

GEOMAGNETIC, GEOCENTRIC, AND GEODETIC COORDINATE TRANSFORMATIONS

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Abstract—A requirement in geomagnetism is to make transformations between geocentric, geodetic (spheroidal), and geomagnetic coordinate systems. A few of these transformations may be included in textbooks, but we have never noted them all presented. This paper provides derivations and subroutines for conversion of spatial coordinates and Cartesian magnetic field components between any combinations of these three coordinate systems.

Key Words: Coordinate transformations, Geomagnetism, Geocentric, Geodetic, Geodesy.

INTRODUCTION

Satellite positions and observations usually are expressed in geocentric coordinates. Most geophysical observations made at, or near, the surface of the Earth are expressed in geodetic coordinates references to an ellipsoidal (spheroidal) Earth. Many geomagnetic phenomena are controlled by the Earth's geomagnetic dipole and are best treated in geomagnetic coordinates. If, like us, you have needed to transform spatial coordinates (and magnetic field components) between these three coordinate systems, then you probably share our frustration in never being able to locate the required equations in the literature. This paper provides equations and FORTRAN-77 subroutines for performing any of the six transformations between these coordinate systems, and includes appropriate transformations for Cartesian components of the magnetic field.

Equations for converting from geocentric to geomagnetic coordinates are presented in many basic textbooks on geomagnetism (e.g. Chapman and Bartels, 1940; Matsushita and Campbell, 1967). Likewise, equations for converting from geodetic to geocentric coordinates are given (e.g. Zmuda, 1971—as used in most computer programs for evaluating spherical harmonic models of the geomagnetic field, such as IGRF). However, the reverse transformations are seldom documented, particularly the conversion from geocentric to geodetic coordinates—perhaps because the equations are inelegant and require the solution of a quartic equation. Some of the widely used subroutines place restrictions on which quadrant(s) can be used and on the sign of the geodetic height (altitude), whereas those presented here apply in all situations provided sensible values of the parameters are entered. Angles are given in degrees unless stated otherwise.

GEOCENTRIC COORDINATES

Geocentric coordinates are spherical polar coordinates with origin at the center of mass of the Earth and polar axis along the northwards rotation axis of the Earth (ON), as illustrated in Figure 1. Longitude is referred to the Greenwich meridian (NGS). A point in space is designated by colatitude (θ), longitude (ϕ) measured eastwards from the Greenwich meridian, and radial distance (r) from the origin. The geocentric latitude of P is angle λ , equal to $90 - \theta$. Unlike the geodetic and geomagnetic coordinate systems, the geocentric system does not require specification of the particular earth model that is used for reference.

GEODETIC COORDINATES

An oblate ellipsoid approximating to the surface of the Earth is used as a reference surface (Fig. 1). The axis of the ellipsoid is the rotation axis of the Earth (ON), and the ellipsoid is defined by its major semiaxis, a , and flattening, f . The minor semiaxis, b , is derived from

$$f = (a - b)/a$$

and ellipticity, e , is given by

$$e^2 = (a^2 - b^2)/a^2. \quad (1)$$

A point P is designated by its geodetic colatitude, ψ , longitude, ϕ , and perpendicular (vertical) height, h , above the ellipsoid. (In practice the local vertical at any point, as measured by a plumb-line, generally will not be exactly parallel to the normal to the ellipsoid.) The geodetic latitude of P is the angle β , equal to $90 - \psi$. Examples of some commonly used reference ellipsoids are given in Table 1.

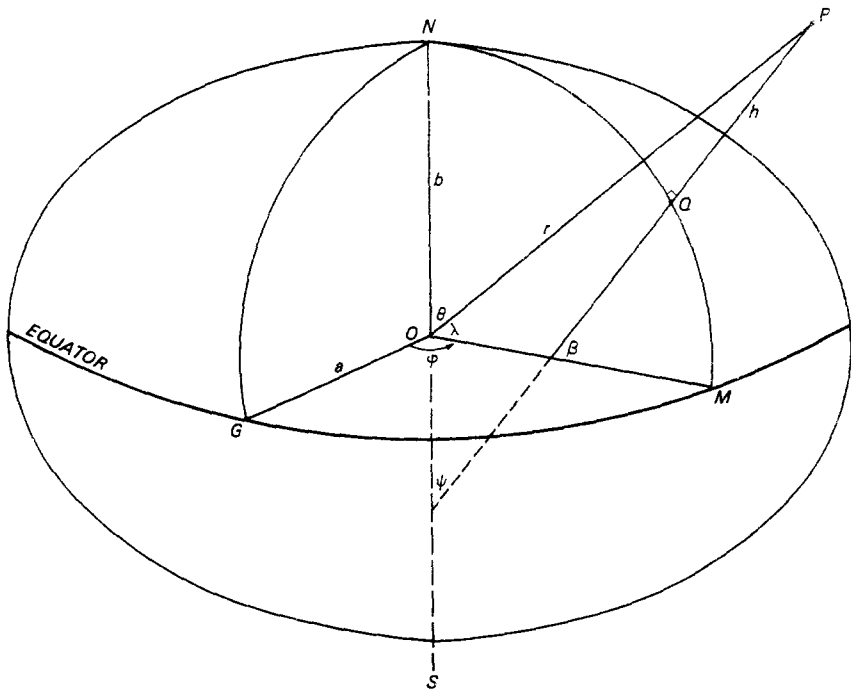


Figure 1. Geocentric and geodetic coordinates illustrated on reference ellipsoid with major semiaxis, a , and minor semiaxis, b . NS is rotation axis of Earth. PQ is normal from point P to surface of ellipsoid. NQM is meridian plane containing P . Geocentric coordinates of P are colatitude, θ (or latitude, λ), longitude, ϕ , and radial distance, r , from center of Earth (O). Geodetic coordinates of P are colatitude, ψ (or latitude, β), longitude, ϕ , and vertical height, h , above ellipsoid ($= PQ$). Geocentric and geodetic longitudes are same, measured eastwards from Greenwich meridian (NGS).

Table 1. Reference ellipsoids			
Name	a (km)	$1/f$	b (km)
Clarke ellipsoid (1866)	6378.206	294.98	6356.58350
International ellipsoid (IAG, Madrid, 1924)	6378.388	297.00	6356.91195
Krassovsky ellipsoid (1942)	6378.245	298.30	6356.86302
IAU ellipsoid (Hamburg, 1964)	6378.160	298.25	6356.77472
Australian national ellipsoid (1966)	6378.160	298.25	6356.77472
South American ellipsoid (1969)	6378.160	298.25	6356.77472
World Geodetic System ellipsoid (1972)	6378.135	298.26	6356.75052

a —Major semiaxis; f —flattening; b —minor semiaxis.

GEOMAGNETIC COORDINATES

The geomagnetic coordinate system (Schmidt, 1918) is a spherical polar coordinate frame with polar axis coincident with the north-pointing, centered-dipole (geomagnetic) axis of the Earth. The zero-longitude meridian is defined as the geomagnetic meridian containing the south geographic pole. Geomagnetic coordinates are distinct from magnetic coordinates (see Matsushita and Campbell, 1967, Appendix 1).

The position of the geomagnetic axis is subject to secular variation, hence when specifying geomagnetic coordinates it is necessary to state the coordinates of the north (or south) geomagnetic pole—usually in geocentric coordinates, which are derived readily from the spherical harmonic Gauss coefficients. Compila-

tions of geomagnetic dipole coordinates have been published by many authors (e.g. Voppel, 1985; Barton, 1989). A widely used standard for the geomagnetic axis is with the north geomagnetic pole at colatitude 11.5° , longitude 291.0°E (e.g. Chapman and Bartels, 1940, table A; Voppel, 1985). Some other values are listed in Table 2.

Table 2. Geocentric coordinates of north geomagnetic pole

Source	Colatitude ($^\circ$)	Longitude ($^\circ\text{E}$)
Matsushita and Campbell (1967)	11.7	291.0
MGST481-1 @ 1980.0	11.198	289.245
DGRF 1980 @ 1980.0	11.194	289.095
IGRF 1985 @ 1985.0	11.018	289.095

THE SUBROUTINES

The following FORTRAN subroutines for performing the transformations are listed in Appendix 2:

DTOCB—geodetic to geocentric
 CTODB—geocentric to geodetic
 CTOMB—geocentric to geomagnetic
 MTOCB—geomagnetic to geocentric.

Letters "C", "D", and "M" denote geocentric, geodetic, and geomagnetic respectively; "B" indicates that the routines include a conversion for the Cartesian magnetic field components—northwards (BN), eastwards (BE), and vertically downwards (BV). It is a simple matter to remove the code for the field components if coordinate transformations alone are required.

Derivations of the equations used are given in Appendix 1. Conversion between geomagnetic and geodetic frames is accomplished by converting to and from geocentric coordinates as an intermediate step.

Double precision (real*8) has been used throughout. The subroutines for converting between geocentric and geodetic coordinates are coded to reduce rounding errors arising from taking powers of large numbers. For the spherical rotations (subroutines CTOMB and MTOCB) single precision would be sufficiently accurate for many purposes.

Spatial coordinates are expressed as colatitude, longitude (positive east of the reference meridian), and either radial distance in kilometers from the center of the Earth (geocentric and geomagnetic frames) or vertical height in kilometers above the reference ellipsoid (geodetic frame). Colatitude, or polar distance, is permitted values within the range 0–180°. Longitude can have any real values. In order to make the subroutine calls symmetrical we have selected to include the coordinates and components that are unchanged by the transformations. The "vertical" direction for any point lies along the normal to the reference surface passing through the point. In practice the local vertical, as measured by a plumb-line, generally will not be exactly parallel to this normal.

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APPENDIX 1

Derivation of Equations

We make use of the equations for solving a spherical triangle with included angles A, B, C (in radians) facing sides a, b, c . For a sphere of unit radius the side lengths are equal to the respective angles in radians subtended at the center of the sphere.

Cosine rule:

$$\cos a = \cos b \cos c + \sin b \sin c \cos A \quad (A1)$$

Sine rule:

$$\frac{\sin A}{\sin a} = \frac{\sin B}{\sin b} = \frac{\sin C}{\sin c} \quad (A2)$$

For a spherical triangle with a right-angle at A :

$$\cos a = \cos b \cos c. \quad (A3)$$

Geodetic to geocentric conversion

The equations derived here are less complicated than those quoted by Zmuda (1971), which are used in the subroutines provided by the U.S. Geological Survey and the British Geological Survey for evaluating spherical harmonic models of the field such as IGRF. Furthermore, our equations give the correct result for points below the surface of the reference ellipsoid (negative heights).

Refer to Figure A1. Geocentric and geodetic longitudes are the same, so the problem is reduced to 2 dimensions. Given the geodetic coordinates (h, ψ) of point P , we wish to determine its geocentric coordinates (r, θ) with respect to the reference ellipsoid with major semiaxis, a , and minor semiaxis, b . We work in Cartesian coordinates, where the "y" axis is the polar axis, the "x" axis lies in the equatorial plane. $Q(X, Y)$ is the point on the reference ellipsoid vertically below $P(X_p, Y_p)$. The geodetic latitude, $\beta = 90 - \psi$.

Because Q lies on the ellipsoid,

$$X^2/a^2 + Y^2/b^2 = 1. \quad (A4)$$

The normal to the ellipsoid through Q is

$$x a^2/X - y b^2/Y = a^2 - b^2$$

which has slope

$$\tan \beta = dy/dx = (a^2/b^2)(Y/X). \quad (A5)$$

From Equations (5) and (6),

$$X = a^2/\sqrt{(a^2 + b^2 \tan \beta)}$$

and

$$Y = (b^2/a^2) \tan(\beta) X$$

hence

$$X_p = X + h \cos \beta; \quad Y_p = Y + h \sin \beta.$$

The geocentric radial distance, r , and geocentric colatitude are determined from

$$r^2 = X_p^2 + Y_p^2; \quad \cos \theta = Y_p/r.$$

To rotate the Cartesian components of the field, we determine the angle, α , between the geocentric and geodetic verticals

$$\alpha = \beta - \lambda = \beta - 90 + \theta. \quad (A6)$$

Then for both northern and southern hemispheres

$$BN_c = BN_d \cos \alpha - BV_d \sin \alpha$$

$$BV_c = BV_d \cos \alpha + BN_d \sin \alpha$$

and

$$BE_d = BE_c$$

where c and d denote geocentric and geodetic respectively.

Geocentric to geodetic conversion

Refer again to Figure A1. Given the geocentric coordinates (r, θ) of point P , we wish to determine the geodetic coordinates (h, ψ) . The Cartesian coordinates of P are

$$X_p = r \sin(\theta); \quad Y_p = r \cos(\theta).$$

Because Q lies on the ellipsoid,

$$X^2/a^2 + Y^2/b^2 = 1 \quad (A7)$$

and P lies on the normal at Q ,

$$X_p a^2/X - Y_p b^2/Y = a^2 - b^2 \quad (A8)$$

it follows that

$$A_0 + A_1 X + A_2 X^2 + A_3 X^3 + A_4 X^4 = 0 \quad (A9)$$

where

$$A_0 = -a^2 X_p^2$$

$$A_1 = 2a^2 e^2 X_p$$

$$A_2 = X_p^2 + Y_p^2 b^2/a^2 - a^2 e^4$$

$$A_3 = -2e^2 X_p$$

$$A_4 = e^4.$$

Equation (A9) can be solved efficiently and safely by using Newton's method, taking as the initial approximation for the root:

$$X_i = 0.5(a + b) \sin \theta.$$

The coefficients of Equation (A9) are real hence the roots are either real or occur as pairs of conjugate complex numbers. In our situation one root is positive, one is negative, and the remaining two are complex. It is evident from Figure A1 that for any position of P this starting value must be sufficiently close to the required (positive) root that Newton's method will converge quickly.

The constant EPS in subroutine CTOGB determines how accurately the root of the quadratic is determined. (In the subroutine the value of EPS is determined by the precision of the computer being used.) The iteration to determine the root is stopped when successive approximations differ by less than EPS. Having determined the root, X , the value of Y is obtained from Equation (A7). The height above the ellipsoid is then

$$h = \sqrt{[(X - X_p)^2 + (Y - Y_p)^2]}$$

The negative square root is taken if $Y > Y_p$. The geodetic latitude, β , is determined from Equation (A5).

The equations hold for points in the northern hemisphere. Points in the southern hemisphere are treated in terms of their mirror images in the northern hemisphere.

To rotate the Cartesian components of the field, we determine the angle between the geocentric and geodetic verticals, α , from Equation (A6). Then for the northern hemisphere

$$BN_d = BN_c \cos \alpha + BV_c \sin \alpha$$

$$BV_d = BV_c \cos \alpha - BN_c \sin \alpha$$

and for the southern hemisphere

$$BN_d = BN_c \cos \alpha - BV_c \sin \alpha$$

$$BV_d = BV_c \cos \alpha + BN_c \sin \alpha.$$

in both situations

$$BE_d = BE_c.$$

Geocentric to geomagnetic conversion

Refer to Figure A2. Given the geocentric colatitude and longitude of the north geomagnetic pole (θ_p, ϕ_p) and the geocentric colatitude and longitude (θ, ϕ) of P , we need to determine the geomagnetic colatitude and longitude (θ_m, ϕ_m) of P . Applying the cosine rule (Eq. [(A1)] to spherical triangle NMP .

$$\cos \theta_m = \cos \theta_p \cos \theta + \sin \theta_p \sin \theta \cos(\phi - \phi_p).$$

Taking the inverse cosine gives θ_m in the correct quadrant. Applying the sine rule [Eq. (A2)],

$$\sin \gamma = \sin \theta \sin(\phi - \phi_p)/\sin \theta_p$$

whence

$$\phi_m = 180 - \gamma = \sin^{-1} [\sin \theta \sin(\phi - \phi_p)/\sin \theta_p].$$

This returns ϕ_m in the first and fourth quadrants only. Equation (A3) is used to test for the obtuse angle situation of the spherical triangle:

if

$$\cos \theta > \cos \theta_p \cos \theta_m$$

then

$$\phi_m = 180 - \sin^{-1} [\sin \theta \sin(\phi - \phi_p)/\sin \theta_p].$$

To rotate the Cartesian components of the field, we determine the angle between the geocentric and geomagnetic "north" by applying the sine rule to triangle NMP ,

$$\alpha = \sin^{-1} [\sin \theta_p \sin(\phi - \phi_p)/\sin \theta_m] \quad (A10)$$

and test to determine the obtuse angle situation:

if

$$\cos \theta_p < \cos \theta \cos \theta_m$$

then

$$\alpha = 180 - \sin^{-1} [\sin \theta_p \sin(\phi - \phi_p)/\sin \theta_m]. \quad (A11)$$

Rotating the field components we get:

$$BN_m = BN_c \cos \alpha - BE_c \sin \alpha$$

$$BE_m = BE_c \cos \alpha + BN_c \sin \alpha$$

and

$$BV_m = BV_c$$

where suffixes m and c denote geomagnetic and geocentric respectively.

Geomagnetic to geocentric conversion

Refer to Figure A2. Given the geomagnetic pole (θ_p, ϕ_p) and the geomagnetic coordinates (θ_m, ϕ_m) , we now wish to determine the geocentric colatitude and longitude (θ, ϕ) of P . Proceeding in much the same way as for the reverse transformation, we get

$$\cos \theta = \cos \theta_p \cos \theta_m + \sin \theta_p \sin \theta_m \cos \gamma.$$

Taking the inverse cosine gives the geocentric colatitude in the correct quadrant.

By the sine rule,

$$\begin{aligned}\sin(\phi - \phi_p) &= \sin \theta_m \sin \gamma / \sin \theta \\ &= \sin \theta_m \sin \phi_m / \sin \theta\end{aligned}$$

whence

$$\phi = \phi_p + \sin^{-1} [\sin \theta_m \sin \phi_m / \sin \theta].$$

This returns ϕ in the first and fourth quadrants only. Testing for the obtuse angle situation,

if

$$\cos \theta_m < \cos \theta_p \cos \theta$$

then

$$\phi = \phi_p + 180 - \sin^{-1} [\sin \theta_m \sin \phi_m / \sin \theta].$$

To rotate the Cartesian components of the field find angle α , from Equations (A10) or (A11), whence

$$BN_c = BN_m \cos \alpha + BE_m \sin \alpha$$

$$BE_c = BE_m \cos \alpha - BN_m \sin \alpha$$

and

$$BV_c = BV_m.$$

APPENDIX 2

Subroutine Listings

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C-----
      subroutine DTOCB (CLATD, ELOND, ALT, BXD, BYD, BZD,
+                      CLATC, ELONC, R, BXC, BYC, BZC)
C
C Geodetic to geocentric conversion for coordinates and field
C I/O double precision (real*8)
C Input (geodetic)
C   CLATD= colatitude (deg), range 0-180
C   ELOND= longitude (degE)
C   ALT = height above spheroid (km), may be negative
C   BXD = north component of field
C   BYD = east component of field
C   BZD = vert component of field
C Output (geocentric)
C   CLATC= colatitude (deg)
C   ELONC= longitude (degE)
C   R = radial distance from geocentre (km)
C   BXC = north component of field
C   BYC = east component of field
C   BZC = vert component of field
C
      implicit real*8 (A-H,O-Z)
      rpd=datan(1.d0)/45.d0
      A = 6378.135d0 ! semi major axis, World spheroid, WGS 1972
      B = 6356.750d0 ! semi minor axis " "
      if (CLATD.lt.0.d0.or.CLATD.gt.180.d0) then ! quit
        write(*,*) 'Colatitude out of range in subroutine DTOCB'
        stop
      endif
C
C Convert to geocentric coordinates
C
      AA = A*A
      BB = B*B
      B2A2= (B/A)**2
      BETA= (90.d0-CLATD)*rpd ! geodetic latitude (rad)
      TANB= dtan(BETA)
      X = A/dsqrt(1.d0+B2A2*TANB*TANB)
      Y = B2A2*TANB*X
      XP= X + ALT*dcos(BETA)
      YP= Y + ALT*dsin(BETA)
      R = dsqrt(XP*XP+YP*YP) ! radial dist. (km)
      CLATC = dacos(YP/R)/rpd ! geocentric colat (deg)
      ELONC = ELOND
C
C Transform field components to geocentric
C
      ALPHA= BETA-(90.d0-CLATC)*rpd ! in radians
      CA = dcos(ALPHA)
      SA = dsin(ALPHA)
      BXC= BXD*CA - BZD*SA
      BYC= BYD
      BZC= BZD*CA + BXD*SA
      return
      end

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-----
      subroutine CTODB (CLATC,ELONC, R ,BXC,BYC,BZC,
+                   CLATD,ELOND,ALT,BXD,BYD,BZD)
c
c   Geocentric to geodetic conversion for coordinates & field
c   Double precision (real*8) throughout
c   The root of the quartic equation is found numerically -
c   EPS is computed according to the precision of the computer
c   A limit, MAXIT, is placed on the number of iterations allowed
c   Input (geocentric)
c     CLATC= colatitude (deg), range 0-180
c     ELONC= longitude  (degE)
c     R    = radial distance from geocentre (km)
c     BXC  = north component of field
c     BYC  = east  component of field
c     BZC  = vert  component of field
c   Output (geodetic)
c     CLATD= colatitude (deg), range 0-180
c     ELOND= longitude  (degE)
c     ALT  = height above spheroid (km)
c     BXD  = north component of field
c     BYD  = east  component of field
c     BZD  = vert  component of field
c
c   Colat=0 and 180 are treated as special cases
c   Colat>90 is calculated from the image point in the 1st quadrant
c
      implicit real*8 (A-H,O-Z)
      rpd=datan(1.d0)/45.d0
      MAXIT=300                      ! max number of iterations allowed
      A = 6378.135d0                 ! semi major axis, WGS 1972
      B = 6356.750d0                 ! semi minor axis
c
c   Special cases
c
      if(CLATC.lt.0.d0.or.CLATC.gt.180.d0) then          ! quit
        write(*,*)'Geocentric colat. out of range in subroutine CTODB'
        stop
      endif
      if(CLATC.eq.0.d0.or.CLATC.eq.180.d0) then          ! at poles
        CLATD=CLATC
        ELOND=ELONC
        ALT= R-B
        BXD=BXC
        BYD=BYC
        BZD=BZC
        return
      endif
c
c   Evaluate EPS, the allowed error for the computed root, according to
c   the precision of the computer
c
      EPS=1.d-5                      ! starting value
100  continue
      if(1.d8+EPS.gt.1.d8) then
        EPS=EPS/2.d0
        goto 100
      endif
      write(*,*)'Precision limit, EPS=',EPS
c
c   Set up equations
c
      if(CLATC.le.90.d0) then          ! 1st quadrant
        IQUAD=1
        CLAT=CLATC*rpd                ! colat in radians
      else                             ! 2nd quadrant
        IQUAD=2
        CLAT=(180.d0-CLATC)*rpd       ! image in 1st quadrant
      endif

      AA= A*A
      BB= B*B
      EE= 1.d0-(B/A)**2                ! eccentricity squared
      XP= R*sin(CLAT)                  ! x coord for geodetic point
      YP= R*cos(CLAT)                  ! y coord  "
      A4= EE*EE
      A3=-2.d0*XP*EE
      A2= XP*XP+(YP*YP)*BB/AA-EE*EE*AA

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      A1= 2.d0*XP*AA-2.d0*XP*BB
      A0=-AA*XP*XP
C
C Find root X by Newton's method
C
      X = 0.5d0*(A+B)*sin(CLAT) ! first approx. root
      write(*,*)
      do IT=1,MAXIT
        P1=A4*X + A3 ! evaluate quartic
        P2=P1*X + A2
        P3=P2*X + A1
        P4=P3*X + A0
        Q1=4.d0*X + 3.d0*A3 ! evaluate derivative
        Q2= Q1*X + 2.d0*A2
        Q3= Q2*X + A1
        G=-P4/Q3
        X= X+G ! new approximation to root
        if(mod(IT,5).eq.0) write(*,*) IT,' Root=',X,G
        if(dabs(G).le.EPS) goto 200 ! close enough
      end do
      write(*,800) MAXIT
800 format(' Root not found after',i4,' iterations in CTODB')
      stop
C
C Find Y
C
200 Y = B*dsqrt(A+X)*dsqrt(A-X)/A
      if(dabs(A-X).le.10.d0*EPS) Y=0.d0
C
C Find the geodetic colatitude and longitude
C
      BETA= datan(Y/(X*(1-EE))) ! geodetic latitude (rad)
      if(IQUAD.eq.1) then
        CLATD=90.d0-BETA/rpd ! colat (deg) in 1st quadrant
      else
        CLATD=90.d0+BETA/rpd ! colat (deg) in 2nd quadrant
      endif
      ELOND=ELONC
C
C Find height, positive or negative
C
      ALT = dsqrt((X-XP)**2+(Y-YP)**2) ! km above ellipsoid
      if(YP.le.Y) ALT=-ALT ! inside ellipsoid, negative
C
C Convert field components into geodetics
C
      ALPHA=BETA+CLAT-90.d0*rpd ! angle between 'verticals'
      if(IQUAD.eq.1) then ! 1st quadrant
        BXD= BXC*dcos(ALPHA) + BZC*dsin(ALPHA)
        BZD= BZC*dcos(ALPHA) - BXC*dsin(ALPHA)
      else ! 2nd quadrant
        BXD= BXC*dcos(ALPHA) - BZC*dsin(ALPHA)
        BZD= BZC*dcos(ALPHA) + BXC*dsin(ALPHA)
      endif
      BYD=BYC
      return
      end

C-----
      subroutine CTOMB (CLATC,ELONC,BXC,BYC,BZC,CLATP,ELONP,
+                      CLATM,ELONM,BXM,BYM,BZM)
C
C Converts coordinates & field from geocentric to geomagnetic
C Double precision (real*8) throughout
C Radial distance from geocentre & Z-field (down) are unchanged.
C Input (geocentric)
C CLATC= colatitude (deg), range 0-180
C ELONC= longitude (degE), any value (+ve east)
C BXC = north component of field
C BYC = east component of field
C BZC = vert component of field
C CLATP= pole colatitude (deg), range 0-180
C ELONP= pole longitude (degE), any value (+ve east)
C Output (geomagnetic)
C CLATM= colatitude (deg)
C ELONM= longitude (degE)
C BXM = north component of field
C BYM = east component of field
C BZM = vert component of field
C

```



```

implicit real*8 (A-H,O-Z)
rpd=datan(1.d0)/45.d0      ! radians per degree

if (CLATC.lt.0.d0.or.CLATC.gt.180.d0.or.
+  CLATP.lt.0.d0.or.CLATP.gt.180.d0) then
  write(*,*) 'Colatitude input out of range in subroutine CTOMB'
  stop
endif
CLATCR= CLATC*rpd
CLATPR= CLATP*rpd
ELONCR= ELONC*rpd
ELONPR= ELONP*rpd
c
c Convert coordinates to geomagnetic (CLATM,ELONM)
c
  A= dcos(CLATPR)*dcos(CLATCR)
  B= dsin(CLATPR)*dsin(CLATCR)*dcos(ELONCR-ELONPR)
  CLATMR= dacos(A+B)
  CLATM = CLATMR/rpd      ! magnetic colatitude
  C= dsin(CLATCR)*dsin(ELONCR-ELONPR)/dsin(CLATMR)
  ELONM= dasin(C)/rpd
  if (dcos(CLATCR).gt.dcos(CLATPR)*dcos(CLATMR)) ELONM=180.d0-ELONM
  if (ELONM.lt.0.d0) ELONM=ELONM+360.d0
c
c Convert field to geomagnetic frame
c
  SA= dsin(CLATPR)*dsin(ELONCR-ELONPR)/dsin(CLATMR)      ! sin alpha
  ALPHA=dsin(SA)/rpd      ! in deg
  if (dcos(CLATPR).lt.dcos(CLATCR)*dcos(CLATMR)) ALPHA=180.0-ALPHA
  CA = dcos(ALPHA*rpd)      ! cos alpha
  BXM = BXC*CA - BYC*SA      ! north (mag)
  BYM = BYC*CA + BXC*SA      ! east (mag)
  BZM = BZC
  return
end

-----
      subroutine MTOCB (CLATM,ELONM,BXM,BYM,BZM,CLATP,ELONP,
+                      CLATC,ELONC,BXC,BYC,BZC)
c
c Converts coordinates & field from geomagnetic to geocentric
c Double precision (real*8) throughout
c Radial distance from geocentre & Z-field (down) are unchanged.
c Input (geomagnetic)
c   CLATM= colatitude (deg), range 0-180
c   ELONM= longitude (degE), any value (+ve east)
c   BXM = north component of field
c   BYM = east component of field
c   BZM = vert component of field
c   CLATP= geomagnetic pole colatitude (deg), range 0-180
c   ELONP= " pole longitude (degE), any value (+ve east)
c Output (geocentric)
c   CLATC= colatitude (deg)
c   ELONC= longitude (degE)
c   BXC = north component of field
c   BYC = east component of field
c   BZC = vert component of field
c
  implicit real*8 (A-H,O-Z)
  rpd=datan(1.d0)/45.d0      ! radians per degree

  if (CLATM.lt.0.d0.or.CLATM.gt.180.d0.or.
+  CLATP.lt.0.d0.or.CLATP.gt.180.d0) then
    write(*,*) 'Colatitude out of range in subroutine MTOCB'
    stop
  endif

  CLATMR=CLATM*rpd      ! in radians
  CLATPR=CLATP*rpd
  ELONMR=ELONM*rpd
  ELONPR=ELONP*rpd

  A= dcos(CLATPR)*dcos(CLATMR)
  B= dsin(CLATPR)*dsin(CLATMR)*dcos(ELONMR)
  CLATCR= dacos(A-B)
  CLATC = CLATCR/rpd
  C= dsin(CLATMR)*dsin(ELONMR)/dsin(CLATCR)
  ELONC= ELONP+dsin(C)/rpd

```

```

      if(dcos(CLATMR).lt.dcos(CLATPR)*dcos(CLATCR)) then
        ELONC=180.d0+2.d0*ELONP-ELONC
      endif
      if(ELONC.lt. 0.d0) ELONC=ELONC+360.d0
      if(ELONC.gt.360.d0) ELONC=ELONC-360.d0
c
c Convert field to geocentric frame
c
      SA= dsin(CLATPR)*dsin(ELONC*rpdc-ELONPR)/dsin(CLATMR)
      ALPHA=dsin(SA)/rpdc ! in deg
      if(dcos(CLATPR).lt.dcos(CLATCR)*dcos(CLATMR)) then
        ALPHA=180.0-ALPHA ! obtuse case
      endif
      CA = dcos(ALPHA*rpdc) ! cos alpha
      BXC = BXM*CA + BYM*SA ! north (mag)
      BYC = BYM*CA - BXM*SA ! east (mag)
      BZC = BZM
      return
      end
c-----

```