

# Clustering solution for flux estimates based on gas bubble stream observations in single beam echo-sounder data using gridded averaging

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## Introduction

Accurate and reliable quantification of underwater gas seepage is of great importance for environmental research as well as for monitoring various sub-sea human activities. Our knowledge about the sensitivity and extent of natural seabed seepage is still very limited and even though human underwater oil and gas exploitation shall decrease in the near future, existing infrastructure as well as new industries considering e.g. seabed carbon capture projects need extensive monitoring methods for leak detection. Inversion tools such as the FlareHunter and ESP-3 which makes it possible to extract seabed gas fluxes using the acoustic backscatter data from single and multi beam echo-sounder data are crucial in this perspective since it provides a method which can cover large areas in a relatively short period of time. Single beam echosounder systems are also relatively cost effective and of limited payload which might make these a viable option for autonomous vehicles. There are, however, several challenges with the current methodology that can be discussed and potentially improved. In this document, I outline a potential caveat to the current clustering methodology used when extracting flow rates from seep sites using single beam echo-sounder data and the ESP-3/FlareHunter software [Velo et al., 2015] and presents a new clustering solution that solves the problem. Everything outlined here is available as Python (and MatLab soon) code at [https://github.com/KnutOlaD/flare\\_clustering](https://github.com/KnutOlaD/flare_clustering), the new clustering method as well.

## What is clustering and why do we cluster?

A typical singlebeam echosounder insonifies a cone-shaped volume of water where the horizontal acoustic footprint  $A_{fp}$  at a given depth  $D$  is given by

$$A_{fp} = \pi \left( D \tan\left(\frac{\theta}{2}\right) \right)^2 \quad (1)$$

where  $\theta$  is the acoustic beam opening angle of the echosounder. For the EK60, which is often used for seep detection and flow rate calculations, the opening angle of the cone is  $7^\circ$ . The resulting horizontal acoustic footprint of the echosounder can therefore be quite large at typical

depths of interest, e.g. at 220 meter the acoustic footprint area is  $A_{fp} = \pi \left(220m \tan\left(\frac{7}{2}\right)\right)^2 = 569m^2$  (see Figure 1a)<sup>1</sup>.

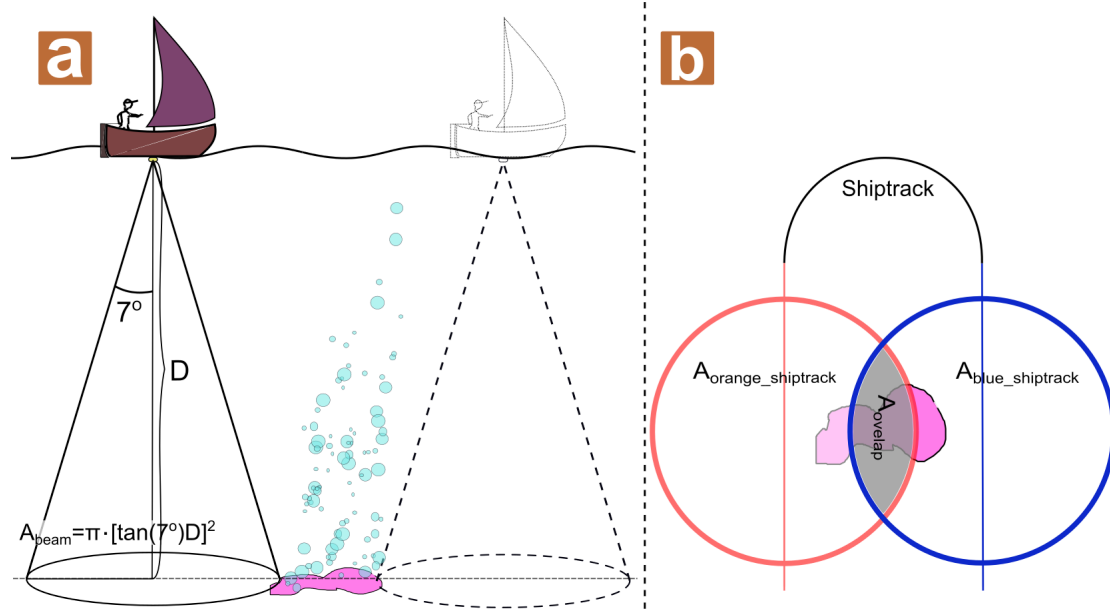


Figure 1: Conceptualized figure showing a) Insonified volume during an echosounder survey and observation of seabed seepage here as a constrained area of seepage and b) top down view of the acoustic footprint illustrating double counting of the same seepage area and an over estimate of the total flow rate during an echosounder survey.

The information used for flow rate estimates in FlareHunter and ESP-3 is the acoustic "target strength" obtained as the logarithm of the summed backscatter cross-section by scatterers in the insonified volume. This data is obtained for each ping and at all depths as given by the vertical resolution  $\Delta d$  of the echosounder. Thus, the insonified volume here refers to a truncated cone with thickness  $\Delta d$  (see illustration in Veloso et al., 2015). The method assumes that the total backscattering cross-section (TS) is a summation of all the backscattering cross-section of scatterers in the truncated cone -volume, which is, after certain post processing steps, assumed to be gas bubbles. The location of the single scatterers within the acoustic footprint of the echosounder is considered to be arbitrary, thus implying that the TS can be produced by several bubble streams with unknown locations within the footprint. In the case where flare observation samples with sufficiently overlapping footprints are obtained, i.e. as in the top-down view in Figure 1b, clustering is applied with the aim to counteract double counting and maximize probability of a best estimate of the total flow rate. In practice, flares that are clustered are treated and counted as seep clusters, instead of individual flares and gets assigned a single flow rate value which is calculated from the individual flow rates of the individual flare observations in the cluster.

<sup>1</sup>The  $596 m^2$  value does not match with  $\sim 760 m^2$  as calculated in Veloso et al., 2015 using the same input values, I don't know why..

## Vanilla clustering solution

This method is included as the `cluster_flowrate_vanilla` function in the `clustering.py` script in the github repository ([https://github.com/KnutOlaD/flare\\_clustering](https://github.com/KnutOlaD/flare_clustering))

The clustering technique suggested by Veloso et al., (2015) assigns an average flow rate to the total area of flares that are sufficiently near each other to be considered a seep cluster. In Veloso et al., (2015), any flare observation with sufficiently overlapping acoustic footprint at the seabed, defined by a threshold value, is clustered. The threshold is typically defined as a center distance of 1.8 times the radius (instead of 2 times) of the averaged acoustic footprint (of the two flare observations). This applies to all flare observations, but it is sufficient that only one flare observation has enough overlap in a cluster implying that a flare cluster can stretch over large areas and contain hundreds of flare observations that have no overlapping area (see Figure 13a in Veloso et al., 2015 and and Figure 2 in this document). Once clustered, Veloso et al., (2015) obtains the flow rate of the cluster by first calculating the average gas flux per unit area  $F_{avg}$  in the cluster as

$$F_{avg} = \frac{1}{K} \sum_{i=1}^K \frac{F_i}{A_i}, \quad (2)$$

where  $F_i$  and  $A_i$  is the flow rate and seabed footprint of flare observation  $i$  in the cluster. The total cluster flow rate  $F_{total}$  is then calculated by multiplying the average flow rate per unit area by the total area of the cluster estimated using a gridded numerical solution

$$F_{total} = F_{avg} A_{cluster} = F_{avg} N \Delta x \Delta y, \quad (3)$$

where  $N$  is the number of grid cells with cell size  $\Delta x \cdot \Delta y$  in the cluster area  $A_{cluster}$ .

## Problems with the vanilla solution?

To put forth a small discussion regarding the vanilla solution, let's consider the four seep observations presented in Figure ... and that they have estimated flowrates of  $F_1 = F_2 = F_3 = 5$  and  $F_4 = 205$  and the same seabed footprint of  $A_1 = A_2 = A_3 = A_4 = 400$ . The seep observations were obtained in a straight line with 80% overlap between  $A_1$  and  $A_2$ , 80% between  $A_2$  and  $A_3$ , and 20% overlap between  $A_3$  and  $A_4$  and were therefore considered a single seep cluster instead of individual seeps (we use areal overlap here for simplicity in the calculations and not radial as is the current norm for defining the threshold for clustering). This means that their individual flow rates were not considered and a common cluster flow rate was calculated. Following the methodology of Veloso et al., (2015) as iterated above, the total flow rate of the cluster can be calculated to

$$F_{total} = \frac{1}{4} \left( \frac{F_1}{A_1} + \frac{F_2}{A_2} + \frac{F_3}{A_3} + \frac{F_4}{A_4} \right) \cdot A_{cluster} = \frac{1}{4} \left( \frac{5}{400} + \frac{5}{400} + \frac{5}{400} + \frac{205}{400} \right) \cdot 880 = 121, \quad (4)$$

where the cluster area  $A_{cluster}$  was calculated as  $A_{cluster} = A_1 + 0.8A_1 + 0.2A_1 + 0.2A_1 = 880$  since  $A_1 = A_2 = A_3 = A_4$  and we already knew the amount of overlapping area (see Figure 2 a).

In this case, the vanilla solution obtains an average of the observations and multiplies it with the total shared area. Since the flow rates are only estimates, this might make sense - however - it is quite obvious that the total flow rate of the cluster is an underestimate. After all, considering

that it's only 20% of the area of flare observation 4 that overlaps, the remaining 80% area of this observation alone should amount to a flow rate of  $0.8 \cdot 205 = 164$ . I believe that the issue here is with violated underlying assumptions that are needed to make the vanilla solution valid. I believe a minimum set of these assumptions are

1. All flare observations in the cluster overlap the same amount (otherwise there needs to be some weighing added to account for this in the averaging operation).
2. Deviations from the cluster mean flow rate is due to random processes and/or that the central limit theorem applies.

Both these assumptions are violated in our example and will also be violated in practically all field observations unless there's only two flares in the cluster. The second assumption is also usually violated in practice. A typical sufficiently large sample size for a random process for the central limit theorem to apply is  $n \sim 30$ . In this perspective, the sample size of a typical cluster is typically too small to assume the central limit theorem.

## New clustering solution by gridded averaging

This method is included as the *cluster\_flowrate\_gridded\_averaging* function in the clustering.py script in the github repository ([https://github.com/KnutOlaD/flare\\_clustering](https://github.com/KnutOlaD/flare_clustering)).

We consider a cluster containing clustered flare observations  $k \in [1..K]$ . How the flare observations are defined as a cluster, i.e. by what technique, is not important, and can be done by the vanilla solution or by a similar technique (e.g. by a overlapping area threshold).

First, we define a grid covering the total cluster area with grid cells  $[i, j] \in [1..N, 1..M]$ , resolution  $\Delta x = \Delta y = \zeta$  and grid cell center locations (relative to the geographic zero reference)  $[i\zeta, j\zeta]$ .

Given that each flare observation footprint area has outer bounds defined by center location  $[k_x, k_y]$  and radius  $R_k$ , we can obtain the vector  $\mathbf{I}_k$  containing the  $[i, j]$  index pairs of all grid cells within the footprint area by including all  $[i, j]$  pairs where both  $i\zeta < k_x + R_k$  and  $j\zeta < k_y + R_k$ .

Furthermore, we can approximate the total area  $A_k$  of flare observation  $k$  by

$$\hat{A}_k = \zeta^2 \sum_{i=1}^N \sum_{j=1}^M \delta_{i,j}, \quad \text{where} \quad \delta_{i,j} = \begin{cases} 1 & \text{when } [i, j] \in \mathbf{I}_k \\ 0 & \text{otherwise} \end{cases} \quad (5)$$

and, assuming uniform seepage for each individual flare observation, we can calculate the flow rate from each grid cell, i.e. the gridded observed gas flux per unit area,

$$\phi_k = \frac{F_k}{\hat{A}_k} \quad (6)$$

where  $F_k$  is the estimated flow rate from the flare observation using FlareHunter/ESP-3.

The total flux from the flare cluster can then be calculated by

$$\Phi = \sum_{i=1}^N \sum_{j=1}^M \phi^\top \cdot \frac{\delta_{i,j}}{\sum \delta_{i,j}} \quad (7)$$

where  $\phi = [\phi_1, \phi_2, \dots, \phi_K]$  is the flux per area of all flare observations in the cluster and  $\delta_{i,j} = [\delta_{i,j_1}, \delta_{i,j_2}, \dots, \delta_{i,j_K}]$  is the delta functions from Eq. 5 for the  $k^{th}$  flare observation in the cluster.

Essentially, we use a numerical grid to calculate the the individually averaged flowrates of all the different configurations of existing areal overlaps. This way we obtain the two aims of clustering without letting flowrates observed at one location impact our estimates at completely different locations. In other words: Where we have overlapping samples, i.e. *only where the footprints are overlapping*, we calculate the average flow rate of the overlapping samples. Otherwise we let the individual estimates remain unchanged.

We can illustrate this in a conceptual manner, disregarding errors associated by the numerical gridding, by the areas  $a_1, \dots, a_{12}$  and associated flowrates  $f_1, \dots, f_{12}$  in Figure 2c.

For instance does area  $a_1$  have no overlap with other observations and the flowrate  $f_1$  is calculated by

$$f_1 = \frac{F_1 a_1}{A_1}, \quad (8)$$

thereby only taking the flowrate observed in flare observation 1 into account. Flowrate of area  $a_8$  on the other hand is calculated as

$$f_8 = \frac{1}{3} \left( \frac{F_2}{A_2} + \frac{F_3}{A_3} + \frac{F_4}{A_4} \right) a_8, \quad (9)$$

here the expression  $\frac{1}{3} \left( \frac{F_2}{A_2} + \frac{F_3}{A_3} + \frac{F_4}{A_4} \right)$  gives the average flowrate per unit area of area  $a_8$  specifically - taking all, but only the observations overlapping in area  $a_8$  into account. The total flowrate is, conceptually, disregarding the numerical errors, in this example calculated as  $\sum_{k=1}^{12} f_k \sim 192$ .

Referring to Figure 1b, we average only in the greyed out area, instead of for the combined area of the two footprints as in the vanilla solution.

The only underlying assumption of this approach (which is also an assumption of the vanilla approach), which it could possibly be useful to also discuss further, is that seepage is uniform within the area of individual flare observations unless we include information from overlapping observations.

## References

- [Veloso et al., 2015] Veloso, M., Greinert, J., Mienert, J., and De Batist, M. (2015). A new methodology for quantifying bubble flow rates in deep water using splitbeam echosounders: Examples from the Arctic offshore NW-Svalbard. *Limnology and Oceanography: Methods*, 13(6):267–287.

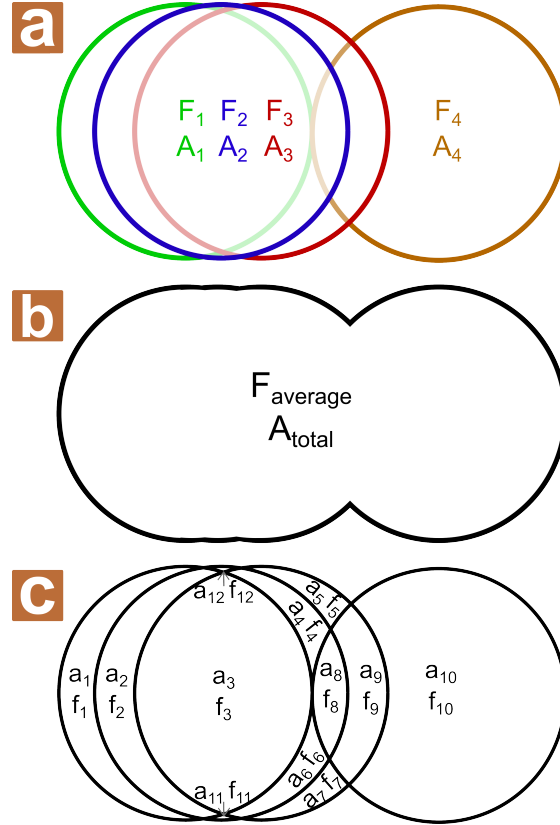


Figure 2: Idealized example of four clustered overlapping flare observations on a line with a) the four different observations and associated areas and flowrates, b) how the Vanilla solution implements averaging in the cluster and c) how the Gridded averaging method implements averaging in the cluster. See text and equations for how the total flux is calculated.