Simpson-Hunter Mechanism Parameterization

Core Components and Initialization

The SimpsonHunterMechanismParam class, implemented in C#, models physical processes related to tidal straining and internal tides in estuarine circulation, focusing on wave-induced Stokes drift, internal tide effects, and tidal straining (Simpson-Hunter mechanism). The class is initialized with the following parameters:

- Number of sigma layers: n_{σ} (vertical grid points)
- Wave amplitude: $a = 0.5 \,\mathrm{m}$ (default)
- Wave period: $T_w = 10.0 \,\mathrm{s}$ (default)
- Internal tide amplitude: $A_{it} = 0.05 \,\mathrm{m/s}$ (default)
- Simpson-Hunter coefficient: $C_{sh} = 0.1$ (default, for tidal straining)
- Gravitational acceleration: $g = 9.81 \,\mathrm{m/s^2}$
- Reference density: $\rho_0 = 1000.0 \,\mathrm{kg/m^3}$

The constructor sets these parameters, allowing customization of wave, internal tide, and tidal straining effects.

Functioning Logic

The class provides three methods to compute contributions to estuarine dynamics:

- 1. ComputeStokesDrift: Calculates the x-direction Stokes drift velocity due to surface waves.
- 2. ComputeInternalTideEffect: Computes vertical velocity perturbations and turbulent kinetic energy (TKE) production from internal tides.
- 3. ComputeTidalStraining: Evaluates the tidal straining effect (Simpson-Hunter mechanism) on the salinity gradient.

Stokes Drift Computation

The ComputeStokesDrift method calculates the Stokes drift velocity in the x-direction using a simplified shallow-water approximation:

- Wave number: $k=\frac{2\pi}{T_w\sqrt{gh}}$, where h is the water depth.
- Depth coordinate: $z = \sigma h$, where σ is the sigma coordinate (0 at bottom, 1 at surface).
- Stokes drift velocity:

$$u_s = a \cdot \frac{2\pi}{T_w} \cdot e^{-2kz} \cdot \sin(\phi_t) \tag{1}$$

where a is the wave amplitude, ϕ_t is the tidal phase, and the velocity is capped: $u_s \in [-0.5, 0.5] \text{ m/s}$ for numerical stability.

Internal Tide Effect

The ComputeInternalTideEffect method computes vertical velocity perturbations and TKE production due to internal tides:

• Buoyancy frequency squared:

$$N^{2} = -\frac{g}{\rho_{0}} \frac{\partial \rho}{\partial z}, \quad \frac{\partial \rho}{\partial z} = \frac{\rho_{2} - \rho_{1}}{\Delta \sigma \cdot h}$$
 (2)

where $\rho_k = \rho_0 + 0.8S_k$, and S_k is the salinity at layer k. N^2 is capped: $N^2 \in [0, 10^{-3}] \, \mathrm{s}^{-2}$.

· Vertical velocity perturbation:

$$w_{\text{tide}} = A_{\text{it}} \sin(\phi_t) \sqrt{N^2} \cos\left(\frac{2\pi z}{h}\right)$$
 (3)

where $w_{\text{tide}} \in [-0.2, 0.2] \,\text{m/s}$.

TKE production:

TKE =
$$0.1A_{\rm it} \left(\frac{\partial w}{\partial z}\right)^2$$
, $\frac{\partial w}{\partial z} = \frac{w_k - w_{k+1}}{\Delta \sigma \cdot h}$ (4)

TKE is capped: TKE $\in [0, 10^{-4}] \, \mathrm{m}^2/\mathrm{s}^3$. No effect is applied at the top layer $(k = n_\sigma - 1)$.

Tidal Straining Effect

The ComputeTidalStraining method models the Simpson-Hunter mechanism, which describes how tidal currents modify salinity gradients:

Salinity gradient:

$$\frac{\partial S}{\partial x} = \begin{cases}
\frac{S_{i+1,k} - S_{i-1,k}}{2\Delta x} & \text{if } 0 < i < n_{\text{cells}} - 1 \\
\frac{S_{i+1,k} - S_{i,k}}{\Delta x} & \text{if } i = 0 \\
\frac{S_{i,k} - S_{i-1,k}}{\Delta x} & \text{if } i = n_{\text{cells}} - 1
\end{cases}$$
(5)

where Δx is the average cell width, and $\frac{\partial S}{\partial x} \in [-10, 10] \, \text{PSU/m}$.

• Tidal straining term:

Straining =
$$-C_{\rm sh}u_t\frac{\partial S}{\partial x}$$
 (6)

where u_t is the tidal velocity, and the term is capped: Straining $\in [-1, 1]$ PSU/s. No effect is applied at the top layer or boundary cells.

Physical and Mathematical Models

The SimpsonHunterMechanismParam class employs the following models:

· Stokes Drift:

$$u_s = a \cdot \frac{2\pi}{T_w} \cdot e^{-2kz} \cdot \sin(\phi_t), \quad u_s \in [-0.5, 0.5]$$
 (7)

• Internal Tide:

$$N^{2} = -\frac{g}{\rho_{0}} \frac{\rho_{2} - \rho_{1}}{\Delta \sigma \cdot h}, \quad N^{2} \in [0, 10^{-3}]$$
(8)

$$w_{\text{tide}} = A_{\text{it}} \sin(\phi_t) \sqrt{N^2} \cos\left(\frac{2\pi z}{h}\right), \quad w_{\text{tide}} \in [-0.2, 0.2]$$
 (9)

TKE =
$$0.1A_{it} \left(\frac{w_k - w_{k+1}}{\Delta \sigma \cdot h} \right)^2$$
, TKE $\in [0, 10^{-4}]$ (10)

Tidal Straining:

Straining =
$$-C_{\text{sh}}u_t\frac{\partial S}{\partial x}$$
, $\frac{\partial S}{\partial x} \in [-10, 10]$, Straining $\in [-1, 1]$ (11)

These models capture wave-induced transport, internal tide-driven vertical mixing, and tidal straining effects on stratification, enhancing the realism of estuarine circulation simulations.