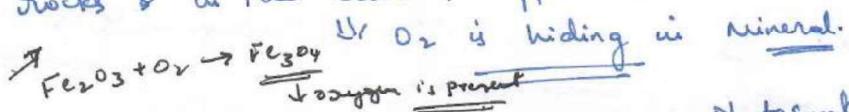


Oxygen in Earth System

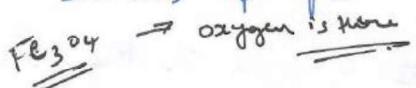
→ Atmospheric O₂ accounts for only a small fraction of the "free" oxygen in the Earth's system
 & where is the rest ??

↓

Much larger quantities of free oxygen are present in the form of oxidized minerals in the Sedimentary rocks & in the Crust & Upper Mantle.



Source of free oxygen



photosynthesis

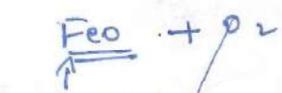
Redox reaction
in the Earth's
mantle.

Photosynthesis

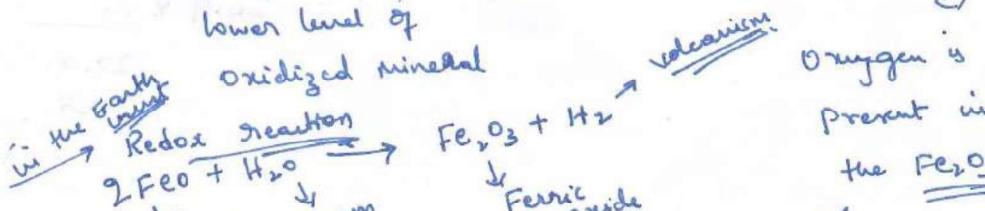


↓ radiation is absorbed → 0.43 to 0.66

Redox reaction (FeO reacts with O₂ due to oxidation to Fe₂O₃ for Fe₂O₃)



lower level of
oxidized mineral



Oxygen is
present in
the Fe₂O₃

The current level of oxidation in the Earth is explained by this Redox reaction. If this has a Redox reaction, we can't back the oxygen
 Photosynthesis is insufficient to explain the prevent oxidation of the earth → we will see this through some worked example

Mass of O_2 required to raise oxidation of Earth.
 units $10^3 \frac{kg}{m^2}$ avg over the surface of
 the earth.

Reservoir	Mass	
Atm O_2	2.353	
Oceans & Sediments	31	
Crust Fe^{3+}	> 100	
Crust CO_2	~ 100	
Crust (other)	> 100	
Mantle Fe^{3+}	> 100	

Some estimates
based on some scientific
arguments & measurements

Ques. # Recconcile the mass of Oxygen in the atm in the preceding table with the Vol Conc given in a composition table.

Soln. Vol' conc of $O_2 \rightarrow 20.9\%$ (From the composition table)

$$\text{Mass conc of } O_2 \rightarrow 20.9 \times \frac{32}{\text{Mol. wt of air}}$$

$$= 20.9 \times \frac{32}{28.91}$$

already estimated

$$\text{consider the } \xleftarrow{\text{Reconcile the}} \text{ Mass of } O_2 \text{ in the air} = 23.12\% \\ \text{in the atm} \approx 1.004 \times 10^4 \frac{kg}{m^2}$$

$$= \frac{23.12}{100} \times 1.004 \times 10^4 \frac{kg}{m^2}$$

$$= 2.353 \times 10^3 \frac{kg}{m^2} \rightarrow \underline{\text{Reconciled}}$$

? Using the data provided in the previous table (carbon determined), estimate the mass of oxygen required to form carbonate deposits in the earth's crust.

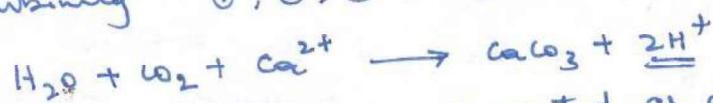
(Inorganic carbon in sedimentary rock is $80,000 \text{ kg/m}^2$ → How much O_2 is required in getting for this $80,000 \text{ kg/m}^2$)

Next question → how much O_2 is released from the photosynthesis.
How are we will you work it out

that is the proxy for the amt released during ↑ photosynthesis [plants, → decay, → die. It proxy is organic carbon.]
if you find out ← $20,000 \text{ kg/m}^2$
the amt of O_2 which has been used

Solⁿ for problem 12 → not a straightforward problem.
Some fundamental chemistry is required.

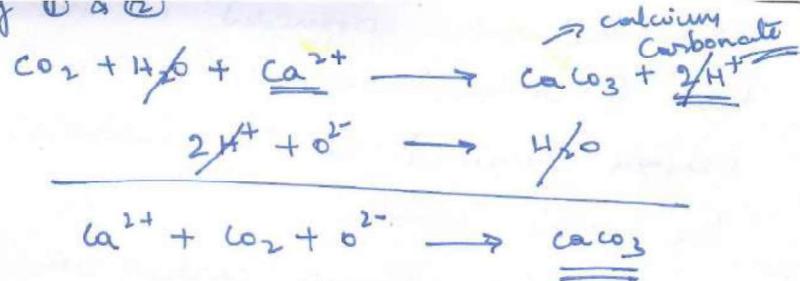
Combining ①, ②, ③ & ⑤



this two hydrogen ions generated as a by-product of this reaction combine with single O ion to form water.



Reading ① & ②



One free oxygen atom is paired with a carbon atom in CO_2 & resides in the earth crust as carbonate. Hence mass of oxygen required to form the carbonate reservoir is

$$8,000 \text{ (kg/m}^2) \times \frac{16}{12}$$

$\approx 10,000 \text{ kg/m}^2$ → This is the basis for writing crust \log in the previous table

Next Problem :- Using the data provided in the previous table, estimate the oxygen liberated by photosynthesis over the lifetime of the earth.

Solⁿ Proxy is the organic carbon in the carbon reservoir. for photosynthesis O_2 ~~ads~~ why.

Plant, respire, decay → die.

finally form as a organic carbon in the crust. → $20,000 \text{ kg/m}^2$

If you find out the amt of O_2 which has been used to form this $20,000 \text{ kg/m}^2$, \rightarrow that is the proxy for the amt released during photosynthesis.

What is the eqn we need to use here??



One oxygen atom with the burial of each carbon atom, one O_2 molecule is released.

Hence Many of oxygen liberated by the photosynthesis is

$$20,000 \text{ kg/m}^2 \times \frac{32}{12}$$

$$= \underline{\underline{53,000 \text{ kg/m}^2}}$$

We got from in-organic $CO_2 \rightarrow$ $100,000 \text{ kg/m}^2$
from photosynthesis \rightarrow $\underline{\underline{53,000 \text{ kg/m}^2}}$
↑ not matching

"Additional mechanism is required to the oxidation of the Earth" \rightarrow which is through the redox reaction.

The principle source of free completed the Introduction
part \rightarrow Quiz portion

For a black body $\frac{I_\lambda}{I}$ or L , we can call as $B_\lambda(T)$

\downarrow
Absorption.

Emissive power reflection, absorption for
of the black body. anything

$$\int_0^\infty E_b(T) d\lambda = \int_0^\infty B_\lambda(T) d\lambda$$

W
m².sr.m.sr.

Black body:

What is that B_λ \rightarrow for all
temp tend to 0, when $T \rightarrow 0$,
tend to ∞ , when $T \rightarrow \infty$
tend to 0, $\lambda \rightarrow 0$
& tend to ∞ , $\lambda \rightarrow \infty$.

which is a great puzzle. Max
Planck came with expression
in 1901 \rightarrow Planck's law.
then he figured out for this
to be true, $E = h\nu$.

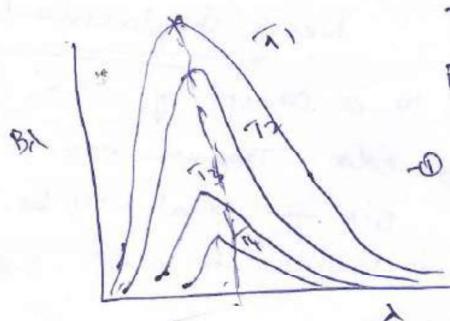
Black body:- is one which absorbs all incident radiation
regardless of its direction & wavelength.

$$B_\lambda(T) = \frac{C_1 \lambda^{-5}}{\pi \left[e^{\frac{C_2}{\lambda T}} - 1 \right]}$$

\rightarrow Planck's distribution

$$C_1 = 3.74 \times 10^{-16} \text{ W/m}^2 \rightarrow \text{First radiation const}$$

$$C_2 = 1.45 \times 10^4 \text{ m.K} \rightarrow \text{Second radiation const}$$



$T_1 > T_2 > T_3 > T_4$
Features of the Planck
distribution

- ① $B_\lambda(T)$ Continuously varies
with λ for every T
"not a discontinuous
fun"
- ② For every temp there
is a peak value of
 B_λ

⑤ This maximum J ~~is~~ with temperature

④ At a given λ , B_λ increases with temperature.

(No two curves for two different temp intersect with each other).

Now: $\frac{\partial B_\lambda(T)}{\partial \lambda} = 0$ & solving we get

$$= \frac{(-1) 5 \lambda^{-6}}{\left[e^{\frac{c_2}{\lambda T}} - 1 \right]} + \frac{(-1) \lambda^5 \cdot e^{\frac{c_2}{\lambda T}} \cdot (-c_2)}{\left[e^{\frac{c_2}{\lambda T}} - 1 \right]^2} \cdot \frac{c_2}{\lambda^2 T}$$

$$\frac{5 \lambda^6}{\left[\frac{c_2}{e^{\frac{c_2}{\lambda T}}} - 1 \right]} = \frac{\lambda^5 \cdot e^{\frac{c_2}{\lambda T}} \cdot c_2}{\left[e^{\frac{c_2}{\lambda T}} - 1 \right]^2} \cdot \frac{c_2}{\lambda^2 T}$$

If you re-arrange this eq:

$$\lambda_m = \frac{2897}{\Delta T} \quad \text{--- ⑥}$$

↗
Wein's displacement law

This also leads us to a concept of color. Today I am wearing blue color. Trouser. 0.6 m is the λ . If you put 0.4 \rightarrow what will be the temp. 6000K. What will happen to me? What is cold then.

The cold u see is not emission, rather it is ~~absorption~~, reflection.
It absorbs all the colors, but it reflects blue color.

Real emission color \rightarrow heat the iron piece. It will be

orange color. \rightarrow it is an emission color.

If nothing is coming out \rightarrow Black (Absence of color).

If it has to radiate black, for 0 to 4000 it has to absorb all the radiation. Radiatively black will be

Visibly black obj.

Use Wein's displacement law Compute the color temperature of the sun for which the λ for the maximum solar emission to be ~~0.475 nm~~. 0.5 Mm

$$\lambda_{max} T = 2898$$

$$0.5 \times 2898 = 1449$$

Some for
the earth is
 $TTE = 288K$
 $\lambda_{max} = 10.1\text{ }\mu\text{m}$

$$T_{sun} = \frac{1449}{5790} \approx 5800K$$

Derivation of Stefan-Boltzmann law from Planck distribution

$$E_b(T) = \int_0^{\infty} \pi B_x(T) d\lambda$$

\hookrightarrow Black body emissive power @ temp (T).

$$E_b(T) = \int_{\lambda=0}^{\infty} \pi \frac{C_1 \lambda^{-5}}{\pi \left[e^{C_2/\lambda T} - 1 \right]} \quad \text{--- (1)}$$

$$\text{Let } \frac{C_2}{\lambda T} = q$$

$$-\frac{C_2}{\lambda^2 T} d\lambda = dn.$$

$$\therefore E_b(T) = \int_{\infty}^0 -\frac{C_1 \lambda^5}{[e^{\lambda/T} - 1]} \cdot \frac{\lambda^2 T}{C_2} dn.$$

The area under the curve remains the same,

but it's $E_b(T)$ =
 labelled as negative.

$$E_b(T) = \int_0^{\infty} \frac{C_1}{C_2 \lambda^3} \frac{T}{[e^{\lambda/T} - 1]} dn$$

$$= \int_0^{\infty} \frac{C_1 T}{C_2 \left[\frac{C_2}{\lambda T} \right]^3 [e^{\lambda/T} - 1]} dn$$

$$= \int_0^{\infty} \frac{C_1}{C_2^4} \frac{T^4 n^3}{[e^{\lambda/T} - 1]} dn \quad \rightarrow \textcircled{2}$$

$$= \frac{C_1 T^4}{C_2^4} \underbrace{\int_0^{\infty} \frac{n^3}{[e^{\lambda/T} - 1]} dn}_0$$

very difficult to integrate.

But you can numerically Integrate. But it will give the result of $\frac{\pi^4}{15}$.

$$E_b(T) = \left(\frac{C_1}{C_2^4} \frac{\pi^4}{15} \right) T^4$$

$$\left[\frac{3.74 \times 10^{-7} \times (3.14)^4}{15 \times (1.438 \times 10^{-2})^4} \right] T^4$$

$$E_b(T) = \sigma T^4 \text{ where } \sigma = 5.67 \times 10^{-8} \frac{\text{W}}{\text{m}^2 \text{K}^4}$$

calculate the equivalent black body Temp (T_E) of the solar photosphere (outermost visible layer of the sun)

based on the following information:

→ The flux density of solar radiation reaching the earth = $F_s = 13.68 \text{ W/m}^2$

→ The earth-Sun distance ('d') is $1.5 \times 10^{11} \text{ m}$.

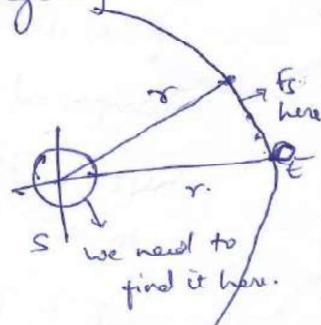
→ R_s of the solar photosphere (radius of the sun) $\rightarrow 7 \times 10^8 \text{ m}$

If you have the flux density at the photosphere, using σT^4 you could estimate the temp.

We know the F_s @ far distance, we need to get this for the photosphere. How are we going to do?

Using inverse square law.

$$F_s \propto \frac{1}{r^2}$$



$$T_{\text{photosphere}} \propto \frac{1}{r_s^2}$$

$$F_s \propto \frac{1}{d^2}$$

$$\frac{F_{\text{photosphere}}}{P_s} = \left(\frac{d}{R_s}\right)^2 \text{ or } \left(\frac{R_s}{d}\right)^{-2}$$

$$F_{\text{photosphere}} = 1368 \left(\frac{7 \times 10^8}{1.5 \times 10^{11}} \right)^{-2}$$

$$= 6.28 \times 10^7 \text{ W/m}^2$$

But $\sigma T_s^4 = F_{\text{photo}}$

$$T_s = \left[\frac{6.28 \times 10^7 \text{ (W/m}^2)}{5.67 \times 10^{-8} \text{ (W/m}^2\text{K}^4)} \right]^{1/4}$$

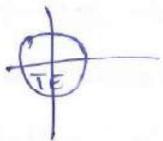
$$T_s = \underline{5769 \text{ K}} \approx \underline{5800 \text{ K}}$$

A small blackbody satellite is orbiting the Earth at a distance far enough away so that the flux density of Earth radiation is negligible, compared to that of Solar radiation.

Suppose that the satellite suddenly passes into the Earth's shadow. At what rate

will it initially cool? The satellite has mass $m = 10^3 \text{ kg}$ & a specific heat $C = 10^3 \text{ J} (\text{kg/K})^{-1}$: it is spherical with a radius $r = 1 \text{ m.}$, & the temp is uniform over its surface

Soⁿ. Consideration of energy balance
Radiative equilibrium:



Under equilibrium : the rate of change of
enthalpy = Incoming - outgoing

Enthalpy is a measure of the energy in a system
& its change is a measure of the heat
added or lost by the system

$$\underline{mc \frac{dT_E}{dt}} = \underline{\text{Incoming} - \text{outgoing}}$$

We need to find out $\underline{\frac{dT_E}{dt}}$

or change in temp.

$$F_E = \sigma T_E^4$$

$$T_E = \left(\frac{F_E}{\sigma} \right)^{1/4}$$

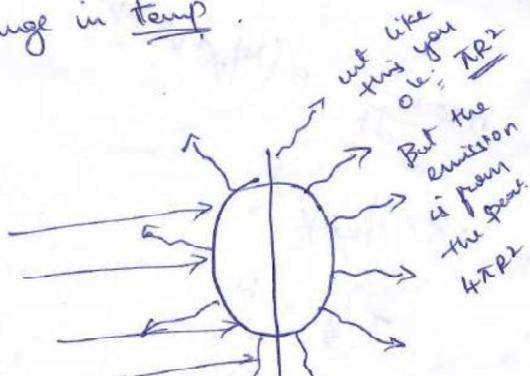
What is F_E ?

$$\text{Emission } (4\pi R^2) \sigma T_{sat}^4 = S \pi R^2$$

$$T_{sat}^4 = \frac{S}{4\pi}$$

$$T_{sat} = \left(\frac{S}{4\pi} \right)^{1/4} \cdot \left(\frac{1368}{4\pi \times 5.67 \times 10^{-8}} \right)^{1/4}$$

$$T_{sat} = 279K$$



When the satellite passes into the Earth shadow, it will no longer be in radiative equilibrium. The flux density of solar radiation abruptly drops to zero while the emitted radiation, which is determined by the temperature of the satellite, drops gradually as the satellite cools by emitting radiation.

$$\therefore mc \frac{dT_E}{dt} = \text{Incoming} - \text{outgoing}$$

$$mc \frac{dT_E}{dt} = -\text{outgoing}$$

$$\frac{dT_E}{dt} = - \frac{(4\pi R^2) \sigma T_E^4 (0.07 K^4) \cancel{J/m^2}}{mc} \frac{W/K}{J/K}$$

$$mc \frac{dT_E}{dt} = (0.07 K^4) \cancel{J/m^2} \frac{W/K}{J/K} = - \frac{4315.086}{mc} \frac{K/s}{J/s}$$

$$\therefore \frac{dT_E}{dt} = -4.315 \times 10^{-3} K/s$$

~~J/s~~

~~J/s~~

$\frac{kg \times 10^3 J}{kg \cdot K} \cdot \frac{W}{J/s}$

~~watts~~

$\frac{W}{s} = \frac{W}{s}$

$10^3 = 10^3$

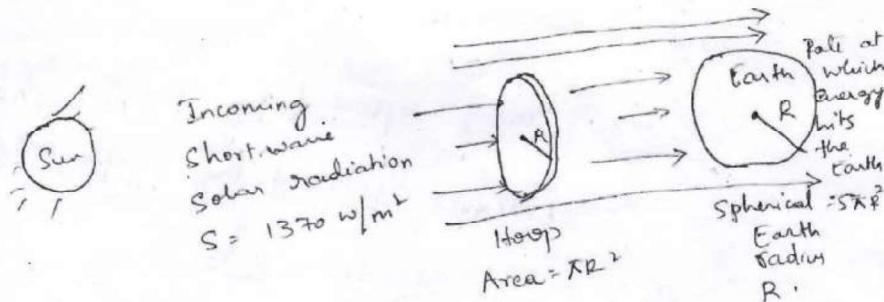
$\therefore 15.5 K/\underline{\text{hr}}$

A Simple Global temperature model

→ The simplest starting point for modeling climate begins with models that focus on factors influencing the single parameter temperature

→ Other factors such as precipitation patterns, winds and storms & so forth are exceedingly important but they are more difficult to approach.

Simple climate model → "Zero-dimensional Model" in which a single average global temperature is obtained that is not a function of time, latitude, longitude.



The basic zero-dimensional energy balance model equates solar energy absorbed by the Earth with the energy that the Earth radiates back to space.

Radiation from the Sun arrives just outside the Earth's atmosphere with an avg annual intensity, called the solar constant, $S = 1370 \text{ W/m}^2$

Rate of Solar energy strikes the Earth

$$= SR^2 \text{ (Watts)}$$

where $S = \text{Solar Const} = 1370 \text{ W/m}^2$

$R = \text{Radius of the Earth (m)}$

Some of the incoming Solar energy that hits the Earth is reflected back into Space. \rightarrow will not contribute to their heating.

The fraction of incoming Solar radiation that is reflected is called the albedo $\rightarrow \underline{\underline{31\%}}$

$\therefore \text{Energy reflected by Earth} = SR^2 \alpha$

$\text{Energy absorbed by Earth} = SR^2(1-\alpha)$

$$\alpha = \text{Earth albedo} = \underline{\underline{0.31}}$$

on the other hand, the earth radiates back the energy \rightarrow in the form of radiation.

Every object radiates the energy at a rate α to its surface area \times its absolute temperature raised to the 4^{th} power.

$$\underline{\underline{\alpha 4\pi R^2 T_e^4}}$$

For this model \rightarrow Assume Earth as a black body.

Black body \rightarrow it radiates as much as any object with the same temperature and area possibly can.

\rightarrow Also assume the earth is an isothermal

\therefore Energy radiated back to the space by

$$\text{Earth} = \sigma A \pi R^2 T_e^4$$

where σ = Stefan-Boltzmann Const $= 5.67 \times 10^{-8}$ $\frac{\text{W m}^2 \text{K}^4}{}$

and T_e = earth's "effective black body" temperature (K)

@ steady state

Rate of energy from sun absorbed =

Rate of energy emitted/radiated back to the space.

$$S \pi R^2 (1 - \alpha) = d \sigma A \pi R^2 T_e^4$$

Solving for T_e

$$T_e = \left[\frac{S(1-\alpha)}{d\sigma} \right]^{1/4}$$

Substituting

$$S = 1370 \frac{\text{W}}{\text{m}^2}, \alpha = 0.31$$

$$\sigma = 5.67 \times 10^{-8} \frac{\text{W}}{\text{m}^2 \text{K}^4}$$

We get

$$T_e = \left[\frac{1370 \text{ W/m}^2 (1 - 0.31)}{4 \times 5.67 \times 10^{-8} \text{ W/m}^2 \text{ K}^4} \right]^{1/4}$$

$$T_e = 254 \text{ K} = \underline{\underline{-19^\circ\text{C}}} \quad \text{--- ①}$$

But the average temperature is 15°C

What is wrong here ??

that means Earth is 34°C more

The Greenhouse effect warmer.

Wien's displacement rule gives the λ at which a blackbody spectrum peaks as a function of its absolute temperature

$$\lambda_{\text{max}}(\text{m}) = \frac{2898}{T(\text{K})}$$

