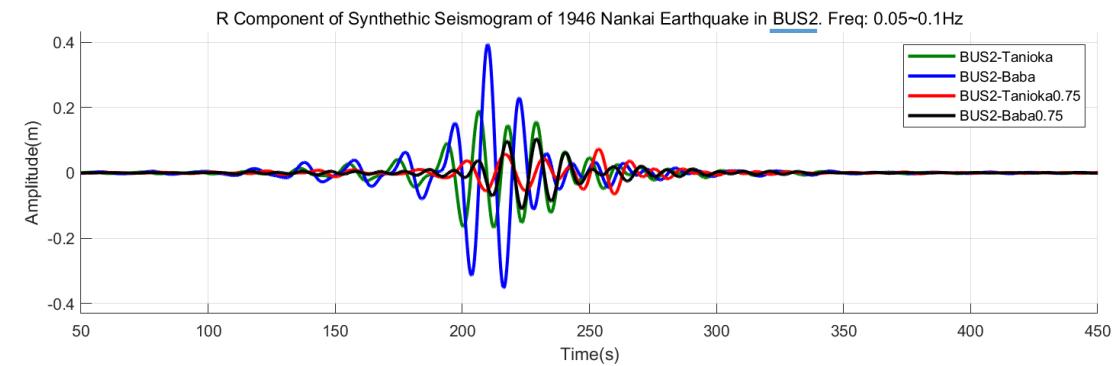
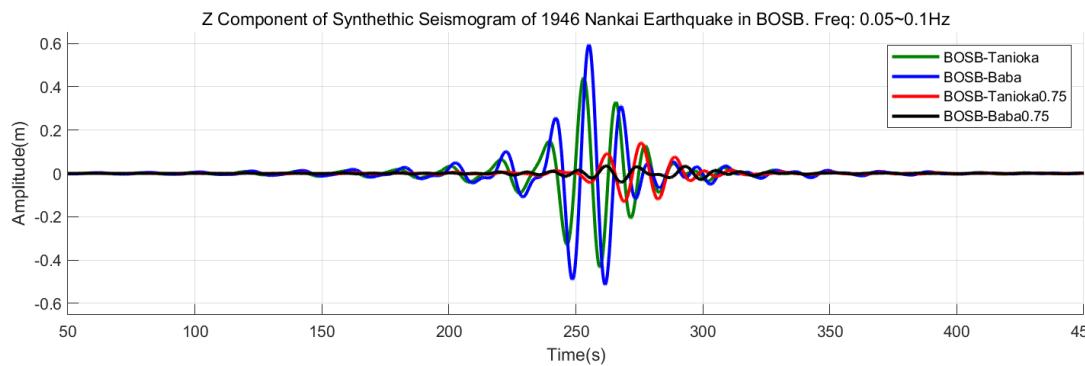
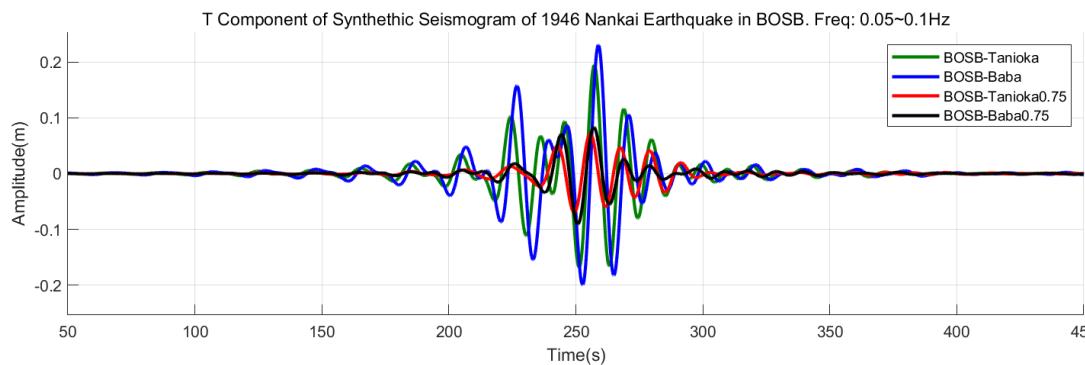
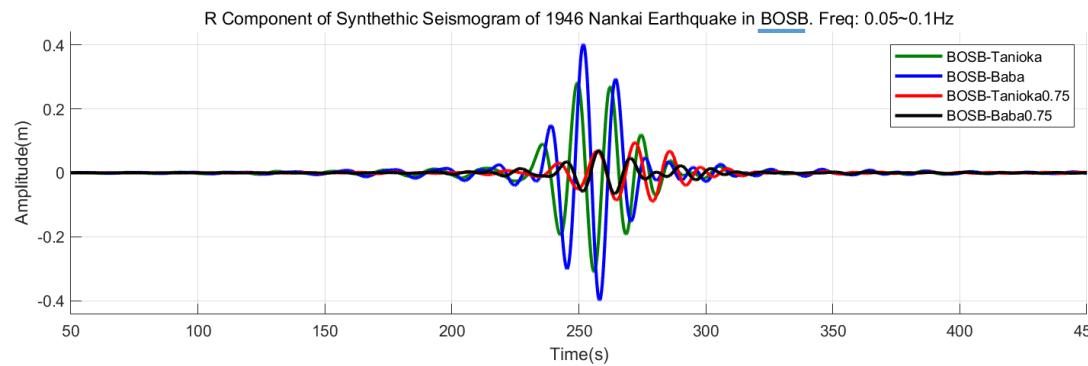
Z Component of Synthetic Seismogram of 1946 Nankai Earthquake in SMKB. Freq: 0.02~0.05Hz

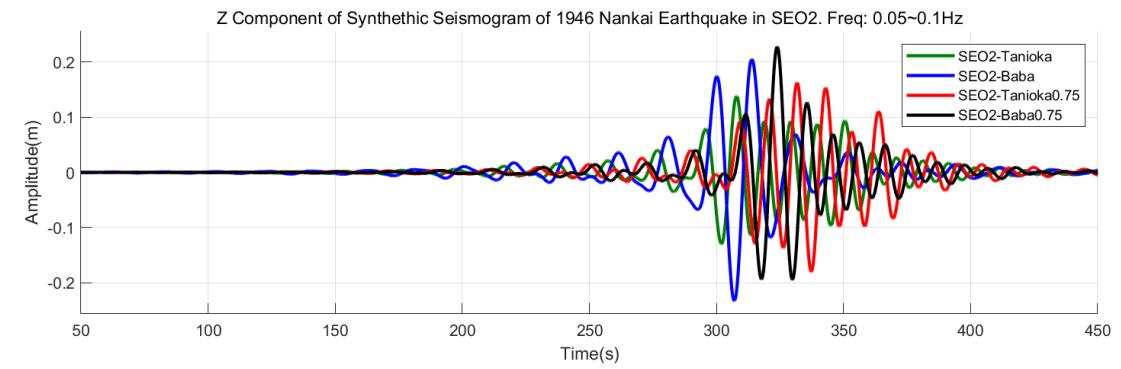
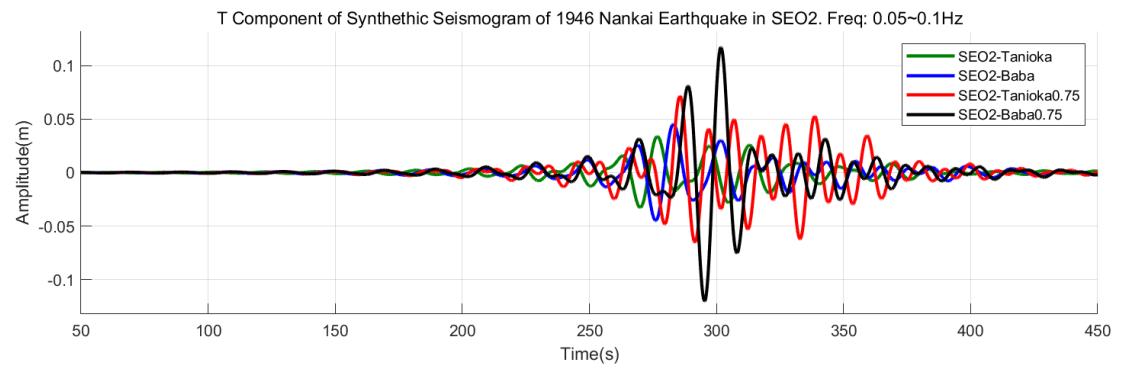
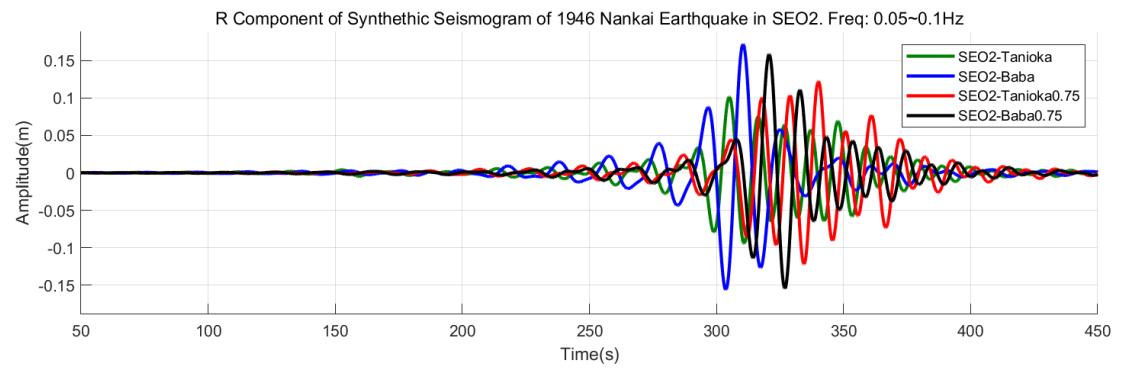
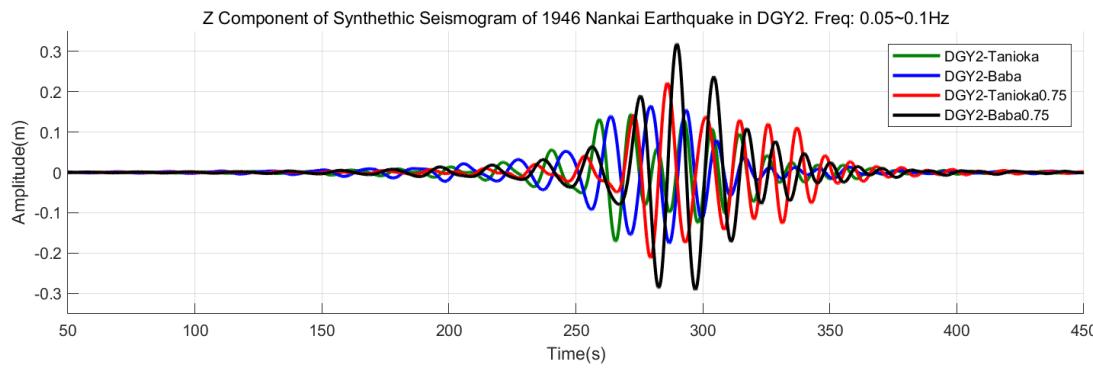
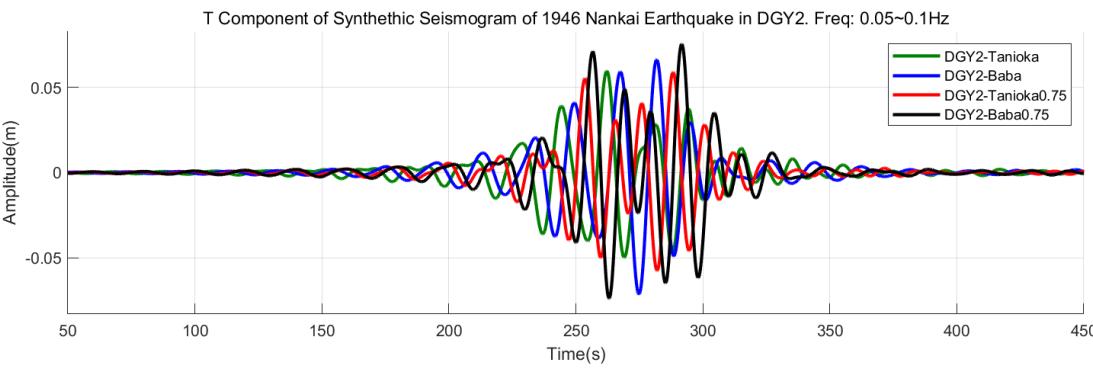
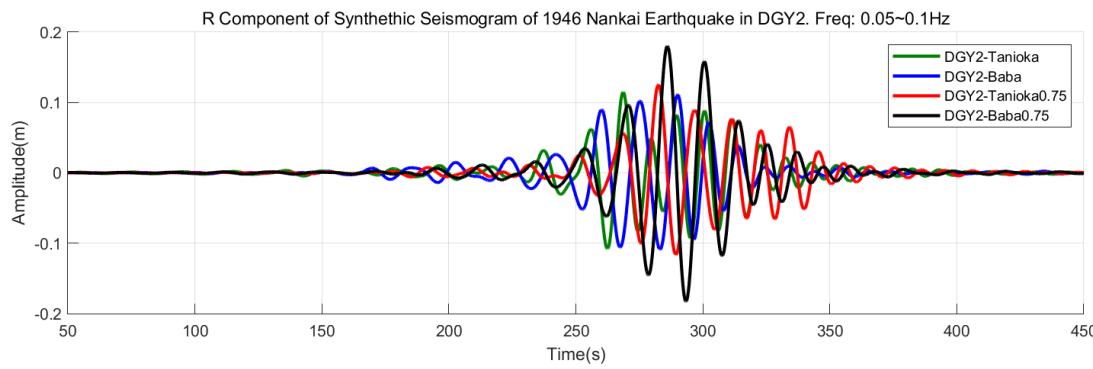


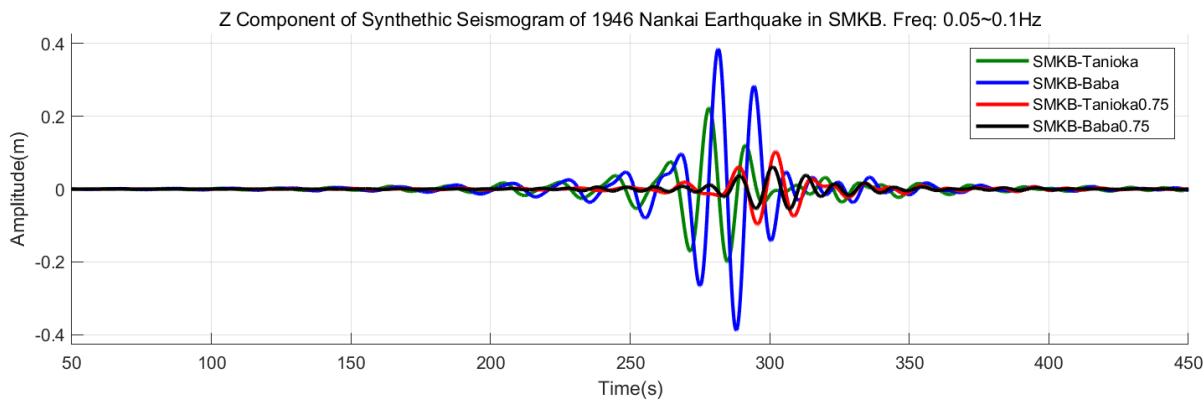
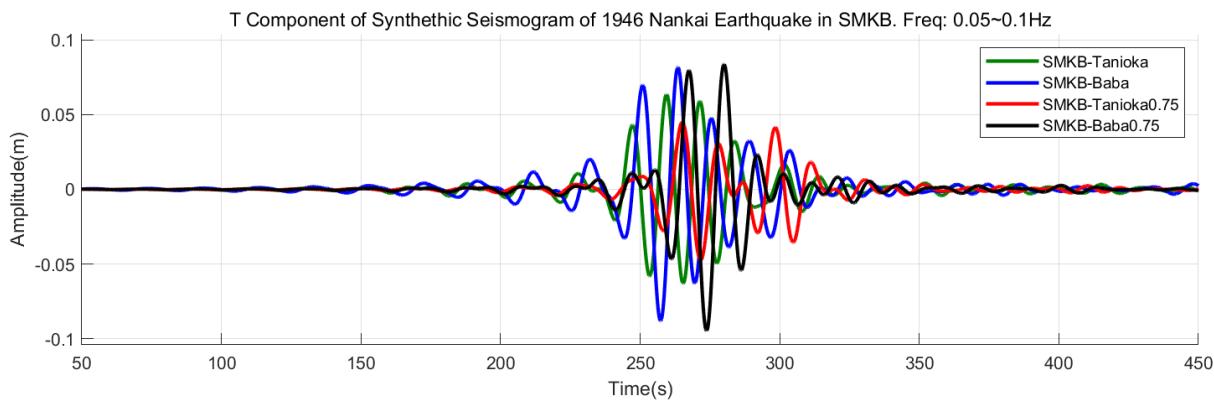
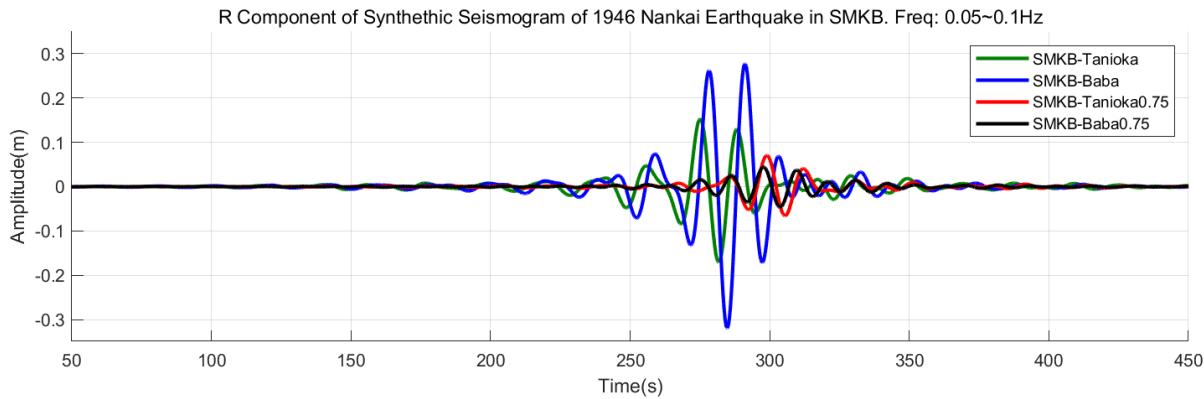
# Synthetic Seismogram Of 1946 Nankai Earthquake (Freq: 0.05~0.1Hz)



## Stations closer to event fit better







# Conclusion

- 2016 Nanaki Earthquake: East Asia Model | 1D-PREM | Observation Data Comparison
  - East Asia model:
    - 0.02Hz~0.05Hz Body Wave & Love Wave: **fit well**
    - High Frequency & Rayleigh Wave: **not fit so well**
  - 1D-PREM:
    - High Frequency Surface Wave: **fit well**
    - Arrive faster than EA
- 1946 Nankai Earthquake: Different Finite Fault Model & Different Rupture Velocity Comparison
  - **Kinematic parameters** are important
- Future Research:
  - **Limitation of kinematic parameters**  
-pseudo-dynamic source model based on statistical analysis on kinematic source parameters of dynamic rupture models
  - **Quantitative analysis** of synthetic seismograms

Table 1. A set of model parameters for pseudo-dynamic source modelling.

Model parameter	Description
1-Point Statistics	$\mu_{\text{slip}}$ Mean slip
	$\mu_{V_r}$ Mean rupture velocity
	$\mu_{V_{\max}}$ Mean peak slip velocity
	$\sigma_{\text{slip}}$ Standard deviation of slip
	$\sigma_{V_r}$ Standard deviation of rupture velocity
	$\sigma_{V_{\max}}$ Standard deviation of peak slip velocity
2-Point Statistics	$a_x$ Correlation length in the along-strike direction (six parameters: slip versus slip, slip versus $V_r$ , slip versus $V_{\max}$ , $V_r$ versus $V_r$ , $V_r$ versus $V_{\max}$ and $V_{\max}$ versus $V_{\max}$ )
	$a_z$ Correlation length in the along-dip direction (six parameters: slip versus slip, slip versus $V_r$ , slip versus $V_{\max}$ , $V_r$ versus $V_r$ , $V_r$ versus $V_{\max}$ and $V_{\max}$ versus $V_{\max}$ )
	$\rho_{\max}$ Maximum correlation coefficient (three parameters: slip versus $V_r$ , slip versus $V_r$ and $V_r$ versus $V_{\max}$ )
	$rD_x$ Response distance in the along-strike direction (three parameters: slip versus $V_r$ , slip versus $V_{\max}$ and $V_r$ versus $V_{\max}$ )
	$rD_z$ Response distance in the along-dip direction (three parameters: slip versus $V_r$ , slip versus $V_{\max}$ and $V_r$ versus $V_{\max}$ )

For autocorrelation,  $\rho_{\max}$  is one and  $rD_x$  and  $rD_z$  are zero by definition.  $rD_x$  and  $rD_z$  are excluded in this study. Thus, 21 model parameters are considered in total (6 for 1-point statistics and 15 for 2-point statistics).

(Seok Goo Song, 2016)

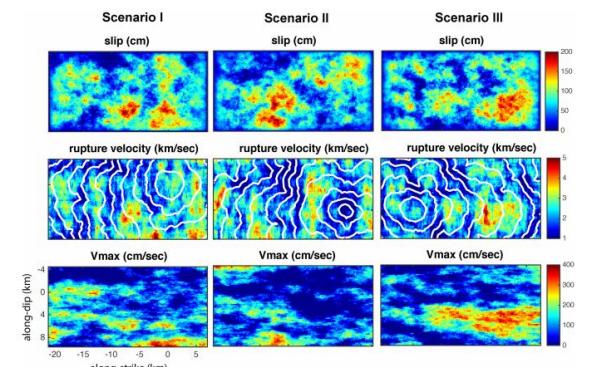


Figure 6. Source modelling examples. The first set of sampled values in Table 4 is used as input source statistics.

감사합니다.