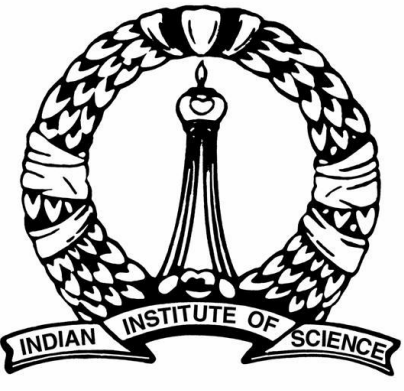


# Time dependent 3-D numerical modeling of cratonic evolution



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## Origin of problem

More than **65-70% of continental crust** was formed around **3 Ga** and **only 5%** of them are preserved today as **cratons** (Blue areas in Fig.2). How did these cratons survive for billions of years?

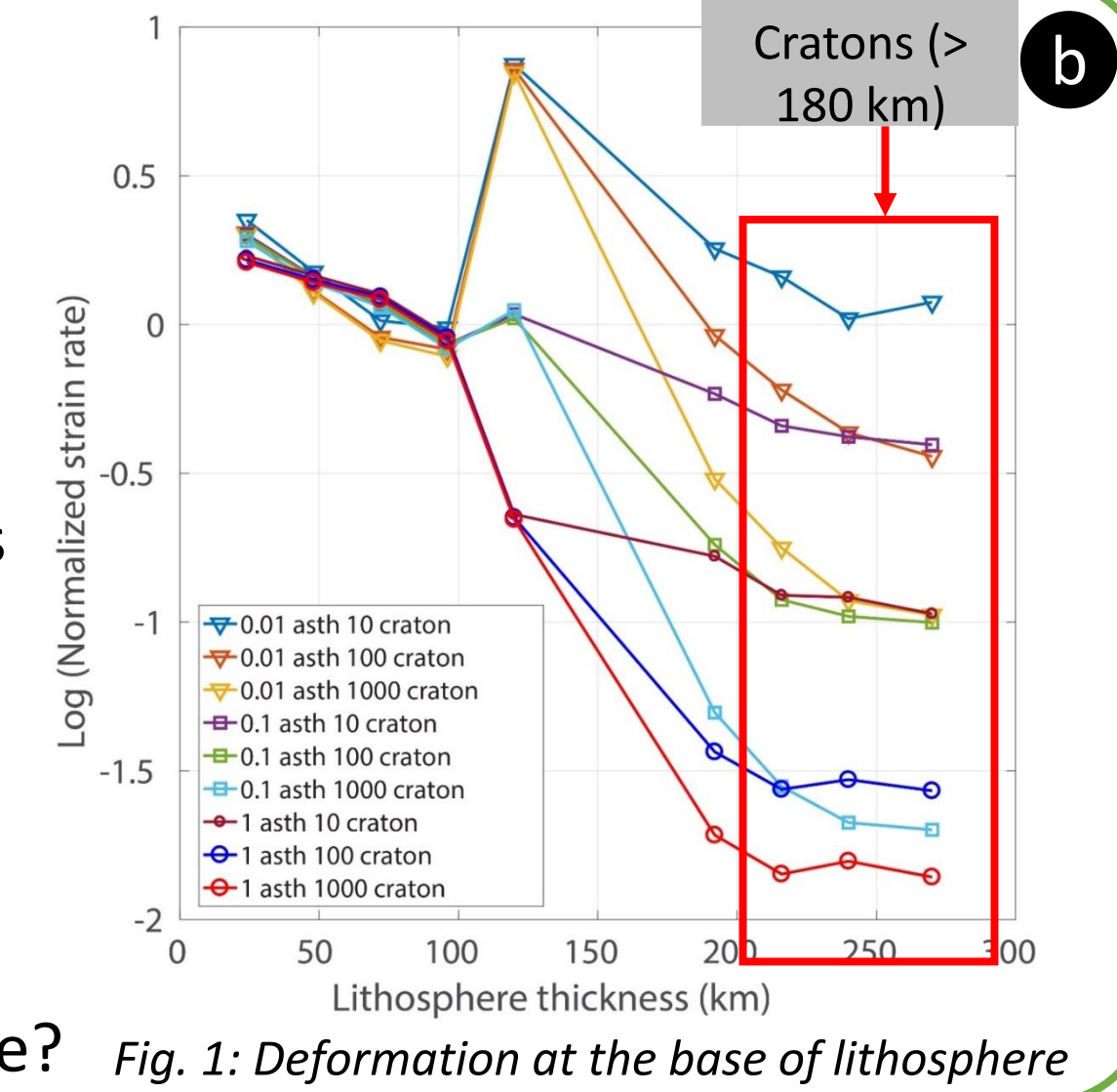
Only **compositional density can not provide** long-term stability of cratons; they need to be **mechanically strong or highly viscous**.

Estimate of craton viscosity remains elusive.

## Earlier findings from instantaneous models

Paul et al. (2019) developed instantaneous mantle convection models driven by density anomalies obtained from tomography. Additionally, they used free slip boundary condition at the surface and core-mantle boundary. Their results showed that with **increasing lithospheric thickness** intensity of **deformation decreases** (Fig. 1). This could potentially stabilize cratons (thickness > 180 km). Moreover, they also found **asthenosphere viscosity** plays an important role in cratons' survival.

**Goals of present study:** 1. Do we observe similar **deformation pattern** in time-dependent models?  
2. Does **asthenosphere viscosity** play a significant role in craton survival?  
3. What is the **minimum requirement of viscosity** for cratons and asthenosphere?



## Present approach using time-dependent models

We **reconstruct** cratons' locations till 409 Ma using finite rotation of Euler poles using Gplates (Fig. 2). CitcomS is used to develop time-dependent mantle convection **models from 409 Ma to present day** using velocity (obtained from GPlates) as surface boundary condition at every 1 Myr. Few results of cratonic evolution are shown below (Figs. 4, 5).

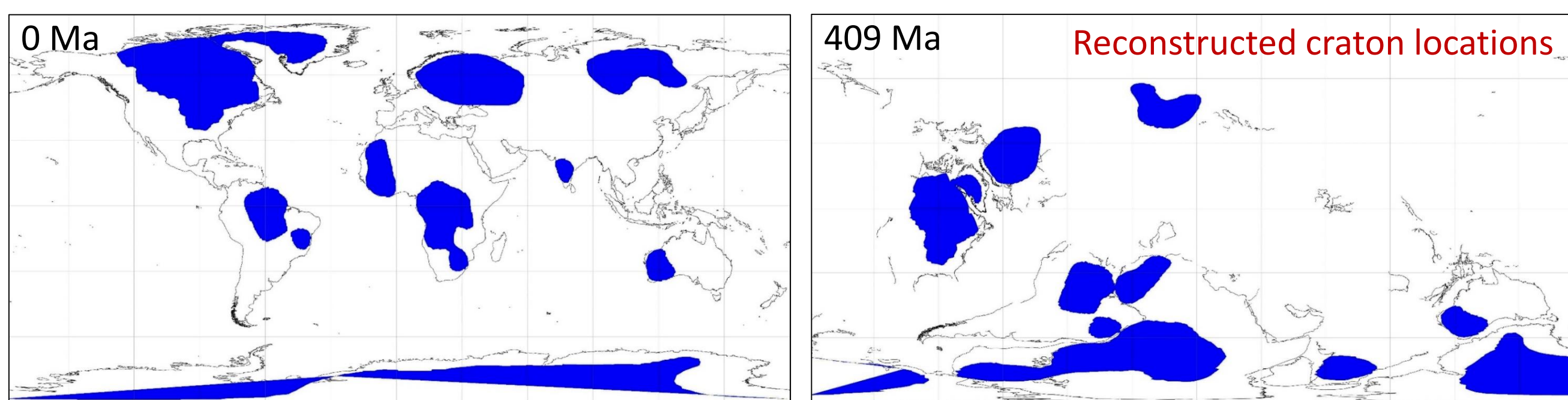


Fig. 2: Present day locations and reconstructed location of cratons

## Viscosity of mantle models

Both **radial and lateral viscosity variation** is incorporated in convection models (Fig. 3) and 9 models of different viscosity combination are produced. **Craton viscosity** vary from **10-1000** times the surroundings;

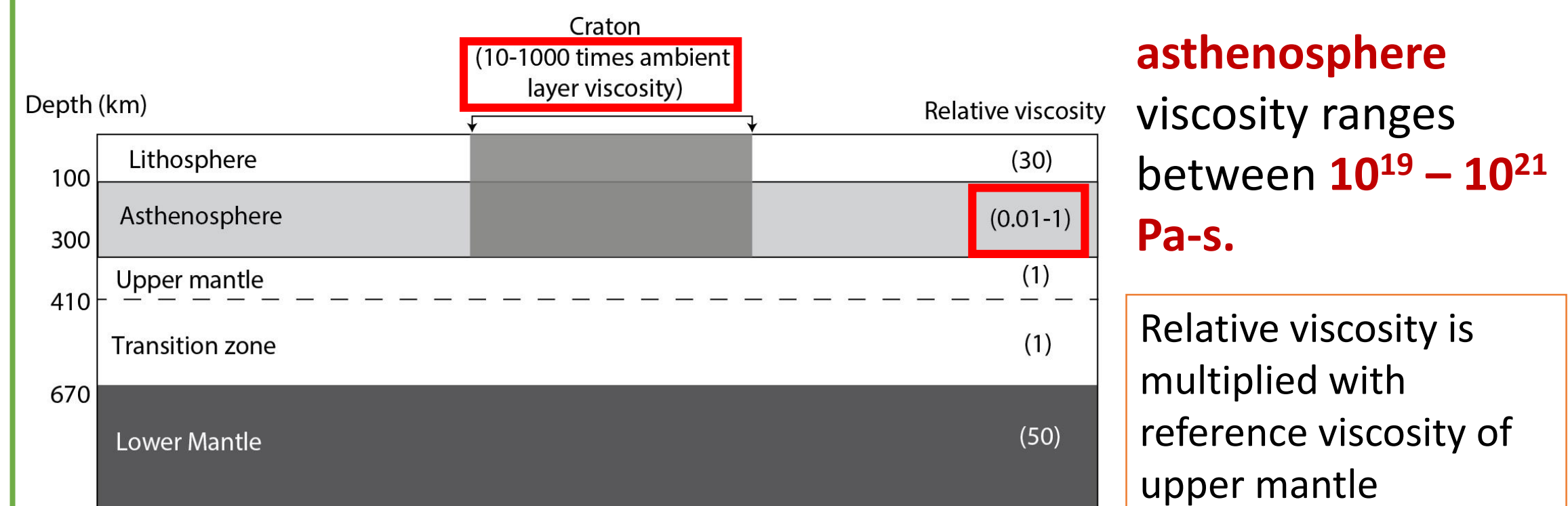


Fig. 3: viscosity structure of mantle model

## Cratonic evolution with time in different models

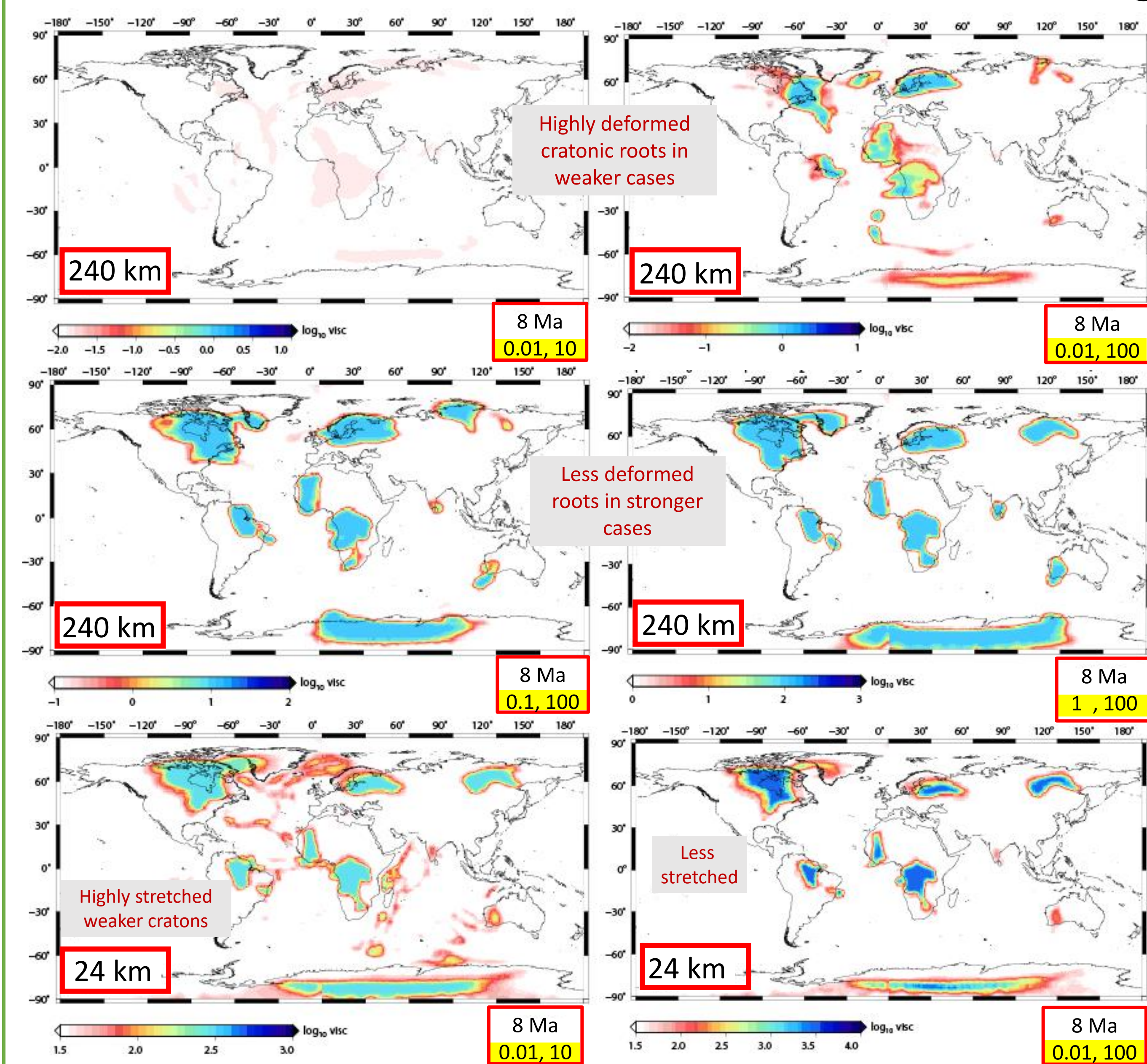


Fig 4 : Viscosity maps of final stage of cratonic evolution (Paul and Ghosh, in review, EPSL)

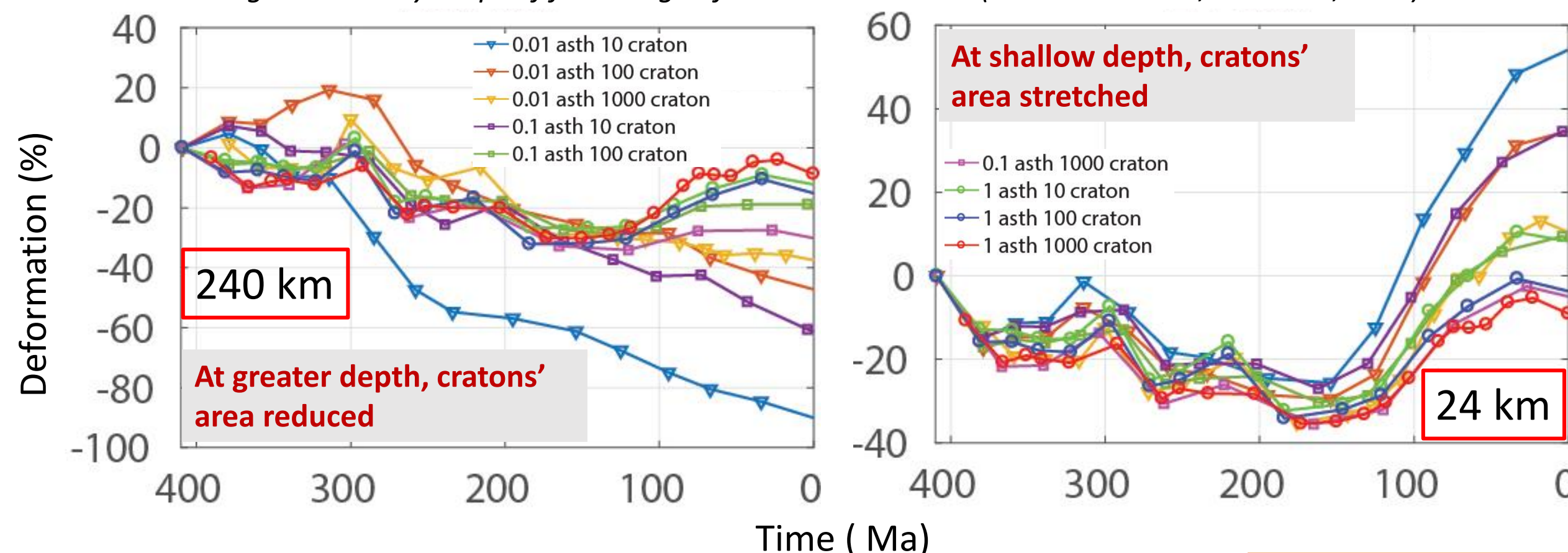


Fig 5: Deformation of cratons with time (Paul and Ghosh, in review, EPSL)

$$\text{Deformation (D)} = 100 \times \left( 1 - \frac{\text{Area of craton at 409 Ma (A}_o\text{)}}{\text{Area of craton at time t (A}_t\text{)}} \right)$$

D > 0, extension of cratons area; D < 0, reduction of cratonic area

Numbers within yellow highlighted boxes indicate asthenosphere and craton viscosity, respectively.

## Mechanism of deformation

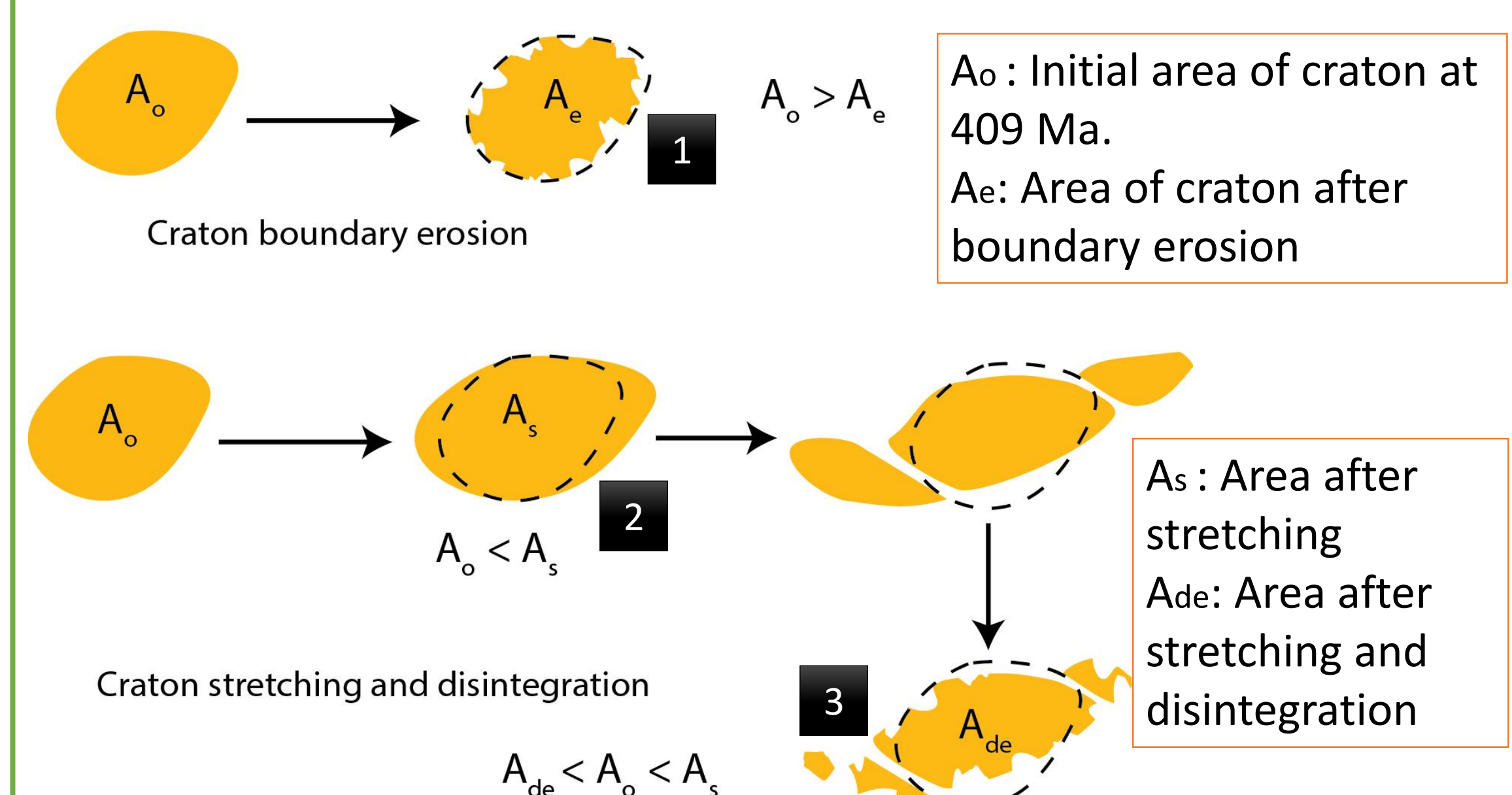


Fig 6. Schematic diagram of craton deformation (Paul and Ghosh, in review, EPSL)

## Viscosity combination for cratons' survival

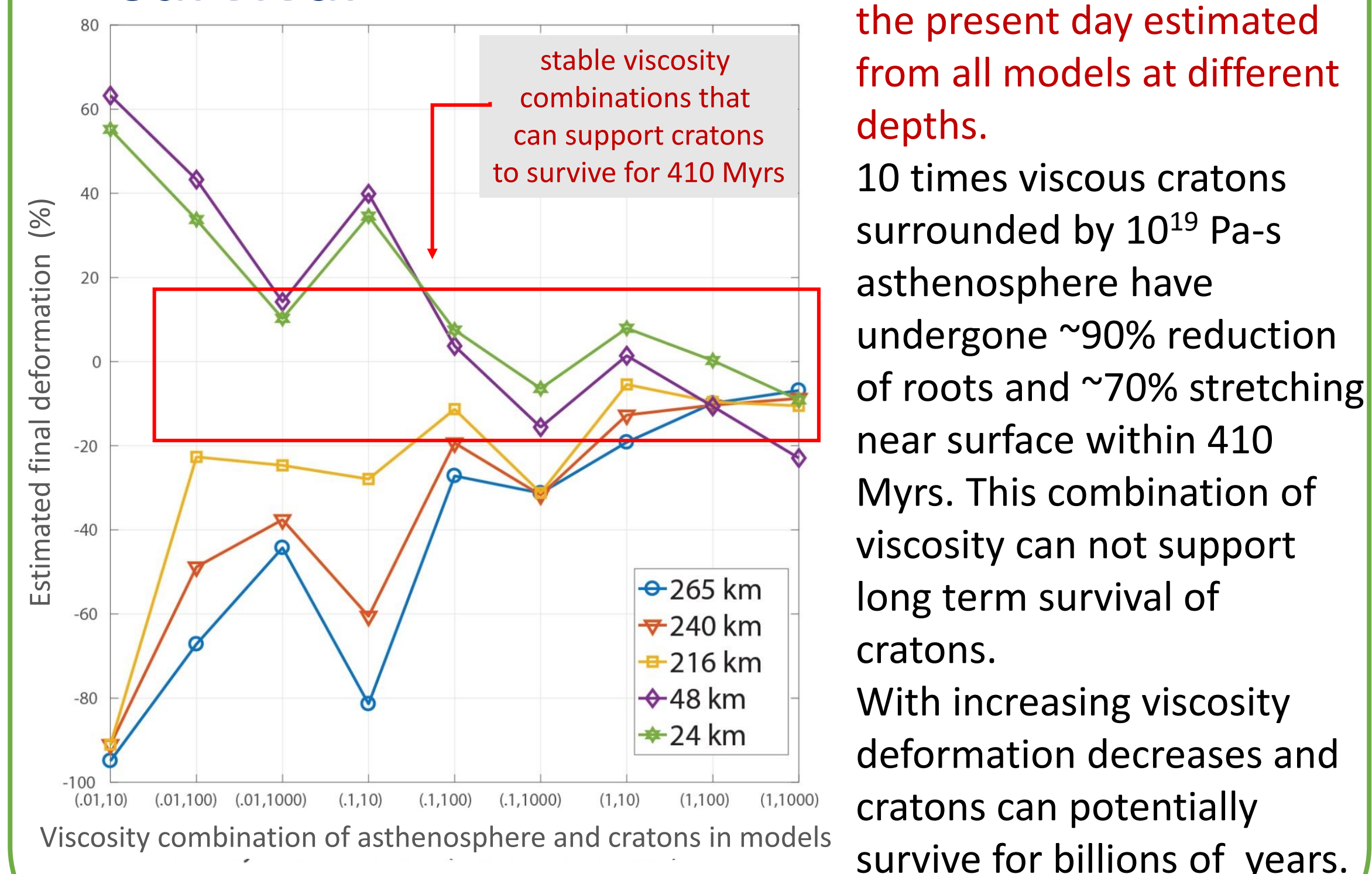


Fig 7. Final deformation of cratons at different depths (Paul and Ghosh, in review, EPSL)

## Conclusions

- Survival of cratons is dependent on the **viscosity of cratons** and as well as that of the **surrounding asthenosphere**.
- Minimum requirement of viscosity: **asthenosphere** viscosity **10<sup>20</sup> Pa-s**, **cratons 100 times** more viscous than surrounding.
- Cratons having the above viscosity structure are at least able to survive till 410 myrs.