Time dependent 3-D numerical modeling of cratonic evolution



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Origin of problem

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More than 65-70% of continental crust was formed around 3 Ga and only 5% of them are preserved today as cratons (Blue areas in Fig.2). How did these cratons survive for billions of years?

Only compositional density can not provide long-term stability of cratons; they need to be mechanically strong or highly viscous.

Estimate of craton viscosity remains elusive.

Earlier findings from instantaneous models

Paul et al. (2019) developed instantaneous mantle convection models driven by density anomalies obtained from tomography. Additionally, they used free slip boundary condition at the surface and core-mantle boundary.

Their results showed that with increasing lithospheric thickness intensity of deformation decreases (Fig. 1). This could potentially stabilize cartons (thickness > 180 km). Moreover, they also found asthenosphere viscosity plays an important role in cratons' survival.

Goals of present study: 1. Do we observe similar deformation pattern In time-dependent models?

- 2. Does asthenosphere viscosity play a significant role in craton survival?
- 3. What is the minimum requirement of viscosity for cratons and asthenosphere?
- O.5

 180 km)

 10.5

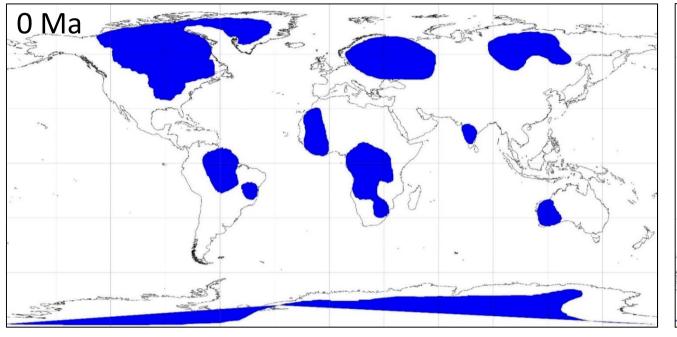
 180 km)

 1

Fig. 1: Deformation at the base of lithosphere

Present approach using time-dependent models ©

We **reconstruct** cratons' locations till 409 Ma using finite rotation of Euler poles using Gplates (Fig. 2). CitcomS is used to develop time-dependent mantle convection **models from 409 Ma to present day** using velocity (obtained from GPlates) as surface boundary condition at every 1 Myr. Few results of cratonic evolution are shown below (Figs. 4, 5).



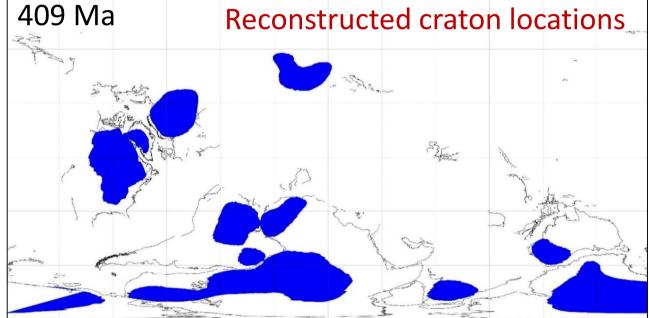


Fig. 2: Present day locations and reconstructed location of cratons

Viscosity of mantle models

Both radial and lateral viscosity variation is incorporated in convection models (Fig. 3) and 9 models of different viscosity combination are produced. Craton viscosity vary from 10-1000 times the surroundings;

asthenosphere

viscosity ranges

Depth (km)

Lithosphere

Asthenosphere

Upper mantle
Transition zone

Lower Mantle

Craton
(10-1000 times ambient layer viscosity)

Relative viscosity
(0.01-1)

(10-1000 times ambient layer viscosity)

Relative viscosity
(10-1000 times ambient layer viscosity)

Relative viscosity
(10-1000 times ambient layer viscosity)

Relative viscosity
(10-1000 times ambient layer viscosity)

(10-1000 times ambient layer viscosity
(10-1000 times ambient layer viscosity)

Relative viscosity
(10-1000 times ambient layer viscosity)

Lithosphere
(10-1000 times ambient layer viscosity)

Asthenosphere
(10-1000 times ambient layer viscosity)

Asthenosphere
(10-1000 times ambient layer viscosity)

Lithosphere
(10-1000 times ambient layer viscosity)

between 10¹⁹ – 10²¹
Pa-s.

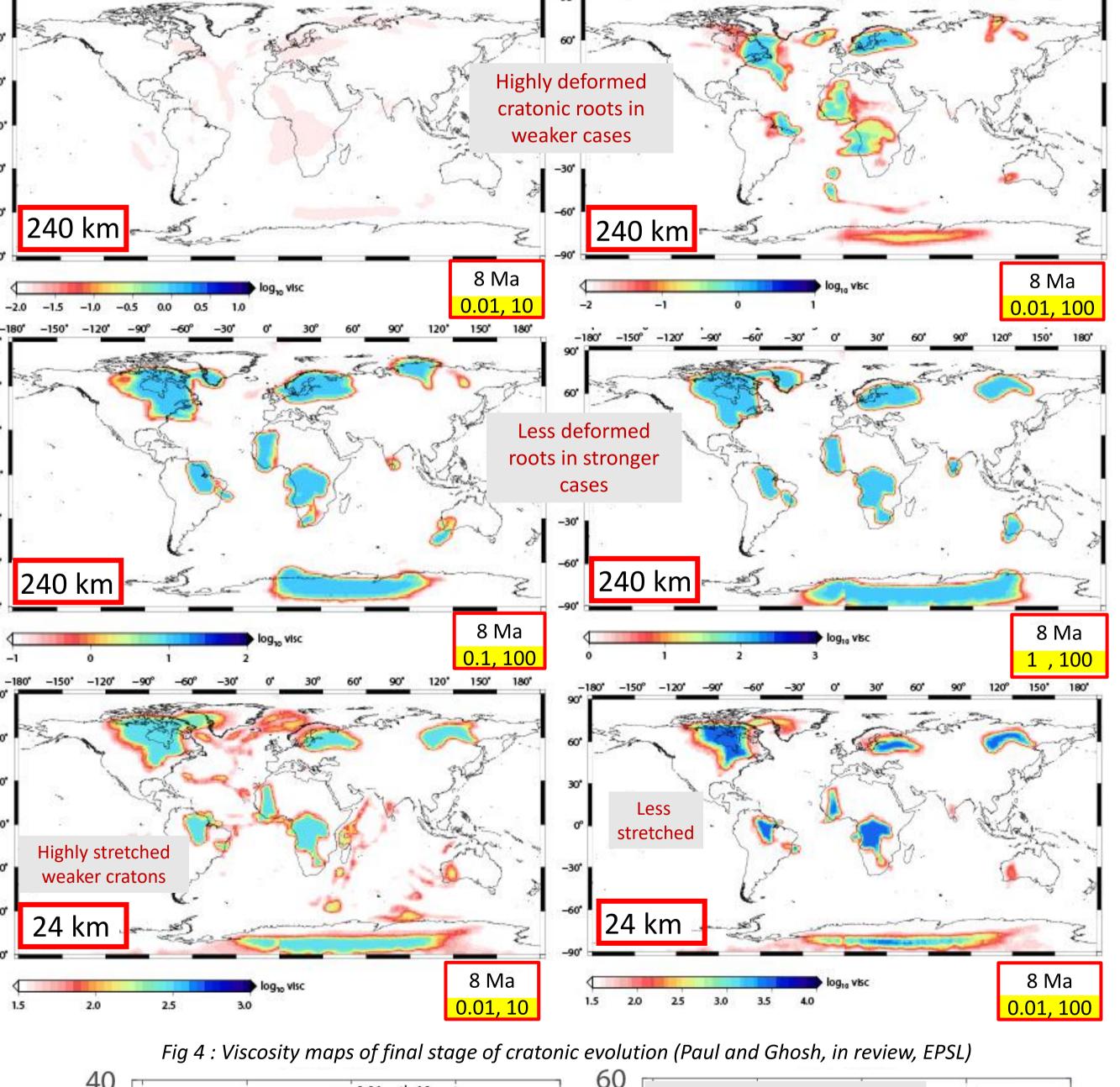
Relative viscosity is multiplied with

reference viscosity of

upper mantle

Fig. 3: viscosity structure of mantle model

Cratonic evolution with time in different models



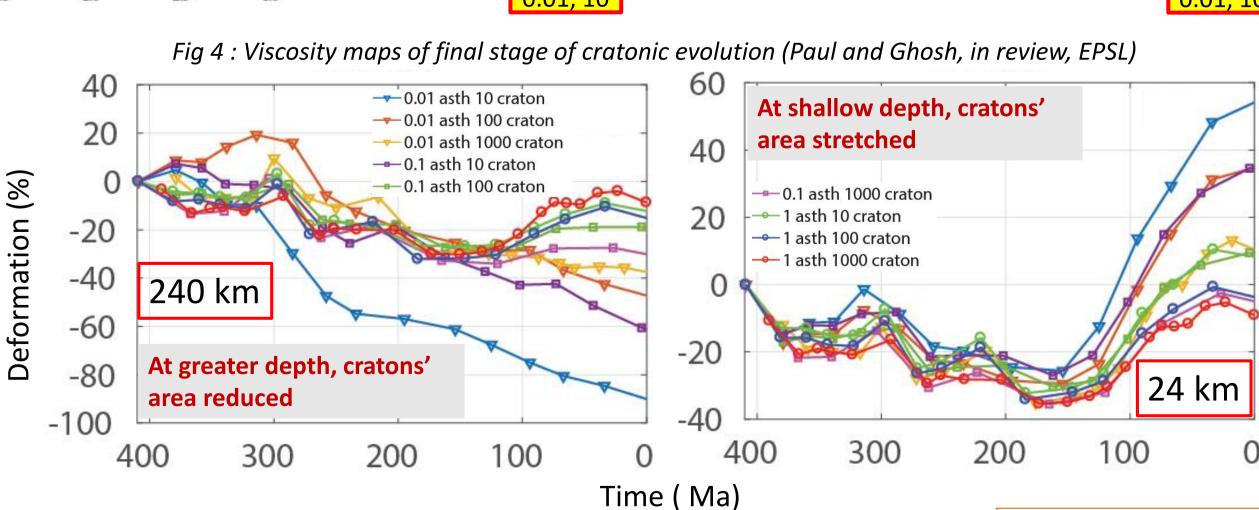


Fig 5: Deformation of cratons with time (Paul and Ghosh, in review, EPSL)

Deformation (D) = 100 \times (1 - Area of craton at 409 Ma (A_o)

Area of craton at time t (A_t)

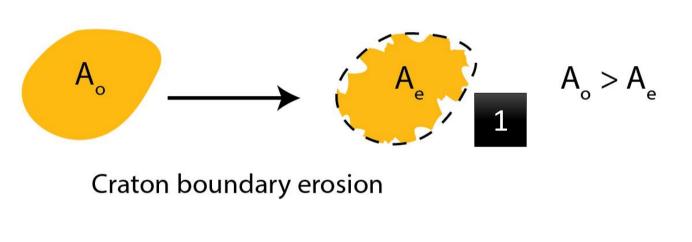
boxes indicate
asthenosphere and
craton viscosity,
respectively.

Numbers within

yellow highlighted

D > 0, extension of cratons area; D < 0, reduction of cratonic area

Mechanism of deformation



Ao: Initial area of craton at 409 Ma.

Ae: Area of craton after boundary erosion

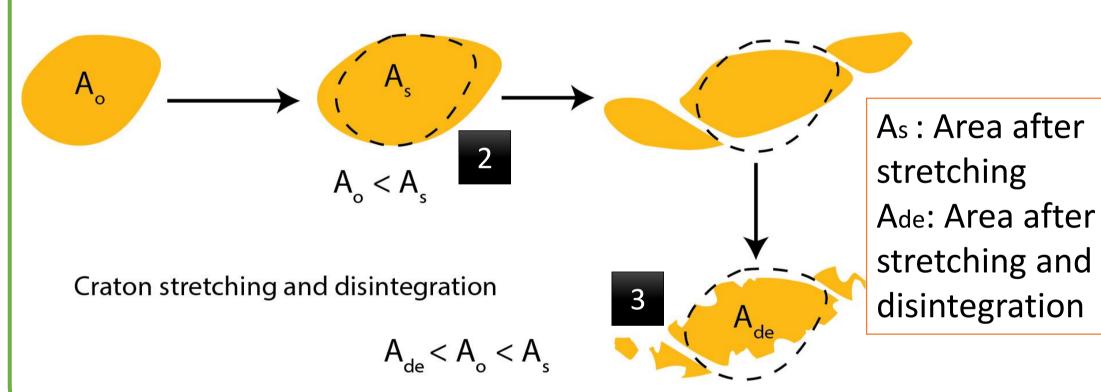


Fig 6. Schematic diagram of craton deformation (Paul and Ghosh, in review, EPSL)

Viscosity combination for cratons'

(.1,10) (.1,100) (.1,1000)

Viscosity combination of asthenosphere and cratons in models

Deformation of cratons at the present day estimated from all models at different depths.

10 times viscous cratons surrounded by 10^{19} Pa-s asthenosphere have undergone ~90% reduction of roots and ~70% stretching near surface within 410 Myrs. This combination of viscosity can not support long term survival of cratons.

With increasing viscosity deformation decreases and cratons can potentially survive for billions of years.

Fig 7. Final deformation of cratons at different depths (Paul and Ghosh, in review, EPSL)

(1,100) (1,1000)

Conclusions



- 1. Survival of cratons is dependent on the **viscosity of cratons** and as well as that of the **surrounding asthenosphere**.
- 2. Minimum requirement of viscosity: asthenosphere viscosity 10²⁰ Pa-s, cratons 100 times more viscous than surrounding.
- 3. Cratons having the above viscosity structure are at least able to survive till 410 myrs.

Reference: 1. J. Paul, A. Ghosh, C. P. Conrad. Traction and strain-rate at the base of the lithosphere: An insight into cratonic survival, *Geophysical Journal International*, 2019.

2. J. Paul and A. Ghosh,.: Evolution of cratons through ages: A time-dependent study, *Earth and Planetary Science Letters, In review.*

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