

CO₂ long-term future after injection

The CO₂, once injected, can undergo changes depending on the nature of the storage formation. This makes the security of the storage even greater. These changes are from free-phase CO₂ to CO₂ dissolved in water to CO₂ forming minerals over thousands of years. These changes are described as trapping mechanisms.

Trapping Mechanisms

The four basic mechanisms which hold the CO₂ in place are stratigraphical/structural, residual, solubility, and mineral trapping. Where the CO₂ is injected into horizontal or gently dipping reservoirs, or into saline formations, it can remain in the reservoir moving very slowly for a long time until eventually it is trapped by residual, solubility or mineral trapping. This is referred to as hydrodynamic trapping.

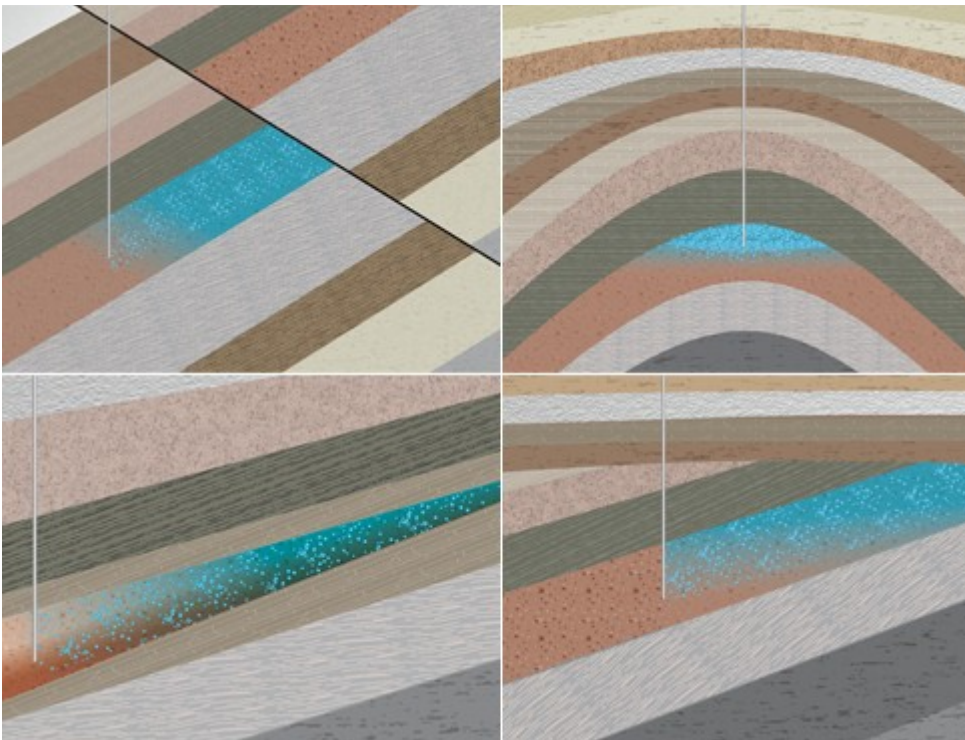
Structural/Stratigraphical Trapping

This type of trapping involves a physical trap. In a structural trap, the CO₂ is trapped at the top of an anticline, or in a tilted fault block. It is kept from further upward movement by the sealing rock (or caprock). There are a variety of geometrical arrangements of reservoir rock/sealing rock pairs. If the geometry of the subsurface does not contain faults or folds, the trapping is stratigraphical. Sometimes a change in rock type within the storage layer will create a closed stratigraphical trap.

When CO₂ is injected, it is not dissolved in formation water. It is said to be free-phase or immiscible. As it is supercritical, it is less dense than the formation water and so it rises upwards. Small bumps in the seal will behave like anticlines and trap the CO₂ until it spills out and continues migrating through the reservoir rock.

The CO₂ will remain trapped by the seal unless the height of the layer of CO₂ creates a capillary pressure that would enable it to enter the seal. Scientists calculate this height based on the pore radius of the sealing rock and the relative densities and surface tensions of CO₂ and water to determine the safe rate of injection into the reservoir.

Where the reservoir has closed structural trapping, it is likely that much of the CO₂ remains in supercritical, free-phase state in the storage reservoir, while the rest is trapped by other mechanisms.



Depth >800m

Clockwise from top left: fault trapping; anticline or stratigraphic trapping; facies change trapping; and unconformity trapping.

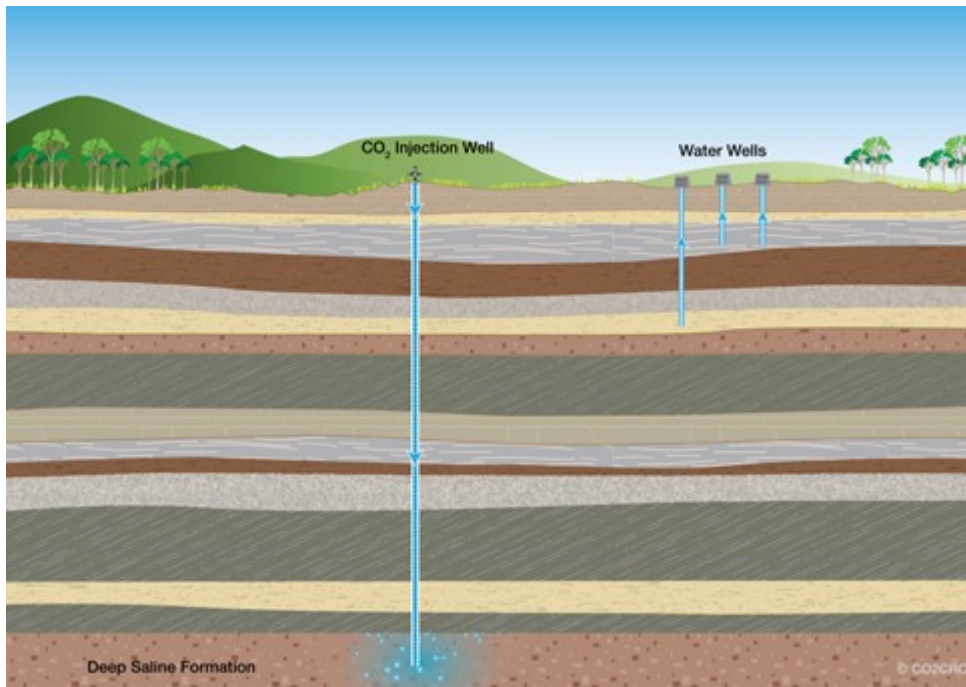
Hydrodynamic trapping

Saline formations generally have low flow velocities (of the order of cm/year). CO₂ that remains free-phase migrates upwards through permeable pathways in the rock. Because CO₂ is less viscous than the formation water (CO₂ is 15-20 times less viscous than brine at 1000m) it moves by a process called [viscous fingering](#). Some CO₂ is dissolved in

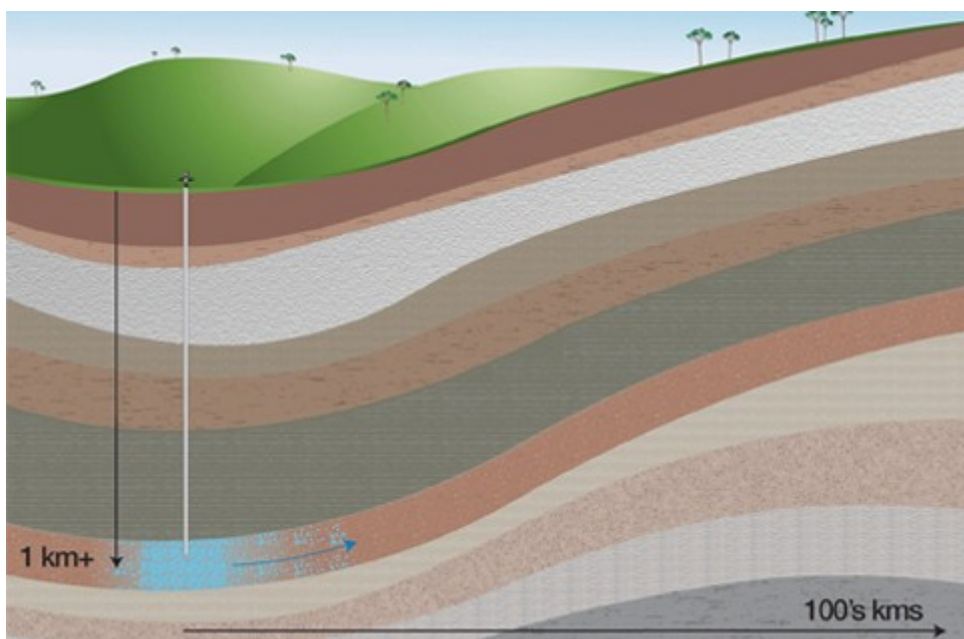
the formation water (see solubility). The dissolved CO₂ migrates with the formation water. The relationship between the dissolved CO₂ and the free-phase CO₂ determines how far the CO₂ will migrate before all of the CO₂ is eventually trapped residually or in solution.

Factors that have a substantial influence on the length of the migration pathway include:

- stratigraphic heterogeneities (eg siltstones, shales and coals in the reservoir rock)
- geochemical reactions
- temperature



Rock formations for geologic storage, such as deep saline formations, would be much deeper than any usable groundwater and separated from that groundwater by thick barriers of impervious rock. These formations generally already proved their effectiveness by keeping highly-salty saline water separate from usable groundwater for millions of years.

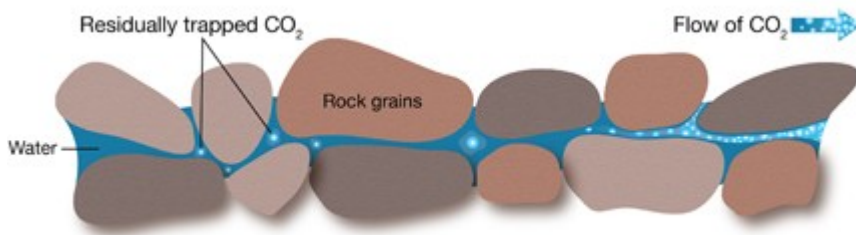


Aquifer trapping

Residual Trapping

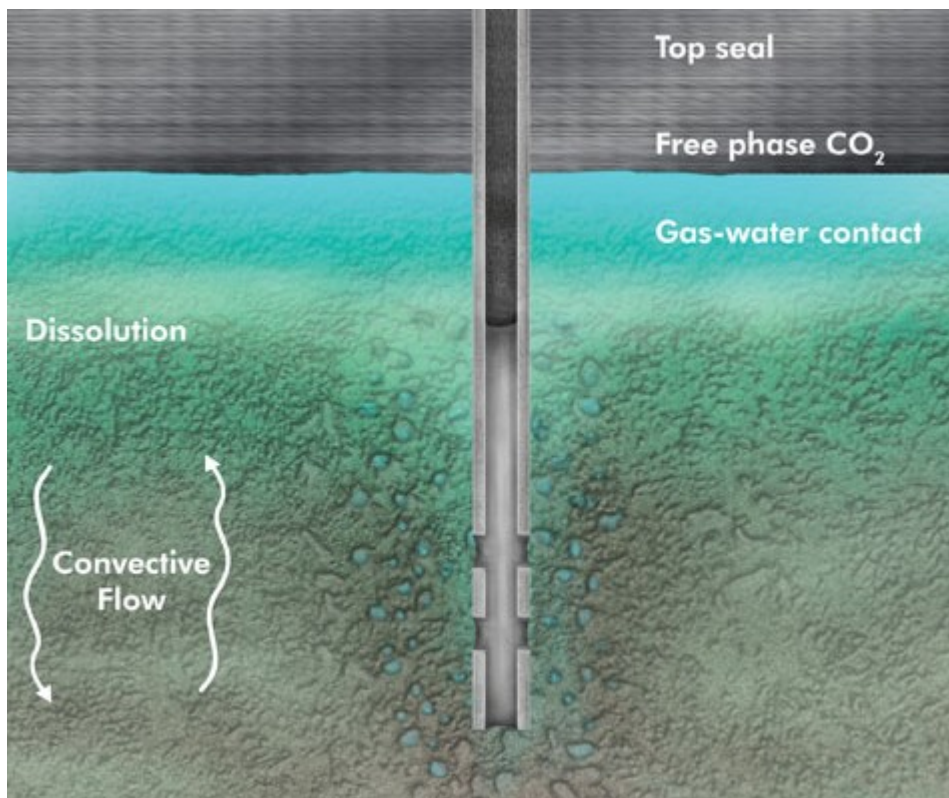
When free-phase CO₂ migrates, it forms a plume. At the tail of this plume, the concentration of the CO₂ falls below

a certain level. It becomes trapped by capillary pressure from the water in the pore spaces between the rock and stops flowing. Over time, this residually trapped CO₂ can dissolve into the formation water.



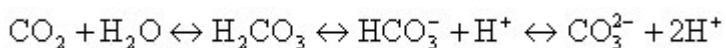
Solubility Trapping

The solubility of CO₂ in water increases with increasing pressure and decreases with increasing temperature and increasing water salinity. As some CO₂ dissolves in water, the water becomes denser, and begins to sink downwards. This allows the CO₂ to become more dispersed in the water, and over time, the amount of CO₂ dissolved in the water can increase. If the reservoir formation is thin, the extent to which mixing occurs through a convective flow is limited.



CO₂ dispersion in reservoir after injection has stopped. Note the capillary trapping around the injection well.

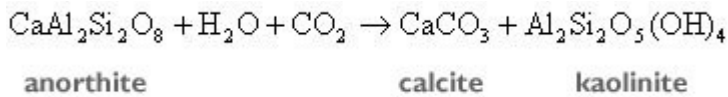
Aqueous CO₂ can form carbonic acid, a weak acid, in the reaction



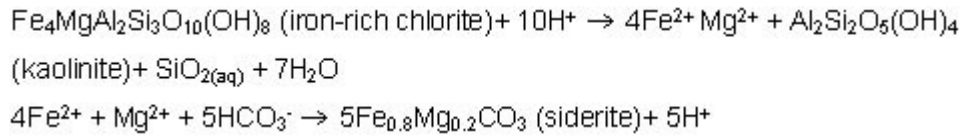
Mineral Trapping

When dissolved CO₂ reacts with the reservoir rock, carbonate minerals can form and precipitate, trapping CO₂ in the most stable form. While there is some reaction in the early years of storage, the time line for this trapping mechanism is over thousands of years. The potential for reaction to form these minerals depends on the composition of the reservoir rock (eg the presence of aluminosilicates), the temperature and pressure of the rock, the chemical composition of the water, the water/rock contact area and the rate of fluid flow through the rock. Mineral trapping is the most secure stage of CO₂ trapping.

One mineral which can be formed is calcite (calcium carbonate) in the following reaction:



A typical mineral reaction that occurs in the natural accumulations of CO₂ in the Otway basin is that of iron-rich chlorite forming kaolinite and siderite. The reaction below shows the formation of siderite:

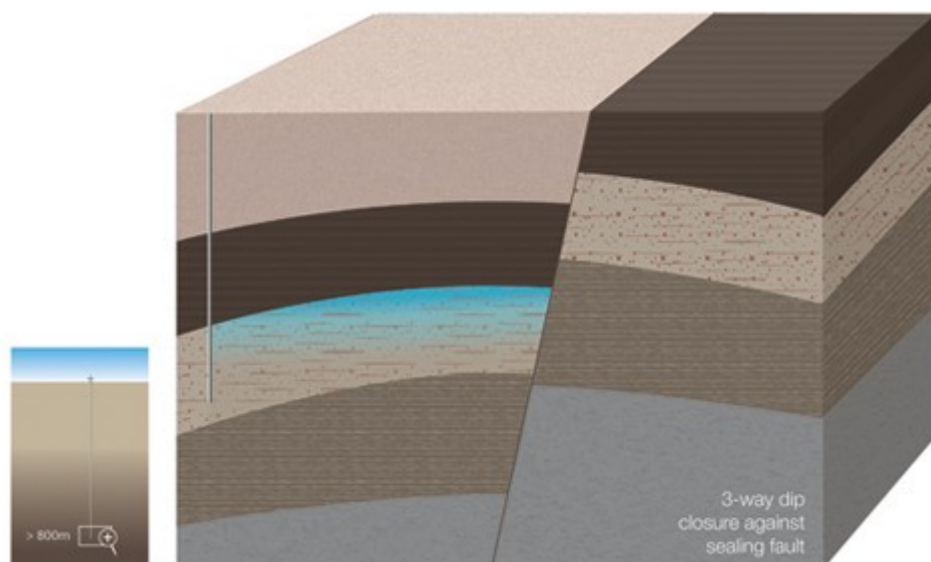
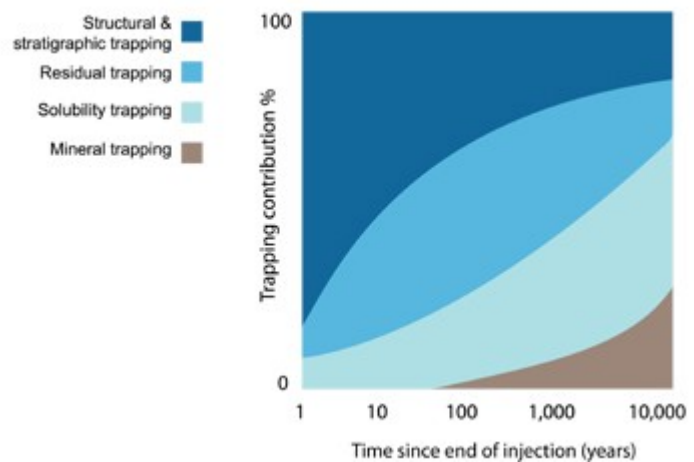


Storage mechanism over time

The way in which the CO₂ is injected and flows through the storage reservoir, together with the time the CO₂ remains in storage, determines the relative proportion of the trapping mechanism over time. In turn, the movement of CO₂ in the reservoir depends on the structure of the reservoir and the composition of the storage rocks and formation waters.

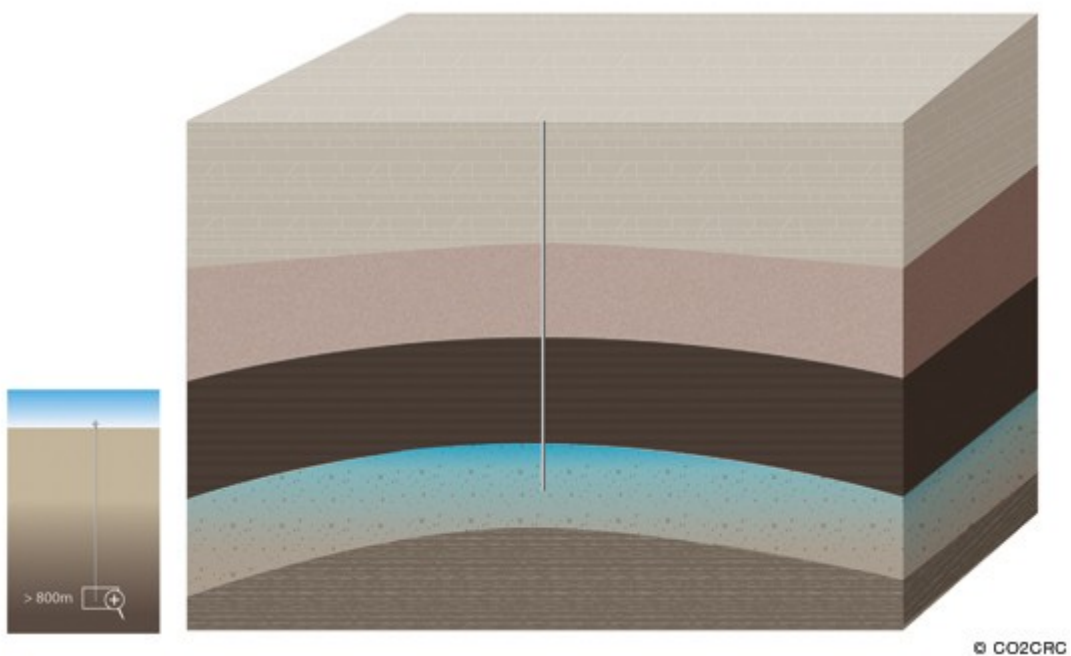
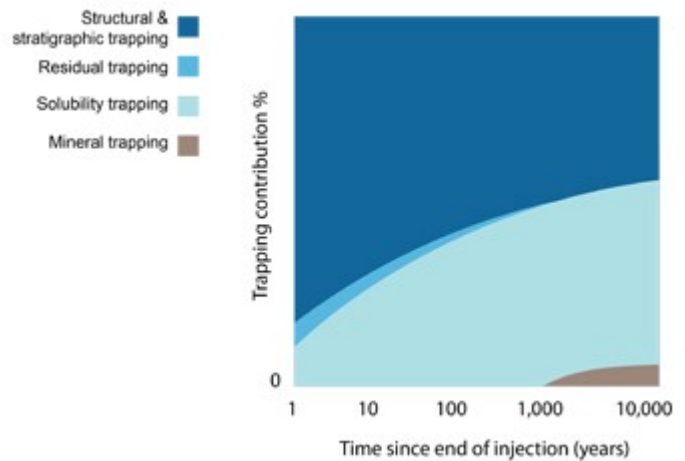
A mix of trapping mechanisms

CO₂ is injected into a thick saline formation of iron/magnesium-rich rock. The formation has a sealing fault on one side. After the CO₂ is injected it moves quickly away from the injection well. Some CO₂ remains free-phase, trapped by the rock structure. Small irregularities in the rock formation increase the migration pathway and residual trapping increasingly comes into play. Over time, significant amounts of CO₂ will dissolve into the formation brine and eventually, mineral trapping will occur.



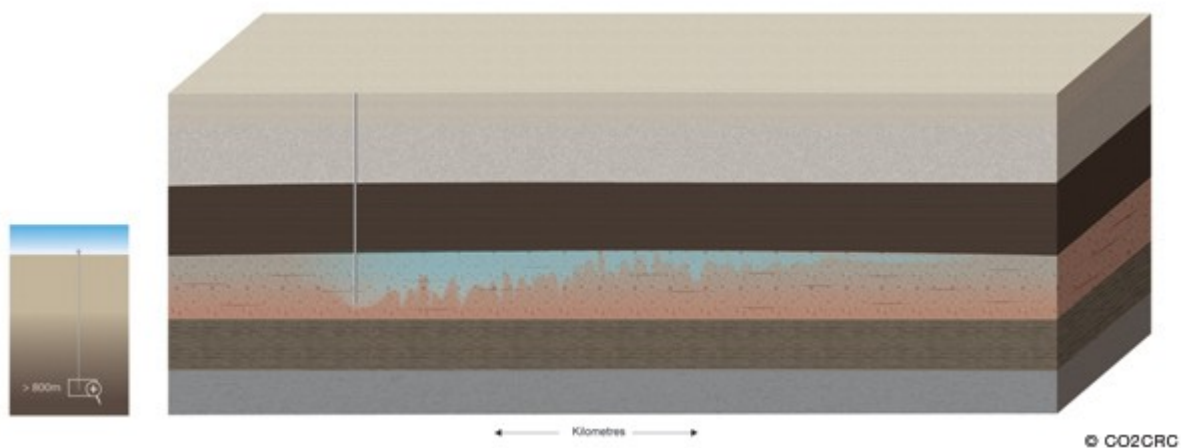
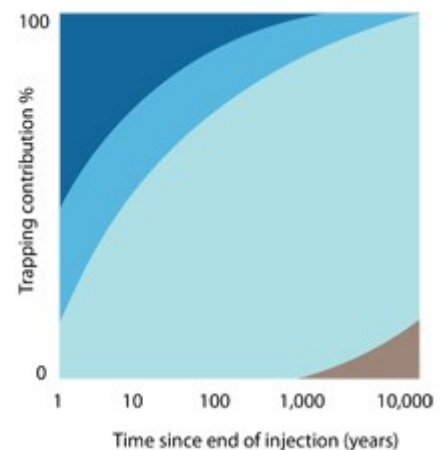
Structural trapping dominates

CO₂ is injected into a thin, silica sand reservoir which is a structural trap. The CO₂ moves to the top of the reservoir, not reacting with the quartz sand and with minor residual trapping. Most of the CO₂ will remain free-phase, with some dissolving into the formation brine over time.



Solubility trapping dominates

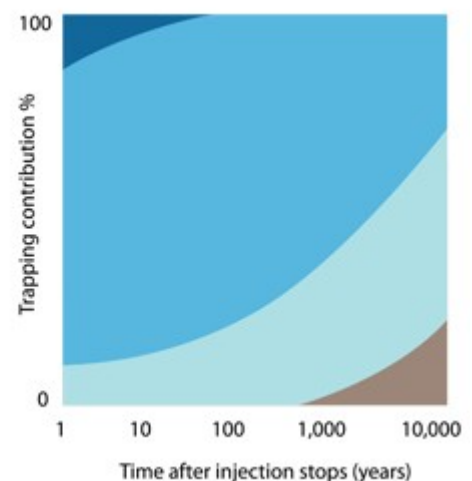
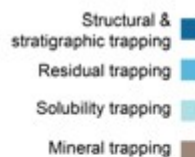
CO₂ is injected into an extensive saline formation of highly permeable rocks. The CO₂ will move easily upwards and spread out below the cap rock, eventually dissolving into the formation brine. There will be minor residual trapping.

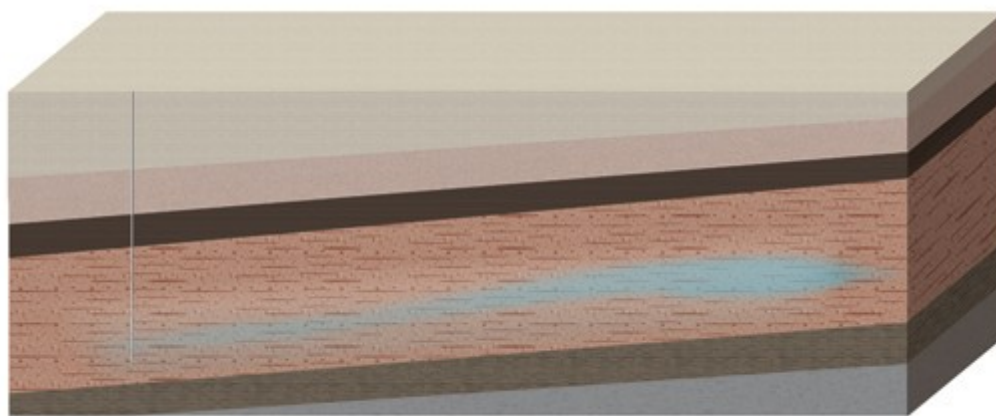


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Residual trapping dominates

CO₂ is injected into a thick, heterogeneous saline formation of rock with mixed pore sizes and low vertical permeability. The CO₂ is prevented from moving directly upwards by "shelves" of very low permeability rock, providing a form of structural trapping within the reservoir. It moves slowly with the formation water, contacting a large number of rock pores in which it becomes residually trapped. The residually trapped CO₂ eventually dissolves in the formation water.





← Kilometres to tens of kilometres →

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